#### **RESEARCH ARTICLE - SOLID EARTH SCIENCES**

# **The earthquake network: the best time scale for network construction**

**Nastaran Lotf[1](http://orcid.org/0000-0001-7258-6117)**

Received: 3 January 2023 / Accepted: 6 June 2023 / Published online: 3 July 2023 © The Author(s) under exclusive licence to Institute of Geophysics, Polish Academy of Sciences & Polish Academy of Sciences 2023

#### **Abstract**

Scientists mapped the seismic time series into networks by considering the geographical location of events as nodes and establishing links between the nodes with diferent rules. Applying the successively defned models to construct the networks of seismic data, a variety of features of earthquake networks are detected (scale-free and small-world structures). Network construction models had changed in detail to optimize the performance of the verifcation of the minimum geographical size defned for the node. In all the studies, people try to use large data sets like years of data to ensure their results are good enough. In this work, by proposing the temporal network construction and employing the small-worldness property for data from Iran and California, we could achieve the minimum time scale needed for the best results. We verifed the importance of this scale by analyzing two signifcant centrality measures (degree centrality and PageRank) introduced in the concept of earthquake network.

**Keywords** Complex networks · Temporal networks · Earthquake networks

### **Introduction**

An earthquake is a sudden motion of a fault that releases an enormous amount of energy and is considered a complex spatiotemporal phenomenon occurring in the earth's crust (Kanamori and Brodsky [2001\)](#page-6-0). Transferring the stress of the movement of one fault to the others results in triggering subsequent events (King et al. [1994;](#page-6-1) Belardinelli et al. [2003;](#page-6-2) Freed [2005\)](#page-6-3). The Omori law (Omori [1894\)](#page-6-4) and the Gutenberg–Richter law (Gutenberg and Richter [1944](#page-6-5)) are empirical laws to characterize the temporal pattern of aftershocks and frequency-energy, respectively. Besides the visible properties, complex interaction exists in the internal of the seismic system (Bak et al. [2002;](#page-6-6) Baiesi and Paczuski [2004](#page-6-7); Gutenberg [2013\)](#page-6-8).

While seismicity is assumed to be a complex phenomenon, the network approach offers a powerful tool for analyzing the dynamic structures of it (Abe and Suzuki [2004a](#page-6-9), [b,](#page-6-10)

Edited by Dr. Rodolfo Console (ASSOCIATE EDITOR) / Prof. Ramón Zúñiga (CO-EDITOR-IN-CHIEF).

 $\boxtimes$  Nastaran Lotfi n.lotf@icmc.usp.br [2005,](#page-6-11) [2006](#page-6-12)). Over the last decade, diferent models proposed to construct the earthquake network (Abe and Suzuki [2004b](#page-6-10); Lacasa et al. [2008](#page-6-13); Rezaei et al. [2017](#page-6-14); Lotf et al. [2018](#page-6-15)). In the simple but basic model introduced by (Abe and Suzuki [2004b](#page-6-10)), the geographical region is divided into small square (cubic) cells, and seismic events with time sequences get connected. Later, (Lacasa et al. [2008\)](#page-6-13) proposed a model to construct the network with a visibility graph. They converted the time series into a graph by inheriting the properties of the series in its structure. They explored periodicity, fractality, chaoticity, and nonlinearity of the seismic time series (Lacasa et al. [2009](#page-6-16); Lacasa and Toral [2010;](#page-6-17) Donges et al. [2013\)](#page-6-18). (Rezaei et al. [2017\)](#page-6-14) introduced the hybrid model, which inherits the bases of the Abe-Suzuki model mixed with a visibility graph. To better capture the evolution of the earthquake network through time, a multiplex network was employed (Lotfi et al. [2018](#page-6-15)). Analyzing the seismic data with a network approach through diferent models helped reveal many features of the seismic activity just by knowing the basic information of magnitude, time of occurrence, and the location of seismic events (Baiesi and Paczuski [2004](#page-6-7); Abe and Suzuki [2004a](#page-6-9); Lotfi and Darooneh [2012](#page-6-19); Abe et al. [2011;](#page-6-20) Lotf and Darooneh [2013\)](#page-6-21). It had been verifed that the earthquake networks that constructed from the seismic data taken from California and Japan (Abe and Suzuki [2004a,](#page-6-9) [b](#page-6-10), [2005](#page-6-11)), Iran (Lotf and Darooneh [2012](#page-6-19), [2013](#page-6-21)),



<sup>1</sup> Instituto de Ciências Matemáticas e de Computação, Universidade de São Paulo, Caixa Postal 668, São Carlos, SP 13560-970, Brazil

<span id="page-1-4"></span>



Chile (Abe et al. [2011](#page-6-22)), Greece (Chorozoglou et al. [2019](#page-6-23)), and Italy (Rezaei et al. [2017\)](#page-6-14) are scale-free and small-world.

Most recent works focused on improving the proposed models to capture the best minimum resolution of the cell size needed for network construction. It means the cell size should be smaller than the specifed limit to be trustable. The main question is how we ensure that the time window, in the scale of dates, months, or years, is large enough for constructing the network. In all the studies done till now, scientists considered the time on such a big scale of years. And the concept of the minimum necessary time window for achieving the best results is missing. In this work, we employ the defnition of temporal network construction and capture the lowest time window essential for network construction. Considering the simplest model for earthquake network construction (Abe-Suzuki model), we found that depending on the region of consideration, the value of the time window threshold would change. We verifed the trustiness of this time window size by analyzing two important centrality parameters, degree centrality, and PageRank. If the time window is small, we miss the information in centrality, and if it is bigger than the threshold, we do not gain extra knowledge than in the threshold time region.

The rest of the paper is organized as follows. In ["Data](#page-1-0)[base](#page-1-0)" section, we provide information about the data sets we employ, and "[Results"](#page-1-1) section is devoted to our results.

# <span id="page-1-0"></span>**Database**

We applied our model for the latest four years of data, Jan 01, 2018, to Dec 31, 2021, for Iran in the range of 24*N* − 44*N* latitude and  $40E - 62E$  longitude with 14062 total events obtained from Iranian Seismological Center,<sup>[1](#page-1-2)</sup> and California in the range of 32*N* − 42*N* latitude and 114*W* − 124*W* longitude with 7575 total events gained from the Northern

### <span id="page-1-1"></span>**Results**

Through diferent models introduced for earthquake network construction, we used the simple model introduced by (Abe and Suzuki [2004b](#page-6-10)). As it is plotted in Fig. [1](#page-1-4), the geographical region is divided into small square cells and having seismic events data ordered by the occurrence time, each square is regarded as one node if an earthquake with any magnitude occurred, and two nodes with consecutive events are connected.

We also divided the seismic data of four years length into small time windows in the following way; in the frst step, we construct the Abe-Suzuki network for the data for the length of one month and study the characteristics of interest. Then, we added the data from the second month to the previous one and reconstructed the network. The process of adding data by time windows of the size of one month continues until the whole 48 months of data are covered. The schematic representation of the temporal network construction is plotted in Fig. [2a](#page-2-0)–c for Iran and Fig. [3a](#page-3-0)–c for California. Figures [2a](#page-2-0) and [3](#page-3-0)a belong to the data of the length of one month for Iran and ten months for California. As the number of events is low, having a sparse network is anticipated. The second Figs. [2b](#page-2-0) and [3b](#page-3-0) are in the middle time when the network is not sparse as the frst month and is not too connected as the last, and Figs. [2c](#page-2-0) and [3c](#page-3-0) represent the networks for the whole four-year data which have very dense connections.

<span id="page-1-3"></span>The second step would be building the adjacency matrix *A* for facilitating the analysis;  $a_{ij} = 1$  if nodes *i* and *j* are

California Earthquake Catalog.<sup>[2](#page-1-3)</sup> In both of the considered data sets, we examined only events with a magnitude larger than 2.5–7. Referring to the analysis done to defne the best cell size for network construction Lotfi and Darooneh ([2013](#page-6-21)), we considered  $N_{cell} = 12.5$  for Iran and  $N_{cell} = 10$ for California.

<span id="page-1-2"></span>

<sup>1</sup> <http://irsc.ut.ac.ir> <sup>2</sup> <http://www.usgs.gov/>



<span id="page-2-0"></span>**Fig. 2 a**–**c** The schematic representation of temporal earthquake networks of Iran for three diferent time steps with time windows of **a** one month, **b** 10 months, and **c** 46 months with fltered data magni-

tudes *>* 4.0, **d**–**f** are the degree distribution of the networks for the three defned networks, respectively, for data with magnitudes *>* 2.5)

connected, and 0 otherwise (Fig. [1\)](#page-1-4). In the network defnitions, the degree of the node is the number of connections a node could have and is calculated from the adjacency matrix  $k_i = \sum_j a_{ij}$ . The degree distribution of the earthquake network of diferent regions obeys a power law indicating the scale-free network characteristics (Abe and Suzuki [2004a,](#page-6-9) [2005](#page-6-11); Lotfi and Darooneh [2012](#page-6-19)). In the scalefree networks, the power law decay of the degrees means that the vast majority of nodes have very few connections, while a few important nodes have a huge number of connections. To check the validity of this characteristic, for each of the above-mentioned networks (Figs. [2](#page-2-0) and [3](#page-3-0)), we plot the degree distribution in Figs. [2d](#page-2-0) and [3d](#page-3-0). One could see that no matter the time length, we would have approximately the power-law distribution with power of  $\gamma$ . This  $\gamma$  varies between 2 and 2.2 for Iran and 1.5 to 1.8 for California.

The other famous characteristic of earthquakes small-worldness (Abe and Suzuki [2004b](#page-6-10); Lotfi and Darooneh [2012\)](#page-6-19). In small-world networks, although most of the nodes are not neighbor together, they could be reached from every other nodes by a small number of links. In other words, in these networks, with *N* nodes and *M* links, the value of the shortest path is similar to the random network with the same numbers of nodes and links, while the clustering coefficient has a larger value. The clustering coefficient of a node  $i$  is the fraction of connection existing among its nearest neighbor nodes to the maximum number of possible links among them. The clustering coefficient of the network would be the average clustering of all nodes:

$$
C_i = \frac{1}{k_i(k_i - 1)} \sum_{j,k} a_{ij} a_{jk} a_{ki}, \quad C = \frac{1}{N} \sum_{i=1}^{N} C_i,
$$
 (1)

where *N* is the total number of nodes in the network. In other words, the clustering coefficient is the tendency of the nodes in the graph to cluster together and has a value  $0 \leq C \leq 1$ . On the other hand, the shortest path is the minimum path length needed to traverse to get from one node to another. The average over all nodes would result in the shortest path of the network:



<span id="page-3-0"></span>**Fig. 3 a**–**c** The schematic representation of temporal earthquake networks of California for three diferent time steps with time windows of **a** 10 months, **b** 19 months, and **c** 46 months with fltered data mag-

nitudes *>* 3.0, **d** to **f** are the degree distribution of the networks for the three defned networks, respectively, for data with magnitudes *>* 2.5)

<span id="page-3-1"></span>



$$
L = \frac{1}{N(N-1)} \sum_{i,j=1,N; i \neq j} d_{ij},
$$
\n(2)

$$
S_w = \frac{C/C_{\text{rand}}}{L/L_{\text{rand}}}.\tag{3}
$$

in which  $d_{ij}$  is the minimum length of the path between two nodes of *i* and *j*.

By having the clustering coefficient and shortest path of the network, Humphries and Gurney [\(2008\)](#page-6-24) introduced a small-worldness metric defned with the averaged clustering coefficient and path length relative to these metrics for random networks. This metric helps to provide an overview of connectivity in the entire network:

 $C_{\text{rand}} = \frac{}{N}$  and  $L_{\text{rand}} = \frac{ln(N)}{ln(})}$  are the values obtained for random networks by randomizing the connections of each earthquake network by keeping the same number of nodes and links. Here,  $\langle k \rangle$  indicates the mean degree of the networks.

The variation of  $S_w$  in time (in the scale of the length of the month) is shown in Fig. [4](#page-3-1). One could see that this



<span id="page-4-0"></span>**Fig. 5** Degree centrality and PageRank for earthquake network of Iran for time windows of **a**, **d** one month, **b**, **e** 10 months, and **c**, **f** 46 months, respectively

value is small for the frst months of consideration. It starts to increase until a threshold and gets stationery later. This behavior could emphasize that until a specifc time window, the variation of the parameters is high. The fuctuations disappear, while one considers a time window that is large enough, and the system gets stationary. The geographical region under consideration and frequency of the seismic event could result in observing diferent values. This value for Iran's data is approximately ten months, while for California it is around 19 months.

To clarify the importance of having the minimum time window, we calculate two of the most important centralities in the concept of earthquake networks and compare them in three diferent time windows. Looking through the literature, one could fnd diferent parameters to calculate the centrality of nodes in the earthquake networks. The simplest and most common centrality that uses the local structure around the nodes is the degree centrality. In Figs. [5](#page-4-0) and [6a](#page-5-0)–c, we plot the degree centrality for three diferent time scales as the following: Figs. [5a](#page-4-0) and [6a](#page-5-0) are for the time window of length 1 month for Iran and 10 months for California. Near the time window of the threshold, we selected the network with the length of 10 months of data for Iran (Fig. [5](#page-4-0)b) and 19 months for California (Fig. [6b](#page-5-0)). And Figs. [5c](#page-4-0) and [6](#page-5-0)c belong to the largest time window (48 months).

The second famous centrality in the concept of earth-quake network is PageRank (Darooneh and Lotfi [2014](#page-6-25); Rezaei et al. [2019\)](#page-6-26). PageRank is an algorithm used to assess the ranks of nodes in a network based on their connections' levels used in the Google search engine for ranking web pages for the frst time (Brin and Page [1998](#page-6-27)). PageRank explained through the random walk. The random walker starts from one node and selects the next one randomly. In this defnition, PageRank of node *i* is the asymptotic probability that the walker meets the node. One could infer that the possibility of reaching one important node is higher



<span id="page-5-0"></span>**Fig. 6** Degree centrality and PageRank for earthquake network of California for time windows of **a**, **d** 10 month, **b**, **e** 19 months, and **c**, **f** 46 months, respectively

than the unimportant ones. Evaluating this centrality is an iterative procedure in which the PageRank of nodes depends on all its neighbors' PageRanks. The following equation describes such a random walking procedure:

$$
PR_{i} = \frac{d}{N} + (1 - d) \sum_{j \in B_{i}} \frac{PR_{j}}{k_{j}^{out}},
$$
\n(4)

in which  $PR_i$  is the PageRank of node *i*,  $B_i$  is the set of nearest neighbors of node *i*, and *kout* is the out-degree of each node. *d* is a constant value (mostly considered as 0.15) defned as the probability of jumping to any vertex. Figures [5](#page-4-0) and [6](#page-5-0)d–f are representing the PageRank of the networks for the three diferent time windows. Taking into account both above-introduced centralities, one could see that in the small-time windows, we could not fnd enough information about the central locations of the regions as it should. If we increase the length of the time window up to the threshold,

the results capture the same central regions as the largest time window.

# **Conclusion**

Recently, diferent models proposed to study the earthquake phenomena to explore the features of this harmful disaster. Although this phenomenon is very complex from a fault and inside earth interactions point of view, it is possible to study it with the complex network with the minimum information: geographical location, time, and magnitude. Among the most famous models proposed, Abe-Suzuki and visibility models, scientists were trying to improve the model's performances. The main idea that got most of the attention from those studying was how they could introduce the best minimum geographical cell size.

Here, we proposed the temporal earthquake network construction for capturing another essential factor of network analysis, the best time window size. We start with constructing an earthquake network in windows of the month length and adding data with a length of one month in each step. We used the most straightforward model introduced by Abe-Suzuki to build our networks. For each constructed network, the small-worldness is evaluated. Studying how this parameter changes by increasing the time window, we could verify the minimum length of time window needed for network construction. This value is smaller than its value in the threshold time window and gets stationary by enlarging the time lengths. This time threshold is diferent for each geographical region as the construction of the earth is diferent. One point of these diferences appears in the frequency of the events on the same time scale. Then, it is a delicate factor to study the minimum and efficient time window size for different geographical regions before the rest of the analysis to ensure obtaining the best results. By considering two famous centralities measures in the concept of earthquake networks (degree centrality, and PageRank), we show that if this size is smaller than the threshold, we will miss the information we should have. If the time window is too large, it does not provide extra information.

**Acknowledgements** N. Lotfi is thankful to the FAPESP (grant with number 2020/08359-1) for the support given to this research.

#### **Declarations**

**Conflict of Interest** The authors declare that they have no known competing fnancial interests or personal relationships that could have appeared to infuence the work reported in this paper.

### **References**

- <span id="page-6-22"></span>Abe S, Pastén D, Muñoz V, Suzuki N (2011) Universalities of earthquake-network characteristics. Chinese Sci Bull 56(34):3697–3701
- <span id="page-6-20"></span>Abe S, Pastén D, Suzuki N (2011) Finite data-size scaling of clustering in earthquake networks. Physica A 390(7):1343–1349
- <span id="page-6-9"></span>Abe S, Suzuki N (2004) Scale-free network of earthquakes. Europhys Lett 65(4):581
- <span id="page-6-10"></span>Abe S, Suzuki N (2004) Small-world structure of earthquake network. Physica A 337(1–2):357–362
- <span id="page-6-11"></span>Abe S, Suzuki N (2005) Scale-invariant statistics of period in directed earthquake network. Eur Phys J B 44(1):115–117
- <span id="page-6-12"></span>Abe S, Suzuki N (2006) Complex-network description of seismicity. Nonlinear Proc Geophys 13(2):145–150
- <span id="page-6-7"></span>Baiesi M, Paczuski M (2004) Scale-free networks of earthquakes and aftershocks. Phys Rev E 69(6):066106
- <span id="page-6-6"></span>Bak P, Christensen K, Danon L, Scanlon T (2002) Unifed scaling law for earthquakes. Phys Rev Lett 88(17):178501
- <span id="page-6-2"></span>Belardinelli M, Bizzarri A, Cocco M (2003) Earthquake triggering by static and dynamic stress changes. J Geophys Res 108(B3)
- <span id="page-6-27"></span>Brin S, Page L (1998) The anatomy of a large-scale hypertextual web search engine. Comput Netw ISDN 30(1–7):107–117
- <span id="page-6-23"></span>Chorozoglou D, Papadimitriou E, Kugiumtzis D (2019) Investigating small-world and scale-free structure of earthquake networks in greece. Chaos, Solitons & Fractals 122:143–152
- <span id="page-6-25"></span>Darooneh AH, Lotfi N (2014) Active and passive faults detection by using the pagerank algorithm. Europhys Lett 107(4):49001
- <span id="page-6-18"></span>Donges JF, Donner RV, Kurths J (2013) Testing time series irreversibility using complex network methods. Europhys Lett 102(1):10004
- <span id="page-6-3"></span>Freed AM (2005) Earthquake triggering by static, dynamic, and postseismic stress transfer. Annu Rev Earth Planet Sci 33:335–367
- <span id="page-6-8"></span>Gutenberg B (2013) Seismicity of the earth and associated phenomena, Read Books Ltd
- <span id="page-6-5"></span>Gutenberg B, Richter CF (1944) Frequency of earthquakes in california. Bull Seismol Soc Am 34(4):185–188
- <span id="page-6-24"></span>Humphries MD, Gurney K (2008) Network 'small-world-ness': a quantitative method for determining canonical network equivalence. PLoS ONE 3(4):e0002051
- <span id="page-6-0"></span>Kanamori H, Brodsky EE (2001) The physics of earthquakes. Phys Today 54(6):34–40
- <span id="page-6-1"></span>King GC, Stein RS, Lin J (1994) Static stress changes and the triggering of earthquakes. Bull Seismol Soc Am 84(3):935–953
- <span id="page-6-13"></span>Lacasa L, Luque B, Ballesteros F, Luque J, Nuno JC (2008) From time series to complex networks: the visibility graph. Proc Natl Acad Sci 105(13):4972–4975
- <span id="page-6-16"></span>Lacasa L, Luque B, Luque J, Nuno JC (2009) The visibility graph: a new method for estimating the hurst exponent of fractional brownian motion. Europhys Lett 86(3):30001
- <span id="page-6-17"></span>Lacasa L, Toral R (2010) Description of stochastic and chaotic series using visibility graphs. Phys Rev E 82(3):036120
- <span id="page-6-19"></span>Lotfi N, Darooneh A (2012) The earthquakes network: the role of cell size. Eur Phys J B 85(1):1–4
- <span id="page-6-21"></span>Lotfi N, Darooneh AH (2013) Nonextensivity measure for earthquake networks. Physica A 392(14):3061–3065
- <span id="page-6-15"></span>Lotfi N, Darooneh AH, Rodrigues FA (2018) Centrality in earthquake multiplex networks. Chaos 28(6):063113
- <span id="page-6-4"></span>Omori F (1894) On the aftershocks of earthquakes. J Coll Sci 7:111–120
- <span id="page-6-14"></span>Rezaei S, Darooneh AH, Lotf N, Asaadi N (2017) The earthquakes network: retrieving the empirical seismological laws. Physica A 471:80–87
- <span id="page-6-26"></span>Rezaei S, Moghaddasi H, Darooneh AH (2019) Pagerank: an alarming index of probable earthquake occurrence. Chaos 29(6):063114

Springer Nature or its licensor (e.g. a society or other partner) holds exclusive rights to this article under a publishing agreement with the author(s) or other rightsholder(s); author self-archiving of the accepted manuscript version of this article is solely governed by the terms of such publishing agreement and applicable law.