## **RESEARCH PAPER**





# **Estimating water erosion using RUSLE, GIS and remote sensing in Wadi‑Qandeel river basin, Lattakia, Syria**

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#### **Abstract**

Soil erosion is the major prominent ecological risk threatening agricultural sustainability in the coastal region of western Syria. The ongoing war conditions in Syria has led to a lack of feld data and measurements related to the spatial evolution of soil loss. Estimating the spatial distribution of potential soil loss is a fundamental procedure in applying the soil conservations measures within the river catchments. The current paper goals to conduct a comprehensive assessment of soil loss risk utilizing revised universal soil loss equation (RUSLE) and remote sensing (RS) data in geographic information system (GIS) environment across the entire Wadi-Qandeel river basin. Results indicate that the annual rate of soil loss in the study area was 93.02 t ha<sup>-1</sup> ya<sup>-1</sup> with a spatial average reaching 58.22 t ha<sup>-1</sup> ya<sup>-1</sup>. Additionally, the soil loss risk map was generated with classification into five susceptible zones: very low (56.44%), low (24.69%), moderate (20.80%), high (2.98%), and very high (2.22%). The present assessment showed a reliable approach to soil erosion rates and categorization of erosion-sensitive zones within the study area. These outcomes can be relied upon to create mitigation procedures for maintaining zones with high and very high soil loss susceptibility under the ongoing war conditions in Syria.

**Keywords** Soil loss · RUSLE · GIS · RS · Conservation · Syria

# **Introduction**

Soil is an essential and pivotal non-renewable resource that presents a broad group of environmental services and goods, especially for human sustainability (Ferrara et al. [2015](#page-8-0); Pal [2016;](#page-8-1) Brevik et al. [2017](#page-7-0); Testi et al. [2010](#page-9-0); Riccardi et al. [2020;](#page-8-2) Abd El-Ghani et al. [2012;](#page-7-1) Hatefard et al. [2021](#page-8-3)). However, a growing body of literature has indicated that soil loss is responsible for about 85% of the degradation of the global agro-ecosystems, and therefore a reduction in nourishment productivity by 17% (Wijesundara et al. [2018;](#page-9-1) Nyesheja et al. [2019;](#page-8-4) Bahir et al. [2021;](#page-7-2) Mountassir et al. [2021a,](#page-7-3) [b](#page-8-5)). Soil erosion by water (SEW) is one of the most severe environmental problems afecting life sustainability and welfare

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worldwide (Zika and Erb [2009\)](#page-9-2). SEW is dynamically generated as a result of spatial integration between physical and human factors, produced a total reduction in the quality of soil health and water resources, therefore a decrease in ecosystem quality and productivity (García-Ruiz [2010;](#page-8-6) Göl [2017;](#page-8-7) Cutini et al. [2015](#page-7-4); Balasubramani et al. [2015](#page-7-5); Mokhtar et al. [2021](#page-8-8); Tang et al. [2015;](#page-9-3) El Mountassir et al. [2021a](#page-7-3), [b](#page-8-5)).

In this regard, Syrian soils are characterized by high susceptibility of water soil erosion risk due to lithology, high rainfall intensities, rough relief, topsoil fertility, degraded vegetation, edaphic, and accelerated human activity, especially in the coastal region of Syria (Mohammed et al. [2020a;](#page-8-9) Abdo [2021\)](#page-7-6). Further, about 18% of the agricultural land in Syria is prone to soil loss hazard that exceeds 100 t ha−1 year−1 in many western mountainous areas (ACSAD [2007;](#page-7-7) Husein and Kalkha [2019](#page-8-10); Abdo [2019\)](#page-7-8). Particularly, SEW in Wadi-Qandeel river basin is the basic environmental hazard that threatens farming, feeding, and secure sustainability. The intensive pattern of rainfall, high runoff peaks, flood, rugged terrain, steep slopes, shallow soil profiles, and sparse vegetation are the major physical criteria that motivate the water erosion in Wadi-Qandeel river basin. At the same time, SEW can be accelerated by the growth of

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anthropogenic activities such as deforestation, overgrazing, urbanization, landuse/landcover change (LULC), intensive cultivation on steep slopes, excessive soil plowing, land abandonment, poor maintenance procedures, military infrastructure, and armed conficts (Emadodin and Bork [2012](#page-7-9); Jafari and Bakhshandehmehr [2016](#page-8-11); Nabiollahi et al. [2017](#page-8-12); Abdo [2018;](#page-7-10) Mohammed et al. [2021](#page-8-13)).

The matter of SEW simulating was the pivot of many spatial evaluations in the coastal area of Syria. Barakat et al. [\(2014](#page-7-11)) and Husein and Kalkha [\(2019\)](#page-8-10) utilized the Coordination of Information on the Environment (CORINE) approach to map the risk of SWE across many coastal watersheds. Mohammed et al. [\(2016](#page-8-14)) predicted SEW by utilizing the Water Erosion Prediction Project (WEPP) model for Lattakia governorate. The Revised Universal Soil Loss Equation (RUSLE) method was applied in the GIS environment to evaluate the SEW in many basins in Syrian coastal region (Abdo and Salloum [2017a,](#page-7-12) [b;](#page-7-13) Almohamad [2020\)](#page-7-14).

Assessing the annual spatial distribution of SEW rates is an urgent need for maintaining soil and water resources at the basins scales. However, utilizing the experimental methods is the most global manners of SEW modeling, particularly in regions with limited information and data. In other words, the integration of remote sensing (RS) data, geographic information system (GIS) environment and RUSLE model is a useful, reliable and accurate method in producing the spatial distribution of SEW map. For the prior matter discussed, the current research will explore the quantities and spatial distribution of SEW in Wadi-Qandeel river basin by using the RS data in calculating RUSLE factors in GIS software, therefore suggesting constructive spatial preservation strategies with suitable applications.

## **Materials and methods**

#### **Study area**

Wadi-Qandeel watershed is one of the coastal river watersheds in Lattakia governorate in the west of Syria (Fig. [1](#page-1-0)). The study basin covers about  $152.23 \text{ km}^2$ , and the highest elevation of 809 m above mean sea level (m a.m.s.l.) (elevation range of 0–809 m a.m.s.l.). This basin bordered Al-Kabeer alshamali river basin to the east, the Mediterranean



<span id="page-1-0"></span>**Fig. 1** Site of Wadi-Qandeel river basin

to the west, Al-arab river basin to the south, Al-Qshish river basin to the north. The study area includes various geological formations, starting from the Cretaceous to the Quaternary. Cretaceous formations composed of Albian, Lower Cenomanian, and Upper Cenomanian with lithological structures of limestone, dolomite, marls, ophiolites, limy marl, and sandy limestone. The basin mainly subjects to the Mediterranean climate pattern: mild and rainy winter and long, dry, and hot summer (Khallouf et al. [2021](#page-8-15)). The average annual precipitation in the catchment varies from 847 to 1301 mm (Fig. [2](#page-2-0)), and most rainfall is concentrated in the winter months (68%). Further, the annual mean in summer temperature is 26° and in winter 15°. The warmest month is August  $(28.6^\circ)$  and the coldest is January (13.8°). High relative humidity prevails throughout the year due to the efects of the Mediterranean water mass. The annual average of humidity is 72%. Study area is characterized by the density of the streams network, mainly refecting the spatial integration between geological, tectonic and climatic characteristics. Wadi-Qandeel drainage network originates at Ballouran village in the Al-shrashir area at an altitude of 802 m above sea level, and flows westward down to its concentration point with the Mediterranean (Abdo [2020\)](#page-7-15). Agriculture and tourism are the basic economic activities of the population. Olives, citrus and agricultural crops are among the most important crops in the basin.



<span id="page-2-0"></span>**Fig. 2** Spatial distribution of rainfall values

#### **RUSLE model parameters**

The universal equation of soil loss RUSLE (Wischmeier and Smith [1978](#page-9-4); Renard et al. [1997\)](#page-8-16) has been commonly used to evaluate the spatial dimension of soil erosion rates in order to generate the conservation goals, with a reasonable scale of validity (Balasubramani et al. [2015](#page-7-5); Djoukbala et al. [2019\)](#page-7-16). RUSLE is a well-established soil erosion empirical model and most of its parameters have undergone profound modifcations since its introduction. RUSLE composes of fve geo-parameters that represent the following inputs: rainfall erosivitity (*R*), soil susceptibility to erosion based on its physical–chemical properties (*K*), terrain (*LS*), vegetation(*C*), and maintenance (*P*), respectively. Meanwhile, RUSLE model has been applied in areas with diferent cases worldwide generally and in the Mediterranean basin environment particularly: Kef et al. [\(2012\)](#page-8-17), in Tunisia; Demirci and Karaburun ([2012](#page-7-17)), in Turkey; Chadli [\(2016\)](#page-7-18), in Morocco; Fagnano et al. ([2012](#page-7-19)), in Italy; in Algeria; Benkadja et al. [\(2015\)](#page-7-20), in Algeria, etc. The average annual soil erosion per unit area is given by the following equation (Eq. [1](#page-2-1)) of RUSLE (Wischmeier and Smith [1978\)](#page-9-4):

<span id="page-2-1"></span>
$$
A = R \times K \times LS \times C \times P, \tag{1}
$$

where *A* is the average annual soil erosion (t ha<sup>-1</sup> year<sup>-1</sup>). *R* is the rainfall erosivity, *K* is the soil erodibility, *LS* is the hill slope length and steepness, *C* is the vegetation factor, and *P* is the support practice. In order to standardize pixel resolution DEM and Landsat imageries resolution, all inputs and outputs for the calculation of erosion risks is in 30-pixel resolution for each sub-factors of RUSLE model. In this regard, RUSLE model was used to estimate SEW as a result of achieving accurate and reliable spatial outcomes for the river basins in the eastern Mediterranean region, the fexibility of its application in areas with limited data, and the possibility of relying on its outputs in formulating strategies for water and soil resources mitigation and conservation measures.

#### **Rainfall erosivity factor (R)**

The rainfall erosivity (R) factor considers the infuence of rainfall kinetic energy and generated runoff on the erosion (Wischmeier and Smith [1978](#page-9-4); Demirci and Karaburun [2012](#page-7-17); Salloum and Abdo [2015](#page-9-5); Djoukbala et al. [2018](#page-7-21); Abdo and Hassan [2018](#page-7-22)). R factor, moreover, is importantly afected by the patterns, spatial–temporal distribution, intensities, momentums, and kinetic energy, raindrops size (Farhan et al. [2013](#page-8-18); Carollo et al. [2018](#page-7-23); Serio et al. [2019](#page-9-6)). According to available monthly rainfall data



(1989–2019) of fve climatic stations established in and around the study basin the map of *R* value was prepared by using Eq. ([2\)](#page-3-0) developed by Wischmeier and Smith [\(1978](#page-9-4)):

$$
R = \sum_{1}^{12} 1.735 \times 10^{(1.5 \log_{10} \left( \frac{P_i}{P} \right) - 0.08188)},\tag{2}
$$

where *R* is a rainfall erositivity factor (MJ mm ha<sup>-1</sup> h<sup>-1</sup> per year); *Pi* is monthly rainfall (mm); *P* is an annual rainfall (mm). Figure [3](#page-3-1) illustrates the spatial distribution of precipitation erosivity. On value, *R* factor values ranged from 621.78 to 1234.55 MJ mm ha−1 h−1 year−1.

#### **Soil erodibility factor (K)**

K factor describes the strong linking of topsoil particles against the rainstorms events, which is commonly acquired by evaluating the physical–chemical topsoil properties of a specifc plot (Wischmeier et al. [1971](#page-9-7); Das et al. [2018](#page-7-24); Saïdi et al. [2014](#page-9-8)). *K* factor is an assessment of the susceptibility of topsoil to detachment and transport by rainfall and runof. It refects the soil erosion rate caused by rainfall erosivity factor (*R*) in each point of the study area (Koirala et al. [2019](#page-8-19)). In the present assessment, soil features were evaluated based on the analysis of soil samples conducted by the National Center for Agricultural Research in Al-Hanadi region. *K* factor map was delineated by utilizing the following equation presented by Wischmeier and Smith [\(1978\)](#page-9-4), Renard et al. [\(1997\)](#page-8-16) and Panagopoulos and Ferreira ([2010\)](#page-8-20):

$$
K = \frac{2.1 \times 10^{-4} (12 - OM) \times M^{1.14} + 3.25(s - 2) + 2.5(p - 3)}{100}
$$
\n(3)

<span id="page-3-0"></span>where *OM* is the organic matter  $(\%)$ , *s* is soil structure class, *p* is permeability class, and *M* is aggregated variable derived from the granular soil texture:  $M = (\%Msilt) \times (\%sit + \%sand)$ , and the modified silt (Msilt) is a percentage of grain size between 0.002 and 0.1 mm. Figure [4](#page-3-2) indicates the spatial distribution of the *K* values, which ranged from 0.021 to 0.032 ton ha  $MJ^{-1}$  mm<sup>-1</sup>.

#### **Slope length and steepness factor (LS)**

Slope Length and Steepness Factors (LS) approach the impact of the relief on the acceleration of soil erosion (Lu et al. [2004\)](#page-8-21). LS factor was generated from two sub-factors: a slope degree parameter (*S*) and a slope–length parameter (*L*); which are extracted from the Digital Elevation Model (DEM) (Hickey [2000](#page-8-22); Boggs et al. [2001\)](#page-7-25). In this context, the LS factor is the most essential causative parameter of overland flow that considers the major reason to soil loss. The relation of soil erosion to slope degrees of hill and mountains area is infuenced by the fora density and soil properties (Koirala et al. [2019\)](#page-8-19). The terrain of the study basin is featured by steep slopes which reach more than 45 degrees, as Fig. [5](#page-4-0) shows. By using the digital elevation model (DEM) with 30 m resolution (ASTER GDEM Validation Team [2009](#page-7-26)) LS factor map was created according to Eq. ([4\)](#page-4-1)



**Fig. 3** Spatial distribution of *R* factor values **Fig. 4** Spatial distribution of *K* factor values

<span id="page-3-2"></span><span id="page-3-1"></span>



$$
LS = (FlowAccumulation \times \frac{CellSize}{22.13})^{0.5} \times (\frac{sinslope}{0.0896})^{1.3}
$$
(4)

where FlowAccumulation is the grid layer of flow accumulation expressed as the number of grid cells, and CellSize is the length of a cell side. *LS* values of the study area are in the range of 0–37.58 as illustrated Fig. [6.](#page-4-2)

#### **Vegetation factor (C)**

Vegetation cover is a complex criterion in soil loss control, because of dissipating the kinetic energy of raindrops, delaying the surface runoff, and enhancing the infiltration capacity (Hu et al. [2015;](#page-8-23) Salloum and Abdo [2016](#page-9-9); Sujatha and Sridhar [2018;](#page-9-10) Acar et al. [2014](#page-7-27)). In this context, *C* factor refects the vegetation condition that can be rapidly varied than other RUSLE parameters (Beskow et al. [2009](#page-7-28)). Landuse/landcover and Normal Diference Vegetation Index (NDVI) are two methods that soil erosion modelling scholars use in calculating *C* values. In the current assessment, the *C* factor values were calculated using the NDVI index which is given by Eq. [5](#page-4-3)

$$
NDVI = (NIR - RED)/(NIR + RED)
$$
 (5)

where NIR is the near-infrared band (band 4, 0.76–0.90  $\mu$ m), and RED is the red band (band 3, 0.63–0.69 µm). Landsat 8 OLI image taken in March 2020 was considered in calculating NDVI values which ranged between−0.3 and 0.83



<span id="page-4-0"></span>**Fig. 5** Spatial distribution of slope values **Fig. 6** Spatial distribution of *LS* factor values

<span id="page-4-2"></span><span id="page-4-1"></span>

<span id="page-4-4"></span><span id="page-4-3"></span>**Fig. 7** Spatial distribution of NDVI values

(Fig. [7](#page-4-4)). The values of *C* factor was calculated using the Eq. [6](#page-5-0)

$$
C_{\text{factor}} = exp\left[-\alpha \frac{NDVI}{(\beta - NDVI)}\right] \tag{6}
$$

where  $\alpha$  and  $\beta$  parameters determine the shape of the NDVI curve. Reasonable results are produced using values of  $\alpha$  = 2 and  $\beta$  = 1. *C* factor values ranged between 0.02 and 1 (Fig. [8](#page-5-1)).

#### **Conservation support practice factor (P)**

Maintenance practice factor  $(P)$  is the ratio of soil loss after a selective support exercise to the corresponding soil loss after up and down farming (Samanta et al [2016](#page-9-11)). However, P factor basically infuences soil erosion by modifying the streaming pattern, degree or orientation of overland flow, and by decreasing the runoff potentials (Ozsoy and Aksoy [2015](#page-8-24)). For cultivated land, the conservation practices included contouring, terracing, strip cropping, and subsurface drainage (Renard et al. [1997](#page-8-16)). *P* factor values range from 0 to 1, the value 0 suggests good conservation support practices and the value 1 suggests poor conservation support practices (Wischmeier and Smith [1978;](#page-9-4) Das et al. [2018](#page-7-24)). Field monitoring indicates the loss of prevention support procedures in the study area. Consequently, *P* factor value for the entire study basin is 1 as proposed by Wischmeier and Smith [\(1978](#page-9-4)).



<span id="page-5-1"></span>

## **Result and discussion**

<span id="page-5-0"></span>Based on the initial information and data entered into the GIS software (ArcGIS 10.2.3), four thematic raster maps were precisely generated representing the spatial factors of RUSLE. *R*, *K*, *LS*, and *C* raster parameters were spatially multiplied in order to map the spatial distribution of potential soil loss per hectare per year at cell level as Fig. [9](#page-5-2) illustrates. In the context of current results, the annual rate of soil loss in the study area ranged from 0 to 93.02 ton ha<sup>-1</sup> year<sup>-1</sup>, with a spatial average reached 58.22 ton ha<sup>-1</sup> year<sup>-1</sup>. Utilizing the *Natural Breaks* method, the output soil erosion map was classifed into fve risk classes: very low (56.44%), low (24.69%), moderate (20.80%), high (2.98%), and very high  $(2.22\%)$  as Table [1](#page-5-3) shows.

The fnding soil loss rate is spatially consistent with the assessments provided by scholars in river basins



<span id="page-5-2"></span>**Fig. 9** Spatial distribution of soil erosion values in Wadi-Qandeel river basin

<span id="page-5-3"></span>**Table 1** Classifcation of soil erosion in study basin

Soil erosion classes	Rate of soil loss class in ton $ha^{-1}$ year <sup>-1</sup>	Area $(km^2)$	%
Very low	< 10	85.92	56.44
Low	$10 - 20$	37.58	24.69
Moderate	$20 - 30$	20.80	13.66
High	$30 - 40$	4.54	2.98
Very high	$40 - 93.02$	3.38	2.22

environment in the eastern Mediterranean as Table. [2](#page-6-0) illustrates. In addition, the outputs of prior literature give the resulting soil erosion rate sufficient validity to be utilized in suggesting practices for spatial conservation of areas with boundaries and critical soil erosion. Meanwhile, Nearing et al. ([1990](#page-8-25)), Irvem et al. ([2007](#page-8-26)), Trabucchi et al. ([2012](#page-9-12)), and Farhan and Nawaiseh [\(2015](#page-8-27)) suggested that 2 to 12 ton  $ha^{-1}$  year<sup>-1</sup> is the acceptable soil loss tolerances limits for the aims of agricultural and economic sustainability in the Mediterranean environment. Further, Ibrahiem ([1986](#page-8-28)) and Kbibo and Nesafi ([1997](#page-8-29)) reported that the tolerable limit of soil loss ranged between 1 to 2.5 t h<sup>-1</sup> year<sup>-1</sup> for the coastal region of Syria, owing to the many geo-factors that infuences the soil formation. Therefore, in light of these limits, it can be confrmed that most of the study basin lands need integrated spatial management of soil loss.

The spatial distributions of high and very high soil loss hazards are mostly concentrated in the northeast and central regions slopes of the study area. Further, these zones are susceptible to the high risk of soil loss as a result of the spatial integration between the kinetic energy of rainfall intensities, runoff and steep slopes. The field checking disclosed that classes with high and very high values of soil erosion are sloping lands. Importantly, vegetation did not reduce the soil loss quantities in those areas as generated erosion map indicated. This result can be demonstrated by the great spatial competition between slope and vegetation factors in favour of the slope factor. Meanwhile, the slope degrees in these areas reaches more than 45 degrees as Fig. [5](#page-4-0) shown. These fndings are consistent with many studies that have pointed out the infuence of slopes on topsoil mobilization (Thomas et al. [2018](#page-9-13); Kayet et al. [2018](#page-8-30)). Therefore, it can be reported that the slope is a pivotal factor in the acceleration of soil loss in the study area.

The current outputs, that targeted the spatial susceptibility of soil loss in the study area, are spatial predictions, and thereby are still questionable. Meanwhile, the absence of adequate data, information and feld measurements across the basin due to diferent afecting reasons, particularly the consequences of the ongoing war, a RUSLE model presented constructive statistical-spatial estimates of potential soil loss risk. These delineated estimates can be useful for decisionmakers in generating strategies for maintaining soils and mitigating areas with high and very high sensitivity to erosion. In this context, among the most important measures that can be proposed in the framework of the erosion mitigation process: (1) adjusting the slopes in areas of high and very high sensitivity by building terraces, (2) enhancing wild vegetation coverage, especially in areas with frequent fres and indiscriminate cutting, (3) investigating the root system efficiency (RSE) of agricultural crops for the proportionality and severity of cultivated slope, (4) developing agricultural techniques, especially ploughing process. The RUSLE application could also be expanded at the regional and national levels as part of the spatial management plans for river basins.

# **Conclusions**

Soil loss represents the most geo-ecological problem threats to food security in the coastal region of Syria. In light of the paucity of spatial data related to soil loss, current research presented a reasonable assessment of the spatial distribution of soil erosion severity in one of the coastal river basins most prone to soil erosion. The goal of this study was fulflled by feeding the GIS environment with multi-source information and data, especially RS data, in the calculation of RUSLE factors. In regard to the present fndings, the generated soil erosion map shows a maximum rate of erosion 93.02 t h<sup>-1</sup> year<sup>-1</sup>. Moreover, the soil erosion amount exceeded the tolerable threshold of soil erosion for the coastal region of Syria (1 to 2.5 t h<sup>-1</sup> year<sup>-1</sup>). These predictions are closely related to the estimates computed in Syria and the Mediterranean countries using the RUSLE model. Also, it was concluded the high infuence of the slope factor in enhancing soil erosion. The spatial integration process between GIS and RS data presented a promising tool for assessing annual

<span id="page-6-0"></span>**Table 2** Some erosion assessments in diferent parts of Mediterranean areas using RUSLE model





rates of soil erosion in the study basin, especially in light of the ongoing war conditions in Syria. Therefore, the results of this study are necessary to manage the risk of soil erosion by taking mitigation and conservation measures. The application of this study's approach at the national level also provides an objective solution for managing areas of the high and very high soil erosion risk, especially in this period which Syria is living the consequences of the ongoing war since 2011.

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**Data availability** The spatial data used to support the fndings of this study are available from the corresponding author upon request.

## **Declarations**

**Conflict of interest** The author declares that they have no confict of interest.

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