Effect of Crab Burrows on $CO₂$ Flux from the Sediment Surface to the Atmosphere in a Subtropical Mangrove Forest on Ishigaki Island, Southwestern Japan

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Abstract

To clarify the effects of crab burrows on variation in sediment $CO₂$ flux in mangrove forest, we measured the traits of crab burrows (density and entrance area size) and the $CO₂$ flux rate from sediment surfaces, in areas with and without burrows, in a subtropical mangrove forest on Ishigaki Island, southwestern Japan. Burrow density and entrance area showed significant differences among seasons (warm, middle, and cool) and mangrove zones (upper-, middle-, and downstream), which may have depended on crab phenology, life cycle, and species composition. The sediment $CO₂$ flux rate was significantly higher at plots with crab burrows $(B+)$ than at those without burrows $(B-)$ in each zone and season. However, standardized sediment CO₂ flux rate by burrow surface area at B+ plots did not differ significantly from that at B− plots. In addition, there were no significant differences in sediment temperature and sediment water content between the two types of plots. Moreover, the level of microbial respiration differed significantly between sediments collected from the deep part and those collected from either the ground surface part or burrow walls. These results suggest that crab burrows increase sediment $CO₂$ flux from the mangrove forest floor by increasing the sediment–atmosphere interface area, thereby inducing a change to aerobic conditions in the sediments around burrows. Therefore, the seasonal and spatial effect of crab burrows on the forest floor should be considered when evaluating sediment $CO₂$ flux and examining the role of the mangrove ecosystem as a carbon sink.

Keywords Coastal ecosystem . Crab burrows . Heterotrophic respiration . Seasonality . Soil respiration

Introduction

Mangrove forests are more effective sinks of atmospheric carbon dioxide $(CO₂)$ than other terrestrial ecosystems because of their high net primary production (NPP) and low heterotrophic

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respiration (HR) (Alongi [2014;](#page-7-0) Barr et al. [2010](#page-7-0); Poungparn and Komiyama [2013\)](#page-8-0). The warm climate in the tropical zone enables the mangroves to fixate atmospheric carbon throughout a year, which could be one of the reasons for high NPP (Komiyama et al. [2008](#page-7-0)). In addition, mangrove sediments are regularly submerged because of tidal fluctuation and mangroves are salt tolerant (Robertson et al. [1991](#page-8-0); Ross et al. [2001\)](#page-8-0). Therefore, water restriction need not be considered as a factor in mangrove forests, and regular submergence might assist high aboveground NPP (Sherman et al. [2003](#page-8-0)). Conversely, the regular submergence may also lead to anaerobic conditions in the pedosphere, resulting in low heterotrophic respiration (Alongi [2014;](#page-7-0) Mall et al. [1991\)](#page-7-0). Generally, $CO₂$ flux from soil surfaces (soil respiration) includes heterotrophic and root respiration in terrestrial ecosystems (Hanson et al. [2000;](#page-7-0) Tomotsune et al. [2013a\)](#page-8-0). However, the $CO₂$ flux from sediment surfaces in the mangrove forest can be considered to equal HR because most metabolic respiration from underground roots (root respiration) is released to the atmosphere through the lenticels on pneumatophores, prop roots, and buttresses above the ground (Scholander et al. [1955](#page-8-0); Tomlinson [1986\)](#page-8-0). Some studies have reported that sediment CO₂ flux ranged from 0.47 to 1.45 µmol CO₂ m⁻² s⁻¹ in tropical and subtropical mangrove forests and was lower than that of terrestrial tropical forests (Poungparn et al. [2009](#page-8-0); Tomotsune et al. 2018). However, our understanding of $CO₂$ fluxes from the mangrove forest floor remains limited compared with the aboveground plant fixation and respiration, and therefore large uncertainties remain in evaluating the ecosystem function of mangrove forests as an atmospheric carbon sink.

Some of the uncertainties can be attributed to the presence of macrobenthos. Macrobenthos are a major biological community of coastal ecosystems, such as sandy and muddy tidal flats (Jordan and Valiela [1982](#page-7-0); Otani et al. [2010](#page-8-0)). They have often been called bioturbators or ecosystem engineers, and they greatly affect the mineralization of organic carbon in sediments (Jones et al. [1994\)](#page-7-0). For example, they directly affect mineralization by feeding the organic matter and also indirectly affect it because the pelleted residue (excretion) is decomposed more easily than the original organic matter (Lee [1997](#page-7-0); Nielsen et al. [2003](#page-7-0)). In addition, their burrowing activities change the physiological, chemical, and biological environments of sediments and then indirectly affect the mineralization process around crustacean burrows (Kristensen et al. [2008](#page-7-0); Laverock et al. [2010\)](#page-7-0). Although these effects would have large impact on $CO₂$ flux in mangrove forests, our knowledge about it, especially the indirect effect, is limited. In mangrove forests, fiddler crabs (Uca spp.) are often present at high density on the sediment surface (Warren [1990\)](#page-8-0) and make cylindrical burrows penetrating to a depth of \sim 20–30 cm during low tides (Dye and Lasiak [1987](#page-7-0)). In addition, the mangrove forest accumulates a large amount of organic carbon in its sediments, which could be the cause of high sediment $CO₂$ flux rate in the pedosphere (Alongi et al. [2012;](#page-7-0) Breithaupt et al. [2012](#page-7-0)). Thus, it can be expected that the presence of burrows would contribute more strongly to variation in the sediment $CO₂$ flux rate in mangrove forests than in other tidal flats.

The sediment $CO₂$ flux rate in mangrove forests is known to be affected by sediment temperature (ST), which implies moderate seasonal variations in the sediment $CO₂$ flux (Chanda et al. [2014;](#page-7-0) Jin et al. [2013\)](#page-7-0). In contrast, the effect of sediment water content (SWC) on sediment $CO₂$ flux is relatively small because the SWC in mangrove forests remains constant and high through the year (Lawton et al. [1981](#page-7-0)). Crustacean burrows might alter the ST and SWC of the sediments around burrows, however, causing changes in the sediment $CO₂$ flux rate. In addition, burrows also increase the sediment–atmosphere interface area and also affect the microbial activity in the sediments around burrows, thus directly increasing the sediment $CO₂$ flux rate (Kinoshita et al. [2003](#page-7-0); Laverock et al. [2010\)](#page-7-0). Moreover, burrow density and entrance area vary seasonally and spatially, reflecting changes in species composition of the macrobenthos community and their phenology (Ishigami et al. [2005;](#page-7-0) Kosuge and Kohno [2010\)](#page-7-0), which also can be expected to affect the annual sediment CO₂ flux at the local scale.

We hypothesized that crab burrows significantly affect variation of the sediment $CO₂$ flux rate through the changing physical (ST, SWC, and interface area) and biological (microbial respiration activity) factors of the sediments. To test this hypothesis, we surveyed the density and entrance area of crab burrows and measured the sediment $CO₂$ flux rate and environmental factors (i.e., ST and SWC) in plots with and without burrows, during three seasons, in three areas along a mangrove river, in Ishigaki Island, southwestern Japan.

Methods

Study Site

The study site was located near the Fukido River on Ishigaki Island, Okinawa Prefecture, southwestern Japan (24°29′ N, 124°13′ E; Fig. [1\)](#page-2-0). The subtropical climate had an annual mean temperature of 24.3 °C and mean precipitation of 2,107 mm between 1981 and 2010 [Ishigakijima Local Meteorological Observatory, Japan Meteorological Agency (ILMO)]. The study area is dominated by two mangrove species: Bruguiera gymnorrhiza and Rhizophora mucronata; R. mucronata is more dominant closer to the river. Tree biomass in the permanent quadrat located near the middle-stream zone was 30 Mg C ha⁻¹ in 2012 (R. mucronata, 25%; B. gymnorrhiza, 75%) (Tomotsune et al. [2018](#page-8-0)). Some respiratory mangrove roots protrude from the sediment surface, and there is no understory on the forest floor. The sediment texture is sandy loam to sandy clay loam (clay content, 10%– 17%). More detailed information about this site is provided in Kinjo et al. ([2005](#page-7-0)) and Kida et al. ([2017](#page-7-0)).

Several taxa of macrobenthos occupy the study site, such as crabs (Uca spp. and Scopimera spp.), mud shrimp (Upogebia spp.), and mudskippers (Periophthalmus spp.) (Kosuge [2013](#page-7-0)). Almost burrows are ones constructed by filterfeeding crabs of Uca spp. and Scopimera spp. Uca spp. inhabited all zones, whereas Scopimera spp. inhabited only the downstream zone in our observations. Although these crabs sometimes take over other burrows, they generally hold one burrow per crab (Ishigami et al. [2005\)](#page-7-0). Their feeding behavior was often observed on the sediment surface, and they hid in their burrows to protect themselves from threats.

Crab Burrow Density and Entrance Area

The line transect method was used to clarify the relationship between sediment $CO₂$ flux rate and burrow density and Fig. 1 Study site located near the Fukido River on Ishigaki Island, Okinawa Prefecture, southwestern Japan (24°29′N, 124°13′E). Light gray and dark gray zones indicate the mangrove forest and river, respectively. Black bars indicate the transect line along which the sediment CO2 flux measurement and survey of the crab burrow density and entrance area were conducted, in the upper-, middle-, and downstream zones (US, MS, and DS, respectively)

entrance area. The burrow density and entrance area could change along the stream zone (Otani et al. [1997](#page-8-0)). Therefore, 200 m transect lines were established in the upper- and downstream (US and DS) zones (Fig. 1). In the middlestream (MS) zone, the line was 125 m in length owing to geographical restrictions. A small quadrat $(50 \times 50 \text{ cm}; \text{ Fig. 2a})$ was established every 5 m along the transect line $(n = 25$ for the 125 m line, and $n = 40$ for the 200 m line), and photographs were taken to describe the burrow density (B_{den}, m^{-2}) and entrance area $(B_{area}, cm²)$, based on image processing using ImageJ (a public-domain Java image-processing program; <http://imagej.nih.gov/ij/>).

Sediment $CO₂$ Flux Rate and Environmental Factors

Sediment $CO₂$ flux measurements were conducted in plots with and without crab burrows (B+ and B− plots, respectively), in every small quadrat (Fig. 2a) and in three seasons (warm, August 2016; middle, November 2016; and cool, January 2017). The closed-chamber method using a portable infrared gas analyzer (IRGA, GMP343; VAISALA, Finland) and vinyl chloride chambers (area, 78.5 cm²; height, 11 cm) (Tomotsune et al. 2017) was used to measure sediment $CO₂$ flux rate (sediment respiration) from the air-exposed sediment surface (not covered with seawater) in the B+ and B− plots $(SR_{B+}$ and SR_B −, respectively, mg CO_2 m^{−2} h^{−1}). One burrow was included in each B+plot, and the respiratory roots of mangroves were not included in it. The measure of SR_{B+} was expected to include direct respiration of crabs inhibiting these burrows. Next, after these measurements, the sediment around a burrow was dug up to confirm the presence of a crab inhibiting the burrow. However, no crabs were observed in any of the plots, which was probably because the crabs had escaped to another burrow when we approached to conduct the measurements.

Fig. 2 (a) Photograph showing the experimental design of the small quadrat, location of crab burrows, and plots with and without burrow (B+ and B− plots, respectively) within the quadrat. (b) Diagram of the sampling of soil sediments. Sediment samples (thickness: about 5 mm) were taken from the crab burrow walls, ground surface (1 cm depth) and deep part of the sediment (from 10 cm depth).

The diameter (B_{di} , cm) and the depth (B_{dep} , cm) of crab burrows within the B+ plots were measured to clarify the effect of the increased sediment–atmosphere interface area created by the burrow on the sediment $CO₂$ flux rate. For cases in which water had accumulated inside any burrows, B_{dep} was defined as the depth to the water surface. The sediment $CO₂$ flux rate in the B+ plots was standardized $(SR_{B + mod})$ by the total area of the sediment–atmosphere interface for consideration of the burrows.

The ST ($^{\circ}$ C) and SWC (%) at a sediment depth of 5 cm were also recorded at the time of the respiration measurement, using a portable temperature sensor (CT-422WR; Custom, Japan) and portable soil water content sensor (SM150T; Delta-T Devices, UK), to clarify their effects on the sediment $CO₂$ flux rate.

The relationships between sediment $CO₂$ flux and ST or SWC were expressed by the following exponential or linear equations:

$$
SR = ae^{bST} \tag{1}
$$

or

$$
SR = cSWC + d \tag{2}
$$

where $a, b, c,$ and d are fitted parameters. These equations were estimated separately for SR_{B-} and SR_{B+} mod. We then calculated the respiration rate at 25 °C (R_{25}) values using Equation (1). In addition, the Q_{10} value (index of the temperature dependency) was calculated to determine the respiratory activity as follows:

$$
Q10 = exp(10b)
$$
 (3)

To clarify the difference in microbial respiration activity among the positions within the sediments around the burrow, sediment samples (\sim 5 mm thickness) were taken from the burrow walls, ground surface part $($ \sim 1 cm depth), and deep part (\sim 10 cm depth). Ten sediments samples per plot were collected from three randomly chosen plots in the MS zone in the middle season (Fig. $2b$) and then composited by the sampling position within plots, to obtain three replicated samples $(n = 3)$. The sampled sediment was put through a sieve (2-mm) mesh size) to remove gravel and fine roots. The microbial $CO₂$ flux (microbial respiration) from the sample at 25 °C was measured by the open-flow method with an infrared gas analyzer (LI-840; Li-COR, USA) in the laboratory (Tomotsune et al. [2013b](#page-8-0)). The microbial respiration was expressed as the dry weight basis of the sediment, dried at 80 °C for 24 h.

Statistical Analyses

All statistical analyses were performed using SPSS software for Windows (SPSS Statistics 17.0; SPSS Inc., Chicago, IL, USA). Statistical comparisons were conducted using two-way

analysis of variance (ANOVA) (season \times zone for B_{den} and B_{area} , burrow \times zone for sediment CO₂ flux rate, ST and SWC) and one-way ANOVA (microbial respiration activity). Tukey multiple comparison was applied as a post-hoc test. Significant differences for all statistical tests were evaluated at $p < 0.05$.

Results

Variations in B_{den} and B_{area}

There were significant interactive effects by season and zone on B_{den} (Table 1). The B_{den} was significantly higher in the MS zone (90–147 m⁻²) than in the US zone (53–93 m⁻²), in all seasons, and B_{den} in the MS and US zones significantly decreased in the cool season (January 2017) (Fig. [3a\)](#page-4-0). In contrast, B_{den} in the DS zone showed no clear seasonal pattern and remained constant throughout the study period.

Significant interactive effects by season and zone were also found for B_{area} (Table 1). It was significantly higher in the DS zone than in other zones in the warm and cool seasons (Fig. [3b\)](#page-4-0). In contrast, the US and MS zones showed no significant differences in B_{area} among the seasons and zones.

Sediment $CO₂$ Flux Rate With/Without Burrows and Environmental Factors

There were no significant interactive effects by burrow and zone on the sediment $CO₂$ flux rate in each season (Table [2\)](#page-4-0). Burrows had a significant effect on the sediment $CO₂$ flux, and SR_{B+} was significantly higher (by 1.1–1.6 times) than SR_{B-} in each season (Fig. [4a, b, c\)](#page-5-0). In addition, the sediment $CO₂$ flux rate showed no significant difference among the zones except for in the middle season.

The B_{di} and B_{dep} of a burrow in the B+ plots varied greatly, from 0.5 to 5.5 cm and from 1.3 to 28.2 cm, respectively. The $SR_{B + mod}$, which was calculated based on these values, showed no significant interactive effects by burrows and zone on the sediment $CO₂$ flux rate in each season (Table [2\)](#page-4-0). In addition, burrows did not have a significant effect on sediment

Table 1 Two-way analysis of variance (ANOVA) (season \times zone) for burrow density and entrance area. Significant differences for all statistical tests were evaluated at the $p < 0.05$ level

	Density F df		\boldsymbol{D}	Entrance area p df	F	
Season	2°		$27.7 \le 0.001$ 2			4.16 0.017
Zone	\mathcal{L}		$35.4 \le 0.001$	$\overline{2}$		$38.3 \le 0.001$
Season x Zone 4		4.08	0.003		2.97	0.020

Fig. 3 Seasonal variations of (a) the burrow density (B_{den}) and (b) the burrow entrance area (B_{di}) , in the upper-, middle-, and downstream zones (US, MS, and DS, respectively)

 $CO₂$ flux, and SR_{B+mod} did not significantly differ from SR_B− in each season (Fig. [4d, e, f\)](#page-5-0).

The ST varied from 19.3 to 33.0 °C and from 19.3 to 33.5 °C in the B+ and B− plots, respectively. The SWC varied from 60.2 to 84.6% and from 60.3 to 84.6% in the B+ and B− plots, respectively. In each season, there was no significant difference in ST or SWC between the B+ and B− plots (data not shown).

The microbial respiration activity showed a significant difference among sampling positions; microbial respiration was significantly higher in the ground surface and burrow wall parts than in the deep part (Fig. [5](#page-5-0)). However, there was no significant difference in microbial respiration in sediment between the ground surface and burrow wall parts.

The SR_B– and SR_{B+mod} positively correlated with ST (R^2) $= 0.238, p < 0.05$ for SR_B ; $R^2 = 0.221, p < 0.05$ for SR_B + mod) and SWC ($R^2 = 0.010$, $p < 0.05$ for SR_B ; $R^2 = 0.029$, $p < 0.05$ for $SR_{B + mod}$) (Fig. [6\)](#page-6-0). The regression curves showed no major difference between SR_{B-} and SR_{B+} mod. Moreover, R_{25} and Q₁₀ were similar between SR_B− (127 mg CO₂ m⁻², 2.23) and $SR_{B + \text{mod}}$ (128 mg CO₂ m⁻², 2.27). In addition, anomalously high values of the sediment $CO₂$ flux rate (positioned far from the regression curve) were occasionally observed for SR_{B-} and SR_{B+} mod.

Discussion

Our results indicate that the crab burrow densities and entrance area sizes varied spatially and seasonally (Fig. 3). A previous study conducted near our study site (Ishigami et al. [2005\)](#page-7-0) reported that the phenology of Uca spp. is different from that of Scopimera spp. in southern Japan; the activity of Uca spp. on the sediment surface is high from August to November, whereas that of Scopimera spp. is high from December to May. In our observations, Uca spp. inhabited all zones, while Scopimera spp. only inhabited the DS zone. Therefore, in the DS zone, these taxa might replace each other throughout the year, and thus there was no seasonal variation in burrow density in that zone (Fig. 3a). Burrow entrance area was positively correlated with crab size (Ishigami et al. [2005](#page-7-0); Otani et al. [1997](#page-8-0)), and our results suggest that the crab communities inhabiting the DS zone included larger-sized individuals. A seasonal variation in crab size is affected by species-specific life cycles and interactions between species. For example, for 1 or 2 years after hatching, most crab species float as plankton (zoea and megalopa larvae) in the seawater and will fix to sediment surfaces in mangrove forests (Shih [2012\)](#page-8-0). This life cycle is also variable by regional climate and is dependent on interactions with other species. Unfortunately, detailed informations about these

Table 2 Two-way analysis of variance (ANOVA) (burrow × zone) for sediment CO₂ flux rate based on ground surface area and modified sediment CO₂ flux rate based on sediment-atmosphere interface area. Significant differences for all statistical tests were evaluated at the $p < 0.05$ level

Fig. 4 Difference in averaged sediment $CO₂$ flux rate between plots without (B−, white bar) and with crab burrows (B+, gray bar), in the upper-, middle-, and downstream zones (US, MS, and DS, respectively), during the (a) warm (August 2016), (b) middle (November 2016), and (c) cool (January 2017) seasons. Panels (d), (e), and (f) show the standardized sediment $CO₂$ flux rate based on the calculated burrow surface area (gray bar) in the warm, middle, and cool seasons, respectively. Error bars indicate standard deviation $(n =$ 25–40)

aspects are limited, especially for our study area. However, the coexistence of two crustacean species with likely different phenology and life cycles may have affected the measure of seasonal variation in burrow entrance area that was observed in the DS zone (Fig. [3b\)](#page-4-0).

An increasing rate of sediment $CO₂$ flux caused by crustacean burrows has been previously reported for coastal

Fig. 5 $CO₂$ flux from ground surface sediment, deep sediment, and crab burrow wall sediment near the middle-stream zone. Error bars indicate standard deviation $(n = 3)$

ecosystems such as sandy and muddy tidal flats (Kinoshita et al. [2003;](#page-7-0) Laverock et al. [2010](#page-7-0); Sasaki et al. [2014](#page-8-0)). In particular, Sasaki et al. (2014) (2014) reported that the sediment $CO₂$ flux rate increased by up to three times with the presence of burrows in a sandy tidal flat along the Seto Inland Sea. Our study is the first report to detect this phenomenon in a natural mangrove forest, as the sediment $CO₂$ flux rate increased by 1.1– 1.6 times with the presence of burrows. The increased sediment $CO₂$ flux can be explained by several physicochemical and biological factors (e.g., Bowden et al. [1998;](#page-7-0) Cleveland et al. [2007;](#page-7-0) Franzluebbers [1999](#page-7-0)). In general, the changes in ST and SWC could explain the variation of microbial respiratory activity and consequently the sediment $CO₂$ flux (Davidson et al. [2000;](#page-7-0) Hanson et al. [2000\)](#page-7-0). However, there were no significant differences in ST and SWC between the B− and the B+ plots in our study, and thus the effects of those factors on the increased sediment $CO₂$ flux are probably negligible. Conversely, the standardized sediment $CO₂$ flux rate by sediment–atmosphere interface area in the B+ plots showed no significant difference as compared with the sediment $CO₂$ flux rate in the B− plots (Fig. 4). Therefore, the higher sediment $CO₂$ flux rate in the B+ plots could be explained by the increased sediment–atmosphere interface area created by the burrows.

Those results, and data showing the active microbial respiration potential of burrow walls to be equivalent to that of the ground surface sediment, indicate that the deep sediment had changed to aerobic conditions through exposure to the atmosphere as a consequence of crustaceans burrowing. Sasaki et al. ([2014](#page-8-0)) reported that the presence of

Fig. 6 Relationships between sediment $CO₂$ flux rate and sediment temperature (ST) or sediment water content (SWC) at 5 cm depth. Black circles and white circles indicate sediment $CO₂$ flux rate in the plots without crab burrows (SR_{B-}) and those with burrows (SR_{B+}), standardized by the surface area of the burrow $(SR_{B + mod})$. Panel (a)

mud shrimp burrows increased the oxidation-reduction potential of the sediments around burrows and consequently increased the $CO₂$ flux from burrow wall sediments in tidal flats. In addition, Kinoshita et al. [\(2003\)](#page-7-0) reported that mud shrimps provided the fresh organic matter with the microbial communities that inhabit burrow walls by taking foods into their burrows and excreting inside, which would change the microbial community and its activities in the burrow walls. According to theses previous reports, it had been expected that the microbial respiration of the burrow wall sediment could be higher than that of the ground surface sediment; however, our results were different to these expectations (Figs. [5,](#page-5-0) 6). One possible reason for this disagreement would be a difference in turnover rates of the burrows. In general, crab burrows are characterized by shallower and simpler shapes as compared with mud shrimp burrows and, consequently, will have a faster turnover (Dye and Lasiak [1987](#page-7-0); Myers [1979\)](#page-7-0). The faster turnover potential would suggest a less-sufficient time to stimulate the microbial community inhabiting burrows created by crabs as compared with the situation for mud shrimp burrows. Therefore, changes in the microbial community would have had a relatively small effect on the sediment $CO₂$ flux rate in our study.

The annual sediment $CO₂$ flux in mangrove forests has been previously estimated at between 0.73 and 2.31 Mg C ha⁻¹ yr⁻¹ (Alongi et al. [2001\)](#page-7-0) and from 1.93 to 2.45 Mg C ha^{-1} yr⁻¹ (Poungparn et al. [2012\)](#page-8-0). However, the previous studies did not describe whether or not burrows were included in the measurement plots for sediment $CO₂$ flux rate. Our results indicate that not including the burrows would underestimate the annual sediment $CO₂$ flux. Moreover, although the direct macrobenthos respiration was not measured in our study, it should have a non-negligible effect on the evaluation of annual sediment $CO₂$ flux. For example, Pülmanns et al. [\(2014\)](#page-8-0) conducted laboratory experiments with coastal crab

indicates the ST relationships based on SR_{B-} (solid line; SR_{B-} $17.05e^{0.080ST}$, $R^2 = 0.240$) and $SR_{B + \text{mod}}$ (broken line; $SR_{B + \text{mod}} =$ 16.61 $e^{0.0818ST}$, $R^2 = 0.262$). Panel (b) indicates SWC relationships based on SR_{B-} (solid line; $SR_{B-} = 7.23$ SWC – 368.78, $R^2 = 0.050$) and $SR_{\text{B + mod}}$ (broken line; $SR_{\text{B + mod}} = 4.31$ SWC – 160.14, $R^2 = 0.029$)

Ucides cordatus and reported that $20\% - 60\%$ of the CO₂ released by burrows originated from crab respiration. Burrow density and structure, which are variable depending on crab density and community, show geographical and locational heterogeneity. For example, some previous studies in mangrove forests reported the range of burrow density at 5.4– 16.4 m⁻² in China (Li et al. [2015\)](#page-7-0) and 75–145 m⁻² in Pakistan (Qureshi and Saher [2012\)](#page-8-0), whereas the range for our study site ($54-147 \text{ m}^{-2}$) was relatively high. In other mangrove forests, Ocypode spp. and Ucides spp., which are lagersized than *Uca* spp., often dominate and construct deeper and more complex burrows (Pülmanns et al. [2014\)](#page-8-0). Therefore, understanding the local traits regarding the different penology and life cycles of the macrobenthos is important for clarifying their roles as accelerators of carbon cycling in mangrove ecosystems through their direct and indirect contributions to sediment $CO₂$ flux.

Previous studies have occasionally observed a high sediment $CO₂$ flux rate in some plots despite the low average sediment $CO₂$ flux rate in mangrove forests (Alongi [2014;](#page-7-0) Tomotsune et al. [2018](#page-8-0)). These studies concluded that the presence of burrows could explain spatial variation of the sediment $CO₂$ flux rate. In our study, however, similar high sediment $CO₂$ flux rates were observed even in the B− plots (Fig. 6), suggesting that other factors are also likely to contribute to the unexpectedly high sediment $CO₂$ flux. For example, fine root respiration at shallow sediment depths might contribute, because metabolic $CO₂$ might leak out from broken and dead roots in the pedosphere. In addition, the high spatial variation of organic matter in sediments might also contribute, because a large amount of fallen litter is transferred by the high tidal range and is often stored around the prop and respiratory roots of mangroves. However, the relationships between these factors and the sediment $CO₂$ flux rate remain unknown at the plot scale, and thus further study is required.

In conclusion, crab burrow density and entrance area varied seasonally and spatially as a result of crab phenology, life cycle, and variations in crab communities. Crab burrows increased the sediment $CO₂$ flux rate mainly through a physical effect, that is, through an increase in sediment–atmosphere interface area. Therefore, it is important to consider the effects of crustacean burrows in the mangrove forest floor when evaluating sediment $CO₂$ flux and when examining the role of the mangrove ecosystem as a carbon sink.

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