#### **RESEARCH PAPER**



# Initiation of flow liquefaction in granular soil slopes: drained versus undrained conditions

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#### Abstract

The mechanism of flow liquefaction in granular soils under drained conditions compared with that under undrained conditions is a subject of great interest. This paper discusses the diverse views in relation to the initiation of instability and the conceptual difficulties brought by the discrepancies. It is shown that the constant shear drained (CSD) stress path is not a conventional stress-controlled test, and overlooking this fact may result in misleading conclusions. The difficulties inherent in the CSD tests have resulted in different criteria for the detection of the onset of instability in experiments and thereafter to different views. Through systematic simulations for the CSD, ACU and ICU stress paths under a range of initial states, it is shown that the instability line in the stress space is state dependent, and its gradient can be well related to the state parameter defined using the critical state concepts. For a given soil, the relationship is nearly unique regardless of the CSD, ACU or ICU stress path, and the critical state concepts and the second-order work criterion work well. Theoretically, it is shown that the conditions for instability under the drained and undrained stress paths are the same, regardless of the constitutive models chosen.

Keywords Flow liquefaction · Granular soil · Instability · Shear behaviour · Stress path

#### List of symbols

- A, B Parameters in Eq. (1)
- $A_{\rm F}$  Fabric anisotropy variable
- [C] Elastoplastic stiffness matrix in Eq. (2)
- *d* State-dependent dilatancy
- $e_0$  Void ratio after isotropic consolidation
- $e_A$  Material constant related to fabric effect
- $e_{\Gamma}$  Critical state parameter in Eq. (8)
- $F^d$  Norm of deviatoric part of fabric tensor
- G Elastic shear modulus
- *K* Elastic bulk modulus
- $K_{\rm p}$  Plastic hardening modulus
- $M_{\rm c}$  Critical state stress ratio
- p' Mean effective stress in standard triaxial setting
- $p_{\rm at}$  Reference pressure (atmospheric pressure)
- *q* Deviatoric stress in standard triaxial setting
- $\varepsilon_{a}$  Axial strain in triaxial test
- $\varepsilon_q$  Deviatoric strain in triaxial test
- $\varepsilon_{v}$  Volumetric strain in triaxial test

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- $\eta$  Stress ratio (q/p')
- $\eta_{IL}$  Stress ratio corresponding to instability line
- $\lambda_c$  Critical state parameter in Eq. (8)
- $\psi$  State parameter defined with reference to CSL
- $\psi_0$  Initial state parameter prior to shearing

# 1 Introduction

Flow liquefaction, also known as static liquefaction or flow slide, is a phenomenon that a granular soil mass, when subjected to a perturbance, undergoes a sudden loss of strength and a rapid development of deformation, accompanied by a quick buildup of pore water pressure. The consequences of flow liquefaction are often dramatic and devastating. Notable cases include the collapse of Fort Peck Dam [6], the flow slides of a submarine berm [31] and the failure of Stava tailings dams [7]. Over the past decades, a great deal of efforts has been made to investigate the mechanisms of flow liquefaction [3, 6, 9, 21, 23, 26, 31, 32, 36, 37, 39], leading to a framework of understanding and analysis that is now widely accepted. The key concepts of the framework

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include (a) the existence of a critical state line in the stress space (Fig. 1), which is the locus of ultimate states of shear failure, termed as steady state or critical state and (b) the existence of a flow liquefaction line or instability line, which links the peak in the stress path and the origin. The two lines in the stress space define a zone of potential instability in which loose granular soil tends to become unstable when a perturbation is imposed. It is worth noting that the instability line is not an intrinsic property but depends on both void ratio and mean effective stress, and its gradient can be related to the initial state parameter ( $\psi_0$ ) of the soil as follows [37]:

$$\eta_{\rm IL} = \left(\frac{q}{p'}\right)_{\rm IL} = \frac{M_{\rm c}}{B} \exp(A\phi_0) \tag{1}$$

where  $M_c$  is the gradient of the critical state locus; A and B are two soil-specific parameters. As shown in Fig. 1, the state parameter is a measure of how far the initial state is from the critical state in terms of density [3]. The implication of Eq. (1) is that a granular soil mass at a medium dense state may also have the potential for liquefaction as long as the confining stress is sufficiently high.

Most of the studies on flow liquefaction have tended to focus on isotropically consolidated specimens subjected to



**Fig. 1** Undrained monotonic loading of loose sand leading to flow liquefaction in the context of critical state concepts (after [39]): **a** q-p' plane and **b** e-p' plane

undrained shear. This condition, however, does not properly replicate in situ stress conditions typical for slopes and dams-in which a static, driving shear stress exists in the element of soil prior to external loading (Fig. 2a). There have been concerns about the potential effect of the initial shear stress. For example Fourie and Tshabalala [18] reported an investigation of a tailings dam failure in South Africa, arguing that the use of the friction angle corresponding to the instability line determined from isotropically consolidated specimens would lead to unreasonably low values of factor of safety. They carried out undrained triaxial tests on isotropically and  $K_0$  consolidated specimens of tailings material, concluding that the instability line determined from  $K_0$  consolidated specimens is positioned steeper in the stress space, thus giving a larger friction angle and a higher factor of safety. This conclusion appears to be counterintuitive, since it suggests that the presence of initial shear stress is *beneficial* to the safety of slopes. More recently, Yang et al. [39] reported new data sets obtained from a specifically designed experimental programme, showing that no notable difference exists between the stress ratios at the onset of instability for isotropically and anisotropically consolidated specimens, as long as their post-consolidation states are similar, and that the critical state concepts are applicable.

Adding to the complexity is that flow liquefaction can also be initiated under *drained* conditions, as evidenced by a number of rainfall-induced slope failures. Among the tragic case histories are the collapse of a coal waste tip in Aberfan (South Wales) in 1966, which killed 144 people including 116 schoolchildren, and the massive slides of a large fill slope in Sau Mau Ping (Hong Kong) in 1972, which claimed 76 lives. There is now a general agreement that a perched water table as a result of rainfall infiltration can be a critical cause for flow slides [8, 35]. In laboratory experiment, the field stress path can be replicated using the so-called constant shear drained (CSD) test in which rise in pore water pressure, due to water infiltration, causes a decrease in mean effective stress at an almost constant shear stress [5, 29]. Since the stress path is largely different from that in conventional undrained triaxial tests, there is growing interest in the mechanical behaviour of soils in the CSD tests [1, 2, 10, 12, 15, 22, 25, 34]. These studies have created valuable data that contribute to the development of fundamental understanding of the problem. Nevertheless, existing views on several critical issues remain rather diverse or controversy. It is noticed that theoretical prediction of the condition for instability using advanced constitutive laws or the notion of loss of controllability [20, 28] is not in agreement with experimental results. Adding to this complexity, several researchers reported that for loose sandy soil specimens, where the state parameter  $(\psi)$  is positive, contraction will not occur along the CSD



Fig. 2 Schematic illustration of stress state changes induced by pore water pressure rise in association with rainfall:  $\mathbf{a}$  field condition and  $\mathbf{b}$  laboratory simulation using constant shear drained (CSD) test

stress path but can occur on the conventional undrained triaxial stress path [22]. Also, the final states of the specimens in many of the reported experiments did not fall close to the corresponding critical state loci established from the conventional triaxial compression tests.

These observed discrepancies are not trivial but bring about significant conceptual difficulties, since they indicate that both instability line and critical state line are stress path dependent. The implication is that the existing analysis framework, established from extensive undrained triaxial tests in conjunction with the critical state concepts, is of little use in the evaluation of flow liquefaction associated with the CSD stress path. In this paper, an attempt is made to address the concerns arising from these discrepancies, by combining theoretical and experimental considerations and paying attention to several fundamental questions: (a) How the mechanisms of flow liquefaction under drained conditions differ from that under conventional undrained conditions? (b) Whether the instability line determined from the CSD tests differs from that determined from the anisotropic or isotropic undrained triaxial tests? (c) Whether the instability line from the CSD tests is dependent on void ratio and effective confining stress? and (d) If yes, can this state dependence be collectively accounted for using the critical state concepts?

# 2 Theoretical and experimental considerations

For the sake of simplicity, discussion is limited to the triaxial setting commonly used in soil mechanics, but it can be generalised to multiaxial conditions without difficulty. In elastoplasticity, the relationship between stress and strain increments can be established for triaxial setting as follows [38]:

$$\begin{cases} dp' \\ dq \end{cases} = \begin{bmatrix} C_{pp} & C_{pq} \\ C_{qp} & C_{qq} \end{bmatrix} \begin{cases} d\varepsilon_{v}^{e} + d\varepsilon_{v}^{p} \\ d\varepsilon_{q}^{e} + d\varepsilon_{q}^{p} \end{cases} = [C] \begin{cases} d\varepsilon_{v} \\ d\varepsilon_{q} \end{cases}$$
(2)

where p' is mean effective stress, and q is deviatoric stress;  $\varepsilon_v$  is volumetric strain, and  $\varepsilon_q$  is deviatoric strain; the superscripts e and p denote elastic and plastic strain components, respectively. The four components of the elastoplastic stiffness matrix [C] can be determined for a given constitutive model. For stress-controlled tests in laboratory, dp and dq are considered control variables; for example dq = 0 and dp' < 0 are prescribed in CSD tests (Fig. 2). If onset of instability is regarded as loss of controllability or uniqueness in the test [20], mathematically, this requires vanishing determinant of the stiffness matrix in Eq. (2), i.e. det[C] = 0. This is exactly the derivation used in several studies [28] to generate the theoretical prediction. One may note that the condition of det[C] = 0 is equivalent to  $K_{p-1}$ = 0, where  $K_p$  is known as plastic hardening modulus in elastoplasticity [38]. Therefore, the condition leading to instability along the CSD stress path is coincident with the condition leading to failure in conventional drained triaxial tests [38]; that is for loose sand, the onset of instability occurs at the critical state, whereas for dense sand, the onset of instability occurs at the peak state.

While the above derivation appears to be theoretically sound, it fails to yield correct prediction for the experimental observation. The reason is that the derivation overlooks an important fact of the CSD tests, that is the test is not a conventional stress-controlled test. In conducting CSD tests on triaxial apparatus, two methods are generally adopted: One is to reduce the confining pressure at a given rate while maintaining the shear stress constant, and the other method is to maintain the confining pressure unchanged while increasing the back pressure. In both methods, the pore water pressure in the specimen is under control such that the volume change of the specimen is not entirely free. Another line of analysing flow liquefaction is to treat it as an instability behaviour and employ the classic Hill's stability criterion [4, 13]. In accordance with Hill's criterion [19] and for triaxial setting, the condition for a material being stable is given as follows:

$$d^2W = \left(dp'd\varepsilon_{\rm v} + dq\,d\varepsilon_q\right) > 0 \tag{3}$$

where  $d^2W$  is the second-order work per unit volume of the material. For a perfect CSD test in which dq = 0, the condition leading to instability becomes  $d\varepsilon_v = 0$ . However, it has been frequently observed in CSD tests that a strictly constant shear stress cannot be achieved [10, 25, 30], due to inherent difficulty. This implies that instability can be initiated under conditions other than  $d\varepsilon_v = 0$ . Depending on the sign of dq, the volumetric strain increment ( $d\varepsilon_v$ ) at onset of instability can be positive or negative in accordance with Eq. (3). Owing to this difficulty and other uncertainty, different criteria have been put forward to detect the onset of instability in CSD tests [10, 22, 25]. As a consequence, diverse views and confusion arise.

It is also noted that in the investigation of the instability behaviour along the CSD stress paths, one often chooses to compare the result with that from isotropically consolidated undrained tests (i.e. ICU tests) and assumes that these ICU tests give a unique instability line. This is not considered reasonable since there is now solid evidence that the instability line from ICU tests is not unique but depends on the post-consolidation state, as discussed earlier. A small change in void ratio can lead to a marked change in liquefaction potential of sand specimens in undrained tests [33, 41]. Furthermore, isotropic consolidation does not replicate the anisotropic stress state in a slope (Fig. 2). A rational and desirable comparison is that the CSD test and the anisotropic consolidation undrained test (ACU test) share a similar post-consolidation state in terms of void ratio, mean effective stress and deviatoric stress prior to shearing. From the experimental point of view, however, it is difficult to obtain systematic data sets that fulfil this requirement, as the behaviour of sandy soils is sensitive to void ratio. In the present study, an alternative strategy is used, which involves systematic simulations of the CSD, ACU and ICU stress paths in the framework of anisotropic critical state plasticity and Hill's stability criterion.

# 3 Critical state plasticity modelling and calibration

A robust critical state-based modelling platform [11, 38] is modified in relation to the effect of fabric anisotropy [24, 42] to improve its predictive capability for a range of responses. The model formulations are briefly presented below for completeness. More details can be found in the relevant references. The yield surface of the model is a wedge in the q-p' plot, and kinematic hardening law is assumed:

$$f = \begin{cases} \eta - \alpha - m = 0 & (\eta > \alpha) \\ \eta - \alpha + m = 0 & (\eta < \alpha) \end{cases}$$
(4)

where  $\eta$  is the stress ratio defined as q/p',  $\alpha$  is back stress and *m* is a constant defining the width of the wedge. The evolution of dilatancy is assumed to depend on state parameter, stress ratio and fabric anisotropic variable as follows:

$$d = d_0 \left[ M_c \exp(n^d \zeta) - \eta \right]$$
<sup>(5)</sup>

where  $M_c$  is the critical state stress ratio,  $d_0$  and  $n^d$  are material constants to be calibrated from experimental data.  $\zeta$  is dilatancy state parameter and is defined as  $\zeta = \psi - e_A(A_F - 1)$ ; here,  $e_A$  is a material constant, and  $A_F$  is a fabric anisotropy variable accounting for the combined effect of loading direction and deviatoric part of the fabric tensor:

$$A_{\rm F} = F^d_{ij} n_{ij} \tag{6}$$

The plastic hardening modulus is defined as follows:

$$K_{p} = \frac{G_{0}h_{0}(1-c_{h}e)}{|\eta-\eta_{in}|} \left(\frac{p'}{p_{at}}\right)^{-1/2} \left[M_{c}\exp(-n^{b}\zeta) - \eta\right]$$
(7)

where  $\eta_{in}$  is the stress ratio, and it is updated at loading reversal state,  $G_0$  is elastic modulus constant,  $h_0$ ,  $c_h$  and  $n^b$ are material constants. The critical state locus in e-p' plane is defined as follows:

$$e_{\rm c} = e_{\Gamma} - \lambda_{\rm c} \left(\frac{p'}{p_{\rm at}}\right)^{\zeta} \tag{8}$$

where  $e_{\Gamma}$ ,  $\lambda_c$  and  $\xi$  are critical state parameters, and they are related to particle characteristics [40]. The degree of fabric anisotropy is assumed to be controlled by the norm of the deviatoric part of fabric tensor ( $F^d$ ), with initial value of  $F_{in}$ , and the fabric evolution law is described by

$$dF^d = c_f \left(1 - F^d\right) \left| d\varepsilon_q^p \right| \tag{9}$$

where  $c_f$  controls the magnitude of evolution, and  $d\varepsilon_q^p$  is plastic deviatoric strain increment.

To calibrate the model parameters, a set of undrained triaxial tests was conducted on Toyoura sand, a uniform quartz sand that has been widely used in the laboratory for various purposes. Its mean particle size  $(D_{50})$  is 0.203 mm, and the coefficient of uniformity  $(C_u)$  is 1.474, with 0% fines content; its particles are subrounded to subangular. All the specimens were prepared by the moist tamping method and were carefully saturated. Figure 3a and b shows the results of four triaxial tests in terms of stress path



Fig. 3 Undrained shear behaviour of isotropically consolidated sand specimens at varying void ratios:  $\mathbf{a}$  and  $\mathbf{b}$  test results;  $\mathbf{c}$  and  $\mathbf{d}$  model simulations

and stress-strain curve; the four specimens were isotropically consolidated to the same stress level but differing void ratios.

It can be seen that under otherwise identical conditions, a small variation in void ratio leads to a markedly different response, from non-liquefiable ( $e_0 = 0.878$ ) to highly liquefiable ( $e_0 = 0.946$ ). The behaviour in between these two cases ( $e_0 = 0.894$ ) is known as limited flow or partial liquefaction, being characterised by a peak strength followed by a limited period of strain softening and then a continuous dilation to high strength. The state corresponding to the local minimum strength, known as the quasi-steady state, is well captured. Using the results of the four triaxial tests, the model is calibrated, and the resulting parameters are given in Table 1. For comparison, the model responses are shown in Fig. 3c and d. Clearly, a good agreement is achieved for the range of responses. Figure 4 shows the critical state locus along with the experimental data in q-p'plane and e-p' plane. Note that all the simulations for the CSD, ACU and ICU stress paths in the present study are

Table 1	Model	parameters	calibrated from	experimental	data
Paramet	er		Variat	ole	Parame

Parameter	Variable	Parameter
Elasticity	$G_0$	225
	ν	0.25
Critical state parameter	$M_{\rm c}$	1.21
	$\lambda_{c}$	0.0225
	$e_{\Gamma}$	0.943
	ξ	0.6
Yield surface	m	0.01
Plastic modulus	$h_0$	24
	$c_h$	0.97
	$n^b$	1.1
Dilatancy	$d_0$	0.704
	$n^d$	3.5
Fabric evolution	$F_{\rm in}$	0.5
	$C_f$	20
	$e_A$	0.045



Fig. 4 Critical state locus (CSL) in a q-p' plane and b e-p' plane



**Fig. 5** CSD behaviour of sand at very loose state ( $e_0 = 0.960$ ,  $p_0' = 200$  kPa): **a** stress path; **b** stress–strain relation; **c** evolution of axial strain and **d** volumetric strain–axial strain relation

produced using the same set of model parameters, without any additional calibration or adjustment.

# 4 Behaviour under CSD and ACU stress paths

#### 4.1 Influence of void ratio under CSD stress path

Figure 5 shows the simulation result for a very loose specimen of Toyoura sand. The specimen is subjected to

isotropic consolidation to  $p_0' = 200$  kPa and  $e_0 = 0.960$ (stage O to A), followed by standard drained triaxial compression to the anisotropic state where q = 200 kPa (stage of A to B), and then subjected to a constant q path by decreasing p'. It is noted that the pattern of axial strain response (Fig. 5c) is quite similar to that observed in experiments. A marked feature of Fig. 5 is that q is perfectly maintained constant throughout the test (Fig. 5a and b). Using Hill's criterion, the state corresponding to vanishing second-order work is marked by triangles in the four plots. Clearly, large axial strain is about to accelerate at this state where the volumetric strain increment achieves zero (i.e.  $d\varepsilon_v = 0$ ). Note that this state is not at the point where the curvature of the axial strain curve is maximum but appears a bit earlier. The instability line in the q-p' plane (Fig. 5a) is markedly below the critical state line, confirming that the onset of instability is well inside the plastic limit.

Figure 6 shows the simulation result for a medium loose specimen ( $e_0 = 0.920$ ) under otherwise identical conditions. While the overall response is similar to that of the very loose specimen, the specimen exhibits a prolonged stage of dilative response upon decreasing the confinement

(Fig. 6d), and the onset of instability is closer to the point where the curvature is maximum (Fig. 6c). The instability line in the stress space is also located closer to the critical state locus, lending clear evidence for its state dependence. The results in Figs. 5 and 6 suggest that the resistance to flow failure can be significantly enhanced by decreasing a small amount of void ratio, and the behind mechanism is the dilatancy potential of the sand. To make the point clearer, Fig. 7 compares the evolution of volumetric strain and second-order work with mean effective stress of the two specimens. Although the two specimens share the same q and p' prior to applying the constant q stress path, the very loose specimen undergoes a shorter and slighter dilation than the medium loose specimen and achieves zero second-order work at a smaller axial strain.

#### 4.2 Influence of void ratio under ACU stress path

For better comparison, a pair of ACU tests is simulated for the two specimens as well, and the results are shown in Figs. 8 and 9, respectively. For the very loose specimen  $(e_0 = 0.960)$ , complete liquefaction is achieved, and its peak strength state exactly corresponds to vanishing



Fig. 6 CSD behaviour of sand at medium loose state ( $e_0 = 0.920$ ,  $p_0' = 200$  kPa): a effective stress path; b stress–strain relation; c evolution of axial strain and d volumetric strain–axial strain relation



Fig. 7 State dependence of CSD behaviour: a volumetric strain evolution and b second-order work evolution



Fig. 8 ACU behaviour of sand at very loose state ( $e_0 = 0.960$ ,  $p_0' = 200$  kPa): a stress path; b stress-strain relation; c evolution of axial strain and d volumetric strain-axial strain relationship

second-order work. The axial strain level at the onset of instability is very small due to the presence of an initial shear stress of as high as 200 kPa, confirming the experimental observations of Yang et al. [39]. By comparison, the medium loose specimen ( $e_0 = 0.920$ ) displays the phenomenon of partial liquefaction or limited flow (Fig. 9),

being characterised by a limited period of strain softening and then continuous dilation to higher strength. The quasisteady state or phase transformation state under anisotropic consolidation is well captured here, see Fig. 9b and c. Furthermore, the gradient of the instability line is found to be largely increased in the case of lower void ratio,



Fig. 9 ACU behaviour of sand at medium loose state ( $e_0 = 0.920$ ,  $p_0' = 200$  kPa): a stress path; b stress-strain relation; c evolution of axial strain and d volumetric strain-axial strain relationship

meaning that the instability line under the ACU conditions is not unique either, but is also state dependent.

#### 4.3 CSD stress path versus ACU stress path

Comparing the results in Figs. 8a and 9a with that in Figs. 5a and 6a, respectively, one may note that the gradient of the instability line under the ACU stress path is nearly the same with its counterpart under the CSD stress path. To confirm this finding, Fig. 10 presents results for two additional specimens which are isotropically consolidated to  $e_0 = 0.960$  and  $p_0' = 200$  kPa, brought to the same anisotropic condition (q = 150 kPa) and then subjected to the CSD and ACU stress path, respectively. The state corresponding to vanishing second-order work is marked by triangles for the CSD test and by circles for the ACU test. It becomes evident that the two points for vanishing second-order work lie on the same line passing through the stress origin (Fig. 10a). Nevertheless, it should be noted that the two specimens display different patterns of secondorder work evolution (Fig. 10c), and the specimen under the ACU stress path tends to lose stability earlier.

In Fig. 11, the responses under the CSD and ACU stress paths are compared in e-p' plane with reference to the critical state locus, for the case of very loose specimen  $(e_0 = 0.960, q = 200 \text{ kPa})$  and the case of medium loose specimen ( $e_0 = 0.920$ , q = 200 kPa). For the former, the specimen undergoes slight dilation in the initial stage as the result of decreasing confining stress, which is followed by a strong contractive response towards the critical state locus. When the specimen is subjected to the ACU stress path, the effective confining stress decreases all the way to the fully liquefied state under the constraint of constant volume. For the medium loose specimen under the constant q path, it undergoes a moderate contractive response upon the initiation of instability, arriving at a state below the critical state locus in the e-p' plane, and then displays a dilative response towards the critical state locus. Under the ACU stress path, the medium loose specimen follows a constant volume path to the left of the critical state locus, and the path then reverses towards the critical state locus. The reversal point corresponds to the phase transformation state. The results shown in Fig. 11 lend support that the critical state concepts work well for both CSD and ACU stress paths and for a range of loose states.



Fig. 10 Model simulations of CSD and ACU tests at very loose state ( $e_0 = 0.960$ ,  $p_0' = 200$  kPa, q = 150 kPa): **a** stress path; **b** stress–strain relation; **c** evolution of axial strain and **d** evolution of second-order work



Fig. 11 Comparison of CSD and ACU behaviour in e-p' plane with reference of critical state line: **a** very loose state ( $e_0 = 0.960$ , q = 200 kPa) and **b** medium loose state ( $e_0 = 0.920$ , q = 200 kPa)

So far, the simulations are for specimens looser than critical state. It is of interest to further explore the response characteristics of specimens denser than critical state in the CSD and ACU stress paths. In doing that, two additional specimens are isotropically consolidated to  $e_0 = 0.910$  and

 $p_0' = 100$  kPa and then brought to the same anisotropic condition (q = 100 kPa). The results are generated using the same set of parameters and are shown in Fig. 12. Along the ACU stress path, the specimen exhibits the limited flow behaviour, characterised by a limited strength reduction



Fig. 12 CSD and ACU behaviour of sand at medium dense state ( $e_0 = 0.910$ ,  $p_0' = 100$  kPa): **a** q-p' plane; **b** stress–strain relation; **c** e-p' plane and **d** volumetric strain–axial strain relation

and subsequently a strong dilative response towards the critical state. For the CSD stress path, the specimen undergoes three stages of response in the e-p' plane, all in the zone below the critical state locus: The first is a dilative response caused by reducing the confining stress, then a very brief and slight contractive response and finally a strong dilation to the critical state with a nearly constant p'. In the q-p' plane, the two specimens share the same instability line which is located very close to the critical state line. Comparing Fig. 12a with Fig. 10a, the influence of density on the location of the instability line is evident. One can expect that when the post-consolidation void ratio is further reduced, the instability line will tend to emerge to the left of the critical state line in the q-p' plane. In such case, there would be no strength reduction or strain softening under the ACU stress path. This implies that the instability of sufficiently dense soils is the mode of dilative failure rather than the diffuse mode for loose soils. In other words, the flow-type failure would not be possible for dense soils. Further discussion of the instability of dilative mode is worthwhile but is outside the scope of this study. Results will be presented in future.

#### 4.4 State parameter-dependent instability line

It is now interesting to examine whether the relationship in Eq. (1), derived from isotropic consolidation undrained tests, is valid regardless of stress paths applied. In doing so, a series of ICU tests is simulated along with the series of CSD and ACU tests. Two examples are shown in Fig. 13: One case is for complete liquefaction and the other for partial liquefaction. The results well confirm that the instability line varies with the post-consolidation state, although the difference in void ratio of the two specimens is only 0.04. In Fig. 14a, the stress ratio at onset of instability is plotted against initial state parameter ( $\psi_0$ ) for all CSD, ACU and ICU simulations (Table 2). Here, the initial state refers to the point after anisotropic consolidation and prior to shearing. It is clear that a unique relationship in the form of Eq. (1) fits the data points quite well, resulting in parameters A and B of -4.85 and 1.08, respectively. If the state parameter corresponding to the onset of instability is used, the goodness of fit can be further improved, as shown in Fig. 14b. The significance of the finding in Fig. 14 is that the state dependence of the instability line is not affected by stress path, and the critical state concepts work



Fig. 13 State-dependent behaviour of sand in ICU test: a stress path and b stress-strain relation



Fig. 14 Unique relationship between stress ratio at onset of instability and state parameter: a initial state parameter and b current state parameter

consistently well. For practical applications, one can use conventional undrained triaxial tests on isotropically consolidated specimens to determine the relationship in Eq. (1) and then apply it to the CSD and ACU stress paths. Moreover, it is worth mentioning that the instability line so determined can also be applied to cyclic loading conditions because of the correspondence between monotonic and cyclic loading behaviours. An excellent example is provided in Fig. 15 where results of two specially designed tests on Toyoura sand are put together: One specimen was subjected to isotropic consolidation and then undrained monotonic loading to flow failure, and the other was anisotropically consolidated to a similar state in terms of void ratio and mean effective stress and then subjected to cyclic loading. One can see that when the cyclic stress path touched on the instability line (i.e. flow liquefaction line) determined from the monotonic loading test, flow failure was initiated in the form of abrupt runaway deformation. It should be noted that the undrained cyclic tests are often

conducted under stress controlled, whereas the monotonic loading tests are often conducted under strain controlled to obtain reliable data after the onset of instability. For stresscontrolled tests, the measurements after the onset of instability may not be reliable and thus should be used with caution in the evaluation of critical states.

## **5** Discussion

Several studies in the literature have concluded that the onset of instability under drained constant shear conditions does not coincide with that under undrained conditions. While the possible reasons for the disagreement have been discussed in the preceding sections, further clarification and verification are desirable. Recalling the incremental stress–strain relationship in Eq. (2), one can readily obtain the following relation for the CSD stress path:

Table 2 Series of CSD, ACU and ICU stress path simulations

Series	$e_0$	$p_0'$ (kPa)	q (kPa)
CSD1	0.920	200	200
CSD2	0.960	200	200
CSD3	0.990	200	200
CSD4	0.910	100	100
CSD5	0.950	300	300
CSD6	0.960	200	150
ACU1	0.920	200	200
ACU2	0.960	200	200
ACU3	0.990	200	200
ACU4	0.910	100	100
ACU5	0.950	300	300
ACU6	0.960	200	150
ICU1	0.910	300	0
ICU2	0.980	500	0
ICU3	0.950	500	0
ICU4	0.895	300	0
ICU5	0.910	100	0
ICU6	0.916	267	0

$$\frac{d\varepsilon_{\rm v}}{d\varepsilon_q} = -\frac{C_{qq}}{C_{qp}} = \omega \tag{10}$$

Apparently, vanishing second-order work requires  $C_{qq} = 0.$ 

On the other hand, the incremental stress–strain relationship in Eq. (2) can be written in an alternative form as follows:

$$\begin{pmatrix} d\varepsilon_{\rm v} \\ dq \end{pmatrix} = \begin{bmatrix} C_{pp}^{-1} & -C_{pq}C_{pp}^{-1} \\ C_{qp}C_{pp}^{-1} & C_{qq} - C_{pq}C_{qp}C_{pp}^{-1} \end{bmatrix} \begin{pmatrix} dp \\ d\varepsilon_q \end{pmatrix}$$
(11)

For undrained conditions, the loss of stability is equivalent to the loss of uniqueness, which requires:

$$\det \begin{bmatrix} C_{pp}^{-1} & -C_{pq}C_{pp}^{-1} \\ C_{qp}C_{pp}^{-1} & C_{qq} - C_{pq}C_{qp}C_{pp}^{-1} \end{bmatrix} = 0$$
(12)

From Eq. (11), one can obtain  $C_{qq} = 0$ . Hence, the condition leading to instability under drained conditions is exactly the same with that under undrained conditions, regardless of the constitutive model chosen. Following the procedure of Yang and Li [38] and for the plasticity model used here, the condition for onset of instability can further be derived as follows:

$$\frac{3GK_p p' - 3KG\eta d \operatorname{sgn}\left(d\varepsilon_q^p\right)}{3G + K_p p' - dK\eta \operatorname{sgn}\left(d\varepsilon_q^p\right)} = 0$$
(13)

where G and K are elastic shear modulus and bulk modulus, respectively, and both depend on the density and confining stress of the soil. This implies the onset of instability is state dependent, which is well consistent with experimental findings [37, 39].

It should be mentioned that the present study has been limited to the triaxial space. This is because most of the experimental data on soil behaviour have been obtained from triaxial tests, primarily triaxial compression tests. To study instability behaviour under more complex stress paths beyond the conventional triaxial space, the formulations need to be generalised to multiaxial conditions. This can be done without much difficulty. Nevertheless, the lack of quality experimental data under such general stress paths for model calibration and for validation of simulations is always a concern. Other interesting issues related to



Fig. 15 Experimental evidence for linkage of monotonic and cyclic loading-induced flow liquefaction: a stress path and b stress-strain relation

instability include, for example the effect of pre-shearing history [14], the effect of fines [27, 41] and the effect of anisotropy [16, 17]. Further research in both experimental and numerical aspects is needed.

## 6 Summary and conclusions

The diverse views in relation to the initiation of instability and flow failure in granular soils under drained and undrained conditions and the conceptual difficulties brought by the discrepancies have been discussed from theoretical and experimental points of view. The main results and findings of the study are summarised as follows:

- (a) The constant shear drained (CSD) stress path is not a conventional stress-controlled test where dq and dp' are prescribed, but is a special one with the volume change under certain control. Overlooking this fact in theoretical analysis may result in the conclusion that loss of instability along the CSD stress path occurs when the plastic hardening modulus becomes zero during loading. This conclusion is misleading and is not supported by the experimental observation that the loss of instability in loose granular soils occurs well within the plastic limit and is a diffuse mode.
- (b) The CSD stress path in laboratory experiments involves several inherent difficulties and uncertainties such that a strictly constant shear stress often cannot be achieved, and the experimental data are often noisy. This is considered the main reason why different criteria were put forward to determine the onset of instability in experimental studies, and why diverse views arose. Owing to the nature of the CSD tests, the response recorded after the onset of instability in experiments may not be reliable and thus should be used with caution in the evaluation of critical states.
- (c) The specifically designed simulations for CSD and ACU stress paths under a range of initial states indicate that the condition leading to instability along the CSD stress path is coincident with that for the ACU stress path, provided that they share the same initial state prior to shearing. While sharing the same instability line in the stress space, they, however, display different evolution patterns of the secondorder work, and loss of instability tends to occur earlier under the ACU stress path.
- (d) For very loose sand that exhibits complete liquefaction under undrained shear conditions, its response to the CSD stress path is characterised by two stages: Initial dilation caused by reducing the confining

stress and subsequent contraction following the onset of instability. For medium loose to medium dense sand that exbibits partial liquefaction or partial strain softening under undrained conditions, the response along the CSD stress path consists of three stages: first dilation, then contraction and finally dilation again towards the critical state.

(e) The state dependence of the instability line is confirmed for CSD, ACU and ICU stress paths using systematic simulations for a range of initial states. This state dependence can be characterised using the initial state parameter defined with reference to the critical state locus. The relationship between the gradient of the instability line and the state parameter can be well described by an exponential function, which is nearly unique for a given soil regardless of the CSD, ACU or ICU stress path. Theoretically, it has shown that the conditions for instability under these different stress paths are exactly the same, regardless of the constitutive models chosen.

Further validation of the above findings using experimental data of both drained and undrained stress paths and extending the study from the compressional regime to the extensional regime are worthwhile. The uncertainty and difficulty involved in the laboratory CSD test, however, needs to be carefully addressed.

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## Declarations

**Conflict of interest** The authors declare no competing interests. All data supporting the findings in this paper are available from the corresponding author upon reasonable request.

# References

- Alipour MJ, Lashkari A (2017) Prediction of sand instability under constant shear drained paths. In: Papamichos E, Papanastasiou P, Pasternak E, Dyskin A (eds) Bifurcation and degradation of geomaterials with engineering applications. Springer, Berlin
- Anderson SA, Sitar N (1995) Analysis of rainfall induced debris flow. J Geotech Eng ASCE 121(7):544–552
- Been K, Jefferies MG (1985) A state parameter for sands. Géotechnique 35(2):99–102
- 4. Borja RI (2006) Conditions for liquefaction instability in fluidsaturated granular soils. Acta Geotech 1:211–224
- Brand EW (1981) Some thoughts on rainfall induced slope failures. In: Proceedings of the 10th international conference on soil mechanics and foundation engineering, pp 373–376
- Casagrande A (1975) Liquefaction and cyclic deformation of sands, a critical review. In: Proceedings of the of 5th Pan-American conference on soil mechanics and foundation engineering, vol 5, pp 79–133

- Chandler RJ, Tosatti G (1995) The Stava tailings dams failure, Italy, July 1985. Geotech Eng ICE 113:67–79
- Chen H, Lee CF, Law KT (2004) Causative mechanisms of rainfall-induced fill slope failures. J Geotech Geoenviron Eng ASCE 130:593–602
- Chu J, Leong WK (2002) Effect of fines on instability behaviour of loose sand. Géotechnique 52(10):751–755
- Chu J, Leong WK, Loke WL, Wanatowski D (2012) Instability of loose sand under drained conditions. J Geotech Geoenviron Eng ASCE 138(2):207–216
- Dafalias YF, Manzari MT (2004) Simple plasticity sand model accounting for fabric change effects. J Eng Mech ASCE 130(6):622–634
- Daouadji A, AlGali H, Darve F, Zeghloul A (2010) Instability in granular materials: experimental evidence of diffuse mode of failure for loose sands. J Eng Mech ASCE 136(5):575–588
- Darve F, Sibille L, Daouadji A, Nicot F (2007) Bifurcations in granular media: macro-and micro-mechanics approaches. C R Méch 335(9):496–515
- Doanh T, Finge Z, Boucq S, Dubujet P (2003) Histotropy of Hostun RF loose sand. In: Modern trends in geomechanics. Springer
- Dong Q, Xu C, Cai Y, Juang H, Wang J, Yang Z, Gu C (2016) Drained instability in loose granular material. Int J Geomech 16(2):04015043
- Dong T, Zheng Y, Liang K, Liu C (2022) Shear strength and shear bands of anisotropic sand. Acta Geotech 17:2841–2853
- Druckrey AM, Imseeh WH, Alshibli KA (2022) Experimental evaluation of the anisotropic critical state theory for sand using 3D fabric evolution data of triaxial experiments. Acta Geotech 17:2193–2211
- Fourie AB, Tshabalala L (2005) Initiation of static liquefaction and the role of K<sub>0</sub> consolidation. Can Geotech J 42(3):892–906
- 19. Hill R (1958) A general theory of uniqueness and stability in elastic-plastic solids. J Mech Phys Solids 6:236–249
- Imposimato S, Nova R (1998) An investigation on the uniqueness of the incremental response of elastoplastic models for virgin sand. Mech Cohesive Frict Mater 3(1):65–87
- Ishihara K (1993) Liquefaction and flow failure during earthquakes. Géotechnique 43(3):351–451
- 22. Junaideen SM, Tham LG, Law KT, Dai FC, Lee CF (2010) Behaviour of recompacted residual soils in a constant shear stress path. Can Geotech J 47(6):648–661
- Lade PV (1993) Initiation of static instability in the submarine Nerlerk berm. Can Geotech J 30(6):895–904
- Li XS, Dafalias YF (2012) Anisotropic critical state theory: role of fabric. J Eng Mech ASCE 138(3):263–275
- Monkul MM, Yamamuro JA, Lade PV (2011) Failure, instability, and the second work increment in loose silty sand. Can Geotech J 48:943–955

- Poulos SJ, Castro G, France JW (1985) Liquefaction evaluation procedure. J Geotech Eng Div ASCE 111(6):772–792
- Rahman MM, Dafalias YF (2022) Modelling undrained behaviour of sand with fines and fabric anisotropy. Acta Geotech 17:2305–2324
- Ramos AM, Andrade JE, Lizcano A (2012) Modelling diffuse instabilities in sands under drained conditions. Géotechnique 62(6):471–478
- Sasitharan S, Robertson PK, Sego DC, Morgensterm NR (1993) Collapse behavior of sand. Can Geotech J 30(3):569–577
- Skopek P, Morgenstern NR, Robertson PK, Sego DC (1994) Collapse of dry sand. Can Geotech J 31(6):1003–1008
- Sladen JA, D'Hollander RD, Krahn J (1985) The liquefaction of sands, a collapse surface approach. Can Geotech J 22(4):564–578
- Vaid YP, Chung EFK, Kuerbis RH (1990) Stress path and steady state. Can Geotech J 27:1–7
- Verdugo R, Ishihara K (1996) The steady state of sandy soils. Soils Founds 36(2):81–91
- Wan RG, Pinheiro M, Guo PJ (2011) Elastoplastic modellingof diffuse instability response of geomaterials. Int J Numer Anal Methods Geomech 35(2):140–160
- Wang G, Sassa K (2001) Factors affecting rainfall-induced flow slides in laboratory flume tests. Géotechnique 51(7):587–599
- Yamamuro JA, Lade PV (1997) Static liquefaction of very loose sand. Can Geotech J 34(6):905–917
- Yang J (2002) Non-uniqueness of flow liquefaction line for loose sand. Géotechnique 52(10):757–760
- Yang J, Li XS (2004) State-dependent strength of sands from the perspective of unified modeling. J Geotech Geoenviron Eng ASCE 130(2):186–198
- Yang J, Liang L, Chen Y (2022) Instability and liquefaction flow slide of granular soils: the role of initial shear stress. Acta Geotech 17(2):1–15
- Yang J, Luo XD (2015) Exploring the relationship between critical state and particle shape for granular materials. J Mech Phys Solids 84:196–213
- 41. Yang J, Wei LM (2012) Collapse of loose sand with the addition of fines: the role of particle shape. Géotechnique 62(12):1111–1125
- Yang ZX, Li XS, Yang J (2008) Quantifying and modelling fabric anisotropy of granular soils. Géotechnique 58(4):237–248

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