



# Geoelectrical investigation of groundwater potential, at Nigerian Union of Teachers Housing estate, Paggo, Minna, Nigeria

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## Abstract

Vertical electrical sounding (VES) was carried out to evaluate the groundwater potentials of Nigerian Union of Teachers Housing estate, Paggo, Minna. The ABEM Terrameter model SAS 4000 was used to collect the subsurface data for the evaluation of groundwater potentials. Six profiles with ten VES stations on each profile were sounded with intervals of 100 m apart. It has a maximum current electrode separation (AB/2) of 100 m. Three and four geoelectric layers were obtained which include: topsoil, weathered layer, fractured layer and fresh basement layer. The stratigraphy of the subsurface shows: topsoil (67.5–835.1  $\Omega$ m), weathered layer (108.0–939.7  $\Omega$ m), fractured layer (118.9–242.0  $\Omega$ m) and the fresh basement layer (1041.0–9704.0  $\Omega$ m). Sixteen VES points were recommended as groundwater potentials of the area having weathered/fractured layer resistivity varying between 65.84 and 454  $\Omega$ m, depths range from 10.61 to 26.37 m, and thickness varies between 9.255 and 24.69 m. The observed frequencies in curve types include 58.33% of H, 26.67% of A, 10% of HA and 5% of AA. A correlation of the borehole log data with the VES was made and is in agreement. Viable boreholes for good portable water should be sited at VES stations A<sub>10</sub> and D<sub>9</sub> with a reasonable thickness of 21.2 and 10.68 m, respectively.

**Keywords** Geoelectric · Evaluation · Vertical electrical sounding · Longitudinal unit conductance · Overburden protective capacity · Groundwater potential

## Introduction

The growth of any community is hinged on the availability of basic amenities such as water, good road network and electricity. The search for sustainable, clean and portable water is a struggle that will never end as it aids in the growth of any community (Salako et al. 2009).

Niger State in north central Nigeria experiences an annual rainfall which ranges from 1200 mm to 1600 mm from the southern part of the state to the northern region. The duration of the rainy season ranges from 120 to 150 days or more from the north to the south (Baimba 1978). The amount of rainfall is usually limited to few months within the year, such that water from surface sources cannot meet the demand for development (Baimba 1978). This inadequacy and pollution of surface water brings about the overdependence on

groundwater from hand dug well and borehole either for individual or public consumption.

Mapping of the subsurface structure and interpretation has provided information about the geologic formations and physical properties of the geologic materials (George et al. 2015a, b). In groundwater exploration, the knowledge of the aquifer characteristics is important as it helps in determination of natural flow of water, depth to bedrock, availability, quantity and quality of the groundwater. Since water occurs both as surface water and groundwater thus forming the hydrological cycle (a continuous re-circulatory movement of waters of the earth).

Groundwater is the water in porous rocks beneath the water table. In other words, it is the water that is contained in aquifers. An aquifer is any geologically deposited material that has the ability to store and transmit significant quantity of water. Groundwater contains dissolved ions and allows current to flow through it due to the conductive nature of the ions. Permeability, porosity, resistivity, thickness of the layers and aquifer yield are subsurface properties that play important role in groundwater movement, availability and potential. Aquifers are found deep beneath the earth's

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surface, and due to their depth of storage and natural filtration through the different subsurface layers (soil horizons), groundwater is relatively pure and grossly protected from surface pollutants. Groundwater is a source of uncontaminated water and is very important for our daily needs like drinking and even agricultural purposes. Groundwater is of major importance to civilization because it is the largest reserve of drinkable water in the regions where humans live. Groundwater may appear at the surface in the form of springs, or it may be tapped by wells/boreholes (Salako et al. 2009). Therefore, to have a sustainable groundwater development for our needs, the study of groundwater potentials becomes very important.

The subsurface is made up of different geologic materials occurring at varying depths. In a basement complex terrain, the lithological unit that can be observed comprises of the weathered layer, weathered/fracture basement and fresh basement (Omorinbola 1984). The porosity and permeability capacity of this subsurface lithology depends on the type of geologic material occupying them (Ojo and Olorunfemi, 1990). Weathered layer that consists of clay would have less porosity, while a highly weathered/fracture basement would be highly porous and permeable (Hazell 1992).

## Geology of the study area

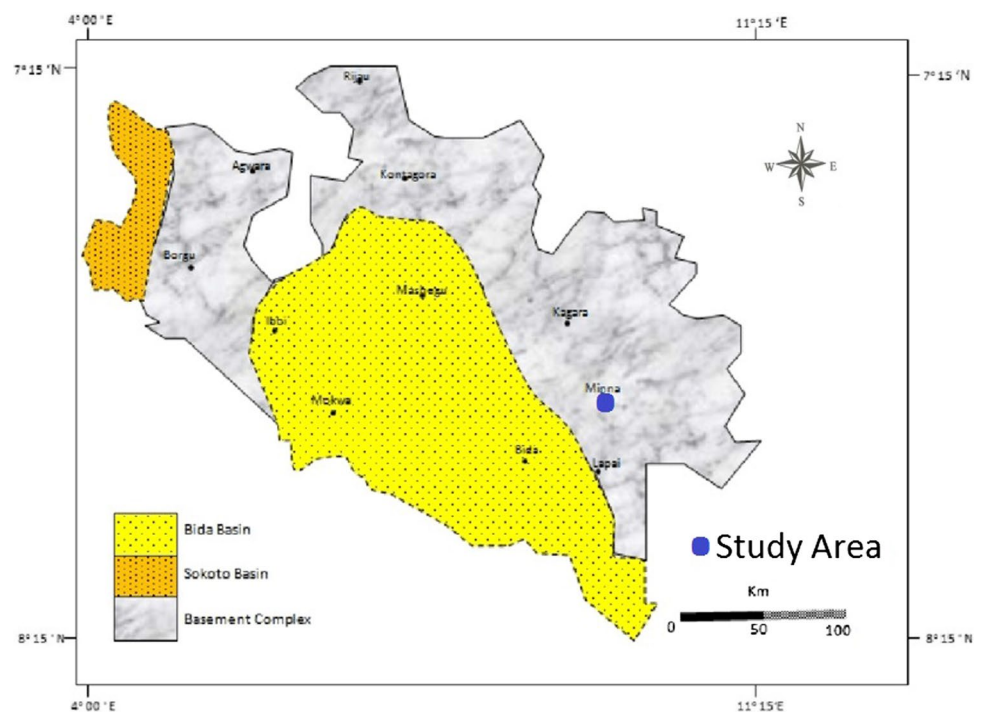
The study area is located at Paggo village in Minna, Niger State. It lies in north central Nigerian Basement Complex with latitude  $9^{\circ}27'N$  and  $9^{\circ}28'N$  and longitude  $6^{\circ}38'E$  and

$6^{\circ}39'E$ . The mean annual temperature is  $22\text{--}25^{\circ}C$ . The area is characterized by both dry and rainy seasons. It has an average rainfall of 1200–1300 mm. The vegetation covering this area is Guinea savannah which is characterized by tall grasses, shrubs and also sparsely distributed trees. Dry season is usually accompanied by dust and fogs. The dry season usually last within the month of November to February followed by the rainy season from April to October. Most of the rocks are granite, gneiss and quartzite. Niger state has two main rock formations like most state that is on the same latitudes which are sedimentary and basement complex rocks (Amadi et al. 2012) (Fig. 1).

## Materials and method of the study

The data were acquired with the ABEM SAS 4000 Terrameter, Global Positioning System (GPS) for taking accurate coordinate of the VES point and elevations, metal electrodes, measuring tape, labelled tag (used in locating station position), hammer (used in driving the electrodes into the ground). The Schlumberger array was adopted. The electrode spread of AB/2 was varied from 1 to a maximum of 100 m. The electrical resistances obtained were multiplied by the corresponding geometric factor ( $k$ ) for each electrode separation to obtain the apparent resistivity. The apparent resistivity was computed using Eq. (1). Sounding data were presented as sounding curves, by plotting apparent resistivity against AB/2. The IPI2win software was then used to obtain the  $n$ -layer model curve for the Schlumberger

**Fig. 1** Geological map of Niger state showing basement complex and sedimentary basin (Amadi et al. 2012)



**Table 1** Resistivity values of rock types (Ajayi and Hassan 1990)

Rock Type	Resistivity ( $\Omega\text{m}$ )
Fadama loam	30–90
Sandy	100–200
Sand and gravel	100–180
Weathered and laterite	150–900
Fresh laterite	900–3500
Weathered basement	20–200
Fractured basement	500–1000
Fresh basement	> 1000

**Table 2** Longitudinal conductance/protective capacity rating (Ogungbemi et al. 2013)

Longitudinal conductance (mho)	Protective capacity rating
> 10	Excellent
5–10	Very good
0.7–4.9	Good
0.2–0.69	Moderate
0.1–0.19	Weak
< 0.1	Poor

sounding curves. This software automatically interprets the Schlumberger sounding curves. The plotted curves reveal the number of layers, thickness, depth and the average resistivity for each layer at different VES points automatically.

$$\rho_a = KR \tag{1}$$

where  $\rho_a$  is an apparent resistivity and the earth resistance ( $R$ ) is given as

$$R = \frac{\Delta V}{I} \tag{2}$$

The geometric factor,  $K$ , is expressed as

$$K = \pi \left( \frac{\left(\frac{AB}{2}\right)^2 - \left(\frac{MN}{2}\right)^2}{MN} \right) \tag{3}$$

The aquifer protective capacity characterization is based on the values of the longitudinal unit conductance of the overburden rock units in the area. The longitudinal layer conductance ( $S$ ) of the overburden at each VES station was obtained from Eq. (4). Table 1 gives the resistivity range with corresponding subsurface earth material (Ajayi and Hassan 1990) and Table 2 represents the protective capacity rating (Ogungbemi et al. 2013)

$$S = \sum_{i=1}^n \frac{hi}{\rho_i} \tag{4}$$

where  $h$  is the saturated thickness of each layer and  $\rho$  is the layer resistivity

The interpretation of the resistivity value for the n-layers was done for 60 VES points. The surfer11 computer programme was then used to produce iso-resistivity contour maps from the obtained data. The VES plots along the various profiles generate the geoelectric sections from where the resistivity variation with depth and thickness was obtained (Fig. 2).

## Results and discussion

The results were summarized in a tabular form, giving information about the average layer resistivity, depth of each layer, thickness and the curve types (Table 3), while geoelectric section where the parameters in Table 3 were obtained is presented in Fig. 2.

### Iso-resistivity contour map for the first layer

The iso-resistivity contoured map for the first layer was generated at an interval of 50  $\Omega\text{m}$  (Fig. 3). The resistivity values range from 50 to 950  $\Omega\text{m}$ . The resistivity value shows that central portion of the area is a fadama loam, while sandy clay and gravels are located around southwest, northwest and north central parts of the area (Table 1).

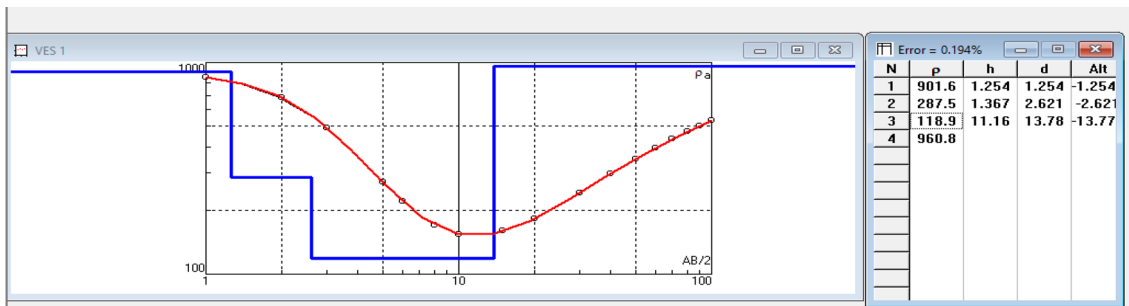
### Iso-resistivity contour map of the second layer

The iso-resistivity map for the second layer was generated at an interval of 50  $\Omega\text{m}$  (Fig. 4). Its resistivity values range from 0 to 1150  $\Omega\text{m}$ . From the map, it is observed that fresh laterite is scattered around the area and weathered laterite concentrates around northeast and central part of the area (Table 1). The weathered/fractured layer resistivity ranges from 65.84 to 454.9  $\Omega\text{m}$ , the depths of these weathered/fractured layers were found to be from 10.61 to 26.37 m and thickness 9.26 to 24.69 m, respectively. In a basement complex terrain, areas with overburden thickness of 15 m and above with fractured layer resistivity of < 1000  $\Omega\text{m}$  are good for groundwater development (Alhassan et al. 2015).

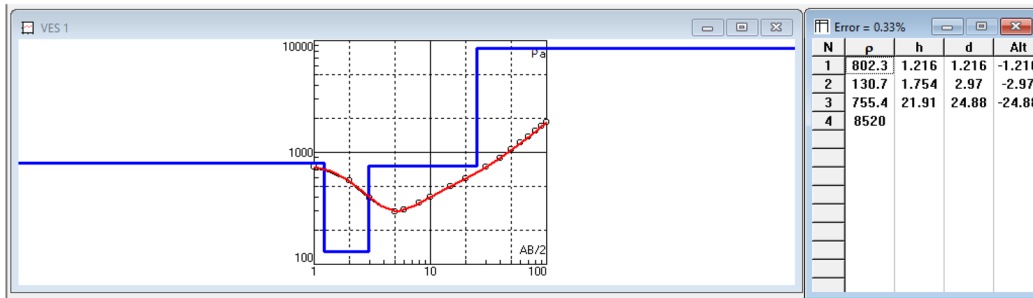
### Iso-resistivity map of the third layer

The iso-resistivity map for the third layer was generated at an interval of 2000  $\Omega\text{m}$  (Fig. 5). The resistivity value ranges from 2000 to 2800  $\Omega\text{m}$ . The resistivity values on this map indicate the fresh basement.

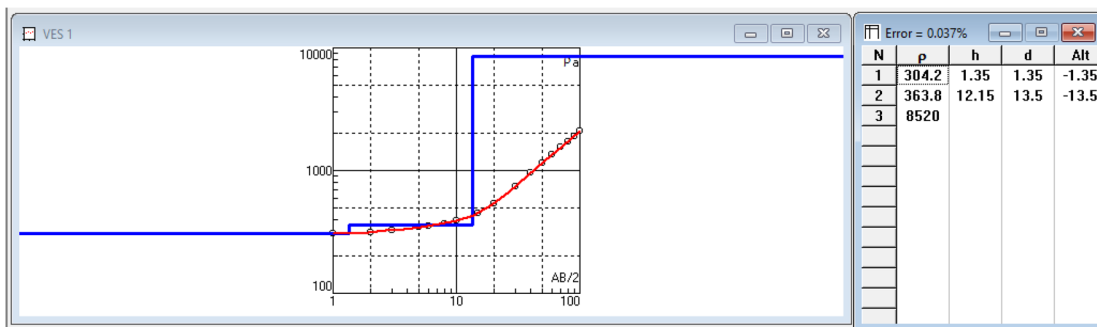
Iso-resistivity map of the fractured basement produced shows the resistivity distribution across the study area which varies from 50 to 950  $\Omega\text{m}$  (Fig. 6). The Isopach map of the fractured basement is produced which gives thickness of



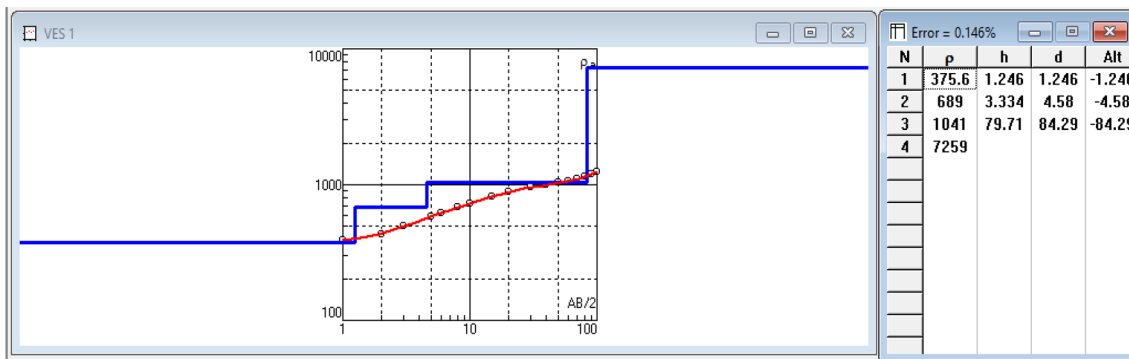
(a) Geoelectric Section (VES Curve A<sub>1</sub>)



(b) Geoelectric Section (VES Curve B<sub>1</sub>)



(c) Geoelectric Section (VES Curve C<sub>1</sub>)



(d) Geoelectric Section (VES Curve D<sub>1</sub>)

**Fig. 2** Geoelectric Section **a** (VES curve A<sub>1</sub>), **b** (VES curve B<sub>1</sub>), **c** (VES curve C<sub>1</sub>), **d** (VES curve D<sub>1</sub>)

the fractured basement across the area. The thickness varies between 1 and 23 m (Fig. 7). Isopach map of weathered layer is also produced, and it presents the thickness of the weathered basement within the study area (Fig. 8).

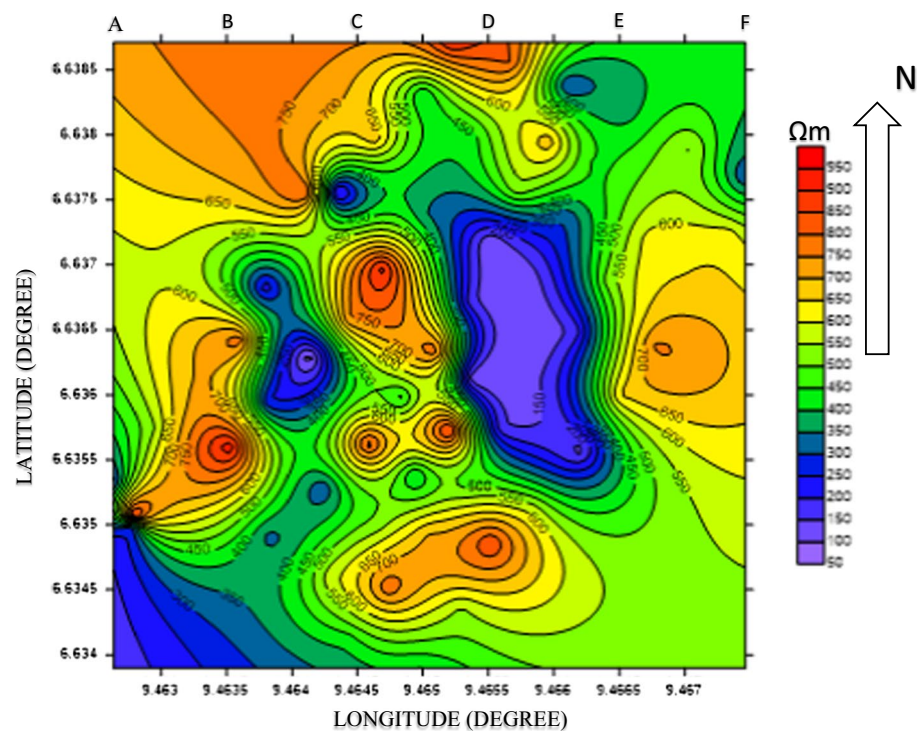
Isopach map of the overburden is produced. The map indicates that thick overburden is located around northwest, north central, north east, south east and south central portions of the area, while the shallow thickness is concentrated

**Table 3** Layer resistivity, depth and thickness

VES station	Latitude (°)	Longitude (°)	No. of layers	Layer resistivity				Layer depth			Layer thickness		
				$\rho_1$	$\rho_2$	$\rho_3$	$\rho_4$	$d_1$	$d_2$	$d_3$	$h_1$	$h_2$	$h_3$
A <sub>1</sub>	09.46521	06.63871	4	901.6	287.5	118.9	960.8	1.254	2.621	13.78	1.254	1.367	11.16
A <sub>2</sub>	09.46500	06.63832	4	399	586	652	7864	2.64	5.49	53.31	2.64	2.86	47.8
A <sub>3</sub>	09.46463	06.63797	3	684.8	112.8	8692		1.222	3.403	∞	1.222	2.181	∞
A <sub>4</sub>	09.46435	06.63756	3	165	743.4	9231		1.222	17.52	∞	1.222	16.3	∞
A <sub>5</sub>	09.46410	06.63720	4	816.9	173.7	367.5	8692	1.104	1.68	26.37	1.104	0.576	24.98
A <sub>6</sub>	09.46379	06.63683	3	213.3	123.6	5060		2.133	3.68	∞	2.133	1.547	∞
A <sub>7</sub>	09.46360	06.63642	3	784.5	127.4	4960		2.082	3.416	∞	2.082	1.334	∞
A <sub>8</sub>	09.46330	06.63602	3	715.7	293.5	3124		2.183	21.4	∞	2.183	19.22	∞
A <sub>9</sub>	09.46294	06.63575	4	539.7	143.4	725.9	1249	1.585	2.056	18.98	1.585	0.471	16.92
A <sub>10</sub>	09.46264	06.63533	3	282.2	108	894		1.25	22.4	∞	1.25	21.2	∞
B <sub>1</sub>	09.46567	06.63856	4	802.3	130.7	755.4	8520	1.216	2.97	24.88	1.216	1.754	21.91
B <sub>2</sub>	09.46538	06.63799	3	431	135	1094		1.272	1.575	∞	1.272	0.303	∞
B <sub>3</sub>	09.46502	06.63739	3	353.9	137.6	9231		1.222	3.825	∞	1.222	2.603	∞
B <sub>4</sub>	09.46469	06.63698	3	923	126.6	8692		1.324	6.791	∞	1.324	5.467	∞
B <sub>5</sub>	09.46443	06.63658	3	741.4	173	8133		1.197	5.791	∞	1.197	4.594	∞
B <sub>6</sub>	09.46413	06.63629	3	67.4	675	9132		1.35	18.3	∞	1.35	17	∞
B <sub>7</sub>	09.46382	06.63595	3	251.1	939.7	9048		1.222	24.33	∞	1.222	23.01	∞
B <sub>8</sub>	09.46353	06.63560	3	941.2	454.9	9231		1.354	10.61	∞	1.354	9.255	∞
B <sub>9</sub>	09.46279	06.63505	3	913.1	108.3	4138		1.199	3.569	∞	1.199	2.37	∞
B <sub>10</sub>	09.46278	06.63501	4	287	111	242	9231	1.2	2.04	17.1	1.2	0.837	15.1
C <sub>1</sub>	09.46609	06.63838	3	304.2	363.8	8520		1.35	13.5	∞	1.35	12.15	∞
C <sub>2</sub>	09.46596	06.63795	3	685	149.2	7864		2.228	6.438	∞	2.28	4.21	∞
C <sub>3</sub>	09.46572	06.63758	3	534.4	136.6	8200		1.297	14.34	∞	1.297	13.04	∞
C <sub>4</sub>	09.46557	06.63718	3	106	148.4	942		1.69	5.26	∞	1.69	3.56	∞
C <sub>5</sub>	09.46535	06.63672	3	181	289	923		1.4	4.69	∞	1.4	3.29	∞
C <sub>6</sub>	09.46510	06.63634	3	786.4	177.2	1970		1.324	5.02	∞	1.324	3.696	∞
C <sub>7</sub>	09.46483	06.63597	3	429.7	503.8	6437		1.38	12.4	∞	1.38	11	∞
C <sub>8</sub>	09.46458	06.63562	3	835.1	186.4	8820		1.297	4.412	∞	1.297	3.115	∞
C <sub>9</sub>	09.46424	06.63528	3	308.5	179.8	8859		1.354	4.778	∞	1.354	3.424	∞
C <sub>10</sub>	09.46383	06.63491	3	332	113	905		1.11	2.5	∞	1.11	1.39	∞
D <sub>1</sub>	09.46655	06.63817	4	375.6	689	1041	7259	1.246	4.58	84.29	1.246	3.334	79.71
D <sub>2</sub>	09.46630	06.63778	3	442.4	124.6	8351		1.311	3.52	∞	1.311	2.209	∞
D <sub>3</sub>	09.46610	06.63736	4	307.1	548.4	662.6	8023	1.585	8.389	14.78	1.585	6.804	6.389
D <sub>4</sub>	09.46587	06.63693	3	146.6	594.2	4062		4.864	26.14	∞	4.864	21.28	∞
D <sub>5</sub>	09.46569	06.63650	3	117	206.6	606		1.251	3.611	∞	1.251	2.36	∞
D <sub>6</sub>	09.46538	06.63613	3	111.9	833.8	5624		6.504	16.26	∞	6.504	9.757	∞
D <sub>7</sub>	09.46520	06.63573	3	887	210	905		1.396	5.975	∞	1.396	4.579	∞
D <sub>8</sub>	09.46494	06.63538	3	385.9	559.5	7864		1.505	35.27	∞	1.405	33.87	∞
D <sub>9</sub>	09.46465	06.63501	3	516.5	65.84	852		1.787	12.47	∞	1.787	10.68	∞
D <sub>10</sub>	09.46429	06.63460	4	580.2	137.8	332	8186	1.26	2.743	17.11	1.26	1.483	14.37
E <sub>1</sub>	09.46704	06.63790	3	555.5	420	8520		3.564	15.22	∞	3.654	11.66	∞
E <sub>2</sub>	09.46672	06.63721	3	617.2	886.8	6373		2.96	23.8	∞	2.96	20.9	∞
E <sub>3</sub>	09.46640	06.63688	3	517.5	161.7	5307		1.236	3.461	∞	1.236	2.225	∞
E <sub>4</sub>	09.46617	06.63649	3	171	112.8	8859		1.354	4.196	∞	1.354	2.842	∞
E <sub>5</sub>	09.46588	06.63613	3	145.1	1199	8595		2.373	10.62	∞	2.373	8.251	∞
E <sub>6</sub>	09.46560	06.63574	3	168.2	326.4	9231		1.405	14.63	∞	1.405	13.22	∞
E <sub>7</sub>	09.46533	06.63533	3	472	644	3587		1.297	12.98	∞	1.297	11.68	∞
E <sub>8</sub>	09.46504	06.63494	3	697.4	116.4	9132		1.354	6.843	∞	1.354	5.489	∞

**Table 3** (continued)

VES station	Latitude (°)	Longitude (°)	No. of layers	Layer resistivity				Layer depth			Layer thickness		
				$\rho_1$	$\rho_2$	$\rho_3$	$\rho_4$	$d_1$	$d_2$	$d_3$	$h_1$	$h_2$	$h_3$
E <sub>9</sub>	09.46473	06.63451	3	786.7	127.2	970.4		1.274	13.68	$\infty$	1.274	12.41	$\infty$
E <sub>10</sub>	09.46443	06.63412	3	423	105.1	923		1.199	2.976	$\infty$	1.199	1.777	$\infty$
F <sub>1</sub>	09.46746	06.63768	3	299	111	961		3.57	5.9	$\infty$	3.57	2.33	$\infty$
F <sub>2</sub>	09.46734	06.63724	3	622.6	137.8	1901		1.18	2.544	$\infty$	1.18	1.364	$\infty$
F <sub>3</sub>	09.46703	06.63681	3	653.7	146.3	2039		1.174	2.304	$\infty$	1.174	1.133	$\infty$
F <sub>4</sub>	09.46677	06.63638	3	755.6	231.9	8692		1.211	16.57	$\infty$	1.211	15.29	$\infty$
F <sub>5</sub>	09.46650	06.63600	3	650.9	281.7	4082		1.557	17.17	$\infty$	1.557	15.61	$\infty$
F <sub>6</sub>	09.46621	06.63557	3	131.3	390.9	8595		2.198	24.08	$\infty$	2.198	21.88	$\infty$
F <sub>7</sub>	09.46585	06.63514	3	562.6	354	1267		1.553	8.521	$\infty$	1.553	6.968	$\infty$
F <sub>8</sub>	09.46552	06.63486	3	852	155.3	8186		2.582	6.624	$\infty$	2.582	4.042	$\infty$
F <sub>9</sub>	09.46529	06.63436	3	546.4	157.5	7387		1.313	2.86	$\infty$	1.313	1.547	$\infty$
F <sub>10</sub>	09.46497	06.63390	3	397.9	180	970.4		1.129	16.64	$\infty$	1.129	15.51	$\infty$

**Fig. 3** Iso-resistivity contour map for the first layer

around extreme north, south west and central parts of the area (Fig. 9).

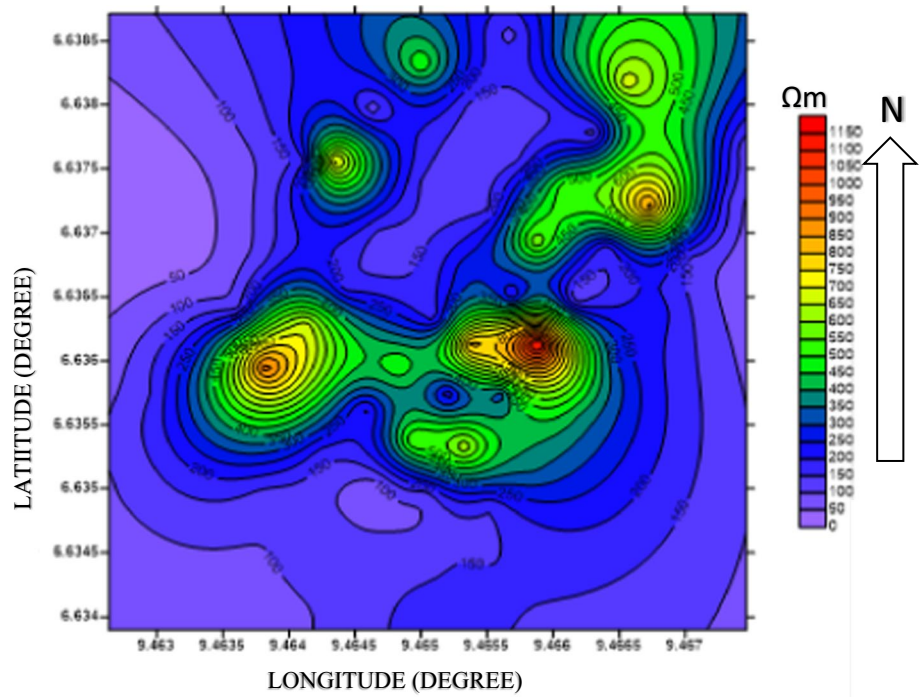
### Aquifer protective capacities evaluation

The study revealed that the area is characterized with poor, weak and moderate protective capacities having a longitudinal conductance ranging from 0.01 to 0.30. The longitudinal conductance values obtained were used to produce the longitudinal contour map which shows the distribution

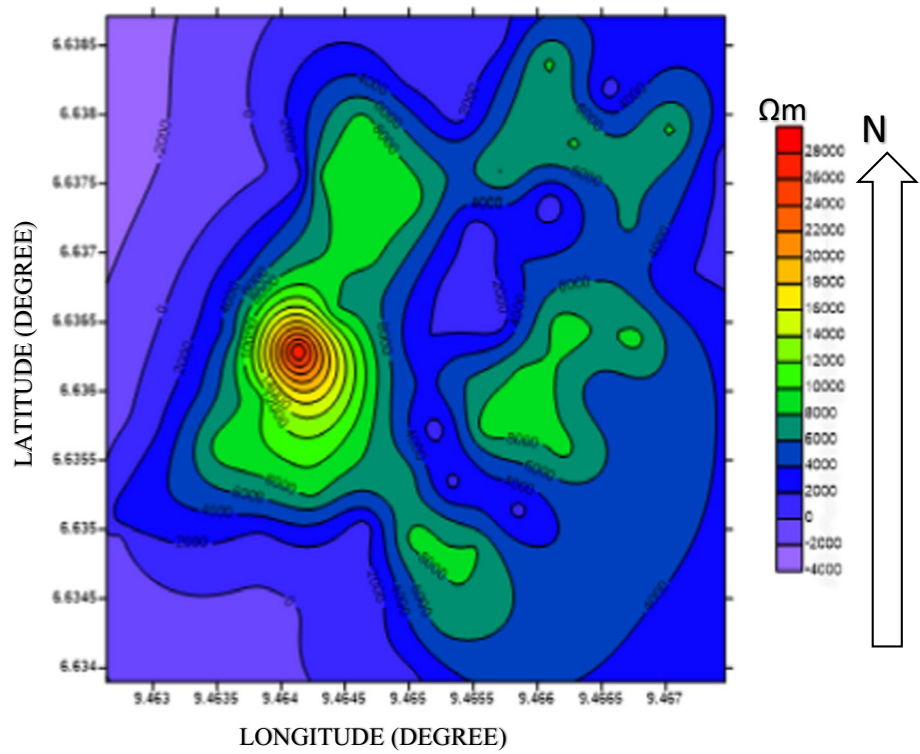
across the area (Fig. 10). The highest longitudinal conductance was seen at VES B<sub>6</sub> (0.30), and the lowest was seen at VES B<sub>2</sub> (0.004).

Table 4 shows 16 VES stations recommended as aquifer potentials of the area. The highest groundwater yield is often obtained from a fractured aquiferous zone or a sub-surface sequence that has a combination of a significantly thick and sandy weathered layer and fractured aquifer (Olorunfemi et al. 1999). A correlation of the nearby bore-hole log with the VES formation (Fig. 11) is in agreement.

**Fig. 4** Iso-resistivity contour map of the second layer



**Fig. 5** Iso-resistivity contour map of the third layer



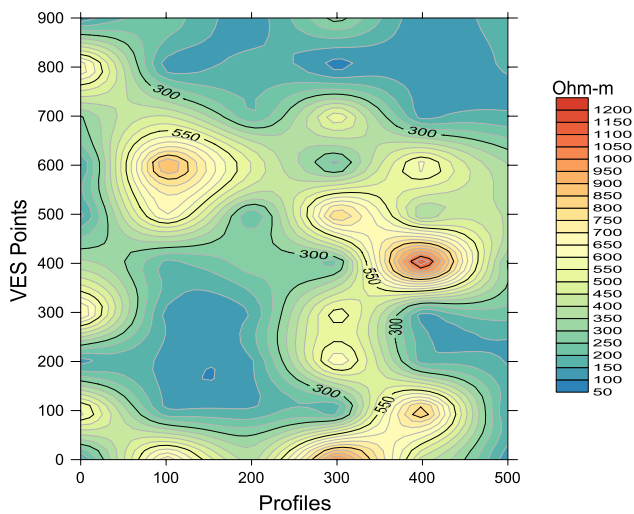


Fig. 6 Iso-resistivity map of the fractured basement

**The curve types**

The curve distribution (Fig. 12) indicates four types. The H-type is the dominant curve with 35% of the area. The weathered/fractured layer in the H-curve type is usually characterized with low resistivity value made up of clayey or sandy clay, and it is usually water saturated and highly porous (Olorunfemi et al. 1999). The A-curve type consists

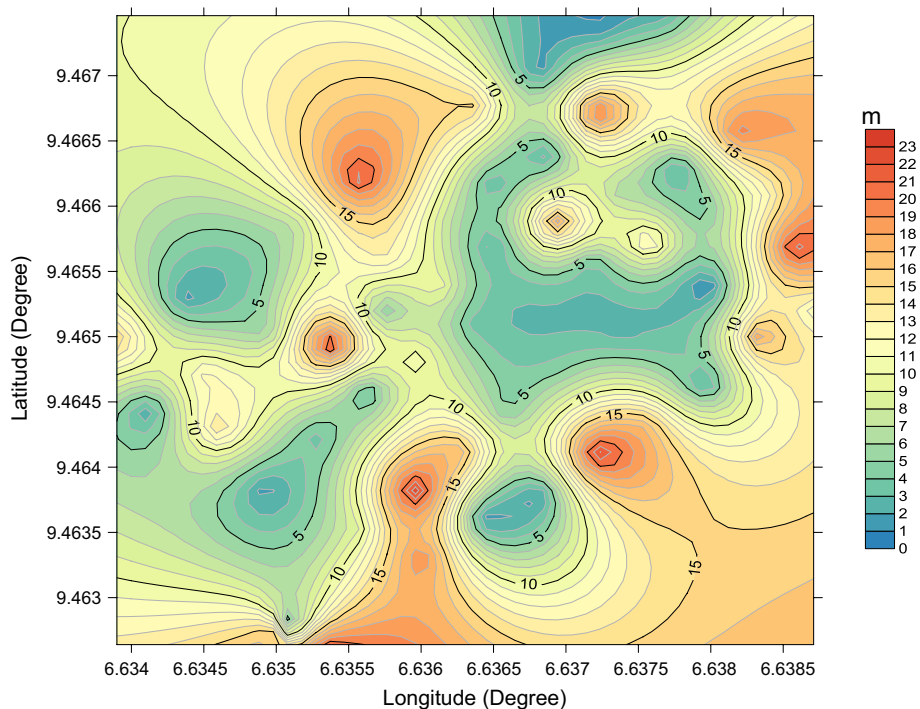
16% of the area, the HA curve type occupies 6%, while the AA occupies 3% of the study area.

**Conclusion**

In the basement complex area, groundwater usually is found in a discontinuous aquifer. When defining the potentials of an aquifer formation, it comes with a tedious exercise due to the complexity properties of the basement rock (Alhassan et al. 2017). Therefore, the groundwater potentials of the study area were classified by employing the electrical resistivity of the fractured layer, the depth of the layer and the thickness of the layer. From the result, it is revealed that the subsurface consists of 3–4 lithologic units which are top layer, weathered layer, fractured layer and fresh basement. The curve types were identified as H, HA, A and AA. The geoelectric and geologic sections for the profile were produced. The aquifer potentials of the area show poor, weak and moderate protective capacities having longitudinal conductance ranging from 0.01 to 0.30 mho.

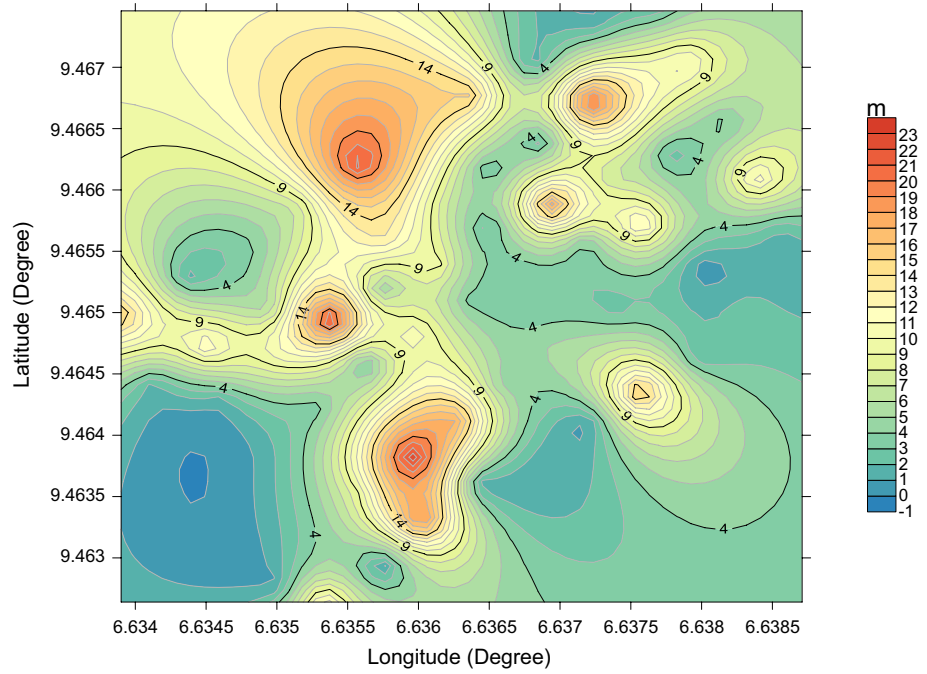
Sixteen VES stations were delineated as groundwater potentials of the study area, having fractured layer resistivity ranging from 65.84 to 454  $\Omega$ m. The depths of these layers are found ranging from 10.61 to 26.37 m and thickness ranging from 9.255 to 24.69 m. VES stations A<sub>10</sub> and D<sub>9</sub> are observed to have fine aquifer with depth

Fig. 7 Isopach map of the fractured basement

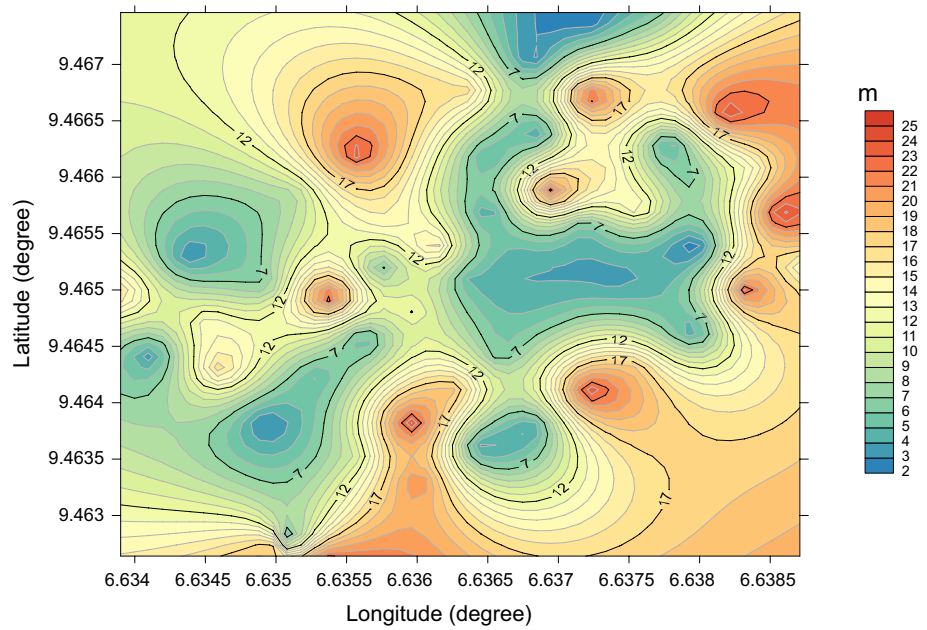




**Fig. 8** Isopach map of the weathered layer



**Fig. 9** Isopach map of the overburden



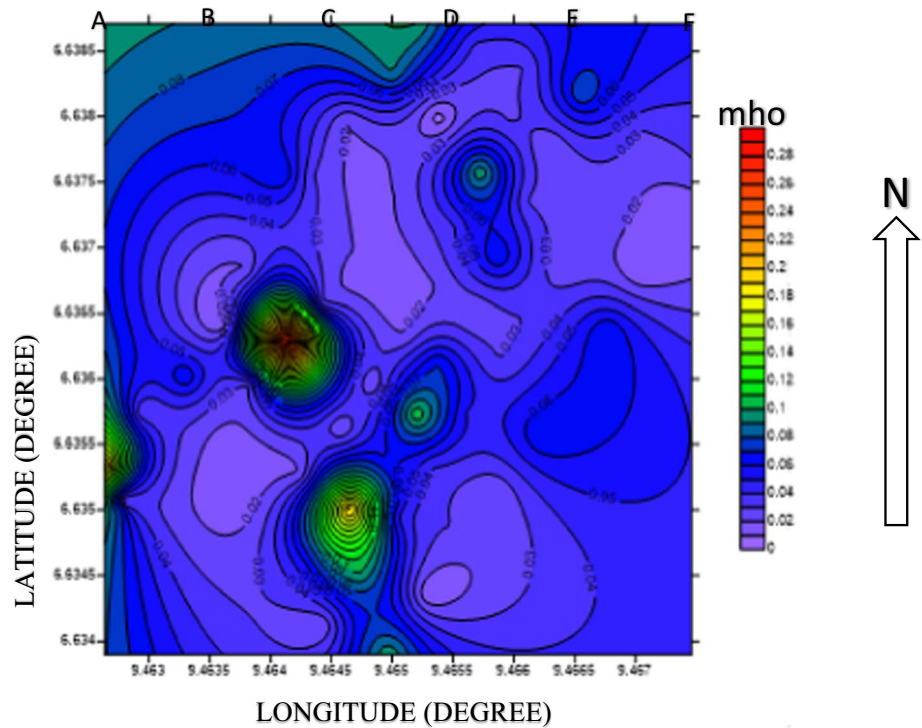
ranging from 12.47 to 22.4 m and also thickness ranging from 10.68 to 21.2 m.

- Areas with poor aquifer protective capacity should be avoided for sinking borehole to reduce leachates infiltration to the ground water

**Recommendations**

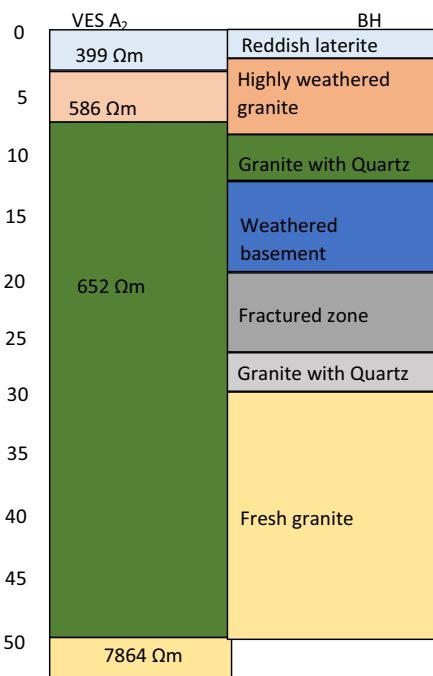
- The VES stations delineated for groundwater potentials should be considered for water development

**Fig. 10** Contour map of the longitudinal conductance distribution in the area

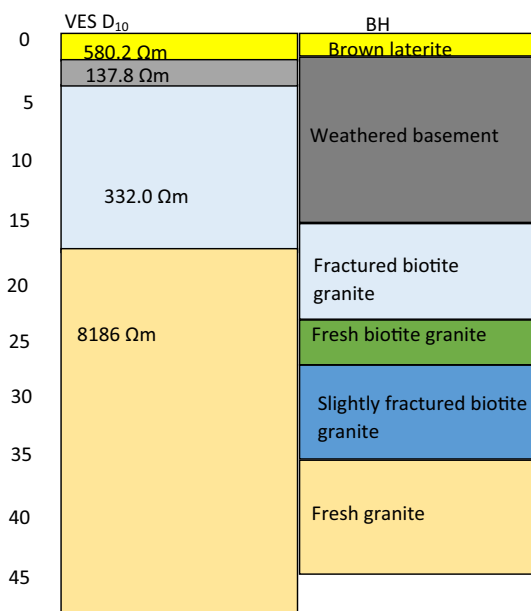


**Table 4** Recommended aquifer potentials of the study area

VES station	Latitude (°)	Longitude (°)	Layer nos	Layer resistivity (Ωm)	Layer depth (m)	Layer thickness (m)	Curve type
A <sub>1</sub>	09.46521	06.63871	4	118.9	13.78	11.16	HA
A <sub>5</sub>	09.46410	06.63720	4	367.5	26.37	24.69	HA
A <sub>8</sub>	09.46330	06.63602	3	293.5	21.40	19.22	H
A <sub>10</sub>	09.46264	06.63533	3	108	22.40	21.20	H
B <sub>8</sub>	09.46353	06.63560	3	454.9	10.61	9.255	H
B <sub>10</sub>	09.46278	06.63501	4	242	17.10	15.10	HA
C <sub>1</sub>	09.46609	06.63838	3	363.8	13.50	12.15	A
C <sub>3</sub>	09.46572	06.63758	3	136.6	13.37	13.04	H
D <sub>9</sub>	09.46465	06.63501	3	65.84	12.47	10.68	H
E <sub>1</sub>	09.46704	06.63790	3	420	15.22	11.66	H
E <sub>6</sub>	09.46560	06.63574	3	326.4	14.63	13.22	A
E <sub>9</sub>	09.46473	06.63451	3	127.2	13.68	12.41	H
F <sub>4</sub>	09.46677	06.63638	3	231.9	16.50	15.29	H
F <sub>5</sub>	09.46650	06.63600	3	281.7	17.17	15.61	H
F <sub>6</sub>	09.46621	06.63557	3	390.9	24.08	21.88	A
F <sub>10</sub>	09.46497	06.63390	3	180	16.64	15.51	H



(a) Correlation of VES data with nearby Borehole Log



(b) Correlation of VES data with nearby Borehole Log

Fig. 11 a, b Correlation of VES data with nearby borehole log

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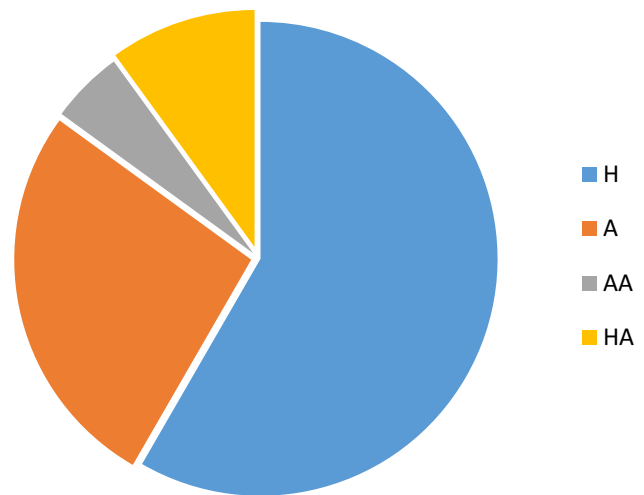


Fig. 12 Chart showing the curve distribution

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