RESEARCH ARTICLE - SOLID EARTH SCIENCES

Seasonal components in GPS displacement time series

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Abstract

We present a simple method for identifcation of seasonal components in GPS displacement time series. The method considers two options for the treatment of long series, considering possible variability in the seasonal phases. The method also proposes a bootstrap scheme to deal approximately with series shorter than one period of the seasonal signal. Tests with synthetic data show that the method gives quite acceptable results, so that results from application to real data, although approximate, should be useful considering the inherent uncertainties in the data.

Keywords GPS · GNSS · Seasonal components · Volcano monitoring

Introduction

Global Positioning System (GPS) measurements are an extremely valuable source of data for a wide range of geophysical studies, many of which had to rely on imprecise or difficult to obtain and sparse data until not very long ago. GPS data is useful for studies ranging from diferent felds of tectonics (e.g. Blewitt and Lavalée [2002](#page-6-0); Davis et al. [2006](#page-6-1); Hefin et al. [2020;](#page-6-2) Ji et al. [2020\)](#page-6-3) to seismology (e.g. Davis et al. [2006](#page-6-1); Wernicke and Davis [2010;](#page-6-4) Gonzalez-Ortega et al. [2014\)](#page-6-5). The present paper deals with application of GPS to studies in volcanology (e.g. Mann et al. [2002](#page-6-6); Bartel et al. [2003;](#page-6-7) Miyagi et al. [2004;](#page-6-8) Puglisi and Bonforte [2004](#page-6-9); Janssen [2007](#page-6-10); Geirsson et al. [2017;](#page-6-11) Larson et al. [2010;](#page-6-12) Lee et al. [2015\)](#page-6-13), where GPS data is used to monitor deformations in volcanic edifces, because these deformations can be related to magma of gas-related processes going on underground, so that GPS measurements are useful for studying

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these volcanic processes and sometimes to anticipate possible eruptive activity.

The subject of the present study is to identify very large periodic (or semi-periodic) components due to seasonal signals (e.g., Dong et al. [2002;](#page-6-14) Ray et al. [2008](#page-6-15); Tregoning and Watson [2009](#page-6-16); Chen et al. [2013;](#page-6-17) He et al. [2017](#page-6-18); Wang et al. [2018](#page-6-19)), the causes of which can be many, e.g., Chanard et al. [2020](#page-6-20), and are outside the scope of this paper.

Because seasonal components present large amplitudes with respect to the amplitudes of background volcanic processes, many methods to identify seasonal components have been proposed (e.g., Davis et al. [2006,](#page-6-1) [2012;](#page-6-21) Freymueller [2009](#page-6-22); Chen et al. [2013,](#page-6-17) [2020;](#page-6-23) Gruszczynska et al. [2017](#page-6-24); He et al. [2017](#page-6-18); Klos et al. [2019](#page-6-25); Ji et al. [2020;](#page-6-3) and many others).

In the present paper, we contribute a simple-minded, intuitive approach to seasonal component identifcation. The analysis program we present identifes each seasonal component, and if desired, outputs a series consisting of the original one minus the seasonal component. If the seasonal component is the object of interest, it can be easily obtained by subtracting the resulting series from the original one.

Method

In what follows, we present a method for identifying seasonal components, whether annual, semiannual, draconitic, etc. We will use a simple method to identify a periodic component within a time series, based on the wonderful capacity

of Fourier analysis to identify a component with a given period, even in the presence of other periods and noise.

Let $d = \{d_i; i = 1, 2, ..., N\}$ be a time series having *N* elements sampled with sampling interval Δt , so that element d_i corresponds to time $t_i = t_1 + (i - 1)\Delta t$, where t_1 is the time of the frst element.

Let the period of the seasonal component to be identifed, and perhaps later eliminated, be *T*, so there will be $N_T = \mathcal{R}(T/\Delta t)$ elements in one period, where \mathcal{R} is the "rounding" function that assigns the integer closest to the real number argument. Hence, we will be really identifying the period $N_T\Delta t$, which we will suppose to be close enough to *T* for practical purposes.

To eliminate the seasonal component two cases are considered. Case one is when series are at least one period long, $N \geq N_T$. It is well known that Fourier analysis can correctly characterize only periods that are submultiples of the total observation time, and other periods present in the sampled series are represented in terms of various submultiple periods. Thus, except for the case $N = kN_T$, where *k* is some positive integer, straightforward application of the Fourier transform to the whole series is not appropriate to determine amplitude and phase of a seasonal component with period *T*.

From a series with *N* elements, $m = N - N_T + 1$ different continuous segments one period long can be chosen, $d^{(j)} = \{d_j, d_{j+1}, \dots, d_{j+N_T-1}; j = 1, \dots, m\}$, and the Fourier transform of each of these series:

$$
D_1^{(j)} = \sum_{q=j}^{j+N_T-1} e^{-i2\pi f_o(q-j)\Delta t} d_q \Delta t,
$$
 (1)

(only one spectral component, the fundamental, is needed) yields an amplitude $A^{(j)}$ and a phase $\psi^{(j)}$ measured with reference to t_i . This phase is changed to a phase $\phi^{(i)}$ referred to the time of the first element of the series t_1 as $\phi^{(j)} = \psi^{(j)} - 2\pi f_o(t_j - t_1).$

Transforms from diferent segments, as many as possible, are used because, as will be discussed below, seasonal signals are usually not invariant over time, which means the amplitude and phase determined from one segment may not be representative of the whole series.

Once amplitudes and phases have been computed, a frst option consists of using the mean values of the *m* amplitudes and phases, A and ϕ , respectively, to define an average seasonal component as:

$$
s_i = \overline{A} + \cos\left[2\pi f_o(t_i - t_1) + \overline{\phi}\right],\tag{2}
$$

which gives good results for non-varying seasonal signals and yields estimates of the spread in \vec{A} and ϕ through their standard deviations. Figure [1](#page-1-0) is an example of this option applied to a segment of a GPS displacement series from Colima volcano.

Fig. 1 A Displacement GPS time series recorded at Colima volcano (blue) and seasonal component from the average phase and amplitude option (red). **B** Resulting displacement after elimination of the seasonal component (black line) and original displacement (thin blue line)

However, it has been observed that the seasonal signal in geodetic time series is not invariant (Tregoning and Watson [2009;](#page-6-16) Davis et al. [2012;](#page-6-21) Chen et al. [2013;](#page-6-17) Ji et al. [2020](#page-6-3); Tucikešić et al. [2020;](#page-6-26) Ruttner [2021\)](#page-6-27), and diferent methods for characterizing the seasonal components are proposed in the papers referenced above.

Our approach to varying seasonal components is based on Eqs. [\(2](#page-1-1)) and ([4\)](#page-1-2) of Davis et al. ([2012\)](#page-6-21), who observe that a small segment of the time signal can be considered as belonging to a seasonal sinusoid if the amplitude does not change too rapidly over the period T , in which case the seasonal *s*(*t*) can be written as:

$$
s(t) = A(t)\cos\theta(t) = A(t)\cos\left[2\pi f_o t + \phi(t)\right],\tag{3}
$$

where $f_o = T^{-1}$, $A(t)$ and $\theta(t)$ are instantaneous amplitude and total phase, respectively, and $\phi(t)$ is the instantaneous phase offset, so that a variation in the instantaneous frequency $f(t)$ implied in Eq. [\(1](#page-1-3)) can be expressed through the variation in $\phi(t)$ as:

$$
f(t) = f_o + \frac{1}{2\pi} \frac{\mathrm{d}\phi(t)}{\mathrm{d}t},\tag{4}
$$

keeping f_o constant.

d*𝜙*(*t*)

Thus, for each of the $d^{(j)}$ series, instantaneous amplitudes $A^{(j)}$ and $\phi^{(j)}$ are determined from the Fourier transform as mentioned above. Since each pair of these values is representative of the whole segment, the values are assigned to the middle element of the series, i.e., *A*_{*j*+*N_T*/2−*k* = *A*^(*j*) and ϕ _{*j*+*N_T*/2−*k* = ϕ ^(*j*), where *k* = 0 for}} even N_T and $k = 1$ for odd N_T , and we have two series of instantaneous parameters, A and ϕ , to substitute in (3). The $N_T/2$ elements at each extreme are "tails" of the method and can either be discarded for series that are long enough so that information from the tails is not needed for interpretation. For short or middle sized series, with lengths only a few times the tail length, the tails can be assigned the closest values, i.e., $A^{(1)}$ and $\phi^{(1)}$ at the beginning, and $A^{(m)}$ and $\phi^{(m)}$ at the end.

The seasonal component is then computed as:

$$
s_i = A_i + \cos[2\pi f_o(t_i - t_1) + \phi_i].
$$
 (5)

Figure [2](#page-2-0) shows an example of the application of this method to a long segment of GPS displacement data from the Colima Volcano, and Fig. [3](#page-2-1) shows the instantaneous phases and amplitudes.

Case two is for series that are shorter than one period $N < N_T$. It is not feasible to characterize a signal with period longer than the total duration from direct application of Fourier analysis, but we propose a bootstrap

> $\overline{\mathbf{x}}$ Ω

> > 2011.5

2012

2012.5

2013

2013.5

 t (year)

2014

2014.5

2015

Fig. 2 A Displacement GPS time series recorded at Colima volcano (blue) and seasonal component from varying phase method (red); dotted lines indicate the tails of the ft. **B** Resulting displacement after elimination of the seasonal component (black line) and original displacement (thin blue line)

Fig. 3 Blue lines are **A** instantaneous phase and **B** amplitude for the seasonal fit shown in Fig. [2.](#page-2-0) Horizontal red lines indicate the mean values and dotted lines are the mean \pm one standard deviation

method to get an approximate characterization, based on the fact that the period of interest is already known (postulated).

First, we compute an incomplete transform with $K = N$ as in (6) :

$$
D_1 = \sum_{q=1}^{K} e^{-i2\pi f_o(q-1)\Delta t} d_q \Delta t, \qquad (6)
$$

which results in approximate amplitude $A^{[0]}$ and phase $\phi^{[0]}$, where [0] indicates that it is the zero'th iteration.

Next, the "missing" data are computed as:

$$
d_i = A^{[y]} + \cos\left[2\pi f_o(t_i - t_1) + \phi^{[y]}\right]; i = N + 1, \dots, N_T, (7)
$$

for $y = 0$. These extrapolated data are appended to the original series, and a new "complete" transform is calculated using (6) with $K = N_T$ to obtain $A^{[1]}$ and $\phi^{[1]}$.

The process is iterated, replacing at each iteration the previously interpolated data by the new ones, until for some

2015.5

y iteration, $A^{[y]} - A^{[y-1]} < \epsilon_A$ and $\phi^{[y]} - \phi^{[y-1]} < \epsilon_A$, where ϵ_A and ϵ_{ϕ} are convergence criteria (we use $\epsilon_A = 0.001$ mm and $\epsilon_{\phi} = 0.001^{\circ}$ or a maximum number of iterations is reached (we use $y_{\text{max}} = 2000$, but for all cases having a reasonable length, estimates converge after at most a few tens of iterations).

After convergence, the seasonal component is computed from:

$$
s_i = A^{[y]} + \cos\left[2\pi f_o(t_i - t_1) + \phi^{[y]}\right]; i = 1, ...N,
$$
\n(8)

and subtracted from the original series.

Figure [4](#page-3-0) shows an example of this bootstrap ftting to the frst 300 elements of the signal shown in Fig. [1](#page-1-0). A comparison of both the ftted seasonal component and the resulting series shows that our approximate iterative ftting works quite well.

Synthetics and discussion

To keep this paper at a reasonable length, only a few examples of application to real data were shown here, but, although results look nice, it is not possible to judge from real data whether seasonal components were correctly identifed, since it is not known a priori which, if any, are present in the data. It is necessary to test the proposed method

with synthetic signals that consist of, or incorporate, known components.

The method was implemented as a MATLAB program, diseas.m, which can be found in [http://cicese.repositori](http://cicese.repositorioinstitucional.mx/jspui/handle/1007/3925) [oinstitucional.mx/jspui/handle/1007/3925](http://cicese.repositorioinstitucional.mx/jspui/handle/1007/3925). Also found at the same address is a complementary program, syntseas.m, which generates synthetic series with one or several diferent periods and corresponding original phases, and the option of varying the phase in time for each period, by adding (or subtracting) a quantity that changes in time as a sine function with a chosen period. The synthetics program can also add to the seasonals a trend, and add random noise, either normally or uniformly distributed, with a chosen standard deviation. Evaluating the efects of a trend in the data is important because it represents a signal with longer "period" than the duration of the time series and, being the double integral of a Dirac delta function, has a Fourier spectrum ∝ *f* [−]2, which means that it can be an important factor in the treatment of short series.

Instructions for the use of the abovementioned programs are included in the corresponding headers, and the programs themselves are user-friendly. Thus, interested readers can use the attached programs to reproduce the results we will now show, and get fgures showing how seasonal components are ftted.

First, we will show a compilation of results from seasonal identification in long series Fig. 5 shows ϵ , the rms fit error

Fig. 4 A Short segment of the displacement GPS time series recorded at Colima volcano shown in Fig. [1](#page-1-0) (blue) and seasonal component from boot-strapping method (red). The yellow lines are the fts for the 9 iterations needed to attain convergence. **B** Resulting displacement after elimination of the seasonal component. Fit and resulting series can be compared with the equivalent segment in Fig. [1](#page-1-0)

Fig. 5 Fit rms error ϵ between true and estimated yearly seasonals for diferent series lengths *N*. Blue lines are the total error for the whole series and dashed red lines are the error excluding the tails

between true and estimated seasonals, for a yearly seasonal with $T = 1$ year, amplitude $A = 5.0$ mm, and initial phase $\phi_0 = 33.0^\circ$, for variousserieslengths*N*. The first seasonal long series model, L0, a sine curve without noise, phase variations, or trend, is not illustrated because the identifcation is exact and errors are zero up to at least six decimal places.

Seasonal model L1 (Fig. [5A](#page-3-1)) illustrates the effect of adding normally distributed random noise with 0.5 mm standard deviation to L0 (random number generator seed=88). The error corresponding to the whole series is plotted as a continuous blue line, and the red dashed line is the error without including the tails. Error increases as the series grows longer, i. e., as more noise is added, but it is always very small, smaller than the noise standard deviation.

Seasonal model L2 (Fig. [5B](#page-3-1)) is L0 with added trend of 1.0mm/yr. The error is null outside the tails and the total error decreases as the series grows and tails become smaller in relation to the total length.

Seasonal model L3 (Fig. [5](#page-3-1)-C) shows the effect of adding varying phase to L0 as:

$$
\phi = \phi_0 + \Delta \phi \sin \frac{2\pi}{\theta} (t - t_1),\tag{9}
$$

where $\Delta \phi$ is the maximum increment (or decrement if negative) that is attained at time $\theta/4$ after t_1 . The example features $\Delta \phi = 20^{\circ}$ and $\theta = 5$ yr. The error for no tails is quite small, and the error with tails decreases as tails get relatively smaller.

Finally, model L4 (Fig. [5](#page-3-1)D) includes both varying phase and trend. Again, trend does not increase the error for the segments without tails, but it produces some error within the tails. Adding noise to L4 (not shown) does not result in a signifcant increase in error.

In the examples shown above, we have added fairly large amounts of noise, phase variation and trend, yet errors remain small, with the largest ones being less than~0.04 of the seasonal range (here 10 mm), and the diference between postulated and adjusted seasonals is quite small.

Next, we use synthetics to test the bootstrap scheme (referred to from here on as BSS) for identifying seasonal components of a given period in short series with durations less than one period. The wisdom of trying to do this may be questioned, because Fourier analysis tells us that it cannot be done exactly. However, it is not uncommon for observed time series to have gaps due mainly to power shortages or instrument malfunction; a problem complicated by the fact that access to GPS stations located on volcanoes may be complicated because of weather factors or because of the volcanic activity itself. Hence, it is not uncommon to have relatively short data segments that may be important to the interpretation of the volcanic activity, and many times some overall curvature in the data segment will indicate the presence of a seasonal component that must be extracted for the data to be useful. Hence, we propose to identify, albeit approximately, seasonal components in short series, and application of the BSS to synthetic series will let us estimate how reliable the results will be depending on the length of the series.

Figure [6](#page-4-0) shows how parameters \vec{A} and ϕ are identified (plots A and B, respectively), the number of iterations to convergence (plot C), and, most important, the rms error between the true seasonal component and the identifed one (plot d), for diferent series lengths *N* from 25 to 364. This example uses a series having the same characteristics as the one used previously, $T = 1$ year, amplitude $A = 5.0$ mm, and initial phase $\phi_0 = 33.0^\circ$; we will not consider varying phases.

Since the most important question is whether the BSS works at all, we consider as model S0 the above mentioned seasonal with no noise or trend added. The results for this

Fig. 6 Estimated amplitude $A(A)$ and phase $\phi(B)$, number of iterations to convergence N_{it} (**C**), and rms fit error between true and estimated seasonals ϵ (**D**) for four different models as a function of series length *N*. Model S0 is a simple sinusoid, model S1 is model S0 plus normally distributed random noise, model S2 is model S0 plus a trend, and model S3 is model S2 plus noise. In (**C**), the approximate N_{it} values for very short series are written rather than plotted in order to see the details for longer series

basic model are shown as blue lines with circles, and it is clear that the BSS works very well even for quite small *N*.

The next model, S1, is model S0 with added normally distributed random noise with 0.5 mm standard deviation, and is plotted as red dashed lines with diamond markers. Although noise slightly distorts the phase and amplitude identifcation, the BSS error is still quite small all the way to lengths ~ 225 samples and is small for lengths as short as \sim 50.

Model S2 is model S0 with added trend of 1.0mm/yr., and is plotted as a black dotted line with triangle markers, and it is clear in Fig. [6](#page-4-0) that a trend superposed on the seasonal components causes large errors in the BSS. The values obtained for A and ϕ differ significantly from the true values, but these diferences appear to cancel somewhat, because the resulting ϵ , although much larger than for the trendless case, is much smaller than the diferences in *A*. The case is pretty much the same for model S3, which is S2 with normal noise added, and curiously the presence of noise reduces slightly the diferences and errors, particularly for small *N* values.

It should be pointed out that even the largest values of $\epsilon \sim 0.289$, occurring for *N* ~ 225, amount only to about 0.029 of the total range of the seasonal signal, which indicates that, although the BSS is an approximate method, it can yield useful results.

For short series that are close in time to other larger series, it should be possible to complement the results of BSS with the seasonal identifcations from the longer series.

A curious feature of BSS is that the number of iterations needed for convergence, N_{it} , appear to depend only on the length of the signal, and are essentially equal for all models, except for extremely short series with 50 or less elements, and even then they difer among models by only a small number of iterations.

The main assumption behind the method presented here is that the seasonal signals can be represented as a sinusoid, an assumption common to most of the papers dealing with seasonal identification (e.g. Blewitt and Lavallée [2002](#page-6-0); Ding et al. [2005;](#page-6-28) Davis et al. [2006,](#page-6-1) [2012](#page-6-21); Bogusz and Figurski [2014](#page-6-29)). This assumption is essential for short series, where variability is hard or impossible to appreciate, but it is relaxed in the variable phase approach for long series, and long real series usually do show at least some variability.

The real signals shown in the examples were low-pass fltered, using a Butterworth flter (Hamming [1977](#page-6-30)) with reference frequency $f_c = 20yr^{-1}$ and filter order $N_B = 4$, before the seasonal identifcation was done; but it turns out that results do not change signifcantly if the identifcation method is applied directly to unfltered signals, and the resulting signals fltered afterwards.

The averages approach works very well for long invariant series or for series longer than one period but short enough so that invariancy is not a major issue. For long variant series, the varying phase approach results in better fts, but for real data, it is impossible to tell whether the ft may be including in it any non-seasonal features.

In any case, since the method employs one-period-long segments, it is imperative that long periods be identifed and eliminated before shorter ones.

Conclusions

We have presented a simple-minded intuitive method for identifying seasonal components in GPS displacement time series. The method uses an approach based on multiple transforms, each of a one-period segment of the time series, appropriate for long series, which can take into account non-stationarity of the seasonal components, and features a scheme to deal with series somewhat smaller than one period of the seasonal component.

Unlike other more sophisticated and complicated methods, the method presented here has the advantage of simplicity, so that it is possible to recognize how the components are identifed in the processed series.

Results are, of course, approximate, but considering the length limitations and the usual noise in the real data it has been applied to, it would not be practical to complicate seasonal signal identifcation to achieve a better, but meaningless, precision.

Two MATLAB programs, one to do the seasonal components identifcation, and another to generate synthetic signals to test the identifcation program and methods, are given in full. Interested readers can test and evaluate the method by themselves using these programs.

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Author contribution FAN contributed to conceptualization, program code, writing—frst draft, and review of subsequent drafts. LÁ-B performed data curation and analysis, writing—contributions to and review of all drafts, and editing. Both authors approved the fnal manuscript.

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Code availability The MATLAB programs, diseas.m and syntseas.m, can be downloaded from [http://cicese.repositorioinstitucional.mx/jspui/](http://cicese.repositorioinstitucional.mx/jspui/handle/1007/3925) [handle/1007/3925](http://cicese.repositorioinstitucional.mx/jspui/handle/1007/3925). No special hardware or software (besides Matlab) are required. Instructions for the use of these programs are included in the corresponding headers.

Declarations

Conflict of interest The authors declare that there are no conficts of interest or competing interests.

References

- Bartel BA, Hamburger MW, Meertens CM, Lowry A, Corpuz E (2003) Dynamics of active magmatic and hydrothermal systems at Taal Volcano, Philippines, from continuous GPS measurements. J Geophys Res 108(B10):2475.<https://doi.org/10.1029/2002JB002194>
- Blewitt G, Lavallée D (2002) Effect of annual signals on geodetic velocity. J Geophys ResSolid Earth 107:ETG9
- Bogusz J, Figurski M (2014) Annual signals observed in regional GPS networks. Acta Geodyn Geomater 11(2):125–131. [https://doi.org/](https://doi.org/10.13168/AGG.2014.0003) [10.13168/AGG.2014.0003](https://doi.org/10.13168/AGG.2014.0003)
- Chanard K, Métois M, Rebischung P, Avouac J (2020) A warning against over-interpretation of seasonal signals measured by the global navigation satellite system. Nat Commun 11(1):1375
- Chen Q, Dam TV, Sneeuw N, Collilieux X, Weigelt M, Rebischung P (2013) Singular spectrum analysis for modeling seasonal signals from GPS time series. J Geodyn 72:25–35
- Chen B, Bian J, Ding K, Wu H, Li H (2020) Extracting seasonal signals in GNSS coordinate time series via weighted nuclear norm minimization. Remote Sens 2020:12. [https://doi.org/10.3390/](https://doi.org/10.3390/rs12122027) [rs12122027](https://doi.org/10.3390/rs12122027)
- Davis J, Wernicke B, Bisnath S, Niemi N, Elósegui P (2006) Subcontinental-scale crustal velocity changes along the Pacifc-North American plate boundary. Nature 441(7097):1131–1134. [https://](https://doi.org/10.1038/nature04781) doi.org/10.1038/nature04781
- Davis JL, Wernicke BP, Tamisiea ME (2012) On seasonal signals in geodetic time series. J Geophys ResSolid Earth 117(B1):B01403. <https://doi.org/10.1029/2011JB008690>
- Ding X, Zheng D, Dong D, Ma C, Chen Y, Wang GL (2005) Seasonal and secular positional variations at eight co-located GPS and VLBI stations. J Geodesy 79:71–81. [https://doi.org/10.1007/](https://doi.org/10.1007/s00190-005-0444-3) [s00190-005-0444-3](https://doi.org/10.1007/s00190-005-0444-3)
- Dong D, Fang P, Bock Y, Cheng MK, Miyazaki S (2002) Anatomy of apparent seasonal variations from GPS-derived site position time series. J Geophys Res Earth 107:ETG-9
- Freymueller J (2009) Seasonal position variations and regional reference frame realization. In: Drewes H (ed) Geodetic reference frames. Springer Berlin Heidelberg. volume 134 of International Association of Geodesy Symposia, pp 191–196. [https://doi.org/](https://doi.org/10.1007/978-3-642-00860-3) [10.1007/978-3-642-00860-3](https://doi.org/10.1007/978-3-642-00860-3) 30
- Geirsson H, d'Oreye N, Mashagiro N, Syauswa M, Celli G, Kadufu B, Smets B, Kervyn F (2017) Volcano-tectonic deformation in the Kivu Region, Central Africa: results from six years of continuous GNSS observations of the Kivu Geodetic Network (KivuGNet). J Afr Earth Sc 134:809–823
- Gonzalez-Ortega A, Sandwell D, Fialko Y, Nava A, Fletcher J, Gonzalez-Garcia J, Lipovsky B, Floyd M, Funning G (2014) El Mayor-Cucapah (Mw 7.2) earthquake: early postseismic deformation from InSAR and GPS observations. J Geophys Res 119(2):1482– 1497.<https://doi.org/10.1002/2013JB010193>
- Gruszczynska M, Klos A, Rosat S, Bogusz J (2017) Deriving common seasonal signals in GPS position time series: By using multichannel singular spectrum analysis. Acta Geodyn Geomater 14(3):267–278

Hamming R (1977) Digital flters. Prentice Hall

He X, Montillet JP, Fernandes R, Bos M, Yu K, Hua X, Jiang W (2017) Review of current GPS methodologies for producing accurate time series and their error sources. J Geodyn 106:12–29

- Hefin M, Donnellan A, Parker J, Lyzenga G, Moore A, Ludwig L, Rundle J, Wang J, Pierce M (2020) Automated estimation and tools to extract positions, velocities, breaks, and seasonal terms from daily GNSS measurements: illuminating nonlinear Salton Trough deformation. Earth Space Sci 7:e2019EA000644. [https://](https://doi.org/10.1029/2019EA000644) doi.org/10.1029/2019EA000644
- Janssen V (2007) Volcano deformation monitoring using GPS. J Spat Sci 52(1):41–54.<https://doi.org/10.1080/14498596.2007.9635099>
- Ji K, Shen Y, Wang F (2020) Signal extraction from GNSS position time series using weighted wavelet analysis. Remote Sens 12:992. <https://doi.org/10.3390/rs12060992>
- Klos A, Bos M, Fernandes R, Bogusz J (2019) Noise-dependent adaption of the Wiener flter for the GPS position time series. Math Geosci 51:53–73
- Larson KM, Poland M, Miklius A (2010) Volcano monitoring using GPS: developing data analysis strategies based on the June 2007 Kilauea Volcano intrusion and eruption. J Geophys Res. [https://](https://doi.org/10.1029/2009JB007022) doi.org/10.1029/2009JB007022
- Lee S-W, Yun S-H, Kim DH, Lee D, Lee YJ, Schutz BE (2015) Realtime volcano monitoring using GNSS single-frequency receivers. J Geophys Res Solid Earth 120:8551–8569. [https://doi.org/](https://doi.org/10.1002/2014JB011648) [10.1002/2014JB011648](https://doi.org/10.1002/2014JB011648)
- Mann D, Freymueller J, Lu Z (2002) Deformation associated with the 1997 eruption of Okmok volcano. Alsk J Geophys Res 107(B4):2072
- Miyagi Y, Freymueller JT, Kimata F, Sato T, Mann D (2004) Surface deformation caused by shallow magmatic activity at Okmok volcano, Alaska, detected by GPS campaigns 2000–2002. Earth Planets Space 56:e29–e32
- Puglisi G, Bonforte A (2004) Dynamics of Mount Etna Volcano inferred from static and kinematic GPS measurements. J Geophys Res 109:B11404. <https://doi.org/10.1029/2003JB002878>
- Ray J, Altamimi Z, Collilieux X, Van Dam T (2008) Anomalous harmonics in the spectra of GPS position estimates. GPS Solut 12:55–64
- Ruttner P (2021) Analysis and prediction of long term GNSS height time series and environmental loading efects. Master's thesis, ETH Zurich
- Tregoning P, Watson C (2009) Atmospheric efects and spurious signals in GPS analyses. J Geophys Res 114:B09403. [https://doi.org/](https://doi.org/10.1029/2009JB006344) [10.1029/2009JB006344](https://doi.org/10.1029/2009JB006344)
- Tucikešić S, Božić B, Mulić M (2020) Absolute time series GNSS point positioning-data cleaning and noise characterization. Tehnički Vjesnik 27(4):1229–1236
- Wang K, Jiang W, Chen H, An X, Zhou X, Yuan P, Chen Q (2018) Analysis of seasonal signal in GPS short-baseline time series. Pure Appl Geophys 175:3485–3509
- Wernicke B, Davis J (2010) Detecting large scale intracontinental slowslip events (SSEs) using geodograms. Seismol Res Lett 81:694– 698.<https://doi.org/10.1785/gssrl.81.5.694>

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