



Dissocladella compressa n. sp., a new Dasycladale (green algae) from the Upper Maastrichtian of Iran

K. Rashidi¹ · F. Schlagintweit²

Received: 13 November 2018 / Accepted: 4 March 2019 / Published online: 28 March 2019
© Saudi Society for Geosciences 2019

Abstract

The new dasycladalean alga *Dissocladella compressa* is described from the lower part of the Tarbur Formation of SW Iran (Zagros Zone). The medium-sized representative of the genus is characterized by rather large primaries displaying slight vertical compression. It occurs in a foraminiferan-algal packstone associated with larger benthic foraminifera typically *Omphalocyclus macroporus* Lamarck, *Loftusia* sp., *Siderolites* cf. *calcitrapoides* Lamarck, and scattered corals. Associated algae include *Trinocladus tripolitanus* Rainieri, *Thrysoporella longa* Radoičić, *Ovulites delicatula* Elliott, accessory udoteaceans, and corallinaceans.

Keywords Dasycladales · Larger benthic foraminifera · Taxonomy · Systematics

Introduction

The Upper Cretaceous Tarbur Formation, named after the village of Tarbur (Fars Province), and cropping out in the SW Zagros Basin, represents a predominantly carbonate lithostratigraphic unit that contains rich microfauna and microflora associated with rudists (James and Wynd 1965). It extends from the northwest to the southeast of the Zagros Basin along the western edge of the imbricated Zagros zone, between the main Zagros Fault and the Sabzposhan Fault to the east (Alavi 2004). The Tarbur Formation overlies and interfingers (towards the southwest) with the Gurpi Formation.

Microflora of the Tarbur Formation is still rather poorly known and consists almost exclusively of calcareous green algae with a clear dominance of Dasycladales. Some taxa of dasycladalean algae were illustrated in some

recently published papers (Abyat et al. 2012, 2015; Afghah 2016; Afghah and Farhoudi 2012; Afghah and Yaghmour 2014; Dehghani et al. 2016; Khosrow Tehrani and Afghah 2004). These determinations, however, are dubious if not incorrect and therefore require further investigations. First results and critical revisions of the microflora of the Tarbur Formation were published recently (Rashidi et al. 2013; Schlagintweit et al. 2016c; Rashidi and Schlagintweit 2018a, 2018b). The present paper deals with the description of a new dasycladale, *Dissocladella compressa* n. sp. Representatives of the genus *Dissocladella* were so far not reported from the Campanian-Maastrichtian interval (see Barattolo 2002, Table 2).

Studied sections

Almost all specimens of *Dissocladella compressa* n. sp. illustrated here are from the type locality, the Rod-Abad section (Fig. 1). More rarely, the new taxon has also been observed in the nearby Mandegan section as well as the Naghan section.

Rod-Abad section (type locality)

Type level of the new species of *Dissocladella* is located in the Rod-Abad section (Fig. 2). From the same locality

Editorial handling: P. Selden

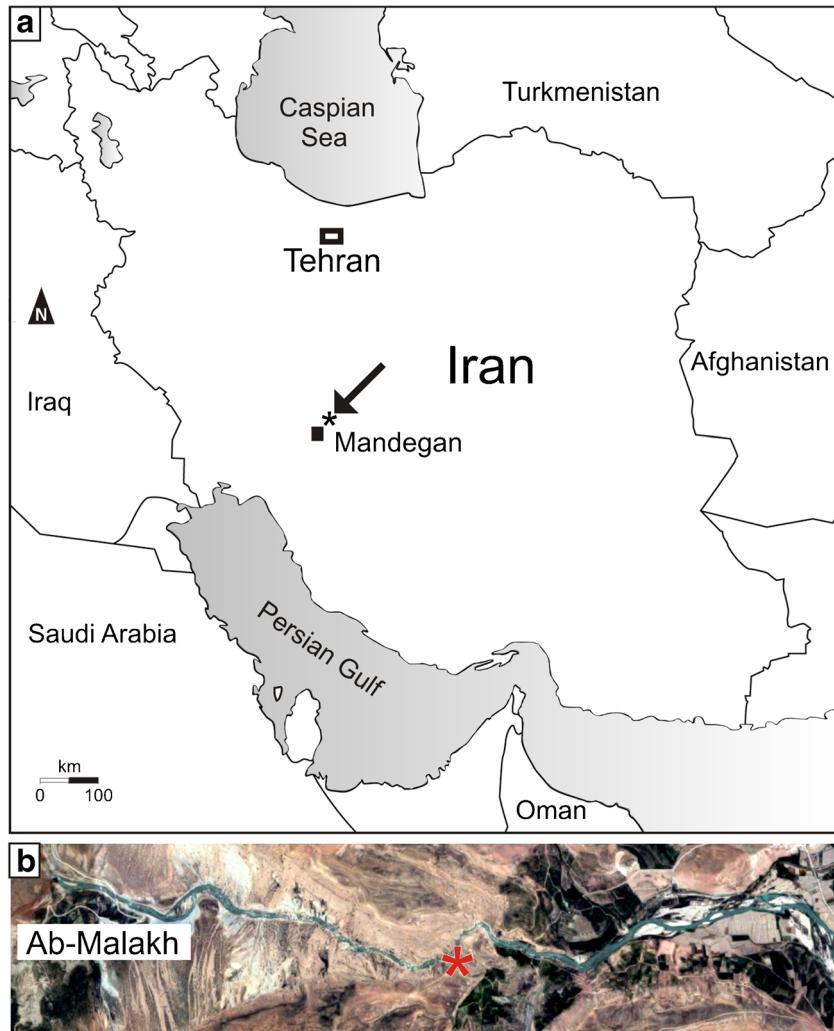
✉ F. Schlagintweit

K. Rashidi
koorosh.rashidi@yazd.ac.ir

¹ Department of Geology, Yazd University, 89195-741, Yazd, Iran

² Munich, Germany

Fig. 1 Location of the Rod-Abad section (star, arrow in **a**; from Consorti and Rashidi 2018, based on Google Maps in **b**)



Elazigina sidereal, a new larger benthic foraminifera (family Rotaliidae) was described recently by Consorti and Rashidi (2018). The section is positioned 6.5 km far to the NE of Mandegan Village, very close to the village of Ab-Malakh (Fig. 1). In the Rod-Abad section, the Tarbur Formation is incompletely exposed (thickness about 70 m) and refers to its basal part. Lithostratigraphically, it represents an equivalent of unit 1 differentiated from the Mandegan section (see below). The lower to middle part of the Rod-Abad section shows four short-term transgressive-regressive pulses expressed by four marly levels that contain rudists in living position. These appear intercalated between massive limestones (Consorti and Rashidi 2018) (Fig. 2). The type level of *Dissocladella compressa* n. sp. is sample TF 3, an algal-foraminiferal packstone (Fig. 3a). The

associated larger benthic foraminifers include *Omphalocyclus macroporus* (Lamarck), *Siderolites* cf. *calcitrapoides* Lamarck, *Loftusia* sp., and *Elazigina siderea* Consorti & Rashidi. Some of the larger benthic foraminifera from the Rod-Abad section are illustrated in Fig. 3b–h. The algal association of TF 3 includes *Dissocladella compressa* n. sp., *Trinocladus tripolitanus* Raineri, *Ovulites delicatula* Elliott, and *Thrysoporella longa* Radović. The Greenwich coordinates of the section base are 31°8'32.46" N and 50°23'51.30" E.

Mandegan section

In the nearby Mandegan section (see Schlagintweit et al. 2016a; Consorti and Rashidi 2018), the Tarbur Formation attains a thickness of almost 280 m from

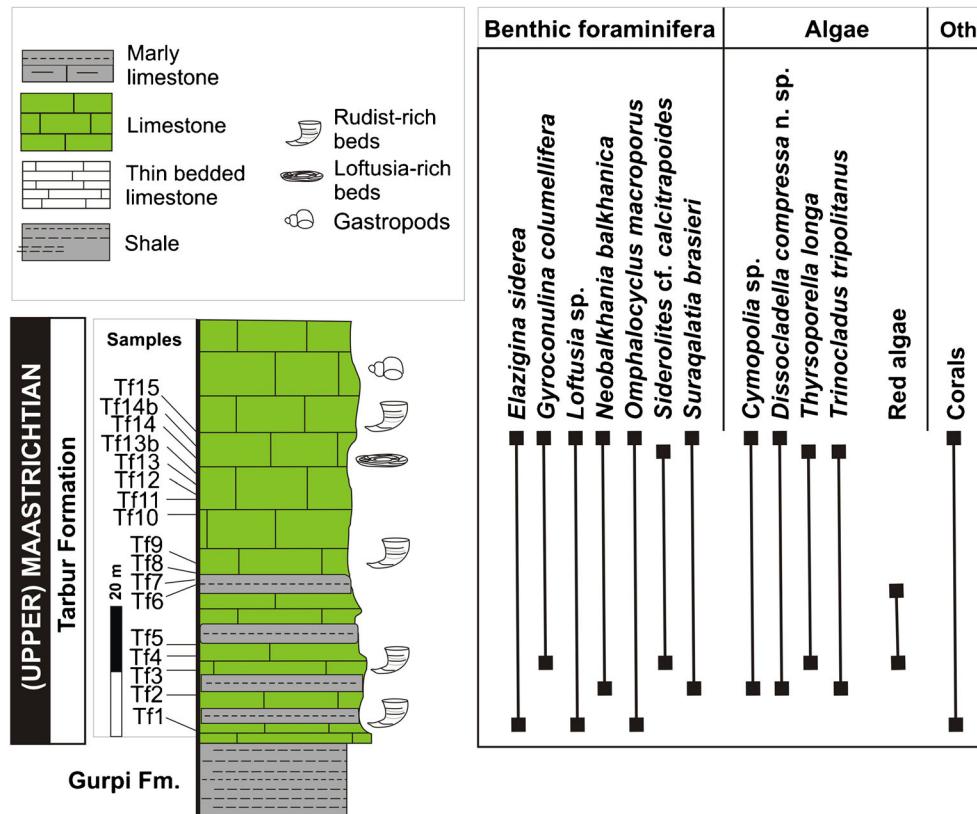


Fig. 2 Vertical distribution of *Dissocladera compressa* n. sp. and selected larger benthic foraminifera in the (late) Maastrichtian of the Tarbur Formation of the Rod-Abad section section

the base (towards the underlying Gurpi Formation) to the top (towards the overlying Pliocene Bakhtiari Formation). Here, *Dissocladella compressa* n. sp. is much less frequent than its type locality. The Greenwich coordinates of the section base are N 31°, 2', 58.13" and E 51°, 24', 34.58".

Naghan section

The Naghan section is located approximately 50 km south west of Naghan town near the Gandomkar Village (see Rashidi and Schlagintweit 2018a). At this locality, the Tarbur Formation is underlain by the Gurpi Formation and overlying by the Paleocene Sachun Formation. The thickness of the Tarbur Formation at the Naghan section is about ~274 m. Here, *Dissocladella compressa* n. sp. is much less frequent than its type locality. The Greenwich coordinates of the section base are 31°47'52" N and 50°32'53" E.

Material and depository

All specimens from *Dissoclarella compressa* n. sp. are from thin sections stored at the Ardakan Payame Noor University, Iran, in the Rashidi collection, under the original sample numbers with the prefixes F, FT, and TF. The numbers correspond to the same sampling levels (F 3 = FT 3) (Fig. 2). From the most prolific sample FT 3, several thin sections have been prepared labeled FT 3-1 to FT 3-11.

Systematic micropalaeontology

Phylum CHLOROPHYTA

Class DASYCLADOPHYCEAE

Oder DASYCLADALES Pascher 1931

Family DASYCLADACEAE Kützing, 1843

Genus *Dissocladella* Pi

ASSOULLET et al., 1978

Figs. 3pars (a), 4(a–i), 5(a–h), and 6(a–i)

Origin of the name The species name refers to the shape of the primaries, being vertically slightly compressed.

Holotype Slightly oblique transverse section illustrated in Fig. 6h on the left side, thin section TF 4-1.

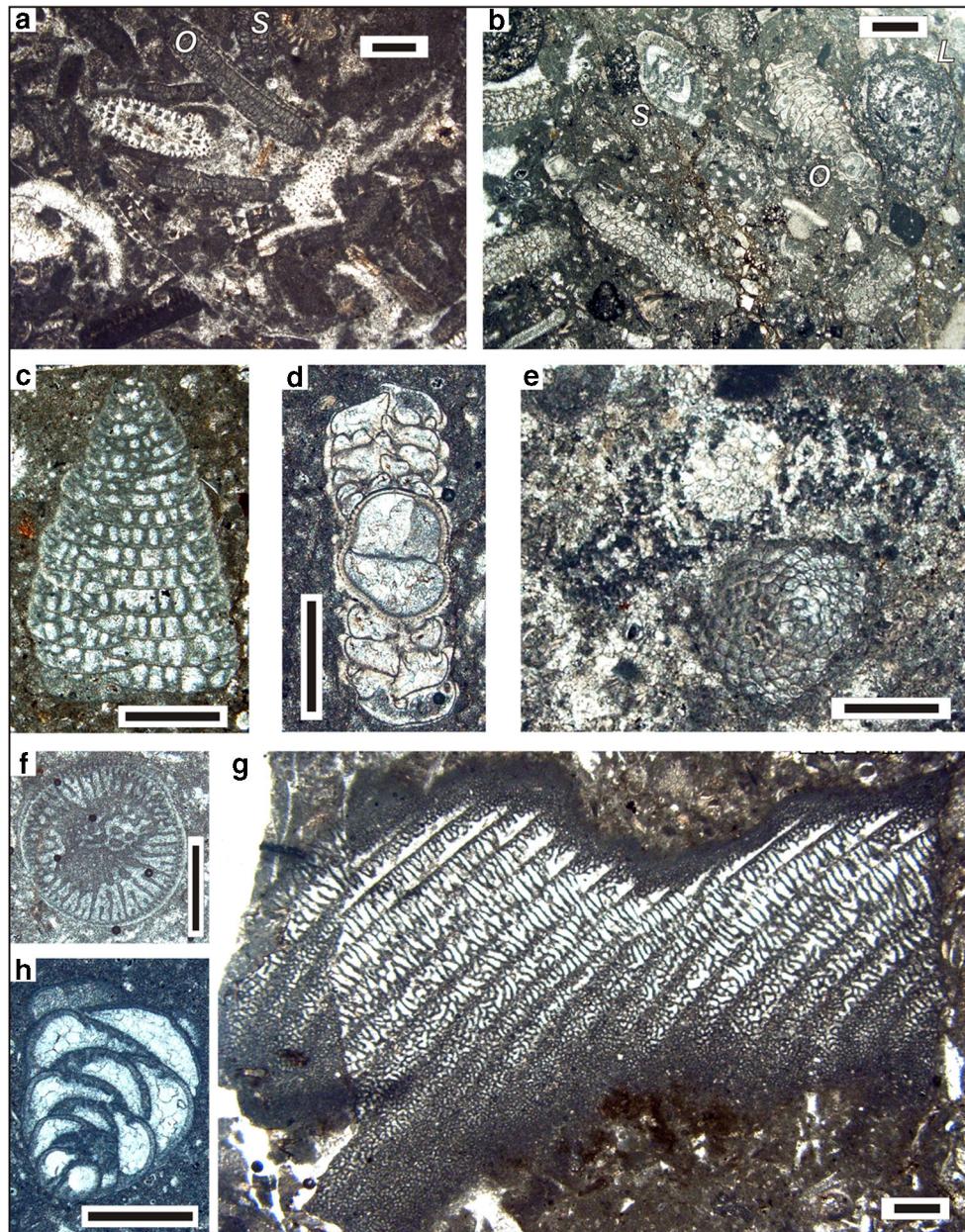
Type locality Rod-Abad section (Figs. 1 and 2).

Type level Mud or grain-supported limestones (wackestones, packstones) with larger benthic foraminifera (e.g., typically association of *Omphalocyclus macroporus*

Lamarck and subordinate *Loftusia* sp.) of the Tarbur Formation.

Description Medium-sized representative of the genus with well calcified cylindrical thallus and main axis on average 42% of the diameter. The inner surface is straight and clearly delineated. Primary laterals numerous, differentiated into a thin proximal stalk, and a main part, subsphaerical to quadrangular-rounded in transverse sections. Due to a slight vertical compression, the primaries are ovoidal in longitudinal

Fig. 3 Microfacies with *Dissocladella compressa* n. sp. (a) and some (larger) benthic foraminifera (b–h) from the (late) Maastrichtian of the Tarbur Formation of the Rod-Abad section, Iran. Foraminiferan packstone with *Loftusia* sp. (L), *Omphalocyclus macroporus* (Lamarck) (L), *Siderolites* cf. *calcitrapoides* (Lamarck) (S). *Dissocladella compressa* n. sp. in b (a, b). *Gyroconulina columellifera* Schroeder & Darmoian, subaxial section (c), and slightly oblique transverse section (f). *Omphalocyclus macroporus* (Lamarck), axial section (d). *Neobalkhania bignoti* Cherchi, Radoičić & Schroeder (above), and *Omphalocyclus macroporus* (Lamarck) (below) (e). *Suragalatia brasieri* Görümüş, Lawa & Nuaimy, tangential section (g). *Minouxia* sp., oblique section (h). Scale bars: 0.6 mm (a, b), 0.5 mm (c–g). Thin sections: F 11 (a), TF 3-11 (b), F 4 (c, d), TF 5-3 (e), FT 11-2 (f), FT 5 (g), and TF 12 (h)



sections. Each primary bears four to five diverging secondaries arising along its distal surface. The latter display a distal swelling thereby touching the neighboring secondaries. In longitudinal sections, the primaries appear slightly shifted between two subsequent verticils, neither clearly alternating nor clearly in line.

Dimensions The main biometric parameters are compiled in Table 1.

Comparisons Among the numerous established species of the genus, several have taxonomically been revised (Ott 1965; Dragastan and Trappe 1986; Barattolo and Bigozzi 1996; Barattolo et al. 2008; Granier 2013), while others are controversially discussed (Raineri 1922; Radoičić et al. 2005; Granier et al. 2017). *Dissocladella compressa* n. sp. represents a medium-sized species clearly different from the large-sized forms, the articulated Paleocene type-

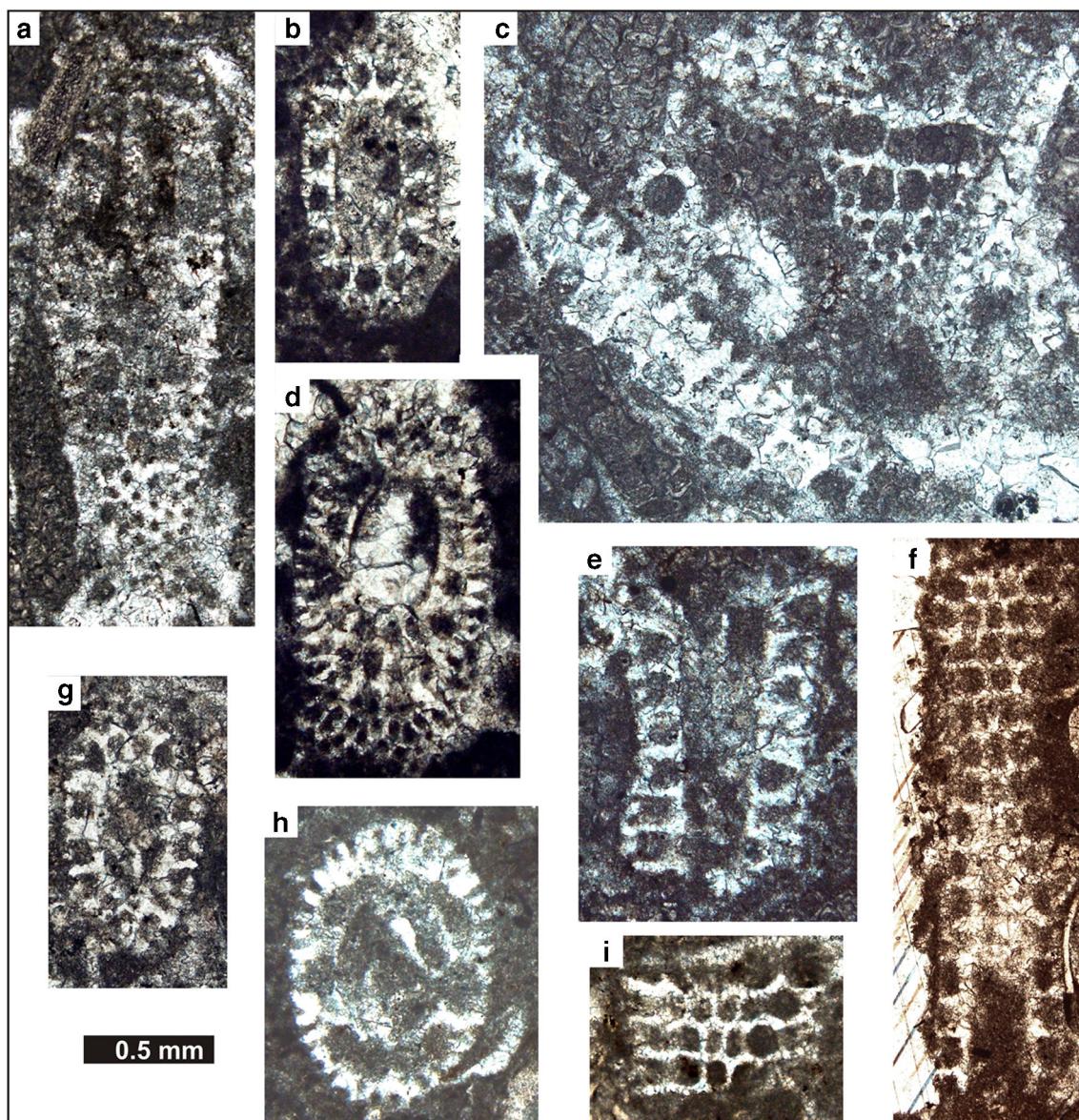


Fig. 4 *Dissocladella compressa* n. sp. from the (late) Maastrichtian of the Tarbur Formation of the Rod-Abad (a–e, i), Mandegan (f), and Naghan sections (h), Iran. Oblique sections (a, b, d, g, h). Fragmentary longitudinal section (left) and tangential section showing primaries and

secondaries (right) (c). Longitudinal section (e). Tangential (upper part) to longitudinal section (lower part) (f). Tangential section (i). Thin sections: TF 3-8 (a), TF 3-11 (b, d, g, i), TF 3 (b), RT 100-2 (f), and NG 113 (h)

species *D. savitriae* Pia, 1936 (in Rao and Pia 1936), or the Lower Cretaceous *D. urgoniana* Dragastan 1989. It should be mentioned that the latter species shows clustered secondaries, meaning arising from a single point at the surface of the primaries (forming a bunch) as in the genus *Montenegrorella* Sokač & Nikler, in Granier and Deloffre 1993 (e.g., Sokač and Nikler 1973; Barattolo 1984; Sokač and Grgasović 2015; Rashidi and Schlagintweit 2018b). Therefore, *D. urgoniana* needs taxonomic revision in our opinion. The widespread Upper Cretaceous *Dissocladiella undulata* Raineri is distinctly smaller than *D. compressa*, also with a small number of primaries per verticil (Raineri 1922; Radoičić et al. 2005;

Schlagintweit and Wilmsen 2014). The same characteristics differentiate the new Iranian species from other Paleogene forms such as *D. lunata* Segonzac, *D. deserta* Elliott, or *D. gracilis* Radoičić (see Table 1 in Radoičić 1991). Last but not least, the vertical compression of the primary laterals of *D. compressa* represents a unique position feature that is here considered a specific criterion.

Recently, another dasycladale also displaying two orders of laterals was described by Rashidi and Schlagintweit (2018b) as *Supgilulummaella tarburensis* from the Tarbur Formation. This larger taxon (D up to 1.68 mm) shows numerous laterals ($w = 14\text{--}20$) inclined to the main axis typical of the genus.

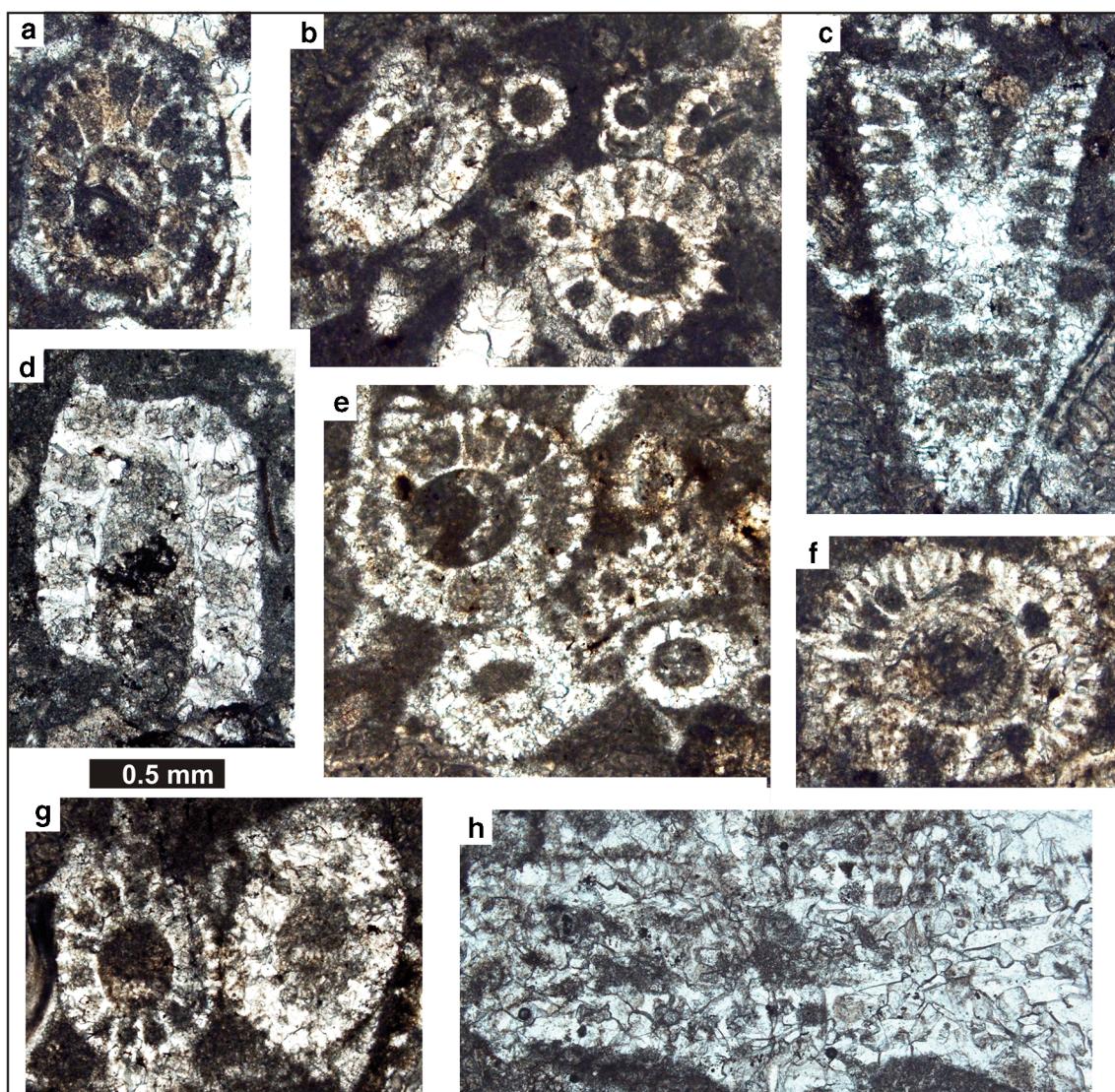


Fig. 5 *Dissocladiella compressa* n. sp. from the (late) Maastrichtian of the Tarbur Formation of the Rod-Abad section, Iran. Oblique sections (a, c, d). Oblique and slightly oblique transverse sections (b). Oblique

transverse sections (e–g). Longitudinal section (h). Thin sections: TF 3–8 (a, c), TF 3–11 (b, e–g), TF 3–10 (d), TF 3 (h)

Remarks In the Rod-Abad section *Dissocladella compressa* may be associated with other alga (Fig. 7). Besides accessory corallinaceans, these include dasycladaleans and not further determined udoteacean al-

gae. Among the dasycladales, we observed more frequently an association of *Dissocladella compressa* with *Trinocladus tripolitanus* Raineri (Fig. 7a, b, d), *Thrysoporella longa* Radoičić (Fig. 7g–l), *Cymopolia*

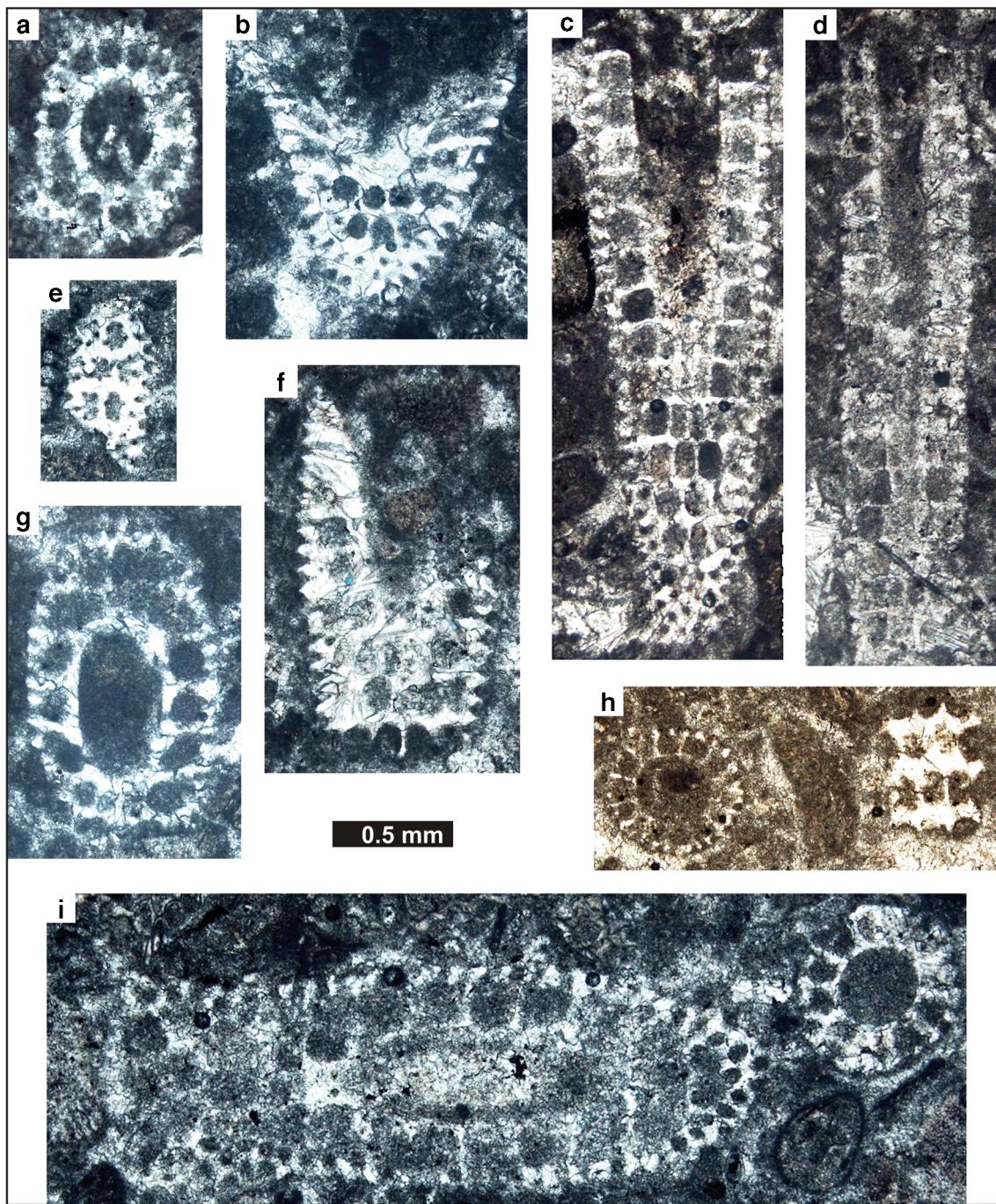


Fig. 6 *Dissocladella compressa* n. sp. from the (late) Maastrichtian of the Tarbur Formation of the Rod-Abad section, Iran. Oblique sections (a, b, g). Longitudinal-tangential sections (c–d, f). Fragmentary tangential section (e). Transverse section left (holotype specimen) and tangential

section (right) (h). Oblique section (left) and slightly oblique transverse section (right) (i). Thin sections: TF 5-4 (a, b, f, g), TF 5-5 (c, e, i), and TF 4-1 (h)

Table 1 Dimensions (in mm; except d/D) of *Dissoclarella compressa* n. sp. from the late Maastrichtian of the Tarbur Formation of Iran (in mm, except ratio d/D). D outer diameter, d inner diameter, pmax maximum

| Sample | D | d | d/D | p | l | p' | l' | h | w |
|----------|-----------|-----------|-----------|-----------|-----------|------------|-----------|-----------|--------|
| NG 113 | 0.75 | 0.37 | 0.5 | 0.18 | 0.19 | 0.08 | 0.08 | — | — |
| Rt 100-3 | — | — | — | 0.16 | 0.16 | — | — | 0.16 | — |
| F 4 | 0.35 | 0.15 | 0.43 | 0.07 | 0.06 | 0.045 | 0.04 | — | 14 |
| F 4 | — | — | — | 0.13 | — | — | — | 0.16 | — |
| FT 11 | 0.48 | 0.26 | 0.54 | 0.09 | 0.08 | 0.06 | — | — | 16 |
| TF 3 | 0.79 | 0.35 | 0.44 | 0.11 | 0.16 | 0.07 | 0.14 | 0.15 | — |
| TF 3 | 0.84 | 0.31 | 0.37 | 0.12 | 0.17 | — | — | 0.16 | — |
| TF 3 | 0.72 | 0.33 | 0.47 | 0.08 | 0.12 | 0.06 | 0.09 | — | 15 |
| TF 3 | 0.55 | 0.25 | 0.45 | 0.1 | 0.13 | — | — | — | — |
| TF 4-1 | 0.47 | 0.21 | 0.45 | 0.08 | 0.1 | — | — | — | 14 |
| TF 4-1 | 0.5 | 0.25 | 0.5 | 0.09 | 0.09 | 0.06 | 0.06 | — | 14 |
| TF 3-8 | — | — | — | 0.075 | 0.13 | 0.06 | 0.16 | — | — |
| TF 3-8 | 0.55 | 0.23 | 0.41 | — | — | — | — | — | 17 |
| TF 3-8 | 0.7 | 0.31 | 0.44 | 0.19 | 0.14 | 0.07 | 0.08 | — | — |
| TF 3-8 | — | — | — | 0.14 | 0.18 | 0.08 | 0.11 | 0.17 | — |
| TF 3-10 | 0.82 | 0.33 | 0.4 | 0.16 | 0.19 | 0.06 | — | 0.18 | — |
| TF 3-11 | 0.87 | 0.41 | 0.47 | 0.13 | 0.17 | 0.05 | 0.1 | — | 19 |
| TF 3-11 | 0.77 | 0.31 | 0.4 | 0.1 | 0.17 | — | — | — | 217–18 |
| TF 3-11 | 0.61 | 0.27 | 0.44 | 0.09 | 0.13 | 0.05 | 0.08 | — | 15 |
| TF 3-11 | 0.63 | 0.24 | 0.38 | — | — | — | — | — | — |
| TF 3-11 | 0.42 | 0.16 | 0.38 | — | — | — | — | — | — |
| TF 3-11 | 0.84 | 0.37 | 0.44 | 0.16 | 0.17 | 0.07 | 0.08 | — | 17 |
| TF 3-11 | 0.53 | 0.2 | 0.38 | 0.1 | 0.1 | — | — | — | — |
| TF 3-11 | 0.76 | 0.31 | 0.42 | 0.11 | 0.13 | 0.08 | 0.12 | — | — |
| TF 3-11 | 0.55 | 0.23 | 0.42 | 0.1 | 0.12 | 0.07 | 0.07 | — | — |
| TF 4-1 | 0.55 | 0.28 | 0.51 | 0.14 | 0.14 | 0.08 | 0.05 | — | 14 |
| TF 4-11 | — | — | — | — | — | — | — | 0.17 | — |
| TF 5-4 | 0.8 | 0.33 | 0.42 | 0.17 | 0.19 | 0.08 | 0.08 | — | — |
| TF 5-4 | 0.63 | 0.27 | 0.42 | 0.1 | 0.12 | 0.08 | 0.08 | — | — |
| TF 5-4 | — | — | — | 0.13 | — | 0.09 | 0.14 | 0.18 | — |
| Tf 5-5 | 0.78 | 0.35 | 0.45 | 0.16 | 0.14 | 0.09 | 0.11 | 0.17 | — |
| FT 3-11 | — | — | — | 0.16 | — | — | — | 0.16 | — |
| Range | 0.35–0.87 | 0.15–0.41 | 0.37–0.54 | 0.07–0.19 | 0.06–0.19 | 0.045–0.09 | 0.04–0.14 | 0.15–0.18 | 14–19 |
| medium | 0.65 | 0.28 | 0.42 | 0.12 | 0.14 | 0.07 | 0.09 | 0.166 | 17 |

sp., and more rarely *Cymopolia tibetica* Morellet & Morellet (Fig. 7c, f). A similar association of *Trinocladus tripolitanus* Raineri with representatives of *Dissoclarella* (here: *D. undulata*) and udoteaceans is reported from the Upper Cretaceous of Lybia (Raineri 1922; Radović et al. 2005) or the Cenomanian of Spain (Schlagintweit and Wilmsen 2014). As in the Tarbur Formation, the microfacies of this algal association correspond to external platform environments with medium to high energy levels. The occurrence of *Trinocladus tripolitanus* in the late Maastrichtian Tarbur Formation

diameter of primary laterals, l length of primary laterals, p'max maximum diameter of secondary laterals, l'max maximum length of secondary laterals, h vertical spacing, w number of primary laterals in a vertical

represents the youngest record of the species (see Barattolo 2002, Table 2). *Thyrsoporella longa* was described by Radović (1990) from the early Paleocene of Iraq and is here recorded for the first time from the late Maastrichtian (see Barattolo 2002).

Stratigraphy

The occurrence of *Siderolites cf. calcitrapoides* Lamarck basal part of the Rod-Abad section is worth mentioning, as this

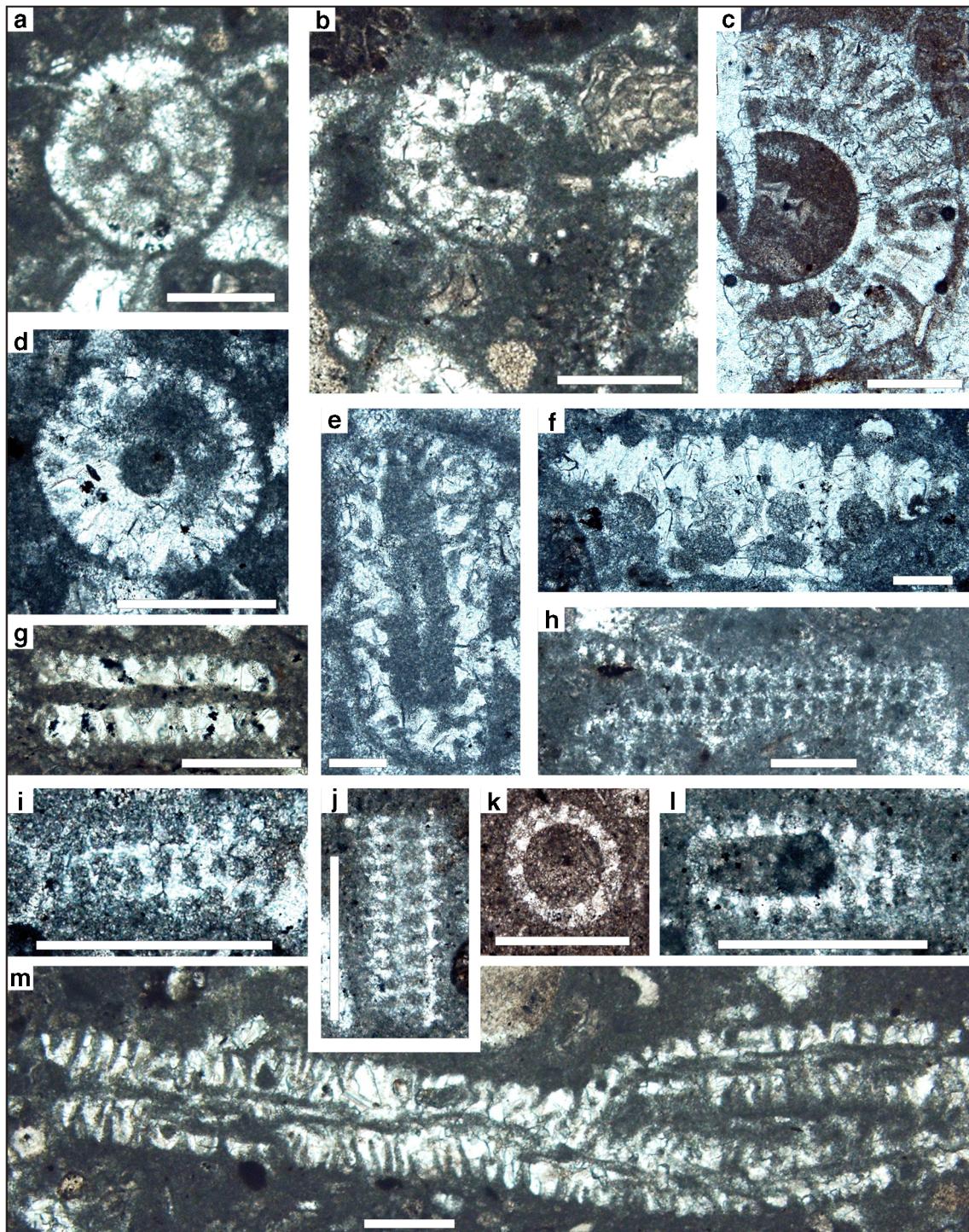


Fig. 7 Calcareous green algae from the (late) Maastrichtian of the Tarbur Formation of the Rod-Abad section, Iran. *Trinocladus tripolitanus* Raineri, transverse sections (a, d), and oblique section (b). *Cymopolia tibetica* Morellet & Morellet, transverse section (c), and fragment (f). Udoteacean alga, oblique section (e). *Thrysoporella longa* Radoičić,

longitudinal section (g), diverse tangential sections (h–j), and slightly oblique transverse section (k). Udoteacean alga showing low-angle branching on the right side (m). Scale bars: 0.3 mm except (k) = 0.2 mm. Thin sections: F 10 (a, b, m), TF 3-1 (c), TF 3-4 (d, f), TF 3-6 (e), F 11 (g), TF 8-1 (h), FT 3 (i, j, l), and TF 14 (k)

taxon has its possible first appearance in the latest early Maastrichtian, or is completely restricted to the upper Maastrichtian (according to Robles-Salcedo et al. 2018).

Other typical Maastrichtian taxa include *Loftusia*, *Gyroconulina*, and *Omphalocyclus* (e.g., Schlagintweit et al. 2016a).

Table 2 Dasycladalea inventory of the Tarbur Formation of the Rod-Abad, Mandegan, and Naghan sections

| Taxa section | Rod-Abad | Mandegan | Naghan |
|--|----------|----------|--------|
| <i>Clypeina</i> sp. | | X | X |
| <i>Cymopolia eochoristosporica</i> Elliott | X | — | X |
| <i>Cymopolia tibetica</i> Morellet & Morellet | X | — | — |
| <i>Dissocladella compressa</i> Rashidi & Schlagintweit | X | X | X |
| <i>Pseudocymopolia anadyomenea</i> (Elliott) | — | X | X |
| <i>Salpingoporella pasmanica</i> Radočić | — | X | X |
| <i>Suppilulumaea tarburensis</i> Rashidi & Schlagintweit | — | X | X |
| <i>Thyrsoporella longa</i> Radočić | X | — | — |
| <i>Trinocladus tripolitanus</i> Raineri | X | X | — |
| <i>Zittelina? arumaensis</i> (Okla) | | X | X |

Conclusive remarks

The description of the new dasycladale *Dissocladella compressa* contributes to the improving knowledge on the microflora of the Tarbur Formation of Iran. It is the first record of the genus from the Maastrichtian. The total range of *Dissocladella* is Upper Triassic (Sokač and Grgasović 1995) to Lutetian (Barattolo 2002). Ongoing studies evidence a greater diversity of larger benthic foraminifera than dasycladaleans. Both may show typical associations. In the case of *Dissocladella compressa*, these are *Omphalocyclus* and *Loftusia* that in the Iranian literature classically define a biozone characteristic for the lower part of the Tarbur Formation (James and Wynd 1965). The late Maastrichtian Rod-Abad section represents another locality lacking Campanian parts of the Tarbur Formation that are so far not proven by hard facts (see Discussion in Schlagintweit et al. (2016a, 2016b, 2016c)).

Acknowledgments The helpful comments provided by two anonymous reviewers are kindly acknowledged.

References

- Abyat A, Afghah M, Feghhi A (2012) Stratigraphy and foraminiferal biozonation of Upper Cretaceous sediments in southwest Sepid Dash, Lurestan, Iran. Austral J Basic App Sci 6(13):18–26
- Abyat A, Afghah M, Feghhi A (2015) Biostratigraphy and lithostratigraphy of Tarbur Formation (Upper Cretaceous) in southwest of Khorram Abad (southwest Iran). Carbonates Evaporites 30(1): 109–118
- Afghah M (2016) Biostratigraphy, facies analysis of Upper Cretaceous – Lower Paleocene strata in South Zagros Basin (southwestern Iran). J Afr Earth Sci 119:171–184
- Afghah M, Farhoudi G (2012) Boundary between Upper Cretaceous and Lower Paleocene in the Zagros Mountain ranges of southwestern Iran. Acta Geol Sin 86(2):325–338
- Afghah M, Yaghmour S (2014) Biostratigraphy study of Tarbur Formation (Upper Cretaceous) in Tang-E Kushk and east of Sarvestan (SW of Iran). J Earth Sci 25(2):263–274
- Alavi M (2004) Regional stratigraphy of the Zagros foldthrust belt of Iran and its proforeland evolution. Amer J Sci 304:1–20
- Barattolo F (1984) Osservazioni su *Suppilulumaea schroederi* n.sp. (alghe Verdi, Dasicladali) del Cenomaniano del Matese (Appennino centrale, Italia). Boll Soc Nat Napoli 92(1983):1–47
- Barattolo F (2002) Late Cretaceous-Paleogene Dasycladaleans and the K/T boundary problem. In: Bucur, I.I., Filipescu, S. (eds.), Research advances in calcareous algae and microbial carbonates, proceedings of the 4th IFAA regional meeting, Cluj Napoca, Romania, 29 August–5 September 2001, pp. 17–40
- Barattolo F, Bigozzi A (1996) Dasycladaleans and depositional environments of the upper Triassic-Liassic carbonate platform of the gran Sasso (central Apennines, Italy). Facies 35:39–54
- Barattolo F, Granier B, Romano R, Ferré B (2008) *Petrascula iberica* (Dragastan & Trappe), *Tersella genotii* Barattolo & Bigozzi, and the relationships of club-shaped dasycladalean algae during Late Triassic-Early Jurassic times. Geol Croatica 61(2–3):159–176
- Consorti L, Rashidi K (2018) A new evidence of passing the Maastrichtian–Paleocene boundary by larger benthic foraminifers: the case of *Elazigina* from the Maastrichtian Tarbur Formation of Iran. Acta Pal Polon 63(3):595–605
- Dehghani M, Vaziri SH, Ahmadi V (2016) Lithostratigraphy and microbiostatigraphy of the Tarbur Formation in the Zangeneh section of interior Fars (folded Zagros), south of Iran. Int J Human Cult Stud 2(4):1872–1884
- Dragastan O (1989) Calcareous algae (new and revised), microproblematicae and Foraminiferida of Jurassic–Lower Cretaceous deposits from the Carpathian area. Rev Esp Micropal 21(1):5–65
- Dragastan O, Trappe J (1986) Some dasyclad algae of the Sinemurian from the north-western Iberian chains (Spain). Paläont Z 60(3–4): 169–179
- Granier B (2013) *Dissocladella haueriviana* Masse in masse et al., 1999 (non masse, 1976), another lower Urgonian Dasycladalean alga revisited. Carnets Géol (Notebooks Geol) Lett 2013/07 (CG2013_L07), 347–355
- Granier B, Deloffre R (1993) Inventaire critique des Algues Dasycladales fossiles. II^e partie – Les Algues Dasycladalees du Jurassique et du Crétacé. Rev Paléobiol 12(1):19–65
- Granier B, Bucur II, Dias-Brito D (2017) About *Trinocladus* Raineri, 1922: when some *Permocalculus* (Gymnocodiacean algae) reveal to be Triploporellacean algae (Revision of the Jesse Harlan Johnson Collection. Part 5). Facies 63:27
- James GA, Wynd JG (1965) Stratigraphic nomenclature of Iranian oil consortium agreement area. AAPG Bull 49(12):2218–2232

- Khosrow Tehrani K, Afghah M (2004) Microbiostratigraphy and microfacies study of the Tarbur Formation in northeast and southeast of Shiraz. *Geosci* 12(53):74–87
- Ott E (1965) *Dissocladella cretica*, eine neue Kalkalge (Dasycladaceae) aus dem Mesozoikum der griechischen Inselwelt und ihre phylogenetischen Beziehungen. *N Jb Geol Pal Mh* 11:683–693
- Radoičić R (1990) Paleogene Dasycladalean algae from the subsurface of the Western Iraqi Desert. *Bull Acad Serbe Sci Arts Cl Sci Math Nat* 32:91–103
- Radoičić R (1991) *Dissocladella gracilis* n. sp. (Dasycladales) from the Thanetian of Majevica Mountain (NE Bosnia, Dinarides). *Rev Paléobiol* 10(2):309–313
- Radoičić R, Conrad MA, Carras N (2005) Observations on *Neomeris cretacea* var. *undulata* Rainieri followed by *Dissocladella bonardii*, n. sp. (Dasycladales, green algae). *Rev Paléobiol* 24(1):311–317
- Raineri R (1922) Algue sifonee fossili della Libia. *Atti Soc Ital Sci nat Mus civ stor nat Milano* 61(1):72–86
- Rao LR, Pia J (1936) Fossil algae from the uppermost Cretaceous beds (the Niniyur Group) of the Trichinopoly district, S. India. *Mem Geol Surv India, Pal Indica, NS* 21(4):1–49
- Rashidi K, Schlagintweit F (2018a) *Cymopolia eochoristosporica* Elliott, 1968 (green alga, Dasycladale) from the upper Maastrichtian of the Tarbur Formation (SW Iran). *Cret Res* 82:99–103
- Rashidi K, Schlagintweit F (2018b) *Zittelina? arumaensis* (Obla, 1995) nov. comb. and *Suppliliumaelia tarburensis* n. sp. (Dasycladales) from the Upper Maastrichtian of Iran. *Arab J Geosci* 11:418
- Rashidi K, Schlagintweit F, Parvizi T, Bamdad L, Saberzadeh B (2013) Late Cretaceous (Maastrichtian) dasycladalean algae from the Naghan area (Zagros Mountains, SW Iran): preliminary results. *Ber Geol BA* 99 (11th Workshop Alpine Geol Stud & 7th IFAA), pp. 114–115
- Robles-Salcedo R, Vicedo V, Caus E (2018) Latest Campanian and Maastrichtian Siderolitidae (larger benthic foraminifera) from the Pyrenees (S France and NE Spain). *Cret Res* 81:64–85
- Schlagintweit F, Wilmsen M (2014) Calcareous algae (Dasycladales, Udoteaceae) from the Cenomanian Altamira Formation of northern Cantabria, Spain. *Acta Pal Romaniae* 10(1–2):15–24
- Schlagintweit F, Rashidi K, Barani F (2016a) First record of *Gyroconulina columellifera* Schroeder & Darmoian, 1977 (larger benthic foraminifera) from the Maastrichtian Tarbur Formation of SW Iran (Zagros Fold-Thrust-Belt). *GeoPersia* 6(2):169–185
- Schlagintweit F, Rashidi K, Babapidour M (2016b) Orbitolinid foraminifera from the Late Maastrichtian of the Tarbur Formation (Zagros Zone, SW Iran). *Acta Pal Romaniae* 12(2):29–46
- Schlagintweit F, Rashidi K, Kivani F (2016c) On the occurrence of *Salpingoporella pasmanica* Radoičić, 2002, (Dasycladales) from the Late Maastrichtian of the Zagros Zone, SW Iran. *Acta Palaeont Romaniae* 12(1):33–42
- Sokač B, Grgasović T (1995) *Dissocladella bystrickyi* n. sp., a new calcareous alga (Dasycladaceae) from Upper Triassic Dolomites of Mt. Medvednica (Northern Croatia). *Geol Croatica* 48(1):1–7
- Sokač B, Grgasović T (2015) *Montenegrella? gracilis* n. sp., a new calcareous alga (Dasycladales) from the Upper Barremian of Mt. Biokovo (Dinarides Mts., Croatia). *Geol Croatica* 68(3):173–178
- Sokač B, Nikler L (1973) Calcareous algae from the Lower Cretaceous of the environs of Nikšić, Crna Gora (Montenegro). *Paleont Jugoslav* 13:7–57