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Estimate the amount of ground water recharge in hard formations, case study: Mashhad, Iran

Hossein Alem¹ · Akbar Esmaeilzadeh Soudejani¹ · Mahmood Fallahi¹

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Abstract

Discovering underground water resources remains a vital and inevitable method for optimizing consumption and also maintaining existing resources, not only in the world in general but also in places like Mashhad in particular which faces increasing need for water. Therefore, the present paper aims to identify potential locations for developing karst formations by calculating infiltration percentage of formations. APLIS method, based on GIS analysis, was used to do analytical hierarchical process. Experimental variables assumed to affect infiltration and storage of rain water include altitude, slope, lithology, soil, preferred infiltration sites, and fracture. Information layers for all the above variables were created, and their proportional weight and combinations were specified. Finally, recharge rates of aquifers were calculated and presented in form of maps which also showed the extent of karst formations for various regions. Results let us see that location-specific average of recharge rate for Mashhad formations. The ability to map the water resources management and determine the water balance components, including estimating the amount of groundwater resource penetration, and the spatial identification of karst sources compared to other existing methods, is higher. Therefore, the map possesses predictive strength for potential regions of development of karstification. Given the recharge rates for various sites and specified karstic regions, we can anticipate a more efficient management with regard to maintaining water resources and managing consumption. Such management would be more economical and less time-consuming.

Keyword Karstification · Recharge · APLIS · GIS · Mashhad

Introduction

Growth of population and unbridled expansion of cities, together with global warming, reducing running water resources, boundless biological pollution, diminishing phreatic zones, have drawn attentions of experts and managements to karstic formations and formations waters. These karstic formations cover 11 percent of Iran, a point of much importance in studying water resources. Studying karstic resources follows the management and maintenance goals of

 Hossein Alem Hosseinalem22@gmail.com
 Akbar Esmaeilzadeh Soudejani akbar.esmaeilzadeh@gmail.com
 Mahmood Fallahi mfallahi37@gmail.com

¹ Department of Earth Sciences, College of Sciences, Shiraz University, Shiraz, Iran these resources. Therefore, identifying these resources is an ineluctable necessity for purposes of consumption management (Aguilera et al. 2009; Allocca et al. 2014). Considering the importance of volume and process of recharge in development of karst, evaluation of recharge was the primary way for identifying potential locations (Allocca et al. 2014).

The recharge of karst aquifers in terms of quantity and spatial distribution depends on various factors such as climate, vegetation. The choice of a suitable method for assessing the amount of water infiltrated into karst areas is debated among researchers. Multi-parameter methods are being developed using the geographic information system tool. Groundwater recharge is an issue that has been systematically raised in various journals, especially since the mid-1980s (Healy 2010). The recharge of karst aquifers in southern Spain has been evaluated by the APLIS method (Andreo et al. 2008). In 2009, groundwater potential in northern Jordan was based on layers of altitude, soil, fault density, geomorphology, and geology using the AHP method



and divided the area into 5 classes "very bad, bad, average, good, and very good" (Awawdeh et al. 2014). In the same year, aerial photographs, geological maps, geomorphology, soil hydrogeological groups, land use and vegetation, and drainage network maps were used to determine the potential of groundwater resources in a region in India and the area was divided into three areas: good, medium, and poor (Nagarajan and Singh 2009). In 2010, a multi-criteria decisionmaking and GIS method was used to determine the appropriate groundwater supply areas in Tunisia (Chenini et al. 2010). In Sri Lanka, the method of geographic information system was used to determine the potential of groundwater nutrition (Senanayake et al. 2016). In the same year, studies were conducted to determine the potential of groundwater in a region of India using multi-criteria decision-making MCDM and AHP method (Kumar et al. 2016). Studies show that methods based on geographic information systems are suitable for producing groundwater potential maps, especially in areas where we lack data (Russo et al. 2015). The results of all these studies have been published in reputable scientific journals.

In geological formations, a usual measuring of recharge is possible by time series in which level of underground water is measured (Bouraoui et al. 1998); heterogeneous karstic rocks, however, make this method unreliable in karstic regions, and therefore, potential locations for development of karst formations cannot be specified (Burke 1995). In measuring recharge rate, conventional methods (evaporation and transpiration, natural isotopic trackers, chemical trackers, artificial trackers, dividing precipitation by time) or numerical methods encounter limitations and problems such as lack of accurate periodic data (Conrad et al. 2004). The key to this problem lies in using methods based on GIS. This method uses variables such as geological characteristics, altitude, slope, lithology, and fracture patterns and is often employed in explaining spatio-temporal distribution of karst recharge (Bakalowicz 2005; Carrasco et al. 2014). Thanks to this method, many advances have been resized in specifying potential regions for karstic development and many researchers have used it for measuring recharge (Duran et al. 2004; Farfán et al. 2010; Fetter 2000; Galvão et al. 2015; Gerner et al. 2012; Hartmann et al. 2014; Jirkama 2007; Kirn et al. 2017; Lee et al. 2006; Marechal et al. 2006). Offering a more accurate and faster analysis, this method facilitates measuring recharge rate of underground water and development of karstification (Kirn et al. 2017). More particularly, IGME (Instituto Geologicoly Minero de Espana), a Spanish geological institute, has developed APLIS, a new analysis for measuring real recharge rate in karstic formations and accurate potential development patterns by considering natural characteristics of the aquifer (altitude, slope, lithology, infiltration, soil) (Carrasco 2005). Unlike the established methods, this method is not dependent on accurate, periodic



data, a fact which makes it much economical in terms of water resource management (Martos-Rosillo et al. 2008). In addition, this method allows for mapping the spatial distribution of natural recharge in the aquifers and development of karstic formations (Martos-Rosillo et al. 2013). APLIS also can perform accurate examination of natural recharge in carbonate minerals (including direct recharge, local concentrated recharge via shallow water holes, and even indirect recharge from the substrate surface waters) (Mejías et al. 2012). APLIS can be regarded as a novel method among numerical formula for calculating recharge that can be called calculation method for aquifer recharge by remote sensing and GIS (Díaz-Guanche et al. 2013). The problems associated with conventional methods are no longer the case here thanks to assessment that considers natural properties of a given karstic aquifer (altitude, slope, lithology, gradual infiltration, fracture, and soil) to evaluate infiltration. APLIS was used in 8 aquifers in Spain. In this trial experiment, results matched actual recharge rates (Moussavi-Harami et al. 1992). The method was tested in a various place worldwide with success: SE Spain (Pedrera et al. 2017 Ramazani Oomali et al. 2008), Greece (Saadat 2014), Cuba (Samper et al. 2005), NE Oman (Scanlon et al. 2002), Mediterranean (Andreo et al. 2008), and southern Spain (Scibek and Allen 2006; Andreo et al. 2015).

This research also examines a spatial identification of regions for developing karsts in Mashhad on a 1:250,000 scale map. The local specifications and lack of periodic data (due to recent severe reduction of water resources) make APLIS a necessary method. The importance of this study is particularly tangible for provision of water resources for drinking and agriculture in the region.

Area under study

General features

The area is situated between eastern latitudes of 58°30' and 60° and northern longitudes of 36° and 37°. It lies in drainage basin of Kashfrud which flows through Binalud and Hezar Masjed mountain ranges (Fig. 1). In topological terms, the area includes both plains and heights, with lowest spot standing 462 m AMSL. Heights are vast, including two distinct ranges with highest point standing 3310 m AMSL. City of Mashhad lies at center of the area under study and is located at 1050 m AMSL. Climatologically, the area is affected by different air masses, giving it a unique climate. Mashhad-Neyshabur plain displays cold, arid climate, while Mashhad-Quchan plain is characterized by cold, semi-arid climate or steppe climate. A small portion of Binalud and Hezar Masjed mountain ranges represents cold, wet climate, while the whole area under study is characterized by



Fig. 1 Geographical coordinates of the area under study

changing but mild climate, tending toward cold and arid conditions and by dry, hot summers and cold, wet winters. Direction of winds is mostly from SE to NW. Maximum temperature during summer is 43 °C, while winters can be as cold as -23° C.

Geological features and structure

Geologically, the area is characterized by deposits from Alborz zone and central Iran (west to the area under study) in Paleozoic period. From the advent of Jurassic sea to the end of Cretaceous period, deposition took place in *Kopedagh* (Shabanian et al. 2010) (Fig. 2). Stratigraphically, there existed a shallow lake in this region in Iran from The Liassic Epoch till early Oligocene Epoch.

Thick consequent marine depositions in the basin extended to 6000 m and Jurassic sediments are conformable to the Oligocene are facing each other. The marine deposits extend to Upper Cretaceous, and only mild Epeirogenesis movements occurred toward the end of the Jurassic and beginning of the Cretaceous. No significant gap has been found in depositing processes in this region, particularly in deeper deposits. The Upper Jurassic units (Mazduran formations) are the main origin of gaseous hydrocarbons. With retreating seas in the Paleocene, much of the region emerged but no considerable folded structures are detected. A regional land subsidence accounts for the reflow of sea into Kopedagh. Structurally, compressive stress stretches from N-NE to S-SE which primarily causes shortening and folding. With later increase in pressure, trust faults have been created in some place (Shabanian et al. 2010). Most of these faults are transcurrent and are characterized by reverse component, and sometimes, normal structures although some thrust faults are also found to have folded structures and are considerably older than transcurrent faults.

Materials and methods

Many methods are available for identifying karstic regions and for calculating recharge rate. The local specifications and lack of periodic data (due to recent severe reduction of water resources) make APLIS a necessary method. However, a systematic identification of determinants in form of base maps, one that aims to maintain and manage consumption of these resources, will prove less time-consuming and more economical. The key to such identification is drawing base patterns in form of base maps which would use APLIS to consider geological, lithological, pedological, tectonic, and topological properties of the area. The mathematical formula in this method involves a combinatory or adaptive layer that displays main infiltration and development parameters of karst in a certain area, because karstic development depends on a host of physical properties of the aquifer although these variables vary in terms of weight and importance in recharge. This means that information layer for each variable is assigned a score according to APLIS ranking system and are then combined





Fig. 2 Geological map of Mashhad: 1:250,000



using APLIS formula. The different multi-principle formulas (such as regression analysis (linear correlation with least squares), ideal point analysis, weighted linear combination) are used to determine the weight of each variable. Recharge rate and, consequently, the development of karst in formations are calculated by combining natural specification of each aquifer: AMSL altitude, slope, lithology, gradual infiltration, and soil. After inserting information layer of each variable into GIS, mean recharge rate and development of karst are determined by APLIS formula (Bakalowicz 2005):

$$R = (A + P + (3 \times L) + (2 \times 1) + S)/0.9$$
(1)

where *R* is mean recharge rate, *A* is altitude, *P* is slope, *L* is lithology, *I* is preferential infiltration, and *S* defines the soil type. The weight of each variable represents its importance in determining recharge rate (Widiastuti 2012). Thus, lithology is three times more important than altitude (AMSL), slope, and soil, while gradual infiltration is two times more important than altitude, slope, and soil.

Results and discussion

Conversion of an issue into a hierarchical structure stands as the most important part of analytical hierarchy process. In identifying karstic regions, our goal is to specify water resources by considering infiltration percentage or recharge in formation. Criteria and sub-criteria include those variables which make options differ. Each variable is characterized by a certain weight, and every region is credited according to criteria. In this way, mean recharge rate (R) for each aquifer (in percent) is used to identify karstic regions by considering mean R for each spatial unit on the map. There are ten ranks for each variable, and each is assigned a number between 1 and 10 in form of an arithmetic progression (Bakalowicz 2005). 1 and 10 indicate minimum and maximum infiltration, respectively. Therefore, in order to determine the penetration (R) in the APLIS equation, firstly, the information layers in the APLIS equation should include: elevation (A), slope (P), lithology (L), preferential penetration (I), and soil type (S) in the GIS environment. Then, the appropriate valuation for each layer of information must be done for use in the APLIS equation. The basis of the valuation is based on the factors influencing the penetration. The following factors have been investigated:

Altitude (A)

This variable is divided into units of 300 MS (Table 1), and each unit is indicated by a score between 1 and 10 (Bakalowicz 2005). Minimum altitude is 462 m; therefore, this variable extends through 6 units (5–10). Given the high altitude of the bulk of the area under study, it is generally explained by large numbers in the rank, with 10 being the frequent indicator (Fig. 3). The layer for this variable and scores for the units show that higher altitudes receive more precipitation, meaning that recharge happens in a larger scale and more karstic formations are achievable. Also, it is understood that no significant difference is visible among units higher than 3100 m AMSL.

Slope (P)

Using GIS software, slope map was produced by inserting DEM map of altitude. The result was ranked in 9 units (Table 2). Score 7 was dropped in scoring this variable. Ranking this variable in varying distance blocks follows the ranking used in *Andles Biological Data System* where slope is divided into 9 ranks and score 7 is omitted (Bakalowicz 2005). Assigned scores become smaller as slope decreases. In other words, sharper slopes are associated with smaller recharge rate (Fig. 4). Therefore, minimum slope is represented with largest score (equal 10). In high altitudes we see 3, 4, and 5, scarcely getting any larger.

Lithology (L)

Karst development often occurs in calcareous and dolomite carbonate formations as well as in evaporation rocks. The existence of any factor such as joint, slot, and fault that causes more water to penetrate these hard formations will increase the development of karst in each formation. But in the APLIS method, any rock that can penetrate is capable of karsting, which is one of the drawbacks of this method. Field observations as well as stratigraphical studies let us use lithological and tectonic properties (Table 3) of the region under study on a scoring scale of 1 to 5 (Fig. 5). Very few fractures in the area are responsible for reducing the effectiveness of this variable,

Table 1 The altitude variable and scoring

Value	3	4	5	6	7	8	9	10
Elevation (m)	462-600	600-900	900-1200	1200-1800	1800-2100	2100-2400	2400-2700	2700-3310





Fig. 3 Altitude map and scoring elevation layers

Table 2 The slope variable and scoring	Value	1	2	3	4	5	6	8	9	10
o o o mg	Slope (%)	>100	76-100	46-76	31-46	21-31	16-21	8-16	3-8	<3

Soil (S)

Table 4 shows kinds of soil and the scoring system. Evolution of soil means thicker formations and smaller grains, hence, less infiltration. Thus and according to field observations, scoring matched the soil typology (Fig. 6). In fact, the high soil thickness due to its evolution on the surface of the earth and the fine grain components of the soil cause delays in the penetration process and consequently evaporation increases.

Preferential infiltration (I)

This layer is the sum of scores for three layers: slope (S), fractures (F), and cover (C).





Fig. 4 Slope map and scoring

 Table 3
 The lithology variable and scoring of this variable

Value	Lithology variable
9–10	Limestone and karst dolomite
7–8	Limestone and fissured and fractured and partly karst dolomite
5–6	Limestone and fissured and fractured dolomite
4	Limestone-marl, sand, and gravel
3	Conglomerate
2	Intrusive and metamorphic rocks
1	Shale, silt, sand

$$I = C + F + S \tag{2}$$

The presence of any joint, fracture, and fault increases the penetration of hardened formations and increases the karst formation of the desired formation. These three variables are scored according to their contribution to gradual infiltration (Table 5). Fractures in three ranks are scored according to the distance from the fracture. This was done with a buffer tool in the GIS environment. Slope was also scored in 9 units where each means a percentage fraction of the slope. Increasing slope increases the speed of runoff and decreases penetration time and decreases penetration. In case of cover, rocky areas got 1, while soil areas got 0. In fact, in determining the factor of penetration preferential, the soil factor is an inhibitory factor in penetration and increasing time for evaporation of water. Eventually, the mean value of these three layers was used to determine the gradual infiltration of various areas (Fig. 7).





Fig. 5 Lithology map and values of the layers

In this way, it was found that gradual infiltration occurs in regions characterized by limestone and mild slopes, making them prone to form karstic structures. In this scoring system, 1 and 10 represent minimum and maximum gradual infiltration, respectively.

Calculating recharge rate (R)

In order to measure recharge rate (percentage) of existing aquifers and, then, to identify potential karstic developments, the information layers were integrated by APLIS method in

Table 4	The soil	variable	and	scoring	of this	variable
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Verti sols	Plano sols	Chromic luvi sols	Histo sols and luvi sols	Solon- chaks	Cam- bisols	Eutric cam- bisols	Calcareous regosols	Arenosols and xerosols	Lep- tosols	Soil variable
1	2	3	4	5	6	7	8	9	10	Value





Fig. 6 Soil map and soil values

 Table 5
 The preferential infiltration variable and scoring of this variable

(Percent)	≤3	3-8	8-16	16-21	21-31	31-46	46-76	76-100	>100
Value	10	9	8	6	5	4	3	2	1
Distance from the main fracture		≤50	50-150	150-300	Litholog	gy S	oil areas	Rocky areas	
Value			10	6	2	Value	0		1

GIS (Fig. 8). For this purpose, the information elevations (*A*), slope (*P*), lithology (*L*), preferential penetration (*I*), and soil type (*S*) are evaluated as a raster, according to the APLIS equation $(R = (A + P + (3 \times L) + (2 \times I) + S) / 0.9)$ combined in the RASTER calculator tool, and the penetration rate (*R*) is obtained in percentage terms.

According to our calculations, minimum, maximum, and mean recharge rate in our study are 15, 96, and 49%, respectively. In terms of location, the mean recharge rate in our study is 49%. This means that it falls between 32% and 54%, the acceptable range suggested by other researches (Bakalowicz 2005).





Fig. 7 Preferential infiltration map and values

Conclusion

In order to locate karsts, the chief determinants were studied in form of digital data in GIS on a map of locations. The results determine the recharge percentage and karstic regions on a 1:250,000 map of Mashhad in NE Province of North Khorasan. The applied method was APLIS hierarchical analysis. Accordingly, parameters such as altitude, slope, lithology, soil, and gradual infiltration (distance from fracture, slope, and presence/absence of soil) were assumed to be influential factors in recharge and development of karsts. The results were classified into recharge (5 classes) and karstic development (5 classes). Considering the correlation between development of karstification and infiltration into formation, we can see that development of karstification means increase in infiltration, while reduction of karstification means less infiltration. Therefore, mean recharge rate (R) in percentage for each spatial unit helps us identify locations with potential for becoming karstification (Willaarts et al.2012) (Table 6).

It was found that locations with limestone and fissured and fractured and partly karst dolomite in high height and





Fig. 8 Approximate recharge rate in APLIS in the area under study

Table 6 Qualitative classification of recharge	Different recharge ranges (%)	20≤	40-20	60-40	80-60	80≥
and spatial development of karstification	Development of karstification	Very low	Low	Moderate	High	Very high

devoid of soil are associated with the chief karstic formations in western and northeastern parts of the area under study. Considering the ability of the mapping obtained in water resources management and determining the water balance components, including estimating penetration percentages (and consequently evaporation and groundwater storage) and identifying the location of the karstic resources compared to other existing methods much higher, the resulting base map, therefore, is used to locate and manage karstic water resources more economically and faster (Zagana et al. 2011).

Compliance with ethical standards

Conflict of interest The authors declare that they have no conflict of interest.

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References

- Aguilera H, Murillo JM (2009) The effect of possible climate change on natural groundwater recharge based on a simple model: a study of four karstic aquifers in SE Spain. Environ Geol 57(5):963–974
- Allocca V, Manna F, De Vita P (2014) Estimating annual groundwater recharge coefficient for karst aquifers of the southern Apennines. Italy Hydrol Earth Syst Sci 18(2):803–817
- Andreo B, Vías J, Durán JJ, Jiménez P, López-Geta JA, Carrasco F (2008) Methodology for groundwater recharge assessment in carbonate aquifers: application to pilot sites in southern Spain. Hydrogeol J 16(5):911–925
- Andreo B, Carrasco F, Durán J.J, Jiménez P, LaMoreaux J (2015) Hydrogeological and Environmental Investigations in Karst Systems. Springer Berlin Heidelberg, Vol. 1, PP: 638.
- Awawdeh M, Obeidat M, Al-Mohammad M, Al-Qudah K, Jaradat R (2014) Integrated GIS and remote sensing for mapping groundwater potentiality in the Tulul al Ashaqif. Northeast Jordan Arabian Journal of Geosciences 7(6):2377–2392
- Bakalowicz M (2005) Karst groundwater: A challenge for new resources. Hydrogeol J 13(2):148–160
- Bouraoui F, Vachaud G, Chen T (1998) Prediction of the effect of climatic changes and land use management on water resources. Phys Chem Earth 23:379–384
- Burke S (1995) Land surface parameterization: regionalized versus distributed approach to groundwater recharge. Phys Chem Earth 20:331–337
- Carrasco F, Valsero JJD, LaMoreaux JW (eds) (2014) Advances in research in karst media. Springer, Berlin
- Conrad J, Nel J, Wentzel J (2004) The challenges and implications of assessing groundwater recharge: a case study - northern Sandveld, Western Cape. South Africa Water SA 30:75–81
- Chenini I, Mammou AB, El May M (2010) Groundwater recharge zone mapping using GIS-based multi-criteria analysis: a case study in Central Tunisia (Maknassy Basin). Water Resour Manage 24(5):921–939
- Díaz-Guanche C, Aldana-Vilas C, Farfán-González H (2013) Mapping Groundwater Vulnerability in Guanahacabibes National Park, Western of Cuba. Management of Water Resources in Protected Areas. Springer, Berlin Heidelberg, p pp.87-94
- Duran Z, Doğru A.G, Toz G (2004) Web-based multimedia GIS for historical sites. Proceedings of the ISPRS Congress, Geo-Imagery Bridging Continents. 35(85): 434-438
- Farfán H, Corvea J.L, De Bustamante I (2010). Sensitivity analysis of APLIS method to compute spatial variability of karst aquifers recharge at the National Park of Viñales, Cuba, Advances in Research in Karst Media, Springer Berlin Heidelberg, Part of the series Environmental Earth Sciences, 19–24.

Fetter C. W (2000). Applied hydrogeology. Prentice hall.

- Galvão P, Halihan T, Hirata R (2015) Evaluating karst geotechnical risk in the urbanized area of Sete Lagoas, Minas Gerais. Brazil Hydrogeology Journal 23(7):1499–1513
- Gerner A, Schütze N, Schmitz GH (2012) Portrayal of fuzzy recharge areas for water balance modelling: a case study in northern Oman. Advances in Geosciences 31(31):1–7
- Hartmann A, Mudarra M, Andreo B, Marın A, Wagenerl T, Lange J (2014) Modeling spatiotemporal impacts of hydro climatic

extremes on groundwater recharge at a Mediterranean karst aquifer. Water Resources Research. 50(8): 6507-6521

- Healy RW (2010) Estimating groundwater recharge. Cambridge University Press, Cambridge
- Jirkama MI, Sykes JF (2007) The impact of climate change on spatially varying groundwater recharge in the grand river watershed (Ontario). J Hydrol 338:237–250
- Kirn, L., Mudarra, M., Marín, A., Andreo, B., & Hartmann, A. (2017). Improved Assessment of Groundwater Recharge in a Mediterranean Karst Region: Andalusia, Spain. In *EuroKarst 2016, Neuchâtel*, Springer, Cham. (pp. 117–125).
- Kumar A, Pandey AC (2016) Geoinformatics based groundwater potential assessment in hard rock terrain of Ranchi urban environment, Jharkhand state (India) using MCDM–AHP techniques. Groundwater for Sustainable Development 2:27–41
- Lee CH, Chen WP, Lee RH (2006) Estimation of groundwater recharge using water balance coupled with base-flow-record estimation and stable-base-flow analysis. Environ Geol 51:73–82
- Marechal JC, Dewandel B, Ahmed S, Galeazzi L, Zaidi FK (2006) Combined estimation of specific yield and natural recharge in a semi-arid groundwater basin with irrigated agriculture. J Hydrol 329:281–293
- Martos-Rosillo S, Pérez-Fernández F, Durán JJ (2008) Estimation of the average annual recharge in the carbonate aquifers of Sierra de Estepa (Seville) using the APLIS method. Geotemas 10:1567–5172
- Martos-Rosillo S, Rodríguez-Rodríguez M, Pedrera A, Cruz-SanJulián JJ, Rubio JC (2013) Groundwater recharge in semi-arid carbonate aquifers under intensive use: The Estepa Range aquifers (Seville, southern Spain). Environmental earth sciences 70(6):2453–2468
- Mejías M, Ballesteros BJ, Antón-Pacheco C, Domínguez JA, Garcia-Orellana J, Garcia-Solsona E, Masqué P (2012) Methodological study of submarine groundwater discharge from a karstic aquifer in the Western Mediterranean Sea. J Hydrol 464:27–40
- Moussavi-Harami R, Robert LB (1992) Geohistory Analysis and Petroleum Reservoir Characteristics of Lower Cretaceous (Neocomian) Sandstones, Eastern Kopet-Dagh Basin. Northeastern Iran AAPG Bulletin 76(8):1200–1208
- Nagarajan M, Singh S (2009) Assessment of groundwater potential zones using GIS technique. Journal of the Indian Society of Remote Sensing 37(1):69–77
- Ramazani Oomali R, Hafezi Moghaddas N, Omidi P, Eftekharnejhad J (2008) A Model for Active Tectonics in Kope Dagh (North-East Iran). World Applied Sciences Journal 3(2):312–316
- Russo TA, Fisher AT, Lockwood BS (2015) Assessment of managed aquifer recharge site suitability using a GIS and modeling. Groundwater 53(3):389–400
- Saadat H, & Mohammadi Z (2013) Hydrogeological characterization of the Asmari limestone aquifer, Anar anticline, Zagros Region, Iran. Carbonates and Evaporites, 1–13.
- Samper F.J, Garcı'a Vera M.A, Pisani B, Varela A, Losada J.A, Alvares D, Espinha Marques J (2005) Aplicacio'n de modelos hidrolo'gicos y sistemas de informacio'n geogra'fica para la estimacio'n de los recursos hı'dricos: GIS-VISUAL-BALAN y su aplicacio'n a cuencas atla'nticas en España (Valiña's) y Portugal (Serra da Estrela). Paper presented at the VII Jornadas de Investigacio'n en la Zona no Saturada del Suelo, A Coruña, 274–296.
- Senanayake IP, Dissanayake DMDOK, Mayadunna BB, Weerasekera WL (2016) An approach to delineate groundwater recharge potential sites in Ambalantota, Sri Lanka using GIS techniques. Geosci Front 7(1):115–124
- Scanlon B, Healy R, Cook P (2002) Choosing appropriate techniques for quantifying groundwater recharge. Hydrogeol J 10(1):18–39



- Scibek J, Allen DM (2006) Comparing modelled responses of two high-permeability, unconfined aquifers to predicted climate change. Globe Planet Change 50:50–62
- Shabanian E, Bellier O, Abbassi MR, Siame L, Farbod Y (2010) Plio-quaternary stress states in NE Iran: Kopeh Dagh and Allah Dagh-Binalud mountain ranges.". Tectonophysics 480(1):280–304
- Tapia F.O, Mora F.E (2004) The Information Society: new horizons for science. Paper presented at the 19th international CODATA conference. Berlı'n. p 13.
- Touhami I, Andreu JM, Chirino E, Sánchez JR, Moutahir H, Pulido-Bosch A, Bellot J (2013) Recharge estimation of a small karstic aquifer in a semiarid Mediterranean region, southeastern Spain using a hydrological model. Hydrol Process 27(2):165–174
- Widiastuti A. P (2012) ZONASI KERENTANAN AIRTANAH BEBAS terhadap PENCEMARAN dengan METODE APLIS

di KECAMATAN WONOSARI KABUPATEN GUNUNG-KIDUL. Journal Bumi Indonesia, 1(2).

- Willaarts BA, Volk M, Aguilera PA (2012) Assessing the ecosystem services supplied by freshwater flows in Mediterranean agroecosystems. Agric Water Manag 105:21–31
- Zagana E, Tserolas P, Floros G, Katsanou K, Andreo B (2011) First outcomes from groundwater recharge estimation in evaporate aquifer in Greece with the use of APLIS method, Advances in the Research of Aquatic Environment, Volume 2 of the series Environmental Earth Sciences, Springer Berlin Heidelberg, PP: 89-96.

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