

New Taxa and New Middle Famennian *Natgorella hirsuta*–*Spinoalius melekessensis* Radiolarian Ecozone of the Volga–Ural Basin

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Abstract—New radiolarian taxa are described from the Lower and Middle Famennian of the Polar Urals and Volga–Urals Basin: three new genera *Natgorella* gen. nov., *Spinoalius* gen. nov., *Diversiacus* gen. nov., and three new species *N. elegans* sp. nov., *N. hirsuta* sp. nov., and *S. melekessensis* sp. nov. New taxa belong to two orders, Entactiniata and Cancelliata, of two classes of spherical radiolarians: Sphaerellaria and Spumellaria. Morphotypes of the outer skeletons of *Spinoalius* gen. nov. and *Diversiacus* gen. nov. are shown to be convergently similar. A new Middle Famennian *Natgorella hirsuta*–*Spinoalius melekessensis* radiolarian ecozone is established.

Keywords: Radiolaria, new ecozone, new taxa, convergence, morphotypes, Late Devonian, Famennian, Polar Urals, Volga–Ural Basin

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HISTORY OF STUDY

In the mid-20th century, Bykova (1955) studied the so-called “primitive” small foraminifera from the Lower and Middle Famennian of the Volga–Ural Basin (Figs. 1a–1p). Bykova (1955, p. 110) for the first time commented on a certain similarity of some *Archaesphaera* and *Parathurammia* with radiolarians: “in the shell structure they are closer to ancient radiolarians than to extant or Mesozoic–Cenozoic planktonic foraminifers.” The only difference between these organisms and radiolarians, according to Bykova, was that their skeleton was calcareous, whereas the radiolarian skeleton is siliceous.

At the same time, it is known that during lithification of the host rock, radiolarian skeletons can be recrystallized, resulting in the either replacement of the primary opal of the skeleton by other minerals (calcite, apatite, fluorite, garnets, pyrite, and gold), or only the spheres remain of the skeleton, and are often misidentified as small primitive foraminifers or algae. All the above suggests that some species from the genera *Archaesphaera* and *Parathurammia* can be considered as radiolarians (Braun and Amon, 1991; Afanasieva, 2000; Vishnevskaya and Sedaeva, 2002; Sedaeva and Vishnevskaya, 2002, 2008; Afanasieva and Amon, 2011, 2012). However, this opinion is not shared by all paleontologists.

Almost a quarter of a century later, undoubtedly Early Famennian radiolarians with a siliceous skeleton were found in the Eletsian Regional Stage of the Pripyat Depression (Figs. 1, V; 2, V; 3, V) and Nazarov

studied them using an electron microscope (Kruchek and Nazarov, 1977; Nazarov, 1988). Later, Early Famennian radiolarians were studied (Afanasieva, 2000) on the Saremboy–Lekkeyaga Borehole in the Timan–Pechora Basin and in the north slope of the Peri-Caspian Basin (Figs. 1, I, III; 2, I, III; 3, I, III) and were found in the Polar Urals and in the Vilva River basin on the western slope of the Middle Urals (Figs. 1, II, IV; 2, II; 3, II, IV) (Afanasieva and Amon, 2011, 2012).

In the Volga–Ural Basin, no confirmed Famennian radiolarians have been recorded. At present, L.I. Kononova during the processing of conodont samples from the Middle Famennian Lebedyanian Regional Substage studied in the Melekess R-1 Borehole, found excellently preserved whole radiolarian shells (Figs. 1, VI; 2, VI; 3, VI).

This paper discusses the results of the study of the first-ever discovered true Middle Famennian radiolarians in the Volga–Ural Basin, and the description of the new Middle Famennian *Natgorella hirsuta*–*Spinoalius melekessensis* radiolarian ecozone.

RESULTS

The New Natgorella Hirsuta–*Spinoalius Melekessensis* Radiolarian Ecozone

A new association of the Middle Famennian radiolarians is established in the Melekess Depression of the Volga–Urals Basin. Radiolarians were found in

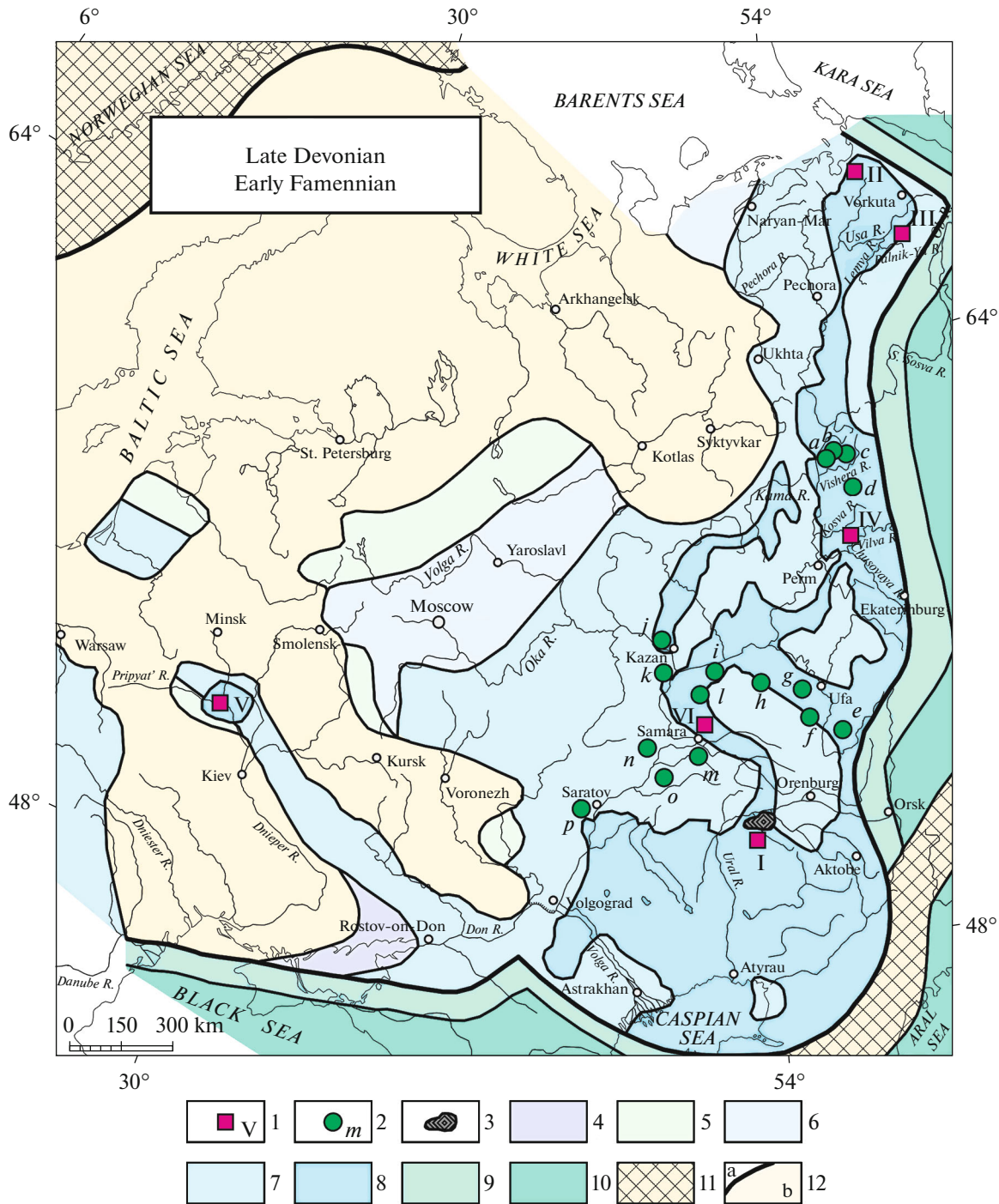


Fig. 1. Distribution of Famennian radiolarians; explanations: (1) locality of the radiolarian ecozones: (I) Peri-Caspian Basin, (II) Polar Urals, Lemva Zone, (III) Timan–Pechora Basin, Saremboy–Lekkeyaga Borehole, (IV) western slope of the Middle Urals, (V) Pripjat Depression, (VI) Volga–Ural Basin (I, III Afanasieva, 2000; II–IV Afanasieva and Amon, 2011, 2012; V Kruchek and Nazarov, 1977; VI this paper); (2) “Uralian” assemblage (Bykova, 1955): (a–d) North Urals: (a) Petrunikha River, (b) Storozhevaya River, (c) Vishera River, (d) Yazva River; (e) South Urals, Sikaza River; (f–h) western Bashkortostan: (f) Elatminki (=Novokonstantinovka), (g) Kargaly River, (h) Tuimazy; (i–k) Tatarstan: (i) Shugurovo Village, (j) Kamskoe Ustie Village, (k) Verkhniy Uslon Village; (l–n) Samara Region: (l) Krasnovka Village, (m) Berezovka Village, (n) Syzran; (o, p) Saratov Region: (o) Pugachev, (p) Teplovka Village; (3) Karachaganak Fossil Reef; (4–10) paleobiogeographic scheme of the Famennian of the Russian Platform and the Greater Urals (Afanasieva, 2000; Afanasieva et al., 2011; Afanasieva and Amon, 2012): (4) low salinity sea; (5–10) normal salinity sea: (5) coastal region, (6) shallow inner shelf, (7) inner shelf, (8) outer shelf; (9) continental slope, (10) paleocean seafloor; (11) active foldbelts; (11) land; (12a) margin of the Russian Platform, (12b) land.

the middle subformation of the Malaya Cheremshana Formation (interval 1940–1980 m) at two levels of the Melekess P-1 Borehole (Sample 7336, interval 1961.9–1963.9 m; Sample 7398, interval 1973.9–1974.7 m)¹ (Figs. 1, VI; 2, VI; 3, VI). The samples were disintegrated in 10% acetic acid solution using the standard procedure for isolating conodonts and radiolarians. Together with radiolarians in Sample 7398, Kononova discovered the conodonts *Palmatolepis glabra lepta* Ziegler and Huddle, 1969. This subspecies ranges from the Late *crepida* Zone to the Late *trachytera* Zone. In Sample 7336, no conodonts were recorded. The age of the interval with radiolarians is emended based on a more diverse conodont assemblage identified in the lower part of the middle subformation (interval 1973.9–1976.5 m): *Palmatolepis glabra lepta* Ziegler and Hudd., *P. glabra pectinata* Ziegler, *P. gracilis gracilis* Branson and Mehl, *P. utahensis* Ziegler and Sandberg, *P. perlobata helmsi* Ziegler, *P. perlobata sigmoidea* Ziegler, *P. perlobata schindewolfi* Müller, *P. inflexa* Müller, *P. rugosa grossi* Ziegler, and *P. marginifera marginifera* Helms. Many of these species and subspecies have large stratigraphic ranges in the Famennian, but *P. utahensis* occurs only in the Late *marginifera* Zone, and *P. inflexa* ranges from the Late *rhomboidea* Zone to the Early *marginifera* Zone. This conodont association occurs almost 2 m below Sample 7398 with radiolarians, and is characteristic of the Early *marginifera* and Late *marginifera* zones, corresponding to the lower part of the Middle Famennian Lebedyanian regional stage (Rodionova et al., 1995).

Thus, the new *Natgorella hirsuta*–*Spinoalius melekessensis* ecozone characterizes part of the middle subformation of the Malaya Cheremshana Formation, which Kononova assigned based on conodonts to the Lebedyanian Regional Substage, which correlates with the Late *marginifera* and *trachytera* zones (Figs. 2, VI; 3, VI) (Fortunatova et al., 2016, 2018).

The taxonomic diversity of the new ecozone is characterized by the presence of six species of radiolarians (Figs. 2; 3): four species of three genera of latticed radiolarians of the class Spumellaria (Figs. 4a, 4b; Pl. 1) and two species of two genera of porous radiolarians of the class Sphaerellaria (Fig. 4c; Pl. 2).

Typical species of the new assemblage include *Natgorella elegans* sp. nov. and *N. hirsuta* sp. nov. (pl. 1), as well as *Spinoalius melekessensis* sp. nov. (Pl. 2, figs. 1–4). The species *Retientactinosphaera magnifica* Afanasieva, 2011, *Polyentactinia rudihispida* Nazarov and Ormiston, 1993 and *Diversiacus zuraevi* (Afanasieva and Amon, 2011) are represented by a few specimens (Fig. 4).

P. rudihispida is found in all studied Lower and Middle Famennian radiolarian localities (Fig. 1, I–VI).

However, this species first appeared in the Timan–Pechora Basin in the Frasnian (Fig. 3, I–VI) (Afanasieva, 2000). The species *D. zuraevi* (Fig. 2, II, VI) and *R. magnifica* (Fig. 3, III, VI) are characteristic only of the Famennian (Afanasieva, 2011; Afanasieva and Amon, 2011). At the same time these species are very important, as they are shared with the new radiolarian association of the Volga–Ural Basin (Middle Famennian, Lebedyanian regional substage) and older Lower Famennian ecozones² (Afanasieva and Amon, 2011, 2012) (Figs. 2, 3):

Tetrentactinia barysphaera–*Holociscus auceps*, Peri-Caspian Basin of Kazakhstan, Karachaganak Fossil Reef, lower Famennian (Afanasieva, 1987, 2000);

Tetrentactinia barysphaera–*Caspiaza spinifera*, Polar Urals, Lemva Zone, Palnik–Yu River, *triangularis* and *crepida* conodont zones (Afanasieva, 2000; Afanasieva and Amon, 2011, 2012);

Tetrentactinia barysphaera–*Retientactinosphaera magnifica*, Timan–Pechora Basin, Saremboy–Lekkeyaga Borehole, Zapadno–Lekkeyaginskaya Borehole-65 (interval 2460–2467 m), *triangularis* and *crepida* conodont zones (Afanasieva, 2000; Afanasieva and Amon, 2011, 2012);

Haplentactinia alekseevi–*Haplentactinia vilvaensis*, western slope of the Middle Urals, Vilva River, *triangularis* and *crepida* conodont zones (Afanasieva and Amon, 2011, 2012);

Tetrentactinia barysphaera–*Ceratoikiscum famennium*, Pripyat Depression of Belarus, Eletsian Regional Substage; conodonts collected alongside radiolarians include: *Palmatolepis glabra* Ulrich and Bassler, *P. minuta* Branson and Mehl., *P. aff. quadrantinoslobata* Sannemann, *Prioniodina smithi* Stauffer, *Scutula venusta* Sannemann (Kruchek and Nazarov, 1977; Nazarov, 1988), characterizing the *rhomboidea* Zone;

Natgorella hirsuta–*Spinoalius melekessensis*, Volga–Ural Basin, Melekess Depression, Lebedyanian Regional Substage, Late *marginifera* and *trachytera* conodont zones.

The radiolarian association of the new ecozone is surprisingly similar to that of the *Haplentactinia alekseevi*–*Haplentactinia vilvaensis* ecozone of the Middle Urals (Afanasieva and Amon, 2011, 2012). The similarity is observed in the presence of the same radiolarian morphotypes, in both cases represented mainly by lattice skeletons, and in the presence of the shared species *Polyentactinia rudihispida*, but the taxonomic composition of the ecozones is different.

Hereby I consider the new Middle Famennian radiolarian association as the *Natgorella hirsuta*–*Spinoalius melekessensis* ecozone, corresponding to the

¹ For a detailed lithological and stratigraphic description of the Melekess P-1 Borehole and paleontological substantiation see Fortunatova et al. (2018).

² “Ecozone is a complex zone, beds in which contains a fossil assemblage reflecting an ecological association during lifetime or the taphonomic features of their oryctocenosis” (*Stratigraphic ...*, 2006, Article VII, 4g).

Stage Substage	Famennian					
	Lower					Middle
	<i>triangularis–crepida</i>				<i>rhomboidea</i>	<i>L. marginifera–trachytera</i>
Conodont zones						
Radiolarian ecozones	I	II	III	IV	V	VI
<i>Bientactinosphaera</i> sp.						
<i>Entactinia micula</i> Foreman, 1963						
<i>Entactinia</i> sp.						
<i>Radiobisphaera domanicensis</i> (Bykova, 1955)						
<i>Astroentactinia</i> sp.						
<i>Borisella praenuntia</i> Nazarov and Ormiston, 1993						
<i>Bientactinosphaera guangxiensis</i> (Li and Wang, 1991)						
<i>Bientactinosphaera obtusa</i> (Hinde, 1899)						
<i>Entactinia oumonhaoensis</i> Wang, 1997						
<i>Entactinia patorovaria</i> Afanasieva, 2000						
<i>Radiobisphaera assidera</i> (Nazarov, 1975)						
<i>Radiobisphaera menneri</i> Afanasieva, 2000						
<i>Radiobisphaera palimbola</i> (Foreman, 1963)						
<i>Entactinia bella</i> Afanasieva and Amon, 2011						
<i>Entactinia bogdanovi</i> Afanasieva, 2000						
<i>Entactinia crustescens</i> Foreman, 1963						
<i>Entactinia parva</i> Won, 1983						
<i>Ornatoentactinia solita</i> Afanasieva, 2000						
<i>Radiobisphaera rozanovi</i> Afanasieva and Amon, 2009						
<i>Astroentactinia biaciculata</i> Nazarov, 1975						
<i>Helioentactinia valavica</i> Nazarov and Ormiston, 1993						
<i>Astroentactinia crassata</i> Nazarov, 1975						
<i>Astroentactinia stellata</i> Nazarov, 1975						
<i>Bientactinosphaera variacanthina</i> (Foreman, 1963)						
<i>Entactinia herculea</i> Foreman, 1963						
<i>Diversiacus zuraevi</i> (Afanasieva and Amon, 2011)						
<i>Bientactinosphaera maslakovae</i> Afanasieva, 2000						
<i>Bientactinosphaera pinica</i> Afanasieva, 2000						
<i>Duplexia spinocurva</i> Afanasieva, 2000						
<i>Bientactinosphaera spinofoliacea</i> Nazarov and Afanasieva, 2000						
<i>Bientactinosphaera symphympora</i> (Foreman, 1963)						
<i>Bientactinosphaera echinata?</i> (Hinde, 1899)						
<i>Bientactinosphaera euthlasta</i> (Foreman, 1963)						
<i>Bientactinosphaera pulcherima</i> (Nazarov and Ormiston, 1993)						
<i>Entactinia cometes</i> Foreman, 1963						
<i>Entactinia monalloeae</i> Foreman, 1963						
<i>Entactinia paula</i> Foreman, 1963						
<i>Entactinosphaera riedeli</i> Foreman, 1963						
<i>Spinoalius melekessensis</i> Afanasieva, sp. nov.						

Fig. 2. Famennian Sphaerellaria: (I–VI) ecozones of Famennian radiolarians: (I) *Tetrentactinia barysphaera*–*Holoeciscus auceps*, Peri-Caspian Basin, Karachaganak Fossil Reef; (II) *Tetrentactinia barysphaera*–*Caspiaza spinifera*, Polar Urals, Lemva Zone; (III) *Tetrentactinia barysphaera*–*Retientactinosphaera magnifica*, Timan–Pechora Basin, Saremboy–Lakkeyaga Borehole; (IV) *Haplentactinia alekseevi*–*Haplentactinia vilvaensis*, western slope of the Middle Urals, Vilva River; (V) *Tetrentactinia barysphaera*–*Ceratoikiscum famennium*, Pripyat Depression; (VI) *Natgorella hirsuta*–*Spinoalius melekessensis*, Volga–Ural Basin, Melekess Depression (I, III) Afanasieva, 2000; (II–IV) Afanasieva and Amon, 2011, 2012; (V) Kruchek and Nazarov, 1977; (VI) Afanasieva, this paper).

Middle Famennian Lebedyanian Regional Substage and correlating with the Late *marginifera* and *trachytera* conodont zones. The new radiolarian ecozone is recognized in the territory of the Volga–Ural Basin. The substantiation of a new ecozone of radiolarians is important for the Devonian stratigraphy of the region for a number of reasons. For instance, this is the first

locality for true Middle Famennian radiolarians, the geochronology of which is supported by conodonts. In addition, the stratigraphic range of radiolarians is expanded from the *triangularis* to *trachytera* conodont zones (Figs. 2, 3). Knowledge on the taxonomic and morphological diversity of radiolarians in the Lower–Middle Famennian ecozones is enhanced. The assem-

Stage Substage Conodont zones Radiolarian ecozones	Famennian					
	Lower					Middle
	<i>triangularis–crepida</i>				<i>rhomboidea</i>	<i>L. marginifera–trachytera</i>
	I	II	III	IV	V	VI
<i>Somphoentactinia somphozona</i> (Foreman, 1963)						
<i>Pluristratoentactinia conspissata</i> Nazarov, 1981						
<i>Tetrentactinia barysphaera</i> Foreman, 1963						
<i>Polyentactinia rudihispida</i> Nazarov and Ormiston, 1993						
<i>Pluristratoentactinia trisphaerata</i> Afanasieva and Amon, 2011						
<i>Tetragregnon sycamorensis</i> Ormiston and Lane, 1976						
<i>Adamasirad cathedrarius</i> Afanasieva, 2000						
<i>Retientactinosphaera clavata</i> Afanasieva, 2011						
<i>Spongentactinia marina</i> Afanasieva and Amon, 2011						
<i>Spongentactinia polaris</i> Afanasieva and Amon, 2011						
<i>Spongentactinia diplostraca</i> (Foreman, 1963)						
<i>Haplentactinia kuzminae</i> Afanasieva and Amon, 2011						
<i>Haplentactinia rhinophyusa</i> Foreman, 1963						
<i>Somphoentactinia multisphaerata</i> Afanasieva and Amon, 2011						
<i>Spongentactinella exilispina</i> (Foreman, 1963)						
<i>Tetragregnon quadrispinosa</i> (Foreman, 1963)						
<i>Russirad kazintsovae</i> Afanasieva, 2000						
<i>Polyentactinia suave</i> (Nazarov, 1977)						
<i>Retientactinosphaera magnifica</i> Afanasieva, 2011						
<i>Haplentactinia alekseevi</i> Afanasieva, 2000						
<i>Haplentactinia vilvaensis</i> Afanasieva and Amon, 2011						
<i>Haplentactinia flagellifera</i> Nazarov and Ormiston, 1993						
<i>Retientactinosphaera unimana</i> (Nazarov, 1977)						
<i>Secuicollacta amoenitas</i> Nazarov and Ormiston, 1993						
<i>Spongentactinella faceta</i> Nazarov and Ormiston, 1993						
<i>Spongentactinia spongites</i> (Foreman, 1963)						
<i>Tetrentactinia gracilispinosa</i> Foreman, 1963						
<i>Natgorella elegans</i> Afanasieva, sp. nov.						
<i>Natgorella hirsuta</i> Afanasieva, sp. nov.						

Fig. 3. Famennian Spumellaria: (I–VI) ecozones of Famennian radiolarians (see Fig. 2).

blages differ in taxonomic composition, but at the same time are dominated by lattice representatives of the class Spumellaria and have the shared species *Retientactinosphaera magnifica* and *Polyentactinia rudihispida*.

Convergent Evolution

E.O. Schmidt introduced the term ‘convergent evolution’ in biology to indicate convergent or parallel emergence of isolated morphological characters or sets of characters, for taxa of various ranks (Mayr, 1970). Convergent evolution is “a process of irreversible acquisition of characters by members of two or more phyletic lineages in connection with their life-style, but not connected to their ancestry” (*Paleontologiya...*, 1995, p. 165).

All convergent characters are secondary developments that emerged independently in the evolutionary process in order to adapt to a specific environment. The convergent similarity of morphotypes is never

deep, and involves only superficial morphological characters. Therefore in cases of convergent evolution even within the same taxonomic group, convergent characters are never exact copies of characters of other organisms. The resemblance of the superficial morphology of the compared taxa is usually partial and relative, rather than the morphology being entirely identical.

The study of the skeletons of the new radiolarian association revealed convergent morphological similarity of morphotypes of the outer appearance of skeletons of the new genera *Spinoalius* gen. nov. and *Diversiacus* gen. nov. (Fig. 5). The convergent similarity of these genera is based on a set of shared characters: (1) outer porous sphere; (2) six main spines, of which five spines are beam-like, and one is four-bladed or three-bladed (pl. 2), (3) rimming of the base of the beam-like spines by four petal-like pores (Pl. 2, figs. 1a, 1b; 7).

Differences are observed in the inner structure of the skeletons, which are characters of the subfamily

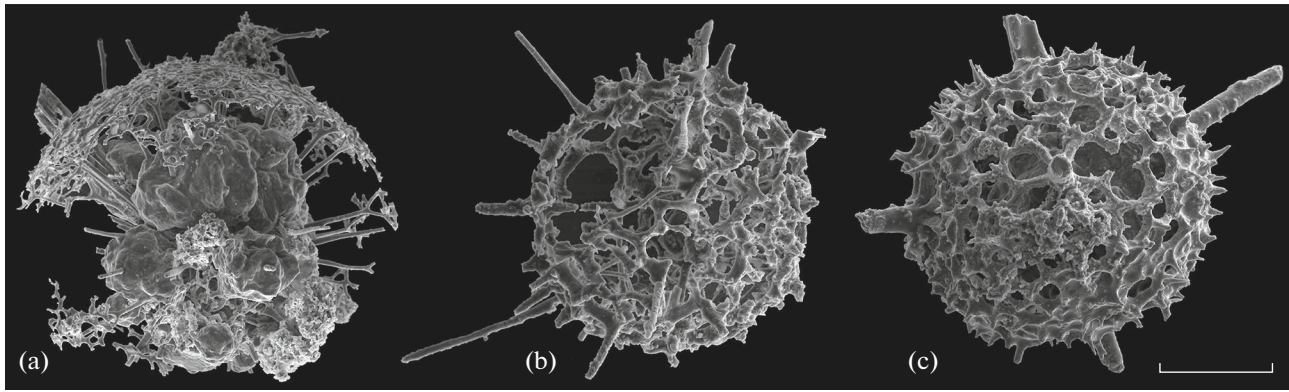


Fig. 4. Radiolarians of the Melekess Depression: (a) *Retientactinosphaera magna* Afanasieva, 2011, specimen PIN, no. 5369/187-L1015 (scale bar = 131 μm); (b) *Polyentactinia rudihispida* Nazarov and Ormiston, 1993, specimen PIN, no. 5369/0132016-019 (scale bar = 71 μm); (c) *Diversiacus zuraevi* (Afanasieva and Amon, 2011), specimen PIN, no. 5369/329-L1024 (scale bar = 92 μm); Upper Devonian, Middle Famennian, Lebedyanian Regional Substage, Late *marginifera* and *trachytera* conodont zones; Melekess P-1 Borehole: (a, b) Sample 7336 (interval 1961.9–1963.9 m); (c) Sample 7398 (interval 1973.9–1974.7 m).

level (Afanasieva, 2000; Afanasieva et al., 2005; Afanasieva and Amon, 2006): the genus *Spinoalius* is characterized by the presence of only one outer sphere (subfamily Entactiniinae) (Fig. 5a, Pl. 2, fig. 2a), whereas the genus *Diversiacus* differs in having two spheres, outer and inner (subfamily Bientactinosphaerinae) (Fig. 5b; Pl. 2, figs. 5, 6, 8).

Thus, in the two subfamilies Entactiniinae and Bientactinosphaerinae, independently similar, geochronologically parallel outer skeletons arose, which may indicate a similar habitat, and similar direction of natural selection. This difference in the morphology of skeletons justifies the recognition of two new genera: *Spinoalius* gen. nov. and *Diversiacus* gen. nov.

MATERIAL

Collections of Famennian radiolarians, nos. 5312 and 5369, are housed in the A.A. Borissiak Paleontological Institute (PIN). Radiolarian taxonomy follows Afanasieva and Amon (2003, 2006) and Afanasieva

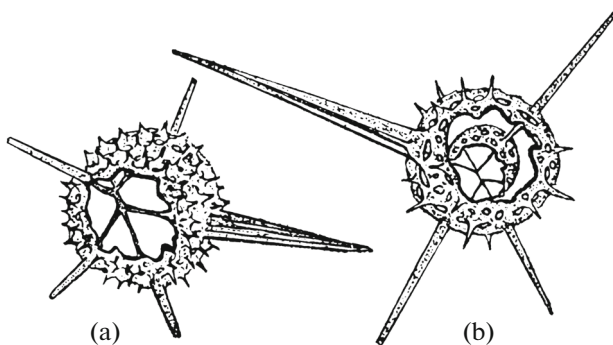


Fig. 5. Convergent similarity of morphotypes of radiolarian outer skeletons: (a) *Spinoalius*, (b) *Diversiacus*.

et al. (2005). Measurements of the radiolarian skeleton and the morphometric analysis of the absolute and relative values of the Paleozoic radiolarian shells followed a standard method (Fig. 6) (Afanasieva, 2000; Afanasieva et al., 2005; Afanasieva and Amon, 2006).

SYSTEMATIC PALEONTOLOGY

PHYLUM RADIOLARIA

SUPERCLASS POLYCYSTINA

CLASS SPUMELLARIA

Order Cancelliata

Family Haplentactiniidae Nazarov in Nazarov and Popov, 1980

Subfamily Pseudorotasphaerinae Noble, 1994

Genus *Natgorella* Afanasieva gen. nov.

Etymology. In honor of the paleontologist and stratigrapher Nataliya V. Goreva; feminine.

Type species. *Natgorella hirsuta* sp. nov.

Diagnosis. Shell spherical, with two lattice shells, formed by entangled rounded bars. Rays of inner six-ray spicule continue as six main rod-like spines with rounded cross-section. Surface of bars of outer spheres with two types of spinules: conical and thin rod-like, bifurcating, in approximately upper third of length.

Species composition. Type species and *N. elegans* sp. nov.

Comparison. This genus differs from the genera *Cancellosphaera* Afanasieva, 2000, *Russirad* Afanasieva, 2000 and *Arcoclathrata* Afanasieva and Amon, 2016, which have a latticed outer shell, a porous inner shell and six main 3-bladed spines, in the two lattice spheres, and six main rod-like spines. The new genus is distinguished from the genus *Retientacti-*

Size in μm	Absolute									Relative						
	D	Ds	dp(ds)	wp(ws)	tw	ts	L	wL	ht	D/Ds	D/tw (Ds/ts)	D/dp (Ds/ds)	D/ht	L/D	L/wL	dp/wp (ds/ws)
Very large, very long, etc.	360.0– 180.1	100.0– 60.1	35.0– 25.1	10.0– 4.1	8.0– 7.1	5.0– 3.1	750.0– 250.1	60.0– 50.1	20.0– 15.1	1.5– 2.5	1.0– 20.0	4.0– 10.0	3.0– 20.0	10.0– 2.1	0.5– 3.0	0.1– 1.5
Large, long, etc.	180.0– 130.1	60.0– 46.1	25.0– 15.1	4.0– 3.1	7.0– 4.1	3.0– 2.1	250.0– 153.1	50.0– 25.1	15.0– 10.1	2.6– 3.5	20.1– 30.0	10.1– 18.5	20.1– 36.0	2.0– 1.0	3.1– 6.0	1.6– 4.0
Small, short, etc.	130.0– 88.1	46.0– 40.1	15.0– 6.1	3.0– 2.1	4.0– 2.1	2.0– 1.0	153.0– 52.1	25.0– 15.1	10.0– 5.1	3.6– 4.0	30.1– 100.0	18.6– 63.0	36.1– 60.0	0.9– 0.4	6.1– 13.0	4.1– 6.0
Very small, very short, etc.	88.0– 20.0	40.0– 35.0	6.0– 1.0	2.0– 1.0	2.0– 1.0	–	52.0– 15.0	15.0– 4.5	5.0– 1.0	4.1– 4.5	100.1– 150.0	63.1– 105.0	60.1– 100.0	0.3– 0.1	13.1– 25.0	6.1– 9.5

Fig. 6. Major absolute and relative morphometric parameters of skeletons of Paleozoic spherical radiolarians (Afanasieva, 2000; Afanasieva et al., 2005; Afanasieva and Amon, 2006): (D) outer sphere diameter, (Ds) primary sphere diameter, (dp) diameter of pores of the outer sphere, (ds) diameter of pore of the primary sphere, (wp) inter-pore width of the outer sphere, (ws) inter-pore width of the primary sphere, (tw) wall thickness of the outer sphere, (ts) wall thickness of the primary sphere, (L) length of main spines, (wL) width of the base of the main spines, (ht) height of spinules.

nosphaera Afanasieva, 2011, which has only one three-bladed spine along with six main rod-like spines.

Natgorella hirsuta Afanasieva sp. nov.

Plate 1, figs. 4–6

E t y m o l o g y. From the Latin *hirsutus* (hirsute).

H o l o t y p e. PIN, no. 5369/327-M; Volga–Ural Basin, Melekess Depression, Melekess P-1 Borehole (interval 1973.9–1974.7 m); Upper Devonian, Middle Famennian, Lebedyanian Regional Substage, Late *marginifera* and *trachytera* conodont zones.

D e s c r i p t i o n. The shell is spherical, very large ($D = 231\text{--}248\ \mu\text{m}$), composed of two lattice shells with six main rod-like spines and numerous spinules. The outer sphere is composed of thin entangled bars ($D/tw = 34.4\text{--}42.8$), forming very large ($D/dp = 8.3\text{--}10.2$) polygonal pores of approximately the same size. The inner sphere is large ($D/Ds = 3.1\text{--}3.3$), relatively thin-walled ($Ds/ts = 23.3\text{--}25.7$) and is 30.3–31.9% of the diameter of the outer spheres. It is formed by entangled bars surrounding very large ($Ds/ds = 5.9$) polygonal pores. Rays of the six-ray spicules are represented by six relatively short ($L/D = 0.7\text{--}1.0$) and thin ($L/wL = 7.8\text{--}13.0$) main rod-like spines. The bars of the outer sphere possess small spinules of two kinds: (1) few, but relatively long ($D/ht_1 = 15.4\text{--}20.1$) conical spinules; (2) very numerous and very long ($D/ht_2 = 2.3\text{--}3.4$) rod-like thin spinules, double-bifurcated at approximately mid-length. The spinules are 29.1–44.4% of the diameter of the outer sphere, and the double bifurcation of numerous thin spinules (Pl. 1, fig. 5b) gives the species a shaggy, scruffy appearance, which is reflected in the species name.

Dimensions in μm and ratios (Fig. 6):

Specimen no.	D	dp	tw	Ds	ds	ts	L	wL	ht ₁	ht ₂
5369/327-M1016 (holotype)	248	21	5	–	–	–	245	20	14	110
5369/327-M1017	241	29	6	77	13	3	221	17	12	70
5369/327-M1023	231	23	6	70	–	3	155	20	15	73

Specimen no.	D/tw	D/dp	L/D	L/wL	D/ht ₁	D/ht ₂	D/Ds	Ds/ds	Ds/ts
5369/327-M1016 (holotype)	42.8	10.2	1.0	12.3	17.7	2.3	–	–	–
5369/327-M1017	34.4	8.3	0.9	13.0	20.1	3.4	3.1	5.9	25.7
5369/327-M1023	38.5	10.0	0.7	7.8	15.4	3.2	3.3	–	23.3

C o m p a r i s o n. This species is distinguished from *N. elegans* sp. nov., characterized by few thin rod-like spinules only occasionally bifurcating in the upper third of their length and numerous, very small conical spines on the surface of the inner sphere, by the presence of very numerous thin rod-like spinules that are frequently double bifurcated.

M a t e r i a l. Three specimens from the type locality.

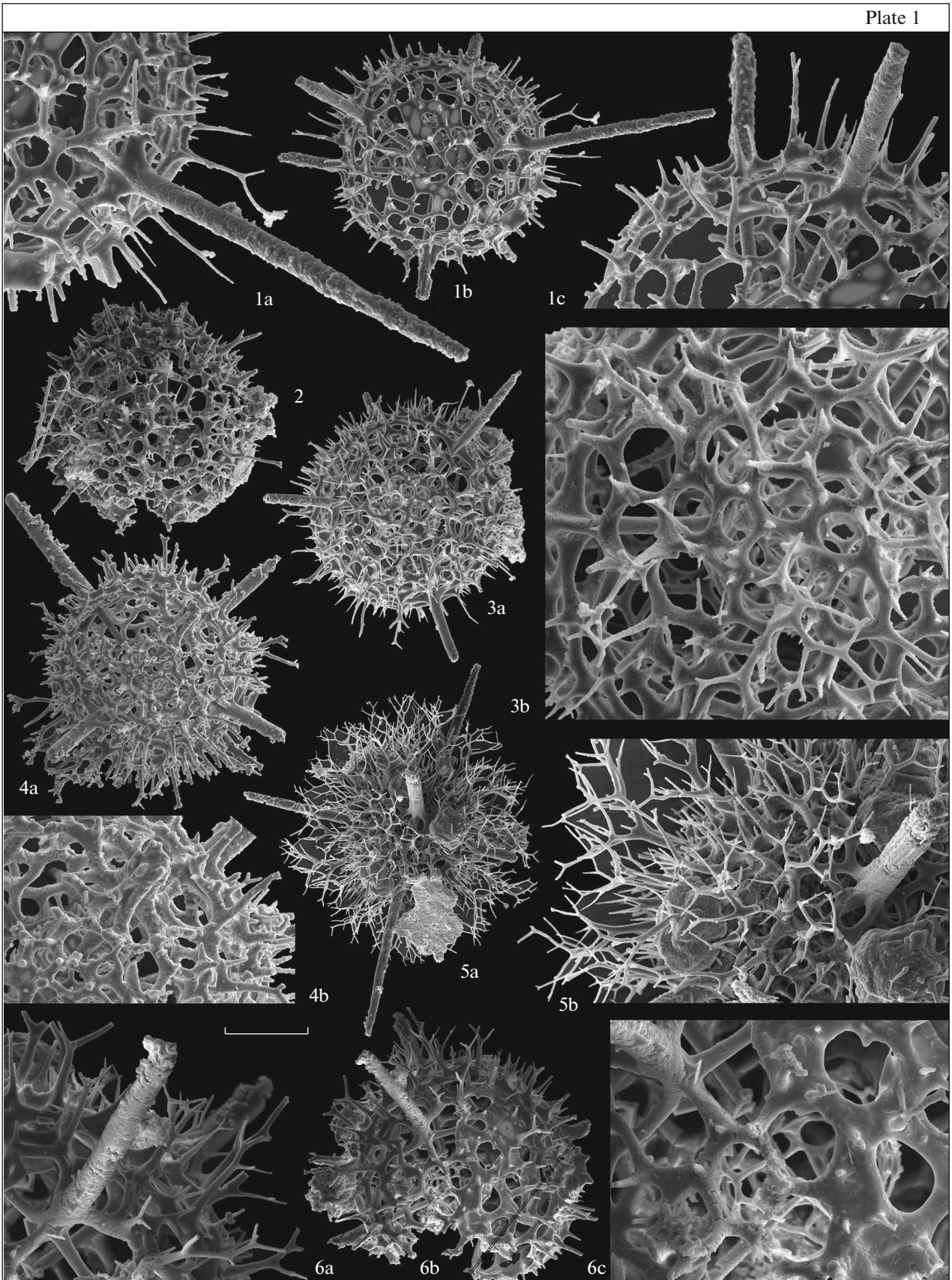
Natgorella elegans Afanasieva sp. nov.

Plate 1, figs. 1–3

E t y m o l o g y. From the Latin *elegans* (elegant).

H o l o t y p e. PIN, no. 5369/328-M1014; Volga–Ural Basin, Melekess Depression, Melekess P-1 Borehole (interval 1961.9–1963.9 m); Upper Devonian, Middle Famennian, Lebedyanian Regional Substage, Late *marginifera* and *trachytera* conodont zones.

D e s c r i p t i o n. The shell is spherical, very large ($D = 206\text{--}219\ \mu\text{m}$), consisting of two lattice spheres with six main rod-like spines and numerous spinules. The outer sphere is composed of thin entangled bars



($D/tw = 36.5-42.0$), forming very large ($D/dp = 7.5-10.3$) polygonal pores of approximately the same size. The inner sphere is large ($D/D_s = 2.8-3.2$), relatively thick-walled ($D_s/ts = 22.7-25.0$), and is 31.1–35.9% of the outer sphere diameter. It was formed by entangled bars surrounding very large ($D_s/ds = 5.7-6.8$) polygonal pores. The surface of the bars of the inner sphere show numerous small conical spinules from 2 to 5 μm long (Pl. 1, fig. 3b). Six main rod-like spines are continuations of the inner six-ray spicule: the spines are relatively short ($L/D = 0.5-0.9$) and thin ($L/wL = 10.2-13.4$). The surface of the bars of the outer sphere possesses two kinds of spinules: (1) very numerous long ($D/ht_1 = 12.2-13.7$) conical spinules; (2) few, but very long ($D/ht_2 = 3.8-4.3$) rod-like thin spinules, sometimes bifurcating in approximately the upper third of their length.

Dimensions in μm and ratios (Fig. 6):

Specimen no.	D	dp	tw	Ds	ds	ts	L	wL	ht ₁	ht ₂
5369/328-M1014 (holotype)	210	28	5	75	11	3	201	15	17	50
5369/328-M1020	219	23	5	68	12	3	—	11	18	58
5369/328-M1021	206	20	5	74	13	3	102	10	15	48

Specimen no.	D/tw	D/dp	L/D	L/wL	D/ht ₁	D/ht ₂	D/Ds	Ds/ds	Ds/ts
5369/328-M1014 (holotype)	42.0	7.5	0.9	13.4	12.4	4.2	2.8	6.8	25.0
5369/328-M1020	36.5	9.5	—	—	12.2	3.8	3.2	5.7	22.7
5369/327-M1021	41.2	10.3	0.5	10.2	13.7	4.3	2.8	5.7	24.7

Material. Three specimens from the type locality.

CLASS SPHAERELLARIA

Order Entactiniata

Family Entactiniidae Riedel, 1967

Subfamily Entactiniinae Riedel, 1967

Genus *Spinoalius* Afanasieva gen. nov.

Etymology. From the Latin *spina* (spine) and *alius* (other); masculine.

Type species. *Spinoalius melekessensis* sp. nov.

Diagnosis. Shell spherical, porous. Rays of the inner six-ray spicules continue into six main spines: one dominant four-bladed spine with a \clubsuit -like cross-section of the base and five rod-like spines with rounded sections of the base. Six petaloids are present near the base of beam-like spines (Pl. 2, figs. 1a, 1b).

Species composition. Type species.

Comparison. This genus differs from the most similar genus *Borisella* Afanasieva, 2000, with a porous skeleton possessing six main beam-like spines, in the presence of a single dominant four-bladed spine with a \clubsuit -like cross-section and five rod-like spines.

Spinoalius melekessensis Afanasieva sp. nov.

Plate 2, figs. 1–4

Etymology. After the locality in the Melekess Depression in the Volga–Ural Basin.

Holotype. PIN, no. 5369/329-M1027; Volga–Ural Basin, Melekess Depression, Melekess P-1 Borehole, interval 1973.9–1974.7 m; Upper Devonian, Middle Famennian, Lebedyanian Regional Substage, Late *marginifera* and *trachytera* conodont zones.

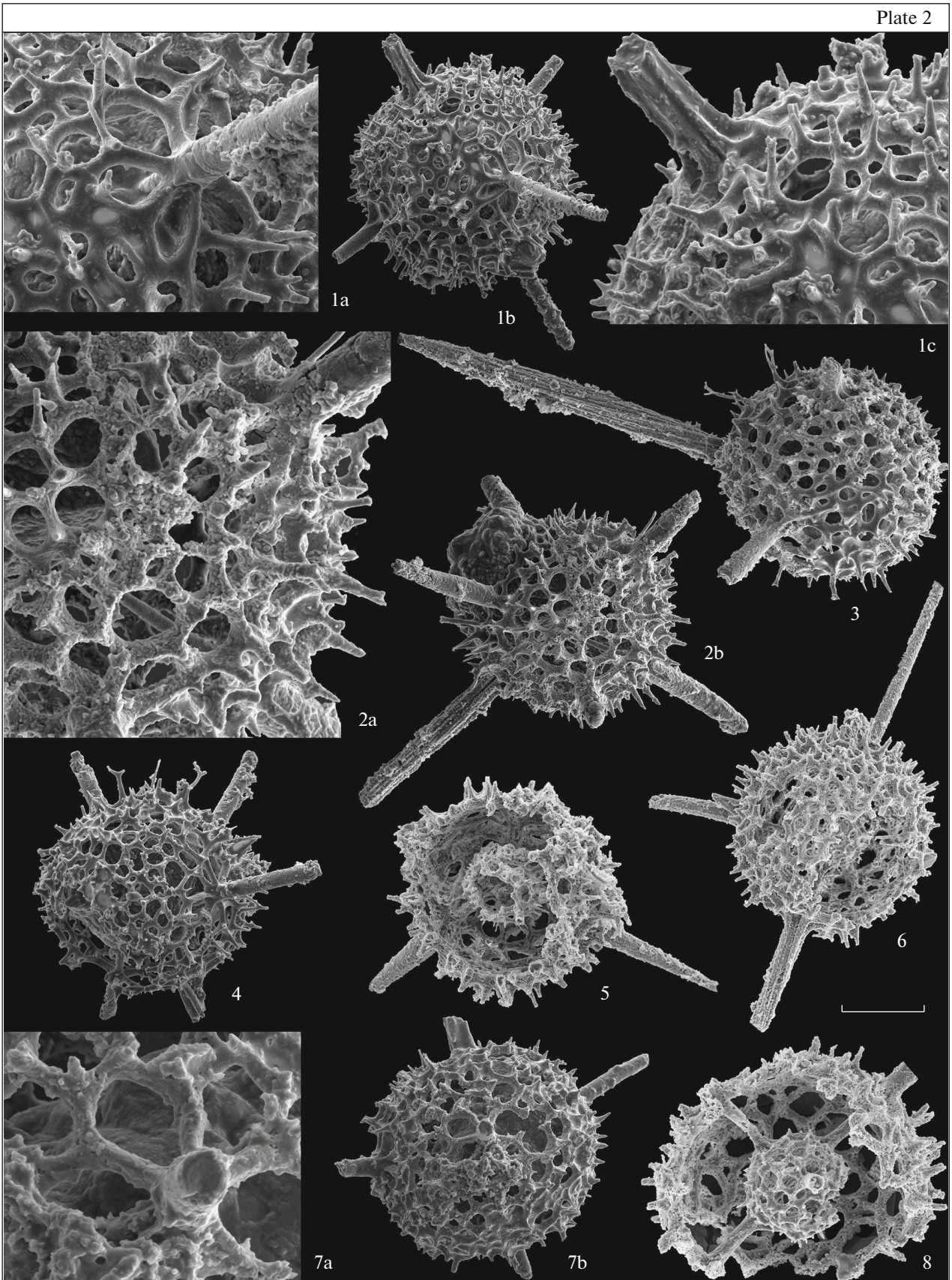
Description. The shell is spherical, very large ($D = 214-252 \mu\text{m}$), composed of one porous sphere. The outer sphere is thick-walled ($D/tw = 19.4-25.6$) with large, rounded pores approximately of the same diameter ($D/dp = 12.2-14.4$). The inter-pore wall are acute and relatively narrow ($dp/wp = 4.5-5.7$). The rays of the inner six-ray spicule continue as six main spines: (1) one dominant four-bladed spine with a \clubsuit -like cross-section, relatively long ($L_1/D = 0.7-1.1$) and thin ($L_1/wL_1 = 7.3-11.7$); (2) five rod-like spines with a rounded cross-section, the spines are short ($L_2/D = 0.4-0.6$) and relatively thick ($L_2/wL_2 = 3.9-6.3$). Six petal-like pores are present near the base of the beam-like spines (Pl. 2, figs. 1a, 1b). The inter-pore wall possess ornamentation in the shape of very high acute spinules, sometimes bifurcating distally (Pl. 2, figs. 3, 4).

Explanation of Plate 1

Upper Devonian, Middle Famennian, Lebedyanian Regional Substage, Late *marginifera* and *trachytera* conodont zones; Volga–Ural Basin, Melekess Depression, Melekess P-1 Borehole: Sample 7336, interval 1961.9–1963.9 m (1–3, 5, 6) and Sample 7398, interval 1973.9–1974.7 m (4).

Figs. 1–3. *Natgorella elegans* sp. nov.: (1) holotype PIN, no. 5369/328-M1014: (1a) fragment (scale bar = 42 μm), (1b) general view (scale bar = 83 μm), (1c) fragment (scale bar = 41 μm); (2) specimen PIN, no. 5369/328-M1020 (scale bar = 87 μm); (3) specimen PIN, no. 5369/328-M1021: (3a) general view (scale bar = 86 μm), (3b) fragment (scale bar = 29 μm).

Figs. 4–6. *N. hirsuta* sp. nov.: (4) specimen PIN, no. 5369/327-M1023: (4a) general view (scale bar = 98 μm), (4b) fragment (scale bar = 49 μm); (5) holotype PIN, no. 5369/327-M1016: (5a) general view (scale bar = 136 μm), (5b) fragment (scale bar = 59 μm); (6) specimen PIN, no. 5369/327-M1017: (6a) fragment (scale bar = 44 μm), (6b) general view (scale bar = 90 μm), (6c) fragment (scale bar = 33 μm).



Explanation of Plate 2

Upper Devonian: (1–4, 7) Middle Famennian, Lebedyanian Regional Substage, Late *marginifera* and *trachytera* conodont zones; Volga–Ural Basin, Melekess Depression, Melekess P-1 Borehole, Sample 7398 (interval 1973.9–1974.7 m); (5, 6, 8) Lower Famennian Substage, *Palmatolepis triangularis* and *P. crepida* conodont zones; Polar Urals, Lemva Zone, Palnik-Yu River: Sample 101/599 (5), Sample 101/594 (6, 8).

Figs. 1–4. *Spinoalius melekessensis* sp. nov.: (1) specimen PIN, no. 5369/329-M1025: (1a) fragment (scale bar = 40 μm), (1b) general view (scale bar = 98 μm), (1c) fragment (scale bar = 40 μm); (2) specimen PIN, no. 5369/329-M1022: (2a) fragment (scale bar = 38 μm), (2b) general view (scale bar = 96 μm); (3) holotype PIN, no. 5369/329-M1027, general view (scale bar = 88 μm); (4) specimen PIN, no. 5369/329-M1026 (scale bar = 88 μm).

Figs. 5–8. *Diversiacus zuraevi* (Afanasieva and Amon, 2011): (5) specimen PIN, no. 5312/160-15958 (scale bar = 79 μm); (6) holotype PIN, no. 5312/160-15718 (scale bar = 98 μm); (7) specimen PIN, no. 5369/160-M1024 (scale bar = 100 μm); (7a) fragment (scale bar = 27 μm), (7b) general view (scale bar = 92 μm); (8) specimen PIN, no. 5312/160-15732 (scale bar = 67 μm).

Dimensions in mm and ratios (Fig. 6):

Specimen no.	D	dp	wp	tw	L ₁	wL ₁	L ₂	wL ₂	ht
5369/329-M1027 (holotype)	230	16	3	9	363	31	86	22	54
5369/329-M1022	231	19	4	11	220	30	136	28	52
5369/329-M1025	252	18	4	13	–	32	125	20	30
5369/329-M1026	214	17	3	9	–	27	102	18	40

Specimen no.	D/dp	D/tw	dp/wp	L ₁ /D	L ₁ /wL ₁	L ₂ /D	L ₂ /wL ₂	D/ht
5369/329-M1027 (holotype)	14.4	25.6	5.3	1.6	11.7	0.4	3.9	4.3
5369/329-M1022	12.2	21.0	4.8	1.0	7.3	0.6	4.9	4.4
5369/329-M1025	14.0	19.4	4.5	–	–	0.5	6.3	8.4
5369/329-M1026	12.6	23.8	5.7	–	–	0.5	5.7	5.4

Remarks. This species differs from species of the most similar genus *Borisella* Afanasieva, 2000, which have six rod-like spines, in the presence of one dominant four-bladed spine with a \blacklozenge -like cross-section and five rod-like spines.

Material. Four specimens from the type locality.

Subfamily Bientactinosphaerinae Afanasieva, 1999**Genus *Diversiacus* Afanasieva gen. nov.**

Etymology. From the Latin *diversus* (diverse) and *acus* (needle); masculine.

Type species. *Bientactinosphaera zuraevi* Afanasieva and Amon, 2011 from the Upper Devonian, Lower Famennian, *triangularis* and *crepida* conodont zones of the Polar Urals, Lebedyanian Regional Substage, Late *marginifera* and *trachytera* conodont zones; Volga–Ural Basin.

Diagnosis. Shell spherical with two porous spheres. Inner frame in the shape of a six-rayed spicule (Pl. 2, fig. 8), the rays of which continue into six main spines. One main spine is three-faceted, with a \blacktriangleleft -like

cross section and lobe-like cross-section of each blade. Five main spines rod-like with rounded cross-section. Six petal-shaped pores near base of rod-like spines (pl. 2, fig. 7).

Species composition. Type species.

Comparison. This genus is distinguished from *Bientactinosphaera* Afanasieva, 2000, which has six main three-bladed spines, by the presence of one main three-bladed spine and five main rod-like spines.

Remarks. Originally, the species *D. zuraevi* was described in the genus *Bientactinosphaera*, based on the presence of two porous spheres (Afanasieva and Amon, 2011). However, skeleton morphology allowed the recognition of a new genus *Diversiacus* gen. nov., with *D. zuraevi* as the type species. In addition, a convergent similarity is established for the external morphology of the skeletons of the new genera *Diversiacus* and *Spinoalius* (Fig. 5). The skeletons of these genera have an outer porous sphere with five main rod-like spines and a one faceted main spine (four-bladed or three-bladed) (pl. 2). Note that the skeletons of the genus *Spinoalius* have (Fig. 5a; Pl. 2, fig. 2a) only one outer sphere (subfamily Entactiniinae), whereas *Diversiacus* differs (Fig. 5b; Pl. 2, figs. 5, 6, 8) in having two spheres: outer and inner (subfamily Bientactinosphaerinae).

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