

Late Miocene Proboscidea (Mammalia) from Macedonia and Samos Island, Greece: preliminary results

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Abstract The abundant Late Miocene proboscidean remains of Greece have never been studied in detail and compared with those of Eurasia in order to determine their taxonomy and their biostratigraphical and palaeoecological significance. The first results of such study are given in this article. During the past decades, several new proboscidean specimens have been added to the old collections, significantly enriching the available material. The Axios Valley (Macedonia, Greece) proboscidean fossils belong mainly to two species of *Choerolophodon*: *C. anatolicus* of early Vallesian age and *C. pentelici* of late Vallesian–Turolian age. *Deinotherium giganteum* is rare and recognized only in the late Vallesian locality Ravin de la Pluie of Axios Valley. A zygodont form has also been identified in the Turolian of Axios Valley, attributed to “*Mammut*” sp. The Late Miocene localities of Nikiti (Macedonia, Greece) revealed several remains of *C. pentelici*, which are similar to the Turolian ones of Axios Valley. The Samos proboscidean collection includes *C. pentelici*, “*Tetralophodon*” *atticus*, “*Mammut*” sp. and *Deinotherium gigantissimum*.

The taxonomy of the Late Miocene peri-Mediterranean *Choerolophodon* is given, and the biostratigraphy and palaeoecology of the Greek Late Miocene proboscideans are discussed.

Keywords Late Miocene · Proboscidea · *Choerolophodon* · Axios Valley · Samos · Greece

Kurzfassung Die obermiozänen Proboscidier aus Griechenland, obwohl sie ziemlich verbreitet sind, wurden niemals im Detail untersucht und mit anderen aus Eurasien verglichen, um ihre Taxonomie, Biostratigraphie und Paläoökologie zu bestimmen. Die ersten Ergebnisse dieser Studie sind in diesem Artikel präsentiert. In den letzten Jahrzehnten kamen viele neue Proboscidier-Funde hinzu und bereichern erheblich das Material der alten Sammlungen. Der Großteil der Proboscidea-Fauna aus dem Axios Tal (Mazedonien, Griechenland) wurde der Gattung *Choerolophodon* mit zwei Arten zugewiesen: *C. anatolicus* aus dem unteren Vallesium und *C. pentelici* aus dem oberen Vallesium-Turolium. Die Art *Deinotherium giganteum* ist selten und wurde nur in der oberen Vallesium Fundstelle Ravin de la Pluie des Axios Tals nachgewiesen. Eine zygodonte Form ist aus dem Turolium des Axios Tals bekannt und wurde der Gattung “*Mammut*” zugewiesen. Aus den obermiozänen Fundstellen Nikiti (Mazedonien, Griechenland) stammen zahlreiche Reste von *C. pentelici*, die denjenigen aus dem Turolium des Axios Tals ähneln. Die Proboscidier-Fauna aus Samos enthält die Arten *C. pentelici*, “*Tetralophodon*” *atticus*, “*Mammut*” sp. und *Deinotherium gigantissimum*. Die Taxonomie der obermiozänen Arten von *Choerolophodon* um den Mittelmeerraum wird vorgestellt und die Biostratigraphie und Paläoökologie der griechischen Proboscidier aus dem Oberen Miozän werden besprochen.

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Schlüsselwörter Oberes Miozän · Proboscidea · *Choerolophodon* · Axios Tal · Samos · Griechenland

Abbreviations

AMNH	American Museum of Natural History (New York, USA)
AMPG	Athens Museum of Palaeontology and Geology, National and Kapodistrian University of Athens (Greece)
BSPG	Bayerische Staatssammlung für Paläontologie und Geologie, München (Germany)
FSL	Faculté des Sciences de Lyon (France)
GPIH	Geologisch-Paläontologisches Institut, Hamburg (Germany)
HGI	Hungarian Geological Institute (Budapest)
LGPU	Laboratory of Geology and Palaeontology, Aristotle University of Thessaloniki (Greece)
MGL	Musée Cantonal de Géologie, Lausanne (Switzerland)
NHMA	Natural History Museum of the Aegean, Mytilinii (Samos Island, Greece)
NHMUK	Natural History Museum United Kingdom (London)
NHMW	Naturhistorisches Museum Wien (Austria)
MNHN	Muséum National d'Histoire Naturelle (Paris, France)
SU	Palaeontology Museum of the Sofia University (Bulgaria)

Introduction

Proboscidean remains are common in the Late Miocene faunas, including several taxa; they are present in all Late Miocene mammal fossiliferous sites of Greece, with relatively abundant material. Although they are common in the Greek fossil record, they have never been thoroughly studied. Since a lot of new material has been recently unearthed, the need has arisen for a more complete study of their taxonomy and for an effort to interpret their relationships and their biostratigraphical and palaeoecological significance. In the present article the proboscidean remains from the Late Miocene localities of Macedonia and Samos Island (Greece) are studied. The material comes from both old collections and new ones, revealed during recent excavations.

The proboscideans of Macedonia studied in the present work come from two areas: the lower Axios Valley and Chalkidiki Peninsula (Fig. 1). A set of fossil mammal localities are known from Axios Valley (Fig. 1), belonging to three different formations, correlated to the whole Late

Miocene. The Axios Valley proboscidean collection is quite rich, including all discoveries of the last ~40 years of field expeditions in the area, and is housed at LGPUT. Additionally, the proboscideans of the old Arambourg's collection, housed at MNHN and described by Arambourg and Piveteau (1929), are included in this study.

The Nikiti localities are situated on Chalkidiki Peninsula (Fig. 1) and have provided a quite rich fossil mammal fauna (Koufos 2006a and refs. cited). Two fossiliferous sites are located in the Nikiti area, belonging to two different formations correlated to the late Vallesian and the early Turolian, respectively (Koufos et al. 1991; Koufos in press and refs. cited). The Nikiti collection includes several remains of proboscideans, housed at LGPUT.

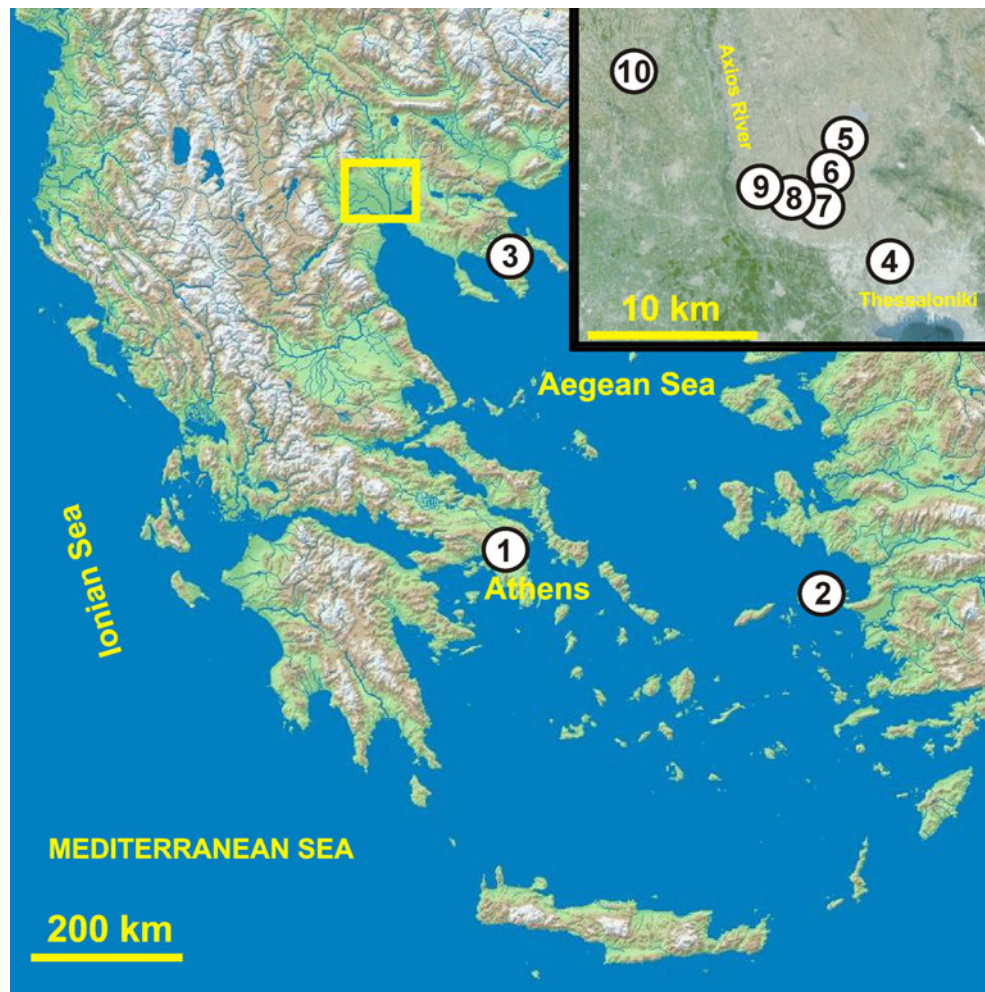
The fossil mammal localities of Samos Island (Fig. 1) are well known since the 19th century. Several old collections from Samos are housed at various museums and institutes and include many interesting proboscidean remains; the main ones (NHMW, MGL, HGI) have been included in the present study. Recent excavations in the Samos fossiliferous sites have provided a new rich mammal collection, housed at NHMA; the proboscidean remains of this new collection are few and have already been described (Konidaris and Koufos 2009). Detailed information on the stratigraphy, the old and new localities, as well as the age of the Samos faunas is given in Kostopoulos et al. (2009) and Koufos et al. (2009).

The majority of the studied proboscidean remains belong to the genus *Choerolophodon*. Study of this material allows us to follow and recognize the evolution of *Choerolophodon* during the whole Late Miocene in the peri-Mediterranean region. Therefore, it seems necessary to begin with the systematics of the Late Miocene choerolophodons of this area. For dental measurements of the studied choerolophodons see “Appendix”.

Taxonomy of the Late Miocene peri-Mediterranean choerolophodons

Taxonomical study of the choerolophodont remains from the Late Miocene of the peri-Mediterranean region has a long history. They were originally found in Pikermi (Attica, Greece) and referred to as *Mastodon pentelicus*, without further description (Gaudry and Lartet 1856). A few years later, the Pikermi material was described in detail under the name *Mastodon pentelici* (Gaudry 1862, 1867). Material belonging to “*Tetralophodon*” *atticus* was also included in this species (Gaudry 1867: pl. 23, figs. 2–4) (Tassy 1985). Subsequently, Schlessinger (1917, 1922) described juvenile choerolophodont material from Samos Island (Greece) and Maragheh (Iran) under the name *Mastodon (Choerolophodon) pentelici*. Since then,

Fig. 1 Map of Greece indicating the location of the fossiliferous sites studied in the text. 1 Pikermi, 2 Samos Island, 3 Nikiti, 4 Pentalophos, 5 Xirochori, 6 Ravin de la Pluie and Ravin des Zouaves 1, 7 Ravin des Zouaves 5, 8 Vathylakkos, 9 Prochoma, 10 Dytiko. *Inset* enlarged area of Axios Valley (localities 4–10). The map was taken from www.shaded-relief.com



all the Late Miocene choerolophodont material has been referred to as *C. pentelici* without any distinction. The first recognition of different evolutionary characters in the dental morphology of the Late Miocene choerolophodonts was given by Gaziry (1976), who studied Turkish material. Later, Tassy et al. (1989) studied a mandible from the Vallesian locality Eşme Akçaköy (Turkey) and recognized differences from the typical *C. pentelici*. The authors distinguished two subspecies, based on mandibular and dental morphology: *C. p. lydiensis* with deflected mandibular symphysis and less expressed choerolophodonty on the molars and *C. p. pentelici* with straight symphysis and more expressed choerolophodonty. A few years later, the study of the choerolophodonts from Kemiklitepe localities (Turkey) indicated two different evolutionary forms of *C. pentelici*: a primitive one in Kemiklitepe-D (KTD) correlated with the early Turolian, MN 11, and a more advanced one from Kemiklitepe A-B (KTA-B) correlated with the middle Turolian, MN 12 (Tassy 1994a).

Meanwhile, a new choerolophodont species, *C. anatolicus*, had been previously erected by Ozansoy (1965)

based on material from Yassiören (Middle Sinap, Turkey). This species has been regarded as synonymous with *C. pentelici* for a long time (Gaziry 1976; Tassy 1994a; Tassy et al. 1989). Its validity was recognized quite later, when new material from Sinap (Turkey) was described (Sanders 2003).

The Axios Valley choerolophodont material spans between biozones MN 9 to MN 13 and supports the separation into two distinct species.

Order Proboscidea Illiger, 1811

Superfamily Elephantoidea Gray, 1821

Genus *Choerolophodon* Schlesinger, 1917

Diagnosis (based on Tassy 1986): Narrow cranium with zygomatic arches slightly remote from the skull; strong development of the facial region, elongated in comparison with the cerebral region; palate extends behind the last functional molar; nasal fossa and orbits redressed above the posterior half (at least) of the last functional molar; crista

orbitotemporalis subvertical or oblique to the base and in front; large and deep symphyseal gutter; upper tusks curve upwards, lack enamel (tusks belonging to juvenile individuals bear an enamel cap), and the cross section is almost circular; lack of lower tusks; lack of dp2, p2, p3/P3, p4/P4; DP2 and dp3/DP3 bilophodont and intermediate molars trilophodont; in the dp3 the first posttrite halflophid posteriorly offset and connected with second pretrite halflophid; in the DP3 the second pretrite halflophid anteriorly offset; choerolophodont teeth: choerodonty, ptychodonty, cementodonty and additionally the intermediate molars and m3/M3 with chevrons of the loph(id)s except for the first one.

Type species: *C. pentelici* (Gaudry and Lartet, 1856) (Late Miocene, Southeastern Europe–West Asia)

Other species included:

C. kisumuensis (MacInnes, 1942) (Early to Middle Miocene, North and East Africa)

C. palaeindicus (Lydekker, 1884) (?Early Miocene, Bugti Hills, Pakistan, South Asia)

C. chioticus Tobien, 1980 (Middle Miocene, Chios Island, Greece, Southeastern Europe)

C. ngorora (Maglio, 1974) (Middle to Late Miocene, Kenya, East Africa)

C. zaltaniensis Gaziry, 1987 (Middle Miocene, Libya, North Africa)

C. anatolicus (Ozansoy, 1965) (Late Miocene, Southeastern Europe–West Asia)

C. corrugatus (Pilgrim, 1913) (Late Miocene, Pakistan, South Asia)

Choerolophodon pentelici (Gaudry and Lartet, 1856)

Synonyms

1856 *Mastodon pentelicus* Gaudry and Lartet: p. 273

1862 *Mastodon pentelici* Gaudry: p. 142

1867 *Mastodon pentelici* Gaudry: pl. 22, figs. 1–3; pl. 23, fig. 1

1911 *Mastodon angustidens* Bakalov: p. 37; pl. 1, figs. 1, 2; pl. 2, figs. 1–3; pl. 3, figs. 1, 2; pl. 4, figs. 1–3

1913 *Mastodon pentelici* Pavlow: p. 36; pl. 7, fig. 4

1914 *Mastodon pentelici* Pavlow: p. 184; pl. 8, fig. 16

1917 *Mastodon (Choerolophodon) pentelici* Schlesinger: p. 181; pl. 23, fig. 1; pl. 24, figs. 1, 2; pl. 25, figs. 1–4; pl. 26, figs. 1, 2; pl. 27, figs. 1–3; pl. 28, figs. 1–3; pl. 29, figs. 1–6; pl. 30, figs. 1–4; pl. 31, figs. 1, 2; pl. 32, figs. 1, 2; pl. 33, fig. 1

1922 *Mastodon (Choerolophodon) pentelici* Schlesinger: p. 142; pl. 20, figs. 1, 2; pl. 21, fig. 1

1924 *Mastodon pentelici* Mecquenem: p. 137; pl. 1, figs. 1–5, 7, 8, 10–14

1929 *Mastodon pentelici* Arambourg and Piveteau: p. 74; pl. 2, fig. 1

1936 *Trilophodon (Choerolophodon) pentelicus* Osborn: p. 263; fig. 205–208

1950 *Mastodon pentelici* Lehmann: p. 221; pl. 22, fig. 109

1952 *Mastodon pentelici* Senyürek: p. 451; figs. 6, 7

1957 *Mastodon pentelici* Čirić: p. 13; pl. 9, 10

1962 *Trilophodon angustidens* Bakalov and Nikolov: p. 98; pl. 48; pl. 49, figs. 1–3; pl. 50, figs. 1–3; pl. 51, figs. 1, 2

1962 *Trilophodon (=Choerolophodon) pentelicus* Bakalov and Nikolov: p. 105; pl. 52

1976 *Choerolophodon pentelici* Gaziry, p. 76; pl. 12, fig. 7

1977 *Mastodon (Choerolophodon) pentelici* Koufos: p. 98; pl. 1, figs. a, b

1980 *Choerolophodon* cf. *pentelici* Thomas et al.: p. 272

1980 *Gomphotherium pentelici* Koufos: p. 93; figs. 34, 35; pl. 11, fig. 1

1980 *Tetralophodon longirostris* Koufos: p. 106; fig. 37; pl. 11, fig. 2

1985 *Choerolophodon (Synconolophus) serridentinoides* Apostolov and Nikolov: p. 126; pl. 3, figs. 1, 2; pl. 4, figs. 1–3

1994 *Choerolophodon pentelici* Tassy: p. 144; p. 146; pl. 1, figs. a, b; pl. 3, figs. a, b

2003 *Choerolophodon* sp. Theodorou et al.: p. 523; fig. 4f

2005 *Choerolophodon pentelici* Tassy: p. 708; figs. 1, 2

2009 *Choerolophodon pentelici* Konidaris and Koufos: p. 140; pl. 1, figs. 1, 2

Type locality: Pikermi, Attica, Greece.

Age: Middle Turolian, MN 12, Late Miocene.

Lectotype: Juvenile skull with DP2–DP4 dex and DP4 sin and associated mandible with dp3–dp4 dex and dp4 sin, MNHN-PIK-3665 (Fig. 2), described by Gaudry (1862: p. 142–152) and figured by Gaudry (1867: pl. 22, figs. 1–3, pl. 23, fig. 1).

Diagnosis: Large choerolophodont species; elongated skull with moderately inclined facial region and low cerebral region; orbits situated at the top of the skull and far behind the last molar in function, the anterior zygomatic process is situated behind the dentition, at the level of the anterior border of the choane, and the whole perinasal area is redressed; distant temporal lines and wide parietal region; the symphysis is situated ventrally at the extension of the horizontal ramus, in both juvenile and adult specimens; anterior border of ascending and horizontal ramus form obtuse angle; in mandibles bearing the m3 there is a large retromolar gap; upper tusks emerge sub-horizontally and outwards and then curve upwards; deciduous teeth are

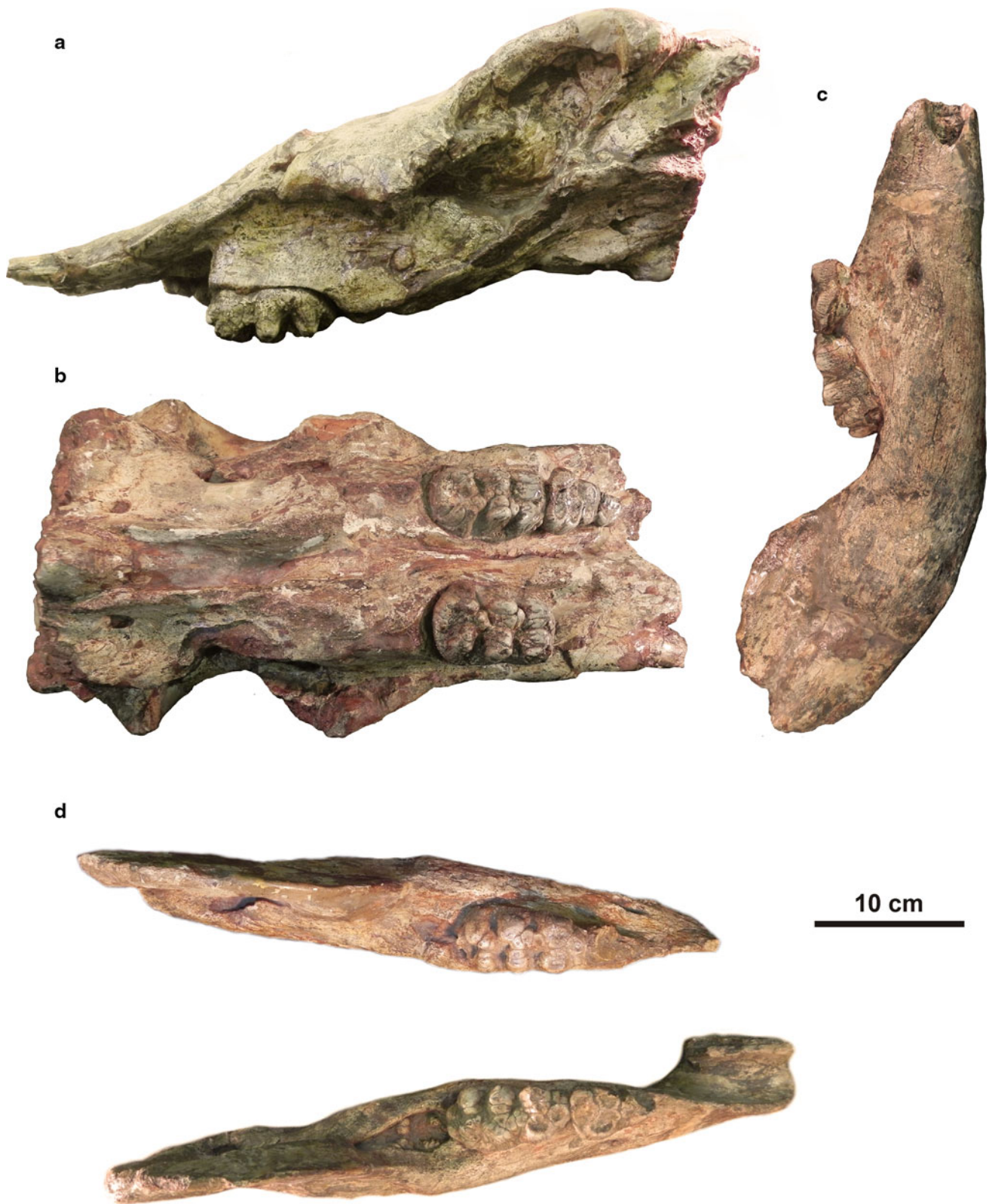


Fig. 2 *Choerolophodon pentelici*, Pikermi (PIK), Attica, Greece, middle Turolian, MN 12. Skull and associated mandible, LECTOTYPE, MNHN-PIK-3665; **a, c** lateral, and **b, d** occlusal view

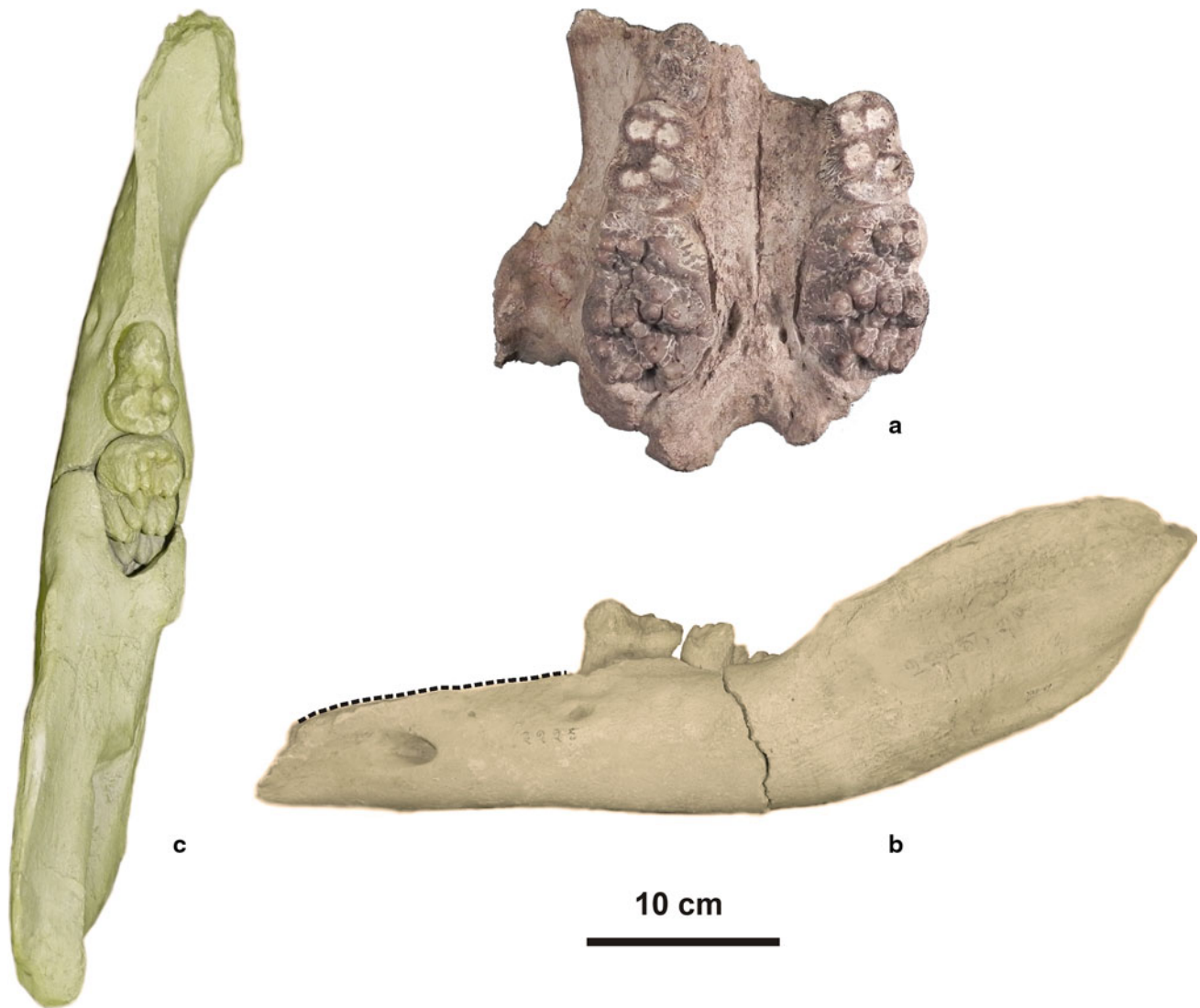


Fig. 3 *Choerolophodon anatolicus*, Yassiören (Middle Sinap), Turkey, Vallesian, MN 9–10. **a** maxilla, LECTOTYPE, MNHN-TRQ-1000, occlusal view; **b, c** left mandibular fragment (cast housed at MNHN-YAS-49, 50), **b** lateral, and **c** occlusal view

large; the posterior cingulum of the DP3 is well developed, and the metacone is separated from the posttrite cusplet of the posterior cingulum; the posterior cingulum of the dp3 is prominent and isolated from the second lophid in the more advanced morphs, forming a third lophid; five lophids or four lophids with an isolated posterior “heel” in the m3; the M3 bears four lophids; well-expressed choerodonty, ptychodonty and cementodonty.

Important localities: Pikermi, Samos Island, Xirochori, Ravin de la Pluie, Ravin des Zouaves 5, Dytiko (Greece), Maragheh (Iran), Kemiklitepe (Turkey), Burgas (Bulgaria).

Choerolophodon anatolicus (Ozansoy, 1965)

1965 *Trilophodon* (*Choerolophodon*) *anatolicus* Ozansoy: p. 47; pl. 9, figs. 1, 2

1976 *Choerolophodon pentelici* Gaziry: p. 46, 49, 52; pl. 5, figs. 5–7; pl. 6, figs. 1–5; pl. 7, figs. 1–4; pl. 8, figs. 1–3; pl. 9, figs. 1–4, 6–9; pl. 10, figs. 1–6; pl. 11, figs. 1–3; pl. 12, figs. 1, 2

1989 *Choerolophodon pentelici lydiensis* Tassy et al.: p. 2146; figs. a, b

2003 *Choerolophodon anatolicus* Sanders: p. 208; figs. 10.5, 10.6a–c, 10.8a–b

Type locality: Yassiören (Middle Sinap), Turkey.

Age: Vallesian, Late Miocene; MN 9 according to Mein (1990), de Bruijn et al. (1992)

Lectotype: Maxilla with DP2–DP4 dex and DP3–DP4 sin (Figs. 3a, 4a). We designate the lectotype of *C. anatolicus*, since Ozansoy (1965: p. 47–48, pl. 9, figs. 1, 2) did not

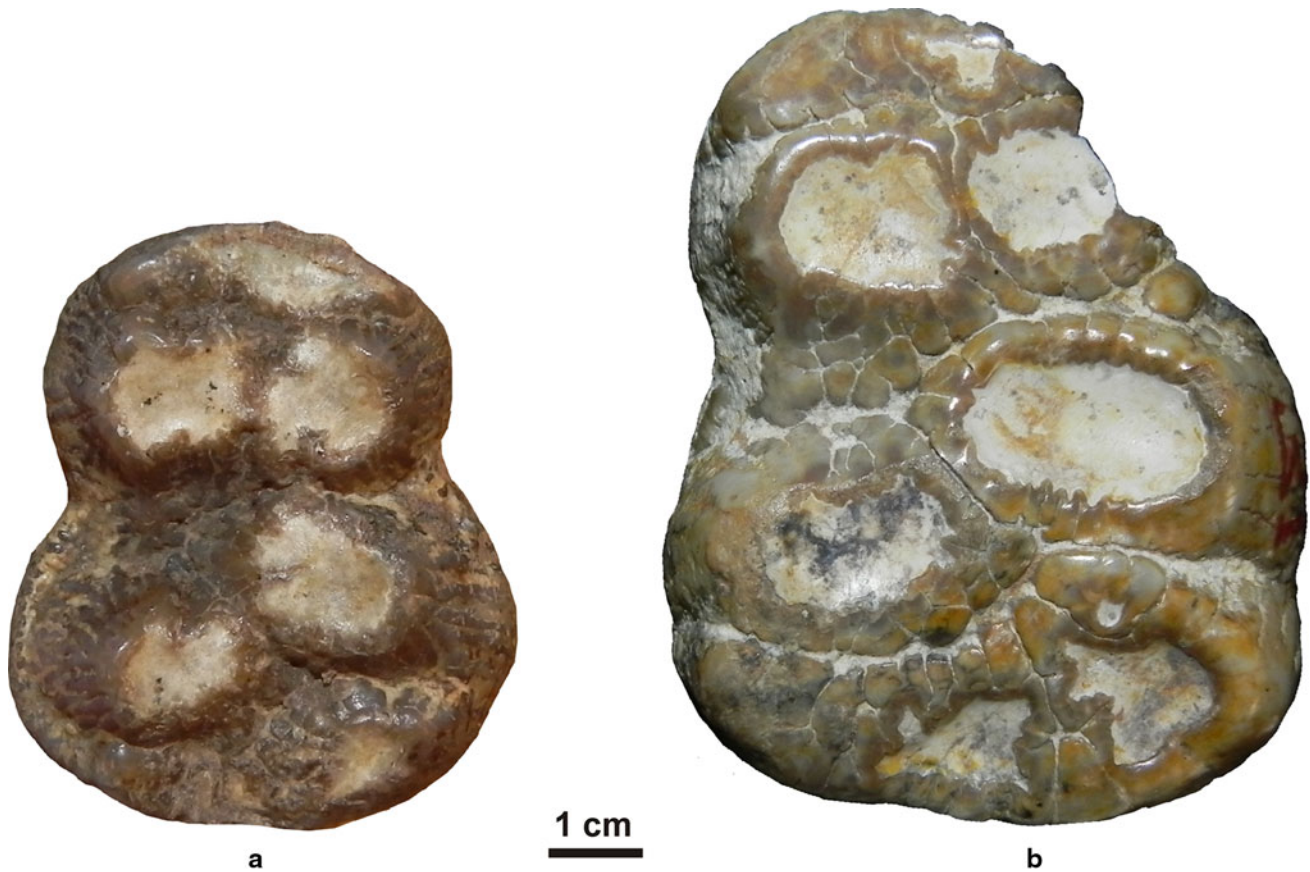


Fig. 4 **a** DP3 dex, *C. anatolicus*, Yassiören (MNHN-TRQ-1000); **b** DP3 sin (reversed), *C. pentelici*, Pikermi (MNHN-PIK-1705)

determine the holotype of this species, and the included material originated from different fossiliferous layers and did not thus belong to the same individual. The maxilla presents diagnostic morphological features which differentiate it from the type species *C. pentelici*. The maxilla is housed at MNHN (MNHN-TRQ-1000).

Diagnosis: Small choerolophodont species; elongated skull with moderately inclined facial region and low cerebral region; orbits situated at the top of the skull and just behind the last functional teeth; anterior zygomatic process situated anterior of the choanae at the level of the last molar; moderate redressed perinasal area; distant temporal lines and wide parietal region; in juvenile specimens the mandibular symphysis is situated ventrally horizontal, while in adult ones it is deflected downwards; ascending and horizontal ramus form obtuse angle, and retromolar gap is present; upper tusks emerge sub-horizontally and outwards and then curve upwards; deciduous teeth are small; in DP3 moderate development of the posterior cingulum and metacone connected to the posttrite cusplet of the posterior cingulum; in dp3 weak development of the posterior cingulum, which is connected to the

second pretrite lophid; m3 with four or five lophids; in some tetralophodont m3 the posterior cingulum is connected to the fourth lophid; M3 with four lophids; moderate to well-expressed choerodonty, ptychodonty and cementodonty.

Differential diagnosis: *C. anatolicus* differs from the type species in the ventrally deflected adult mandibular symphysis, the more anteriorly situated orbit and anterior zygomatic process and the generally less redressed perinasal area, the weaker posterior cingulum in the dp3/DP3 connected to the second loph(id) (Fig. 5), the smaller deciduous teeth and the variable presence of tetralophodont m3's with posterior cingulum connected to the fourth lophid (primitive character observed in earlier choerolophodonts and absent in *C. pentelici*); the molars tend to be smaller in size.

In agreement with Tassy (1983) and Tassy et al. (1989), the ventral margin of the mandibular symphysis must be considered as a distinctive character among choerolophodont species. *C. anatolicus* presents a mandibular symphysis that is horizontal in juvenile individuals (Fig. 3b) and deflected in adult ones (Tassy et al. 1989: fig. b).



Fig. 5 DP3 and dp3 of *Choerolophodon* from various localities of Greece. **a** *C. anatolicus*, LGPUT-PNT-93, DP3 sin (reversed); **b** *C. pentelici*, LGPUT-XIR-21, DP3 sin (reversed); **c** *C. pentelici*,

LGPUT-NIK-1605, DP3 dex; **d** *C. anatolicus*, LGPUT-PNT-158, dp3 sin (reversed); **e** *C. pentelici*, LGPUT-RPI-259, dp3 dex; **f** *C. pentelici*, LGPUT-NIK-1590, dp3 dex

According to Tassy et al. (1989), in elephantiformes the rostrum of juvenile specimens is horizontal, regardless of the deflection in adult ones; for example, Tassy (1994b) mentions the straight ventrally mandibular symphysis in juvenile specimens of *Gomphotherium angustidens* (Cuvier, 1817) and the deflection in adult ones. Tassy et al. (1989) proposed that the absence of inclination in adult *C. pentelici* could be attributed to pedomorphosis and cannot be interpreted as retention of a primitive character.

Important localities: Sinap-Middle Member, Eşme Akçaköy, Akin, Kayadibi (Turkey), Pentalophos (Greece).

Axios Valley localities

The fossiliferous sites of Axios Valley were discovered in the beginning of the 20th century; the first collection is housed at MNHN and is labelled as “Salonique” (Arambourg and

Piveteau 1929). This collection originated from several stratigraphic horizons, and its mixing hampers comparisons. The new series of excavations in Axios Valley started in 1972 (de Bonis et al. 1973) and provided a large number of fossils, including numerous proboscidean ones. The study of the Late Miocene deposits of Axios Valley suggests three different Formations (Koufos in press and refs. cited):

1. Nea Mesimvria Formation, which includes the localities Pentalophos 1 (PNT), Xirochori 1 (XIR), Ravin de la Pluie (RPI) and Ravin des Zouaves 1 (RZ1)
2. Vathylakkos Formation, which includes the localities Ravin des Zouaves 5 (RZO), Prochoma 1 (PXM), Vathylakkos 1, 2, 3 (VLO, VTK, VAT)
3. Dytiko Formation, which includes the localities Dytiko 1, 2, 3 (DTK, DIT, DKO).

More information on the stratigraphy, fauna and age of the Axios Valley Late Miocene mammal localities is given in Koufos (2006a) and Koufos (in press and refs. cited).

Locality Pentalophos 1 (PNT): The locality Pentalophos has revealed a rich proboscidean material, including juvenile and adult specimens. All of them show typical choerolophodont features and can be attributed to *Choerolophodon*. Previous references on the presence of *Tetralophodon longirostris* (Kaup, 1832) (de Bonis and Koufos 1999; Koufos 2006a) were in fact based on molars belonging to adult individuals of *Choerolophodon*. The dp3/DP3, which are used by several authors (e.g. Tassy 1994a) for the definition of evolutionary stages among Late Miocene Eurasian choerolophodonts, show the most primitive characters among the Axios Valley material (Fig. 5a, d). The morphology of these teeth fits with that of *C. anatolicus* from Yassiören (Middle Sinap, Turkey) described by Ozansoy (1965) (Figs. 3, 4a) and with that of unpublished material from this locality housed at MNHN (MNHN-YAS-24, 44, 46). Moreover, the dimensions of all the PNT juvenile teeth (dp3/DP2-DP4) are close to those of *C. anatolicus* (Ozansoy 1965; Sanders 2003) (Fig. 6) and confirm their specific attribution. Concerning the adult material, an unworn m3 (PNT-154) with four lophids resembles in morphology and size those from Eşme Akçaköy, Turkey (GPIH-T.K.-Nr. 1792 and cast MNHN-EA-1) described by Gaziry (1976) and Tassy et al. (1989). Unfortunately, an available senile mandible (PNT-156), bearing an almost worn m3 (possibly with five lophids), lacks the symphysis. Based on:

1. The more primitive morphology of the juvenile material from Yassiören [MN 9, according to Mein (1990), de Bruijn et al. (1992) and Sanders (2003)]
2. The teeth and mandibular characters of the material from Eşme Akçaköy [early Vallesian (Sickenberg et al. 1975; Mein 1990; de Bruijn et al. 1992; Fortelius 2011)]

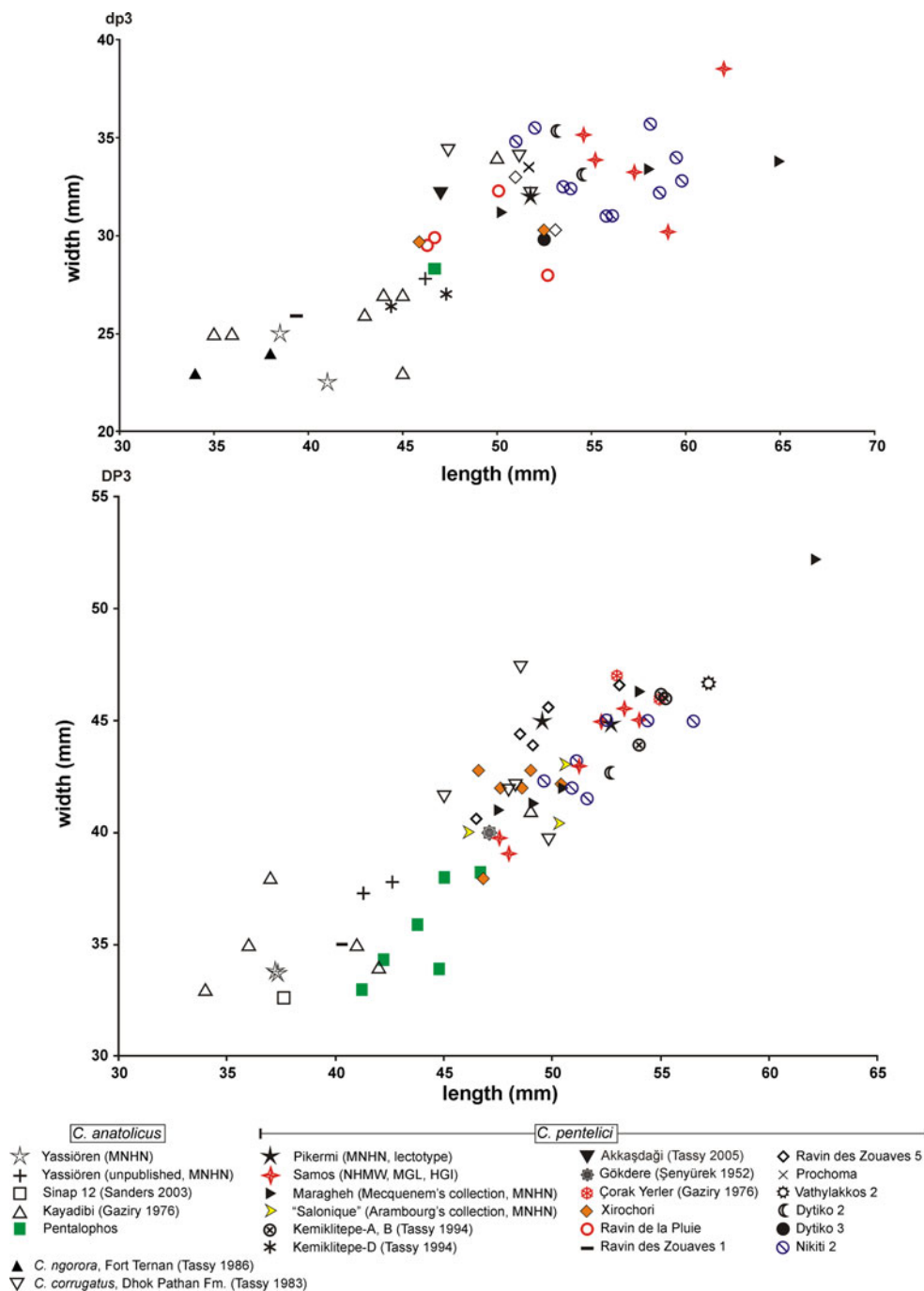
3. The more primitive morphology of the adult skull from Akin (Turkey) [Vallesian (Gaziry 1976; Tassy 1985; Sickenberg et al. 1975)]
4. The PNT material and its age,

we recognize *C. anatolicus* as a valid species and we agree with Sanders (2003) in including in this species *C. p. lydiensis*. The available data suggest that the biostratigraphic range of this species is restricted to the early Vallesian (MN 9) and not extended to the early Turolian as suggested by Sanders (2003).

As mentioned by Koufos (in press), the Pentalophos fauna is the most peculiar Late Miocene one. Since the PNT choerolophodont material shows more primitive characters than the XIR one, a locality which corresponds to the lower part of late Vallesian, ~9.6 Ma (see below), an older age (early Vallesian) is possible for Pentalophos. This age is in agreement with data from the rest of the fauna; for instance, the PNT giraffids and hipparions are more primitive than the XIR and RPI ones (Geraads 1989; Koufos 2000; de Bonis and Bouvrain 2003), and the PNT aardvark *Orycteropus pottieri* Ozansoy, 1965 is similar to the Vallesian one from Middle Sinap (de Bonis et al. 1994). The choerolophodont material from Pentalophos gives thus additional evidence for an early Vallesian age for the locality.

Locality Xirochori 1 (XIR): The locality XIR is also placed into the Nea Mesimvria Fm, and as mentioned above, its choerolophodonts bear more evolved characters than those from the PNT material, which places them closer to the typical species *C. pentelici*. The dp3/DP3 (Fig. 5b) clearly have a more advanced morphology and stoutness in the posterior cingulum, and the dimensions of the juvenile teeth have on average an intermediate position between Pentalophos and the Turolian samples (Fig. 6). The only available m3 (XIR-6) bears five lophids and is morphologically and metrically similar to the m3 from Sinap 83, Turkey [described by Sanders (2003: fig. 10.7); estimated age ~9.45 Ma according to Kappelman et al. (2003)], being larger than the teeth from Yulafli (Turkey, MN 10) described by Geraads et al. (2005). A juvenile skull (XIR-23) presents slightly more primitive morphology in comparison with the type material of *C. pentelici* from Pikerimi in the less redressed perinasal area, whereas a juvenile mandible (XIR-20) shows similar morphology to mandibles of *C. pentelici*. Based on the dp3/DP3 morphology and size, the XIR choerolophodont material can be determined as *C. pentelici*. The XIR fauna is correlated with the late Vallesian (MN 10), and the magnetostratigraphic record suggests an estimated age of ~9.6 Ma (Koufos in press and refs. cited). The dental morphology of the XIR *Choerolophodon* and the age of the locality suggest that it corresponds to the most primitive form of *C. pentelici*.

Fig. 6 Scatter diagrams (length–width) of the lower and upper third deciduous premolars (dp3/DP3) of *Choerolophodon* from various localities



Localities Ravin de la Pluie (RPI) and Ravin des Zouaves 1 (RZ1): Both fossiliferous sites are located in the upper part of Nea Mesimvria Fm and belong to the same horizon (Koufos in press and refs. cited). The RPI choerolophodons (Fig. 5e) are similar to the XIR ones, corresponding to the same evolutionary stage and suggesting a late Vallesian (MN 10) age for the former locality. The RZ1 material includes a juvenile skull and the associated mandible (both bearing the complete juvenile dentition). The

dental dimensions of these specimens are rather small (Fig. 6), but unfortunately both dp3/DP3 are totally worn and do thus not provide any morphological data. However, the morphology of the skull and the mandible are comparable to the already known material of *C. pentelici* from Pikermi and Samos. Apart from the choerolophodont fossils, an isolated deinotherere p4 is also known from RPI. Its morphology and size ($74.7 \times 51.1+$ mm) are similar to those of the Montredon specimens (France, late Vallesian) and the

large specimens from Dinotheriensande (Germany, Vallesian) [described by Tobien (1988) and Gräf (1957) respectively], permitting its attribution to *Deinotherium giganteum* Kaup, 1829. The previously mentioned presence of *Tetralophodon* in RPI (Koufos 1980; de Bonis and Koufos 1999; Koufos 2006a) is not confirmed. The rest of the fauna is diverse in taxa and indicates a late Vallesian (MN 10) age; the magnetostratigraphic record suggests an estimated age of ~9.3 Ma (Koufos in press and refs. cited).

Locality Ravin des Zouaves 5 (RZO): This locality is situated in the lower part of the Vathylakkos Fm and includes several proboscidean remains. The choerolophodont material is quite rich, containing mostly juvenile specimens. The dp3/DP3 show advanced morphology, similar to that of the Pikermi and Samos teeth, and all deciduous teeth are large, showing dimensions similar to the Turolian samples from Pikermi, Samos, Akkaşdağı, Kemiklitepe A, B (Fig. 6). Therefore, both morphology and dimensions of the RZO juvenile material support a Turolian age. A sub-adult mandible with the symphysis situated ventrally at the extension of the horizontal ramus confirms the attribution of the material to *C. pentelici*.

Apart from *Choerolophodon*, a mammutid species has already been identified by Koufos (1980) as “*Mammut tapiroides*”. It concerns a juvenile skull with clearly developed zygodont characters in the deciduous dentition, thus permitting its attribution to “*Mammut*” [for the use of quotation marks for the generic name see Markov (2004)]. The particular skull and its teeth are similar to those of a juvenile individual from Pikermi (NHMUK-M 10104), described by Tassy (1985) and identified as *Zygodon* cf. *borsoni* (Hays, 1834). The RZO skull is also similar (comparison only based on the provided photo) with a juvenile skull from Belka (Ukraine, MN 12) identified as *Turicius turicensis* (Schinz, 1824) by Korotkevich (1988: pl. 25, fig. b). Furthermore, the dental elements of the RZO skull are similar to those of a juvenile maxilla from Halmypotamos (Greece, MN 12) described by Melentis (1967: pl. 6, figs. 1–3) as “*Mastodon (Zygodon) tapiroides*”. As has been demonstrated by Markov (2008), the Turolian “*Mammut*” differs from the Pliocene “*M.*” *borsoni* and belongs to “*M.*” *obliquephus* (Mucha, 1980). This species was originally erected for a mandible with a long symphysis and “*borsoni*” teeth from Romanovka (Ukraine), of uncertain age. A mandible from the Turolian locality of Ahmatovo (Bulgaria) (Nikolov and Kovačev 1966: pl. 2, fig. 1), which is similar to that of “*M.*” *obliquephus*, reinforces the validity of the taxon. The

distinctive character between “*M.*” *obliquephus* and “*M.*” *borsoni* is the length of the mandibular symphysis. Since no mandible is known from RZO, the studied skull is referred to as “*Mammut*” sp. The RZO fauna is older than the Pikermi one, and it is correlated to the early Turolian (MN 11); magnetostratigraphic study suggests an estimated age of ~8.2 Ma (Koufos in press and ref. cited).

Locality Prochoma (PXM): The locality PXM is placed in the Vathylakkos Fm. Its proboscidean remains include some specimens of *Choerolophodon*. The large size of the lower deciduous teeth (Fig. 6), which are metrically similar to the type material of *C. pentelici*, as well as the size of two m1's, which is similar to that of the *C. pentelici* m1's from Maragheh (Iran, Mecquenem's collection, MNHN-MAR-4006, 4143), suggest the attribution of the PXM species to *C. pentelici*. The PXM fauna is correlated to the middle Turolian (MN 12), and the magnetostratigraphic record suggests an estimated age of ~7.4 Ma (Koufos in press and refs. cited).

Locality Vathylakkos 2 (VTK): This locality is also situated in the Vathylakkos Fm. The proboscideans are scanty, including only two juvenile maxillary fragments. The morphology of the deciduous teeth, especially of the DP3, which is the largest among the studied ones (Fig. 6), confirms the presence of *C. pentelici* in the locality. The fauna of VTK is also dated to the middle Turolian (MN 12) (Koufos in press and refs. cited).

Localities Dytiko 2 (DIT) and Dytiko 3 (DKO): These localities are placed in the Dytiko Fm and provided some proboscidean remains. Choerolophodonts are represented by very juvenile specimens. The morphology and size of their dp3/DP3 (Fig. 6) are similar to those of *C. pentelici* from Samos and Kemiklitepe A, B, allowing their attribution to this species. The sole specimen from DKO is a hemimandible preserving only the dp3. The morphology of the mandible, as well as the morphology and dimensions of the dp3 (Fig. 6), indicate *C. pentelici*. The complete Dytiko fauna is younger than the Pikermi one and is correlated to the late Turolian (MN 13) (Koufos in press and refs. cited).

Arambourg's collection (MNH): Arambourg and Piveteau (1929) list *Mastodon pentelici* from the localities Ravin X (R. X), Ravin des Zouaves, Ravin R. Ar., Ravin C, Ravin de Vatilük ou Ravin G (=Vathylakkos 3, VAT), Ravin de Konikovo (=DTK). Unfortunately, the authors do not provide details for the exact localities of the specimens. There is only one illustration (Arambourg and Piveteau

1929: pl. II, fig. 1) of a juvenile mandible bearing the dp4's (MNHN-SLQ-2), which in the MNHN collection is labelled as originating from R. X. The symphysis is situated at the extension of the horizontal ramus, and the dimensions of the dp4's are similar to the known Turolian ones. Additionally, a DP4 (MNHN-SLQ-3) labelled as coming from Ravin X is close metrically to the type material of *C. pentelici* from Pikermi, indicating a Turolian age. The rest of the R. X fauna also suggests a Turolian age for the locality (Koufos 2006a). Other proboscidean remains of the Arambourg's collection lack any locality indication; however, they can be attributed to *C. pentelici* due to the dimensions of the teeth (Fig. 6) and the morphology of the DP3 (e.g. MNHN-SLQ-1122).

Nikiti localities

The fossiliferous localities of Nikiti (Chalkidiki Peninsula, Macedonia) were discovered in 1990 (Koufos et al. 1991). There are two localities, Nikiti 1 (NKT) and Nikiti 2 (NIK), which are located in the Nikiti Formation (Koufos in press and refs. cited). The NKT represents the oldest one, dated to the late Vallesian, MN 10 (Kostopoulos and Koufos 1999; Koufos 2000; in press), whereas NIK, situated about 20 m above the latter, is dated to the early Turolian (Kostopoulos and Koufos 1999; Vlachou and Koufos 2002). The locality NIK provided a rich mammalian fauna, the overall study of which is in progress. No proboscidean material is known from NKT, apart from some postcranial elements.

Locality Nikiti 2 (NIK): The sole known proboscidean species from NIK is *C. pentelici*, represented by numerous juvenile dental remains. The morphology (Fig. 5c, f) and the size of the dp3/DP3 (Fig. 6) are typical for the Turolian choerolophodons. Moreover, several juvenile mandibles have the same morphological features as those of the type material of *C. pentelici* from Pikermi. Apart from these remains, an isolated tusk belonging to an adult individual was found about 100 m away from NIK at the same level. The tusk is long, strongly curved, lacks enamel and has circular cross section. The strong curvature clearly separates it from the straighter tusks of “*Mammut*” and *Tetralophodon* and suggests its attribution to *Choerolophodon*.

Samos localities

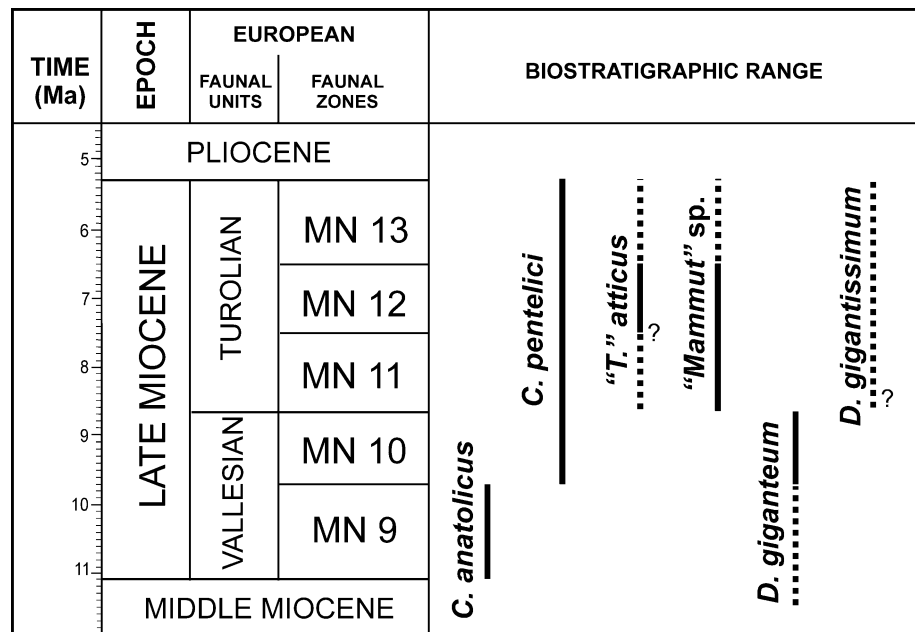
The mammal localities of Samos Island are among the richest Late Miocene ones, and several scientists have unearthed important material (e.g. E. Forsyth Major, B. Brown). Most specimens of the old collections lack

stratigraphic indications, and their comparison and dating are thus difficult. During a recent field campaign, which started in 1993, numerous fossils have been unearthed, and extensive stratigraphic work has been carried out aiming at correlation of the various old and new fossiliferous sites (Koufos and Nagel 2009). The most recent stratigraphic interpretations for Samos are given by Kostopoulos et al. (2009). The fossiliferous localities of Samos are situated in the Mytilinii Formation, which is dated to Turolian. The proboscidean remains from the recent excavations come from two sites, Mytilinii-1A (MTLA) and Mytilinii-1B (MTLB), dated to the middle Turolian, MN 12 (Koufos et al. 2009).

Forsyth Major (1894), in his catalogue for the MGL collection of Samos, referred to the taxa *Mastodon pentelici*, *Mastodon turicensis* and *Deinotherium* sp., originating from “Adriano” site. A detailed study of the Samos choerolophodons was given by Schlesinger (1917, 1922), who invented the subgenus *Choerolophodon* and ascribed several juvenile skulls and mandibles to *Mastodon (Choerolophodon) pentelici*, which are housed at NHMW and HGI. Osborn (1936) described a partial skull belonging to *C. pentelici*, housed at AMNH. Later, Lehmann (1950) described some more material stored at BSPG, including two juvenile maxillary fragments of *C. pentelici*, as well as a skull and a mandible that were attributed to *Mastodon longirostris* [all the BSPG proboscidean material from Samos was destroyed during World War II (pers. communication Heissig 2011)]. Bernor et al. (1996) reported *D. giganteum*, *M. borsoni*, *C. pentelici* and *Stegotetralodon grandincisivus* (Schlesinger, 1917) from the Main Bone Beds of the Mytilinii Formation. Tassy (2005) mentioned that the latter authors listed *S. grandincisivus* instead of “*T.*” *atticus* (Wagner, 1857) for Pikermi, which might be the case also for Samos. Recently, Konidaris and Koufos (2009) described *C. pentelici* from MTLB and *Zygodolophodon turicensis* possibly from MTLA.

Apart from the certain presence of *C. pentelici* in Samos, the more detailed studies during the past few years provided additional results. The collection housed at MGL includes the species *C. pentelici* and *D. gigantissimum* Stefanescu, 1892. The latter species is represented by a partial right mandible with the dp3 (MGL-S 1048) and the erupting dp4 (MGL-S 380), associated to each other and forming the horizontal ramus, and a left mandibular fragment with the unerupted dp4 (MGL-S 381). Possibly both specimens belong to the same individual, but this cannot be confirmed. The presence of *Zygodolophodon turicensis* or any other mammutid in the MGL collection has not been confirmed. The size of the deinothere dp3 (57.6 × 42.3 mm) places it among the largest known so far from Europe, surpassing the Late Miocene dp3's from Montredon (France, FSL-210393),

Fig. 7 Biostratigraphic distribution of the Late Miocene proboscidean taxa of Greece mentioned in the text. The “?” in “*T.* *atticus*” and *D. gigantissimum* is placed because both species are known from the old collections of Samos, thus indicating a general Turolian age. However, “*T.* *atticus*” is known from Pikermi (type locality) dated to MN 12



Prottes (Austria, Huttunen 2002), Kayadibi (Turkey, Gaziry 1976) and Nessebar (Bulgaria, SU-190, 191). The only larger dp3 is known from the Turolian locality of Baltavár (Hungary, HGI-Ob-3150). The deinotherium material from Nessebar was originally ascribed to *Prodeinotherium* by Bakalov (1914) and Bakalov and Nikolov (1962), but Markov (2004, 2008) records it as juvenile specimens of *D. gigantissimum*. Accepting the validity of *D. gigantissimum*, as demonstrated by Tarabukin (1974) and Markov (2008), and keeping in mind the large dimensions of the Samos specimen, we allocate it to this species.

As far as the tetralophodonts from Samos housed at BSPG are concerned, their direct observation is impossible as they have been destroyed. Based on the description and illustration of Lehmann (1950: pl. 22, fig. 108), their attribution to “*T.* *atticus*” seems plausible [for the use of quotation marks for the generic name see Tassy (2005)]. It should be mentioned, however, that the morphology of the juvenile teeth of “*Mastodon*” *grandincisivus* [an amebelodont, which reached tetralophodonty in parallel to tetralophodont gomphotheres (Tassy 2005)] remains unclear. Moreover, this species shares with “*T.* *atticus*” the same geographic and biostratigraphic distribution, coexisting with it in several localities (Markov 2008). Therefore, a more detailed work is necessary for this taxonomic problem. Apart from the BSPG material, an unpublished juvenile left mandible is housed at NHMW (NHMW-V 49), labelled as *C. pentelici*. Unfortunately the juvenile dentition is inside the mandibular bone, preventing any observation of its morphology, but the mandible has a strong symphysis, which bears a lower

juvenile tusk. The morphology of the mandible excludes any affinities with other elephantoids present at Samos (*Choerolophodon*, “*Mammut*”) and indicates similarities with “*T.* *atticus*”.

The other issue concerning the Samos proboscideans is the mammutid material. The only known published material is an isolated dp4 described by Konidaris and Koufos (2009). Direct studies of comparative juvenile material, including specimens of *Z. turicensis* from Simorre (France, MNHN-Si 11) and Innersdorf (Germany, BSPG-1916 I 1), the Pikermi material (MNHN-PIK-3613, NHMUK-M 10104) and “*M.* *borsoni*” from Valčedram (Bulgaria, SU) [this specimen was originally identified as *Z. tapiroides* by Bakalov and Nikolov (1962) but later as “*M.* *borsoni*” by Markov (2004)], permit the attribution of the Samos dp4 to “*Mammut*”, based on the clearly developed zygodont characters, the height of the crown in the interlophids on the lingual side [for this character see Tassy (1985)] as well as the similar dimensions to the dp4 from Pikermi (NHMUK). Therefore, the dp4 of Samos should be referred to as “*Mammut*” sp.

Biostratigraphic and palaeoecological remarks

The current knowledge on the Late Miocene proboscideans from Macedonia and Samos can provide useful biostratigraphic information. The species *C. anatolicus*, present in the locality Pentalophos, characterizes the early Vallesian (Fig. 7). No other locality of this age with large mammals is known from Greece (Koufos 2006a). During the late

Table 1 Proboscidean taxa from various Late Miocene localities of Greece

Localities	Axios Valley localities										Samos localities			Nikiti
	PNT	XIR	RPI	RZ1	RZO	PXM	VTk	DIT	DKO	R.X	SAM	MTLA	MTLB	NIK
European faunal units	Early Vallesian		Late Vallesian		Early Turolian	Middle Turolian	Late Turolian	Late Turolian		Turolian	Turolian	Middle Turolian	Early Turolian	
European faunal zones	MN 9	MN 10		MN 11	MN 12	MN 13		?	?		MN 12		MN 11	
<i>Deinotherium giganteum</i>			+											
<i>Deinotherium gigantissimum</i>											+			
<i>Choerolophodon anatolicus</i>	+													
<i>Choerolophodon pentelici</i>		+	+	+	+	+	+	+	+	+	+		+	+
“Mammut” sp.					+							+		
“Tetralophodon” <i>atticus</i>											+			

PNT Pentalophos, XIR Xirochori, RPI Ravin de la Pluie, RZ1 Ravin des Zouaves 1, RZO Ravin des Zouaves 5, PXM Prochoma, VTK Vathylakkos 2, DIT Dytiko 2, DKO Dytiko 3, R. X Ravin X, SAM Samos old collections (includes BSPG, HGI, MGL and NHMW), MTLA Mytilinii-1A, MTLB Mytilinii-1B, NIK Nikiti 2

Vallesian, two proboscideans were present, *C. pentelici* and *D. giganteum*, associated at Ravin de la Pluie (Fig. 7; Table 1). The absence of a tetralophodont elephantoid can be attributed either to the scarcity of localities of this age or to limited sampling (which is not the case, since RPI has brought to light a rich collection) or to ecological reasons. The coexistence of *D. giganteum* and *T. longirostris* was common during the Vallesian of Central and Western Europe [*Choerolophodon* did not migrate to this area (Koufos 2003)], but very scarce in Southeastern Europe (where *Choerolophodon* is the dominant proboscidean species); for example, the Middle Sinap collection does also not include any tetralophodont elephantoid. However, in Yulafli (Turkey), Geraads et al. (2005) mark the coexistence of *Choerolophodon*, *Tetralophodon* and *Deinotherium*. The authors suggest a wet and probably forested landscape for the area, whereas Koufos (2006b) considers more open conditions during the whole Vallesian for the more southern part of Europe. It appears possible that different ecological conditions during this period prohibited the migration of tetralophodons to South and of choerolophodons to Central and Western Europe.

During the Turolian, the typical form of *C. pentelici* is present at all Greek localities (Fig. 7; Table 1) and is by far the most abundant proboscidean species, indicating that it was well adapted to this region. The dental microwear of *Choerolophodon* from Axios Valley is under study and is expected to provide more data on the dietary preferences of this proboscidean. The situation is more “balanced” in the

Pikermi proboscidean assemblage, where more material and taxa are present (the material housed at AMPG is under study). In RZO, *C. pentelici* coexisted with “*Mammut*” (Table 1). At Samos, *C. pentelici*, “*Mammut*” sp., “*T.*” *atticus* and *D. gigantissimum* were present. Noticeable is the absence of “*T.*” *atticus* and *Deinotherium* at the Turolian localities of the Axios Valley. “*T.*” *atticus* is known from the middle Turolian localities of Pikermi (Wagner 1857, Gaudry 1862) and Akkaşdağı (Tassy 2005). Furthermore, the coexistence of *Choerolophodon*, “*T.*” *atticus* and *Deinotherium* is known from several Turolian localities, including the close geographically localities of Bulgaria (e.g. Hadjidimovo) (Markov 2008 and refs. cited), indicating that these species occupied the wider region of Southeastern Europe. Therefore, most probably, the absence of “*T.*” *atticus* and *Deinotherium* from the Turolian localities of the Axios Valley can be attributed to the limited sampling, since the proboscidean specimens of the middle Turolian localities Prochoma and Vathylakkos are rather few, particularly when compared with the rich samples of Pikermi, Samos and Hadjidimovo. In any case, although proboscideans seem to be a significant palaeoecological indicator, their remains are usually scarce, and the absence of a fossil taxon does not necessarily mean that it was not included in the fauna.

All these Turolian species did not survive until the Pliocene. Schmidt-Kittler et al. (1995) list *C. pentelici* from Maramena (Serres Basin, Greece), which is dated, according to these authors, at the Turolian/Ruscianian

boundary (MN 13/MN 14). This is probably one of the last occurrences of *Choerolophodon pentelici*.

Results and conclusion

The preliminary study of the proboscideans from the Late Miocene of Macedonia and Samos Island provided the following results:

- The study of choerolophodons from the Axios Valley localities (which cover the whole Late Miocene) permitted the recognition of different evolutionary stages for the genus *Choerolophodon*. Moreover, the distinction at specific level was confirmed with the most primitive form belonging to *C. anatolicus* and the advanced one to *C. pentelici*. This allowed the taxonomic revision of the Late Miocene peri-Mediterranean choerolophodons.
- At all studied localities, the dominant proboscidean is *Choerolophodon*. *C. anatolicus* was recognized at Pentalophos and characterizes the early Vallesian (MN 9). The most primitive form of *C. pentelici* is recognized at Xirochori, Ravin de la Pluie and Ravin des Zouaves 1 (late Vallesian, MN 10). The typical and advanced form of *C. pentelici* is traced at the localities Ravin des Zouaves 5, Vathylakkos 2, Prochoma, Dytiko 2, 3, Samos (old collections), Mytilinii-1B and Nikiti 2 (Turolian). Consequently, *C. pentelici* characterizes the late Vallesian-Turolian (MN 10-MN 13).
- The mammutid material from Ravin des Zouaves 5 (early Turolian, MN 11) and Mytilinii-1A (middle Turolian, MN 12) is attributed to “*Mammut*” sp.
- “*T.*” *atticus* is rare and recognized only in Samos (old collections), thus only a general Turolian age can be assumed for the species. The type locality of the species, Pikermi, is dated to middle Turolian.
- The deinotheres are represented by two species. *Deinotherium giganteum* was traced at the late Vallesian (MN 10) locality Ravin de la Pluie. *Deinotherium gigantissimum* is present only at Samos (old collections), thus only a general Turolian age can be assumed for the species.
- During the Vallesian, three proboscidean species were present: *C. anatolicus* during the early Vallesian, and *C. pentelici* and *D. giganteum* during the late Vallesian. The absence of *Tetralophodon* at the Vallesian

localities of Greece can probably be attributed to ecological reasons. The open environments of South-eastern Europe prohibited the migration of *Tetralophodon* in the area, whereas the more closed ecological conditions of Western and Central Europe were favourable for this elephantoid.

- During the Turolian, four proboscidean species were recognized: *C. pentelici*, “*Mammut*” sp., “*T.*” *atticus* and *D. gigantissimum*. The absence of the two latter species from the Turolian localities of Axios Valley is possibly attributed to limited sampling, as they are present in the neighbouring areas (e.g. Pikermi, Samos, Hadjidimovo).

New material from the broader region is needed to clarify the taxonomy, evolution and relationships, especially of the rarer proboscidean taxa. The study of the Pikermi material (AMPG) will probably shed more light on these issues. Moreover, several palaeoecological methods should be applied (e.g. dental microwear, isotope analysis) to understand the ecological niche of these proboscideans.

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Appendix

See Table 2.

Table 2 Dental measurements (in mm) of *Choerolophodon* from various Late Miocene localities of Greece

	PNT		XIR		RPI		RZI		RZO		VTK		PXM	
	L	B max	L	B max	L	B max	L	B max	L	B max	L	B max	L	B max
DP2														
NS	2		1		3		1		4		1			
V	26.6–27.3	19.5–22.9	28.7	23.2	29–30.8	20.6–23	27	21.6	29–34.5	24.1–27	33	23		
M	27	21.2			29.8	21.8			32	25.3				
SD	0.5	2.4			0.9	1.2			2.3	1.3				
dp3														
NS	1		2		4		1 (worm)		2				1	
V	46.7	28.3	45.9–52.5	29.7–30.3	46.3–52.7	28–32.3	39.4	25.9	51–53.1	30.3–33			51.7	33.5
M			49.2	30	49	30			52.1	31.7				
SD			4.7	0.4	3	1.8			1.5	1.9				
DP3														
NS	6		6				1 (worm)		5		1			
V	41.2–46.7	33–38.2	46.6–50.4	38–42.8			40.3	35	46.5–53.1	40.6–46.6	57.2	46.7		
M	44	35.6	48.2	41.6					49.4	44.2				
SD	2	2.2	1.5	1.8					2.4	2.3				
dp4														
NS	–		3		3		2		1				3	
V			(61)–(68)	(37)–45.1	68.2–82	(39.2)–45.4	69.3–71.6	39.6–41.6	(66)	(46)			73.8–76.7	42.5–44
M			65.3	39.9	76.4	41.9	70.5	40.6					75.5	43.2
SD			3.8	4.5	7.3	3.2	1.6	1.4					1.5	0.8
DP4														
NS	3		5		1		1		2		1			
V	(65.4)–70	(47.3)–50.5	67.6–78.5	(37)–53.7	73.6	(51)	61.5	42.4	72.5–78.7	(53.7)–55.2	(76)	51.5		
M	67.4	48.4	71.4	47.1					75.6	54.5				
SD	2.4	1.8	4.2	6.8					4.4	1.1				
m1														
NS	–		1		1				1				1	
V			92.5	55.8	103.6	57.8			91.8	53.4			95.1	55
M														
SD														
M1														
NS	–		1		1		1							
V			81.7	56.3	(78.2)	(53.2)	80	(45)						
M														
SD														

Table 2 continued

	PNT		XIR		RPI		RZ1		RZO		VTK		PXM	
	L	B max	L	B max	L	B max	L	B max	L	B max	L	B max	L	B max
m2														
NS	2		-		-		-		1		-		-	
V	(104)-109	61.8-62							(130)	(62)				
M	106.5	61.9												
SD	3.5	0.1												
M2														
NS	-		-		1		-		-		-		-	
V					107	69								
M														
SD														
m3														
NS	3		1		-		-		-		-		-	
V	(148)-174	63.8-81.8	187	74.6										
M	163.0	74.9												
SD	13.5	9.7												
	DIT		DKO		R.X		SAM		MTLB		NIK			
	L	B max	L	B max	L	B max	L	B max	L	B max	L	B max		
DP2														
NS	1		-		-		7		1		7			
V	32	24.3					30.6-35.3	22-27	34.6	23.2	33-36	22.2-27.8		
M							33.2	24.6			34.3	24		
SD							1.6	1.9			1.1	2		
dp3														
NS	2		1		-		5		-		10			
V	53.1-54.5	33-35.2	52.5	29.8			54.6-62	30.2-38.5			51-59.8	31-35.7		
M	53.8	34.1					57.6	34.2			55.8	33.2		
SD	1	1.6					3	3			3.1	1.7		
DP3														
NS	1		-		-		6		-		7			
V	52.6	42.5					47.6-54	39-45.5			49.6-56.5	41.5-45		
M							51.5	42.8			52.4	43.4		
SD							2.7	2.8			2.4	1.6		

Table 2 continued

	DIT		DKO		R.X		SAM		MTLB		NIK	
	L	B max	L	B max	L	B max	L	B max	L	B max	L	B max
dp4	-		-		2		3		-		2	
NS					77.6-78.3	(41)-(42)	(67.4)-82	(40.5)-47			75.8-76.5	44.3-44.7
V					78	41.5	73.5	43.4			76.2	44.5
M					0.5	0.7	7.6	3.3			0.5	0.3
SD												
DP4												
NS	1		-		1		6		-		-	
V	76.6	(39)			72.5	48.6	(65.8)-73	(47)-53				
M							69.4	50.2				
SD							2.7	2.3				
m1												
NS	-		-		-		-		-		-	
V												
M												
SD												
M1												
NS	-		-		-		2		1		-	
V							89.4-94.1	60.1-60.4	(86.8)	(68)		
M							91.75	60.25				
SD							3.3	0.2				
m2												
NS	-		-		-		-		-		-	
V												
M												
SD												
M2												
NS	-		-		-		-		1		-	
V									(126.3)	(62.15)		
M												
SD												

L length, B max maximum breadth, NS number of specimens, V variation (minimum-maximum value), M mean value, SD standard deviation; measurements in parentheses indicate that the complete measurement was not available. For locality abbreviations see Table 1 footnote

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