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Assessment of radiological hazards and effective dose from natural radioactivity in rock samples of Hassan district, Karnataka, India

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Abstract

The natural primordial radionuclide activity concentrations in different types of rock samples in and around Hassan district were measured using gamma spectrometry-based high-purity Germanium detector. The average activity concentration of ²²⁶Ra, ²³²Th, and ⁴⁰K radionuclides in rock samples varies from 29.4 ± 1.2 to 83.7 ± 1.7 , 37.9 ± 1.2 to 198.4 ± 2.2 , and 346.1 ± 9.0 to 1024.4 ± 19.4 Bq kg⁻¹ with an average value of 50.9 ± 1.4 , 79.6 ± 1.5 , and 609.2 ± 12.8 Bq kg⁻¹, respectively. The radiological hazard indices such as radium equivalent activity, gamma index, alpha index, external and internal health hazard indices, absorbed dose rate and annual effective doses were estimated. The average value of radium equivalent activity is less than the criterion limit of 370 Bq kg⁻¹. The average indoor and outdoor annual effective dose values for rock samples are lower than the world average recommended value of 1 mSv y⁻¹. The study concludes that except granite rocks, all the rock samples analyzed are safe when used as construction materials and do not pose any significant radiation hazards.

Keywords Radionuclides · Health hazard indices · HPGe · Equivalent effective dose · Radium equivalent activity · Rocks

Introduction

The study of distribution of radionuclides in soil, rock and building materials is most important to know the health risk due to radiations coming from these materials (Damla et al. 2011). Major portion of the background radiation is coming from the primordial radionuclides such as ²²⁶Ra, ²³²Th, and ⁴⁰K which are present in the earth's crust (soil, rocks) and building materials. These radionuclides are not uniformly distributed; hence, the knowledge of their distribution in soil, rock, and building materials plays an important role in radiation protection and measurement (International Commission on Radiological Protection 1999). The radionuclides present in rocks are the sources of external and the internal radiation exposures in dwellings. The external exposure

J. Sannappa sannappaj2012@gmail.com is due to gamma radiations originated from terrestrial and extra-terrestrial sources. Internal exposure is due to the ingestion of ²²⁶Ra, ²³²Th, and the radionuclides produced by the decay of these as well as ⁴⁰K through dietary intake and inhalation of radioactive inert gases, radon (222Rn, a daughter product of ²²⁶Ra) and thoron (²²⁰Rn, a daughter product of ²²⁴Ra), and their short-lived products. The specific activities of ²²⁶Ra, ²³²Th, and ⁴⁰K in rocks and the building raw materials mainly depend on geological and geographical conditions as well as geochemical characteristics of those materials (UNSCEAR 1993). Weathering of the earth's crust is the ultimate mechanism for the release of primordial radionuclides into the soil, which constitutes the principal sources of natural background radiation (Rangaswamy et al. 2016). The activity concentrations of primordial radionuclides such as ²²⁶Ra, ²³²Th, and ⁴⁰K in the soil samples largely depend on the mineral compositions of rock from which the soil is originating as well as the type of formation and transfer processes that are involved (International Commission on Radiological Protection 1984; Ningappa et al. 2008). The radiological impact from the natural radioactivity is due to radiation exposure of the body by gamma rays and irradiation of lung tissues from inhalation of radon and its progeny. From the health risk point of view, it is necessary to know the dose limits of public exposure and to measure

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the natural environmental radiation level due to the distribution of radionuclides in ground, air, water, foods, building interiors, etc. (El-Taher 2010). Therefore, it is important to measure the activity concentration of radionuclides in rock samples used as building materials for assessing the radiological risks to human health and for the use and management of these rocks.

It is very important to estimate the radionuclide activity concentration in rock samples because the local rocks are raw materials used for construction purposes. The indoor radon, thoron, and their progeny concentrations depend on the existence of natural radioactivity in the underlying soil and rock, and building materials. People spend about 80% of their time in indoor environment; therefore, it is important to assess the radiation hazards arising due to the use of rocks in construction of dwellings. Many researchers are being carried out to map radionuclides in rocks, natural radioactivity levels, and gamma dose rate all over the world. However, currently there were no available data from this study area on the natural radioactivity levels in rocks and radiological hazards assessment. This study has been conducted and reported for the first time on natural radioactivity measurements in rocks of Hassan district. The objective of the present study was to measure the activity concentrations of primordial radionuclides in different types of rocks in and around Hassan district to estimate the radiological risk, and secondly to generate baseline data of natural radioactivity in rock samples of the study area. Hence, this study will serve as baseline data for future natural radioactivity monitoring in rocks from Hassan district.

Geology of the study area

Hassan district is located on the border of the Western Ghats, in the southern part of Karnataka state. Figure 1 shows the study area of the Hassan district, Karnataka state, India. It lies between 12° 13' and 13° 33' north latitudes and 75° 33' and 76° 38' east longitudes. Hassan district is considered as Field Museum of Geology, since it consists of varieties of rocks as well as minerals. Geologically, Hassan district forms a part of Precambrian terrain. The major litho units are granitic gneiss, dunite, pyroxenite, peridotite, amphibolite, serpentinite, gabbroic rocks, granulites and horn-blende dykes with titaniferous magnetite, associated pegmatite



Fig. 1 The geological map of the study area

rocks and metamorphic rocks, steatite, Dharwarian schists and dyke rocks. The most important rock formations of the district are Nuggehalli and Holenarasipura schist belts. These two schist belts are the hosts of number of mineral deposits such as chromite, titaniferous magnetite, chalcopyrite, kaolin, asbestos, quartz, granites, gabbroid, ultramafic igneous rocks, dolerites, diorites, bauxite, magnesite, green granites (dykes), garnets, micas, kyanite, columbite, and tantalite.

Materials and methods

Collection and preparation of rock samples

Samples of rocks, such as granites, gneiss, greenstone, pegmatite, dolerites, dykes, pyrites, and schist, quartzite, were collected in the exposed study area. The surface rock was broken with a hammer. The broken rock pieces (gravel size) were collected. Finally, each individual sample was thoroughly mixed separately and composited from which a final sample of about 2 kg was taken and sealed in a polyethylene bag; date of sampling, sample description and location were noted on the sample container and also in a master sample register. The rock samples were powdered and sieved through 150 µm sieves and these samples were transferred to a porcelain dish and oven dried overnight at 110 °C (Rangaswamy et al. 2016). Finally, the prepared samples were weighted and sealed in a 250-ml plastic container and kept for a month before counting by gamma spectrometry to ensure that the radioactive equilibrium was reached between ²²⁶Ra, ²²²Rn, and its progeny.

Estimation of ²²⁶Ra, ²³²Th, and ⁴⁰K in rock samples

The activities of ²²⁶Ra, ²³²Th, and ⁴⁰K in the rock samples were determined by gamma-ray spectrometry employing a 41% relative efficiency n-type low-background HPGe detector having a composite carbon window (Canberra, USA). This detector can be used for the measurement of gamma energies from 5 to 10 MeV. The composite carbon window allows about 90% gamma energies to pass through. The detector was enclosed in a 10-cm thick-graded lead shield (Model 747, Canberra, USA) to reduce the background. The Canberra, DSA-1000 (which consists of HV bias supply, ADC and 16K MCA) and GENIE-2000 softwares were used for data acquisition and analyses. The detector efficiency calibration was performed using the IAEA quality assurance reference materials: RG U-238, RG Th-232, RG K-1, and SOIL-6. The standard material and samples were taken in containers of same size and type so that detection geometry remained the same. Samples were counted long enough to reduce the counting error. The ²²⁶Ra activity was evaluated from the weighed mean of the activities of three photopeaks of 214 Bi (609.3 keV, 1129.3 keV and 1764.5 keV) after applying Compton corrections. In the case of 232 Th, one photopeak of 228 Ac (911.2 keV) and two photopeaks of 208 Tl (583.1 keV and 2614.5 keV) were used in the same way. The activity of 40 K was derived from its 1460.8 keV gamma line. The activity of radionuclides (in Bqkg⁻¹) was calculated using the following relation (IAEA/RCA 1989):

Activity (Bq kg⁻¹) =
$$\frac{(S \pm \sigma) \times 100 \times 1000 \times 100}{EWA}$$
 (1)

where *s* is the net count per second under the photopeak of intensity, σ is the standard deviation of *s*, *E* is the photo-peak efficiency of the detector (%), *A* is the gamma abundance (%) of the radionuclides, and *W* is the mass of the sample (g).

Results and discussion

The average activity concentrations of the primordial radionuclides, such as ²²⁶Ra, ²³²Th, and ⁴⁰K, in the rock samples collected from different regions in and around Hassan district were summarized in Table 1. The mean value of ²²⁶Ra, ²³²Th, and ⁴⁰K radionuclide activity concentration in rock samples varies from 29.4 ± 1.2 to 83.7 ± 1.7 , 37.9 ± 1.2 to 198.4 ± 2.2 , and 346.1 ± 9.0 to 1024.4 ± 19.4 Bq kg⁻¹ with an average value of 50.9 ± 1.4 , 79.6 ± 1.5 , and 609.2 ± 12.8 Bq kg⁻¹, respectively.

The natural radioactivity levels in rock samples vary from different sampling sites due to the variation of concentrations of ²³⁸U, ²³²Th, and ⁴⁰K in the geological formation. The lower activity concentration of ²²⁶Ra, ²³²Th, and ⁴⁰K in rock samples was observed at Kunduru, Palya, Doddamagge, Belur, and Hagare. This is due to the fact that the type of rocks in these locations are mainly schist, dolerite (green color), and diorites which contain lower activity concentration of ²²⁶Ra, ²³²Th, and ⁴⁰K (Rafique et al. 2014). Slightly higher activity concentration of ²²⁶Ra, ²³²Th, and ⁴⁰K in rocks was observed in Ramanathapura, Byrapura, Halebeedu, Hethur, Neggalahally, Honnavally, Belavadi, and Chiknayakanahally. This is because these regions were attributed by gneiss, green granites, and altered gray granites, which contain slightly higher activity concentrations of ²²⁶Ra, ²³²Th and ⁴⁰K. The highest average activity concentrations of ²²⁶Ra, ²³²Th, and ⁴⁰K in rock samples were found at Arehally, Arsikere, Javgal, Channaraypatna, Shravanabelagola, and Nuggehally. The types of rocks in these locations are granites, pegmatite-intruded pink granites and metamorphic rocks, and the regional rocks of chromite mineralization are pegmatite veins housed in granite rocks, both facies enriched in uranium-bearing minerals (Srinivasa et al. 2015). These rocks contain higher radioactivity concentrations as compared to other rock types (United Nations

Table 1 The average activity concentration of 226 Ra, 232 Th, and 40 K in different types of rock samples collected from in and around Hassan district

| Sl No. | Locations | Rock type | Activity of radionuclides in Bq kg ⁻¹ | | | | | |
|--------|---------------------|------------------------|--|-------------------|-------------------|--------------------------------------|--|--|
| | | | ²²⁶ Ra | ²³² Th | ⁴⁰ K | ²³² Th/ ²²⁶ Ra | | |
| 1 | Palya | Schist | 30.7 ± 1.2 | 42.8 ± 1.3 | 421.6 ± 9.5 | 1.39 | | |
| 2 | Kunduru | Schist | 29.4 ± 1.2 | 37.9 ± 1.2 | 391.1 ± 9.1 | 1.29 | | |
| 3 | Doddamagge | Gray granite | 36.5 ± 1.3 | 54.8 ± 1.4 | 346.1 ± 9.0 | 1.5 | | |
| 4 | Ramanathpura | Gray granite | 42.9 ± 1.3 | 42.1 ± 1.2 | 564.5 ± 8.6 | 0.98 | | |
| 5 | Arsikere | Pink granite | 83.7 ± 1.7 | 198.4 ± 2.2 | 785.7 ± 17.1 | 2.37 | | |
| 6 | Javgal | Metamorphic | 69.6 ± 1.5 | 96.9 ± 1.7 | 698.8 ± 13.9 | 1.39 | | |
| 7 | Byrapura | Quartzite | 46.4 ± 1.3 | 65.3 ± 1.5 | 612.5 ± 13.4 | 1.41 | | |
| 8 | Belur | Green granite | 36.9 ± 1.2 | 78.9 ± 1.6 | 486.2 ± 11.6 | 2.14 | | |
| 9 | Halebeedu | Green granite | 43.7 ± 1.3 | 41.4 ± 1.3 | 516.5 ± 12.2 | 0.95 | | |
| 10 | Hagare | Black granite | 32.9 ± 1.2 | 54.1 ± 1.4 | 589.7 ± 8.0 | 1.65 | | |
| 11 | Channarayapatna | Pink granite | 63.3 ± 1.6 | 139.4 ± 1.9 | 615.8 ± 17.3 | 2.27 | | |
| 12 | Thagaduru | Chromite | 51.9 ± 1.3 | 79.4 ± 1.6 | 423.9 ± 9.6 | 1.53 | | |
| 13 | Nuggehally | Titaniferous magnetite | 61.3 ± 1.5 | 132.8 ± 1.9 | 536.7 ± 9.2 | 2.1 | | |
| 14 | Shravanabelagola | Gray granite | 74.7 ± 1.6 | 124.7 ± 1.8 | 641.2 ± 13.2 | 1.67 | | |
| 15 | Malnad Eng college | Pegmatite | 66.9 ± 1.4 | 98.7 ± 1.7 | 982.3 ± 17.8 | 1.72 | | |
| 16 | MCF | Gray granite | 57.0 ± 1.4 | 87.2 ± 1.6 | 956.1 ± 18.5 | 1.3 | | |
| 17 | Gorur | Gray granite | 53.8 ± 1.4 | 72.7 ± 1.4 | 604.5 ± 12.4 | 1.35 | | |
| 18 | Holenarasipura | Ultra basic igneous | 61.6 ± 1.5 | 78.6 ± 1.5 | 786.2 ± 17.0 | 1.46 | | |
| 19 | Arehally | Pegmatite | 78.7 ± 1.7 | 131.6 ± 1.6 | 1024.1 ± 19.4 | 1.67 | | |
| 20 | Neggalahally | Gneiss | 40.7 ± 1.3 | 44.4 ± 1.3 | 546.8 ± 12.9 | 1.09 | | |
| 21 | Hethur | Gneiss | 39.1 ± 1.2 | 48.9 ± 1.3 | 513.8 ± 12.2 | 1.25 | | |
| 22 | Honnavalli | Gray granite | 43.8 ± 1.3 | 58.4 ± 1.4 | 490.6 ± 12.7 | 1.33 | | |
| 23 | Nonavinakere | Gray granite | 55.7 ± 1.4 | 83.7 ± 1.6 | 589.7 ± 13.3 | 1.5 | | |
| 24 | Chikkanayakanahally | Gray granite | 45.7 ± 1.3 | 62.9 ± 1.5 | 347.9 ± 9.2 | 1.38 | | |
| 25 | Belavadi | Quartzite | 38.1 ± 1.2 | 49.6 ± 1.4 | 783.3 ± 17.0 | 1.3 | | |
| 26 | Kadur | Gray granite | 46.1 ± 1.3 | 58.2 ± 1.4 | 642.3 ± 13.5 | 1.27 | | |
| 27 | Devanur | Gray granite | 52.5 ± 1.4 | 72.3 ± 1.5 | 348.2 ± 9.1 | 1.38 | | |
| | Minimum | | 29.4 ± 1.2 | 37.9 ± 1.3 | 346.1 ± 9.0 | 0.95 | | |
| | Maximum | | 83.7 ± 1.7 | 198.4 ± 2.2 | 1024.4 ± 19.4 | 2.37 | | |
| | Average | | 50.9 ± 1.4 | 79.6 ± 1.5 | 609.2 ± 12.8 | 1.5 | | |
| | Median | | 46.2 ± 1.3 | 64.1 ± 1.4 | 520.5 ± 12.6 | 1.39 | | |

Scientific Committee on the Effects of Atomic Radiation (UNSCEAR) 2000; Rangaswamy et al. 2016; Ningappa et al. 2008). The higher ²²⁶Ra, ²³²Th, and ⁴⁰K activity concentrations in the granite rocks may be due to their natural mineralogical compositions (El-Arabi 2007). Mineralogically, the granite gneisses are formed mainly through the intergrowth of quartz, k-feldspar, biotite, muscovite, silicate, and chlorite (Rangaswamy et al. 2016; Rafique et al. 2014). The radionuclides in the rock samples are directly linked with the minerals such as zircon, monazite, thorite, uranothorite, iron oxides, and fluorite. Zircon usually contains uranium and thorium concentrations which ranged from 0.01 to 0.19 and 1 to 2%, respectively (Cuney et al. 1987). Uranium in iron oxides is first trapped by adsorption (Speer et al. 1981). The high uranium content in the mineralized

granite and pegmatite is attributed to the ability of iron oxide in them on adsorbing uranium (Shahul et al. 2014). The studied locations consist of variety of rock systems and these rocks consist of varying amount of primordial radionuclide activity concentrations. Among the studied rock samples, granites and pegmatite-intruded pink granites contain higher activity concentration of primordial radionuclides because of their mineral compositions.

The activity concentration of the radionuclides in samples follows the trend, ${}^{40}\text{K} > {}^{232}\text{Th} > {}^{226}\text{Ra}$. The higher activity concentrations of ${}^{40}\text{K}$ in the rock samples are possibly because principal potassium-bearing minerals in granites are feldspar, orthoclase, muscovite, biotite and mica. The abundance of potassium to some extent is proportional to silica content of the rocks. Since silica content of granite is high,

consequently the associated radioactivity due to ⁴⁰K is also high (Ningappa et al. 2008). The ratio of 232 Th and 226 Ra was in the range of 0.80–1.96 with a median of 1.37. The median value can be used as an indicator of the relative occurrence of uranium and thorium. Activity concentration of thorium is higher than that of radium at all locations because this area is surrounded by pegmatite rocks intruded in granitic rocks (Ningappa et al. 2008). The average ²²⁶Ra, ²³²Th, and ⁴⁰K activity concentrations of rock samples obtained from this study are found to be higher than the worldwide average activity concentration value of 50 Bq kg^{-1} for 226 Ra and 232 Th, and 500 Bg kg⁻¹ for 40 K, respectively (United Nations Scientific Committee on the Effects of Atomic Radiation (UNSCEAR) 2000). The results of the present investigation were compared with the average radioactivity concentration in the rocks of other regions of the world (Table 2).

Correlation between the combinations of radionuclides was carried out to understand the existence of these radionuclides together at a particular location. The correlation between the activity concentrations of 226 Ra and 232 Th in the rock sample is shown Fig. 2. In this case, the regression was found to be linear and positive. A good positive correlation coefficient (R^2) of 0.76 between 226 Ra and 232 Th was observed. It is as expected, since 226 Ra and



Fig. 2 Correlation between 226 Ra and 232 Th activity concentrations of rock samples

²³²Th come from natural decay series of ²³⁸U and ²³²Th. A good positve correlation between ²²⁶Ra and ²³²Th may also due to the fact that these isotopes formed at the formation of the earth in equal concentrations.

| Country | Radioactivity cor | ncentrations (Bq kg ⁻¹) | References | |
|---------------------------------|-------------------|-------------------------------------|-------------------|--|
| | ²²⁶ Ra | ²³² Th | ⁴⁰ K | |
| World | 50 | 50 | 500 | (United Nations Scientific Committee on the Effects of Atomic Radiation (UNSCEAR) 2000) |
| Indian average | 32 | 64 | 400 | (United Nations Scientific Committee on the Effects of Atomic Radiation (UNSCEAR) 2000) |
| Indian | 14.8 | 18.3 | 370 | (Mishra and Sadasivan 1971) |
| Brazil | 31 | 73 | 1648 | (Anjos et al. 2005) |
| Egypt | 15.64 ± 1.14 | 14.46 ± 1.08 | 405.7 ± 29.5 | (Harb et al. 2008) |
| Nigeria | 21.1-129 | 424-150 | 64.5-882 | (Tufai et al. 1992) |
| Ghana | 1.2-40 | 3.3-117.5 | 13.5-1510 | (Ahmed et al. 2006) |
| Turkey | 9.4-27.3 | 11.7-23.4 | 275.4-689.2 | (Ahmed et al. 2006) |
| Nepal | 17-100 | 24-260 | 32–541 | (Wallova et al. 2010) |
| Cyprus | 1-588 | 1–906 | 50-1606 | (Tzortzis et al. 2003) |
| Azad Kashmir, Pakistan | 28.46 ± 0.45 | 48.63 ± 1.12 | 666.7 ± 9.39 | (Muhammad et al. 2013) |
| Bangalore rural district | 93.2 | 306.2 | 1074.4 | (Ningappa et al. 2008) |
| Granite region of Karnataka | 32.4-165 | 32.4–530 | 400-1250 | (Sannappa et al. 2010) |
| North Karnataka | 52.76 | 71.51 | 1035 | (Kerur et al. 2010) |
| Ramanagara and Tumkur districts | 41.08 ± 2.12 | 86.26 ± 2.94 | 869.29 ± 3.78 | (Rangaswamy et al. 2016) |
| Mysore city | 52.9 ± 1.3 | 73.8 ± 1.8 | 750.1 ± 9.8 | (Sannappa et al. 2003) |
| Hassan district | 50.94 ± 1.4 | 79.55 ± 1.5 | 606.20 ± 12.8 | Present study |

Table 2 Comparison of the average activity of ²²⁶Ra, ²³²Th, and ⁴⁰K in rock samples of Hassan district with different environs of the world

Radiological characterization

To highlight the radiological risk related to the radionuclides present in the rocks, it is necessary to calculate and assess the radium equivalent activity, gamma index, alpha index, outdoor and indoor radiation hazard indices, air-absorbed dose rates and annual effective doses. Table 3 summarizes the calculated average values for the above parameters for rock samples.

Estimation of radium equivalent activity (Ra_{eg})

The radium equivalent activity (Ra_{eq}) was computed in this work to assess the gamma radiation risk to human beings. This compares the specific activity of the samples containing different amounts of ²²⁶Ra, ²³²Th, and ⁴⁰K (Amrani and Tahtat 2001). The radium equivalent concept allows a single index or number to describe the gamma output from different mixtures of 226 Ra, 232 Th, and 40 K in rock samples from different locations. It is accepted that 370 Bq kg⁻¹ of 226 Ra, 259 Bq kg⁻¹ of 232 Th, and 4810 Bq kg⁻¹ of 40 K produces the same gamma-ray dose rate. Since the distribution of radionuclides is not uniform, the exposure due to the radionuclides is estimated in terms of radium equivalent radioactivity. This index was calculated using the formula (Amrani and Tahtat 2001; Rangaswamy et al. 2016):

$$Ra_{eq} = C_{Ra} + 1.43C_{Th} + 0.07C_{K},$$
(2)

where C_{Ra} , C_{Th} and C_K are the specific activities of ²²⁶Ra, ²³²Th, and ⁴⁰K in Bq kg⁻¹, respectively.

The Ra_{eq} due to radionuclides in rock samples varied from 113.76 to 427.80 Bq kg⁻¹ with an average value of

 Table 3
 Calculated radiation hazard indices such as absorbed dose, alpha index, gamma index, radium equivalent activity, external hazard index, and equivalent effective dose for rock samples

| Sl No. | Locations | $I_{\boldsymbol{\gamma}}$ | I_{α} | $\operatorname{Ra}_{eq}(\operatorname{Bq} kg^{-1})$ | H _{ex} | H _{in} | D (nGy h ⁻¹) | $E_{\text{In}} (\text{mSv y}^{-1})$ | $E_{\text{Out}} (\text{mSv y}^{-1})$ | $E_{\text{Total}} (\text{mSv y}^{-1})$ |
|--------|---------------------|---------------------------|--------------|---|-----------------|-----------------|--------------------------|-------------------------------------|--------------------------------------|--|
| 1 | Palya | 0.46 | 0.15 | 124.39 | 0.34 | 0.42 | 57.63 | 0.28 | 0.07 | 0.35 |
| 2 | Kunduru | 0.42 | 0.14 | 113.76 | 0.31 | 0.39 | 52.8 | 0.26 | 0.06 | 0.32 |
| 3 | Doddamagge | 0.51 | 0.18 | 141.39 | 0.38 | 0.48 | 64.34 | 0.32 | 0.08 | 0.39 |
| 4 | Ramanathpura | 0.54 | 0.21 | 146.54 | 0.4 | 0.51 | 68.77 | 0.34 | 0.08 | 0.42 |
| 5 | Arsikere | 1.53 | 0.42 | 427.8 | 1.16 | 1.38 | 191.2 | 0.94 | 0.23 | 1.17 |
| 6 | Javgal | 1.02 | 0.34 | 277.28 | 0.75 | 0.94 | 128.1 | 0.63 | 0.16 | 0.79 |
| 7 | Byrapura | 0.69 | 0.23 | 186.94 | 0.5 | 0.63 | 86.42 | 0.42 | 0.11 | 0.53 |
| 8 | Belur | 0.68 | 0.18 | 187.21 | 0.51 | 0.61 | 85 | 0.42 | 0.1 | 0.52 |
| 9 | Halebeedu | 0.52 | 0.22 | 142.58 | 0.39 | 0.5 | 66.69 | 0.33 | 0.08 | 0.41 |
| 10 | Hagare | 0.48 | 0.16 | 132.56 | 0.36 | 0.45 | 59.95 | 0.29 | 0.07 | 0.37 |
| 11 | Channarayapatna | 1.11 | 0.3 | 308.07 | 0.83 | 1 | 138.2 | 0.68 | 0.17 | 0.85 |
| 12 | Thagaduru | 0.71 | 0.26 | 198.13 | 0.54 | 0.68 | 89.64 | 0.44 | 0.11 | 0.55 |
| 13 | Nuggehally | 1.15 | 0.31 | 317.56 | 0.86 | 1.03 | 144.3 | 0.71 | 0.18 | 0.88 |
| 14 | Shravanabelagola | 1.09 | 0.37 | 302.31 | 0.82 | 1.02 | 136.5 | 0.67 | 0.17 | 0.84 |
| 15 | Malnad Eng college | 1 | 0.28 | 271.1 | 0.73 | 0.89 | 125.5 | 0.62 | 0.15 | 0.77 |
| 16 | MCF | 0.99 | 0.33 | 267.33 | 0.72 | 0.9 | 124.6 | 0.61 | 0.15 | 0.76 |
| 17 | Gorur | 0.74 | 0.27 | 204.21 | 0.55 | 0.7 | 93.93 | 0.46 | 0.12 | 0.58 |
| 18 | Holenarasipura | 0.83 | 0.27 | 226.53 | 0.61 | 0.76 | 105 | 0.52 | 0.13 | 0.64 |
| 19 | Arehally | 1.26 | 0.39 | 345.75 | 0.93 | 1.15 | 158.6 | 0.78 | 0.19 | 0.97 |
| 20 | Neggalahally | 0.54 | 0.2 | 146.15 | 0.39 | 0.5 | 68.35 | 0.34 | 0.08 | 0.42 |
| 21 | Hethur | 0.55 | 0.19 | 148.61 | 0.4 | 0.51 | 69.03 | 0.34 | 0.08 | 0.42 |
| 22 | Honnavalli | 0.6 | 0.22 | 165.09 | 0.45 | 0.56 | 75.97 | 0.37 | 0.09 | 0.47 |
| 23 | Nonavinakere | 0.8 | 0.28 | 220.76 | 0.6 | 0.75 | 100.9 | 0.49 | 0.12 | 0.62 |
| 24 | Chikkanayakanahally | 0.58 | 0.23 | 162.38 | 0.44 | 0.56 | 73.59 | 0.36 | 0.09 | 0.45 |
| 25 | Belavadi | 0.64 | 0.19 | 169.59 | 0.46 | 0.56 | 80.36 | 0.39 | 0.1 | 0.49 |
| 26 | Kadur | 0.66 | 0.23 | 179.02 | 0.48 | 0.61 | 83.34 | 0.41 | 0.1 | 0.51 |
| 27 | Devanur | 0.65 | 0.26 | 182.7 | 0.49 | 0.64 | 82.45 | 0.4 | 0.1 | 0.51 |
| | Minimum | 0.42 | 0.14 | 113.76 | 0.31 | 0.39 | 52.8 | 0.26 | 0.06 | 0.32 |
| | Maximum | 1.53 | 0.42 | 427.8 | 1.16 | 1.38 | 191.2 | 0.94 | 0.23 | 1.17 |
| | Average | 0.78 | 0.25 | 210.1 | 0.57 | 0.72 | 96.29 | 0.47 | 0.12 | 0.59 |
| | Median | 0.68 | 0.23 | 186.94 | 0.5 | 0.63 | 85 | 0.42 | 0.1 | 0.52 |

210.10 Bq kg⁻¹. The average Ra_{eq} value of all the rock samples is less than the criterion limit of 370 Bq kg⁻¹ (Amrani and Tahtat 2001) and as such these rocks do not pose any radiological hazard when used for the construction purpose.

External and internal health hazard indices (H_{ex} and H_{in})

The concept of the hazard indices was used to assess the potential radiological risk associated with human. Gamma radiation emitted by the radionuclides concern is an estimate of external hazard index that is determined using equation (Beretka and Mathew 1985; UNSCEAR 2008):

$$H_{\rm ex} = \frac{C_{\rm Ra}}{370} + \frac{C_{\rm Th}}{259} + \frac{C_{\rm K}}{4810} \tag{3}$$

²²²Rn is a gaseous product of the decay of radium (²²⁶Ra). It is short lived and constitutes a major source of internal radiation exposure (Farai and Vincent 2006). The internal exposure of living cells to radon and its daughter products is referred to as internal hazard index (H_{in}) and was estimated using equation (UNSCEAR 2008):

$$H_{\rm in} = \frac{C_{\rm Ra}}{185} + \frac{C_{\rm Th}}{259} + \frac{C_{\rm K}}{4810},\tag{4}$$

where C_{Ra} , C_{Th} , and C_K are the activity concentrations of ²²⁶Ra, ²³²Th, and ⁴⁰K in Bq kg⁻¹, respectively. The value of these indices must be less than 1 to keep the radiation hazards insignificant.

The calculated internal (H_{in}) and external (H_{ex}) radiation hazard indices of the rock samples are found to vary from 0.39 to 1.38 with an average value of 0.72 and 0.31 to 1.16 with an average value of 0.57, respectively. The average internal and external hazard indices values of the rock samples are less than unity and there is no potential radiation hazard due to the use of the studied materials for construction purposes.

Gamma index (I_v)

The gamma index (I_{γ}) , which is expressed as the sum of fractions of the measured activity concentrations in the sample, has been defined by the European Commission (European Commission (EC), Radiation Protection 112 1999) and it is given below

$$I_{\gamma} = \frac{C_{Ra}}{300} + \frac{C_{Th}}{200} + \frac{C_K}{3000}$$
(5)

The gamma index is correlated with the annual dose rates due to excess external gamma radiation caused by superficial material. The calculated gamma index (I_{γ}) values of all the studied rock samples varied in the range of 0.42–1.53 with an average value of 0.78. The value of I_{γ} must be less than 1 to keep the radiation hazard insignificant to the general population. The calculated average I_{γ} values of the rock samples are far below the criterion limit of unity as per European Commission of Radiation Protection reports (European Commission (EC), Radiation Protection 112 1999). Therefore, the rocks were found to be safe and did not pose any serious radiation hazard to the population living in this area.

Alpha index (I_α)

Alpha index (I_{α}) is the assessment of the excess alpha radiation due to the radon inhalation originating from the building materials (Megumi et al. 1998; Stoulos et al. 2003). The alpha index was determined by the following formula:

$$I_{\alpha} = \frac{C_{Ra}}{200} \tag{6}$$

Alpha index $(I_{\alpha}) \leq 1$ corresponds to a ²²⁶Ra activity concentration ≤ 200 Bq kg⁻¹. When the activity concentration of ²²⁶Ra in a rock exceeds the value of 200 Bq kg⁻¹ (I_{α}>1), it is possible that the radon exhalation from this material could cause indoor radon concentration exceeding 200 Bq m⁻³ (Righi and Bruzzi 2006). The calculated alpha index (I_{α}) values for rocks are found in the range of 0.14–0.42 with an average value of 0.25. From the result, it is concluded that the activity concentrations of ²²⁶Ra in all the samples are less than the recommended upper limit of 200 Bq kg⁻¹. The alpha index values for all rock samples are less than unity. This reveals that the radon exhalation from the rocks would cause indoor concentration less than 200 Bq m⁻³.

Gamma-absorbed dose rate in air

The absorbed dose rates (*D*) due to gamma radiation in air have been calculated based on the guidelines provided by UNSCEAR (United Nations Scientific Committee on the Effects of Atomic Radiation (UNSCEAR) 2000). The UNSCEAR has given the dose conversion factors for converting the activity concentrations of 226 Ra, 232 Th, and 40 K into doses (nGy h⁻¹ per Bq kg⁻¹) as 0.462, 0.604, and 0.0417, respectively. Using these factors, the external gamma-absorbed dose rate (*D*) due to rocks is calculated using the formula (United Nations Scientific Committee on the Effects of Atomic Radiation (UNSCEAR) 2000)

$$D(nGy h^{-1}) = 0.462C_{Ra} + 0.604C_{Th} + 0.0417C_{K}.$$
 (7)

For rock samples, the calculated results of absorbed dose rates range from 52.80 to 191.20 nGy h^{-1} with an average value of 96.29 nGy h^{-1} which is higher than the population-weighted average value of global primordial radiation of 55 nGy h^{-1} (United Nations Scientific Committee on the Effects of Atomic Radiation (UNSCEAR) 2000).

Annual effective dose (AED)

The indoor annual effective dose rates (*E*) are evaluated using the conversion coefficient from the absorbed dose in air to the effective dose (0.7 SvG y⁻¹) and the indoor occupancy factor (0.8) implied that 80% of time is spent indoors, as proposed in UNSCEAR (UNSCEAR 1993). Thus, the annual effective dose (mSv y⁻¹) received by a building occupant due to the activity in the rocks and building materials was estimated using the following equation:

$$E_{\rm In} = D \times 8760 \times 0.8 \times 0.7 \times 10^{-6}.$$
 (8)

Similarly, the outdoor annual effective dose in rock samples is estimated from the net absorbed gamma radiation dose rate (*D*) by taking into account the outdoor occupancy factor 0.2 and the conversion factor from absorbed dose rate in air to effective dose 0.7 SvG y⁻¹ for the adults. The E_{Out} is calculated using the following equation proposed by UNSCEAR (United Nations Scientific Committee on the Effects of Atomic Radiation (UNSCEAR) 2000):

$$E_{\rm Out} = D \times 8760 \times 0.2 \times 0.7 \times 10^{-6},\tag{9}$$

where E_{Out} is the outdoor annual effective dose in mSv y⁻¹.

The calculated indoor annual effective dose rates of the rocks samples are found to vary from 0.26 to 0.94 mSv y⁻¹with an average value of 0.47 mSv y⁻¹. Similarly, the calculated outdoor annual effective dose rate of rock samples varies from 0.06 to 0.23 mSv y⁻¹ with an average value of 0.12 mSv y⁻¹. The total annual effective dose rates for rocks are found to vary from 0.32 to 1.17 mSv y⁻¹with an average value of 0.59 mSv y⁻¹. From the results, it is clear that the average annual equivalent effective dose rate for rock samples is within the safety limit (1 mSv y⁻¹) as proposed by UNSCEAR (UNSCEAR 2008) and as recommended by the International Commission on Radiological Protection for the individual members of the public (International Commission on Radiological Protection 1999).

Conclusion

The gamma-ray spectrometry technique has been used to determine the ²²⁶Ra, ²³²Th, and ⁴⁰K activity levels in rock samples. The results reveal that the average activity concentrations of ²²⁶Ra, ²³²Th, and ⁴⁰K in the rocks are slightly higher than the world average values. The highest average activity was observed in granite rocks. The estimated radiological hazard indices such as radium equivalent activity, gamma index, alpha index, external hazard index, internal hazard index, and annual effective dose rate are well within the world average values. The average outdoor

gamma-absorbed dose rates are slightly higher than the population-weighted average value of global primordial radiation of 55 nGy h⁻¹. Except pink granites from Arsikere, the average indoor and outdoor annual effective dose values in rock samples were found to be lower than the safety limit (1 mSv y⁻¹) as proposed by UNSCEAR (UNSCEAR 2008). From the radiological point of view, all the rock samples (except granites, pegmatite-intruded pink granites) of the study area are safe for construction purposes. The data obtained in this study will serve as baseline data for assessing the radiation exposure of the residents.

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