

CO₂ Injection into Saline Carbonate Aquifer Formations II: Comparison of Numerical Simulations to Experiments

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Abstract Sequestration of carbon dioxide in geological formations is an alternative way of managing extra carbon. Although there are a number of mathematical modeling studies related to this subject, experimental studies are limited and most studies focus on injection into sandstone reservoirs as opposed to carbonate ones. This study describes a fully coupled geochemical compositional equation-of-state compositional simulator (STARS) for the simulation of CO₂ storage in saline aquifers. STARS models physical phenomena including (1) thermodynamics of sub- and supercritical CO₂, and PVT properties of mixtures of CO₂ with other fluids, including (saline) water; (2) fluid mechanics of single and multiphase flow when CO₂ is injected into aquifers; (3) coupled hydrochemical effects due to interactions between CO₂, reservoir fluids, and primary mineral assemblages; and (4) coupled hydromechanical effects, such as porosity and permeability change due to the aforementioned blocking of pores by carbonate particles and increased fluid pressures from CO₂ injection. Matching computerized tomography monitored laboratory experiments showed the uses of the simulation model. In the simulations dissolution and deposition of calcite as well as adsorption of CO₂ that showed the migration of CO₂ and the dissociation of CO₂ into HCO₃ and its subsequent conversion into carbonate minerals were considered. It was observed that solubility and hydrodynamic storage of CO₂ is larger compared to mineral trapping.

Keywords CO₂ injection · Aquifer · Calcite deposition · Permeability and porosity alteration · Numerical model

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Nomenclature

k_f	Instantaneous absolute permeability
k_0	Initial absolute permeability
c	Kozeny–Carman-power
s_{ki}	Reactant stoichiometric coefficient of reaction k
s'_{ki}	Product stoichiometric coefficient of reaction k
s_w	Irreducible water saturation
p_d	Entry pressure, MPa
rk	Rate of reaction k
rrk	Constant part of rk
C_i	Concentration of component i in void volume (ppm)
E_{ak}	Temperature dependence of rk
H_{rk}	Reaction enthalpy
R	Gas constant ($\text{Pa} \cdot \text{m}^3/\text{mol} \cdot \text{K}$)
T	Temperature (K)
λ	Pore-size distribution index of Brooks–Corey model
ϕ	Instantaneous effective porosity
ϕ_0	Initial effective porosity
φ_f	Fraction of void porosity occupied by fluids
φ_s	Fraction of void porosity occupied by solid particles

1 Introduction

Injection of CO_2 into aquifers results in variety of strongly coupled physical and chemical process as multiphase flow, dissolution–deposition kinetics, solute transport, hydrodynamic instabilities due to displacement of less viscous CO_2 with more viscous brine (viscous fingering), capillary effects and upward movement of CO_2 due to gravity (gravity override) (Reichle et al. 1999). Reactions among the formation rock, the aquifer fluid and CO_2 may lead to change in the formation permeability and the effective porosity, thus the storage capacity of the formation.

Through the injection of CO_2 in carbonated deep saline aquifers, some changes in rock properties are expected. Change in the rock permeability and effective porosity result from dissolution of rock minerals, transportation and later precipitation of them. Continuous dissolution of reactant minerals alters the concentration of aquifer fluid, thus in later times leading to precipitation of product phases. While dissolution of rock minerals increases the formation permeability, precipitation of those minerals leads to decrease in the formation permeability and the effective porosity (Lichtner 1996; Xu et al. 2004; Izgec et al. 2007).

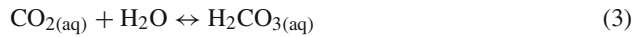
At the CO_2 front where CO_2 is dissolved in water, minerals such as calcite may dissolve readily, leading to an increase in permeability and porosity along the flow channel. This leads to a higher flow rate and increased dissolution, forming what is known as wormholes. It is also known that from the various field applications of enhanced oil recovery that CO_2 has been known to reduce injectivity in some cases but to increase permeability near injection wells in others, such as in carbonate reservoirs.

For a carbonate system kinetically controlled reactions could be defined as (Omole and Osaba 1983; Snoeyink and Jenkins 1980):

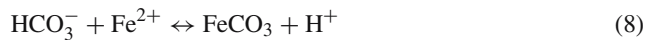
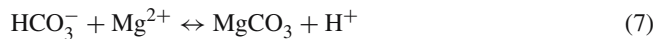
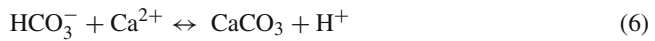


In general, atmospheric or subsurface CO_2 dissolves in water and generates a weak carbonic acid, which subsequently dissociates into HCO_3^- and CO_3^{2-} according to reaction steps given

as:



The dissolved bicarbonate species react with divalent cations to precipitate carbonate minerals. Formation of Ca, Mg, and Fe(II) carbonates are expected to be the primary means by which CO₂ is immobilized (Gunter et al. 1997).



Mechanisms by which a precipitate reduces permeability include deposition of the carbonate particles precipitated from the saturated solution on the pore walls due to attractive forces between the particles and the surfaces of the pores, individual particles blocking pore throats, and several particles bridging across a pore throat (Pruess and Xu 2001). In a carbonate formation major cause of reduction in rock properties is precipitation of Ca(HCO₃)₂ and NaCl. Pressure drop through the flow paths affects the precipitation rate, thus leads to variations in rock properties by changing the solubility of substances. Assuming there is Darcian flow in the porous media, it can be said that there is a linear relationship between the pressure drop and the axial distance in the direction of flow. Considering this relationship and solute transport concept, it should be expected that permeability increases in near well bore region and then gradually decreases through the flow direction (Omole and Osaba 1983). In the case of multiphase flow and/or heterogeneous and fractured formations, dramatic deviations from the linear pressure decrease are reported (Matthai and Belayneh 2004). This, in turn may affect the permeability alteration trends. It is previously reported that permeability decline caused by only calcite deposition in the porous bed can reach to 90% of the initial permeability, depending on solution composition, initial permeability, temperature, and flow rate and solution injection period (Moghadasi et al. 2005). On the other hand, some researchers (Omole and Osaba 1983) reported increase in the permeability of dolomite cores by 3.5–5% after similar CO₂ treatments while reduction in permeability was observed in other experiments. Those results suggest that the process strongly depends on the distribution of the rock minerals.

As discussed-above, injection of CO₂ in geologic formations may give rise to a number of physical and chemical phenomena, such as miscible or immiscible displacement of native fluids, dissolution of injected fluids into reservoir fluids, changes in effective stress with associated porosity and permeability change and the possibility of inducing seismic activity, chemical interactions between fluids and solids, and nonisothermal effects. Johnson et al. (2001) simulated the large scale injection of CO₂ currently being produced at Statoil's Sleipner facility in the Norwegian sector of the North Sea (Korbol and Kaddour 1995). They analyzed the coupled processes and mechanisms that lead to structural, solubility, and mineral trapping, and quantified the relative effectiveness of the distinct sequestration processes

as a function of key reservoir properties. [White et al. \(2001\)](#) applied reactive chemical transport modeling to simulate mineral sequestration of CO₂ in saline reservoirs underlying the Colorado Plateau. [McPherson and Lichtner \(2001\)](#) used a mathematical sedimentary basin model, including multiphase flow of CO₂, groundwater, and brine, to evaluate resident times in possible aquifer storage sites and migration patterns and rates away from such sites in the Powder River Basin of Wyoming. Numerical simulation of CO₂ disposal in aquifers has also been studied by [Doughty and Pruess \(2004\)](#), [Xu et al. \(2004\)](#), [Allen et al. \(2005\)](#), [Kumar et al. \(2005\)](#), and [Lagneau et al. \(2005\)](#). In these studies, local equilibrium is assumed to govern the distribution of aqueous chemical species. A critical comparison of simulation codes that could be used to model injection of CO₂ in geologic formations was presented by [Pruess et al. \(2002\)](#). They concluded that a considerable number of numerical simulation codes (i.e., CMG's GEM, Geoquest's Eclipse, TOUGH2, etc.) were capable of simulating, in realistic, quantitative detail, the important flow and transport processes that would accompany geologic sequestration; however, the hydromechanical test problem was solved by only one code (TOUGH-FLAC).

In this study, we propose an alternative numerical model developed using CMG's STARS that considers physical phenomena including (1) thermodynamics of sub- and supercritical CO₂, and PVT properties of mixtures of CO₂ with other fluids, including (saline) water; (2) fluid mechanics of single and multiphase flow when CO₂ is injected into aquifers; (3) coupled hydrochemical effects due to interactions between CO₂, reservoir fluids, and primary mineral assemblages; and (4) coupled hydromechanical effects, such as porosity and permeability change due to the aforementioned blocking of pores by carbonate particles and increased fluid pressures from CO₂ injection. Unlike the aforementioned numerical simulation studies permeability change as a result of dissolution and deposition of the carbonate particles precipitated from the saturated solution on the pore walls, individual particles blocking pore throats, and several particles bridging across a pore throat and adsorption of CO₂ is considered in the developed numerical model. The reaction model's heterogeneous mass transfer (source and sink) terms were applied to the nonequilibrium capture and release of calcite particles as well as salt particles by the porous rock. First details of the model will be given. Then calibration of the model using CT monitored experiments will be introduced. Sensitivity analyses of deposition and dissolution reactions, as well as other important simulation parameters will be discussed.

2 Numerical Modeling Approach

2.1 Main Features

CMG's STARS ([CMG 2003](#)) compositional advanced processes and thermal finite difference reservoir simulator is capable of simulating many types of chemical additive processes such as surfactant and polymer injection, using cartesian or cylindrical grid and porosity models (single porosity or double porosity) in both laboratory and field scale. STARS is previously used to simulate several processes including CO₂ injection into coal bed methane reservoirs ([Law et al. 2001](#)), modeling phase behavior including the effect of pressure and temperature on asphaltene precipitation during primary production ([Kohse et al. 2000](#)) and fully implicit thermal simulation for multicomponent fluid flows ([Siu et al. 1990](#)). [Law et al. \(2001\)](#) stated that the predicted CH₄ production rates and compositions in the produced gas streams predictions by STARS and ECLIPSE were in general agreement for pure CO₂ injection into a

coal bed methane reservoir. This indicates that STARS has a very similar performance even though it uses a very different modeling approach compared to ECLIPSE.

One conservation equation for each chemical component for which a separate accounting is desired, and two separate phases (water and gas) along with equations describing phase equilibrium between phases are used to model injection of CO₂. There exists a set of these equations for each region of interest, which is usually a discretized grid block. Lastly, there is an equation describing the operating condition of each injection and production well. An implicit time weighing scheme is used for the individual components of the model consisting of flow, transport, and geochemical reaction. A fully implicit approach, which simultaneously solves the transport and the reaction equations was used. No flow boundary conditions are used. CO₂ was defined as real gas and its solubility in water is taken to be proportional to CO₂ partial pressures at pressures of a few bars, but increases only very weakly with pressure beyond 100 bars (Spycher et al. 2003). A Henry's law formulation with fugacity correction was used. Supercritical dilute aqueous CO₂ thermodynamical properties adapted from Sengers et al. (1992) are input in a tabular fashion. Water viscosity and density is taken to be function of temperature and salinity only. The dissolution and deposition reactions given by Eq. 1 through 5 are implemented and treated separately. The reaction model's heterogeneous mass transfer (source and sink) terms were applied to the nonequilibrium capture and release of calcite particles as well as salt particles by the porous rock. These particles captured by the porous medium can cause permeability reductions (blockage) in a manner similar to equilibrium mass transfer to the rock (adsorption). The reaction source/sink term is given as:

$$V \sum_{k=1}^{nr} (s'_{ki} - s_{ki}) \bullet rk \quad (9)$$

where s_{ki} and s'_{ki} represent the reactant and product stoichiometric coefficients of reaction k . Equation 9 proceeds at the rate of rk moles per day per reservoir volume. This relationship has one degree of freedom, which is a proportionality factor. The quantities s_{ki} and s'_{ki} can be multiplied by an arbitrary factor "a", but rk must be divided by "a" so that the source/sink terms remain. Usually the factor "a" is chosen such that $s_{ki} = 1$ for the main reacting component. The kinetic model determines the speed of reaction rk . The general expression is given by:

$$rk = r_{rk} \cdot \exp\left(\frac{-E_{ak}}{RT}\right) \cdot \prod_{i=1}^{nc} C_i \quad (10)$$

where r_{rk} is reaction rate constant that specifies the frequency of the reaction, E_{ak} is the activation energy; E_{ak} that determines the temperature dependence of reaction rate, rk . While the enthalpies of reaction can be characterized between well-defined limits and can even be calculated from first principles, the observed activation energies can vary dramatically. This is because certain components in the rock surface may act as catalysts. The concentration factor for reacting component i is given as follows:

$$C_i = \varphi_f \cdot \rho_j \cdot s_j \cdot x_{ji} \quad (11)$$

where j is the phase in which component i is reacting, and x_{ji} represents water or gas mole fractions, φ_f is fluid porosity, ρ_j is density, and s_j is saturation. Concentration factor for the solid component is given by

$$C_i = \varphi_v c_i \quad (12)$$

where φ_v is void porosity (ratio of void volume to gross volume) and c_i is the concentration of component i in void volume. The void porosity can be occupied by fluids (represented by fluid porosity, φ_f) and the deposited calcite particles (represented by solid porosity, φ_s). Fluid porosity and void porosity is related with the following relationship.

$$\varphi_f = \varphi_v = \left[1 - F_{\text{fluid}} - \sum_i \left(\frac{c_i}{\rho_{si}} \right) \right] \quad (13)$$

where F_{fluid} is the fraction of void volume occupied by fluid components. It then requires that the permeability depends upon reaction rate constants, to account for the changes in capture efficiency as the calcite particle size to pore throat size ratio changes. To specify the dependence of permeability on chemical reactions and nonequilibrium mass transfer an effective permeability reaction rate scaling factor table (i.e., Table 2) was used. Once porosity as a function of time is calculated permeability is adjusted accordingly. Thus, permeability change was controlled by reaction frequencies (1/min-kPa) of the dissolution and deposition reactions and Kozeny–Carman coefficient (c) given by Eq. 14 (CMG 2003).

$$k_f = k_0 \left(\frac{\phi}{\phi_0} \right)^c \left(\frac{1 - \phi_0}{1 - \phi} \right)^2 \quad (14)$$

The rate of propagation of in situ created calcite particles are strongly affected by their interaction with the rock matrix. These interactions can be chemical (e.g., ion exchange) or mechanical (e.g., blockage) or some combination of mechanisms. The capture levels can depend on fluid concentrations, temperature and rock type (e.g., permeability). A phenomenological description of these phenomena, wherein a set of constant temperature adsorption isotherms (adsorption level as a function of fluid composition) are input, is adapted. These isotherms can be either in tabular form or in terms of the well-known Langmuir isotherm correlation.

$$AD = \frac{A \cdot z}{1 + Bz} \quad (15)$$

Where z is some fluid component composition, A and B are generally temperature dependent Langmuir constants. The maximum adsorption level associated with this formula is A/B . Permeability alteration often accompanies adsorption (especially if adsorption is of mechanical, blockage type). The simulator accounts for this via region dependent resistance factors (RRF) which allow correlation of local permeability with local adsorption levels. It is assumed that only single-phase flow paths are altered. Thus for example, the gas phase permeability reduction factor is defined as

$$RKG = 1 + (RRF - 1) * AD / ADMAXT \quad (16)$$

RKG varies between 1.0 and a maximum of RRF as adsorption level increases. The mobility of the gas phase is divided by RKG , thus accounting for blockage. Ideally, not only the adsorption maximum but the rate of increase of adsorption with fluid composition should be known in order to fit the two Langmuir parameters A and B . If this is not reported, as is often the case, one must use the fluid composition at the adsorption maximum to indirectly determine this second factor.

From the aforementioned discussion, critical simulation parameters can be summarized as follows:

1. Dissolution and deposition reaction frequencies;
2. Kozeny–Carman coefficient;

3. Reaction rate scaling factor;
4. Blockage effect of particles;
5. Adsorption rate of CO₂; and
6. Initial concentration of solid phase carbonate and aqueous phase bicarbonate, aqueous phase sodium bromide particles.

For a typical problem cumulative mass balance errors at the ending time were less than 1% for all components. In order to achieve this level of accuracy and to prevent large material balance errors due to insufficient accuracy of iterative matrix solution, a precision parameter that is the ratio by which the mean equation residual must be reduced from its initial value before the solution is accepted or used. When used with an automatic time-step cutting control algorithm this method ensured accurate solutions with 10–15 Newton cycles per time-step. As the size of a problem increases, the relative amount of time spent in the various STARS component routines will change. The time spent in the EOS routine increases roughly, linearly with the size of the problem (number of equations solved). Solution time of the linear equations increases roughly proportional to the problem size raised to an exponent of about 1.4–1.6 (using iterative solvers) or 3 for direct solvers. In typical, large STARS runs, on the order of 60% of the execution time is spent in the linear equation solver and a further 30% is spent setting up the Jacobian matrices.

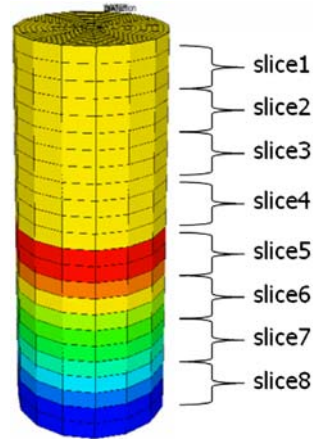
2.2 Comparison with Experiments

Computerized Tomography (CT) experiments reported by [Izgec et al. \(2007\)](#) were used to calibrate the numerical model. In these experiments heterogeneous carbonate cores drilled from a carbonate formation were used. Experiments modeled fast near well bore flows (60 ml/min) as well as slow reservoir flows (3 ml/min) at differing temperature (18, 35, and 50°C) and differing orientations (vertical and horizontal). For vertically aligned experiments epoxy coated core plugs of 10.7 cm long and 4.72 cm in diameter as opposed to 7 cm long and 3.81 cm ones in horizontal experiments were used. In all experiments porosity values along the core plugs were measured at the end of each CO₂ injection period using a computerized tomography scanner. Absolute permeability to NaBr brine at differing concentrations (0, 2.5, 5, and 10%) was calculated using Darcy's law. NaBr brine as opposed to NaCl brine was used in the experiments since NaBr brine provided excellent CT number contrast with that of the CO₂ that enabled tracking of porosity changes as a result of calcite dissolution and deposition. Details of the experiments are reported elsewhere ([Izgec et al. 2007](#)).

Radial grid block system with $14 \times 25 \times 24$ blocks was used to model the laboratory experiments (Fig. 1). Initial porosities obtained from CT scans were designated to each block corresponding to a slice. The missing porosities in between the slices were distributed using an inverse-distance squared distribution function. Initial permeabilities of each grid were assumed constant. Brooks–Corey CO₂–water relative permeabilities and Brooks–Corey capillary pressures ($\lambda = 0.4$, $P_d = 0.11$ MPa, $S_{wirr} = 0.13$) were used. Note that low pore size distribution index, λ indicates greater heterogeneity.

Calibration of the simulation model was conducted by changing the following parameters: reaction frequencies of the deposition and dissolution reactions, Kozeny–Carman coefficient, reaction rate scaling factor, blockage effect of particles, adsorption rate of CO₂ and initial concentrations of solid phase carbonate and aqueous phase bicarbonate, aqueous phase NaBr. It was observed that matches could be obtained by setting dissolution and deposition reaction frequencies to 3500 and 550 and Kozeny–Carman coefficient to 6.5. For horizontal-oriented experiments the initial concentration of the dissolved bicarbonate (1 g mole/cm³) was 10 times higher compared to vertical cases. Figures 2 to 4 give the comparison of experimental

Fig. 1 Distribution of slice porosities in core-scale model



and numerical permeability values for vertical and horizontal oriented heterogeneous and vertical oriented homogenous core plug experiments with varying conditions. For all cases, it was observed that the simulations followed the permeability alteration trend; however, the magnitude of the changes could not be matched exactly. This mismatch could be attributed to the inappropriate representation of the heterogeneous nature of the carbonate core plugs used in the experiments. The results could be further improved by incorporating detailed porosity data. However, it is uncertain whether a simulation model constructed to represent a carbonate can sufficiently capture the essential topological features of these heterogeneous carbonate samples. Nevertheless, results show that composition of the fluids initially present in the core plug and reaction frequencies of the reactions play important roles in fluid–rock–gas interaction leading to changes in rock properties. [Xu et al. \(2004\)](#) reported similar results for sandstones. Blockage effect of calcite and salt particles also play a major role in porosity, thus in permeability alteration trend. Finally, it can be concluded that the calibrated model is capable of reflecting the natural behavior of the CO₂ injection process for all cases and conditions.

Fig. 2 Calibrated model results for a heterogeneous vertical oriented core plug experiment

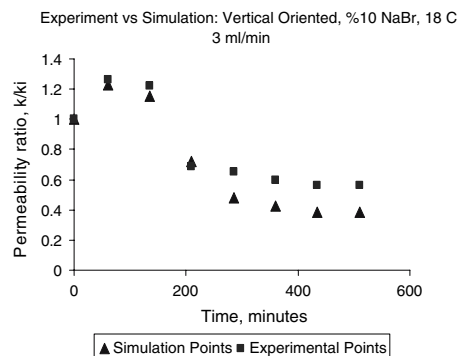


Fig. 3 Calibrated model results for a heterogeneous horizontal oriented core plug experiments

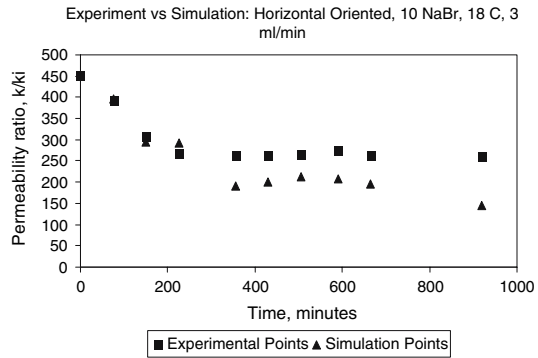


Fig. 4 Calibrated model results for homogenous vertical oriented core plug experiments

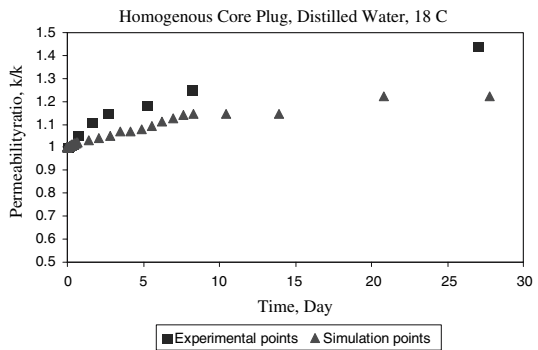


Table 1 Simulation parameters

Number of grids	14 × 15 × 24
Horizontal permeability, md	90
Vertical permeability, md	20
Average effective porosity, %	23.8
Core pressure, kPa	101
Core temperature, °C	18
Injection rate, kg/min	2.507 × 10 ⁻³

2.3 Problem Setup

In order to show the abilities of the model data from one of the matched experiments (Table 1) was used. In this experiment, CO₂ was injected in a continuous manner (3 cc/min) from the bottom of the simulation model. It was observed that chemical reactions occurred preferentially in the center of the core where CO₂ injection is performed (Fig. 5). At this location, the sizes of the grids are relatively smaller than the ones located near the borders. Dissolution and increase in permeability were observed especially at the inlet face of the core plug while at some grid blocks permeability impairment was observed. At the bottom and the top of the core some permeability decreases were also observed. As the injection continued CO₂–rock–brine interaction resulted in various nonuniform dissolution patterns and in some cases re-precipitation and permeability reduction. Thus, it is an unstable dissolution process leading to different dissolution regimes. This unstable dissolution process creates preferential flow patterns so-called wormholes.



Fig. 5 Time dependent (a) permeability (md) change, (b) CO₂ saturation evolution, (c) brine density (kg/cm³) change, (d) adsorbed CO₂ (ppm), and (e) concentration of bicarbonate (g-mole/cm³) (bottom). Model results of experiment 3 at 10, 30, 65, 105, 145, and 180min (left to right of each image)

A cutting plane at the middle of the core was used to examine the profiles on a vertical plane (Fig. 5). It was observed that injected CO₂ moved toward to top of the core while some amount was dissolved in brine. Figure 5 shows the vertical movement of gaseous CO₂, caused by buoyancy forces, and time-dependent gas saturation distribution. It could be seen that after the vertical movement, free phase gas accumulated at the upper part of the core for a while. Then, CO₂ started to dissolve into water. Thus, gas saturation at the top portion of the core decreased. It is known that dissolution of CO₂ into water increases the density of the brine (Pruess and Xu 2001). Change in water mass density by time observed in the simulation shows that as the CO₂ dissolves in brine, the density of brine increases (Fig. 6). This goes on until CO₂ can not be dissolved any more in the brine. Water density increased with time as free phase gas amount was decreasing. As the mass free phase gas at the upper portion of the core dissolved into brine, density of the fluid increased and started to migrate downwards and replaced by unsaturated brine. Results of core-scale simulation reveals that adsorption of CO₂ is less pronounced compared to other trapping mechanisms. Adsorption of CO₂ takes place where CO₂ is in free gas phase. Figure 5 shows the concentration of bicarbonate ions in brine with time. As seen from this figure solid concentration of the bicarbonate in the core plug continuously increased. This means dissolution takes place at the flow path of the CO₂, especially near the inlet. These results are in accord with aforementioned experimental observations (Izgec et al. 2007).

2.4 Sensitivity Analyses

2.4.1 Chemical Reaction Rate Sensitivity

A sensitivity analysis regarding reaction rates can give some understanding of its impact on the overall geochemical behavior of the system. The results of a sensitivity analysis of the effect of forward reaction (dissolution) and backward reaction (re-precipitation) frequencies of Eq. 6 on change in rock properties are given in Fig. 7 (forward and backward reactions). In this figure, average change of permeability in the entire simulation grid is plotted. These results reveal that as the reaction frequency of forward reaction (dissolution) increases the amount of dissolved particles, the total amount of particles that are blocking the throats increases. This leads to decrease in permeability. Meanwhile dissolution of rock leads to

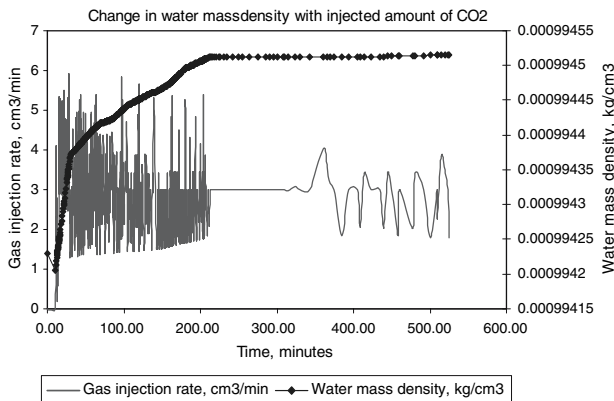


Fig. 6 Time dependent brine density change during the continuous injection of sub-critical CO₂

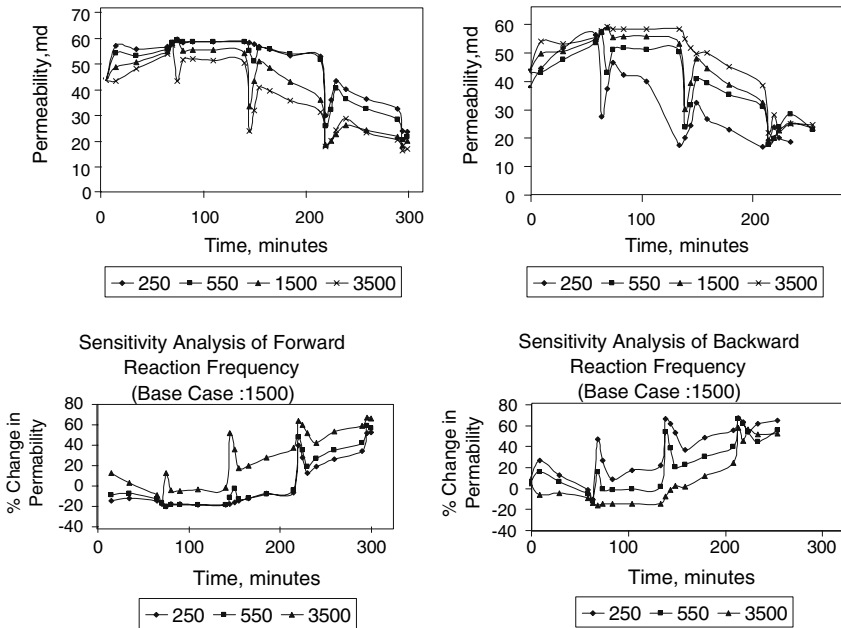


Fig. 7 Sensitivity analysis of forward (left) and backward (right) reaction frequencies

some increase in permeability. But it could be speculated that the effect of particles blocking the pore throats is stronger compared to the permeability increase due to dissolution of rocks. On the other hand, it is observed that increase in backward reaction (re-precipitation), frequency is accompanied with increase in permeability. Backward reaction frequency affects the precipitation rate of insoluble carbonate precipitates. Thus, a decrease in the amount of dissolved particles leads to a decrease in the blockage probability of individual pore throats. During some time-steps, reductions in permeability due to the precipitation of insoluble carbonate precipitates are observed for a number of individual grid blocks. Precipitation of these species leads to decrease in porosity, thus permeability. It could be speculated that the effect of decrease in total amount of particles blocking the pore throats is stronger compared to the permeability reduction due to precipitation and mineralization. The results of sensitivity analysis also reveal that doubling the forward reaction frequency may shift permeability alteration trend about (–) 10% from its previous value for some time-steps. On the other hand, doubling the backward reaction frequency may lead to a 30% increase in permeability of an individual point on permeability alteration trend. As indicated by [Xu et al. \(2004\)](#), CO₂ sequestration by matrix minerals varies considerably with rock type. They further reported that the accumulation of carbonates in the rock matrix lead to a considerable decrease in porosity that in turn adversely affected permeability. As a final remark, the porosity, pore size distribution, and the pore geometry of the rock may also affect the permeability alteration.

2.4.2 Kozeny–Carman Power

Instantaneous permeability values are obtained using these definitions with a Kozeny–Carman type relationship (Eq. 14). Kozeny–Carman type equations were previously used by [McCume et al. \(1979\)](#), [Itoi et al. \(1987\)](#) and [Lichtner \(1996\)](#). [Pange and Ziauddin \(2005\)](#) used a modified

Kozeny–Carman approach to model experimental dissolution patterns observed in carbonates during acidizing process. The application of Eq. 14 with exponents 3, 5, and 12 delineates the possible range of k – ϕ relations (Kühn 2004). The exponent 3 represents clean formations with relatively smooth shaped grains. An exponent of approximately 5 has been determined for anhydrite precipitation found in rock samples of deep geothermal aquifers from Northern Germany. The mineral deposit developed in this case in geological time scales. On the contrary, an exponent of 12 has been determined in core flooding laboratory experiments, representing the technical time scale, where anhydrite dissolved and subsequently precipitated within a temperature front (Bartels et al. 2002). Note that permeability calculated from porosity using Kozeny–Carman type of equation assumes that tortuosity is constant. In practice, however, as carbonic acid dissolves calcite and then precipitates again, the tortuosity should change continuously. For the model developed through this study, permeability is changing continuously because of blockage effects of species produced through chemical interactions among rock–fluid–CO₂. Model calculates instantaneous porosity using void and fluid porosity concepts. The results of a sensitivity analysis of the effect of Kozeny–Carman power, c , on change in rock properties are given in Fig. 8. Results reveal that increasing the Kozeny–Carman power four times may result in up to 43% change in individual points on the permeability alteration trend.

Fig. 8 Sensitivity analysis of Kozeny–Carman power

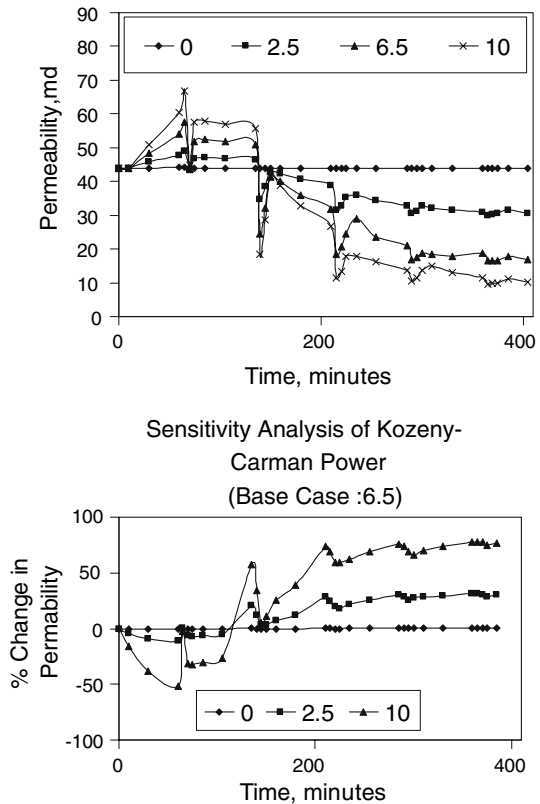


Table 2 Dependence of reaction or mass transfer rate on permeability

Effective permeability, md	Reaction rate factor, 1/min
10	2.5
20	1
750	0.75
1000	0.5

2.4.3 Blockage Effect of Particles

In computerized tomography monitored CO₂ injection experiments in carbonate core plugs conducted by [Izgec et al. \(2007\)](#), in some parts of the core plugs, a sharp decrease in permeability was observed. They attributed the permeability decrease to the blockage effect of salt and bicarbonate particles. Thus they concluded that salt may precipitate too. The model developed in this study involves the non equilibrium blockage by captured solid (nonfluid) components. Particles may block pore throats through two different mechanisms: bridging and plugging that can be estimated by two empirical relationships ([Hibbeler et al. 2003](#)). If pore diameter is less than two times particle diameter, then a bridge will form. If the particle diameter is between 1/3rd and 1/7th the size of the pore diameter then it will plug. Using the relationship that the square root of permeability is approximately equal to pore diameter in microns, plugging range of calcite particles can be calculated. For the heterogeneous carbonate core plugs for which the CO₂ injection experiments were modeled, the pores and throats have small (0.64 μm) to large values (10.65 μm) ([Izgec et al. 2007](#)). Thus, the small calcite particles with a particle diameter larger than 0.64 μm could plug the smaller pores along the flow path for all cores except for the homogeneous core plug. Several simulation tests were conducted to achieve an acceptable blockage effect relation in order to calibrate the model. Table 2 summarizes the optimum reaction flow restrictions due to blockage effect of dissolved particles for several permeability values.

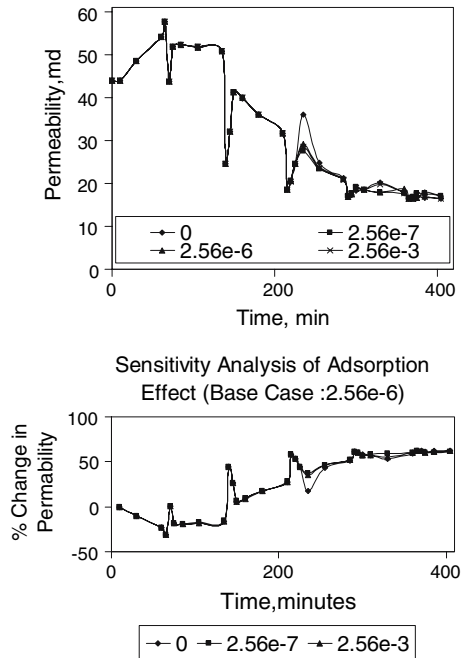
2.4.4 Adsorption Rate of CO₂

Adsorption of CO₂ was modeled via compositional dependence specified by the aforementioned Langmuir isotherm coefficients. As seen from the results, absence or presence of adsorption of CO₂ in carbonates has a small effect on permeability alteration trend (Fig. 9). [Wang and Thomson \(1995\)](#) showed that calcite decomposition at high temperatures (i.e., 440 and 560°C) is enhanced when steam is present due to steam's adsorptive properties which are faster and more significant than CO₂ adsorption. They reported that the hypothesis of a quantitative Langmuir type model is compatible with the kinetic data. [Zijlma et al. \(2000\)](#) reported similar results for the influence of H₂O and CO₂ on the reactivity of limestone for the oxidation of NH₃ at high temperatures. Although the temperature range considered in this study is much lower, the results are in accord with the aforementioned studies. Further research is required on this topic in order to generalize the findings.

2.4.5 Initial Concentrations of Species

The effect of initial concentrations of species should be studied to understand and model the process. In Eqs. 1–5, the concentration of CO₂ determines how much bicarbonate will be produced since water and the carbonate amounts in a carbonate reservoir is much larger compared to the concentration of CO₂. The solubility of carbon dioxide in water and hence its

Fig. 9 Sensitivity analysis of CO₂ adsorption

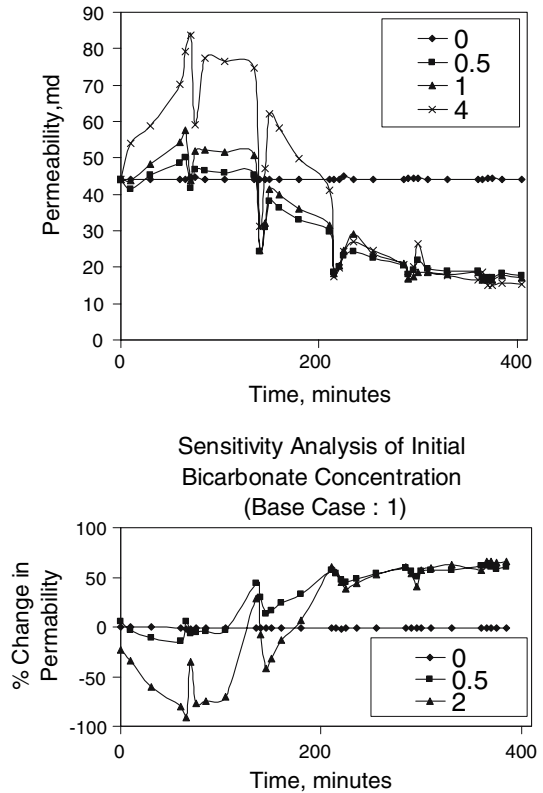


concentration increases with increase in pressure, and decreases with increase in temperature. The solubility of CO₂ is also reduced if the water contains some dissolved solids. Another concept is related to mobilization velocity of particles. Even before the beginning of injection of CO₂, some small particles will already be present in the porous medium. When CO₂ and/or water is injected into the porous medium and the water velocity reaches its mobilization velocity, these particles may move and cause some plugging, hence reducing the permeability of the porous medium. The results of a sensitivity analysis of the effect of initial concentrations of soluble bicarbonate on change in rock properties are given in Fig. 10. Permeability may drastically change if the initial concentration of bicarbonate is high. Thus mobilization of dissolved particles is the effective concept for the permeability alteration behavior.

3 Discussion

Due to numerous uncertainties and approximations such as chemical complexity of injection induced water–CO₂–rock interaction processes, current state-of-the-art numerical models cannot give a complete quantitative prediction of geochemical evolution of CO₂ injection. For example, since the chemical reaction kinetics of heterogeneous reactions are scale and history-dependent, kinetics cannot be reliably quantified. The lack of sufficiently detailed published reaction kinetics data further complicates the problem. In order to avoid adverse effects from CO₂ separating into liquid and gas phases in the injection system, geologic disposal of CO₂ into aquifers would be made at supercritical pressures; however, only a few studies do examine reactions in host rock in response to addition of CO₂ under supercritical conditions. In geochemical modeling studies that incorporate kinetic rate laws

Fig. 10 Sensitivity analysis of initial bicarbonate concentration



(Perkins and Gunter 1995; Gunter et al. 2000) and one study combining experiment and modeling (Gunter et al. 1997), dissolution of silicate minerals in a brine and precipitation of carbonate are reported. In the latter study, however, experiments consisted of month-long batch reactions at 105°C and 90 bars in which no silicate dissolution textures or new major reaction products were observed. With few exceptions (i.e., Gunter et al. 1997; Kaszuba and Janecky 2000; Kaszuba et al. 2001, 2003), there is no published experimental evaluation of geochemical reactions that occur within a supercritical CO₂-brine-aquifer system at reservoir temperature and pressure. For supercritical CO₂ injection at 200°C and 200 bars, Kaszuba et al. (2003) reported that the reaction among supercritical CO₂, brine, and rock exhibits relatively rapid kinetics that are similar to rates measured in systems where gaseous CO₂ is injected. However, numerical modeling is an excellent resource for analyzing and evaluating long term CO₂ sequestration and geochemical performance. A critical evaluation of modeling results can provide useful insight into sequestration mechanisms and controlling geochemical conditions and parameters. Based on the aforementioned observations one could expect that supercritical CO₂ injection will produce similar results to those presented in this study, but further research is required to verify this conclusion.

4 Conclusions

Results of core scale numerical modeling study showed that:

1. STARS can be used for numerical modeling of CO₂ sequestration in deep saline aquifers.
2. Either a permeability improvement or a permeability reduction can be obtained through the injection of CO₂ into carbonate aquifers. The trend of change in rock properties is very case dependent because it is related to distribution of pores, brine composition and as well the thermodynamic conditions.
3. Precipitation process may change permeability and porosity drastically.
4. Composition of the fluids initially present in the formation and reaction frequencies of dissolution and deposition play important roles in fluid–rock–CO₂ interaction leading to change in rock properties. In this regard, adsorption is less significant compared to others.
5. For the short period of simulation time considered in this paper hydrodynamic (free gas phase) and solubility (dissolved) trapping is more pronounced compared to mineral trapping of CO₂ and adsorption of CO₂. Adsorption of CO₂ takes place where CO₂ is in free gas phase.

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