RESEARCH ARTICLE

Metal extractability patterns to evaluate (potentially) mobile fractions in periurban calcareous agricultural soils in the Mediterranean area—analytical and mineralogical approaches

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Abstract A set of periurban calcareous agricultural Mediterranean soils was spiked with a mixture of Cd, Cu, Pb and Zn at two levels within the limit values proposed by current European legislation, incubated for up to 12 months, and subjected to various one-step extraction procedures to estimate mobile (neutral salts) and potentially mobile metal fractions (complexing and acidic extraction methods). The results obtained were used to study metal extractability patterns according to the soil characteristics. The analytical data were coupled with mineralogical investigations and speciation modelling using the Visual Minteq model. The formation of soluble metal-complexes in the complexing extracts (predicted by the Visual Minteq calculations) led to the highest extraction efficiency with complexing extractants. Metal extractability patterns were related to both content and composition of carbonate, organic matter, Fe oxide and clay fractions. Potentially mobile metal fractions were mainly affected by the finest soil fractions (recalcitrant

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organic matter, active lime and clay minerals). In the case of Pb, scarce correlations between extractable Pb and soil constituents were obtained which was attributed to high Pb retention due to the formation of $4PbCO_3 \cdot 3PbO$ (corroborated by X-ray diffraction). In summary, the high metal proportion extracted with complexing agents highlighted the high but finite capacity to store potentially mobilizable metals and the possible vulnerability of these soils against environmental impact from metal accumulation.

Keywords Calcareous soil · One-step extractions · Speciation modelling · Mineralogical composition · Soil pollution

Introduction

Soil degradation is a serious problem in Europe, mainly driven or exacerbated by human activity. The number of potentially contaminated sites in the European Union-25 is estimated at approximately 3.5 million, with soil pollution from diffuse sources being recognized as one of the major soil threats by the European Union Soil Thematic Strategy (European Commission 2006). This is especially true in metropolitan areas subjected to risks from pollution due to a strong industrial development and spreading urbanization, leading to an increase of "urban elements" in soils (Ba, Cd, Co, Cu, Mg, Pb, Sb, Ti and Zn) (de Miguel et al. 1997). This scenario is typical of periurban agricultural soils which are affected and likely to become more so in the future unless sustainability criteria are applied (Iram et al. 2012; Peris et al. 2007). Based on these premises, there is a potential risk of metal contamination of groundwater and of accumulation in food crops which is directly related to the mobility of metals in soils.

The mobility of metals can be greatly decreased in calcareous Mediterranean soils where low rainfall and high evapotranspiration lead to metal accumulation in the first few centimetres of the soil. This metal accumulation capacity, enhanced by the high metal retention provided by carbonates, can however be affected by other capacitycontrolling properties such as texture, content and type of organic matter, soil pH-redox conditions and the content of oxides of Fe, Al and Mn. Metals accumulated and retained could be potentially mobilizable by organism activity (Sayyad et al. 2010) and/or by changes in the conditions in the environment (Bolan et al. 2003). The high but finite capacity to store potentially mobilizable metals suggests that these soils should be considered to be the most vulnerable (Batjes 2000). This high vulnerability which allows metals to move from one environmental echelon to another, combined with the complexity of the interaction between metals and the constituents of calcareous soils, highlights the importance of the study of the soil properties affecting the mobility (and thus the availability) of metals in these soils.

It is generally accepted that to assess the mobility and availability of metals in soils it is not sufficient merely to analyse the total metal content in soils, nor is this a useful tool for determining potential risks, as soils are dynamic systems. A very delicate equilibrium exists among the metal pseudo-total content (inactive and inert), the metal mobile fraction (effective soluble, very active, bioavailable) and the metal mobilizable/potentially mobile fraction (potentially bioavailable, leachable and partly active) (Gupta et al. 1996). Using chemical-extracting media of different strengths, it is possible to differentiate between these fractions. The combination of indirect methods (analytical data from metal extractions and speciation modelling of the soil extracting solution) with direct methods (soil mineralogical investigations) is a useful tool for elucidating soil properties and constituents as well as the chemical processes that affect metal extractability, mobility and availability (Ettler et al. 2007; Pérez-Esteban et al. 2013; Sipos et al. 2008).

The objective of the study was to investigate the mobile and potentially mobile fractions of Cd, Cu, Pb and Zn using a set of extraction methods of different strengths, speciation modelling using Visual Minteq Model, and mineralogical investigations in several metal-spiked periurban calcareous agricultural soils representative of the Mediterranean area. Since the role of soil constituents in metal desorption processes is not well defined for these soils, we placed particular emphasis on the soil constituents and properties affecting metal extractability patterns.

Materials and methods

Study area, soil characteristics and sampling

The study area is part of a periurban axis that combines agricultural activity and the main residential and industrial uses in the Madrid region; it is exposed to atmospheric deposition from transportation and industrial activities. The soil samples come from different plots in the "El Encín" Agricultural Research Station (Alcalá de Henares, Madrid, Spain), at an altitude of 588 m. located on the Henares River on guaternary sediments (IGME 1990). These alluvial sediments have led to an ancient calcaric Fluvisol (Moreno Merino 1998) which presents Anthric characteristics today (FAO 2006) mainly as a result of agricultural use. The average annual temperature is 13 °C, average annual rainfall is 401 mm year⁻¹ and potential evaporation is about 760 mm year⁻¹. The site is typical of a Mediterranean pluviseasonal-oceanic bioclimate on an upper meso-Mediterranean low dry bioclimatic belt (Rivas-Martínez and Rivas-Saenz 2009).

A set of ten agricultural soil samples (M1, M2, M3, M4, M5, L1, L2, VL1, VL2 and VL3), differing in their carbonate content, amount of organic matter and textural class, was selected according to information from previous studies (de Santiago-Martín et al. 2013; Lafuente et al. 2008). Sampling was done at randomly selected points. To avoid potential bias, 30–40 kg was taken from each sampling point (0– 30 cm) and homogenized. Soil samples were air-dried and passed through a 2-mm sieve.

Experimental design

The environmental impact from metal accumulation in periurban agricultural areas was simulated by the addition of a multi-elemental salt solution of metals to the soil samples. Cd, Cu, Pb and Zn were selected owing to their different speciation, mobility and extractability in soils. Three containers (40 cm wide \times 59 cm long \times 21 cm high) of 10 kg each were set for each sample: one un-spiked sample with the addition of distilled water, and the other two spiked at two different concentration levels using nitrate salts of metals in aqueous solution: low level (Tt1) $(3 \text{ mg kg}^{-1} \text{ of } \text{Cd}+140 \text{ mg kg}^{-1} \text{ of } \text{Cu}+300 \text{ mg kg}^{-1} \text{ of}$ Pb+300 mg kg⁻¹ of Zn) or high level (Tt2) (20 mg kg⁻¹ of $Cd+875 mg kg^{-1} of Cu+600 mg kg^{-1} of Pb+2000 mg kg^{-1}$ of Zn) within the limits proposed by current European legislation (Directive 86/278/EEC). Each soil sample and the corresponding solution was mixed and left to equilibrate for a period of 12 months at room temperature without cover or drainage. During this equilibration period, the soils were air-dried, mixed and rewetted with distilled water in cycles of about 2 weeks, in order to favour metal redistribution processes into the soil matrix. The appropriate amount of water was added to bring the soil samples to the each estimated field capacity (Jalali and Khanlari 2008). Distribution processes play a key role in determining the extractability and availability of metals in soils (McLaughlin 2001).

At the end of the equilibration period (12 months), duplicates were randomly removed from each un-spiked and metal-spiked soil sample for extraction of the metals by means of one-step extraction methods. Different extracting solutions have been used to assess mobile and potentially mobile fractions of metals, both in equilibrium with inactive and inert fractions (Gupta et al. 1996). The mobile fraction was estimated with 0.01 M CaCl₂, 1 M MgCl₂, 0.1 M NaNO₃ and 1 M NH₄NO₃ methods, and the potentially mobile fraction was estimated with 5 mM DTPA, 0.05 M EDTA, 0.5 M HNO₃, 0.11 M HAc, 10 mM LMWOA (a mixture of low molecular weight organic acids) and 1 M NH₄Ac methods. The analytical data were coupled with mineralogical investigations (Sipos et al. 2008), measurement of the pH in extracts and speciation modelling using the Visual Minteq model (Ettler et al. 2007; Pérez-Esteban et al. 2013). In order to evaluate the soil sorption capacity, a sorption test was conducted prior to the incubation experiment and the distribution coefficients were calculated.

Analytical methods

All chemicals were obtained from analytic grade reagents of Merck (Germany). All glassware used was pre-washed with an aqueous solution of $HNO_3 0.1$ % for 24 h and rinsed with de-ionized type I water.

Soil physicochemical parameters

According to ISRIC (2002) methods, the following parameters were determined: soil pH in a 1:2.5 soil to water ratio, equivalent CaCO₃ (ECC) according to the acid neutralization method, total organic C by the Walkley-Black wet oxidation procedure, particle-size distribution by Robinson's pipette method, cation exchange capacity (CEC) by the ammonium acetate method, and crystalline and amorphous Fe and Mn oxide contents by dithionitecitrate extraction, in addition to acid oxalate extraction. The active equivalent CaCO₃ finely divided or "active lime" (AL) was determined with NH₄-oxalate as described by Drouineau (1942). We used the two-step acid hydrolysis procedure with H₂SO₄ to determine the recalcitrant pool of organic matter (RP) (Rovira and Vallejo 2000). Total Ca, Mg, Fe, Cd, Cu, Pb, and Zn contents of the soil samples were determined after wet digestion (150 °C, 6 h) with a mixture of nitro-perchloric-hydrofluoric acids under highpressure conditions using PHAXE 2000 reactors (SISS 1985). Fe, Mn, Ca, Mg, Cd, Cu, Pb and Zn concentration in the corresponding extracts was quantified by atomic absorption spectroscopy, AAS (Analytikjena NovAA 300). All samples were extracted and analysed in duplicate. Quantification limits in milligrams per litre were: Ca=1, Mg=0.05, Fe=0.5, Mn=0.2, Cd=0.2, Cu=0.2, Pb=0.5 and Zn=0.1.

Soil mineralogical analyses

Mineral composition of soil samples was examined by Xray diffraction (XRD) using an EQ 0434520 31 02 (X'Pert MPD) diffractometer with Cu K α radiation operated at 45 kV and 40 mA. All XRD patterns were recorded with a dwell time of 1 s and 0.04° 2 θ step. Un-spiked and metal spiked (Tt1 and Tt2 levels) soil samples (<2 mm) were examined on randomly oriented powders. Separation of the clay fraction of un-spiked soils (<2 µm) was performed by sedimentation in aqueous suspension previous NH₄OH dispersion and posterior H₂O₂ treatment. The clay fraction in air-dried samples after saturation with Mg was calcinated at 550 °C for 2 h and after an ethylene glycol solvation. Abundance of soil minerals in both fine earth and clay fractions was assessed semi-quantitatively (Bish 1994).

One-step extraction methods of metals

The procedures to determine extractable metals are described as follows: 0.01 mol L^{-1} CaCl₂ solution, 1:5 w/v for 2 h (Van Ranst et al. 1999); 1 mol L^{-1} MgCl₂ solution, 1:8 w/v for 1 h (Tessier et al. 1979); 0.1 mol L^{-1} NaNO₃ solution, 1:2.5 w/v for 2 h (Gupta and Aten 1993); 1 mol L⁻¹ NH₄NO₃ solution, 1:2.5 w/v for 2 h (DIN 1995; Legislation Germany); $0.005 \text{ mol } \text{L}^{-1}$ diethylene triamine pentaacetic acid (DTPA) solution in 0.01 mol L^{-1} CaCl₂ solution and 0.01 mol L^{-1} triethanolamine (TEA), 1:2 w/v for 2 h (Lindsay and Norwell 1978); 0.05 mol L^{-1} ethylene diamine tetra-acetic acid (EDTA) solution, 1:10 w/v for 1 h (Quevauviller et al. 1996); 0.5 mol L⁻¹ HNO₃ solution, 1:5 w/v for 30 min (Van Ranst et al. 1999); 0.11 mol L^{-1} CH₃COOH solution (HAc), 1:40 w/v for 16 h (Rauret et al. 1999); 0.01 mol L^{-1} low molecular weight organic acids (LMWOA) solution consisting of acetic, lactic, citric, malic and formic acids with a molar ratio of 4:2:1:1:1, 1:10 w/v for 16 h (Feng et al. 2005); and 1 mol L⁻¹ CH₃COONH₄ solution (NH₄Ac) buffered by HAc at pH 7, 1:30 w/v for 4 h in column (Van Ranst et al. 1999). In all cases, except for NH₄Ac extractions, the samples and the extraction solution were shaken in a vibrator agitator (Vibromatic, Selecta) at 400 opm. The supernatant of each extraction was centrifuged at 3,500 rpm for 15 min and then filtered (low ash filters, 5–7 μ m). Dilutions were made with the corresponding extraction solution. Cd, Cu, Pb and Zn concentration in the extracts was quantified by AAS.

The computer program Visual Minteq v.3.0 (Gustafsson 2011) was used to predict the metal speciation in the extracts of three representative soil samples (M1, L1 and VL1 soils). We used the database that comes by default in the Visual Minteq program. The measured values of pH, metal concentration in extracts and the concentration of the extractant applied were used as the input data. Cd, Cu, Pb, and Zn were simultaneously entered into the code, and the calculations were performed for each individual extract.

Sorption capacity of metals

The metallic solution and each soil sample (1:2 w/v) were mechanically shaken for 24 h, and the supernatant of each extraction was centrifuged (3,500 rpm, 15 min) and then filtered. Concentration of Cd, Cu, Pb and Zn in the supernatant was quantified by AAS. The distribution coefficients (K_d) of the metals were calculated (Lafuente et al. 2008) using the equation: $K_d = (M_{sorbed})/(M_{solution})$, where M_{sorbed} is the amount of sorbed metal per unit weight of soil (milligrammes per kilogramme) and $M_{solution}$ is the amount of metal in solution per unit volume of liquid (milligrammes per litre).

Statistical analysis

Significance of differences of the means (n=10) of the relative metal extractability was investigated by means of one-way ANOVA using a post hoc test (Tukey). The homogeneity of variances was verified by Levene test. Correlation coefficients were calculated for metal-spiked (Tt1 and Tt2 level) soil samples (n=10), for each contamination level) to relate both the Kd values and the amount of metals extracted with the different methods employed to the soil physicochemical parameters (Pearson's correlation test) and the mineralogical composition (Spearman's correlation test). Since mineralogical properties are semiquantitative variables, we used a nonparametric correlation test, the Spearman's test. Analyses were conducted using SPSS (Statistical Package for the Social Sciences) v.17 (SPSS, Inc.) software.

Results and discussion

Soil characteristics

Soil physicochemical parameters

The main results of the physicochemical analyses from unspiked soils are shown in Table 1. All soil samples showed pH values above 8. The ECC content ranged from moderate (M1, M2, M3, M4 and M5 soils) to low (L1 and L2 soils) and very low (VL1, VL2 and VL3 soils). Differences were

	TACHOH	OIBallic	IFACHOR			OXIDE ILECTION	Faruci	e-size a	ISUDDUDO	ц	CEC	lotal	content			
kg ⁻¹	AL	TOC g kg ⁻¹	MO	RP % TO	Č C/N	Cry-Fe g kg ⁻¹	CS g kg ⁻¹	FS	Silt	Clay	${\rm cmol_c}~{\rm kg}^{-1}$	Fe g kg [_]	¹ Ca	Mg	Cu mg kį	Pb
06	77	9	10	69	8	8	111	603	126	161	7	18	94	11	8	14
48	53	12	20	48	12	6	114	569	124	193	6	19	94	11	12	56
25	33	18	31	99	10	12	23	592	215	170	11	26	41	11	10	24
18	23	15	26	54	11	11	95	605	146	154	6	20	36	6	11	24
17	39	10	17	57	14	8	166	567	129	139	6	17	66	6	15	24
06	21	18	32	72	11	13	159	590	78	172	7	24	40	7	12	26
00	42	8	13	75	10	10	112	462	167	259	13	22	34	6	6	22
2	15	12	21	71	11	11	45	459	168	328	19	26	10	9	10	24
7	15	12	21	74	11	12	66	386	172	344	20	27	13	9	13	21
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 Table 1
 Physicochemical parameters of the un-spiked soil samples

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observed in organic matter (OM) content and particle-size distribution within each group of soil samples. Thus, the OM content ranged from low to high, in all cases showing a high proportion of RP. The textural classes of soils ranged from sandy-loam to sandy-clay loam, with fine sand being the most abundant size fraction. Crystalline Fe oxide content was low (~10 g kg⁻¹). In general, the total Ca and Mg content varied in line with the carbonate fraction. The total Cu, Pb and Zn content was similar to that obtained by other authors for agricultural soils in the Mediterranean area (Jiménez Ballesta et al. 2010; Micó et al. 2006; Peris et al. 2007) and in no case exceeded the levels established by the European Union (Directive 86/278/EEC). The total Cd content was lower than the quantification limit.

Mineralogical composition

Examinations by XRD of the un-spiked soil samples indicated high contents of quartz and phyllosilicates in varying proportions (Table 2). The content in feldspars was always very low. Calcite was the predominant mineral in the carbonate group. Dolomite was not detected in L2, VL1 and VL3 soil samples. The analysis of XRD of the metal-spiked soils showed the presence of lead oxide carbonate (4PbCO₃·3PbO; JCPDS No. 00-017-0732), an intermediate oxycarbonate confirmed by the reflections at 0.364, 0.286, 0.267, 0.215 and 0.165 nm in samples originally characterized as containing calcite and dolomite in their composition (mainly M1, M2 and M3 soils at Tt2 level). Figure 1 shows the XRD patterns from un-spiked and metal-spiked (at Tt2 level) samples of M1 soil. Table 2 shows the mineralogical composition of the clay fraction of un-spiked soils. The most common phyllosilicate was illite, identified as a trioctahedral mineral based on the intensities of the 001/002 basal reflections (Douglas 1985). The minerals 1:1 of the kaolinite group displayed lower percentages. Vermiculite was occasionally present (M1 and VL1 soils).



Fig. 1 Patterns of X-ray diffraction of the randomly orientated from un-spiked and metal-spiked (Tt2 level) M1 soil sample. *Pb* lead oxide carbonate ($4PbCO_3 \cdot 3PbO$)

With regard to non-phyllosilicate minerals, it is worth noting the presence of quartz, calcite and dolomite in varying proportions.

Metal sorption capacity

The sorption test showed that the metal sorption capacity of the selected soils was very high, ranging from 94 to 100 % for Cu, Pb and Zn (data not shown), in agreement with the results obtained previously in the study area (de Santiago-Martín et al. 2013). The Cd sorption capacity, at about 64– 92 %, was lower than for other metals, and the lowest percentage was found in VL3 soil, with the highest coarse sand (CS) content. These results indicate a priori that these soils would be able to buffer an impact from metal contamination. Nevertheless, high metal extraction percentages were obtained with the one-step extraction methods used, as discussed below. The distribution coefficients (K_d) calculated highlighted that Pb was in all cases the most retained metal (Online resource 1), Cd was the metal with the lowest K_d value, and Cu and Zn presented intermediate values. In

Fraction	<2 mm	l					<2 µr	n				
Soil Sample	2:1 L	1:1 L	Q	F	С	D	V	Ι	K	Q	С	D
M1	++	Tr	++++	++	+++	++	++	+++	+	++	++++	tr
M2	+	tr	++++	+	+++	+	nd	++	+	++	++++	tr
M3	++	+	++++	++	++	tr	nd	++++	++	++++	++	+
M4	+	tr	++++	+	+	tr	nd	++++	+	++++	++	nd
M5	+	tr	++++	+	+	tr	nd	+++	+	+++	++++	nd
L1	++	tr	++++	++	++	tr	nd	++	+	++	++++	+
L2	+	tr	++++	++	++	nd	nd	+++	+	+++	++++	nd
VL1	++	tr	++++	++	tr	nd	+++	++++	+	++++	nd	nd
VL2	+	tr	++++	+	tr	tr	nd	++++	+	++++	nd	nd
VL3	+	tr	++++	++	nd	nd	nd	++++	+	++++	nd	nd

Table 2Mineralogical compo-sition in un-spiked soil samples

Number of plus sign is proportional to abundance: ++++ most abundant, + least abundant, *tr* trace and *nd* not detected

L layer, Q quartz, F feldspars, C calcite, D dolomite, V vermiculite, I illite, K kaolinite

order to elucidate what soil constituents affect metal sorption capacity, a Pearson's correlation analysis was performed. Table 3 shows that negative and significant correlations were obtained between sand fractions and K_d values in all cases, pointing out the key role of the finest mineral fractions in affecting metal sorption patterns.

Extraction efficiency and metal speciation in extracts

Values of Cd, Cu, Pb and Zn concentration obtained in single chemical extractions in the un-spiked and metalspiked soil samples are shown in Fig. 2, as a percentage of total metal. The absolute values of the extractable metals, expressed in milligrammes per kilogramme, are shown in Online resource 2. It was observed that the highest proportion of metals extracted in metal-spiked soil samples was in all cases with complexing and acidic extractants, and the lowest was with neutral salts. The high metal proportion extracted with complexing agents indicates a high potential metal mobility in these soils.

The Visual Minteq thermodynamic calculations were useful in determining the metal speciation in the extracts, which depends on the pH and chemical composition of the solutions (Online resource 3). The pH values of the extracts are shown in Table 4. In the case of acids and chelating agents, the pH values of soil extracts were lower in VL1 soil (with a lower ECC content) than in M1 and L1 soils.

The high proportion of metal extractable with EDTA (60-82 %) was attributed to its high chelation stability constant (Meers et al. 2007). Similarly, the pH range of our soils favoured the formation of soluble complexes with metals in the EDTA and DTPA extracts (up to 100 % of Me- $EDTA^{2-}$ and Me-DTPA³⁻). Nevertheless, it should be taken into account that the very high percentage of negatively charged complexes could lead to lower metal extraction in soils with a higher OM content, due to re-adsorption processes in the organic matter (Ettler et al. 2007; Peters 1999). In HAc extractions, the speciation was dominated by Me-Acetate⁺, Me-(Acetate)₂ and Me²⁺. The highest proportion of Me-(Acetate)₂ was in M1 soil, and the highest rate of free metal ion was found in VL1 soil, which was attributed to the different pH value of the soil extracts (Table 4). The speciation was more complex with LMWOA extractions. Metals were present as Me-Citrate⁻ (up to 97 %) due to the higher chelation stability constant of citric acid than of the other organic acids in the mixture employed. The lower stability constants of the citrate complexes with Cd and Pb than with Cu and Zn (Barton and Abadía 2006) resulted in a lower Cd and Pb extraction percentage at Tt2 level. Moreover, at this level, we observed the appearance of other Cd and Pb species in minor proportions (free metal ion, Me-Malate (aq), Me-Lactate⁺ and Me-Acetate⁺) due to different complexing abilities in response to pH (Qin et al. 2004).

Table	1 2413011 5 201101	מיוחוו החלווועות	TILS CALCULATION	octive transfer p	and secondination	ar paramensis	AUTO VALUE	o UI VU, VU, I		me novide-m	oamproo at 111	allu 112 17773	
Level	Parameter	ECC	AL	Tot-Ca	Tot-Mg	MO	RP	Cry-Fe	CS	FS	Silt	Clay	CEC
Tt1	Cd K _d	0.539	0.453	0.563	0.550	0.130	0.043	0.379	-0.452	0.070	0.287	0.133	-0.158
	Cu K _d	-0.295	-0.063	-0.195	-0.139	-0.563	0.316	-0.074	-0.171	-0.643^{*}	0.259	0.614	0.657^{*}
	Pb K _d	-0.033	-0.002	0.290	-0.174	-0.357	0.164	-0.349	0.319	-0.072	-0.09	-0.130	-0.130
	$\operatorname{Zn} \mathrm{K}_{\mathrm{d}}$	-0.351	-0.242	-0.219	-0.256	-0.180	0.200	0.370	-0.433	-0.729^{*}	0.452	0.798^{**}	0.747^{*}
Tt2	$Cd K_d$	0.346	0.507	0.394	0.652*	0.194	0.386	0.221	-0.777**	0.062	0.733*	0.148	0.015
	$Cu K_d$	0.100	0.296	0.113	0.351	0.071	0.213	0.306	-0.821^{**}	-0.175	0.669*	0.450	0.324
	Pb K_d	0.229	0.240	0.151	0.337	0.278	0.209	0.478	-0.646^{*}	-0.059	0.458	0.316	0.097
	${ m Zn}~{ m K_d}$	-0.160	-0.008	-0.129	0.282	0.278	0.239	0.476	-0.862**	-0.364	0.830^{**}	0.574	0.509
ECC ec exchang	uivalent CaCO ₃ , e capacity	AL active lime	e, Tot-Ca and	Tot-Mg total C	Da and Mg, Ol	<i>M</i> organic mat	ter, RP recal	citrant pool, (<i>Cry-Fe</i> crystallin	e Fe-oxide, CS	coarse sand, H	⁷ S fine sand, C	EC cation

p<0.05; p<0.01; p<0.01; p<0.001 (n=10), statistical significance at these probability levels



Fig. 2 Relative Cd, Cu, Pb and Zn extractability (%) with DTPA, EDTA, HNO₃, HAc, LMWOA, NH₄Ac, CaCl₂, MgCl₂, NaNO₃, and NH₄NO₃-methods in un-spiked (*white rectangle*) and metal-spiked soil samples at Tt1 (*grey rectangle*) and Tt2 levels (*black rectangle*). The *boxplots* show the lower quartile, the median and the upper quartile,

with whiskers extending to the most extreme data point (n=10). Different letters indicate significant differences between extraction methods at p < 0.05 after one-way ANOVA. Extractable Cd was lower than the quantification limit in un-spiked soil samples

In NH₄Ac extracts, acetic complexes were the dominant species of Cd, Pb and Zn, which were mainly present as Me-(Acetate)₃⁻, Me-(Acetate)₂ (aq) and Me-Acetate⁺. The dominant species in the case of Cu were amine complexes.

Very high amounts of metals were extracted with the HNO_3 method (Fig. 2). In general, free metal ion was the predominant species despite the large difference in the pH of the extractant solution from the various soils (Table 4). This could indicate that the main processes affecting metal desorption with the HNO_3 extraction method could be either the dissolution of some of the soil components, or the metal

displacement by excess H^+ from the exchangeable complex (Vidal et al. 2004).

Different patterns were obtained in the case of neutral salt extractions. Metal speciation in CaCl₂ and NaNO₃ extracts was dominated by the free ionic form, except for Cu and Pb in NaNO₃ extracts which were in a hydroxocomplex form at pH 7.8–8.1. Moreover, it should be noted that a similar metal percentage was extracted in un-spiked and in metalspiked soils, suggesting that metal redistribution within the contact time of this study (12 months) could reduce their concentration in exchangeable positions (McLaughlin

Soil	Level	DTPA (pH 7.3)	EDTA (pH 4.6)	HNO ₃ (pH 0.5)	HAc (pH 2.9)	LMWOA (pH 2.9)	NH ₄ Ac (pH 7.0)	CaCl ₂ (pH 7.5)	MgCl ₂ (pH 6.7)	NaNO ₃ (pH 5.9)	NH ₄ NO ₃ (pH 4.7)
M1	Un-spiked	6.9	6.7	5.0	5.2	7.1	7.4	7.9	7.6	8.1	7.4
	Tt1 level	7.1	6.5	4.3	5.2	6.8	7.3	7.8	7.5	7.9	7.3
	Tt2 level	7.2	6.7	3.9	5.1	6.5	7.3	7.3	7.3	7.5	7.1
L1	Un-spiked	6.9	6.4	1.8	4.7	7.1	7.2	7.7	7.5	7.8	7.2
	Tt1 level	6.9	6.4	1.8	4.7	6.7	7.3	7.5	7.2	7.7	7.1
	Tt2 level	6.9	6.4	1.7	4.6	6.2	7.2	7.2	7.0	7.2	6.9
VL1	Un-spiked	6.8	4.4	0.5	3.8	7.0	7.0	7.7	7.2	7.7	7.1
	Tt1 level	6.9	4.4	0.6	3.8	6.9	7.1	7.7	7.1	7.8	7.2
	Tt2 level	6.9	4.3	0.6	3.8	6.6	7.1	7.4	7.0	7.5	7.1

 Table 4
 Values of pH of the extraction solutions and the individual extracts from un-spiked and metal-spiked (Tt1 and Tt2 levels) M1, L1 and VL1 soil samples

2001). Metals in MgCl₂ extracts were mainly in the form of chlorocomplexes. In the case of NH₄NO₃ extracts, amine complexes were the dominant species of Cd, Cu and Zn, while Pb formed mainly nitric complexes. It should be noted that the calculated ionic strengths for MgCl₂ and NH₄NO₃ extract exceeded 1 mol L^{-1} , so the results of this calculation must be considered with caution. The low metal extraction percentage with neutral salts indicates that sorption to exchange positions is not the main metal retention mechanism in these soils but the specific adsorption. Moreover, it is worth noting that competition mechanisms can be established between the alkaline-earth cations, major cations in these soils and the cations of the extractants. In contrast, high amount of metals were extracted with MgCl₂ which was attributed to its higher salt concentration, the combined effect of complexation by chloride and metal displacement from the exchangeable complex by Mg²⁺ (Meers et al. 2007), and to slight carbonate dissolution (Gleyzes et al. 2002).

Metal extractability patterns

Metal extractability patterns in metal-spiked soils were studied according to the soil physicochemical characteristics (Tables 5, 6, 7, 8), and by mineralogical investigations. In this regard, it was observed in general that soil properties and components affecting metal extractability patterns depend on the metal being studied, as discussed below. The extractions with EDTA are an exception, showing a general extractability pattern for all metals. In all cases, significant negative correlations were observed between EDTA-extractable metals and OM content, mainly with the recalcitrant fraction, showing that re-adsorption processes could occur in organic material.

Cadmium extractability patterns

The key role of the soil carbonate fraction in Cd extractability was evident from the correlation coefficients calculated. Cadmium extractability patterns, estimated with MgCl₂, NH₄NO₃, DTPA and LMWOA methods, showed significant and negative correlations with all the soil constituents measured in relation to the carbonate fraction (ECC, AL, total Ca-Mg contents and the relative proportion of calcite and dolomite) (Table 5). This indicates that the soil carbonate fraction govern both mobile and potentially mobile fractions of Cd in these soils. Due to the contact time of the incubation experiment (12 months), it would produce a slow diffusion of soluble Cd species into the crystal defects and pores of the lime, resulting in stable bonds (Buekers et al. 2007), and hence decreasing Cd extractability. The positive and significant correlations obtained between acidic extractants and ECC and AL contents could corroborate this hypothesis, since these extractants total or partially dissolve the carbonate, thereby releasing Cd into the solution (Vidal et al. 2004). This pattern has been observed by other authors when studying Cd desorption by fibrous minerals, i.e. palygorskite and sepiolite group (Shirvani et al. 2007), and Mn oxides (Zaman et al. 2009). Despite the mineralogical study showing no evidence of formation of Cd-carbonate or oxide, probably due to the fact that the Cd concentrations used were below the quantification limit of the XRD equipment (<2 %), surface precipitation phenomena can occur.

On the other hand, negative and significant correlations at Tt2 level were obtained between DTPA-extractable Cd and the crystalline Fe oxide and OM contents, and between EDTA-extractable Cd and RP content. The formation of stable organomineral associations, favoured by the high affinity of Cd for high molecular weight organic acids, can explain these patterns (Prokop et al. 2003). In view of these results, it could be concluded that potentially mobile Cd, estimated with EDTA and DTPA, is affected by crystalline Fe oxides and OM. Nevertheless, readsorption processes of negatively charged Me-DTPA and EDTA complexes in the organic matter should be taken into account (Ettler et al. 2007).

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Level	Extraction method	ECC	AL	Tot-Ca	Tot-Mg	MO	RP	Cry-Fe	Clay	CEC	Calcite <2 mm	Dolomite
Tt1	DTPA	-0.609	-0.811**	-0.739*	-0.608	0.565	-0.050	0.542	0.410	0.469	-0.539	-0.479
	HNO ₃	0.699*	0.605	0.491	0.573	0.188	-0.035	0.345	-0.091	-0.287	0.365	0.374
	LMWOA	-0.549	-0.663*	-0.459	-0.783^{**}	0.092	0.056	-0.242	-0.122	0.072	-0.663*	-0.492
	$MgCl_2$	-0.714*	-0.773 **	-0.721*	-0.737*	0.167	-0.092	0.175	0.490	0.586	-0.533	-0.610
	$NaNO_3$	-0.471	-0.422	-0.432	-0.636*	-0.134	0.089	-0.515	-0.317	-0.048	-0.149	-0.400
Tt2	DTPA	0.159	0.311	0.521	-0.069	-0.715*	0.151	-0.946^{***}	-0.383	-0.364	-0.167	0.046
	EDTA	-0.523	-0.313	-0.455	-0.117	-0.205	-0.649*	0.027	0.492	0.613	-0.477	-0.452
	HNO ₃	-0.083	-0.191	-0.114	0.328	0.793^{**}	0.472	0.461	-0.161	-0.099	0.118	-0.216
	HAc	0.669*	0.641^{*}	0.528	0.291	0.101	0.119	-0.195	-0.738*	-0.775**	0.421	0.426
	LMWOA	-0.646*	-0.514	-0.594	-0.654*	-0.090	-0.279	-0.300	-0.170	0.140	-0.384	-0.662*
	$\rm NH_4Ac$	-0.365	-0.523	-0.486	-0.595	0.312	-0.231	0.301	0.037	0.031	-0.458	-0.754*
	$CaCl_2$	-0.519	-0.512	-0.483	-0.710*	-0.156	-0.326	-0.398	-0.325	-0.051	-0.489	-0.385
	$MgCl_2$	-0.672*	-0.678*	-0.814^{**}	-0.571	0.367	-0.147	0.617	0.637*	0.719*	-0.607	-0.610
	NaNO ₃	-0.563	-0.543	-0.531	-0.718*	-0.157	-0.359	-0.381	-0.277	0.018	-0.632	-0.492
	NH_4NO_3	-0.680*	-0.653*	-0.731*	-0.810^{**}	-0.085	-0.508	-0.157	-0.057	0.220	-0.650*	-0.649*
ECC eq. $*_{p < 0.05}$	uivalent CaCO ₃ , <i>AL</i> activ ;; ** <i>p</i> <0.01; *** <i>p</i> <0.00	ve lime, $Tot-Cc$ 1 $(n=10)$, stati	a and Tot-Mg to stical significar	otal Ca and Mg, nce at these prob	<i>OM</i> organic m ability levels	latter, RP recal	citrant pool, C	<i>ry-Fe</i> crystalline	: Fe-oxide, <i>CE</i>	C cation excha	nge capacity	

lable 0	Significance correlation	analyses calcu	ulated between	soil physicoch	emical parame	eters (Pearson	s correlation test) and extractable	e Cu m metal-sj	piked soil samj	oles at 111 and	1t2 levels
Level	Extraction method	ECC	AL	Tot-Ca	Tot-Mg	MO	RP	Cry-Fe	CS	FS	Clay	CEC
Tt1	DTPA	-0.033	-0.133	-0.053	-0.411	-0.091	0.242	-0.533	0.906***	0.404	-0.652*	-0.492
	HAc	0.355	0.477	0.453	0.033	-0.746*	0.086	-0.787^{**}	0.525	0.326	-0.442	-0.421
	$\rm NH_4Ac$	0.751^{*}	0.819^{**}	0.882^{**}	0.583	-0.590	0.108	-0.638*	0.144	0.555	-0.504	-0.633*
	$MgCl_2$	0.648*	0.452	0.489	0.321	0.178	-0.220	-0.145	0.405	0.722*	-0.753*	-0.870 **
	NaNO ₃	0.374	0.091	0.224	0.017	0.490	0.045	0.006	0.465	0.693*	-0.704*	-0.751*
	NH_4NO_3	0.802^{**}	0.836^{**}	0.826^{**}	0.444	-0.533	0.117	-0.568	0.344	0.586	-0.582	-0.727*
Tt2	DTPA	-0.527	-0.486	-0.429	-0.662*	-0.233	-0.327	-0.520	0.700*	0.059	-0.349	-0.042
	EDTA	-0.629	-0.391	-0.638*	-0.478	-0.416	-0.924^{***}	-0.043	0.085	-0.711*	0.536	0.656^{*}
	HAc	0.162	0.302	0.174	-0.258	-0.427	-0.304	-0.634^{*}	0.621	0.400	-0.569	-0.454
	$CaCl_2$	-0.438	-0.503	-0.418	-0.648*	-0.022	-0.225	-0.371	0.784^{**}	0.199	-0.473	-0.202
	$MgCl_2$	-0.331	-0.464	-0.400	-0.614	0.115	-0.174	-0.247	0.764*	0.309	-0.543	-0.286
	$NaNO_3$	-0.444	-0.466	-0.404	-0.642*	-0.091	-0.234	-0.423	0.767**	0.194	-0.455	-0.179
<i>ECC</i> equestion exchange	nivalent CaCO ₃ , <i>AL</i> acti capacity	ve lime, Tot-C	a and Tot-Mg	total Ca and M	lg, <i>OM</i> organi	c matter, RP	recalcitrant pool,	Cry-Fe crystall	ine Fe-oxide, C	S coarse sand,	FS fine sand,	CEC cation

*p<0.05; **p<0.01; ***p<0.001 (n=10), statistical significance at these probability levels

At Tt2 level, significant and positive correlations were obtained between $MgCl_2$ -extractable Cd and clay content and CEC, showing that exchangeable processes also play a relevant role at this contamination level. The significant and negative correlations between HAc-extractable Cd and clay content and CEC may be an artefact of carbonate, as the soils with a lower proportion of clay generally present higher carbonate contents in this study.

Copper extractability patterns

There were significant and positive correlations between NH_4Ac and NH_4NO_3 -extractable Cu and ECC, AL and total Ca contents at Tt1 level (Table 6), showing that carbonate fraction affects Cu extractability patterns with these extraction methods. Since the mineralogical study showed no evidence of crystallized Cu species, sorption processes, others that precipitation ones, are affecting these patterns, probably adsorption ones. In this scenario, ammonium ions could displace the Cu adsorbed from carbonates. The highest correlation coefficients (p < 0.01) obtained for AL indicated that the finest carbonate fraction could play a greater role in affecting Cu extractability.

On the other hand, significant and negative correlations were obtained between DTPA, MgCl₂ and NaNO₃extractable Cu and the clay content and CEC (Tt1 level) and between HAc-Cu and the crystalline Fe oxide content (whatever the level). This indicated that even though AL should affect Cu extractability patterns, other soil constituents present in the finest soil fraction <2 µm, such as Fe oxides and phyllosilicates, were implied. The enhanced adsorption capacity of illite minerals in the presence of Fe (Sipos et al. 2008), together with the high ability of Cu^{2+} to displace the Fe^{2+} , can explain these patterns. Nevertheless, the significant and positive correlations observed between EDTA-extractable Cu and CEC at Tt2 level show that other sorption mechanisms, such as exchangeable processes, also can take place (Sayen and Guillon 2009).

Similarly, significant and negative correlations were also obtained between HAc-Cu and OM content at Tt1 level, and between EDTA-Cu and RP at Tt2 level. The higher Cu electronegativity, the predominance of positively charged Cu hydroxocomplexes (characteristics of the pH range studied) and the degree of humification of the soil samples (C/N~11) favour the formation of stable associations between Cu and humic substance. These results indicate that the potential mobility of Cu may be governed by both inorganic and organic fractions through organomineral associations (Besnard et al. 2001), probably humate–Fe oxide associations (Sayen and Guillon 2009; Sipos et al. 2008) at both levels of contamination.

Level	Extraction method	ECC	AL	Tot-Ca	Tot-Mg	RP	Clay	CEC	Dolomite <2 mm
Tt1	DTPA	-0.330	-0.232	-0.327	-0.680*	0.030	0.122	0.211	0.150
	HNO ₃	0.683*	0.590	0.444	0.457	0.041	-0.369	-0.454	0.369
	HAc	0.718*	0.742*	0.492	0.354	0.130	-0.366	-0.445	0.876**
	LMWOA	-0.330	-0.332	-0.371	-0.650*	0.081	-0.393	-0.161	-0.060
	MgCl ₂	-0.430	-0.400	-0.436	-0.703*	0.201	-0.310	-0.052	-0.099
	NaNO ₃	-0.638*	-0.494	-0.595	-0.642*	-0.102	0.047	0.313	-0.177
	NH ₄ NO ₃	-0.467	-0.245	-0.433	-0.198	0.145	0.672*	0.740*	-0.102
Tt2	DTPA	-0.555	-0.345	-0.369	-0.571	-0.272	0.008	0.233	-0.043
	EDTA	-0.475	-0.204	-0.574	-0.463	-0.914***	0.460	0.573	-0.077
	HAc	0.537	0.579	0.227	0.144	-0.402	-0.472	-0.477	0.640*
	LMWOA	-0.425	-0.337	-0.423	-0.651*	-0.334	-0.328	-0.069	-0.088
	CaCl ₂	-0.462	-0.392	-0.469	-0.710*	-0.606	-0.116	0.057	-0.352
	MgCl ₂	-0.620	-0.596	-0.657*	-0.778**	-0.402	-0.167	0.126	-0.336
	NaNO ₃	-0.525	-0.434	-0.445	-0.705*	-0.343	-0.302	-0.020	-0.193
	NH ₄ NO ₃	-0.526	-0.482	-0.476	-0.695*	-0.322	-0.332	-0.045	-0.237

 Table 7
 Significance correlation analyses calculated between soil physicochemical parameters (Pearson's correlation test) and mineralogical composition (Spearman's correlation test) and extractable Pb in metal-spiked soil samples at Tt1 and Tt2 levels

ECC equivalent CaCO₃, AL active lime, Tot-Ca and Tot-Mg total Ca and Mg, RP recalcitrant pool, CEC cation exchange capacity

p < 0.05; p < 0.01; p < 0.001 (n=10), statistical significance at these probability levels

These results therefore highlight that the combined action of the finest organic and inorganic fractions govern the Cu extractability patterns in these calcareous agricultural soils.

Lead extractability patterns

Few correlations were obtained for extractable Pb (Table 7). At level Tt1, significant and positive correlations were obtained between acidic extractants and the ECC and AL contents, suggesting the total or partial dissolution of the carbonate fraction (Vidal et al. 2004). Nevertheless, significant correlations were not obtained at Tt2 level, despite the formation of lead oxide carbonate observed (Fig. 2). The lack of correlations with soil physicochemical parameters at this level led us to study the relationship with soil mineralogical composition. The significant and positive correlations obtained between the HAc-extractable Pb and the relative proportion of dolomite (<2 mm) at both levels suggest that the Pb extractability pattern with acidic extraction methods is mainly affected by the dolomite phase. However, extractable Pb with other extraction methods with lower strength (DTPA, LMWOA, MgCl₂, NH₄NO₃ and NaNO₃) showed significant and negative correlations with the total Mg content. Rangel-Porras et al. (2010), studying the mechanism for retention of Pb ions in presence of calcite, observed the nucleation of micro-precipitates of Pb on the calcite surfaces. In our case, the pH range of our soils, combined with the incubation time of our experiment (12 months), favours these precipitation processes on the surfaces of Mg-rich clay minerals, such as dolomite, explaining the negative correlations obtained. These results therefore suggest that the potential mobility of Pb depends to some degree on the turnover of inorganic C. Nevertheless, the intensities of reflections obtained from XRD were not high, showing the low crystallinity degree. Hence, precipitation mechanisms in carbonate fraction cannot be regarded as the only process affecting Pb retention in these soils. As previously reported, adsorption and complexing processes could play a key role in affecting Pb retention (Bradl 2004).

The predominant species in the pH range in the study, $Pb(OH)^+$, form stable bonds with OM and clay, probably through organomineral associations regulating the retention of Pb. However, the slow Pb desorption from organic matter, in contrast to its rapid initial sorption (Strawn and Sparks 2000), explain the scarce correlations obtained with clay and OM. This sorption mechanism may not affect Pb extractability patterns or its potential mobility as high sorption does not necessarily lead to high bond strength at the same time. On the other hand, NH_4NO_3 -extractable Pb correlated positively and significantly with clay content and CEC at Tt1 level, suggesting that a fraction of Pb could be in exchange positions at this level.

Zinc extractability patterns

The Pearson's correlation coefficients calculated showed that particle-size distribution was the main soil characteristic

Table 8 extractab	Significance correlatio	n analyses cal oil samples at 7	culated betwee ft1 and Tt2 lev	en soil physico /els	ochemical par	ameters (Pearsc	n's correlation	test) and mine	ralogical compo	sition (Spearma	ın's correlatic	
Level	Extraction method	Tot-Mg	MO	CS	FS	Silt	Clay	CEC	Mica-illite <2 µm	Kaolinite	Quartz	Dolomite
Tt1	DTPA	-0.530	0.064	0.836**	0.246	-0.710*	-0.506	-0.298	-0.039	-0.406	0.039	-0.220
	EDTA	-0.388	-0.346	-0.197	-0.352	0.258	0.353	0.529	0.658*	-0.058	0.526	-0.468
	HAc	0.193	0.212	0.425	0.802^{**}	-0.536	-0.815^{**}	-0.821^{**}	-0.540	-0.058	-0.415	0.220
	LMWOA	-0.582	-0.088	0.856^{**}	0.323	-0.753*	-0.573	-0.343	-0.230	-0.522	-0.197	-0.165
	$\rm NH_4Ac$	-0.030	0.184	0.644*	0.727*	-0.827^{**}	-0.753*	-0.790**	-0.434	-0.290	-0.421	0.275
	$CaCl_2$	-0.618	0.152	0.803^{**}	0.324	-0.753*	-0.532	-0.309	-0.211	-0.406	-0.151	-0.055
	$MgCl_2$	-0.683*	0.164	0.765**	0.161	-0.683*	-0.383	-0.147	-0.020	-0.290	0.086	-0.138
	NaNO ₃	-0.680*	-0.050	0.813^{**}	0.162	-0.674*	-0.427	-0.166	-0.039	-0.406	-0.026	-0.110
	NH_4NO_3	-0.517	-0.032	0.858^{**}	0.409	-0.848^{**}	-0.605	-0.426	-0.368	-0.522	-0.421	0.083
Tt2	DTPA	-0.662^{*}	-0.130	0.806^{**}	0.159	-0.663*	-0.426	-0.164	-0.007	-0.406	0.007	-0.165
	EDTA	-0.108	-0.634*	0.233	-0.490	0.034	0.270	0.253	0.171	-0.522	0.224	-0.881^{**}
	HAc	-0.234	0.019	0.552	0.523	-0.517	-0.657*	-0.587	0.678*	0.522	0.737*	-0.275
	LMWOA	-0.615	0.097	0.744*	0.166	-0.680*	-0.373	-0.194	0.020	-0.058	-0.053	-0.028
	$\rm NH_4Ac$	-0.184	0.090	0.683*	0.539	-0.780^{**}	-0.629	-0.602	0.164	-0.406	0.164	-0.275
	$CaCl_2$	-0.667*	-0.113	0.776^{**}	0.160	-0.630	-0.420	-0.148	-0.388	-0.290	-0.434	0.220
	$MgCl_2$	-0.674*	0.014	0.718*	0.106	-0.555	-0.365	-0.081	-0.026	-0.290	0.026	-0.193
	$NaNO_3$	-0.659*	-0.199	0.731^{*}	0.116	-0.570	-0.376	-0.087	0.184	-0.058	0.243	-0.083
	NH_4NO_3	-0.657*	-0.179	0.737*	0.142	-0.592	-0.393	-0.109	0.250	-0.290	0.243	-0.275
Tot-Mg t	otal Mg, <i>OM</i> organic mi	atter, CS coarse	e sand, FS fine	sand, CEC ca	tion exchange	capacity						

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p<0.05; p<0.01; p<0.01; p<0.001 (n=10), statistical significance at these probability levels

affecting Zn extractability patterns, and no significant correlations were obtained with other soil physicochemical parameters, except the general case of EDTA and OM (Table 8). Thus, significant and negative correlations were obtained between extractable Zn with almost all extractants and clay and silt contents, suggesting that the soil fractions with a higher adsorption capacity are primarily responsible for governing both mobile and potentially mobile Zn fractions. From the mineralogical study, we obtained positive correlations between the relative proportion of kaolinite, illite and quartz (<2 µm), and HNO₃ and EDTA-extractable Zn. This indicates that the presence of these minerals, with a low specific surface area, favours a higher Zn extractability. The significant and negative correlations obtained between extractable Zn with several extractants and the total Mg content, mainly at Tt2 level, indicate that lower Zn extractability may be linked to the presence of Mg-rich minerals (Lafuente et al. 2008), probably due to the similarity of ionic radius between Zn²⁺ and Mg²⁺. Similarly, significant and negative correlations were obtained between EDTA-extractable Zn and the dolomite relative proportion ($<2 \mu m$) at the Tt2 level. Therefore, from these results, it can be deduced that the Zn extractability patterns are affected by Mg-rich minerals, probably dolomite, and it is expected that Mg ions were released into the soil solution by displacement of Zn ions. In previous works (data not shown), we observed an increase in the content of Mg in soil solution in the metal-spiked soils at Tt2 level. These results therefore highlight the fact that the clay content alone cannot explain the Zn-extractability patterns, since soil mineral composition plays a significant role in controlling its extractability.

Conclusions

Despite carbonate fraction typically conferring a high metal sorption capacity, this fraction alone cannot explain the metal extractability patterns in the metal-spiked soils studied, under the experimental conditions used. Instead, the content and composition of some soil fractions (mainly organic matter, Fe oxides, and clay) have been shown to govern metal extractability patterns, highlighting the need to consider the intrinsic soil characteristics of each polluted site. In particular, the finest organic and inorganic soil fractions played the main role in affecting the both mobile and potentially mobile fractions of metals. In this respect, the combination of indirect and direct methods employed in this study has proved to be an essential tool for investigating soil constituents affecting (potentially) mobile fractions of Cd, Cu, Pb and Zn in these calcareous agricultural soils. The high metal extraction percentage, mainly with chelating agents, highlights the potential vulnerability of these soils against an environmental impact from metal accumulation in periurban agricultural areas. Results are also useful for remediation purposes since the application of chelating agents to agricultural soils becomes to be a generalized practice.

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