

Historical maps and satellite images as tools for shoreline variations and territorial changes assessment: the case study of Volturno Coastal Plain (Southern Italy)

I. Alberico¹ $\mathbf{D} \cdot \mathbf{G}$. Cavuoto¹ \cdot V. Di Fiore¹ \cdot M. Punzo¹ \cdot D. Tarallo¹ \cdot N. Pelosi¹ \cdot L. Ferraro¹ \cdot E. Marsella¹

Received: 9 March 2017 /Revised: 25 September 2017 /Accepted: 16 October 2017 /Published online: 16 November 2017 \oslash Springer Science+Business Media B.V. 2017

Abstract Anthropic pressure has caused severe variations of Mediterranean coastal areas currently hosting about 480 million people. The replacement of natural land covers with crops and urban environment coupled with the reduction of sediment supply to the coast, subsidence, Relative Sea Level Rise and the high frequency of storm events, have caused severe shoreline erosion. In this paper, we stress the key role of historical maps, topographic maps and free satellite images to forecast the rates of coastline changes and to recognize the main features of past landscapes as tools for risk reduction. This data was recorded into a Geographical Information System dedicated to coastal management that allows to compare different coastal zones and elaborate maps. The analysis was applied to the case study of Volturno Coastal Plain (VCP), extending from the town of Mondragone to Patria Lake (Campania Region, Southern Italy). Indeed, the intense territorial modification that occurred between the seventies and eighties, coupled with the exposure to coastal erosion, make the VCP a good test area. The spatial analysis algorithms allowed to outline the main features of past landscapes and how they changed from roman times to present while the coastal evolution (erosion, accretion) and possible future coastal trend was assessed with the Digital Shoreline Analysis System (DSAS) software. Starting from the Bourbon domain, the reclamation caused the first great territorial change (e.g. wetlands were transformed into agricultural lands, regimentation of surperficial water) but the negative effects of antrophic pressure, as the intense urbanization of

the coastal belt, emerged in the seventies of the last century. The shoreline position was defined for 9 time intervals (from 1817 to 2012) as the ratio of the distance between two shorelines and the relative elapsed time. Moreover, for the 1957– 1998 and 1998–2012 time windows, the shoreline trends were calculated with the weighted linear regression method. The first trend pointed out an intensive erosional phase (mean value: 5 m/yr) for a wide sector close to the Volturno River mouth, the eroded sediment nourished the beaches of other coastal sectors. This phase was related to the reduction of River sediment supply mainly due to the construction of the Ponte Annibale dam on the Volturno River. The second (1998–2012) showed an alternation of erosion and accretion sectors due to a sediment starved condition to deltaic zone and to a sequence of 52 sea protection works in the Gaeta Gulf. Furthermore, the regression values of more recent time interval, was assumed as a scenario to draw the probable shoreline position in 2022. The overlay of this shoreline on the Technical Maps of Campania Region at 1:5000 scale highlighted the urban area that could be exposed to damages.

Keywords Historical analysis . Coastal trend . Natural landscape . Natural hazards . Coastal management

Introduction

The presence of fresh water, easily cultivable lands and natural resources offered by the sea exposed the coastal zones to an uncontrolled urbanization.

Currently, the world coastal population density (99.6 people per $km²$) is more than double of inland areas (37.9 people per km²), (Millenium Ecosystem Assessment [2005\)](#page-18-0). In European coastal regions (Nomenclature Des Unitès Territoriales Statistiques – NUTS3; EUROSTAT [1988](#page-17-0)), this

 \boxtimes I. Alberico ines.alberico@iamc.cnr.it

¹ Istituto per l'Ambiente Marino Costiero, CNR, Calata Porta di Massa, 80133 Napoli, Italy

value is on average 10% higher than inland, only in some countries it is more than 50% higher (EEA Report [2006](#page-17-0)). However, the population density map for European coastal zone (NUTS3) shows a value varying from 100 to 500 habitants per square kilometer. Higher values generally characterize the Southern Europe.

The presence of one-third of the population (480 million people) in the Mediterranean coastal areas has caused strong territorial tranformations in the last centuries (SOER [2015\)](#page-18-0). The replacement of wetlands with crops, roads, and urban environment coupled with the changes in sediment supply to the coast (about 45% of the sediments are retained behind dams or extracted from rivers (UNEP/MAP [2009](#page-18-0)) are the main causes of shorelines realignment. Moreover, artificial structures, the destruction or degradation of sea grass beds and dune vegetation, the extraction of gas and water may also affect the cycling and redistribution of sediment in neighboring coastal areas (UNEP/MAP [2009\)](#page-18-0). Approximately onefourth of the European Mediterranean coastline is subject to erosion.

In the last centuries, the Sele and Volturno-Garigliano coastal plains, the two largest sandy coastal areas of the Campania Region, as many other Italian coastal plains (Cocco and Pippo [1988](#page-17-0); Cocco and Magistris [1988;](#page-17-0) Cocco and Iuliano [1999](#page-17-0); Bellotti [2000;](#page-16-0) Pranzini [2001;](#page-18-0) Aucelli et al. [2009;](#page-16-0) Alberico et al. [2011](#page-16-0)), have suffered erosion.

Starting from the analysis of historical documents and free satellite images, the main aim of the present work was to identify the main features of past landscapes, territorial changes, shoreline variations and shoreline trends (del Pozo and Anfuso [2008](#page-17-0); Dickson et al. [2007;](#page-17-0) Dolan et al. [1980;](#page-17-0) Dolan et al. [1991](#page-17-0); Gibeaut et al. [2003](#page-17-0); Hapke et al. [2006;](#page-17-0) Jimenez et al. [1997](#page-17-0); Leatherman [1983](#page-18-0); Morton et al. [2004;](#page-18-0) Pajak and Leatherman [2002](#page-18-0); Romagnoli et al. [2006](#page-18-0); Alberico et al. [2011\)](#page-16-0). This knowledge can be considered essential for natural risk reduction. In fact, past land-uses represent a guide for environmental restoration (Sutton-Grier et al. [2015](#page-18-0) and reference therein) while the shoreline variations and shoreline trends point out the coastal changes and the possible future shoreline position.

The Volturno Coatal Plain (VCP) was chosen as a case study because it was exposed to erosional hazard and had been deeply modified by antrophic pressure. The urban environment currently occupies the greatest part of the coastal belt.

The coastal erosion at VPC followed two main events: the reclamation works, realized in the lower part of this plain during the Bourbon domain, and the construction of Ponte Annibale dam on the Volturno River in the fifties, both events caused a reduction of river sediment supply to the coast. After the seventies, the extensive urbanization of coastal zones and the construction of several coastal defense structures caused the

alternation of accretion and erosion. Moreover, the trend calculated for the 1998–2012 time window, was used to identify the possible position of shoreline in 2022. The overlay of this shoreline over Technical Maps of Campania Region offered the possibility to identify the urban areas possibly exposed to risk in the next future.

The data and the procedures implemented for this study were integrated into the VPC coastal management framework (Alberico et al. [2017](#page-16-0)). Taking into account that coastal erosion is not the only threat to coastal zones, this framework is flexible and can easily be upgraded and integrated with other environmental data related to natural hazards (marine inundation, salt water intrusion, storm events), often exacerbated by Global Warming (IPCC [2014](#page-17-0)) and land-uses.

Geological and geomorphological setting

The Volturno Coatal Plain is an integral part of Campanian Plain ($\sim 5700 \text{ km}^2$, therein CP), which stretches for several kilometers inland from the coast and is bounded toward the Tyrrhenian Sea by a narrow coastal strip (2–0,15 km) running from the eastern border of Massico Mountain to Patria Lake. The continental shelf is about 16 km wide in front of the Volturno River and narrows southward, the shelf break is located at about 125 m depth (Iorio et al. [2014\)](#page-17-0). The convexity of the isobaths in front of Volturno River at 100–150 m depth, identify the ancient river mouth shifted seaward during the last glacial cycle (Pennetta et al. [1998](#page-18-0)) (Figs. [1](#page-2-0)a–c).

According to several geological reconstructions (Ippolito et al. [1973;](#page-17-0) Aprile and Ortolani [1978](#page-16-0); Ortolani and Pagliuca [1986;](#page-18-0) Cinque et al. [1987;](#page-17-0) Cinque et al. [2000;](#page-17-0) Bruno et al. [2000\)](#page-17-0), the CP represents a morphostructural basin that has been eveloping since Late Pliocene along the palaeo-Tyrrhenian margin of the Apennines fold-thrust belt as a consequence of the post-orogenic phases and extensional processes (stretching and thinning of the continental crust) linked to the Tyrrhenian rift-basin (e.g., Doglioni [1991](#page-17-0); Patacca and Scandone [2007;](#page-18-0) Ferranti and Oldow [1999](#page-17-0); Milia and Torrente [2003\)](#page-18-0) (Fig. [2\)](#page-3-0).

The CP boundaries are defined at the surface by NW–SE, NE–SW and E–W trending faults with polyphasic history, which define several half-grabens in subsurface (from northwest to southeast: Volturno, Patria Lake, Campi Flegrei-Acerra and the Naples Bay basins) and downthrow the Meso-Cenozoic Apennine units forming the basin flanks and basement by about 2000– 5000 m (Mariani and Prato [1988](#page-18-0); Milia and Torrente [2003](#page-18-0); Milia et al. [2013\)](#page-18-0).

Fig. 1 The Volturno Coastal Plain is bounded by Massico Mountain to the northwest and by Patria Lake to the south (a). The location of the study area is illustrated at Regional (b) and National scale (c)

Starting from the early Pleistocene, the CP has recorded the onset of rapid tectonic subsidence and intense infilling of continental and marine deposits (mainly constituted by alluvial, fluvial, swampy, lagoonal, and paludal deposits) interfingered with the products of intense explosive eruptions related to the Somma-Vesuvio district and Campi Flegrei supposedly controlled by regional faults (De Vivo et al. [2001;](#page-17-0) Milia and Torrente [2003;](#page-18-0) Romano et al. [1994;](#page-18-0) Cinque et al. [1997](#page-17-0)).

The late Pleistocene-early Holocene sea-level rise (ca. 15 ky–6 ky BP) promoted the rapid flooding of lower VP followed by the maximum coastline retreat, which from 6 km reached to 10 km inland with respect to the present shore (Amorosi et al. [2012](#page-16-0); Sacchi et al. [2014](#page-18-0)). Since ca. 6.5 ky BP, a progressive change from transgressive to regressive conditions marked the establishment of a coastal progradational trend (ca. 4 ky BP) causing the seaward migration of the adjacent plain and the growth of Volturno River delta (Amorosi et al. [2012;](#page-16-0) Barra et al. [1989;](#page-16-0) Cinque et al. [1997\)](#page-17-0). At about 2 ky cal BP, the continuous progradation favored beach, marsh and lagoonal environments which still persist along the present coastal zone (Ruberti et al. [2014](#page-18-0); Amorosi et al. [2012](#page-16-0); Sacchi et al. [2014\)](#page-18-0).

At present, the coastline is fairly straight, except for the prominent Volturno River mouth, where a large swamp area as the remnants of a series of beach ridges and swales elongated in NW/SE direction is preserved. This coast is classifiable as dissipative-type and wave-dominated (Bellotti [2000\)](#page-16-0). The surface circulation in the Tyrrhenian Sea is characterized by complex patterns, dominated by both (semi) permanent recirculation and transient features, often hindering the identification of a consistent mean flow even at the large (basin) scale, and specifically in its south-eastern sector (Rinaldi et al. [2010;](#page-18-0) Iacono et al. [2013](#page-17-0)). The situation is further complicated in the coastal areas, and particularly in the Gulf of Gaeta, where the local circulation is influenced mostly by wind direction and variability, as evidenced by recent studies by Iermano et al. [\(2012](#page-17-0)) based on both numerical modelling and satellite remote sensing. In particular, their analyses suggest that the recurrence of periods of wind reversal and fluctuating freshwater discharge rates may substantially modify the local circulation and crossshelf exchange close to the river mouths.

Materials and methods

The orthophoto maps, satellite images, historical and topographic maps are the data sources used to evaluate the main features of past landscapes and the shoreline variation at VCP from 1817 to 2012 (Table [1](#page-4-0)).

The main natural elements (i.e. wetlands, pinewoods, dunes, beaches) characterizing the past landscapes were digitized and designed as part of coastal management framework published by Alberico et al. [\(2017\)](#page-16-0) (Fig. [3](#page-4-0)).

The shoreline variations were assessed with the Digital Shoreline Analysis System (DSAS) software (Thieler et al. [2009](#page-18-0); Kumar et al. [2010](#page-18-0); Abuodha and Woodroffe [2010;](#page-16-0)

Fig. 2 (a) Schematic structural map of Campanian Plain and interpretative cross section of the Campanian margin redrawn from Milia and Torrente [\(1999\)](#page-18-0). (b) Simplified geological-geomorphological map of the Volturno Plain based on the reinterpretation of survey works of

Alberico et al. [2011;](#page-16-0) Kuleli et al. [2011;](#page-18-0) Aiello et al. [2013](#page-16-0); Albuquerque et al. [2013;](#page-16-0) Kaliraj et al. [2015;](#page-17-0) Nguyen and Woodroffe [2016\)](#page-18-0). The software ensures a repeatable procedure, favors a systematic update and allows a comparison between different coastal zones.

For this analysis, the geo-referencing procedure, seasonal shoreline variations and the accuracy in identifying and digitizing the land–water interface (Crowell et al. [1991;](#page-17-0) Moore [2000](#page-18-0); Morton et al. [2004](#page-18-0); Romagnoli et al. [2006](#page-18-0); Thieler and Danforth [1994](#page-18-0)) are sources of potential errors affecting the shoreline variations assessment.

In accordance with del Pozo and Anfuso ([2008](#page-17-0)) the shoreline position was defined as the water line at the time of the historical documents or satellite images because the study area is located in a microtidal environment $(\pm 20 \text{ cm})$. Moreover,

1:100.000 scale Geological sheets n°171, 172, 184 (ISPRA – Servizio Geologico d'Italia [1966;](#page-17-0) [1967;](#page-17-0) [1968\)](#page-17-0). Ancient shorelines are derived from Amorosi et al. ([2012](#page-16-0)); Sacchi et al. ([2014](#page-18-0)); Ortolani and Pagliuca ([1999](#page-18-0)); Cocco et al. ([1992](#page-17-0))

for the same reason we considered the daily and seasonal water line position exposed to a very low uncertainty and thus the accuracy of shoreline detection is based on the accuracy of geo-referencing process.

The historical and topographic maps were georeferenced according to the technical map of Campania Region at 1:5000 scale that resides in the WGS84 coordinate system. The process involves the assignment of geographical coordinates, as read on the technical map of the Campania Region, to the various control points (road junctions, historic monuments, geodetic points) identified on non-geo-referenced maps. The accuracy associated to this procedure, expressed by RMSE (Route Mean Square Error) is shown in Table [1.](#page-4-0)

The positional accuracy of GoogleEarth® imageries required the comparison of georeferenced points, defined on

Table 1 Topographic maps, ortophoto maps and satellite images employed for the study of territorial changes and shoreline variation at Volturno Coastal Plain. For each documents are reported: scale and Root Mean Square Errors (RMSE) of georeferencing process

Date	Input data	Data Source holder	Scale	Uncertainty (RMSE, m)
1817	Topographic and hydrographic map of areas close to Naples	Historical map of IGM	1:25,000	30
1860	The map of the kingdom of Naples	Historical map of IGM	1:20,000	20
1936	IGM topographic map	Topographic map	1:25,000	12
1957	IGM topographic map	Topographic map	1:25,000	4
1987	IGM topographic map	Topographic map	1:25,000	8
1989	Orthophoto map	Ministry for the Environment, Land and Sea	1:10,000	5.7
1997	Orthophoto map	Ministry for the Environment, Land and Sea	1:10,000	6.2
1998	Technical map	Campania Region	1:5000	2
2004	Google Earth image			6.81
2010	Google Earth image			2.67
2012	DAEDALUS satellite image	Consorzio TELAER STA	Pixel size:1.3 m	4.8

the technical Map of Campania Region, with corresponding points as displayed by the GoogleEarth® application (Potere [2008](#page-18-0); Casella et al. [2010;](#page-17-0) Becek et al. [2011a](#page-16-0), [2011b](#page-16-0); Kazimierz et al. 2011). To this aim, the coordinates of 16 control points, mainly corresponding to building corners, are compared to define the RMSE of the Google Earth imagery

Fig. 3 Flow chart illustrating the main steps of procedures adopted in the present work for the territorial changes analysis and the shorelines variation assessment (right side of image). A simplified representation of framework (left part of image) dedicated to the antrophic pressure assessment (Alberico et al. [2017\)](#page-16-0) and the key role of data integration are illustrated

(Table 2). A similar procedure was applied to assess the accuracy of orthophoto maps of the Italian Environmental Ministry and Daedalus satellite image (Table 2).

For the shoreline analysis, 10 shorelines, 170 transects (lines along which the shoreline variations and trends are calculated), a baseline (the starting point for all transects), the spacing and length of transects and the shoreline parameters (e.g. shorelines date, uncertainty in position) were used as input data. The baseline was traced back to all shorelines; it matches with a more smoothed shape of the shoreline in 1998. The transects were spaced every 200 m; their intersection with the shorelines identify the coastal points used in DSAS to calculate, for the single transect, both the rate of shoreline changes and the shoreline trend (Figs. [4](#page-6-0)a–c).

The shoreline changes were calculated, for each pair of consecutive shorelines, as the ratio of the distance between two shorelines and the elapsed time (Fig. [4](#page-6-0)b) (Addo et al. [2008](#page-16-0); Alberico et al. [2011;](#page-16-0) Houseret al. 2008; Leyland and Darbya [2008;](#page-18-0) Thieler and Danforth [1994](#page-18-0)). These measurements allowed to reveal the shoreline behavior for nine time intervals.

The shoreline trend was evaluated by applying the weighted linear regression method. All the coastal points available along the single transect (Fig. [4c](#page-6-0)) contributed to the trend assessment with a weight related to the accuracy of shoreline position in detail, the lower the uncertainty (low georeferencing error) the higher the weight (Genz et al. [2007\)](#page-17-0). The coastal trend was calculated for the 1950–1998 and 1998–2012 time windows. The latter was also used to identify the shoreline position for the year 2022. To this aim, assuming that the driving factors were unmodified, the rate of erosion calculated for each transect was multiplied by ten years. This shoreline was then superimposed on the Technical Map of Campania Region at 1:5000 scale to identify the urban zones possibly exposed to erosion in the next future.

The environmental change of coastal area from roman times to present

During the Roman Empire (27 B.C. - 476 A.D.), the plain close to the Volturno River, named Campania Felix, became the core of a rich agricultural resource. In this period several infrastructures were built; the Domitiana road and the bridge connecting the Volturno River sides (Canino [1981](#page-17-0)). After the fall of the Western Roman Empire (476 A.D.), the agricultural crisis and the barbarian invasions caused an abandonment of plain. This event was probably associated to the unset of Late Antique Little Ice Ages from 536 to ca. 660 AD. This cooling phase, spanning most of the northern hemisphere, was considered by Büntgen et al. [\(2016](#page-17-0)) and also recognized in Margaritelli et al.

Table 2 List of points used in the georeferencing process of Google Earth (2004, 2010), orthophoto maps (1989) and Daedalus image (2012). For each point, the differences in coordinates (DX and DY) between the Technical map, used as master map, and the other documents and the Root Mean Square are listed

Fig. 4 The shorelines detected from the 1817 to 2012 were superimposed to Digital Elevation Model of Volturno Coastal Plain (upper left side of image "a''). Three detailed maps show the shoreline variation at Volturno River (A), Agnena (B) and Regi Lagni mouths (C). In figure "b'', the map illustrates a subset of 170 transects

[\(2016](#page-18-0)), as an additional environmental factor contributing to change in social reorganization in Europe and Asia. During the Norman domain (1130–1189 A.D), the Campania was enclosed in the Regno di Sicilia and dominated by Altavilla, Hohenstaufen, Angioini and Aragonesi. During the Aragonese dynasty and under Alfonso V Napoli became a capitol of the Mediterranean area while a decline occurred as a consequence of high tax pressure applied by Carlo V during his domain. Only after the Austrian domain, the ascension to the throne of Bourbon dynasty (1735–1860) favoured the improvement of human well-being thanks to the promulgation of economic and legislative reforms.

In the topographic map of 1816, the VCP is a natural land and four main rivers are present in the area: a) the Volturno River in the central part of the plain, b) the Savone and Regia Agnena

realized along the Volturno Coastal Plain to assess the shoreline variation, the red transect illustrates the methodology followed to assess the rate of shoreline change between two shorelines consecutive in time. In figure "c'', the shoreline trend from 1957 to 1998 is pointed out

rivers in the northern territories, they ending their path in swamp areas close to the coast and c) the Clanio River in southern territories, which ran into the Patria Lake through the Vena channel (Figs. [5](#page-7-0)a, b) (Savarese [1856\)](#page-18-0). The reclamation of Clanio River, renamed as Regia Agnena, started in 1592 by will of Conte Di Miranda and continued under the Conte Lemos, from 1599 to 1603, ending in 1610 under the Conte di Benavente. The morphology of territories to the left of Voturno River was described as a plain with very low slope while a variable morphology characterized the right side. The large amount of sediment discarded by the Volturno River favoured a rapid accretion of beach at the *Bagnane* mouth causing the development of homonym canal parallel to the coast (Fig. [5b](#page-7-0)). In time the river reconstructed a new mouth to flow directly at the sea. This process repeated itself several times (Afan de Rivera [1847](#page-16-0)).

Fig. 5 The maps illustrate the main streams and swamp areas characterizing the Volturno Coastal Plain at the beginning of nineteenth century

In the cartography of 1817, Mondragone, S. Nicola, S. Angelo and San Vito urban centers characterized the northern area of VCP while a small group of houses, representing the source of modern Castel Volturno town, were located on the right side of the *Volturno* River (Fig. [6](#page-8-0)). A few farms were spread over the territory and the Domitiana road, joining Castel Volturno and Capua, is detectable. A pine forest characterized the area very close to the shoreline and dune systems, with elevation ranging from 3 to 5 m were present on the right side of the Volturno River. In the map of 1822, the swamp zones labelled Pantano di Mondragone, Pantano di Castello and Il Pantano and the Regi Lagni canals were well detectable (Fig. 5b). In 1831, Ferdinando II established the authority "Amministrazione" Generale di Bonificazione" to reclaim these areas. In 1837, the hydraulic structures of Regi Lagni was improved. In 1841, the reclamation continued in the lower part of VCP and about 54.754 moggia (1 moggio in the study area = c.a. 3300m²) became agricultural lands. As first step, the Regia Agnena, used to drain the territory located between the Volturno and Savone rivers, was modified and its channel was extended to the sea. In 1842, a large pond located in the lowland area labelled $\lq e$ Bagnane", between the Volturno River and Regia Agnena, was infilled thanks to the construction of a diversion (Bartolotti canal)

from the Voturno River; this construction worked very well and thus in 1845 a new canal (de' Militari) was built to infill this swamp area. In 1849, only the Napoli and Caserta Provinces supported the reclamation contributing to the construction of about 72 km of main canals, 254 km of secondary canals, 150 km of trenches and the new mouth of Clanio stream (today known as Regi Lagni).

Starting from 1850, the reclamation continued with the landfill of lowland areas; il Bosco, il Pantano di Mondragone, le Terre di Bartolotti and Pantano di Castello well visible in fig. 5b. After 1855, the lack of preservation works favoured damages at the left bank of Clanio River and most of the territories close to the Regi Lagni were flooded (Ciasca [1928](#page-17-0)). In 1917, the Institution of "Opera Nazionale Combattenti" restarted with reclamation work; several water pumps were placed in the lower part of VCP, the Regi Lagni mouth was rebuilt and the left river bank of Volturno River was restored.

In 1860, two main elements structured the territory: a wide pine forest, namely "Paneta di Mondragone" and "Paneta di Castel Volturno", located on both sides of the Volturno River mouth and the Regi Lagni channel in the southern area (Fig. [6\)](#page-8-0).

In 1907, Castel Volturno preserved its extension while Mondragone grew towards the coast and the road network

Fig. 6 The figures illustrates the main features characterizing the past landscape and its evolution through time. Only the principal elements were drawn, location of other elements or localities are indicated by a number. (1) dunes, (2) Pineta di Mondragone, (3) Pineta di Castel

Volturno, (4) Scuderia locality Villaggio Agricolo locality, (5) Scatozza locality, (6) Masseria Pineta d'Ischitella, (7) I Variconi, (8) Pinetamare locality, (9) Punta dei Tamari locality. The dotted red line is the Domitiana road

became denser than in the previous years (Fig. 6). After thirty years, Mondragone reached the coast and the Cirio factory was built at about 263 m from the shoreline. On the right side of the Volturno River, a cluster of buildings labelled Villaggio Agricolo was built at about 900 m from the shoreline and the road network become denser than in the past.

In 1936, the Consorzio in Destra del Basso Volturno deemed it fundamental to build the Nuova Inalveazione del Torrente Savone channel, realized from 1949 to 1954, to collect the waters of mountain rivers of the south-eastern sector of Roccamonfina, of the eastern slope of the Monte Massico and of north-eastern slope of Monte Maggiore. After World War II, the lower part of Regi Lagni and of Regia Agnena were improved and several depressed areas were infilled. The territorial reorganization of lower part of VCP ended in the fifties.

In the 1957, a group of houses were built north of Cirio factory at about 130 m from the coast (Mondragone municipality); single buildings lined the *Domitiana* road and new buildings rose in the Scatozza locality. In the locality "Masseria Pineta d'Ischitella" an increase of houses and of a single bathhouse very close to the coastline are evident (Fig. 6).

The quantitative analysis of urban growth for the last 200 years, pointed out that the VCP was affected by an urban expansion lower than 1% from 1817 to 1936. This value reached 5% at the end of 1957, although territorial transformation reached its peak between 1957 and 1984. In this period the urban expansion increased by about 80% and the seaside tourist centre of Pinetamare was also built (Alberico et al. [2017](#page-16-0)) (Fig. [7](#page-9-0)). This urban expansion occurred in areas very close to the shoreline between Punta dei Tamari and the Savone stream; southwards the urban environment was sparse, it became dense and closer to the coastline (about 400 m) at Volturno River. To the south of Scatozza locality, the urban environment was dense along the coastal belt with the only exception of Variconi Oasis. From Scuderia locality to Regi Lagni the urbanized zones were farther from the shoreline (850 m) than later areas. Southward, a wide

urbanized area overtook the Domitiana road while on the coast a big touristic village Pinetamare (also known as Villaggio Coppola) with a little harbor, was being built. In the Masseria Pineta d'Ischitella locality an increase of urban zones that reaches the Patria Lake was recorded (Fig. [8\)](#page-10-0). In the following years, the percentage of new urbanized areas decreased by 12% because major urbanization had already saturated the area. After the nineties, urbanized areas increased by very low percentages (lower than 2%) similar to that characterizing the VCP before the sixties (Fig. 7) probably due to the publication of Campanian Regional Territorial Plan (2008), Caserta Provincial Territorial Plan ([2012](#page-17-0)) and Hydro-geomorphological Setting Plan of Basin Authority of Liri-Garigliano-Volturno [\(2010\)](#page-17-0).

Shoreline changes analysis

Starting from the cartography of the 1817, the first map useful for a quantitative evaluation, the shoreline variation was assessed for nine time intervals (Fig. [9\)](#page-11-0). Prior to this data, thanks to the discovery of the ancient road "Via Domitiana", of ruins of coastal villas (villae maritimae) and of the thermal settlements of "*Aquae Sinuessanae*", the shoreline of Roman times (III – II century a.C.) was shaped by Cocco et al. [\(1992\)](#page-17-0). The position of the coastline was very close to the present shoreline in the Sinuessa locality (north of Mondragone) while it was about 2 km behind the current shoreline at Castel Volturno locality. The progradation continued through time and in 1600 d.C. the shape of the shoreline mainly coincided with the alignment of coastal watchtowers. From this date to 1809 a coastal progradation, with a mean value of 1 m/ year, characterized the study area (Cocco and Pippo [1988](#page-17-0)).

From 1817 to 1860, the VCP with the exception of the coastal sector between the *Voturno* River to *Agnena* channel was in erosion. This state was probably due to the infill of lowland area (reclamation) that caused a reduction in river sediment supply to the coast, mainly between 1840 and 1850, and to the long shore current that accumulates the sediment on the updrift side of Volturno mouth forming a progradational ridge and swale sequence of the strand plain (Fig. [9](#page-11-0) and Table [3](#page-12-0)).

From 1860 to 1936 the whole shoreline was in accretion with the highest rates characterizing the zones close to the Volturno River mouth (Fig. [9](#page-11-0) and Table [3](#page-12-0)). From 1936 to 1957 the shoreline was still in accretion and only the Volturno River mouth was in erosion with a mean rate of −1.52 m/yr. (Fig. [9](#page-11-0) and Table [3](#page-12-0)). The erosion focused on the delta apex is likely to be the first signal of contraction of the sedimentary load due to the construction of the dam "Ponte Annibale" on the Volturno river in 1952, although the possibility that erosional hot spot is the result of focused effects of episodic wave attack during the later period of observation, cannot be ruled out.

In the beginning of the sixties, human pressure deeply modified the VCP. From 1957 to 1987, the areas close to the Volturno River mouth were in erosion with a mean rate of −3.6 m/yr., while northward the shoreline was in accretion with a mean rate of 1.7 m/yr. (Fig. [9](#page-11-0) and Table [3\)](#page-12-0). During this period, the erosion of the river mouth due to the reduction of sediment yields from the Volturno River is more evident than in the past while the other zones show low rate of accretion. Moreover, the coastline was stable to the north of Patria Lake while shoreline accretion characterized the Regi Lagni zone. This status was probably due to the breakwater built to protect the Pinetamare harbor in the beginning of the seventies. This structureinfluenced the local hydrodynamic regime; the sediment, transported southward by marine currents, stopped on its northern side favoring the accretion of beaches whereas the sediment reduction in the south caused erosion. To reduce erosion, several groins were built at the beginning of eighties, these structures reduced the erosion at Villaggio Pinetamare but shifted this process southtowards.

In only two years, from 1987 to 1989, the coastal zone exposed to erosion was much wider than in previous years and extended from Villaggio Agricolo to Panetelle locality with erosion rates varying from −16 m/yr. to −8 m/yr. (Fig. [9](#page-11-0) and Table [3\)](#page-12-0). This intense erosional phase was probably the result of a sea storm reaching the areas between Castel Volturno and Mondragone in 1987 (Mimmi [1987](#page-18-0)). The coastal sector close to Regi Lagni mouth was in accretion with a mean rate of 16,59 m/yr.

Fig. 8 Urban growth at Volturno Coastal Plain from 1817 to 2012. The municipality boundaries of Castel Volturno and Mondragone are illustrated in the image at the top of figure. The urban growth along the shoreline is shown in the two images at bottom of figure (from Alberico et al. [2017;](#page-16-0) modified)

From 1989 to 1998, the areas close to left and right banks of Volturno River were in erosion with a mean rate of −3.7 m/ yr. and −6.1 m/yr., respectively; while to the north of Volturno River (mean rate 2 m/yr) and close to Regi Lagni mouth (mean rate of 4.8 m/yr) coastal accretion prevails (Fig. [9](#page-11-0) and Table [3\)](#page-12-0). The erosional status of apex and the accretion of the sides is a typical rotational behavior, easily visible along many deltas and barrier systems. The sediment yield is significantly reduced thus inducing a first phase of cannibalization, in the

form of erosion of the more vulnerable parts of the delta (foreland apex). The waves erode the apex while sediments are distributed along the sides, thus the mass balance is maintained.

From 1998 to 2004 the left bank of Volturno River was still in erosion with a mean value of −3.1 m/yr. Southward for a sector of about 2 km long, the shoreline was in accretion with a mean rate of 4.4 m/yr.; lower values characterized the area of Regi Lagni mouth and Pinetamare locality. Southward, an

Fig. 9 The rate of shoreline changes between two shorelines successive in time is illustrated for nine time intervals. The location of transects is shown in the bottom right of the figure

erosional mean rate of −2 m/yr. was active at Ischitella locality. This condition was a consequence of groins built along the coast to the north of this locality. To the north of Volturno River, a 2.5 km long coastal sector was in erosion with a mean rate of −2.1 m/yr. followed by a wide zone, located beyond the Agnena channel, characterized by a mean accretion of 1.9 m/yr. (Fig. 9 and Table [3\)](#page-12-0).

From 2004 to 2010 the left bank of Volturno River was in accretion while all the other sectors, with the exception of a few narrow zones, were in erosion. From 2010 to 2012 the VCP was characterized by coastal sectors in erosion alternated with those in accretion (Fig. 9 and Table [3](#page-12-0)).

We also analyzed the shoreline trends for two time intervals: 1957–1998, 1998–2012 (Figs. [10](#page-14-0)a, b). The first shoreline trend showed a more intensive erosional phase for a wide sector close to the Volturno River mouth, and accretion in the other coastal zones. The second shoreline trend showed sectors in erosion alternating with those in accretion (Figs. [10a](#page-14-0), b). From south to north, the first coastal sector in erosion (mean value: −1.47 m/yr) was located to the north of the Patria Lake, while moving towards the Volturno mouth, three sectors characterized by accretion (mean value: 2.07 m/yr), erosion (mean value: −2.49 m/yr) and accretion (mean value: 2.96 m/yr) were pointed out (Fig. [10a](#page-14-0)). As far as the Volturno River is concerned, the left bank was in erosion (mean value: −1.47 m/yr) while the right one was in accretion (mean value: 1.41 m/yr). Continuing towards the north, the area immediately after the right side of the Volturno River Mouth shows a mean rate of −2.14 (Fig. [10](#page-14-0)b) while the coast is stable from there to *Mondragone*.

Furthermore, the latter trend was used to identify the possible location of the shoreline in 2022, 10 years can be considered a long term to integrate this hazard into territorial management plans (Fig. [11\)](#page-15-0). The erosion rate is similar for all sectors (about -2 m/yr). Only the sector "L" was characterized by higher values and an urban environment very close to the shoreline.

Table 3 Shoreline variations for the nine time intervals investigated. For each interval, several sectors (group of transects) with same shoreline variation (erosion or accretion) were recognized and for each one the mean, maximum and minimum value of erosion and accretion are reported

Table 3 (continued)

Discussion and conclusive remarks

In the frame of supporting the risk reduction, this work illustrates the key role of historical maps, ortophoto maps and free satellite images to collect data on the natural elements of past landscapes, territorial evolution and shoreline changes. The intense urbanization, that reaches the shoreline in several zones, and the exposure to many natural hazards (coastal erosion, marine inundation, fluvial flooding, salt water intrusion, subsidence) make the Volturno Coastal Plain (VCP) a good case study.

The shoreline variations for nine time intervals (1817- 2012 time span) and the coastal trend for 1957–1998 and 1998–2012 time windows were assessed. The latter trend was also assumed as a scenario to identify the possible shoreline position for the year 2022. This data is integral part of the Coastal Management framework (Alberico et al. [2017\)](#page-16-0) which already hosts data (e.g. urban areas, population, municipality boundary, censual district scale, historical documents) and contains procedures to identify the coastal zones where land use changes can be functional to risk reduction. The noticeable flexibility of this framework allows to register many other types of environmental data (e.g. climatic, land-use, soil erosion, human activities) as well as hazard data on both natural (e.g. floods, earthquakes, salt water intrusion) and anthropic (fires, explosion, diffusion of pollutants) events.

In the last centuries, the shoreline evolution at VCP was similar to other coastal areas along the Tyrrhenian and Adriatic coasts. The erosion prevailed at the river mouth while the areas close to the delta were generally in accretion or stable thanks to the redistribution of eroded sediment from the delta apex (Aucelli et al. [2009;](#page-16-0) Bellotti [2000:](#page-16-0) Pranzini [2001](#page-18-0) Alberico et al. [2011](#page-16-0)). During the 1817–1860 time span the coast was in erosion; this condition was probably a consequence of reclamation works realized during the Bourbon domain that caused a reduction of river sediments supply to the coast. In the following years and until the sixties, the coastline of VCP was in progradation; only a narrow zone close to the Volturno River mouth was in erosion, likely due to the reduction of river sediments supply to the coast after the construction of the Ponte Annibale dam on Volturno River. Starting from these years, the construction of 3 new dams, 10 artificial barrages and the extraction of sediments from the Volturno Hydrographic Basin caused the reduction of sediment necessary for beach nourishment in the time interval 1962–1992. Moreover, the construction of 52 sea protection works, the

Fig. 10 The coastal trend for 1957–1998 (a) and 1998–2012 (b) time windows. For the latter time window, the map illustrates the shoreline trend of 12 zones superimposed to the shaded relief of Volturno Coastal Plain. The table reports the maximum, minimum and mean values for the single zone

Fig. 11 The red line illustrates the possible position of coastline for the year 2022 in the four sectors (L, H, F, B) (for location see also Fig. [10](#page-14-0)). The map on the bottom right side of figure illustrates the location of these sectors along the Volturno Coastal Plain while the table shows the rate and

the statistic parameters showing the accuracy of the model. For the sector "N" also in erosion, the available data does not allow the identification of a trend value

urban environment growth and the destruction and degradation of dune vegetation favored the increase of instability conditions in the VCP.

The coastal dynamics changed along the investigated period; the delta recorded the decrease of sediment supply during the progressive changes from progradation, stabilization toward initial disequilibrium (rotational behavior) to the advanced disequilibrium (small-scale cannibalization). As reported in Fig. [9](#page-11-0), there is a clear shift in sediment budget: all the pre-II World War conditions up to the post-II World War time (1957) indicated a clear accretionary behavior, linked to the large amount of sediment supplied by the river. From 1957 to 1998 interval, a classical rotational behavior of both sides of the delta, possibly due to intervention in the Volturno basin occurred (i.e. construction of dams and barrages and sediments quarrying from 1962 to 1992). The critical time of radical change of the erosional/depositional style is represented by the alternating erosional/depositional cells in the more recent time span 1998–2012. This could be considered an example of short-scale cannibalization, occurring when a deltaic coastline is sediment starved (Simeoni et al. [2007](#page-18-0)).

The erosional rate, assessed from 1998 to 2012, was assumed as a reference value to draw a reasonable location of the coastine for the year 2022. The overlay of this shoreline on the Technical Maps of Campania Region allowed the identification of urban areas possibly exposed to erosion in the future. In detail, the rate of −2 m/yr, typifying three (sectors: L, H, F) out of four sectors in erosion, could expose to damages several buildings close to the coast (Fig. [11\)](#page-15-0).

The rapid growth of the urban environment without rules in the sixties (Frallicciardi and Cerisano [2013](#page-17-0)), lead to the present-day territorial status: a northern sector characterized by a wide urbanized environment (1,5 km depth from the coast) close to the shoreline, separated by Mondragone urban center by a narrow sector that still preserves natural features; a southern sector with a Natural Reserve (Variconi Oasis) and urbanized environment, except for Villaggio Coppola, far from the shoreline. For these areas, according to Alberico et al. (2017) a risk reduction through the natural environment restoration could be functional because the greatest part of urban environment close to the coast is inhabited all year long. In this frame, the past land-use (the Mediterranean maquis, pinewood, belts of psammophyte and halophyte vegetation, wetlands areas and ponds) was considered a good guide to identify the measures to make this environment resilient.

The flexibility of the system here proposed also ensures the possibility to archive shorelines from a monitoring activity making it possible to analyze the shoreline with a detailed spatial and temporal scale at a fixed time interval.

Acknowledgements Financial support for this study has been provided by Project PON03 I-AMICA (High technology infrastructure for the Climate- Environmental Integrated Monitoring; <http://www.i-amica.it>), Development Objective OR4.4 - Study and monitoring of the interface processes between the biosphere, hydrosphere and functionality of coastal ecosystems and RITMARE Flagship Project, funded by MIUR (NRP 2011–2013). The suggestions of two anonymous reviewers, whom the authors gratefully appreciate, greatly improved an early version of the manuscript.

References

- Abuodha PAO, Woodroffe CD (2010) Assessing vulnerability to sealevel rise using a coastal sensitivity index: a case study from southeast Australia. J Coast Conserv 14:189–205. [https://doi.org/10.](https://doi.org/10.1007/s11852-010-0097-0) [1007/s11852-010-0097-0](https://doi.org/10.1007/s11852-010-0097-0)
- Addo KA, Walkden M, Mills LP (2008) Detection, measurement and prediction of shoreline recession in Accra, Ghana. ISPRS J Photogramm Remote Sens 63:543–558
- Afan de Rivera C (1847) Memoria intorno al bonificamento del bacino inferiore del Volturno. Stamperia del Fibreno, Napoli. Harvard's library collections
- Aiello A, Canora F, Pasquariello G, Spilotro G (2013) Shoreline variations and coastal dynamics: a space time data analysis of the Jonian littoral, Italy. Estuar Coast Shelf Sci 129:124–135
- Alberico I, Amato V, Aucelli PPC, D'Argenio B, Di Paola G, Pappone G (2011) Historical shoreline evolution and recent shoreline trends of Sele plain coastline (southern Italy). The 1870–2009 time window. J Coast Res 28(6):1638–1647. [https://doi.org/10.2112/](https://doi.org/10.2112/%20JCOASTRES-D-10-00197) [JCOASTRES-D-10-00197](https://doi.org/10.2112/%20JCOASTRES-D-10-00197)
- Alberico I, Iavarone R, Angrisani AC, Castiello A, Incarnato R, Barrra R (2017) The vulnerability indices as tools for natural risk reduction. The Volturno coastal plain case study. J Coast Conserv. [https://doi.](https://doi.org/10.1007/s11852-017-0534-4) [org/10.1007/s11852-017-0534-4](https://doi.org/10.1007/s11852-017-0534-4)
- Albuquerque M, Espinoza J, Teixeira P, de Oliveira A, Corrêa I, Calliari L (2013) Erosion or coastal variability: an evaluation of the DSAS and the change polygon methods for the determination of erosive processes on sandy beaches. In: Conley DC, Masselink G, Russell PE, O'Hare TJ (ed) proceedings 12th international coastal symposium (Plymouth, England). Journal of coastal research 65:1710-1714
- Amorosi A, Pacifico A, Rossi V, Ruberti D (2012) Late quaternary incision and deposition in an active volcanic setting: the Volturno Valley fill, southern Italy. Sediment Geol 282:307–320. [https://doi.org/10.](https://doi.org/10.1016/j.sedgeo.2012.10.003) [1016/j.sedgeo.2012.10.003](https://doi.org/10.1016/j.sedgeo.2012.10.003)
- Aprile F, Ortolani F (1978) Nuovi dati sulla struttura profonda della Piana Campana. Boll Soc Geol It 97:591–608
- Aucelli PPC, Iannantuono E, Rosskopf CM (2009) Evoluzione recente e rischio di erosione della costa molisana (Italia meridionale). Ital J Geosci (Boll Soc Geol It) 128(3):759–771
- Barra D, Bonaduce G, Brancaccio L, Cinque A, Ortolani F, Pagliuca S, Russo F (1989) Holocene geological evolution of Sarno river coastal plain (Campania). Mem Soc Geol Ital 42:255–267
- Becek K, Ibrahim K, Darussalam B (2011a) On the positional accuracy of the GoogleEarth® imagery. In: FIG working week (bridging the gap between cultures), Marrakech, Morocco, 18−22 may 2011
- Kazimierz Becek, Khairunnisa Ibrahim, Brunei Darussalam (2011b) On the positional accuracy of the GoogleEarth® imagery. Fig Working Week (Bridging the Gap between Cultures), Marrakech, Morocco, 18–22 May 2011
- Bellotti P (2000) Il modello morfo-sedimentario dei maggiori delta tirrenici italiani. Boll. Soc. Geol. It. 119:777–792
- Bruno PPG, Di Fiore V, Ventura G (2000) Seismic study of the '41st parallel' fault system offshore the Campanian– Latial continental margin, Italy. Tectonophysics 324(1):37–55
- Büntgen U, Myglan VS, Charpentier Ljungqvist F, McCormick M, Di Cosmo N, Sigl M, Jungclaus J, Wagner S, Krusic PJ, Esper J, Kaplan JO, de Vaan MAC, Luterbacher J, Wacker L, Tegel W, Kirdyanov AV (2016) Cooling and societal change during the late antique little ice age from 536 to around 660 AD. Nat Geosci 9:231– 236. <https://doi.org/10.1038/ngeo2652>

Canino A (1981) Guida d'Italia: Campania. Touring club italiano, Milano

- Casella V, Franzini M, Padova B, Spalla A, Bruneo MG (2010) L'accuratezza della georeferenziazione delle immagini Google Earth: un caso-studio sul Comune di Pavia. Atti 14a Conferenza Nazionale ASITA, Brescia, 9–12 novembre 2010
- Caserta Provincial Territorial Plan (2012) [ftp://ftp.provincia.caserta.it/](ftp://ftp.provincia.caserta.it/pub/Ptc%20Caserta/PTCP) [pub/Ptc%20Caserta/PTCP/](ftp://ftp.provincia.caserta.it/pub/Ptc%20Caserta/PTCP)
- Ciasca R (1928) Storia delle bonifiche del Regno di Napoli. Laterza, Bari Cinque A, Alinaghi HH, Laureti L, Russo F (1987) Osservazioni preliminari sull'evoluzione geomorfologica della Piana del Sarno (Campania, Appennino meridionale). Geogr Fis Din Quat 10:161– 174
- Cinque A, Aucelli PPC, Brancaccio L, Mele R, Milia A, Robustelli G, Romano P, Russo F, Santangelo N, Sgambati D (1997) Volcanism, tectonics and recent geomorphological change in the bay of Napoli. Supplementi Geografia Fisica Dinamica Quaternaria 3:123–141
- Cinque A, Ascione A, Caiazzo C (2000) Distribuzione spazio-temporale e caratterizzazione della fagliazione quaternaria in Appennino meridionale. In: Galadini F, Meletti C, Rebez A (eds) Le ricerche del GNDT nel campo della pericolosità sismica 1996–1999. Roma, Italy, pp 203–218
- Cocco E, Iuliano S (1999) L'erosione del litorale in sinistra foce Sele (Golfo di Salerno): dinamica evolutiva e proposta di intervento a difesa e tutela della spiaggia e della pineta litoranea di Paestum. Il Quaternario 12(2):125–140
- Cocco E and De Magistris MA (1988) Evoluzione storica e recente del litorale di Paestum (Golfo di Salerno). Mem Soc Geol Ital 41:697– 702
- Cocco E and De Pippo T (1988) Tendenze evolutive e dinamiche delle spiagge della Campania e della Lucania. Mem Soc Geol Ital 41:195– 204
- Cocco E, Crimaco L, de Magistris MA (1992) Dinamica ed evoluzione del litorale campano-laziale: variazioni della linea di di riva dall'epoca romana ad oggi nel tratto compreso tra la foce del Volturno e Torre S. Limato (Mondragone). Atti 10° congresso AIOL 1992, 543-555
- Crowell M, Letherman SP, Buckley MK (1991) Historical shoreline change: error analysis and mapping accuracy. J Coast Res 7(3): 839–852
- De Vivo B, Rolandi G, Gans PB, Calvert A, Bohrson WA, Spera FJ, Belkin HE (2001) New constraints on the pyroclastic eruptive history of the Campanian volcanic plain (Italy). Mineral Petrol 73:47– 65
- Dickson ME, Walkden MJ, Hall JW (2007) Systemic impacts of climate change on an eroding coastal region over the twenty-first century. Clim Chang 84(2):141–166
- Doglioni C (1991) A proposal of kinematic modelling for W-dipping subductions possible applications to the Tyrrhenian-Apennines system. Terra Nova 3:423–434
- Dolan R, Hayden B, May P, Suzette R (1980) The reliability of shoreline change measurements from aerial photographs. Shore and Beach 48: 22–29
- Dolan R, Fester MS, Holme SJ (1991) Temporal analysis of shoreline recession and accretion. J Coast Res 7(3):723–744
- EEA Report (2006) The changing faces of Europe's coastal areas No 6/2006. ISSN 1725-9177
- EUROSTAT (1988). <http://ec.europa.eu/eurostat>
- Ferranti L, Oldow JS (1999) History and tectonic implications of lowangle detachment faults and orogen parallel extension, Picentini Mountains, southern Apennines fold and thrust belt, Italy. Tectonics 18:498–526
- Frallicciardi AM and Cerisano F (2013) Vecchie e nuove questioni di bonifica nella Piana del Volturno. In: Società Geografica Italiana (ed) Trasformazioni del paesaggio in aree di bonifica dell'Italia centro-meridionale in epoca post-unitaria, Roma, pp 117–142
- Genz AS, Fletcher CH, Dunn RA, Frazer LN, Rooney JJ (2007) The predictive accuracy of shoreline change rate methods and alongshore beach variation on Maui, Hawaii. J Coast Res 23:87–105
- Gibeaut JC, Waldinger R, Hepner T, Trembla TA, White WA, Xu L (2003) Changes in bay shoreline position, west bay system, Texas. Report of the Texas coastal coordination council pursuant to national oceanic and atmospheric administration award no. NA07OZ0134 GLO contract number 02-225R
- Hapke CJ, Reid D, Richmond BM, Ruggiero PJ (2006) List, National assessment of shoreline change part 3: Historical shoreline change and associated coastal land loss along sandy shorelines of the California coast. U.S. Geological Survey Open-file Report 2006– 1219
- Hydro-geomorphological Setting Plan of Basin Authority of Liri-Garigliano-Volturno (2010) [http://www.autoritadibacino.it/index.](http://www.autoritadibacino.it/index.php?option=com_content&view=article&id=139&Itemid=379Accessed) [php?option=com_content&view=article&id=139&Itemid=](http://www.autoritadibacino.it/index.php?option=com_content&view=article&id=139&Itemid=379Accessed) [379Accessed](http://www.autoritadibacino.it/index.php?option=com_content&view=article&id=139&Itemid=379Accessed) 28 July 2016
- Iacono R, Napolitano E, Marullo S, Artale V (2013) Seasonal Variability of the Tyrrhenian Sea Surface Geostrophic Circulation as Assessed by Altimeter Data. J Phys Oceanogr 43:1710–1732
- Iermano I, Liguori G, Iudicone D, Buongiorno Nardelli B, Colella S, Zingone A, Saggiomo V, Ribera d'Alcalà M (2012) Filament formation and evolution in buoyant coastal waters: observation and modelling. Prog Oceanogr 106:118–137
- Iorio M, Capretto G, Petruccione E, Marsella E, Aiello G, Senatore MR (2014) Multi-proxy analysis in defining sedimentary processes in very recent prodelta deposits: the northern Phlegraean offshore example (eastern Tyrrhenian margin) rend. Fis Acc Lincei. [https://doi.](https://doi.org/10.1007/s12210-014-0303-3) [org/10.1007/s12210-014-0303-3](https://doi.org/10.1007/s12210-014-0303-3)
- IPCC (2014) Climate change 2014: synthesis report. Contribution of working groups I, II and III to the fifth assessment report of the intergovernmental panel on climate change [Core writing team, Pachauri RK and Meyer LA (ed)]. IPCC, Geneva, Switzerland, 151 pp
- Ippolito F, Ortolani F, Russo F (1973) Struttura marginale tirrenica dell'Appennino Campano: reinterpretazione di dati di antiche ricerche di idrocarburi. Mem Soc Geol It 12:227–250
- ISPRA Servizio Geologico d'Italia (1966) Carta Geologica d'Italia in scala 1:100.000 - Foglio n.172 "Caserta". Litografia artistica cartografica, Firenze
- ISPRA Servizio Geologico d'Italia (1967) Carta Geologica d'Italia in scala 1:100.000 - Foglio n.184 "Isola d'Ischia-Napoli". Litografia artistica cartografica, Firenze
- ISPRA Servizio Geologico d'Italia (1968) Carta Geologica d'Italia in scala 1:100.000 - Foglio n.171 "Gaeta". Istituto italiano d'arti grafiche, Bergamo
- JAM d P, Anfuso G (2008) Spatial approach to medium-term coastal evolution in south Sicily (Italy): implications for coastal erosion management. J Coast Res 24(1):33–42
- Jimenez JA, Sanchez-Arcilla A, Bou J, Ortiz MA (1997) Analyzing short-term shoreline changes along the Ebro Delta (Spain) using aerial photographs. J Coast Res 13(4):1256–1266
- Kaliraj S, Chandrasekar N, Magesh NS (2015) Evaluation of coastal erosion and accretion process along the southwest coast of Kanyakumari, Tamil Nadu using geospatial techniques. Arab J Geosci 8(1):239–253
- Kuleli T, Guneroglu A, Karsli F, Dihkan M (2011) Automatic detection of shoreline change on coastal Ramsar wetlands of Turkey. Ocean Eng 38:1141–1149
- Kumar A, Jayappa KS, Deepika B (2010) Application of remote sensing and GIS in change detection of the Netravati and Gurpur river channels, Karnataka, India. Geocarto Int 25:397–425
- Leatherman S (1983) Shoreline mapping: a comparison of techniques. Shore and Beach 51:28–33
- Leyland J, Darbya SE (2008) An empirical–conceptual gully evolution model for channeled sea cliffs. Geomorphology 102:419–434
- Margaritelli G, Vallefuoco M, Di Rita F, Capotondi L, Bellucci LG, Insinga DD, Petrosino P, Bonomo S, Cachof I, Cascella A, Ferraro L, Florindo F, Lubritto C, Lurcock PC, Magri D, Pelosi N, Rettori R, Lirer F (2016) Marine response to climate changes during the last five millennia in the central Mediterranean Sea. Glob Planet Chang 142:53–72
- Mariani M, Prato R (1988) I bacini neogenici costieri del margine tirrenico: approccio sismico-stratigrafico. Mem Soc Geol Ital 41: 519–531
- Millennium Ecosystem Assessment (2005) Ecosystems and human wellbeing: Synthesis. Island Press, Washington DC, p 155
- Milia A, Torrente MM (1999) Tectonics and stratigraphic architecture of a pery-Tyrrhenian half-graben (bay of Naples, Italy). Tectonophysics 315:301–318
- Milia A, Torrente MM (2003) Late quaternary volcanism and transtensional tectonics in the Naples Bay, Campanian continental margin. Italy Miner Petrol 79:49–65
- Milia A, Torrente MM, Massa B, Iannace P (2013) Progressive changes in rifting directions in the Campania margin (Italy): new constrains for the Tyrrhenian Sea opening. Glob Planet Change 109:3–17
- Mimmi V (1987) Il maltempo devasta il sud danni per decine di miliardi. [http://ricerca.repubblica.it/repubblica/archivio/repubblica/1987/01/](http://ricerca.repubblica.it/repubblica/archivio/repubblica/1987/01/13/il-maltempo-devasta-il-sud-danni-per.html) [13/il-maltempo-devasta-il-sud-danni-per.html](http://ricerca.repubblica.it/repubblica/archivio/repubblica/1987/01/13/il-maltempo-devasta-il-sud-danni-per.html). Accessed 10 October 2017
- Moore L (2000) Shoreline mapping techniques. J Coast Res 16(1):111– 124
- Morton RA, Miller TL, Moore LJ (2004) National assessment of shoreline change part 1: historical shoreline changes and associated coastal land loss along the US Gulf of Mexico. US Geol Surv Open File Rep:2004–1043
- Nguyen TTX, Woodroffe CD (2016) Assessing relative vulnerability to sea-level rise in the western part of the Mekong River Delta in Vietnam. Sustain Sci 11:645–659. [https://doi.org/10.1007/s11625-](https://doi.org/10.1007/s11625-015-0336-2) [015-0336-2](https://doi.org/10.1007/s11625-015-0336-2)
- Ortolani F, Pagliuca S (1986) Relazioni tra struttura profonda ed aspetti morfologici e strutturali della fascia tirrenica dell'Appennino campano. Atti Riun. Ann. Gr. Naz. Geogr. Fis. e Geomorj, Amalfi (SA)
- Ortolani F, Pagliuca S (1999) Evoluzione geologica e geomorfologica olocenica e problemi geoambientali delle pianure costiere antropizzate della Campania. Giorn Geol 61:235–237
- Pajak MJ, Leatherman S (2002) The high water line as shoreline indicator. J Coast Res 18(2):329–337
- Patacca E, Scandone P (2007) Geology of the southern Apennines. In: Mazzotti A, Patacca E, Scandone P (eds) Results of the CROP project sub-project CROP-04 southern Apennines (Italy).

Bollettino della Società Geologica Italiana (Italian journal of geoscience) special issue 7, pp 75–119

- Pennetta M, Valente A, Abate D, Boudillon G, De pippo T, Leone M, Terlizzi F (1998) Influenza della morfologia costiera sulla circolazione sedimentazione sulla piattaforma continentale Campano-Laziale tra Gaeta e Cuma (Italia Meridionale). Boll Soc Geol It 117:281–295
- Potere D (2008) Horizontal positional accuracy of Google Earth's highresolution imagery archive. Sensors 8:7973–7981
- Pranzini E (2001) Updrift river mouth migration on cuspate deltas: two examples from the coast of Tuscany (Italy). Geomorphology 38: 125–132
- Rinaldi E, Buongiorno Nardelli B, Zambianchi E, Santoleri R, Poulain PM (2010) Lagrangian and Eulerian observations of the surface circulation in the 1175 Tyrrhenian Sea. J Geophys Res 115, [C04024](https://doi.org/C04024). <https://doi.org/10.1029/2009JC005535>
- Romagnoli C, Mancini F, Brunelli R (2006) Historical shoreline changes at an active island volcano: Stromboli, Italy. J Coast Res 22(4):739– 749
- Romano P, Santo A, Voltaggio M (1994) Geomorphological evolution of Volturno river plain during late quaternary (middle-late Pleistocene– Holocene). Il. Quaternario 7:41–56
- Ruberti D, Vigliotti M, Di Natale M, Ivaldi R (2014) Recent evolution of a delta plain and a coastal zone: the Volturno delta system (southern Italy). Seconda Università degli Studi di Napoli; Giornate Scientifiche di Ateneo 2014
- Sacchi M, Molisso F, Pacifico A, Vigliotti M, Sabbarese C, Ruberti D (2014) Late-Holocene to recent evolution of Lake patria, South Italy: an example of a coastal lagoon within a Mediterranean Delta system. Glob Planet Chang 117:9–27
- Savarese G (1856) Bonificamento del Bacino Inferiore del Volturno. Dalla Stamperia Reale, Napoli
- Simeoni U, Fontolan G, Tessari U, Corbau C (2007) Domains of spit evolution in the Goro area, Po delta, Italy. Geomorphology 86: 332–348
- SOER (2015) The European environment-state and outlook 2015. A comprehensive assessment of the European environment's state, trends and prospects, in a global context, European Environment Agency. <http://www.eea.europa.eu/soer>
- Sutton-Grier AE, Wowk K, Bamford H (2015) Future of our coasts: the potential for natural and hybrid infrastructure to enhance the resilience of our coastal communities, economies and ecosystems. Environ Sci Policy 51:137–148
- Thieler ER, Danforth WW (1994) Historical shoreline mapping (II): application of the digital shoreline mapping and analysis systems (DSMS/DSAS) to shoreline change mapping in Puerto Rico. J Coast Res 10:600–620
- Thieler ER, Himmelstoss EA, Zichichi JL, Ergul A (2009) The digital shoreline analysis system (DSAS) version 4.0 - an ArcGIS extension for calculating shoreline change. Open-file report. US geological survey report no. In: 2008–1278 [http://woodshole.er.usgs.gov/](http://woodshole.er.usgs.gov/project-pages/dsas/version4/) [project-pages/dsas/version4/](http://woodshole.er.usgs.gov/project-pages/dsas/version4/)
- UNEP/MAP (2009) State of the environment and development in the Mediterranean. URL: [http://www.planbleu.org/actualite/uk/](http://www.planbleu.org/actualite/uk/soed2009_Uk.html) [soed2009_Uk.html](http://www.planbleu.org/actualite/uk/soed2009_Uk.html)