ORIGINAL PAPER

Recent alluvial fan developments in Muğla (SW Turkey)

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Received: 22 May 2020 / Accepted: 23 April 2021 / Published online: 4 May 2021 \odot Saudi Society for Geosciences 2021

Abstract

The Muğla Polje (SW Turkey), which formed as a result of the Pleistocene activity of the Muğla Fault, is an example of a rift basin with two alluvial fans (the Muğla Fan, 3.33 km² and the Düğerek Fan, 1.26 km²). This study focuses on the development of these fans through the evaluation of aerial photos, Google-Earth images, the measurements of sedimentary sections, lithological examinations, mineralogical discriminations, and in situ observations. Fractured Jurassic-Cretaceous limestone and upper Miocene-Pliocene clastics have been observed in drainage areas of both fans. In the drainage area of the Muğla Fan, where low strength upper Miocene-Pliocene clastics are abundant, the river incision is higher. As a result, higher amount of sediment supply led to formation of the larger Muğla Fan. Small scale tectonic activities and rainy periods led to the formation of poorly sorted, angular-subangular, and coarser-grained pebblestone-cobblestone deposits, whereas, tectonically quiescent and dry periods were represented by the finer-grained sediments at the apex and upper fan of both fans. They pass into the finer-grained deposits in the downdip direction. Recently, the Muğla and Düğerek fans surface has been covered by urbanization. Man-made channel was built to protect humans from floods, which starts from the apex and extends beyond the distal fan in Polje center. Thus, sediment and water required for fan development have been all cut. Both fan development were initially controlled by nature, such as tectonic activities and climate, and now by humans. Therefore, this study recommends the constant monitoring of the impact of human interference on fan development while regulating the feeding mechanism of fan basin interiors, natural balance of local environment, and urban lives.

Keywords Graben . Polje fillings . Human interference . Recent sedimentation

Introduction

Alluvial fans are in conical-semi conical aggradational sedimentary bodies and arcuate in plain view between the upland catchments-sediment source and depositional basin (Hooke [1967;](#page-18-0) Bull [1972a](#page-17-0), [b;](#page-17-0) Blair [1999,](#page-17-0) [2003](#page-17-0); Coulthard et al. [2002;](#page-17-0) Blair and McPherson [2009\)](#page-17-0). They can be found at the edges of the river or fault controlled valleys (Wasklewicz and Scheinert [2016](#page-19-0)). Lithologically, fan deposits contain mainly coarse-grained and poorly sorted sediments (Blair and McPherson [1994,](#page-17-0) [2009](#page-17-0)).

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Five major factors influence fan processes: (i) the lithology of bedrock in a catchment area, (ii) the shape of catchment area, (iii) the neighboring environment, (iv) the climate and (v) the tectonism and active faults (Schumm [1977](#page-19-0); Coulthard et al. [2002;](#page-17-0) Crosta and Frattini [2004](#page-17-0); Nichols [2005](#page-18-0); Pope and Wilkinson [2005](#page-19-0); Blair and McPherson [2009](#page-17-0); Wasklewicz and Scheinert [2016;](#page-19-0) Ventra and Clarke [2018\)](#page-19-0). Tectonic is considered as the main controlling factor on location, morphology of fans through creating a setting, relief, and accommodation for the alluvial fans (like in the Muğla, SW Turkey) (Ventra and Clarke [2018\)](#page-19-0). Structural-morphological obstacle and to a lesser extent other factors such as river and human impact may prevent building of the regular fan geometry (Crosta and Frattini [2004](#page-17-0)). The climate is considered as a controlling factor in erosion, water discharge, vegetation covers, depositional processes (including sheet and channelized fluvial flow), and debris flows (Hooke [1968;](#page-18-0) Denny [1967;](#page-17-0) Starkel [1983;](#page-19-0) Crosta and Frattini [2004;](#page-17-0) Harvey et al. [2005;](#page-18-0) Ventra and Clarke [2018\)](#page-19-0). The alluvial fans are widely developed in arid climatic conditions; however, change in climate may trigger fan

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formations in various sizes (Harvey et al. [2005](#page-18-0); Ventra and Clarke [2018](#page-19-0)). Some urbanization is located on alluvial fans. Thus human activities such as channels-embankment should affect the fan development, which may be added as the sixth controlling factor on alluvial fan environments (Constante et al. [2010](#page-17-0); Walstra et al. [2010;](#page-19-0) Ortega et al. [2014](#page-19-0); Wasklewicz and Scheinert [2016](#page-19-0); Roberts et al. [2019](#page-19-0)).

The alluvial fans are one of the important landforms in Turkey (Kuzucuoğlu et al. [2019a\)](#page-18-0). Recently, they develop under the dry climatic condition in eastern Turkey and mild climatic condition in western Turkey (Kazancı and Roberts [2019;](#page-18-0) Kuzucuoğlu et al. [2019a](#page-18-0); Özsayın et al. [2019;](#page-19-0) Şaroğlu and Güngör [2019\)](#page-19-0). They can be found in different ages and in diverse tectonic regimes such as normal fault and strike slip fault in Turkey (Brückner [2019](#page-17-0); Kayan [2019;](#page-18-0) Kuzucuoğlu [2019;](#page-18-0) Kuzucuoğlu et al. [2019a,](#page-18-0) [2019b](#page-18-0); Ocakoğlu [2020](#page-19-0); Öner [2019](#page-19-0); Özsayın et al. [2019;](#page-19-0) Sarıkaya et al. [2015](#page-19-0); Şaroğlu and Güngör [2019\)](#page-19-0). SW Turkey hosts four different types of basins during the Neogene-Quaternary; piggyback basin, pull-apart basin along oblique strike-slip faults, intermontane basin, and graben (Görür et al. [1995](#page-18-0); Bozkurt and Mittwede [2005;](#page-17-0) Sözbilir [2005](#page-19-0); Alçiçek [2007](#page-17-0); Gürer et al. [2013;](#page-18-0) Seyitoğlu and Işık [2015;](#page-19-0) Elmas et al. [2019\)](#page-17-0). The alluvial fan, colluvial, and fluvial deposits are found in the Çameli Basin in the SW of Muğla (Alçiçek et al. [2005](#page-17-0); Alçiçek [2007\)](#page-17-0); in the Gediz and Küçük Menderes Grabens (Süzen et al. [2006](#page-19-0); Rojay et al. [2005\)](#page-19-0); in the Manisa Basin of western end of the Gediz Graben (Özkaymak et al. [2013](#page-19-0)); and in the Söke-Milet Basin in western end of the Büyük Menderes Graben in the N of Muğla (Sümer et al. [2012\)](#page-19-0).

The Muğla Polje is a 12-km-long, 6-km-wide NW-SE trending, normal fault-controlled basin and filled with the post Miocene deposits (Gürer et al. [2011](#page-18-0), [2013\)](#page-18-0). Only the basic characteristics of the Muğla and the Düğerek fans in the Muğla Polje were mentioned in previous studies (Figs. [1](#page-2-0) and [2;](#page-4-0) Gürer et al. [2011](#page-18-0), [2013](#page-18-0); Gül [2015](#page-18-0); Küçükuysal et al. [2018\)](#page-18-0). No study has addressed the detailed sedimentological properties of these alluvial fans. How the human interference affect the fan development is less clear. The main purposes of this study are to describe the geometry of the alluvial fans based on aerial photos and Google-Earth images, to define the lithological-facies properties according to field study and borehole data, and to discuss the artificial impact on the fan developments.

Geological setting

The lithology of the region is of Permo-Carboniferous to Recent (Fig. [1\)](#page-2-0). Permo-Carboniferous schist, Triassic-Eocene dolomite-limestone, Lower-Middle Triassic metasandstone-metaconglomerate, and Oligocene conglomerate-sandstone are exposed in the south of the study area. The investigated fans are located in the north part of the Muğla Polje that includes closed karstic depression and polje plain (Kayan [1979;](#page-18-0) Pehlivan [1993](#page-19-0); Güner [2001\)](#page-18-0). The north of the Muğla Polje is bounded by the active fault zone (Gürer et al. [2011,](#page-18-0) [2013](#page-18-0)) namely the Muğla-Yatağan Fault Zone (Şaroğlu et al. [1987](#page-19-0)) or Muğla Fault (Karabacak [2016\)](#page-18-0).

The Gereme Formation, the Yatağan Formation, the Quaternary alluvium, and colluvium are exposed in the northern part of the Muğla Polje (Atalay [1980;](#page-17-0) Konak et al. [1987;](#page-18-0) Göktaş [1998;](#page-18-0) Gürer et al. [2011,](#page-18-0) [2013;](#page-18-0) Gül [2015](#page-18-0)). The northern rocky mountains are formed by the Jurassic-Cretaceous (Jcr), grey-colored, fractured limestones, and dolomites of the Gereme Formation (Figs. [1](#page-2-0) and [2;](#page-4-0) Göktaş [1998](#page-18-0); Gürer et al. [2011](#page-18-0), [2013](#page-18-0)). The Yatağan Formation (upper Miocene-Pliocene) composed mainly of red-brown colored muddy conglomerate and mudstone alternations, and composed to a lesser extent light grey limestone-marl, tuff and travertine intercalations, which were created by horizontal-flat topography on the Jurassic-Cretaceous limestone (Atalay [1980;](#page-17-0) Konak et al. [1987;](#page-18-0) Aktimur et al. [1996;](#page-17-0) Göktaş [1998;](#page-18-0) Gürer et al. [2011,](#page-18-0) [2013\)](#page-18-0). Narrow colluvial wedges, which composed of very poorly to poorly-sorted, angular to subangular, and clay to boulder sediments, have linear extensions in front of the northern faulted margin of the Muğla Polje (Gül [2015\)](#page-18-0). The alluvial fans include angular to subrounded, poorly-sorted, clay to cobble-sized sed-iments (Gürer et al. ([2011,](#page-18-0) [2013\)](#page-18-0). To the Polje center, both colluviums and alluvial fan deposits laterally pass into the subrounded to rounded, poorly to moderately sorted, dominantly red-colored, clay and gravel-sized sediments bearing alluviums (Kayan [1979](#page-18-0); Aktimur et al. [1996](#page-17-0); Göktaş [1998;](#page-18-0) Gürer et al. [2011](#page-18-0), [2013](#page-18-0); Gül [2015;](#page-18-0) Küçükuysal et al. [2018](#page-18-0)).

Material and methods

The studied fans are located \sim 2 km east of the Muğla district (Figs. [1](#page-2-0) and [2](#page-4-0)). Throughout the province of Muğla, a Mediterranean climate prevails (Köppen-Geiger classification), with warm-dry summers and cold-rainy winters (annual average precipitation of 1101 mm and annual average temperature of 15.3°; [https://tr.climate-data.org/asya/tuerkiye/mugla/](https://tr.climate-data.org/asya/tuerkiye/mugla/mentese-923599/) [mentese-923599/](https://tr.climate-data.org/asya/tuerkiye/mugla/mentese-923599/) Access date: 25.01.2019).

The aerial photos and Google-Earth images were used to determine the alluvial fan boundary and settlement changes covering the last 45 years. Digital Elevation Model (DEM) and Google-Earth images were used to describe the geometry, transverse, and longitudinal profiles of fans. The lithological properties of the fan deposits were examined by field studies and the borehole data from the report of the Menteşe District of Muğla Metropolitan Municipality Microzonation (MMMM [2017](#page-18-0)). Results of sieve analysis and Unified Soil Classification System (USCS [1985\)](#page-19-0) were used for sediment classification based on Folk [\(1954\)](#page-18-0) system.

A sedimentary section was measured to determine the lithological variations in the main channel feeding the Düğerek Fan. Samples from selected beds of the Düğerek Fan (D-1) and the Yatağan Formation (D-2) were used to compare the mineralogical properties of the fan deposit and its possible source lithology. Cation exchange capacity (CEC) and Fourier transform infrared spectroscopy (FTIR) analyses, due to their fast-analytical capability and applicability in the present laboratory conditions, were selected to compare the fine fraction mineralogy of both samples (D-1 and D-2).

Cation exchange capacities of the samples were determined based on the Methylene Blue Test. The methods of Jones [\(1964\)](#page-18-0) and Çokca and Birand ([1993](#page-17-0)) were followed for the preparation of Methylene Blue (MB) solution and the determination of the CEC of the fine fractions of the samples. MB adsorption of the clay-water mixture was monitored until the complete exchange reactions happen between clay surface and the MB solution. The CEC values were calculated based on Çokca and Birand ([1993\)](#page-17-0) with the provided formula (1):

$$
(C.E.C.) = 100/f*Vcc*NMB = (meq/100gr)
$$
 (1)

where f stands for a dry weight of clay sample (7.5 gr), Vcc is the amount of methylene blue diluted added to the clay sample, NMB is the normality of the MB solution $(\text{standard} = 0.028)$.

Infrared (IR) spectra of the two samples were recorded in transmittance mode as a function of wavenumber using a Thermo Scientific Nicolet iS10 Smart iTR from 4000 to 400 cm^{-1} with a number of scan 32 and resolution 16 cm^{-1} . Attenuated total reflection FTIR (ATR-FTIR) technique was selected for FTIR analysis. FTIR-spectroscopy enables to determine the minerals, non-crystalline phases as well as organic matter present in the samples. As stated in Vaculíková and Plevová, (2005), absorption of infrared radiation is also affected by the crystallinity of the minerals and their grain sizes.

Characteristics of the fans

Muğla Fan

Morphology The Muğla Fan covers approximately 3.33 km^2 , and its perimeter is roughly 7 km. It has almost radial distribution in plain view (Figs. 1, [2,](#page-4-0) and [3](#page-5-0)). It extends longitudinally 2.2 km (maximum) towards SW direction. An altitude of the fan is 700 m in apex part and 625 m in the distal part. The fan, located in front of the normal faulted hills, formed in the upper Miocene-Pliocene clastics (Fig. [3](#page-5-0)). The river, bounded with 70–80° inclined valley walls, incises into this older clastics and feeds the Muğla Fan (Fig. [2\)](#page-4-0). A 850-m high hill is located at the apex part of the fan, skirts of the hill has a 45–

Fig. 1 (A) Main tectonic features of Turkey (modified from Savaşçın and Oyman [1998\)](#page-19-0). (B) General geological map of the study area (modified from Aktimur et al. [1996](#page-17-0); Gül et al. [2013](#page-18-0); Akbaş et al. [2011](#page-17-0); and [http://](http://yerbilimleri.mta.gov.tr/anasayfa.aspx)

[yerbilimleri.mta.gov.tr/anasayfa.aspx:](http://yerbilimleri.mta.gov.tr/anasayfa.aspx) access date 10.04.2019). (C) The columnar section of the study area (modified from Aktimur et al. [1996;](#page-17-0) Göktaş [1998](#page-18-0); Gül [2015\)](#page-18-0)

C System	Serie	Group	Format	Lithology	Explanations
Quaternary			Alluvial Fan Deposits (Qaf) Colluviums (Qac) Alluviums (Qal)	\circ . \circ \circ оŚ a Qac 'Qaç Qaf Qal _c	Qal: gravel, sand, mud Qaf: Conglomerate, breccia, mud Qac: Breccia, mud
	Jp. Miocene- Pliocene	Muğla Group	Yatağan Formation		Disconformity Conglomerate, Sandstone, Mudstone Disconformity
	ow. Miocene Oligocene?	Akçay Group	Kerme Formation		Conglomerate, Sandstone
Cretaceous Jurassic-		Ören Group	Gereme Formation		Disconformity Limestone Dolomite
Triassic	Lower- Middle		Karaova Formation	\circ \circ \circ \circ \circ \circ \circ \circ \circ \circ \circ \circ \circ	Metasandstone Metaconglomerate
Eocene Triassic \blacksquare		Marçal Group			Tectonic Contact Dolomite Limestone
arboniferous Permo-		Kavaklıdere Group			Disconformity Schist Low.: Lower Up.: Upper Unscaled

Fig. 1 (continued)

80° slope. The transverse sections of the Muğla Fan are planoconcave upwards geometry in the middle part, and wedge shaped at the edges. The apex part of the fan overlies the Yatağan Formation with angular unconformity. However, this relation was not clearly observed during the field studies due to soil cover and artificial fillings. The longitudinal section shows wedge shape of the fan deposits that laterallygradually pass into the alluvial-colluvial deposits (Fig. [3](#page-5-0)). The upper part of Muğla Fan deposits is topographically higher than the alluvial deposits (Figs. [2](#page-4-0) and [3](#page-5-0)). The older settlements of the Muğla City are located on top of the Yatağan Formation, and Quaternary alluvial fan depositscolluviums (Fig. $3(A)$). Recent settlements occupy the all fan deposits (Figs. [2](#page-4-0) and $3(B)$). After heavy rains (during winter), flash flooding waters directly flow to the western part of the Muğla Fan, cause local shallow ponding in a days. The pond area is larger than in the Düğerek area. Depending on the continuity of precipitation and temperature, ponding can survive a few weeks to one or two months.

Sedimentological properties The borehole data reveals that the proximal part of the fan contains the GM-GP (moderately to poorly graded gravels according to USCS ([1985](#page-19-0)) – poorly sorted gravels) and the msG (muddy sandy Gravels and Gravel according to Folk [\(1954\)](#page-18-0) classification). The GP-GM (poorly to moderately graded gravels according to USCS [\(1985](#page-19-0)) – moderately to well sorted gravels), the msG and mG (muddy sandy Gravels and muddy Gravel according to Folk [\(1954\)](#page-18-0) classification) increase towards the mid fan and distal fan (Fig. [4\)](#page-6-0). Locally, as in M-8 borehole (MMMM

Fig. 2 (A) The Google-Earth view of the Düğerek and Muğla Fans in the Muğla Polje (Google-Earth [2020\)](#page-18-0). (B) The DEM view of the Düğerek (blue arrow) and Muğla Fans (green arrow; modified from Gül et al. [2013](#page-18-0))

[2017\)](#page-18-0), the GW (well graded-poorly sorted gravels) and the G (gravels) type sediments were detected. The mud size sediment bearing the CL (low plasticity Clay) – the gM or (g)sM (gravely mud or slightly gravely sandy mud) type sediment, the SM (silty sand), and the gM (gravely Mud) type sediments were also recorded (Figs. [4](#page-6-0) and [5\)](#page-7-0).

Due to the settlements, roads, parks, artificial fillings, and soil covers, there are not almost any field exposures of the Fig. 3 (A) The aerial photo (1976) of the Mu ğla Fan, green arrows indicate approximate outer boundary of the Mu ğla Fan (modified from courtesy of the General Directorate of Mapping; 127, 1976, 153 21 3 78). (B) The Google-Earth view of the Mu ğla Fan (Google-Earth [2020\)](#page-18-0). (C-D) Longitudinal (A-A ') and transverse (B-B ') cross-sections of the Mu ğla Fan

Fig. 4 Some borehole information's belongs to the different part of the Muğla Fan compiled from microzonation report of Menteşe District of Muğla Metropolitan Municipality (MMMM [2017](#page-18-0))

Muğla Fan except some basement excavations of the constructions (Fig. $5(A)$ $5(A)$). Figures 5 (B) and (C) indicate the presence of > 5 m thick in the GP-GM (poorly to moderately graded gravels– moderately to well sorted gravels), the msG and the mG (muddy sandy gravels and muddy gravel). These are overlain by the 50-cm-thick muddy soil and later by the artificial filling. Figures 5 (D) and (E) point out roughly 9 m thick in the GP-GM (poorly to moderately graded gravels– moderately to well sorted gravels), the msG and the mG (muddy sandy gravels and muddy gravel). These are overlain by the alternation of GM – mG type gravel and paleosoil. All gravels are overlain by the recent soil with an erosive contact. Manmade channel including some stairs for decreasing the flow velocity is located at the apex of the fan (Fig. $5(A)$). NE to SW directed channel carries the excess water and sediments directly to the center of the Polje.

Mineralogical properties The thin sections of the Yatağan Formation include only micrite and calcite mineral of limestone fragments (Zeybek [2017\)](#page-19-0). Soil samples from distal part of the Muğla Fan contain kaolinite mineral based on CEC and XRD analysis (Küçükuysal et al. [2018](#page-18-0)).

Structural properties The Muğla Fault was not clearly detected in this region (Figs. [1\(B\)](#page-2-0) and [2\)](#page-4-0). Gül et al. ([2016](#page-18-0)) reported that Yatağan Formation inclined to NW with an angle of 20°– 30°. They also defined various joint sets (73°/179° Dip/Dip direction; 71°/334°, 67°/174°, 74°/149°; 60°/160°, 72°/197°, 72°/8°, 74°/331°). Those joint sets led to densely fractured rocks at the apex of the Muğla Fan (Fig. [5\(A\)\)](#page-7-0).

Düğerek Fan

Morphology The Düğerek Fan covers 1.26 km^2 area, and its perimeter is 4.53 km long, which longitudinally extends to a maximum of 1.2 km towards the SW direction. It has a radial distribution in down dip direction (Figs. [1,](#page-2-0) [2](#page-4-0), and [6](#page-9-0)). The transverse sections of the Düğerek Fan illustrate the plano-

Fig. 5 Field photos of the Muğla Fan. (A) The apex part of the Muğla Fan is following the valley incised into the upper Miocene-Pliocene clastics of the Yatağan Formation. Man-made channel directly transfers the water from upper fan to the distal fan. Green arrows show a previous sediment

route that recently occupied by buildings. (B, C) The basement photo from 200 m south of the M5 borehole. (D, E) The basement photo from 500 m northwest of the M9 borehole

concave upward geometry in the mid-part and wedge shape at the edges. An altitude of the fan is 750 m in apex part and extend to 620 m in the distal part. The fan is located in front of the normal faulted hills formed by the Jurassic-Cretaceous limestone and upper Miocene-Pliocene clastics. The rivers, bordered by 50–60° inclined valley walls, incised into older rocks and fed the Düğerek Fan (Fig. [2](#page-4-0)). A high of 950 m hills are located at the apex part of the fan; skirts of the hills have a 40–70° slope. Nearly horizontal and horizontal bedded, poorly sorted breccia of the fan deposits overlie the Yatağan Formation with angular unconformity (Figs. [6](#page-9-0), [7,](#page-10-0) [8](#page-11-0), and [9\)](#page-12-0). The wedge shaped fan deposits laterally and gradually passes into the alluvial-colluvial deposits (Fig. [6](#page-9-0)). The Düğerek Fan deposits are topographically higher than the surrounding alluvial deposits (Figs. [2](#page-4-0) and [6](#page-9-0)). The older settlements of the Düğerek district were located on top of the upper fan (Fig. $6(A)$). Recent settlements occupy the all fan deposits (Figs. [2](#page-4-0) and $6(B)$). After heavy rains (during winter), flash flooding waters directly flow to the Polje center to the west of the Düğerek Fan and cause local ponding in a few hours. Ponding time, depending on the continuity of precipitation and temperature, can reach a few weeks to one or two months.

Sedimentological properties The Yatağan Formation (15°/ 310°, Dip/Dip direction) crops out below of the proximal part of the Düğerek Fan (Fig. [8\)](#page-11-0). An alternation of 1–3 m thick, very poorly poorly sorted, clast supported breccia (pebblestone-cobblestone) levels and 5–20 cm thick marlmudstone-calcilutite- rarely tuff levels were found in the Yatağan Formation. The gravels of the breccia levels are limestone fragments of the Gereme Formation and their sizes vary between 5 mm and 40 cm. The interstitial matrix of the gravels is the fine-grained sand and mud-sized sediments (Figs. [8](#page-11-0) and [9\)](#page-12-0).

The Düğerek Fan starts with deformed, very poorly sorted, clast-supported breccia formed by subangularangular limestone fragments of the Gereme Formation and lesser extent rounded limestone fragments of the Yatağan Formation (maximum clast size (MCS): \sim 35– 40 cm, average clast size $(ACS): ~ 8–10$ cm) and crudely bedded units on footwall of the Muğla Fault (Figs. [7,](#page-10-0) [8](#page-11-0), [9,](#page-12-0) and [10\)](#page-13-0). A 5-m section was measured from the southern side of the fault scarp in the continuation of the fan (Figs. [9](#page-12-0) and $10(A-C)$ $10(A-C)$). This part of the fan contains horizontally bedded breccia levels, whose thickness varies between 10 and 25 cm. The bottom part of the section is composed of relatively durable breccia level, including very poorly sorted, clast-supported, or matrix-supported cobblestone levels (Fig. $9(D)$). The middle and upper parts of the section include poorly durable, matrix-supported, muddy breccia is covered with limestone fragments (Fig. $9(B)$). The red colored muddy matrix quantity increases towards the upper levels (Fig. $9(C)$).

Through the mid-fan and distal-fan, the coarser proximal fan deposits pass into the poorly-sorted, matrix-supported muddy breccia including limestone fragments in 5 mm to 10 cm size (Fig. [7\)](#page-10-0). The amount of mud in these levels is considerably higher than the proximal part (Fig. [7](#page-10-0)). The borehole data points out that the proximal part of the fan contains the GM-GW (moderately to well-graded gravels according to USCS [\(1985\)](#page-19-0) – poorly sorted gravels) and, the msG and G (muddy sandy gravels and gravel according to Folk [\(1954\)](#page-18-0)). The GP-GM (poorly to moderately graded gravels are according to USCS [\(1985\)](#page-19-0) – moderately to well sorted gravels) the msG and mG (muddy sandy gravels and muddy gravel according to Folk [\(1954](#page-18-0))) increased towards the mid fan and outer fan (Fig. [7\)](#page-10-0).

NE-SW directed man-made channel is extending from the apex part to the distal fan and ends within the alluvial deposits. This channel carries both the sediment and water, especially during the flash flood time. Excess sediment in channels cleaned periodically by the excavator and accumulated on the channel bank (Fig. $10(C)$). The Düğerek Fan deposits pass into the alternations of mudstone and sub-rounded gravelbearing conglomerate (Fig. [10\(D\)](#page-13-0)) (Küçükuysal et al. [2018\)](#page-18-0) in 1 km south of the apex-proximal part.

Mineralogical properties The mineralogical composition of red colored mudstone part of the Yatağan Formation was determined by CEC and FTIR analysis (Fig. $9(E)$). CEC value for D-2 was determined as a 5.22 meq/100g. Based on Grim [\(1968](#page-18-0)), the CEC value falls in a CEC range of kaolinites. Figure [9](#page-12-0) (E) shows the FTIR spectra of D-2 with wellresolved bands of kaolinites in OH stretching vibrations of surface hydroxyl groups $(3742, 3675, 3650 \text{ cm}^{-1})$ and one stretching vibration of inner hydroxyl groups (3618 cm^{-1}) (Saikia and Parthasarathy [2010;](#page-19-0) Djomgoue and Njopwouo [2013\)](#page-17-0). Additionally, 875 cm^{-1} band corresponds to Al-OH bending vibration of kaolinite (Saikia and Parthasarathy [2010\)](#page-19-0). Vibrations at 1744, 1457, and 875 cm $^{-1}$ are assigned to calcite (Vaculíková and Plevová [2005](#page-19-0); Saikia and Parthasarathy [2010](#page-19-0)). Quartz was defined with Si-O bands at 1048 cm $^{-1}$ together with the doublet at 778–798 cm $^{-1}$ (Vaculíková and Plevová [2005\)](#page-19-0). The existence of low intensity bands in $600-400$ cm⁻¹ range (Saikia and Parthasarathy 2010) and 1649 cm⁻¹ vibration figure out the presence of hematite. The results of D-2 sample revealed that the existence of kaolinite as dominant clay mineral associated with quartz, calcite, and to a lesser extent hematite (Fig. $9(E)$). Calculated CEC value for D-1 (mud part of the Düğerek Fan) (6.72 meq/100 g) and almost the same FTIR-ATR spectrum as $D-2$ (Fig. $9(E)$).

Structural properties The active Muğla Fault has created several fault scarps (Fig. $10(C)$) that cut Jurassic-Cretaceous limestone, upper Miocene-Pliocene clastics (Fig. $10(A)$), and

Fig. 6 (A) The aerial photo (1976) of the Düğerek Fan, blue arrows indicate the approximate outer boundary of the Düğerek Fan (modified from courtesy of the General Directorate of Mapping; 137, 1976, 153.21 3.78). (B) The Google-Earth view of the Düğerek Fan (Google-Earth [2018\)](#page-18-0). (C, D) Longitudinal (A-A') and transverse (B-B') cross sections of the Düğerek Fan

Quaternary alluvial fans (Fig. $10(C)$) in the apex part. Figures [11](#page-14-0) (A) and (B) show the topography and recent view of the Düğerek Fan. Two rivers cut the Gereme and Yatağan formations in the drainage feeding region of the Düğerek Fan, which merge in the proximal-apex part of the fan. Then, a single river valley cut and erodes the fault plane, the Yatağan Formation and the older Düğerek Fan deposits (Figs. [2,](#page-4-0) [6](#page-9-0), and $11(A)$). Pre-Pliocene topography of the northern Muğla region includes rugged topped Jurassic-Cretaceous limestone and nearly flat topped upper Miocene-Pliocene clastics (Fig. $11(C)$). Subsequent normal faulting cuts both formation and led to the formation of the Muğla Polje (Fig. $11(D)$) and numerous fault scarps as shown in Figs. [8,](#page-11-0) [10](#page-13-0) (A), and 10 (C). Then, rivers eroded the older formations and carry the sediments to both formations and created colluviums and alluvial fan deposits, Fig. $10(B)$ (Fig. $11(E)$). The latest activities' boundary fault cut all three units and present recent form of the Polje (Fig. [11\(F\)](#page-14-0)).

Discussion

The Muğla and Düğerek Fans have similar properties such as radial shape in plain view (Figs. [1](#page-2-0) and [2](#page-4-0)), poorly sorted gravel bearing deposits (Figs. [4](#page-6-0) and 7), occupied by settlements (Figs. $5(A)$ and $11(B)$), and limitation in development due to man-made activities (Figs. $5(A)$ and $10(C)$). However, they are also some significant size difference, even in close distance, ranging from feeding from identical provenance, development under the same climatic condition to tectonic activities. The Muğla Fan has been fed by the relatively big stream. This big stream has a large drainage area and is deeply incised into upper Miocene-Pliocene clastics and to a lesser extent Jurassic-Cretaceous limestone (Figs. [1](#page-2-0) and [2](#page-4-0)). The Düğerek Fan has been fed by two small rivers which has two small branches. These branches have small drainage area and are relatively surficial/incised into Jurassic-Cretaceous limestone and to a lesser extent the upper Miocene-Pliocene clastics (Figs. [1](#page-2-0) and [2](#page-4-0)). Strength of the upper Miocene-Pliocene clastics is lower than the Jurassic-Cretaceous limestone. The Muğla Fan is three times bigger than the Düğerek Fan (Figs. [1](#page-2-0) and [2\)](#page-4-0). Thus, more sediments have been obtained from the Muğla Fan drainage area. Other differences are related with tectonic activity in the apex of fans; for example, the two relatively small segments of the Muğla Fault, which are

Fig. 7 Some borehole information's belonging to the different parts of the Düğerek Fan compiled from the microzonation report of Menteşe District of Muğla Metropolitan Municipality (MMMM [2017](#page-18-0))

Fig. 8 (A) The cross-section of apex of the Düğerek Fan, whose deposits overlie the Yatağan Formation with angular unconformity. (B) Sketch diagram of the cross-section given in A

located at the apex of the Düğerek Fan (Fig. [11\(E\)](#page-14-0)). Deformation of fan sediments was determined in the Düğerek Fan, while it is not observed in the Muğla Fan.

Tectonic activity and climate were suggested as the two main controlling factors on alluvial fan development (Pope and Wilkinson [2005](#page-19-0); Chen and Guo [2017;](#page-17-0) Bouzari [2020](#page-17-0)). Tectonic processes control the boundary, irregularity of the bottom topography, and the size of the sedimentary basin (Harvey et al. [2005](#page-18-0); Gül et al. [2013](#page-18-0); Gül [2015;](#page-18-0) Bouzari [2020\)](#page-17-0). Both tectonic process and climate govern the amount-input-transportation of sediment and water (Harvey et al. [2005](#page-18-0); Gül et al. [2013;](#page-18-0) Gül [2015](#page-18-0); Houa et al. [2020](#page-18-0); Blackburn et al. [2021\)](#page-17-0). In addition to these factors, gravity, host rock type, drainage pattern, and human impact were all proposed as the other controlling factors for the development of alluvial fans (Coulthard et al. [2002](#page-17-0); Crosta and Frattini [2004;](#page-17-0) Nemec and Kazancı [1999;](#page-18-0) Harvey et al. [2005](#page-18-0); Shukla [2009](#page-19-0); Sanders [2010](#page-19-0); Walstra et al. [2010;](#page-19-0) Villacorta et al. [2019\)](#page-19-0).

Tectonic

Normal fault (numerous example from Turkey have been listed in the introduction section), thrust fault (Shukla [2009;](#page-19-0) Bouzari [2020\)](#page-17-0), and strike slip fault (Şaroğlu and Güngör [2019\)](#page-19-0) activities can lead to form of alluvial fans. Ocakoğlu [\(2020\)](#page-19-0) emphasized that the tectonic activity initiated a rapid geomorphologic change in the high angle normal faulted Büyük Menderes Graben. Polje evolution in the Isparta Angle with marsh and alluvial fillings in the east of the Muğla Fan was developed due to tectonic subsidence, karstification, and climatic changes (Deynoux et al. [2005;](#page-17-0) Kuzucuoğlu et al. [2019a\)](#page-18-0). Development of the tectonokarstic plain in Central Anatolia with a river deposits, alluvial

Fig. 9 (A) Measured section of the proximal part of the Düğerek Fan. (B) General field view of the location of the cross-section. (C) Close-up view of the fan deposits. (D) Close view of the contact between the Düğerek

fan, and lacustrine deposits was related with tectonic activity, karstification, and river erosion (Kuzucuoğlu et al. [2019a\)](#page-18-0).

Indo-Gangetic Plain in N of India consists of numerous laterally coalescing mega alluvial fans (e.g., Gaula Fan, 30 km long) that developed under the thrust tectonic activity of the basin margin (Shukla [2009\)](#page-19-0). Thrust tectonic along the Central and Eastern Alborz Mountains in the Eyvanekey, Garmsar, Semnan areas, northern Iran led to formation of the 131 alluvial fans (Bouzari [2020\)](#page-17-0). Larger primary fans initiated with under the effect of the high tectonic subsidence and uplifting, secondary fans developed on first fans under the effect of the minor faults (Bouzari [2020\)](#page-17-0). Normal tectonic activity initially controlled the compartment of the rifting, then basement topography, stratigraphic architecture and spatial distribution of the depositional basin in Guaizihu sag, Yingen-Ejinaqi Basin in China (Houa et al. [2020\)](#page-18-0). Guaizihu sag contains graben-half grabens separated by local highs and alluvial fan and fan deltas (Houa et al. [2020](#page-18-0)). Fault segments

Fan and the Yatağan Formation clastics. (E) FTIR spectra of the muddy matrix of breccia in proximal part of the Düğerek Fan (Quaternary, D-1) and mudstone Yatağan Formation (upper Miocene-Pliocene; D-2)

created topographical highs acts as sediment source for alluvial fans (Houa et al. [2020\)](#page-18-0).

Moreover, highly jointed rocks allow more physical weathering as well as higher amounts of sediments (Pope and Wilkinson [2005\)](#page-19-0). Sequences with upward fining characterize the tectonically quiescent periods, while the coarsening upward sediments typify the intense tectonic activity period (Fidolini et al. [2013](#page-17-0)). Similarly, breccia level in colluvial wedge can be formed as a result of the tectonically active periods, while subsequent muddy parts represent the quiet period (Gül [2015\)](#page-18-0). Thus, the alluvial sediment ages can be used for the interpretation of the activation age of the faults (Sarıkaya et al. [2015\)](#page-19-0). The asymmetric profile of the Pliocene Eşen Graben in SW Anatolia was the result of intense tectonic activity and sediment input (Öner [2019](#page-19-0)).

Tectonic activities are the primary controlling factor for both the Muğla and Düğerek fan developments. Muğla and surrounding region have been developed as a depressional

Fig. 10 (A) Two faults cut the Yatağan Formation and Düğerek Fan at the proximal part of the fan (faulting at the lower level and erosion cause the hanging valley). (B) General field view of the deformed Düğerek Fan deposits that attached to the footwall of the Muğla Fault Plane (Jcr: Jurassic-Cretaceous Limestone; Ty: Yatağan Formation; Qaf: Düğerek Fan deposits). (C) Man-made irrigation channel extends from the apex

to the distal fan. Heavy rainfall at the mid December 2020 carries sediments to downdip direction from apex. Those sediments have been cleaned by the excavator and pour on to embankment. (D) The fan deposits pass into the subrounded-rounded limestone gravel bearing the muddy Gravel, and soils of the alluvial deposits at 1 km southwest of apex

basin and filled with clastics during the post-Pliocene (Kayan [1979\)](#page-18-0), or late Miocene (Gürer et al. [2011,](#page-18-0) [2013](#page-18-0)). Intense tectonic activities during the Pleistocene created half graben-Muğla Polje (Kayan [1979](#page-18-0)). Those tectonic activities also led to some local highs that separated the Muğla and Düğerek Fans (e.g., by Houa et al. [2020](#page-18-0)). They also shaped a small, suitable, flat, and unhampered environment for fans. Complex tectonic history of the region also produced closely jointed rocks. Sezer [\(2003\)](#page-19-0), Karabacak [\(2016\)](#page-18-0) and Türe [\(2017\)](#page-19-0) have emphasized the seismic activity of the Muğla Fault since ancient times. Perpendicular to the basin margin faults, structural alignments created weakness zone along small and large valleys, which transport the sediment down to dipper areas. Fault segments act as a sediment source that similar to Houa et al. [\(2020\)](#page-18-0) observation. Big-scale tectonic activities caused sediment tilting and throwing like in the Düğerek Fan (Fig.

[11\(](#page-14-0)C–F)). Small scale tectonic activities with ground shaking have triggered the evolution, removing, and initiation of the downdip transportation of unstable–loose sediments on hinterland-source rocks. The tectonically active periods resulted in the deposition of coarse-grained sediments, while quiescent periods bring along the development of relatively finergrained, muddy, and matrix-supported sediments. Due to the conical shape of the fans, both fans were evaluated as an unconfined fan and aggrading fan (Crosta and Frattini [2004;](#page-17-0) Scheinert et al. [2012](#page-19-0)).

Climate controls both physical and chemical weathering of rocks and transportation of sediments (Clemmensen et al. [2001;](#page-17-0) Mackel et al. [2003;](#page-18-0) Hartley et al. [2005;](#page-18-0) Löwner et al.

Fig. 11 Geological evolution of the Dü ğerek Fan. (A) The Yata ğan Formation (upper Miocene-Pliocene; Ty) overlies the Gereme Formation limestone (Jurassic-Cretaceous; Jcr) with angular unconformity during late Miocene-Pliocene (Kayan [1979\)](#page-18-0). (B) Field view of the Dü ğerek Fan. After the heavy rainfall at the mid December 2020, surface waters and percolated waters were accumulating at the Polje center. (C) After deposition of the Yata ğan Formation, the Mu ğla Fault was active, and led to formation of the Mu ğla Polje (Pleistocene, Kayan [1979](#page-18-0); end of Pliocene Gürer et al. [2011](#page-18-0) , [2013\)](#page-18-0). (D) The Dü ğerek Fan deposits (Plio-Quaternary; Qaf) overlie the Yata ğan Formation with angular unconformity. (E) Quaternary activity of the Mu ğla Fault (Türe [2017\)](#page-19-0) was also cut the Dü ğerek Fan deposits. (F) Continuous erosion and deformation have been supplying the recent topographical view of the Dü ğerek Fan and surroundings

 $\mathbf b$

[2005;](#page-18-0) Boggs Jr [2006](#page-17-0); Blackburn et al. [2021\)](#page-17-0). Climatic conditions were suggested as a main control on sediment supply, lithofacies, and sedimentary features in Guaizihu sag, Yingen-Ejinaqi Basin in China (Houa et al. [2020\)](#page-18-0). Similarly, changing climate controlled the sediment supply, fluctuating water budget, lateral migration of channel of the Gaula Fan in N India (Shukla [2009\)](#page-19-0). Humid climate promoted the gravel size sediments carried by the powerful rivers with a steeper gradient (fan expansion cycle A of Gaula Fan), while dry phase represented by debris flow sedimentation deposited due to catastrophic rains (cycle B) (Shukla [2009\)](#page-19-0). Alluvial fan sediment characteristics including permeability organized the surface water flow, infiltration to the ground, and so hydrogeological condition (Blackburn et al. [2021](#page-17-0)). Those water also is feeding the ephemeral stream and lake at the end of the alluvial fan (Blackburn et al. [2021](#page-17-0)). Moreover, the freeze-thaw weathering and wetting-drying cycles under the climatic effects are the main physical weathering processes (Boggs Jr [2006\)](#page-17-0).

Gül ([2015](#page-18-0)) proposed that recent climatic conditions of the Muğla (including rainy winter and spring season; daily and seasonal temperatures differences) and closely jointed rock, promote the disintegration of source rock (physical weathering), and therefore controls the sediment quantity. The coarser-grained fan sediments signify intense rainy seasons. On the contrary, dry-arid climatic conditions resulted in less amount of water and sediment inputs, and compose the fine-grained sand-silt-clay dominated sediments. CEC values (Fig. $9(E)$) confirmed the mineralogical compositions and the dominant clay mineral phase of the Düğerek Fan (D-1) to be identical to that Yatağan Formation (D-2). Presence of kaolinite as predominant clay mineral (low CEC values and intense spectral vibrations) in the bulk compositions of samples reveals the intense weathering and subsequent leaching. Zhang et al. ([2016](#page-19-0)) proposed that kaolinite product of highly hydrolytic weathering under the effect of warm-humid climate with a minimum perennial temperature ~ 15 °C like in the Muğla. The hematite reddening represents mild to intense hydrolysis conditions due to the high internal drainage in such karstic nature of the carbonate host rocks (Boero and Schwertmann [1989\)](#page-17-0). Fan properties affect the hydrogeological condition, such as surface flow and ground water circulation. They cause short-term ponding after heavy rainfall-flashflood.

Gravity

Slope with gravity effect diverts the sediment and regulates the sediment transportation type (debris flow, grain flow, rock fall) in recent sedimentation (Miall [1977](#page-18-0); Sanders [2010;](#page-19-0) Gül [2015\)](#page-18-0). Facies A (matrix supported cobble conglomerate), facies B (pebble conglomerate), facies C (pebbly sandstone), facies D (fine-grained sandstone), and facies E (mudstone) were separated in the Fangyan Formation in Danxia landform in Jianglangshan geopark in SE China based on lithology, sedimentary structures, bed thickness, and geometry (Chen and Guo [2017\)](#page-17-0). Stream-flow dominated sediments were responsible from formation of the Fangyan Formation along the mountain fronts (Chen and Guo [2017](#page-17-0)). The poorly sorted, matrix-clast supported, disorganized granule-cobble conglomerates were alternated mudstone and sandy mudstone (in lesser amount) in alluvial fan environment of the Guaizihu sag, Yingen-Ejinaqi Basin, China (Houa et al. [2020](#page-18-0)). Disorganized matrix-clast supported conglomerates were interpreted as a product of the pseudoplastic (or viscous) debris flow or hyperconcentrated flood flow on the proximal fan, while sandy conglomerates and pebbly sandstones interpreted as a density flows or cohesionless debris flow (Houa et al. [2020\)](#page-18-0). Four facies (disorganized gravel, horizontal bedded gravel, cross bedded gravel, sandy gravel and mottled sandy silt in proximal fan; horizontal bedded gravel, cross bedded gravel, laminated silty mud, mottled sandy silt in midfan; cross bedded sand, laminated silty mud, mottled sandy silt in distal fan; marix supported gravel, mottled sandy silt) were defined in 30-km-long Gaula Fan sequence in N India (Shukla [2009\)](#page-19-0).

Similarly, high slope of both fan drainage areas and highly inclined fault scarp at the Polje margin led to form a debris flow dominated products such as poorly sorted, angularsubangular gravel (pebble-cobble breccia, GW-GM-GP types, G-msG types) at the apex and upper fan area. Several studies confirmed that sediment grain size is decreasing to downdip direction in alluvial fan (Alçiçek et al. [2005](#page-17-0); Arzani [2005;](#page-17-0) Rojay et al. [2005;](#page-19-0) Shukla [2009](#page-19-0); Sümer et al. [2012;](#page-19-0) Özkaymak et al. [2013;](#page-19-0) Houa et al. [2020\)](#page-18-0). Similarly, the coarser deposits of Muğla and Düğerek fans pass into the pebblefine cobble bearing muddy breccia (GP-GM types, msG-mG types) (waning flood deposits), which are mostly soft, reddish brown mudstones with desiccation cracks and plant debris, and product of waning flow at middle and distal part of the fan. Both fan deposits pass into fluvial clastics including pebble-cobble conglomerate and thick red colored mudstone. These sediments are probably the product of hyperconcentrated debris flow and debris flow (Hartley et al. [2005;](#page-18-0) Harvey et al. [2005](#page-18-0); Houa et al. [2020](#page-18-0)).

Anthropogenic causes

Alluvial and colluvial sediments offered preferable settlement areas due to agriculturally fertile soil and soft-easily workable ground for the construction (Bauziene [2002](#page-17-0); Mackel et al. [2003;](#page-18-0) Constante et al. [2010](#page-17-0); Villacorta et al. [2019\)](#page-19-0). Ventra and Clarke [\(2018\)](#page-19-0) emphasized that the topographic changes, inundation, and run out in the length of sediment transport are complicated in urbanized areas on fan surfaces. Humans settled on the fan surfaces are under the risk of flood and debris flow hazards (Ortega et al. [2014](#page-19-0); Ventra and Clarke [2018\)](#page-19-0). An

intensive agriculture, reforestation, and urbanization change the water velocity, upstream channelizing-scouring, bank undercutting, which decrease the sediment load that triggers the channel incision, and induces entrenchment and narrowing the fan (Ortega et al. [2014\)](#page-19-0). Man-made levee, ancient irrigation channels, and dam construction played an important role in the rapid fan development and maintaining low gradient of the fan in the Lower Khuzestan (SW Iran) (Walstra et al. [2010\)](#page-19-0). Wasklewicz and Scheinert [\(2016\)](#page-19-0) showed the channelization effect on the embryonic-telescopic fan development in a few years. The capital city of Peru, Lima, is also located on top of the alluvial fan, 31% of the Peruvian live on it (Villacorta et al. [2019\)](#page-19-0). Earthquakes, floods, debris flows, and landslides caused a significant death toll in this region (Villacorta et al. [2019\)](#page-19-0). Rimac river channel migration and sediment transportation has been affected by the intense human intervention (Villacorta et al. [2019\)](#page-19-0). Sediments in man-made channels decrease the river drainage efficiency, thus settlements near the present river channels under the risk of fast moving floods (Villacorta et al. [2019](#page-19-0)). Roberts et al. [\(2019\)](#page-19-0) found catchment erosion during the intense human land use that decrease forest area, more sediment supply to the Nar Lake in the Central Anatolia. SW Turkey includes several grabens filled by the Neogene to Quaternary deposits (Alçiçek et al. [2005;](#page-17-0) Bozkurt and Mittwede [2005;](#page-17-0) Sözbilir [2005](#page-19-0); Rojay et al. [2005;](#page-19-0) Süzen et al. [2006](#page-19-0); Alçiçek [2007](#page-17-0); Sümer et al. [2012](#page-19-0); Özkaymak et al. [2013;](#page-19-0) Seyitoğlu and Işık [2015](#page-19-0); Sümer et al. [2018;](#page-19-0) Elmas et al. [2019\)](#page-17-0). Similar to the Muğla Polje, several settlements were marked on top of the alluvial fan, fluvial deposits such as the Söke town (Sümer et al. [2012\)](#page-19-0), Ödemiş, Tire, Kiraz, Bayındır towns (Rojay et al. [2005\)](#page-19-0) in north of the Muğla.

Early settlements in the Muğla region were in the northern rocky area, especially in ancient times (Fig. $3(A)$, $5(A)$, and [6\(A\)](#page-9-0); Güner [2001](#page-18-0)). The aerial photos in 1976 reveal that the distribution of the settlements concentrated on the apex part of the fan deposits (Figs. $3(A)$ and $6(A)$). However, after the establishment of a university, the population drastically increase to over a hundred thousand ([https://www.nufusu.com/](https://www.nufusu.com/ilce/mentese_mugla-nufusu) [ilce/mentese_mugla-nufusu,](https://www.nufusu.com/ilce/mentese_mugla-nufusu) Access date 11.05.2020). These settlements cover the whole alluvial fan deposits and extend to the alluviums (Figs. [3,](#page-5-0) [5](#page-7-0), [6](#page-9-0), and [11\)](#page-14-0). Man-made channel construction (Figs. $5(A)$ and $10(C)$), extension of the previous channels and asphalt-impermeable roads directed the surface water flow into the Polje center (Pehlivan [1993](#page-19-0)). Thus, excess water has accumulated in Polje center especially during the winter-autumn season, then this water discharged into underground karstic system via dolines (Üstün et al. [2017\)](#page-19-0). After heavy rainfall in mid-December 2020, coarse sediment accumulated in man-made channel in the eastern part of the Düğerek Fan (Fig. [10\(C\)](#page-13-0)). However, surface water inside Polje are immediately ponded in the west of the fan (Fig. [11\(B\)\)](#page-14-0). Man-made channels now carry the sediment and water further away from the fan to the Polje center. The water

initially accumulating in the channel and lowland agriculture area is then leaked to the underground, while the sediments are captured in the channel, and were later cleaned by the excavator in the embankment (Fig. $10(C)$). Hydrogeological properties of both fans, ponding evolution time after rainfall, duration of ponding, and settlements effect on them are required continuous monitoring and researching for sustainable agriculture and irrigation, and for avoiding the negative effects of flashfloods.

Earlier formation of alluvial fans of the Muğla Polje was controlled by tectonic and climate, and recently, the development or renewing of fans have been stopped due to human interference. While the disadvantages created by these new conditions have been proven in different places (e.g., Ortega et al. [2014](#page-19-0); Ventra and Clarke [2018](#page-19-0); Villacorta et al. [2019\)](#page-19-0), this situation has not yet been revealed in the Muğla polje. It requires comprehensive studies in terms of the agriculture, hydrogeology, and socio-economic impact.

Conclusion

The Muğla Polje is one of the good examples of grabens in SW Anatolia, which were filled by the Miocene - Quaternary deposits. Recent sedimentation in the Polje gives an example of the tectonism and climate relation. The Muğla Polje contains two alluvial fans (the Muğla and Düğerek fans) developed in front of the normal faulted basin margin. Both fans unconformably overlie the upper Miocene-Pliocene clastics rocks and Jurassic-Cretaceous limestone. They also laterally gradually pass into the colluvial close to the basin margin and alluvial deposits to the basin center. Tectonism and climate were determined as the main nature controlling factor for these alluvial fans at the beginning. Tectonic activities promote the disintegration of the bedrocks (Jurassic-Cretaceous limestone and upper Miocene-Pliocene clastics) and irregular basement topography of the Polje. The basal topography of Polje and sediment input point lead to form a two separate alluvial fans. The seismically active periods triggered the downdip movement of sediments and created coarse-grained (pebblestone and cobblestone) sediments of the fan. The subsequent tectonic cessation period was represented by the finer-grained-mud dominated levels. Excess and powerful water input during rainy periods carry coarser-grain sediment. Thus, this promote the coarse-grained (pebblestone and cobblestone) sediments of the fan. Those sedimentological properties also affect the hydrogeological properties and surface-subsurface flow in downdip direction. These flows cause ponding at the end of the fans. Fan deposits serve topographically higher and suitable area for settlements, subsequently the urbanization, manmade channels, and road cover fan deposits. Thus, human interference has now cut the sediment and water input from the apex to the downdip fan parts. At least, human impact

allowed that flashflood can reach to ponding area in short time. However, other effects such as on agriculture, irrigation are unclear. This study suggests constant monitoring of the alluvial sedimentations to generate and balance between the local environment and urban life.

Acknowledgements The authors of this manuscript would like to thank the anonymous reviewers and the editors for their invaluable comments which significantly improved the quality of the paper. The efficient use of FTIR facility in Chemistry Department at Mugla Sitki Kocman University is highly acknowledged. The authors would like to thank Mr. Iliya Bauchi Danladi for English editorial checking of the paper.

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