#### **ORIGINAL ARTICLE**



# Geochemical characterization and evaluation of groundwater suitability for domestic and agricultural utility in semi-arid region of Basara, Telangana State, South India

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#### **Abstract**

Hydrogeochemical investigations were carried out in semi-arid region of Basara to estimate the quality of groundwater for its suitability for domestic and agricultural purposes. For this region 34 groundwater samples were collected in different locations and analyzed for various ions, viz., Na<sup>+</sup>, Ca<sup>2+</sup>, Mg<sup>2+</sup>, K<sup>+</sup>, Cl<sup>-</sup>, HCO<sub>3</sub><sup>-</sup>, SO<sub>4</sub><sup>2-</sup>, CO<sub>3</sub><sup>2-</sup>, HCO<sub>3</sub><sup>-</sup>, NO<sub>3</sub><sup>-</sup> and F<sup>-</sup> to assess the water chemistry with sodium absorption ratio, %Na, residual sodium carbonate, magnesium hazard. The nitrate and fluoride concentrations were above the maximum permissible limit, while calcium, sodium, potassium and chloride were found below the desirable limits in most of the groundwater samples. The Wilcox diagram illustrates that 59% of the samples belong to excellent to good category, while the US Salinity Laboratory diagram indicates medium salinity/low sodium content in 64.70% of samples. In general, the geochemistry of groundwater in Basara region is influenced by the water rock processes through percolation and dissolution of rock forming minerals, while calculated values of saturation index for Anhydrite, Aragonite, Artinite, Brucite, Calcite, Fluorite, Gypsum, Dolomite and Magnesite of the groundwater samples were less than zero, indicating under-saturation. Chadha rectangular diagram for geochemical classification and hydrochemical processes of groundwater for Basara provinces indicates 50% of Na<sup>+</sup>–Cl<sup>-</sup>, 29% of Ca<sup>2+</sup>–Mg<sup>2+</sup>–Cl<sup>-</sup> and 18% of the water samples concentrate in the category of Na<sup>+</sup>–HCO<sub>3</sub><sup>-</sup> type.

 $\textbf{Keywords} \;\; \text{Hydrochemistry} \cdot \text{Groundwater quality} \cdot \text{SAR} \cdot \% \text{Na} \cdot \text{RSC} \cdot \text{Basara} \cdot \text{South India}$ 

#### Introduction

In recent times, there has been a tremendous increase in demand for freshwater and water shortage in arid and semiarid regions due to population increase, urbanization, industrialization and intense agricultural activities in many parts of the world (Raju et al. 2011). Due to inadequate supply of surface waters, most of the places in India are depending mainly on groundwater resources for drinking and domestic, industrial, and irrigation uses. Hence, safe drinking water is a basic need of human development, health and well-being

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and hence, an internationally accepted human right (WHO 2011). Moreover, water has been viewed as an infinite and bountiful resource; water today defines human, social and economic development. The alarming rate of population growth, evolving industrial society, advances in technology, and the existing trend of depletion of groundwater resource has raised some serious environmental problems. The quality of drinking water has increasingly been questioned from health point of view for many decades (Vinod et al. 2009). Hence, knowledge on hydrochemistry of freshwater is important to assess the quality of groundwater in especially rural areas that influences the suitability of groundwater for domestic, irrigation, and industrial needs. For this reason, groundwater is a very precious natural resource. The severity of environmental problems related to groundwater varies from place to place depending on the geology, hydrologic, climatic conditions and geochemical factors that influence. However, groundwater quality studies with reference to drinking and irrigation purposes in different regions have been carried out, viz., Anantapur (Gowd 2005), Tamil Nadu



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(Subramani et al. 2005), Uttarakhand (Jain et al. 2010), Punjab (Kumar et al. 2007), Sri Lanka (Nishanthiny et al. 2010), Iran (Aghazadeh and Mogaddam 2010; Narany et al. 2015), Nigeria (Ackah et al. 2011), and China (Zhang et al. 2012).

Groundwater in India accounts for about 80% of domestic water requirements and more than 45% of total agricultural water, irrigating 39 million hectares (Kumar et al. 2005). Groundwater is also the single largest and most productive source of water for irrigation in India (Kinzelbach et al. 2003). Mikayilov and Acar 1998 has concluded that the relationship between agricultural practices and groundwater pollution was established. Earlier studies have identified different types of groundwater pollution like geogenic or anthropogenic source, major ions, nitrogen pollutants, trace elements, agricultural pesticides, dissolution or weathering of the rocks and soil (Adimalla and Venkatayogi 2017; Narsimha and Sudarshan 2017a, b; Narsimha and Sudarshan 2013; Ding and Zhang 2002; Islam and Singhal 2004; Elhatip et al. 2003; Kara et al. 2004; Krishna and Govil 2004; Narsimha et al. 2013a; 2013b; McBride 2004; Guo et al. 2007).

The present study area is located in the southwestern part of the Adilabad district and forms part of the Survey of India toposheet 56 F/13. The area lies between the North latitudes 18° 51' and 18° 59' and East longitudes 77° 54' and 77° 59' (Fig. 1) covering an area of 186 km<sup>2</sup>. The study area covers main villages Basara, Mudhol, Melapur, Bederelli, Aljapur, Yamcha, Pedda Kirgul, Chinna Kirgul, Oni, Ruvi, Shetpalli, Bamangaon, Gannora, Vitholi, Vitholi Tanda and Karegaon. The Basara region occupied by Archaean crystalline rocks is represented by pink and gray granites. These granites are composed of quartz, feldspars has major minerals, biotite, fluorite and apatite as accessory minerals (Narsimha and Sudarshan 2017b). The mean daily minimum and maximum temperature is 15 and 29 °C during December and 28 and 46 °C during May. The normal annual rainfall in the district is 1157 mm (Narsimha and Sudarshan 2017b). The groundwater is the major source of water for domestic, drinking, and agricultural needs of the area. As yet, the geochemistry of groundwater and its appropriateness for drinking and agricultural purposes in the semi-arid region of Basara have not been studied in great detail. For this reason, a detailed hydrogeochemical study and classification of the groundwater was carried out to evaluate their suitability

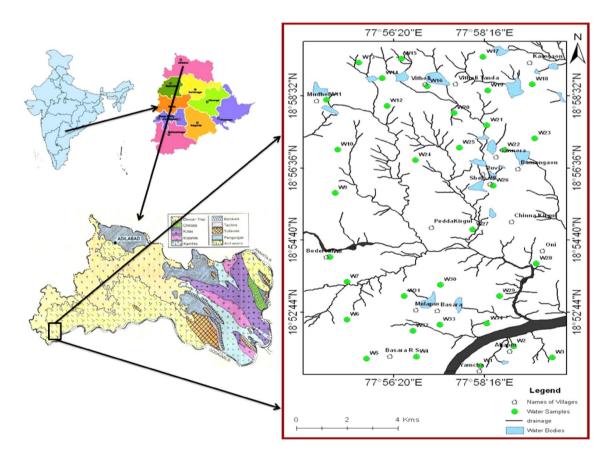


Fig. 1 Location map of the Basara area and its groundwater locations along with drainage pattern from Adilabad district, Telangana State, South India



for various purposes, viz., domestic/drinking water and irrigation/agricultural purposes.

#### Materials and methods

Systematic groundwater sampling was carried out in 34 representative bore wells from the Basara region in Adilabad District to get the baseline information regarding its quality (Fig. 1). The entire study area has only bore wells ranging from 30 to 60 m depth. The samples were collected after 10 min of pumping and stored in polyethylene bottles. Immediately after sampling, pH, electrical conductivity (EC), and total dissolved solids (TDS)

were measured in the field itself. Other major parameters magnesium (Mg<sup>2+</sup>), calcium (Ca<sup>2+</sup>), sodium (Na<sup>+</sup>), potassium (K<sup>+</sup>), chloride (Cl<sup>-</sup>), carbonate (CO<sub>3</sub><sup>2-</sup>), bicarbonate (HCO<sub>3</sub><sup>-</sup>), fluoride (F<sup>-</sup>), nitrate (NO<sub>3</sub><sup>-</sup>) and sulfate (SO<sub>4</sub><sup>2-</sup>) were analyzed according to the standard procedures described in standard methods for the examination of groundwater (APHA 1999; Table 1). The groundwater samples were collected and analysis was completed within 72 h in the Hydrogeochemical lab at Applied Geochemistry Department, Osmania University, Hyderabad. In addition to this, US Salinity Laboratory diagram, Chadha diagram and Wilcox diagram were employed to classify and determine the suitability of groundwater for irrigation by correlating sodium absorption ratio/electrical conductivity and percent sodium/electrical conductivity, respectively (Table 1). Eventually saturation indices for Anhydrite, Aragonite, Artinite, Brucite, Calcite, Fluorite, Gypsum,

Table 1 Instrumental, titrimetric and calculation methods were used for chemical analysis of groundwater samples from Basara region, Telangana State

Parameters	Characteristics	Analytical method	Reagents	Unit	References
General	pН	pH/EC/TDS meter	pH 4, 7 and 9.2	_	APHA (1999)
	Electrical conductivity	pH/EC/TDS meter	Potassium chloride	μS/cm	APHA (1999)
	Total dissolved solids (TDS)	Calculation	EC X (0.55-0.75)	mg/L	Hem (1991)
	Total hardness (as CaCO <sub>3</sub> )	EDTA titrimetric	EDTA, ammonia buffer and Eriochrome Black-T (EBT) indicator	mg/L	APHA (1999)
Major cations	Calcium (as Ca <sup>2+</sup> )	EDTA titrimetric	EDTA, sodium hydroxide and murexide	mg/L	APHA (1999)
	Magnesium (as Mg <sup>2+</sup> )	Calculation	MgH = TH-CaH; Mg = MgH X Eq.Wt of $Mg \times Normality of EDTA$	mg/L	APHA (1999)
	Sodium (as Na <sup>+</sup> )	Flame photometric	Sodium chloride (NaCl) and KCl	mg/L	APHA (1999)
	Potassium (as K <sup>+</sup> )	Flame photometric	NaCl and KCl	mg/L	APHA (1999)
Major anions	Bicarbonates (HCO <sub>3</sub> <sup>-</sup> )	Titrimetric	Hydrosulfuric acid (H <sub>2</sub> SO <sub>4</sub> ), phenolphthalein and methyl orange	mg/L	APHA (1999)
	Chloride (Cl <sup>-</sup> )	Titrimetric	Silver nitrate (AgNO <sub>3</sub> ), potassium chromate	mg/L	APHA (1999)
	Sulfates (SO <sub>4</sub> <sup>2-</sup> )	UV-visible spectrophotometer	HCl, ethyl alcohol, NaCl, barium chloride, sodium sulfate	mg/L	АРНА (1999)
	Nitrate (NO <sub>3</sub> <sup>-</sup> )	UV-visible spectrophotometer	Potassium nitrate (KNO3), Phenol disulfonic acid, ammonia	mg/L	APHA (1999)
	Fluoride (F <sup>-</sup> )	ISE (ion selective electrode; Thermo Orion)	TISAB III and NaF	mg/L	APHA (1999)
Irrigation water	Sodium absorption ratio (SAR)	$\frac{Na^{+}}{\sqrt{\left(Ca^{2^{+}}+Mg^{2^{+}}\right)/2}}$	-	meq/L	Richards (1954)
	Residual sodium carbonate (RSC)	$(CO_3^- + HCO_3^-) - (Ca^{2+} + Mg^{2+})$	-	meq/L	Eaton (1950)
	Percent sodium (%Na)	$\frac{Na^{+}+K^{+}}{\left(Ca^{2^{+}}+Mg^{2^{+}}+Na^{+}+K^{+}\right)} \times 100$	-	%	Wilcox (1955)
	Magnesium hazard	$Mg^{2+} \times 100/(Ca^{2+} + Mg^{2+})$	_	meq/L	Szabolcs and Darab (1964)



Dolomite and Magnesite of the groundwater samples are obtained using Visual MINTEQ (3.1) interactive.

#### **Results and discussion**

The pH of water is a very important indication of its quality and provides important piece of information regarding types of geochemical equilibrium (Hem 1985). The pH of the analyzed samples varies from 6.47 to 8.03 with a mean value of 7.09 in the study area, indicating alkaline nature (Tables 2, 3). pH values of all the collected samples are well within the safe limit as prescribed by

WHO (1990). Though pH has no direct effect on human health, all biochemical reactions are sensitive to variation of pH (Subba Rao and Krishna Rao 1991). The electric conductivity (EC) values for study area ranges from 235 to 2118  $\mu$ S/cm at 25 °C with an average value of 595.24  $\mu$ S/cm (Tables 2, 3). The higher EC may be attributed to high salinity and high mineral content at the sampling site (Vinod et al. 2009). It depends upon temperature, concentration and type of ions present in groundwater and also concentration of EC increases with the concentration of TDS. Total dissolved solids (TDS) in water comprise all inorganic salts including carbonate, bicarbonate, chloride, fluoride, sulfate, phosphate, nitrate, calcium,

Table 2 Analytical data for the groundwater samples from the study area (EC in μS/cm, remaining mg/L, except pH)

Sample ID	Longitude	Latitude	pН	EC	TDS	TH	CO <sub>3</sub> <sup>2-</sup>	HCO <sub>3</sub>	Cl-	SO <sub>4</sub> <sup>2-</sup>	NO <sub>3</sub>	Ca <sup>2+</sup>	Mg <sup>2+</sup>	Na <sup>+</sup>	K <sup>+</sup>	F <sup>-</sup>
BGW-1	77° 58′ 11″	18° 51′ 17″	6.61	392	251	160	0	85	128	42.5	22	42	13	76	4	0.37
BGW-2	77° 58′ 53″	18° 51′ 42″	6.61	863	552	325	0	152	340	75	60	104	16	122	25	0.57
BGW-3	77° 59′ 42″	18° 51′ 31″	7.02	314	201	150	0	110	64	25	17.5	38	13	47	3	0.74
BGW-4	77° 56′ 49″	18° 51′ 32″	6.57	235	150	115	0	85	50	5	4	30	9	15	8	0.46
BGW-5	77° 55′ 46″	18° 51′ 29″	7.24	549	351	40	21	140	96	75	38	10	4	154	2	2.2
BGW-6	77° 55′ 21″	18° 52′ 32″	7.61	863	552	80	18	350	113	150	40	20	7	250	3	2.22
BGW-7	77° 55′ 21″	18° 53′ 33″	7.52	2039	1305	650	12	134	638	375	24	105	10	420	4	0.91
BGW-8	77° 54′ 59″	18° 54′ 13″	6.47	2118	1355	850	0	98	1010	300	80	339	1	370	4	0.06
BGW-9	77° 55′ 6″	18° 55′ 56″	6.77	1177	753	500	33	128	110	250	38	125	22	120	5	0.48
BGW-10	77° 55′ 9″	18° 57′ 6″	7.62	392	251	35	12	135	53	120	42	10	2	161	4	2.56
BGW-11	77° 54′ 55″	18° 58′ 27″	8.03	471	301	85	15	128	50	80	40	8	16	121	4	2.48
BGW-12	77° 56′ 12″	18° 58′ 16″	7.3	392	251	115	18	125	89	50	35	38	5	118	4	1.19
BGW-13	77° 55′ 36″	18° 59′ 26″	6.9	471	301	160	15	98	124	50	15	36	17	102	1	1.12
BGW-14	77° 56′ 6″	18° 59′ 1″	7.53	314	201	110	15	120	43	7.5	32	26	11	63	2	1.13
BGW-15	77° 56′ 30″	18° 59′ 32″	7.62	314	201	150	12	110	74	12.5	13	40	12	49	2	0.96
BGW-16	77° 57′ 3″	18° 58′ 48″	6.78	549	351	250	18	49	266	17.5	60	66	21	95	2	0.94
BGW-17	77° 58′ 14″	18° 59′ 35″	6.97	314	201	85	15	61	89	7.5	4	32	1	46	2	0.54
BGW-18	77° 59′ 17″	18° 58′ 51″	7.27	314	201	35	15	110	64	2.5	130	10	2	119	2	4.33
BGW-19	77° 58′ 20″	18° 58′ 41″	6.52	549	351	270	15	61	227	25	0.8	66	25	82	3	0.43
BGW-20	77° 57′ 37″	18° 58′ 5″	7.42	471	301	75	18	125	96	27.5	42	18	7	129	3	1.46
BGW-21	77° 58′ 19″	18° 57′ 45″	7.07	314	201	95	18	79	42	5	56	24	8	66	3	1.43
BGW-22	77° 58′ 41″	18° 57′ 5″	7.28	471	301	150	21	103	120	37.5	35	20	24	110	3	2.44
BGW-23	77° 58′ 80″	18° 56′ 85″	6.57	627	401	235	9	67	287	20	56	68	16	120	7	0.27
BGW-24	77° 56′ 48″	18° 56′ 49″	7.16	863	552	210	24	116	467	50	60	36	29	312	4	2.03
BGW-25	77° 57′ 44″	18° 57′ 9″	7.21	471	301	90	30	120	89	51	48	28	4	139	4	1.25
BGW-26	77° 58′ 27″	18° 56′ 8″	7.18	314	201	125	15	85	85	15	2	38	7	61	4	0.76
BGW-27	77° 58′ 1″	18° 54′ 57″	6.89	392	251	185	12	134	74	10	27	38	21	43	4	0.86
BGW-28	77° 59′ 22″	18° 54′ 2″	6.91	706	452	160	15	165	227	32.5	35	42	13	169	5	0.78
BGW-29	77° 58′ 35″	18° 53′ 10″	7.15	627	401	210	30	183	252	25	32	85	6	120	39	0.26
BGW-30	77° 57′ 20″	18° 53′ 28″	6.84	392	251	170	18	104	156	37.5	8	50	10	92	4	0.75
BGW-31	77° 56′ 34″	18° 53′ 10″	6.71	549	251	315	24	122	308	17.5	60	102	14	94	16	0.71
BGW-32	77° 56′ 45″	18° 52′ 14″	7.73	235	150	95	30	73	89	27.5	25	22	10	89	5	0.54
BGW-33	77° 57′ 19″	18° 52′ 24″	6.68	627	401	310	0	98	234	50	18	104	12	107	5	0.11
BGW-34	77° 58′ 19″	18° 52′ 26″	7.44	549	351	120	18	250	110	50	42	18	18	191	4	1.02

BGW Basara groundwater



Table 3 Concentrations of ions and their comparison with the WHO

Chemical parameters	Expressed	Minimum	Maximum	WHO (1990)		Sample numbers exceeding	
				Desirable limit	Permissible limit	permissible limit	
pH	Units	6.47	8.03	6.5-8.5	9.2	Nil	
EC	μS/cm	235	2118	750	1500	6, 22	
TDS	mg/L	150	1355	500	1500	Nil	
TH	mg/L	35	850	100	500	7, 8, 9	
HCO <sub>3</sub> <sup>-</sup>	mg/L	49	350	_	_	_	
$CO_3^{2-}$	mg/L	0	33	_	_	_	
Cl-	mg/L	42	1010	250	600	7, 8	
$SO_4^{2-}$	mg/L	200	600	2.5	375	Nil	
NO <sub>3</sub>	mg/L	0.8	130	_	45	2, 8, 16, 18, 21, 23, 24, 25, 31	
F <sup>-</sup>	mg/L	0.06	4.33	0.5	1.5	5, 6, 10, 11, 18, 22, 24	
Na <sup>+</sup>	mg/L	15	420	50	200	6, 7, 8, 24	
$K^+$	mg/L	1	39	_	_	_	
Ca <sup>2+</sup>	mg/L	8	339	75	200	8	
$Mg^{2+}$	mg/L	1	29	30	150	Nil	

magnesium, sodium, and potassium (Sawyer 1994). Total dissolved solids of the groundwater, in the study area, vary from 150 to 1355 mg/L with an average of 377.82 mg/L (Tables 2, 3). The groundwater of the study area has been

classified based on TDS values, according to the procedure suggested by US Geological Survey 2000. It is clear from Table 4 that 94.11% of groundwater samples fall into freshwater category, whereas 5.88% of groundwater

Table 4 Classification of groundwater based on irrigational water quality parameters

Parameter	Classification	Range	Number of samples	% of samples	References
TDS	Fresh water	< 1000	32	94.11	US Geological Survey (2000)
	Slightly saline	1000-3000	2	5.88	
	Moderately saline	3000-10,000	Nil	Nil	
	High saline	10,000-35,000	Nil	Nil	
TH	Safe	< 75	3	8.82	Sawyer et al. (2003)
	Moderate-hard	75–150	12	35.29	
	Hard	150-300	13	38.23	
	Very hard	> 300	6	17.64	
Salinity hazard	Excellent (C1)	Up to 250	2	5.88	Richards (1954), Todd (1980)
	Good (C2)	250-750	27	79.41	
	Fail/medium (C3)	750-2250	5	14.7	
	Poor/bad (C4)	> 2250	Nil	Nil	
Alkali hazard	Excellent (S1)	Up to 10	31	91.17	Richards (1954), Todd (1980)
	Good (S2)	10–18	3	8.82	
	Fail/medium (S3)	18–26	Nil	Nil	
	Poor/bad (S4 & S5)	> 26	Nil	Nil	
Percent sodium (%Na)	Safe	< 60	20	59	Eaton (1950)
	Unsafe	> 60	14	41	
RSC	Good	< 1.25	31	91	Richards (1954)
	Doubtful	1.25-2.50	9	9	
	Unsuitable	> 2.50	Nil	Nil	
Magnesium hazard	Suitable	< 50	30	88	Szabolcs and Darab (1964)
	Unsuitable	> 50	4	12	



samples fall into slightly saline category (Table 4). The acceptable limit of total hardness (TH) (as CaCO<sub>3</sub>) is 300 mg/L, which can be extended up to 600 mg/L in case of non-availability of any alternate water source (WHO 1990). The total hardness as CaCO<sub>3</sub> equivalents in the study area ranges from 35 to 850 mg/L with mean value of 197.35 mg/L (Tables 2, 3). Sawyer et al. (2003) classified groundwater is given in Table 4, according to which 35.29% of samples belongs to moderate—hard type, and 38.23 and 17.64% of water belong to hard and very hard type, respectively.

## Carbonate (CO<sub>3</sub><sup>2-</sup>) and bicarbonate (HCO<sub>3</sub><sup>-</sup>)

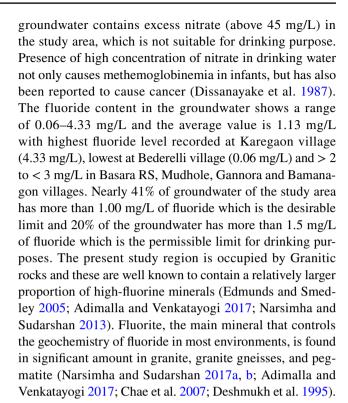
Carbonate ( $CO_3^{2-}$ ) and bicarbonate ( $HCO_3^{-}$ ) in the ground-water is varying from 0 to 33 mg/L and 43 to 183 mg/L with an average value of 15.17 and 104.44 mg/L, respectively (Tables 2, 3). The content of  $HCO_3^{-}$  has no known adverse health effects and all groundwater samples lie within the desirable limit of 300 mg/L (WHO 1990; Table 3).

## Chloride (CI<sup>-</sup>) and sulfate (SO<sub>4</sub><sup>2-</sup>)

The chloride concentration in the groundwater samples vary from 42 to 1010 mg/L with an average value of 182.06 mg/L (Tables 2, 3). The desirable limit of chloride in potable water is 250 mg/L and the permissible limit is 1000 mg/L (WHO 1990), while 23.52% of the groundwater samples exceed the desirable limit, only one sample in the study area exceed the permissible limit. Excess of Cl<sup>-</sup> in the water is usually taken as an index of pollution and considered as tracer for groundwater contamination (Loizidou and Kapetanios 1993). Shand (1952) indicated that residual water contained in pores of granites or within the crystals of rocks may contain chloride. Chloride is higher (1010 mg/L in Bederelli village) due to leaching of upper soil layers by industrial and domestic activities and dry climates; furthermore, higher concentration of Cl<sup>-</sup> in drinking water causes a salty taste and has a laxative effect in people not accustomed to it (Bhardwaj and Singh 2011). The sulfate concentration of the groundwater samples of the study area ranges from 2.5 to 375 mg/L with an average value of 56.985 mg/L, which is below the permissible limit of 400 mg/L (WHO 1990).

#### Nitrate (NO<sub>3</sub><sup>-</sup>) and fluoride (F<sup>-</sup>)

Nitrate contamination in groundwater is one of the major issues in water quality studies (Schilling and Wolter 2007; Narsimha and Sudarshan 2017a). Nitrate concentration in the groundwater of the study area varies from 0.4 to 80 mg/L with mean 22.068 mg/L (Tables 2, 3). The highest nitrate level is recorded at Bederelli village (80 mg/L), Aljapur, Vitholi thanda and near Basara (60 mg/L), 20.58% of the

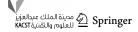


#### Sodium (Na<sup>+</sup>) and potassium (K<sup>+</sup>)

Sodium is one of the most important constituents of groundwater because its higher concentration, more than 200 mg/L, makes the water unsuitable for domestic use and causes severe health problems like hypertension, congenial diseases, kidney disorders and nervous disorders in human body (Vinod et al. 2009; Narsimha and Sudarshan 2017a). Sodium concentration in groundwater ranges from 15 to 810 mg/L with an average of 146 mg/L (Tables 2, 3). According to WHO (1990) guidelines, the maximum admissible limit is 200 mg/L and 15% of the samples are above this limit. The concentration of potassium in natural water is generally less than 10 mg/L (WHO 1990). The potassium concentration in groundwater in the study area varies from 1 to 39 mg/L with an average value of 5.71 mg/L. The permissible limit of K<sup>+</sup> for drinking water is specified as 12 mg/L as per WHO (1990) and 8.82% of the samples are above this limit.

### Calcium (Ca<sup>2+</sup>) and magnesium (Mg<sup>2+</sup>)

Calcium in the groundwater of the study area is varying from 8 to 339 mg/L with an average value of 59 mg/L (Tables 2, 3). The presence of large amount of carbon dioxide may increase the solubility of calcium up to 200–300 mg/L in the presence of bicarbonate (Hem 1985; Karanth 1989). The maximum admissible limit for calcium is 200 mg/L (WHO 1990). Only 6% of groundwater of the study area has Ca<sup>2+</sup>



concentration above the permissible limit suggested by WHO (1990) (Table 3). Magnesium in the groundwater of the study area is varying from 1 to 29 mg/L and the average value is 11.94 mg/L (Tables 2, 3). The required permissible limit of magnesium in groundwater for drinking purpose is 30 mg/L (WHO 1990) and the concentrations are found to be within the permissible limits (Tables 2, 3).

# Geochemical diagrams/classification of groundwater

Hydrochemical diagrams are aimed at facilitating interpretation of evolutionary trends, particularly in groundwater systems, when they are interpreted in conjunction with distribution maps and hydrochemical sections. For these purposes, Chadha (1999), Gibbs (1970), Soltan (1999) and hydrogeochemical saturation index states have been emphasized to better understand the groundwater system.

#### Chadha diagram

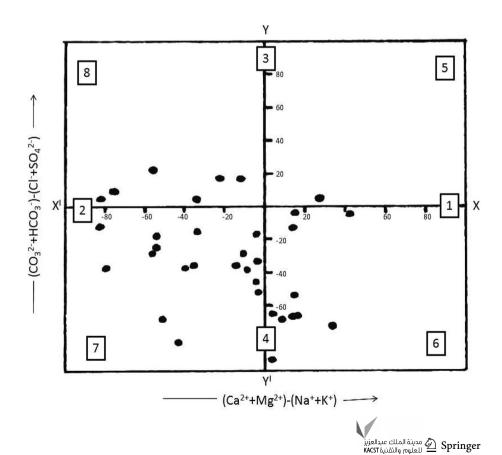
The term "hydrochemical facies" is used to describe occurrence modes of groundwater in an aquifer that differs in their chemical composition. Classification of geochemical facies and interpretation of chemical data of groundwater can be classified on the basis of dominant ions using the Chadha rectangular diagram (Chadha 1999). In the proposed diagram (Fig. 2), the differences in milliequivalent percentage

between alkaline earths ( $Ca^{2+} + Mg^{2+}$ ) and alkali metals  $(Na^+ + K^+)$ , expressed as percentage reacting values, are plotted on the x-axis, and the differences in milliequivalent percentage between weak acidic anions  $(CO_3^{2-} + HCO_3^{-})$ and strong acidic anions  $(Cl^- + SO_4^{2-})$  are plotted on the y-axis. Results of chemical analyses were plotted on the proposed diagram to test its applicability for geochemical classification of groundwater and to understand the hydrochemical processes shown in Fig. 2. The majority of groundwater samples (50%, Fig. 2) in the study area fall under the subfield of Na<sup>+</sup>-Cl<sup>-</sup> water type with salinity problems both in irrigation and domestic uses. The 29% of the groundwater samples fall under the subfield of alkaline earths exceeding alkali metals and strong acidic anions exceeding weak acidic anions (Ca<sup>2+</sup>-Mg<sup>2+</sup>-Cl<sup>-</sup> type), such waters have permanent hardness. The remaining 18% of the water samples concentrate in the category of Na<sup>+</sup>-HCO<sub>3</sub><sup>-</sup> type (Fig. 2); such waters release residual sodium carbonate while irrigating and cause foaming problems when used for domestic purposes. Only 3% of groundwater samples fell in alkaline earths and weak acidic anions that exceed both alkali metals and strong acidic anions (Ca<sup>2+</sup>-Mg<sup>2+</sup>-HCO<sub>3</sub><sup>-</sup> type) such waters may have temporary hardness.

#### Gibbs diagram

Gibbs diagrams for cations  $(Na^+ + K^+)/(Na^+ + K^+ + Ca^{2+})$  and anions  $(Cl^-)/(Cl^- + HCO_3^-)$  of the groundwater samples

Fig. 2 Geochemical classification and hydrochemical parameters of groundwater (after Chadha 1999). 1. Alkaline earths exceed alkali metals; 2. alkali metals exceed alkaline earths; 3. weak acidic anions exceed strong acidic anions; 4. strong acidic anions exceed weak acidic anions; 5. Ca<sup>2+</sup>-Mg<sup>2+</sup>–HCO<sub>3</sub> water type with temporary hardness; 6. Ca<sup>2+</sup>-Mg<sup>2+</sup>–Cl<sup>-</sup> water type with permanent hardness; 7. Na<sup>+</sup>-Cl<sup>-</sup> water type with salinity problems both in irrigation and domestic uses; 8. Na<sup>+</sup>-HCO<sub>3</sub> water type causes foaming problems in domestic use



**Fig. 3** Mechanism controlling groundwater chemistry of the Basara  $\triangleright$  region (Gibbs 1970). **a** TDS vs (Na<sup>+</sup> + K<sup>+</sup>):(Na<sup>+</sup> + K<sup>+</sup> + Ca<sup>2+</sup>) and **b** TDS vs (Cl<sup>-</sup>:Cl<sup>-</sup> + HCO<sub>3</sub><sup>-</sup>) were showing rock, evaporation and precipitation dominance

were plotted separately against TDS (Gibbs 1970; Ramesam and Barua 1973; Adimalla and Venkatayogi 2017; Viswanathaiah et al. 1978; Narsimha and Sudarshan 2017a, b) to assess the rock dominance, evaporation dominance and precipitation dominance in groundwater chemistry (Fig. 3). The present study reveals that the groundwater chemistry majorly influenced by the rock dominance, which indicates that the foremost mechanism of weathering of rocks forming minerals due to the entire province of occupied by granitic gneisses. Eventually, few groundwater locations were inclined towards evaporation which indicates influence of anthropogenic activity is prevalent, leading to increase in concentration of sodium, chloride and also TDS (Adimalla and Venkatayogi 2017; Narsimha and Sudarshan 2017b).

#### **Soltan classification**

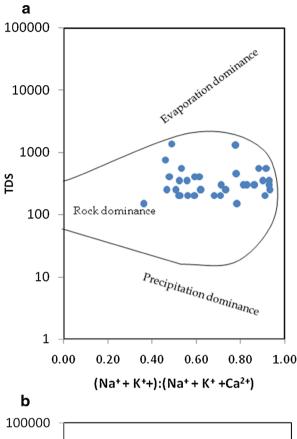
Soltan (1999) has clearly classified groundwater into two types, which are base-exchange indices (r1) and meteoric genesis indices (r2) as shown below.

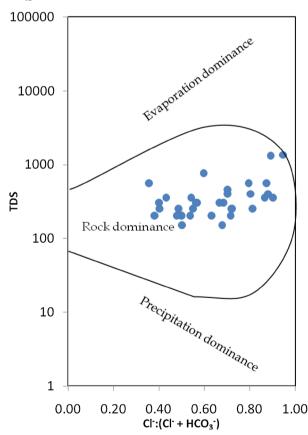
$$r1 = (Na^+ - Cl^-)/SO_4^{2-}$$

 $r2 = [(K^+ + Na^+) - Cl^-]/SO_4^{\ 2-}$ , where Na<sup>+</sup>, K<sup>+</sup>, Cl<sup>-</sup> and  $SO_4^{\ 2-}$  concentrations are expressed in meq/L. If r1 < 1 and r2 < 1, the groundwater sources are of Na<sup>+</sup>– $SO_4^{\ 2-}$  and deep meteoric type, respectively, while r1 > 1 and r2 > 1 indicates the sources are of Na<sup>+</sup>– $HCO_3^-$  and shallow meteoric type, respectively, and also presented in Table 5. Based on Soltan classification 59% of groundwater belongs to Na<sup>+</sup>– $SO_4^{\ 2-}$ , rest of the samples is of Na<sup>+</sup>– $HCO_3^-$ , while 56% of groundwater belongs to deep meteoric and 44% of are of shallow meteoric type.

#### **Hydrogeochemical saturation index states**

The saturation state of minerals in the groundwater can be expressed by the saturation index (SI). The saturation index (SI) is vital to envisage the subsurface mineralogy from groundwater data without collecting the samples of the solid phase and analyzing the mineralogy (Deutsch 1997). Moreover, Frenkel (1984) has concluded that arid and semi-arid regions of surface and subsurface water was nearly saturated with respect to CaCO<sub>3</sub>, and also contains high concentrations of sulfate and calcium which could be precipitated in the soil. Visual MINTEQ (3.1) has widely been used to estimate the saturation index (SI) of the groundwater with respect to mineral phases (Gustafsson 2012; Nagaraju et al. 2014; Appelo and Postma 1996). The solubility products of several minerals which







**Table 5** Groundwater classification according to base-exchange (r1) and meteoric genesis index (r2) criteria

Sample no.	<u>r</u> 1	Water type	r2	Water type
BGW-1	- 0.34	Na <sup>+</sup> -SO <sub>4</sub> <sup>2-</sup>	- 0.23	Deep meteoric
BGW-2	-2.74	$Na^{+}-SO_{4}^{2-}$	- 2.33	Deep meteoric
BGW-3	0.46	Na <sup>+</sup> -SO <sub>4</sub> <sup>2-</sup>	0.61	Deep meteoric
BGW-4	- 7.28	Na <sup>+</sup> -SO <sub>4</sub> <sup>2-</sup>	- 5.32	Deep meteoric
BGW-5	2.56	Na <sup>+</sup> –HCO <sub>3</sub> <sup>-</sup>	2.59	Shallow meteoric
BGW-6	2.46	Na <sup>+</sup> –HCO <sub>3</sub> <sup>-</sup>	2.49	Shallow meteoric
BGW-7	0.03	Na <sup>+</sup> -SO <sub>4</sub> <sup>2-</sup>	0.05	Deep meteoric
BGW-8	- 1.98	Na <sup>+</sup> -SO <sub>4</sub> <sup>2-</sup>	- 1.97	Deep meteoric
BGW-9	0.41	Na <sup>+</sup> -SO <sub>4</sub> <sup>2-</sup>	0.43	Deep meteoric
BGW-10	2.20	Na <sup>+</sup> –HCO <sub>3</sub> <sup>-</sup>	2.25	Shallow meteoric
BGW-11	2.31	Na <sup>+</sup> -HCO <sub>3</sub>	2.37	Shallow meteoric
BGW-12	2.52	Na <sup>+</sup> –HCO <sub>3</sub> <sup>-</sup>	2.62	Shallow meteoric
BGW-13	0.90	Na <sup>+</sup> -SO <sub>4</sub> <sup>2-</sup>	0.93	Deep meteoric
BGW-14	9.78	Na <sup>+</sup> –HCO <sub>3</sub> <sup>-</sup>	10.11	Shallow meteoric
BGW-15	0.17	Na <sup>+</sup> -SO <sub>4</sub> <sup>2-</sup>	0.37	Deep meteoric
BGW-16	- 9.25	Na <sup>+</sup> -SO <sub>4</sub> <sup>2-</sup>	- 9.11	Deep meteoric
BGW-17	- 3.26	Na <sup>+</sup> -SO <sub>4</sub> <sup>2-</sup>	-2.94	Deep meteoric
BGW-18	64.77	Na <sup>+</sup> –HCO <sub>3</sub> <sup>-</sup>	65.75	Shallow meteoric
BGW-19	- 5.45	Na <sup>+</sup> -SO <sub>4</sub> <sup>2-</sup>	- 5.30	Deep meteoric
BGW-20	5.07	Na <sup>+</sup> –HCO <sub>3</sub> <sup>-</sup>	5.20	Shallow meteoric
BGW-21	16.20	Na <sup>+</sup> –HCO <sub>3</sub> <sup>-</sup>	16.93	Shallow meteoric
BGW-22	1.79	Na <sup>+</sup> –HCO <sub>3</sub> <sup>-</sup>	1.89	Shallow meteoric
BGW-23	- 6.91	Na <sup>+</sup> -SO <sub>4</sub> <sup>2-</sup>	- 6.48	Deep meteoric
BGW-24	0.38	Na <sup>+</sup> –SO <sub>4</sub> <sup>2–</sup>	0.48	Deep meteoric
BGW-25	3.33	Na <sup>+</sup> –HCO <sub>3</sub> <sup>-</sup>	3.43	Shallow meteoric
BGW-26	0.82	Na <sup>+</sup> -SO <sub>4</sub> <sup>2-</sup>	1.15	Shallow meteoric
BGW-27	- 1.04	Na <sup>+</sup> -SO <sub>4</sub> <sup>2-</sup>	-0.55	Deep meteoric
BGW-28	1.40	Na <sup>+</sup> –HCO <sub>3</sub> <sup>-</sup>	1.59	Shallow meteoric
BGW-29	- 3.63	Na <sup>+</sup> –SO <sub>4</sub> <sup>2–</sup>	- 1.71	Deep meteoric
BGW-30	-0.51	Na <sup>+</sup> -SO <sub>4</sub> <sup>2-</sup>	-0.38	Deep meteoric
BGW-31	- 12.62	Na <sup>+</sup> -SO <sub>4</sub> <sup>2-</sup>	- 11.50	Deep meteoric
BGW-32	2.38	Na <sup>+</sup> –HCO <sub>3</sub>	2.60	Shallow meteoric
BGW-33	- 1.87	Na <sup>+</sup> -SO <sub>4</sub> <sup>2-</sup>	- 1.75	Deep meteoric
BGW-34	5.00	Na <sup>+</sup> -HCO <sub>3</sub>	5.10	Shallow meteoric

might be expected to precipitate from irrigation waters are provided in supplementary material Table 1. If SI = 0, there is equilibrium between the mineral and the solution; SI < 0 reflects under saturation (dissolution) with respect to that particular mineral, and SI > 0 oversaturation. The calculated values of SI for Anhydrite, Aragonite, Artinite, Brucite, Calcite, Fluorite, Gypsum, Dolomite and Magnesite of the groundwater samples ranged from -4.13 to -1.08, -3.13 to -1.19, -12.34 to -8.84, -7.87 to -6.29, -2.99 to -1.05, -2.92 to -0.61, -3.88 to

**Table 6** Saturation indexes of minerals minimum and maximum for groundwater in the semi-arid region of Basara, Telangana State, South India

Minerals name	Minimum	Maximum
Anhydrite	- 4.13	- 1.08
Aragonite	- 3.13	- 1.19
Artinite	- 12.34	- 8.84
Brucite	- 7.87	- 6.29
Calcite	- 2.99	- 1.05
Dolomite	- 6.15	- 2.49
Epsomite	- 6.85	- 4.11
Fluorite	- 2.92	- 0.61
Gypsum	- 3.88	- 0.83
Halite	- 7.65	- 5.06
Huntite	- 16.83	- 9.72
Hydromagnesite	- 29.88	- 20.65
Lime	- 22.59	- 21.16
Magnesite	- 4.51	- 2.59
Mirabilite	- 9.77	- 5.29
Natron	- 13.34	- 9.42
Nesquehonite	- 7.30	- 5.38
Periclase	- 12.35	- 10.77
Portlandite	- 12.60	- 11.16
Thenardite	- 11.21	- 6.72
Thermonatrite	- 15.29	- 11.37
Vaterite	- 3.56	- 1.61

-0.83, -6.15 to -2.49 and -4.51 to -2.59, respectively (Table 6), which clearly shows all SI values were less than zero, indicating under-saturation (Fig. 4).

# Chemical characterization of hydrogeochemical data

Chemical characterization of groundwater schemes developed are mainly based on the concentrations of various predominant cations and anions or on the interrelationship of ions and also a number of techniques and methods have been developed to interpret the chemical data. In this study, the groundwater quality with respect to irrigation/agricultural is assessed by the following methods:

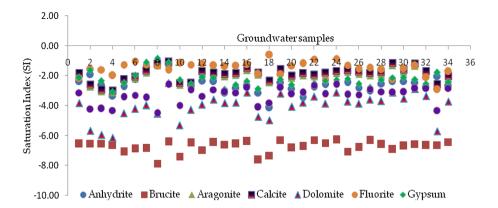
- 1. Sodium adsorption ratio (SAR).
- 2. Electrical conductivity and percent sodium (%Na).
- 3. Residual sodium carbonate (RSC).
- 4. Magnesium hazard.

#### Sodium adsorption ratio (SAR)

Excess sodium in water produces the undesirable effects of changing soil properties and reducing soil permeability.



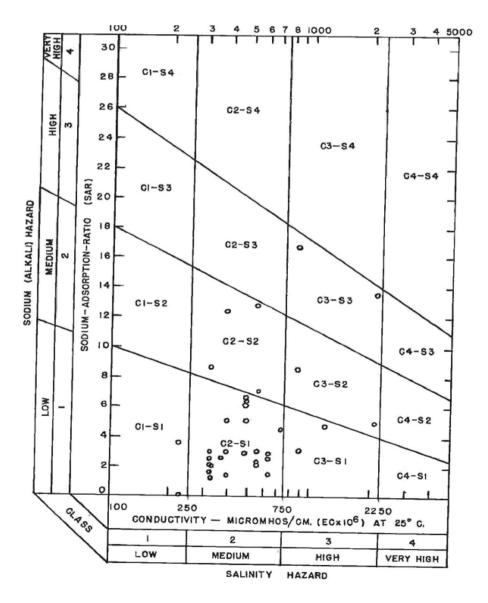
Fig. 4 Saturation indices of groundwater samples with respect to Anhydrite, Brucite, Aragonite, Calcite, Dolomite, Gypsum and Fluorite



Hence, for considering the suitability for irrigation the assessment of sodium concentration is essential. The sodium or alkali hazard in the use of water for irrigation is determined by the absolute and relative concentration of cations. The relative activity of sodium ion in the exchange reaction with soil

is expressed in terms of sodium adsorption ratio (SAR). The SAR, which indicates the effect of relative cation concentration on Na<sup>+</sup> accumulation in the soil, is used for evaluating the sodicity of irrigation water. The sodicity hazard of water is generally described by the SAR:

Fig. 5 Classification of irrigation water quality, with respect to salinity hazard and sodium hazard (after US Salinity Laboratory 1954)





$$SAR = \frac{Na}{\sqrt{(Ca + Mg)/2}}.$$

The plot of the analytical data on the US Salinity Laboratory (1954) in which the EC is taken as salinity hazard and SAR as an alkalinity hazard (Fig. 5) shows that 64.70% (Table 4) fall into the category of C2S1, indicating medium salinity and low alkalinity water which can be used for irrigating most of the soil and crops with little danger of exchangeable sodium. 12% of samples which falls into C2S2 class indicate medium salinity medium sodium hazard. 5.88% of samples which falls into C1S1 class indicate low salinity and low alkalinity water. It can be used for irrigation for most soil and crops with less negative impact (Richards 1954). 5.88% of the water samples falls into C3S1 class, which shows a high salinity hazard and low alkalinity hazard. However, 5.88% of sample falls into C3S2 type (high salinity and medium sodium hazards). These types of water may be used on coarse-textured or organic soils with good permeability (Karanth 1989). 5.88% of samples fall into C3S3 class (Fig. 5; high salinity and high sodium hazards).

#### Electrical conductivity and percent sodium

EC and Na concentrations are important in classifying irrigation water. The EC values vary from 235 to 2118 µS/cm (Tables 1, 2). High salt content (high EC) in irrigation water leads to formation of saline soil. Salinity, on the irrigated lands is the major cause of loss of production, and it has adverse environmental impacts on irrigation. Saline conditions severely limit the choice of crops and adversely affect crop germination and yields. It is important that all evaluations regarding irrigation water quality are linked to the evaluation of the soils to be irrigated. Sodium concentration is important in classifying irrigation water because sodium reacts with soil to reduce its permeability. In all natural waters, percent sodium is a parameter to evaluate its suitability for agricultural purposes (Wilcox 1955); sodium combining with carbonate forms alkaline soils, while sodium combining with chloride forms saline soils. Either type of sodium-enriched soil will support little or no plant growth (Todd 1980). The sodium percentage (%Na) in the study area ranges from 27.90 to 91.46. A high sodium percentage causes deflocculation and impairment of the tithe and the permeability of soils (Karanth 1989).

$$\% Na = \frac{Na^{+} + K^{+}}{(Ca^{2^{+}} + Mg^{2^{+}} + Na^{+} + K^{+})} \times 100.$$

The chemical quality of groundwater samples was studied by plotting analytical data (Wilcox 1955) relating EC and sodium percent (Fig. 6) that show that out of the 34 groundwater samples, 59% of the samples belong to excellent to good category followed by 6% of the samples belonging to good to permissible category, 26% belonging to permissible to doubtful, and 9% of the samples belonging to doubtful to unsuitable category for irrigation uses (Table 4). The agricultural yields are generally low in lands irrigated with waters belonging to doubtful to unsuitable category. This is probably due to the presence of excess sodium salts, which cause osmotic effects on soil-plant system.

#### Residual sodium carbonate (RSC)

Residual sodium carbonate (RSC) also influences the suitability of groundwater for irrigation uses. Residual sodium carbonate can be estimated by subtracting the quality of alkaline earths  $(Ca^{2+} + Mg^{2+})$  from the carbonate  $(CO_3^{2-} + HCO_3^{-})$ . When the sum of carbonates is in excess of calcium and magnesium, there may be a possibility of complete precipitation of calcium and magnesium (Ragunath 1987). As a result, the relative proportion of sodium in the water is increased in the form of sodium carbonate, and this excess, denoted by RSC, is calculated as follows (Eaton 1950; Ragunath 1987):

$$RSC = (CO_3^{2-} + HCO_3^{-}) - (Ca^{2+} + Mg^{2+}).$$

The concentration of ions is expressed in meg/L. According to the US Department of Agriculture, water having more than 2.50 meg/L of RSC is not suitable for irrigation purposes. Based on the RSC values 91.11% of groundwater samples fall under good category and remaining 8.82% fall under doubtful category (Table 4).

#### Magnesium hazard (MH)

Generally, calcium and magnesium maintain a state of equilibrium in most waters. Calcium and magnesium do not behave equally in the soil system; magnesium deteriorates soil structure particularly when water is sodium dominated and highly saline. A high level of Mg<sup>2+</sup> is usually due to the presence of exchangeable Na<sup>+</sup> in irrigated soils. In equilibrium, more Mg2+ present in water will adversely affect the soil quality, rendering it alkaline, resulting in decreased and adversely affected crop yields. Szabolcs and Darab (1964) proposed a magnesium hazard (MH) value for irrigation water and an MH is calculated by the following equation.

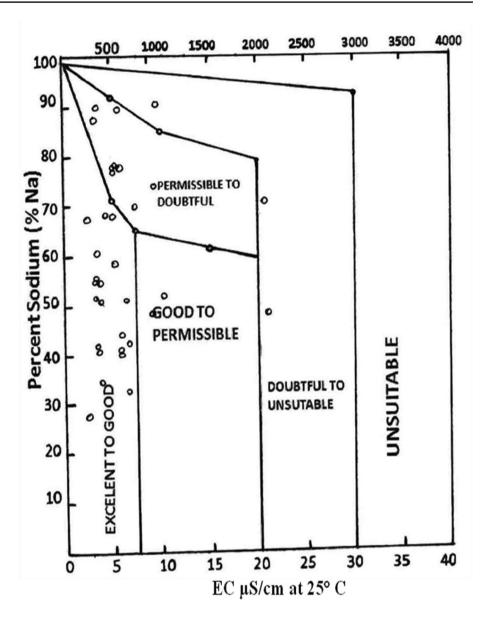
Magnesium hazard (MH) = 
$$Mg^{2+} \times 100 / (Ca^{2+} + Mg^{2+})$$
,

where all ions are expressed in meq/L. A magnesium ratio of more than 50 is considered to be harmful and unsuitable for irrigation use. This would be adversely affecting the crop yield, as soils become more alkaline. The magnesium hazard values of the study area range from 0.48 to 76.73 with an average value of 32.06 (Table 4). Majority



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Fig. 6 Classification of irrigation water quality with respect to electrical conductivity and percent sodium (after Wilcox 1955)



of groundwater samples (88%) are within magnesium hazard of 50 and remaining 12% of groundwater samples were above 50 which are considered as harmful and unsuitable for irrigation use.

#### **Conclusions**

The main aim of this study was to evaluate the groundwater quality for irrigation and domestic application in Basara region, Adilabad district, South India. The entire study area occupied by Archaean crystalline rocks is represented by pink and gray granites. Groundwater in the Basara region is mostly alkaline in nature. The hydrochemical analysis data revealed that the highest nitrate was in the Bederelli village (80 mg/L), Aljapur, Vitholi thanda and near Basara

(60 mg/L), while high fluoride was at Karegaon village (4.33 mg/L), Mudhole (2.5 mg/L), Gannora and Bamanagon (3.1 mg/L), lowest at Bederelli village (0.06 mg/L). Calcium, sodium, magnesium and sulfate were within the permissible limit, but chloride is higher at Bederelli village (1010 mg/L). However, 59% and 56% of groundwater samples were shown Na<sup>+</sup>-SO<sub>4</sub><sup>2-</sup> and deep meteoric signatures respectively, which were typically influenced by the groundwater chemistry. The interesting and most significant finding of this study was that almost all groundwater samples were under-saturated with respect to Anhydrite, Aragonite, Artinite, Brucite, Calcite, Fluorite, Gypsum, Dolomite, Epsomite, Halite, Hydromagnesite, Lime, Mirabilite, Natron, Nesquehonite, Periclase, Portlandite, Thenardite, Thermonatrite, Vaterite and Magnesite. Based on Gibbs, groundwater samples from the Basara region show



rock dominance indicating prevalent water—rock interaction through percolation is also wide and some of the samples are inclining towards evaporation. Most of the groundwater samples fall into the C2S1 dominance in US Salinity Laboratory diagram indicating medium salinity and low alkalinity water which is useful for irrigation in almost all types of soils. Na%, RSC and magnesium hazard reveal that in groundwater samples 59% belong to excellent to good category, 91.11% fall under good category and 88% are within magnesium hazard of 50, which indicates its suitability for irrigation in all types of soils, respectively.

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