SHORT NOTE



# June 19 2015 Rainfall Event Over Mumbai: Some Observational Analysis

Sudipta Sarkar<sup>1,2</sup>  $\cdot$  Ramesh P. Singh<sup>3</sup>

Received: 8 August 2015 /Accepted: 16 February 2016 /Published online: 4 April 2016  $\odot$  Indian Society of Remote Sensing 2016

Abstract Some of the major metropolitan centers in the world are highly susceptible to flash floods and major disruptions, owing to sudden and excessive rainfall events. The city of Mumbai, India's financial capital, suffered one such event on 19 June, 2015. This was a second event of such nature, following the landmark event of 26 July, 2005. Such extreme rainfall events are often brought about by certain rapidly developing, local disturbances, which if actively monitored, may be provide important information that can be of great use for early warning to civic authorities and emergency planners. In this paper, we have analyzed a number of different meteorological and remotely sensed parameters, a few days before the actual event, to track the development and eventual culmination of a "perfect storm" that affected Mumbai and left the city tattered. We show how regional upper layer disturbance patterns are developed, induced by warming of sea-surface temperature (SST) and sustained by instability in the atmospheric boundary layers to quickly develop into massive cyclonic storms.

Keywords India . Rainfall . Mumbai . Local-disturbance . Sea-surface temperature . Geopotential height . Wind shear

 $\boxtimes$  Sudipta Sarkar sdptsarkar2@gmail.com

- Science Systems and Applications Inc., 10210 Greenbelt Rd, Suite 600, Lanham, MD 20706, USA
- School of Life and Environmental Sciences, Schmid College of Science and Technology, Chapman University, One University Drive, Orange, CA 92866, USA

#### Introduction

On 19 June 2015, the city of Mumbai experienced an extreme rainfall event within 24 hours. The rainfall measured at Mumbai, in Santa Cruz and Colaba was respectively, around 283 mm and 209 mm (source: Indian Meteorological Department Automated Weather Stations – AWS). This rainfall within 24 hours was one of the highest recorded in Mumbai after the extreme rainfall event of 26 July 2005. At Chhatrapati Shivaji International airport (Fig. [1a](#page-1-0)) (WMO station) total rainfall of  $\sim$ 283 mm was recorded on 19 June 2015 (Fig. [1b\)](#page-1-0), the third highest total rainfall on a day since 2000. This was followed by another extreme rainfall event on 23 June when several weather stations in the city recorded rainfall in the range 220–273 mm (Skymet AWS: [http://www.skymetweather.com/content/](http://www.skymetweather.com/content/weather-news-and-analysis/monsoon-rains-recorded-by-skymet-in-mumbai/) [weather-news-and-analysis/monsoon-rains-recorded-by-skymet](http://www.skymetweather.com/content/weather-news-and-analysis/monsoon-rains-recorded-by-skymet-in-mumbai/)[in-mumbai/](http://www.skymetweather.com/content/weather-news-and-analysis/monsoon-rains-recorded-by-skymet-in-mumbai/)). Over the entire month of June, Mumbai received the second highest total rainfall in a decade, after June of 2013.

Extreme rainfall events ( $>50$  mm) in a day are common along the western coast of India, during the summer monsoon period (Gadgil et al. [2004](#page-6-0)). But such extreme rainfall events cause frequent flash floods, especially in low lying urban cities, throughout the world due to growing built up areas and lack of a good drainage system (Carter [1961;](#page-6-0) Hollis [1975;](#page-6-0) Beecham and Chowdhury [2012\)](#page-6-0). Mumbai city, with population of more than 19 million (as of 2013), has grown in all respect, built up areas and new infrastructure. During summer monsoon, water logging/flash flood is a common problem affecting daily lives of people. During the last such event on 26 July 2005, Mumbai received more than 450 mm of rainfall, severely affecting the daily life of people and causing financial loss of well beyond \$2 billion (Prasad and Singh [2005;](#page-6-0) Bohra et al. [2006\)](#page-6-0).

Even given the history of such rainfall events, the heavy rainfall in the month of June 2015 was unexpected especially with the progress of a strong El-Nino that was expected to

<sup>&</sup>lt;sup>1</sup> NASA Goddard Space Flight Center, 8800 Greenbelt Rd, Greenbelt, MD 20771, USA

<span id="page-1-0"></span>Fig. 1 a) The location of the WMO station located in Chhatrapati Shivaji airport in Mumbai. (map courtesy: Google maps). b) The daily-accumulated total rainfall (mm) during 1 January 2000 – 6 July 2015, obtained from the WMO station located in Chhatrapati Shivaji airport in Mumbai



cause poor monsoon and lower than average rainfall (Sikka [1980;](#page-7-0) Pant and Parthasarathy [1981](#page-6-0); Rasmusson and Carpenter [1983\)](#page-7-0). Although several discrepancies in the relationship between El-Nino and Indian monsoon has been discussed (Kumar et al. [1999](#page-6-0); Sarkar et al. [2004](#page-7-0); Kumar et al. [2006\)](#page-6-0), still such an extreme rainfall (June 2015) was not expected in view of the ongoing El-Nino event. The city was totally unprepared for this event and there were heavy losses in terms of workdays lost, hindrance to mass transit systems, damage to properties and even loss of life.

In this paper, we have analyzed a number of satellite and meteorological parameters to study the development of parameters relating to extreme rainfall, off the western coast of India, close to Mumbai. Understanding of such rapidly developing local fluctuations provide an insight in developing forecasts of extreme rainfall events to alert civic bodies and emergency personnel, to face such an event. Similar efforts to explain the 2005 event based on observation data (Prasad and Singh [2005](#page-6-0)) as well as modeling (Bohra et al. [2006](#page-6-0); Kumar et al. [2008](#page-6-0); Sahany et al. [2010\)](#page-7-0) provided valuable insights. Bohra et al. [\(2006\)](#page-6-0), have pointed out deficiencies in the current weather prediction model being used by the Indian agencies and have suggested several improvements to fix the model. Kumar et al. ([2008](#page-6-0)) and Sahany et al. [\(2010\)](#page-7-0), have discussed in detail the factors that may have triggered such events. Another study, using the WRF model, to understand the catastrophic rainfall over Uttarakhand in June 2013 (Kumar et al. [2015](#page-6-0)) has also highlighted the importance of timely prediction and warning for better disaster mitigation.

### **Materials**

Extreme rainfall events are associated with the land-ocean– atmosphere coupling. Understanding of such coupling is complex and requires various types of data related to land, ocean, atmosphere and meteorological parameters.

<span id="page-2-0"></span>The global merged Infrared (IR) brightness temperature data (BT) are considered for the cloud-top temperature, during and around the period of anomalous rainfall event of 19 June 2015. This data is obtained from the Climate Prediction Center (CPC) archive of US National Weather Service. We have used the full resolution (4 km) IR ( $\sim$ 11 micron) brightness temperature that covers 60 N-60S and is derived from all available geostationary satellites (GOES-8/10, METEOSAT-7/5 and GMS) (Janowiak et al. [2001](#page-6-0)), we will refer this data as CPC IR BT.

We have also considered BT data from the thermal infrared band (10.5–12.5 micron) (TIR) of Very High Resolution Radiometer (VHRR) onboard the Kalpana-1 (METSAT) satellite, launched by Indian Space Research Organization (ISRO). The half-hourly data have been analyzed at  $\sim$ 12 PM local time, daily, we have obtained this data from the MOSDAC [\(http://www.mosdac.gov.in/\)](http://www.mosdac.gov.in/) meteorological data archive of India.

SST anomalies have been taken from the NOAA  $0.25^{\circ}$  x  $0.25<sup>0</sup>$  optimum interpolated dataset. This dataset has been derived by blending in different observations (Reynolds et al. [2007\)](#page-7-0) from multiple platforms (satellites, ships, buoys).

We have also used the Global Forecast System (GFS) model data on a  $0.25^{\circ} \times 0.25^{\circ}$  lat-long grid, from the National Centers for Environment Prediction (NCEP), USA. GFS is a coupled weather forecast model that is run by NCEP globally at a base horizontal resolution of ~13 kms for the first 10 days which drops to  $\sim$ 27 km thereafter. Vertically, the model is



190 210 230 250  $\frac{1}{170}$  $\frac{1}{27}$  $\frac{1}{290}$ 



305 300 295 290 285 280 275 270 265 260 255 250 245  $23$ JUN 15JUN 17JUN 19JUN  $21$ JUN  $25$ JUN 13JUN<br>2015

Fig. 2 a) The CPC IR BT (Kelvins), collected at 12:30 local time, from 18–23 June 2015 (going from left to right, from top panel to lowest panel). **b**) Brightness temperature on 19 June (*left*) and time series of

Brightness temperature (in Kelvins, right) from 13– 25 June of 2015, derived from Kalpana-1 TIR, showing changes in BT over Mumbai. The city of Mumbai is marked with a black square

<span id="page-3-0"></span>divided into 64 layers and temporally, model forecast is given out every hour for the first 12 hours, three hourly thereafter till day 10 and every 12 hours after that. The GFS forecasts are made at four time steps every day (00z, 06z, 12z and 18z) and forecasts are made up to 16 days in the future. GFS analyses, that describe the state of atmosphere at the time of model initialization, are given out at 00z and 12z every day.

The daily rainfall data over the Santa Cruz (Chhatrapati Shivaji airport) have been taken from WMO station data archive, through their global web interface [\(http://gis.ncdc.noaa.](http://gis.ncdc.noaa.gov/map/viewer/%23app=clim&cfg=cdo&theme=daily&layers=0001&node=gis) [gov/map/viewer/#app=clim&cfg=cdo&theme=](http://gis.ncdc.noaa.gov/map/viewer/%23app=clim&cfg=cdo&theme=daily&layers=0001&node=gis) [daily&layers=0001&node=gis](http://gis.ncdc.noaa.gov/map/viewer/%23app=clim&cfg=cdo&theme=daily&layers=0001&node=gis)).

#### **Results**

Figure [2a](#page-2-0) shows the daily CPC IR BT in and around Mumbai. The city of Mumbai is shown with a black square in this and all subsequent figures. Figure [2b](#page-2-0) shows the BT for 19 June, from the Kalpana-1 TIR channel. The sudden drop in the IR brightness temperature/cloud top temperature (around 250 K from an average of  $\sim$ 290 K) is clearly seen on 19 June (Fig. [2b](#page-2-0)), from the Kalpana-1 TIR BT time series. This sudden drop in BT on 19 June 2015 reflects rapid vertical uplift and advective cooling associated with build up of an intensive storm and lifting.

This strong advective cooling is also seen from Fig. 3 that shows the lifted index on 18 June. Lifted index is a measure of the difference between 500 mb ambient temperature and that of a rising parcel of air. A negative lifted index is thus an indication of developing instability in the atmosphere. A strong negative lifted index (−3 K and lower) is seen to hover over Mumbai and surrounding areas from Fig. 3. The development of a storm system and intensifying cyclone pattern is further seen from Fig. [4.](#page-4-0) The 500 mb geopotential height

Fig. 4 Plots of different meteorological parameters as derived from the  $\blacktriangleright$ GFS 12z analysis. Each set of figures represents days during 18 – 23 of June 2015, from left to right and from top to bottom. a) 500 mb geopotential height (shaded, in gpm) with wind stream contours at 925 mb. b) The vertical shear, calculated between 200 and 850 mb. c) Absolute vorticity at 500 mb  $(x10^5)$  hertz. The city of Mumbai is marked with a black square in all the figures

(Fig. [4a\)](#page-4-0) variation from 18–23 June 2015, shows the development of a strong depression, just off the west coast of India, near the southern tip of Gujarat. The positioning of this depression is seen to change from 19 June 2015 (Fig. [4a](#page-4-0) upper panel center) to 23 June 2015 (Fig. [4a](#page-4-0) lower panel right). The 500 mb depression is observed very close to the city of Mumbai, on 19 June causing heavy rainfall.

A strong positive vorticity is also seen to develop at 500 mb at the same time, indicating strong counter-clock wise wind pattern and storms at the upper level (Fig. [4c](#page-4-0)). Again the positioning of the positive vorticity pattern and its adjacency to the city of Mumbai is noteworthy. What further aids in the build up of this strong cyclonic storm system is the low tropospheric wind shear of 15 knots or less on 19 and 20 June 2015 (Fig. [4b](#page-4-0)). The vertical wind shear increases between 21 and 22 June 2015 and is likely responsible for the break up of the single cell storm pattern to two cells as clearly seen from the geopotential height on 23 June 2015 (Fig. [4a\)](#page-4-0).

Furthermore, the position of the depression at 500 mb and location of the maximum positive vorticity shows Mumbai and its adjoining areas on the eastern or right hand side of this developing and strong counter clock wise wind pattern, indicating strong upper level uplift over this area, corroborating similar observation seen from the analysis of CPC and Kalpana-1 BT (Fig. [2](#page-2-0)).

The SST anomaly plot from 16 to 19 June 2015 (Fig. [5a](#page-5-0)) shows a strong positive anomaly all along the western coast



Fig. 3 The best of 4-lifted index from GFS on 18 June 12z analysis (Kelvin)

<span id="page-4-0"></span>

-20 -10 0 10 15 20 25 30 35 40 45 50 60 70

<span id="page-5-0"></span>Fig. 5 a) SST anomaly pattern (in deg. C) during 16–19 June (left to right and from top to bottom panel). b) Time series of weekly SST anomaly since first week of 2000 until the week of 5 July 2015



and especially off the southern tip off Gujarat. This pattern is seen to develop from 16 June 2015 onwards; a peak is seen on 19 June 2015 and slowly dissipates thereafter (not shown in Figure). The weekly SST anomaly from the first week of 2000, over a point close to the Mumbai coast (19 N, 72E) (Fig. 5b), shows the SST to be the second highest in the last one decade.

Such anomalous SST of almost 1.5 degree or more led to the development of a convective cell off the southern coast of Gujarat, in the Arabian Sea. A strong upward draft developed on 18 June, as seen from the strong low-level wind patterns (Fig. [4a](#page-4-0)). This convective cell intensified on 19 June and moved close to Mumbai and intense rainfall occurred over Mumbai on 19 June. Strong temperature contrast between 500mb ambience and the convective parcel (Fig. [3\)](#page-3-0), weak

tropospheric shear and strong surface winds enhanced the heat transfer from the ocean.

Strong negative vertical velocity (Omega) values with high levels of atmospheric moisture content in the air and strong upward moisture advection is also seen from Fig. [6a](#page-6-0) and [b.](#page-6-0) This pattern is found during 18 to 23 June, showing the persistence of the vertical cell through continued moisture advection and moist adiabatic lapse rate. After 19 June this convective cell shifted westward towards deeper Arabian Sea and became even stronger. At this point the cell split into two, aided partly by the increasing tropospheric wind shear on 21 and 22 June. Since the main center of the storm shifted further off the coast of Mumbai, even though rainfall occurred until 23 June but the heaviest rain occurred on 19 June when all the meteorological parameters were ideal for the perfect storm.

<span id="page-6-0"></span>Fig. 6 a) Vertical velocity or Omega (mb/hr) at 700 mb (shaded) with contours of relative humidity (%) at 850 mb, derived from GFS 12z analysis on 18 of June and 23 June. b) Vertically integrated moisture convergence calculated between surface and 300 mb pressure level (kg/s) for the same time and days as in (a). Convergence is shown as positive



# **Conclusions**

Though the general trends in Indian summer monsoon rainfall have been actively studied and a number of favorable parameters responsible for storm have been identified, but still the physical process and the key role of some of the parameters are not fully well understood. Moreover local fluctuations due to orography, local disturbances and short-term atmospheric oscillations can cause some local to regional anomalous rainfall pattern that can further hamper our ability to model and predict rainfall anomalies. Hence monitoring of real time weather information and meteorological variables is of crucial importance, especially for large cities like Mumbai where sudden rainfall and major water logging can completely throw the entire city off gear. Our study shows the necessity of an integrated approach to study of local and regional weather pattern that is still lacking in many parts of India and in many of the major metropolitan centers.

Acknowledgments The authors are grateful to the two anonymous reviewers for their comments and suggestions which have helped us to improve the earlier version of the paper. We are grateful to MOSDAC ([http://www.mosdac.gov.in/\)](http://www.mosdac.gov.in/) for providing meteorological data from Indian satellites to use in the present study.

## References

Beecham, S., & Chowdhury, R. (2012). Effects of changing rainfall patterns on WSUD in Australia. Journal of Water Management, Institution of Civil Engineers UK, 165(5), 285–298.

- Bohra, A. K., Basu, S., Rajagopal, E. N., Iyenger, G. R., Das Gupta, M., Ashrit, R., & Athiyaman, B. (2006). Heavy rainfall episode over Mumbai on 26 July 2005: assessment of NWP guidance. Current Science, 90(9), 1188–1194.
- Carter, R. W. (1961). Magnitude and frequency of floods in suburban areas. U.S. Geological Survey Professional Paper, 424-B, 9–11.
- Gadgil, S., Vinayachandran, E. N., Francis, P. A., & Gadgil, S. (2004). Extremes of Indian summer monsoon rainfall. Geophysical Research Letters, 31(12), L12213. doi[:10.1029/2004GL019733.](http://dx.doi.org/10.1029/2004GL019733)
- Hollis, G. E. (1975). The effects of urbanization on floods of different recurrence intervals. Water Resources Research, 11(3), 431–435.
- Janowiak, J. E., Joyce, R. J., & Yarosh, Y. (2001). A real–time global half–hourly pixel–resolution infrared dataset and its applications. Bulletin of American Meteorological Society, 82(2), 205–217.
- Kumar, K. K., Rajagopalan, B., & Cane, M. (1999). On the weakening relationship between the Indian Monsoon and ENSO. Science, 284(5423), 2156–59. doi:[10.1126/science.284.5423.2156.](http://dx.doi.org/10.1126/science.284.5423.2156)
- Kumar, K. K., Rajagopalan, B., Hoerling, M., Bates, G., & Cane, M. (2006). Unraveling the mystery of Indian monsoon failure during El Nino. Science, 314(5796), 115–19. doi[:10.1126/science.](http://dx.doi.org/10.1126/science.1131152) [1131152](http://dx.doi.org/10.1126/science.1131152).
- Kumar, A. J., Dudhia, R., Dev Niyogi, R., & Mohanty, U. C. (2008). Analysis of the 26 July 2005 heavy rain event over Mumbai, India using the Weather Research and Forecasting (WRF) model. Quarterly Journal of the Royal Meteorological Society, 134(636), 1897–1910. doi[:10.1002/qj.325](http://dx.doi.org/10.1002/qj.325).
- Kumar, P., Shukla, B. P., Sharma, S., Kishtawal, C. M., & Pal, P. K. (2015). A high-resolution simulation of catastrophic rainfall over Uttarakhand, India. Natural Hazards, 80(2), 1119–1134.
- Pant, G. B., & Parthasarathy, S. B. (1981). Some aspects of an Association between the Southern Oscillation and Indian Summer Monsoon. Archives For Meteorology, Geophysics, and Bioclimatology Series B, 29(3), 245–52. doi[:10.1007/bf02263246](http://dx.doi.org/10.1007/bf02263246).
- Prasad, A. K., & Singh, R. P. (2005). Extreme rainfall event of July 25– 27, 2005 over Mumbai, West Coast, India. Journal of the Indian Society of Remote Sensing, 33(3), 365–70. doi:[10.1007/](http://dx.doi.org/10.1007/bf02990007) [bf02990007.](http://dx.doi.org/10.1007/bf02990007)
- <span id="page-7-0"></span>Rasmusson, E. M., & Carpenter, T. H. (1983). The relationship between eastern equatorial pacific sea surface temperatures and rainfall over India and Sri Lanka. Monthly Weather Review, 111(3), 517–528. doi: [10.1175/1520-0493\(1983\)111<0517:trbeep>2.0.co;2.](http://dx.doi.org/10.1175/1520-0493(1983)111%3C0517:trbeep%3E2.0.co;2)
- Reynolds, R. W., Smith, T. M., Liu, C., Chelton, D. B., Casey, K. S., & Schlax, M. G. (2007). Daily high-resolution-blended analyses for sea surface temperature. Journal Of Climate, 20(22), 5473–5496. doi:[10.1175/2007jcli1824.1.](http://dx.doi.org/10.1175/2007jcli1824.1)
- Sahany, S., Venugopal, V., & Nanjundiah, R. S. (2010). The 26 July 2005 heavy rainfall event over Mumbai: numerical modeling aspects.

Meteorology and Atmospheric Physics, 109(3–4), 115–128. doi: [10.1007/s00703-010-0099-3.](http://dx.doi.org/10.1007/s00703-010-0099-3)

- Sarkar, S., Singh, R. P., & Kafatos, M. (2004). Further evidences for the weakening relationship of Indian rainfall and ENSO over India. Geophysical Research Letters 31 (13), doi: [10.1029/2004gl020259](http://dx.doi.org/10.1029/2004gl020259).
- Sikka, D. R. (1980). Some aspects of the large scale fluctuations of summer Monsoon Rainfall over India in Relation to fluctuations in the planetary and regional scale circulation parameters. Journal of Earth System Science, 89(2), 179–195. doi[:10.1007/bf02913749](http://dx.doi.org/10.1007/bf02913749).