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Groundwater pollution risk assessment using a calculated contamination index and geostatistical analysis: Jerba Island case study (southeast of Tunisia)

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Abstract

In this study, the groundwater contamination degree in the unconfined coastal aquifer Jerba under arid climate (Southeastern Tunisia) was investigated, and a new index, groundwater contamination index (GCI), was calculated and combined with the geostatistical analysis and the self-organizing maps (SOM). The groundwater samples used in this research were collected from 79 groundwater wells during the dry season (August–October 2014). Hydrochemical parameters (Cl⁻, Br⁻, Na⁺, Ca₂⁺, Mg²⁺, K⁺, SO₄²⁻, NO₃⁻, NO₂⁻, HCO₃⁻, Li⁺, and F⁻) and indicator bacteria of fecal contamination (total coliforms, thermotolerant coliforms, and *Escherichia coli*) were analyzed. The GCI was calculated based on selected indicator parameters notably NO₃⁻ and Li⁺, measured seawater fraction based on chloride balance, and fecal bacteria tracers. Geostatistical modeling was used for assessing and mapping groundwater contamination degrees. Ordinary Kriging was adopted for spatial interpolation to study the spatial pattern of the groundwater hydrochemical variables over the island using GIS software ArcGIS 10.1. The SOM method was adopted to analyze the relationship between ions and identify processes controlling groundwater salinization. According to the GCI, most of the unconfined aquifer (76%) comes under the significant pollution zone (high to moderate pollution), and the other areas (24%) are defined as areas with low degrees of pollution. The self-organizing maps (SOM) indicated that Cl⁻, Br⁻, and Na⁺ emanate mainly from seawater intrusion and Mg²⁺, Ca²⁺, and Li⁺ are mostly derived from rock-water interactions. Results show that the new index is robust and gives the best classification of groundwater quality.

Keywords Groundwater contamination degree · Kriging · Geostatistic · Coastal aquifer · Jerba Island

Introduction

In regions under arid and semiarid climate conditions worldwide, groundwater is a major source of water supply (Patel et al. 2020; Carol et al. 2021; Silva et al. 2021). The identification of key processes governing groundwater mineralization is based on geochemical tracers such as nitrogenous elements (Souid et al. 2017; Solgi and Jalili 2021; Zhang et al. 2021); conservative elements

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Obtaining the highest possible compatibility between "ideal" and "possible" scenarios is one of the most difficult aspects of any environmental management system (Barilari et al. 2020). Groundwater quality and risk assessments include a variety of quantitative methods, such as mathematical groundwater flow modeling and the transportation and the evaluation of pollution plume. These strategies will be of limited utility to groundwater managers who need to achieve proactive and timely groundwater protection across large aquifers (Somaratne et al. 2013). There are other qualitative procedures for assigning a label based on land use (Zaporozec 2004) as well as intermediate methodologies known as index methods, which include aquifer vulnerability assessment tools (e.g., DRASTIC, GALDIT, and GOD coupled to geophysical information system), and water quality index (Nsabimana et al. 2021). Methodological approaches that integrate numerous groundwater quality variables in a specific index to assess pollution degrees have found widespread uses in environmental science and engineering (Wagh et al. 2018; Herojeet et al. 2020).

In the Mediterranean regions, salinization is the major concern for coastal aquifers because of climate change and overexploitation (Nisi et al. 2022). In addition to their relative scarcity, Tunisia's groundwater resources, especially those of the shallow aquifers, are often of poor quality. Thus, about 12% of groundwater resources are characterized by salinity lower than 3 g/l. The overexploitation results in the degradation of the groundwater quality with an increase in the salinization and contamination risks. Overexploitation also leads to the lowering of the piezometric levels. In addition, the various forms of pollution that affect the unconfined aquifers make the groundwater sometimes unfit for domestic use (Elloumi 2016). Moreover, Tunisian groundwater resources have been exploited excessively in recent years, causing the depletion of reserves of many aquifers. The volume extracted from aquifers is estimated at 480 Mm³/year, or 24% of total withdrawals of groundwater, which is hardly sustainable. According to the 2004 national inventory, there were 750 sources of contamination with urban, industrial, and agricultural origins that are likely to seriously affect groundwater supplies. The inventory notes more than 40 urban water treatment plants that do not comply with the Tunisian standard of discharge into the aquatic environment. The food industry is the primary source of pollution, followed by the sectors of textile and clothing, rubber, and plastic (Besbes et al. 2014).

Jerba unconfined aquifer is a coastal aquifer under an arid climate. It is influenced by several factors: climate conditions, anthropic activities, and seawater intrusion. This aquifer was decreed as a safeguard area of groundwater resources (Decree N^o. 85–1108 of 29 August 1985-Tunisia). The previous hydrogeological studies carried out on the shallow aquifer of Jerba Island focus on the geochemical processes of groundwater salinization based on geochemical and statistical modeling and the assessment of groundwater piezometry (Kharroubi et al. 2012; Telahigue et al. 2018; Souid et al. 2017, Souid et al. 2020). Such studies demonstrate the deterioration of the groundwater geochemical quality. The lack of a sustainable groundwater exploitation management strategy has worsened the situation on the island. The novel contribution of this work is to propose an empirical index based on a contamination tracer to facilitate the understanding of the aquifer contamination processes.

The aim of the present research is to study the groundwater contamination degree of an arid coastal aquifer known by the scarcity of water resources and an aquifer system subjected to many natural and human pressures. The specific objectives were to identify the causes of groundwater pollution for a better monitoring and recognize the threats to groundwater supply, based on a contamination index (combining chemical and bacteriological tracers of contamination) and geostatistical approaches. Jerba Island's unconfined aquifer in the southeast of Tunisia is used as a study case.

Methodology and data processing

Location of the study area

The area of research is in Tunisia's southeast between latitude 33°38' and 33°56'N and longitude 10°43' and 11°04'E. It occupies an area of about 514 Km² (Fig. 1). The annual average rainfall is 250 mm, and the annual average evaporation is 1082 mm (Souid et al.2020). Surface waters mainly consist of rainwater stocked in special tanks and used for drinking and domestic purposes. The island's topography is flat, with an average altitude of 50 m. The island is covered by quaternary sediments composed, as shown in the hydrogeological cross-section (Fig. 1), by a thick-bedded Mio-Plio-Quaternary sand and clay overlying bedrock downwards and upwards by the gypseous crusts attributed to Villafranchian. The shoreline of the island is draped by the marine facies of Holocene and Pleistocene. The Mio-Pliocene marl and clay act like an aquitard, while the Mio-Pliocene sand represents the basal aquifer (Teissier 1967; Jedoui 2000; Bouaziz et al. 2003; Yahyaoui 2012).

The underground of the island encloses a phreatic aquifer and a deep confined aquifer (Fig. 1), within the quaternary sediments separated vertically by a thick clay layer that represents the aquitard stratum. The unconfined aquifer, which is mainly composed of lenticular sandy levels of Quaternary upper Pleistocene and Quaternary Holocene, is the most used aquifer level by the locals for both domestic and agricultural purposes (Kharroubi et al. 2012; Yahyaoui 2012; Souid et al. 2017). The unconfined aquifer in the island is mostly recharged by the vertical rainwater infiltration and surface water (irrigation water and wastewater) infiltration. The aquifer depth ranges between 3 and 50 m with a mean depth of 30 m. The unsaturated zone is mainly made up of Mio-Plio-Quaternary sandy loam, silt loam, and sand layers. Overexploitation has led to the continuous decline in the aquifer's piezometric level (Souid et al. 2017, 2020).

Physicochemical and bacteriological analyses

A total of 79 groundwater points were sampled, after purging for 5 min to verify that the groundwater samples were reflective of the in situ conditions, during the summer season



Fig. 1 a Geological map of the of the Jerba Island; b hydrogeological cross-section

(August–October 2014). Samples for the enumeration of fecal contamination tracers were stored in a sterile glass container (200 ml). For physicochemical analyses, groundwater samples were kept in polyethylene bottles (500 ml). During the

groundwater sampling, EC, salinity, and pH were measured on the field using a Consort 933 multi-parameter. Cations (Na⁺, K⁺, Ca²⁺, Mg²⁺, Li⁺) and anions (Cl⁻, Br⁻, NO₃⁻, NO₂⁻, SO₄²⁻, HCO₃⁻, F⁻) were analyzed using chromatography (ionic liquid chromatograph Metrohm 850). Prior to analysis, the filtration of water samples through cellulose membrane (pore size: 0.45 um) was carried out. Samples were kept cool and transported to the laboratory of the Higher Institute of Water Sciences and Techniques, University of Gabes, Tunisia. To assess the quality control measures of chemical analyses, ionic balance errors (within 10%) were calculated.

The membrane filtration standard (ISO 9308–1 2000) was adopted to enumerate total (TC) and thermotolerant coliforms (Thc). The seeding of these bacteria was carried out on the bacterial culture medium Tergitol TTC Agar during 42 h of incubation. The incubation temperature is 37 °C for TC and 40 °C for Thc. The enumeration results were expressed by CFU 100 ml⁻¹ unit. Total coliform colonies were transplanted on a tryptophan reagent to identify *Escherichia coli* bacteria. The identification of *E. coli* in aqueous media was confirmed after an incubation time of 24 h at a temperature of 44 °C. *E. coli* bacteria cause indole production, which can be determined by adding Kovac's reagent. Bacteriological analyses were performed at Tunisia's Public Health Ministry's Regional Hygiene Laboratory in Jerba.

Geostatistical analysis

A histogram and standard QQPlot tools were used to verify the normality of the distribution pattern in the geochemical dataset. Normal QQPlot is a graphical normality assessment method to indicate univariate normality (Johnston et al. 2001), investigate the distribution of data, and look for outliers. For that, The Shapiro-Wilk and Kolmogorov-Smirnov tests were performed to assess the normality of the dataset. The normality tests are used in addition to the graphical normality evaluation (Elliott and Woodward 2007). These statistical tests compare the scores in the samples to a Gaussian distribution of scores, which presents a normally distributed set, with the same mean and standard deviation. If the "P" value of the tests is more than 5%, the null hypothesis is accepted and "sample distribution is Gaussian." If the test is significant ("P" value of the tests is less than 5%), the distribution is non-Gaussian (Ghasemi and Saleh Zahediasl 2012). The Kolmogorov-Smirnov test is an empirical distribution function that compares the test distribution's theoretical cumulative distribution function to the data empirical distribution function (Oztuna et al. 2006). The Shapiro-Wilk test is more efficient than the Kolmogorov-Smirnov test since it is based on the correlation between the data and the matching normal scores (Peat and Barton 2005).

Semivariograms are used to quantify the spatial autocorrelation (dependence) between sampled points. The semivariogram is based on the simple geographic dogma: samples that are close to each other are more similar on the geochemical behavior than the more distant samples. This principle is called autocorrelation (statistical and spatial relationships between samples). Experimental semivariogram and experimental covariance were calculated based on Eqs. (1) and 2, respectively:

$$\gamma(h) = \frac{1}{2n(h)} \sum_{i=1}^{n(h)} (Z(x_i) - (Z(x_{i+h}))^2$$
(1)

where n(h) is the pair number of the studied data within a specified range of equidistance and direction. When the variables $Z(x_i)$ and $Z(x_{i+h})$ are auto-correlated, the estimated value $Z^*(x_0)$ (the result of Eq. 4) is small, and thus, the pair of points is spatially uncorrelated. Based on the analysis of the empirical variogram (experimental variogram), a suitable mathematical variogram model (e.g., Gaussian, spherical, exponential) is then adjusted, usually by weighted least squares, and variogram parameters (range, nugget, and sill) are then used in the Kriging approach (Zare-Mehrjardi et al. 2010).

If the random variable pair x_i and x_{i+h} can take on the values $Z(x_i)$ and $Z(x_{i+h})$ for i = 1, ..., n with equal probabilities, $p_i = \frac{1}{n}$, then the covariance can be expressed in terms of means $Z(x_i)$ and $Z(x_{i+h})$ as:

$$C(h) = \frac{1}{n(h)} \sum_{i=1}^{n(h)} \left[(Z(x) - Z(xi))(Z(x+h) - Z(xi+h)) \right] \quad (2)$$

For a model that provides accurate unsampled point estimations, the standardized mean (SM) error should be close to 0, the root-mean-square error (RMSE) and average standard error (ASE) should be as small as possible (comparing obtained values for each tested model), and the RMSS should be close to 1 (Johnston et al. 2001). The smallest RMSE indicates the most accurate predictions. The RMSE was derived according to Eq. (3) (Zare-Mehrjardi et al. 2010):

$$RMSE = \sqrt{\frac{1}{N} \sum_{i=1}^{N} (Z(x_i) - Z^*(x_i))^2}$$
(3)

 $Z(x_i)$ is the measured observation at the position x_i , $Z^*(x_i)$ is the estimated observation at the position x_i , and N is the sample number.

Ordinary Kriging was adopted to assess the spatial pattern of the studied groundwater hydrochemical variables over the island. The GIS software ArcGIS 10.1 was used for spatial interpolation in this work. Kriging may be considered the most efficient method for prediction (interpolation) at unsampled areas. This geostatistical approach permits the study of spatial autocorrelation between the investigated variables. One of the most significant advantages of the Kriging method is that it provides for the estimation of the error interpolation values of the regionalized variable when there are unsampled sites, providing a measure of the estimation accuracy and reliability of the variable's spatial distribution (Burgess and Webster 1980). The general equation of the Kriging estimator is:

$$Z^*(x_0) = \sum_{i=1}^n \lambda_i Z(x_i) \tag{4}$$

For each groundwater sample, the hydrochemical data were inserted into a digital database and GIS, using ArcGIS v10.1.

For spatial distribution maps, a histogram and a normal QQPlot analysis were performed. It was observed that the data are approximately normally distributed. Table 1 shows the results of the normality tests. It is clear that all variables, for both tests, have a *P* value less than 0.05, which indicates a non-Gaussian distribution of data, except for the GCI parameter. Log transformation was applied for the parameters that have a non-normal distribution to obtain a normal distribution.

Prediction model performances were assessed by cross-validation. Table 2 presents the selected model characteristics for the studied groundwater quality parameters. The mathematical models of semivariograms (Gaussian, exponential, circular, or spherical models) were tested for salinity, GCI, and seawater fraction. The models provide information about the spatial distribution including the Kriging interpolation input parameters (Boudibi et al. 2021a; Hatvani et al. 2021). The adequate model for interpolation is selected referring to the calculated error values (Table 2). The SM error is close to 0 for all studied parameters, the RMSE and ASE are low comparing to obtained values for each tested model, and RMSS is closer to the RMSE. The cross-validation results clearly indicated that circular and spherical semivariogram models are the fitted models to the spatial interpolation of the dataset.

Calculated groundwater contamination index (GCI)

A contamination index is a useful tool that can be applied to assess groundwater pollution. Various indices have been used by researchers to assess groundwater contamination (Adimalla et al. 2020; Herojeet et al. 2020; Serra et al. 2021). The groundwater contamination index provides a single value which summarizes the big dataset of quantity parameters and represents data in a simple way. GCI is an effective approach for assessing and mapping the degree of groundwater pollution (Kumar 2014). In this work, a new index is formulated and proposed to evaluate the intensity of groundwater pollution in the Jerba shallow aquifer. The different steps involved in the research process are shown by the flow chart (Fig. 2).

The calculated contamination index was derived from the combination of chemical and bacteriological parameters. The combination of chemical and bacteriological tracers helps to highlight pollution from natural and anthropogenic sources (Fig. 3). The selection of the parameters integrated in the calculation of the index was based on the indicators most involved in the geochemical signature of the groundwater. This means that a specific tracer has been chosen for each contamination process to assess the contamination degree of the aquifer. In this study, lithium and nitrate concentrations, seawater fraction, and fecal indicator bacteria (TC and ThC coliforms and Escherichia coli) were used as tracer parameters to measure the pollution degree. Weights are assigned to each parameter chosen for the proposed index depending on a number of factors, including health risk and importance in assessing groundwater quality (Herojeet et al. 2020). Pollution degrees ranged into six classes according to the concentration of each parameter (Table 3). For each parameter, the values were grouped by importance to three clusters:

- Low pollution degree cluster: this class includes the low values. For this class, the pollution degree is considered low.
- Medium pollution degree cluster: this group is composed of medium values. The pollution level of this cluster is medium.
- High pollution degree cluster: high values of each parameter are included in this cluster.

Tests of normality								
	Kolmogo Significar	mirnov wit	th Lilliefors	Shapiro–	Wilk			
	Statistic	df	P value	Interpretation	Statistic	df	P value	Interpretation
Salinity	0.118	79	0.008	NGD	0.958	79	0.011	NGD
Seawater fraction	0.118	79	0.009	NGD	0.929	79	0.000	NGD
GCI	0.115	79	0.061	GD	0.967	79	0.126	GD

GD, Gaussian distribution (P > 0.05); NGD, non-Gaussian distribution (P < 0.05)

Table 2Experimentalsemivariograms and fittedmodels of GCI, salinity, andseawater fraction

Table 1 Normality test results

Parameter	Semivariogram	RMSE	ASE	SM	RMSS	Range	Nugget	Partial sill
GCI	Circular	3.07	2.93	0.01	1.04	8.6×10^{3}	0.001	1.23
Salinity	Spherical	1.48	1.32	0.01	1.1	6.63×10^{3}	0.5	4.36
Seawater fraction	Circular	4.68	3.98	0.01	1.16	9.09×10^{3}	8.44	37.27



To calculate the GCI, three steps were required. In the first step, each parameter (lithium, nitrate, seawater fraction, total and thermotolerant coliforms and E. Coli) was assigned a relative weight (class) ranging from 1 to 6 based on its contents. The lowest values received the lowest weight (1), while the highest values received the highest weight (6). In the following step, the weight for each parameter was calculated to determine its relative contribution to the process of

index (GCI)

groundwater contamination. At the end of the process, the contamination index is calculated by adding the pollution degree classes assigned to each parameter. The groundwater contamination index was calculated as the weighted sum of all contamination class degrees (Eq. 5):

$$GCI = \sum_{i=1}^{n} C_i \tag{5}$$

where C_i is the single contamination class degree of each tracer parameter.

Results and discussion

Hydrogeochemistry

Basic statistics of the studied parameters (Table 4) show that the standard deviation values are low for all variables (parameters) compared to the means. The dataset distribution assessment based on skewness and kurtosis demonstrates that the fluoride distribution is right-skewed (negative skewness). The rest of the hydrogeochemical parameters are right-skewed (positive skewness). The statistical distribution form is platykurtic (kurtosis < 0) for pH, EC, salinity, HCO_3^- , Cl^- , Br^- , SO_4^{-2-} , Na^+ , and Ca^{2+} and leptokurtic (kurtosis > 0) for F^- , NO_3^- , NO_2^- , K^+ , Mg^{2+} , and Li^+ .

Salinity rates range from 8.8 in the shoreline to 0.5 g/l in the inner regions of the island. The hydrochemical fingerprint of Mio-Pliocene–Quaternary Jerba aquifer is characterized by Na-Cl groundwater facies type (Fig. 4). For a coastal aquifer in an arid climate, the high levels of Cl⁻ (the average concentration is about 1738.56 mg/l) and Na⁺ ions (with a mean concentration of 1107.96 mg/l) were due to the seawater impact on groundwater hydrogeochemical signature. Referring to the average concentrations, chloride is considered as the most abundant ionic species in the chemical composition of analyzed groundwater samples. Chemical contamination tracers like NO₃⁻, NO₂⁻, F⁻, and K⁺ are plentiful in the aquifer. The highest concentration was measured

Table 3 Classes of the parameters used to calculate GCI

for nitrate with average values of 75.97 mg/l. Li⁺ concentrations of groundwater points ranged between 1 and 14 mg/l.

The correlation between salinity and chemical elements was assessed using Pearson correlation (for normal distribution) to establish the relationship between ions and aquifer abnormal salinization. A P value of < 0.01 is considered to be significant. The correlation matrix (Table 5) measures the strength of the correlation among groundwater mineralization, Ca²⁺, Mg²⁺, K⁺, Li⁺, Cl⁻, and Br⁻. These findings demonstrate the importance of these ions in the groundwater salinization processes of Jerba unconfined aquifer. The feature maps generated by the application of unsupervised machine learning technique, i.e., SOM, confirm the obtained results (Fig. 5). The maps (grid) of the studied variables (salinity, Ca²⁺, Mg²⁺, K⁺, Li⁺, Cl⁻, and Br⁻) of the dataset display the trends of statistical correlation between ions and mineralization. The similar neuron patterns in the grids (the coloring of the map) indicate the positive correlation between studied parameters. The Kohonen map clustering method proves the significant correlations among salinity, halogens (Cl⁻ and Br⁻), alkaline earth metals (Ca²⁺ and Mg²⁺), and alkali metals (Na⁺, K⁺ and Li⁺).

The selected groundwater contamination index parameters

Lithium

The maximum content of Li^+ ion amounts to 15.69 mg/l. The abundance of this alkaline metal is assigned to groundwater-aquifer rock interactions related to cation-exchange reactions caused by seawater mixing. With reference to the hydrogeological study by Souid et al. (2018) on this coastal aquifer, lithium can be used as a salinization, resulting from the seawater invasion process, chemical proxy. As a result of the ion exchange mechanism, the content of this alkaline metal significantly increases corresponding to the seawater fraction. The use of this chemical element as a salinization ion tracer justifies its selection for the calculation of the index.

Pollution degree	Low		Medium		High		
Class	C _{Li}	C _{NO3}	$C_{ m sf}$	C _{TC}	$C_{ m thC}$	C _{Ec}	
	1	2	3	4	5	6	
Lithium concentration (mg/l)	>2	2–4	4–6	6–8	8-10	>10	
Nitrate concentration (mg/l)	>10	10-20	20-30	30–40	40-50	> 50	
Seawater fraction	>0	0–5	5-10	10–15	15-20	>20	
Total coliforms (UFC)	<100	100-1000	1000-10,000	10,000-100,000	>100,000	-	
Thermotolerant coliforms (UFC)	Absent	<10	10-100	100-1000	1000-10,000	> 10,000	
E. coli	Absent	Present	Present	Present	Present	Present	

Table 4 Statistical characteristics of the hydrochemical parameters

Parameters	Min	Max	Mean	Variance	S.D	Kurtosis	Skewness
рН	6.90	8.19	7.52	0.07	0.27	-0.13	0.16
Salinity (mg/l)	514	8882	3799.43	4,327,173.71	2080.19	-0.63	0.49
CE (mS/cm)	1.04	15.26	6.86	12.53	3.54	-0.71	0.41
TDS (mg/l)	665	9760	4387.14	5,125,075.63	2263.66	-0.71	0.41
HCO ₃ ⁻ (mg/l)	136.89	664.42	401.84	16,561.06	128.69	-0.69	0.37
F ⁻ (mg/l)	3.95	6.83	6.01	0.27	0.52	4.70	-1.67
Cl ⁻ (mg/l)	136.57	4319.83	1738.56	1,161,230.60	1077.60	-0.54	0.58
Br ⁻ (mg/l)	8.75	12.32	10.35	0.68	0.83	-0.80	0.19
NO ₃ ⁻ (mg/l)	39.32	196.89	75.97	1487.00	38.56	0.65	1.16
$NO_2^{-}(mg/l)$	14.02	80.85	23.99	129.05	11.36	7.08	2.13
$SO_4^{2-}(mg/l)$	56.31	1178.56	569.44	63,284.62	251.56	-0.02	0.24
Na ⁺ (mg/l)	223.42	2601.97	1107.96	332,198.75	576.37	-0.64	0.47
K ⁺ (mg/l)	12.05	154.65	56.23	862.54	29.37	0.41	0.76
Ca^{2+} (mg/l)	68.98	754.56	301.48	23,567.01	153.52	-0.20	0.61
Mg^{2+} (mg/l)	20.63	437.97	149.58	7378.86	85.90	0.41	0.75
Li ⁺ (mg/l)	1.12	15.70	5.68	9.03	3.00	0.37	0.74

Min, minimum; Max, maximum; S.D., standard deviation



Fig. 4 Groundwater facies based on (Ca + Mg) - Na + K)VS. $HCO_3-(Cl+SO_4)$

Seawater fraction (fsea)

In this study, the seawater fraction is calculated based on Cl⁻ content (Eq. 6). Considered as a conservative tracer, chloride is unaffected by water/rock reactions.

$$f_{sea} = \frac{C_{Cl,sample} - C_{Cl,fresh}}{C_{Cl,sea} - C_{Cl,fresh}}$$
(6)

where $C_{Cl,sample}$ is the measured chloride content of the sample, $C_{Cl,sea}$ is the chloride content of the Mediterranean sea sample, and $C_{Cl, fresh}$ represents the chloride content of the freshwater (the lowest analyzed salinity in groundwater samples).

Calculated seawater percentage varies between a maximum value of 18.73% and a minimum value of -0.83% (Fig. 6). The estimated seawater fraction spatial distribution map was used to define the spatial expansion of seawater intrusion in the island (Fig. 7). The low seawater percentages are within the central and the eastern parts of the study area.

The groundwater salinity increase is due to intensive exploitation caused by overpumping, which leads to seawater intrusion into coastal aquifers, promoting vertical exchanges between the shallow aquifer and deep confined aquifer characterized by brackish groundwater. Based on GIS interpolation approaches, we estimated the geographical extent of groundwater salinization. Figure 7 shows that the spatial distribution of groundwater salinity is mapped to the piezometric level. Areas characterized by high piezometric levels, considered as recharging areas, have low salinity values. On the other hand, areas of low piezometric levels show high salinity values. The map clearly shows that around the shoreline, in which the groundwater level is as low as 0 m below ground level (mbgl), the trend indicates that depletion has occurred at a very fast rate.

Nitrate

Nitrate is naturally present in water, coming from the natural cycle of organic nitrogen degradation. The high levels of nitrate in groundwater are due to agricultural activities (irrigation and organic fertilizer like manure), livestock, or industrial and domestic sewage (Li et al. 2019; He et al. 2022a, 2022b). In fact, 95% of the nitrogen flow is in the form of nitrates, the organic and ammonia forms remaining insignificant in agricultural inputs (Koné et al. 2009;

Table 5 Correlation matrix

	Salinity	HCO_{3}^{-}	F	Cl-	Br-	NO_3 -	NO_2^-	SO_4^{2-}	Li^+	Na^+	K^{+}	Ca^{2+}	Mg^{2+}
Salinity	1.00												
HCO ₃ -	0.29	1.00											
F-	0.11	0.14	1.00										
Cl	0.99	0.26	0.12	1.00									
Br-	0.79	0.39	0.14	0.79	1.00								
NO ₃ -	0.46	0.14	-0.12	0.47	0.36	1.00							
NO ₂ -	0.44	0.28	-0.06	0.43	0.40	0.91	1.00						
SO4 ²⁻	0.64	0.01	-0.04	0.54	0.40	0.15	0.21	1.00					
Li ⁺	0.98	0.38	0.13	0.96	0.79	0.39	0.40	0.60	1.00				
Na^+	0.99	0.37	0.13	0.98	0.80	0.46	0.45	0.59	0.97	1.00			
\mathbf{K}^+	0.98	0.38	0.13	0.96	0.79	0.39	0.40	0.60	1.00	0.97	1.00		
Ca^{2+}	0.98	0.36	0.13	0.96	0.79	0.39	0.40	0.64	0.99	0.98	0.99	1.00	

Fig. 5 SOM maps for Cl⁻, Br⁻, Na⁺, Ca²⁺, Mg²⁺, and Li⁺







Ma et al. 2009; Boudibi et al. 2021a). The Jerba shallow aquifer shows high levels of NO_3^- . The measured concentration of nitrate in sampled wells ranges from 39.32 to 196.89 mg/l. By comparing the nitrate levels with the standards set by the World Health Organization (WHO), which sets a threshold of 50 mg/l for NO_3^- , we note that 60.75% of the groundwater samples are beyond the standards and therefore unfit for human consumption. Referring to Fig. 8, NO_3^- concentration showed clear spatial variability. The absence of a sewer system in rural and suburban areas of the island contributes to the increase of nitrate concentration. In Jerba Island, rural domestic sewage is primarily composed of toilet effluent and greywater (water used in washing, bathing, and kitchen).

Coliforms (total and thermotolerant) and E. coli

Bacteriological analysis revealed that 94.9% of groundwater samples had TC densities greater than 10 CFU 100/ ml. Except for three wells that were deeper than 50 mbgl, all samples were tested positive for Thc and *E. coli*. The wastewater discharge in rural regions of the island generates a high bacteriological load. Fecal pollution is instituted by means of septic tanks, which contribute to input fecal germs directly in the aquifer and via surface pollution leaching. In addition, agricultural practices by using manure as natural fertilizer participate in the propagation of bacteriological contaminants. Domestic waste released into abandoned wells also participates in aquifer contamination.

The groundwater pollution risk map based on GCI

The GCI values range from 15 to 32 during the dry season. This index is classified into three categories, which include GCI < 20 (low degree of groundwater contamination), 20 < GCI < 25 (moderate to high degree of groundwater

contamination), and GCI>25 (high degree of contamination index). The groundwater pollution risk map (Fig. 9) shows that the high index values belong to the south and northeast regions, while the low index values are located in the western part of the island. The contaminated parts of the aquifer are affected by several sources of pollution and overexploited. The spatial distribution maps of GCI calculated for 54 wells confirm the contamination of the coastal areas, especially the northern and western parts of the island that are considered as areas of low topography and piezometry. The decrease of the piezometric level in the aquifer generates the inversion of the hydraulic gradient. This fact promotes the transfer of seawater and contaminated waters to the aquifer. The overexploitation of the Mio-Pliocene-Quaternary Jerba aquifer has resulted in a number of drawdown cones, a lower aquifer, and depressions that allow seawater to enter the aquifer (Souid et al. 2018). Souid et al. (2018, 2020) used the traditional geochemical investigation and a simple assessment of the contamination of the Jerba island, but the comparison the GCI elaborated in this study and the other methods (Kharroubi et al. 2012; Telahigue et al. 2018) showed that it is robust and gives the best classification of groundwater quality which is very faithful to the reality of the Island. It includes several parameters that represent the pressures exerted on the aquifer by different components (seawater intrusion, nitrate pollution, and fecal contamination). The calculated GCI values will identify the extent of contamination by combining the different sources of pollution existing on the island. In addition, the usefulness of this index is to assess the groundwater contamination degree and the identification of the major contamination mechanism, by area, depending on the weights assigned to each parameter.

Jerba shallow aquifer is a coastal aquifer under an arid climate, which is influenced by several contamination processes. However, punctual sources of pollution like septic tanks and abandoned wells used as uncontrolled garbage







Fig. 8 Spatial distribution of NO₃⁻ concentration



Fig. 9 Spatial distribution pattern of calculated GCI

dumps are considered as direct sources of groundwater pollution. Previous studies carried out by Souid et al. (2017) have shown severe fecal pollution of the aquifer due to the septic tanks. Agricultural practices also led to an increase in nitrate contents and fecal bacteria by the infiltration of irrigation water loaded with fertilizers and germs. The exploited groundwater quantities of unconfined aquifer on Jerba Island reached 3.65 Mm³/year with available resources not exceeding 3.47 Mm³/year. This deficit balance between available resources and exploitation with a low rate of groundwater recharge, due to the aridity of the climate, aggravates the situation and promotes seawater intrusion processes.

Conclusion

This study presents an approach for the hydrogeological pollution assessment of arid coastal shallow aquifers. Results of the groundwater contamination index (GCI) in this research demonstrated that groundwater pollution at several sites of the island is very intense. The GCI was calculated by admitting class intervals for lithium and nitrate contents, calculated marine intrusion fraction, and fecal tracers. Assessment of results based on the spatial distribution of pollution degree illustrates that the proposed index reveals that 75.91% of the aquifer comes under a significant pollution zone (the coastal regions of the island). The groundwater pollution risk map shows the areas that are most affected by pollution and have high levels of nitrate, lithium, fecal germs, and a high marine intrusion rate. Areas characterized by high GCI are subject to the combined effect of anthropogenic activities and sweater intrusion. The overexploitation of the aquifer leads to a groundwater balance deficit between exploitation and available resources. These results confirm the impact of two major processes involved in aquifer pollution: seawater intrusion, a process that affects coastal areas and is activated by overpumping, and the anthropogenic impact favored by septic tanks and wells used as dumping grounds.

Declarations

Conflict of interest The authors declare no competing interests.

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