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Geochemical processes regulating groundwater chemistry with special reference to nitrate and fluoride enrichment in Chhatarpur area, Madhya Pradesh, India

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Abstract The present study focuses on the hydrogeochemical composition of groundwater in Chhatarpur area with special focus on nitrate and fluoride contamination, considering the fact that groundwater is the only major source of drinking water here. Carbonate and silicate mineral weathering followed by ground water-surface water interactions, ion exchange and anthropogenic activities are mainly responsible for high concentrations of cations and anions in the groundwater in the region. The average concentration of nitrate and fluoride found in 27 samples is 1.08 and 61.4 mg/L, respectively. Nitrate enrichment mainly occurs in areas occupied with intense fertilizer practice in agricultural fields. Since the area is not dominated by industrialization, the possibility of anthropogenic input of fluoride is almost negligible, thus the enrichment of fluoride in groundwater is only possible due to rock-water interaction. The highly alkaline conditions, which favor the fluorite dissolution, are the main process responsible for high concentration of fluoride.

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Faculty of Science and Environment, Mahatma Gandhi Chitrakoot Gramodaya Vishwavidyalaya, Chitrakoot 485780, Madhya Pradesh, India **Keywords** Chhatarpur District · Groundwater quality · Nitrate · Fluoride · Hydrogeochemistry

Introduction

The groundwater resources are under huge risk due to speedy population growth, industrialization, intensive agriculture, urbanization and changing land use pattern, which in turn is endangering the water resources in terms of quality and quantity. The variation in the groundwater quality in an area is a function of physical and chemical parameters. These parameters are greatly influenced by geological formations and anthropogenic activities (Subramani et al. 2005). Groundwater chemical signature mainly depends on the interaction pattern with aquifer minerals or by internal mixing among different groundwater along flow paths in the subsurface (Domenico 1972; Datta et al. 1996; Belkhiri et al. 2011). The groundwater quality is also strongly affected by change in land use/land cover pattern (Basnyat et al. 1999; Roth et al. 1996; Osborne and Wiley 1988). Therefore, assessment of groundwater quality is needed to ensure its safer use (Vijith and Satheesh 2007). The pesticides and fertilizers used in agriculture to increase the production causes degradation of water quality (Carpenter et al. 1998; Matson et al. 1997; Kolpin 1997; Griffith 2001).

Groundwater contamination is a global problem that has severe impacts on human health as well as on ecosystem and economy (Addiscott et al. 1991). This problem is mainly because of the ineffective environmental conservation policies as well as areas with intensive industrialization, use of fertilizers, mining, etc. To minimize the impacts of groundwater pollution, regular monitoring of groundwater quality is needed. Dissolved nitrogen in the form of nitrate (NO₃) is the most common contaminant of groundwater. The source of nitrate in groundwater originates from nonpoint sources such as leaching of chemical fertilizers, animal manure and sewage discharges. Nitrate enrichment (more than 45 mg/L) in drinking water is an environmental health concern because it increases the risks of methemoglobinemia in infants. It is found that 21 out of 29 states of India have localized nitrate pollution in groundwater (CGWB 2010).

High concentration of fluoride (more than 1.5 mg/L) in groundwater or surface water is also a case of great concern because it causes fluorosis (Alagumuthu and Rajan 2010). The main sources of fluoride in groundwater are rocks containing fluoride minerals such as fluorspar, cryolite, muscovite, biotite, fluorite and fluorapatite are the most common. (Handa 1975; Wenzel and Blum 1992; Msonda et al. 2007). Mica is also one of the potential sources of fluoride where it is found as fluorapatite (Singh et al. 2011). The content of fluoride in groundwater can be credited to the weathering/leaching of fluoride containing minerals (Singh et al. 2011) as well as pH, temperature and concentration of calcium and bicarbonate ions in water (Alagumuthu and Rajan 2010), where human interference like mining and industrial activities are almost negligible. Thus, the areas where such rock types are dominant probability of fluoride concentration can be questionable (Sarma and Rao 1997; Nordstrom et al. 1989; Banks et al. 1995; Gizaw 1996; Frengstad et al. 2001). In the case of India, there are several locations in the States of Andhra Pradesh, Gujarat, Karnataka, Madhya Pradesh, Rajasthan, Chhattisgarh, Haryana, Orissa, Punjab, Haryana, Uttar Pradesh West

Fig. 1 Location of the study area with the location of sampling points showing groundwater sample collected Bengal, Bihar, Delhi, Jharkhand, Maharashtra, and Assam where the fluoride in groundwater exceeds 1.5 mg/l (CGWB 2010).

The main objective in the study is to assess groundwater quality of Chhatarpur District of Madhya Pradesh, India, to determine the factors affecting groundwater quality as well as identification of the main geochemical processes controlling the groundwater quality.

Study area

The study area is located on the central portion of Bundelkhand plateau in Madhya Pradesh state of India. It is bordered by Uttar Pradesh state and lies between $24^{\circ}06'$ – $25^{\circ}20'$ latitudes and $79^{\circ}59'$ – $80^{\circ}26'$ longitudes (Fig. 1). The climate of study area is semi-arid type, although this is highly variable depending on the monsoon. The area has mean annual maximum and minimum temperatures of 32.7 and 18.1 °C, respectively. The annual rainfall in the area is about 1,060 mm, with the highest precipitation (about 90 % of the total amount) occurring between June and September.

Geology of the area

The geology plays an important role in the groundwater quality determination as different geological units have different weathering rates and process (Duan et al. 2002). Figure 2 shows the geology of the study area. Geologically, the area is covered by Bundelkhand granite (older than 2.6



Fig. 2 Geology of Chhatarpur



billion years), Bijawar group (2.6-2.4 billion years), Vindhyan supergroup (1.4-0.9 billions years), and alluvium (Auden 1933). Bundelkhand granites are frequently associated with pegmatites and quartz reefs (Ahmad 1984). Most of the part of the study area is covered by Bundelkhand granite in northern and north central part. Bundelkhand granites are overlain by the Bijawar group of rocks (Bhattacharya and Bhattacharya 1976). The alluvium is present in north eastern part of the area. The Bijawar group of the area consists of terrigenous sequence of basal carbonates and shales with green schists or pillow basalts, chloritic shales, ferruginous quartzites and banded iron formations. The Vindhyan basin is a very large intra-cratonic sedimentary basin, which contains a sequence of sandstone-shale-limestone (Rao et al. 2005). Vindhyan super group have been broadly divided into four major groups, namely the Semri (lower), Kaimur, Rewa and Bhander (upper) based on lithological similarities and into Lower and Upper Vindhyans on the basis of major tectonic (major unconformity) evidences (Mallet 1869). The upper Kaimur group rocks of the Vindhyan supergroup are represented by sandstones all over the Vindhyan basin. The sandstones are red to greyish-white, medium to fine grained, compact and highly jointed. Lower Vindhyan formations are represented by Tirohan limestone, Baghain sandstone; diamond-bearing conglomerate of Kaimur group is present in Panna District (Bhattacharya and Bhattacharya 1976; Avtar et al. 2011).

Land use/land cover of the area

The land use/land cover map was generated using unsupervised classification of Landsat TM data of February, 2009. The study of land use/land cover is important to analyze the spatial-temporal characteristics of land use/land cover change and the spatio-temporal variability of groundwater quality (Jinfeng and Quanjun 2009). The area was classified into 10 classes, viz. dense forests, current fallow land/agriculture area, land without scrap, degraded forest, land with scrub, fallow land, bare exposed rock, water body, built-up area and water logged area. Most of the area is covered by forests fallowed by agricultural area (Fig. 3).

Materials and methodology

A total of 27 water samples were collected in April 2008 in polypropylene bottles (Tarsons) from shallow hand pumps, dug-wells and tube wells. Table 4 in the appendix shows the location of the sampling points. At the time of sampling, bottles were thoroughly rinsed 2–3 times with sample water to be taken. The non-conservative and physical parameters (pH, temperature, electrical conductivity) were measured in situ with their respective probes. Bicarbonate (HCO₃⁻) was analyzed using titration method with sulfuric acid. Other anions (Cl⁻, SO₄²⁻, NO₃⁻ and F⁻) have been analyzed by ion chromatography and major cations (Na⁺, Ca²⁺, Mg²⁺

Fig. 3 Land use/land cover of Chhatarpur



and K^+) were analyzed by atomic absorption spectrometer (AAS). In order to discern regional trends and to identify chemical processes, it is necessary to make distinction between water types and the classification of samples into different groups (Stuyfzand 1999). Presentation of geochemical data in the form of graphical charts such as the Piper diagram is helpful to recognize and differentiate various water facies (Piper 1944). In this study the Piper diagram is drawn using the AqQA software (RockWare).

In order to understand the origin of groundwater mineralization and the processes responsible for chloride, fluoride and nitrate enrichment in groundwater in the study area, saturation index (SI) was calculated using PHREEQC code (Parkhurst and Appelo 1999). SIs are used to evaluate the degree of equilibrium between water and respective mineral. Changes in saturation state are useful to distinguish different stages of hydrochemical evolution and help to identify which geochemical reactions are important in controlling water chemistry (Coetsiers and Walraevens 2006; Langmuir 1997). The SI of a mineral can be obtained using following equation (Garrels and Christ 1965).

 $SI = \log_{10}(K_{IAP}/K_{SP})$

where SI is the saturation index, K_{IAP} is the ion activity product of particular solid phase and K_{SP} is the solubility product of the phase. Based on the value of the SI, the saturation states are recognized as saturation (equilibrium; SI = 0), unsaturation (dissolution; SI < 0) and oversaturation (precipitation; SI > 0). In other words, SI < 0 which specifies case of undersaturation could reflect the character of water from a formation with insufficient amount of the mineral for solution or short residence time and SI > 0 specifies that the groundwater is oversaturated with respect to the particular mineral and therefore incapable of dissolving more of the mineral. Such an index value reflects groundwater discharging from an aquifer containing ample amount of the mineral with sufficient resident time to reach equilibrium.

The matrix of hydrogeochemical data obtained was subjected to multivariate analytical techniques. These techniques help to simplify and organize large data sets in order to make useful generalizations. The data set was subjected to correlation analysis using Spearman's rank coefficient which is based on the ranking of the data and not their absolute value [Statistical Package for Social Sciences (SPSS) software package (Version 20)] (Kumar et al. 2010).

Results and discussion

Hydrochemical characteristics

A statistical summary of the analytical results (average, minimum, maximum, standard deviation and range) for each water-quality characteristic analyzed is shown in Table 1. Slightly alkaline nature of water with high EC are the combined effect of the high concentration of dissolved solids in the groundwater which can be attributed to high reactivity

 Table 1
 Statistical summaries of hydrogeochemical parameters of ground water samples

Parameters	Average	Minimum	Maximum	SD	Range
рН	7.5	7.2	8.0	0.2	0.8
EC (µs/cm)	773.5	392.0	1,500.0	277.5	1,108.0
HCO_3^- (mg/L)	209.3	98.0	367.0	68.3	269.0
Cl^{-} (mg/L)	79.8	11.0	241.0	54.8	230.0
$\mathrm{SO_4}^{2-}$ (mg/L)	31.7	0.0	85.0	18.0	85.0
NO ₃ ⁻ (mg/L)	61.4	8.0	192.0	45.7	184.0
F^{-} (mg/L)	1.1	BDL	2.4	0.6	1.3
TH as CaCO ₃ (mg/L)	303.3	150.0	825.0	138.7	675.0
Ca ²⁺ (mg/L)	67.0	24.0	164.0	32.4	140.0
Mg^{2+} (mg/L)	33.7	6.2	158.1	28.6	152.0
Na ⁺ (mg/L)	44.4	16.0	83.0	15.0	67.0
K ⁺ (mg/L)	3.0	0.5	26.0	4.9	25.5

BDL below detection limit

of the soil, local variation in soil type, multiple aquifer system and agricultural activities in the area. Among anions the trend was $HCO_3^- > Cl^- > NO_3^- > SO_4^{2^-} > F^-$, while for cations the trend was $Ca^{2+} > Na^+ > Mg^{2+} > K^+$. The average meq/L ratio trend of $Ca^{2+}+Mg^{2+}/Na^{+}+K^{+}$ is 3.1 indicates the dominance of carbonate rock weathering in the Chhatarpur area. Relation between Tz+ (here Tz+ defined as the sum of all major cations, i.e. Ca²⁺, Na⁺, Mg²⁺ and K^+) versus (Ca²⁺ + Mg²⁺) and (Na⁺ + K⁺) is shown in Fig. 4. The $(Ca^{2+}+Mg^{2+})$ versus Tz+ plot for the groundwater yields a linear trend (with $R^2 = 0.96$ and calculated p value = 0.000014) indicating that $Ca^{2+} + Mg^{2+}$ accounts for most of the cations. In contrast, the trend between Na⁺ + K⁺ and Tz+ is poor (with $R^2 = 0.27$ and calculated p value = 0.104). These results infer that the contribution of cations via alumino-silicate weathering is low in comparison to carbonate weathering.



Fig. 4 Scatter plot showing the relation between Tz^+ versus $Ca2^+ + Mg2^+$ and $Na^+ + K^+$. Here Tz^+ represents sum of the all major cations, i.e. Ca^{2+} , Na^+ , Mg^{2+} and K^+

Contaminations for chemical species

The statistical analysis for the comparison of the values of different chemical species in the groundwater samples from the standard given by WHO and ISO for desirable limit and permissible limit has been given in Table 2. The concentrations of cations, such as Na⁺, K⁺, Ca²⁺ and Mg^{2+} all were within the maximum permissible limits for drinking, except for a few samples in the case of magnesium. However, for some chemical species there is a great difference between the percentage of samples which were found beyond permissible limit (PL) and the percentage of samples which was found beyond desirable limits (DL) (Table 2). For major anions like HCO_3^- and Cl^- , concentrations lie within permissible limits except sporadic enrichment of Cl⁻. Generally, high HCO₃⁻ concentration is indicative of carbonate weathering and decomposition of organic matter; however, higher concentrations of Cl⁻ can be attributed to the input from sewage effluents in the village areas. For minor anions likes NO3⁻ and F⁻ concentration are of main concern from the health perspective as 48 and 20 % of the water samples had high concentrations of nitrate and fluoride, respectively, compared to WHO permissible limit. With an average concentration of 61.4 mg/L for NO₃⁻, groundwater samples clearly show an indication for the influence of fertilizer use in surrounding agricultural areas as well as microbial mineralization in the groundwater. Usually the ammonia within nitrogen fertilizers is transformed into nitrate by the process of nitrification, and the possible reaction responsible here is shown:

$NH_3 + 2O_2 \rightarrow NO_3^- + H^+ + H_2O.$

Furthermore, the higher concentration of nitrate may also be a result of the presence of *E. coli*, *Staphylococcus aureus*, *Proteus vulgaris*, *Salmonella typhi*, and *Pseudomonas aeruginosa* in the fecal matter, as reported by Saha and Kumar (2006). This seems reasonable as there is a severe lack of a proper sanitation system in the district. Further, animal waste sources are also a significant contributor of nitrate to groundwater, especially within mixed land use. Some part of the study area is occupied by industrial and urban land use; therefore, some nitrate leaching from landfill sites and industrial effluents cannot be neglected (Kundu and Mandal 2009; Kumar et al. 2010).

Fluoride is one of the main trace elements in groundwater and generally occurs as a natural constituent. Its pollution and health aspect was also reported earlier by Chaurasia et al. (2007). Weathering of bedrock containing fluoride minerals (like biotite, muscovite) is generally responsible for high concentration of fluoride in groundwater. Generally in acidic medium, i.e. groundwater with low pH, fluoride remains adsorbed in clay due to its low solubility. However, in alkaline medium, i.e. groundwater

Parameter	WHO standard (1993)		ISO:10500: 1991 standard		% of sample exceeding		Health implications	
	Desirable limit (DL)	Maximum permissible limit (PL)	DL	PL	Desirable limit	Permissible limit		
рН	7-8.5	9.2	6.5-8.5	No relax	_	_	Taste	
EC (µS/cm)	500	1,400	500	1,000	93	7 (20)	Gastro-intestinal irritation	
Na ⁺ (mg/L)	-	200	-	_	_	_	Scale formation	
Ca ²⁺ (mg/L)	75	200	75	200	26	_	Scale formation	
Mg ²⁺ (mg/L)	50	150	30	100	10 (52)	(4)	Encrustations in water supply structure	
K ⁺ (mg/L)	-	200	-	200	_	_	Interfere with nerve impulse	
Cl ⁻ (mg/L)	200	600	250	1,000	4	_	Salty taste	
SO_4^{2-} (mg/L)	200	400	200	400	_	_	Laxative effect	
NO_3^- (mg/L)	-	45	45	100	(48)	48 (15)	Methemoglobinemia	
F ⁻ (mg/L)	-	1.5	1.0	1.5	(44)	20 (20)	Fluorosis	

Table 2 Different guideline values for the drinking water with percentage of water samples beyond permissible limit (PL) and the resulting health implications

Values in parenthesis are the percentage of sample having values higher than limits of drinking standard of ISO standard

with high pH, the hydroxyl (OH⁻) group replaces the exchangeable fluoride from minerals like biotite/muscovite and results in its enrichment. The possible reaction responsible for muscovite is shown below (Singh et al. 2011):

$$\begin{aligned} & \operatorname{KAl}_2[\operatorname{AlSi}_3\operatorname{O}_{10}]\operatorname{F}_2 + 2\operatorname{OH}^- \to \operatorname{KAl}_2[\operatorname{AlSi}_3\operatorname{O}_{10}][\operatorname{OH}_2] \\ & + 2\operatorname{F}^- \end{aligned}$$

The distribution pattern of NO_3^- and F^- is shown in Figs. 5 and 6, respectively. When we compare these distribution patterns from the land use–land cover pattern of the study area in Fig. 3, we found that nitrate enrichment corresponds well to the area dominated by agricultural practices; however, fluoride enrichment mainly corresponds to the land without scrub and with very sparse amount of forest/agriculture.

The major ions of water analyses are depicted in the Piper Trilinear equivalence diagram (Fig. 7). Based on the geochemical facies and their percentage, water samples can be classified into five groups, viz. (1) Ca-HCO₃ type (40 %), (2) Mg-HCO₃ type (37 %), (3) Ca-Cl type (15 %), (4) Mg-Cl type (4 %), and (5) Na-HCO₃ type (4 %). Focusing on the Piper diagram of cations, a clear migration tendency from bicarbonate pole to chloride pole is evident, which may indicate water quality shift from pristine environment to the system with much anthropogenic activities.

A correlation matrix of the physiochemical parameters for the water samples is shown in Table 3. The correlation matrix (with statistical significance level at the 0.05) showed good to moderate positive correlation of pH with EC, Ca^{2+} and HCO_3^{-} . A good to moderate correlation was observed between the following pairs: Ca²⁺-Mg²⁺, Ca²⁺-HCO₃⁻, Ca²⁺-SO₄²⁻, Ca²⁺-F⁻, Mg²⁺-HCO₃^{-,} Na⁺-K⁺, Na⁺ $-NO_3^-$, K⁺ $-NO_3^-$, Cl⁻ $-SO_4^{2-}$, Cl⁻ $-NO_3^-$ and SO₄²⁻ -NO₃⁻. Also there is an interesting negative correlation found between HCO_3^- -NO₃⁻. The pairwise relationships between physiochemical parameters show that water quality is mainly controlled by carbonate weathering. There is presence of dolomite as well, which acts as a common source for both calcium and magnesium. The presence of minerals like fluorspar, apatite and fluorapatite are also indicated by a good correlation between calcium and fluoride. Anthropogenic sources, especially excess use of fertilizers, may control the geochemistry at point scales and can be shown by high correlation between chloride, sulphate and nitrate. Negative correlation between bicarbonate and nitrate is a strong indicator of microbial denitrification process (Chkirbene et al. 2009).

Origin of groundwater mineralization in Chhatarpur area

Speciation modeling was performed using program PHRE-EQC (Parkhurst and Appelo 1999) to check the possibility of solubility control for fluorite and other species. Most samples are undersaturated with respect to fluorite, gypsum, dolomite, calcite, aragonite and anhydrite except at some of the locations. On the other hand, groundwater is strongly undersaturated with respect to halite.

The positive values of the calculated SI with respect to calcite, dolomite at few sampling points suggest their oversaturation in respect to those minerals which mainly

Fig. 5 Spatial distribution pattern of nitrate





Fig. 6 Spatial distribution of fluoride

has a geological origin derived from the abundant carbonate weathering points in the plain (Fig. 8). Thus, calcite as well as dolomite might act as the dual origin of Ca^{2+} and

 HCO_3^- in these samples. This explanation for calcite it well described by Appelo and Postma (2005), as shown below through following equation:





Table 3 Correlation matrix for chemical constituent of ground water samples

	pH	EC	HCO ₃ ⁻	Cl^{-}	$\mathrm{SO_4}^{2-}$	NO_3^-	F^{-}	PO4 ³⁻	Ca ²⁺	Mg ²⁺	Na ⁺
EC	0.6										
HCO_3^-	0.7	0.7									
Cl^{-}	0.4	0.5	0.5								
SO_4^{2-}	-0.2	0.4	-0.1	0.6							
NO_3^-	-0.2	0.5	-0.4	0.7	0.7						
F^{-}	-0.1	0.3	0.2	0.2	0.1	0.1					
PO_4^{3-}	0.2	0.1	0.1	-0.1	0.1	0.1	0.3				
Ca ²⁺	0.6	0.8	0.7	0.8	0.6	0.4	0.6	0.1			
Mg^{2+}	0.1	0.5	0.6	0.2	0.4	0.2	0.1	0.1	0.6		
Na ⁺	0.1	0.5	0.3	0.5	0.3	0.5	0.1	0.1	0.4	0.1	
K^+	-0.1	0.1	0.1	0.1	0.1	0.6	0.1	0.1	0.1	0.1	0.5

Values equal to or exceeding critical values are in bold font

$CaCO_3 + CO_2 + H_2O \leftrightarrow Ca^{2+} + 2HCO_3{}^-$

Furthermore, gypsum and anhydrite show same saturation pattern to that of calcite and dolomite suggesting that calcium is also being supplied through gypsum and anhydrite. Saturation of fluorite minerals at some points suggests contact of alkaline water (supported by high pH value in the study area) with fluoride-bearing minerals. Basically, fluorite dissolution might be catalyzed by bicarbonates produced through hydrolysis of aluminosilicate minerals which is well supported by localized granitic or sandstone dominant aquifers in this study area. The possible reaction responsible here is shown as following equations:

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$$\begin{split} &2CaMgFeAl_2Si_3O_{12}(Pyroxene)+8CO_2+0.5O_2+11H_2O\\ &\rightarrow 2Clay\ minerals\ (e.g.,\ kaolinite)+2Fe(OH)_3+2Ca^{2+}\\ &+2Mg^{2+}+8HCO_3^-+2SiO_2 \end{split}$$

 $CaF_2 + 2HCO_3^- \rightarrow CaCO_3 + 2F^- + H_2O + CO_2$

So using saturation index, it is concluded that weathering, dissolution and hydrolysis are the main processes determining water quality at spatial scale.

Conclusion

The groundwater quality of the study area has a primary problem of nitrate and fluoride contamination. The groundwater chemistry is mainly controlled by natural conditions of rock-water interaction, i.e., geogenic processes along with the intense agricultural activities at some point. The alkaline nature of groundwater triggers replacement of the exchangeable fluoride from minerals like biotite/muscovite and results in its enrichment. Another possible mechanism is that granitic rocks, which contain abundant fluoride-bearing minerals, upon weathering release it to the dissolved phase in groundwater. The groundwater mineralization in the study area is not a homogenous process, but it is related to different sources and dynamics with space variation. The results of this study make clear that deep groundwater monitoring with numerical simulation for future water quality is important for understanding contaminant migration.

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Appendix

Table 4	Additional	information	of	ground	water	sampling	points
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No.	Sample ID	Latitude (degree decimal)	Longitude (degree decimal)	Well depth (m)
1	ch1	24.509	79.091	32
2	ch2	24.477	79.275	30
3	ch3	24.547	79.269	30
4	ch4	24.299	79.287	28
5	ch5	24.650	79.498	24
6	ch6	24.692	79.369	22
7	ch7	24.562	79.533	26
8	ch8	24.481	79.723	25
9	ch9	24.588	79.976	28
10	ch10	24.900	79.591	23
11	ch11	24.395	79.385	25
12	ch12	24.021	79.725	22
13	ch13	25.027	79.799	24
14	ch14	24.798	79.477	23
15	ch15	24.996	79.651	25
16	ch16	25.268	80.194	24
17	ch17	25.186	80.287	21
18	ch18	25.067	80.194	20
19	ch19	25.139	80.006	22
20	ch20	25.226	79.346	18
21	ch21	25.054	79.450	22
22	ch22	25.115	79.387	18
23	ch23	25.045	79.865	16
24	ch24	24.750	79.959	25
25	ch25	24.792	79.795	24
26	ch26	24.100	79.931	22
27	ch27	24.868	79.720	24

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