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Foliar and litter needle carbon and oxygen isotope compositions relate to tree growth of an exotic pine plantation under different residue management regimes in subtropical Australia

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Abstract

Background and aim Significant differences in tree growth were observed in an exotic pine plantation under different harvest residue management regimes at ages 2-10 years. However, the variations in tree growth between residue management treatments could not be explained by soil and foliar nutrient analyses, except by potassium (K) concentration. Therefore, this study determined the carbon isotope composition (δ^{13} C) and oxygen isotope composition (δ^{18} O) of current and archived foliar samples from the exotic pine plantation to establish relationships with foliar K concentration and tree growth indices as a means to determine changes in stomatal conductance (g_s) and photosynthetic rate (A_{max}) or water use efficiency (WUE), and therefore understand the variations in tree growth across treatments.

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Present Address: S. S. Tutua Zai Na Tina Centre for Organic Systems, Kastom Gaden Road, East Honiara, Solomon Islands *Methods* The harvest residue treatments were: (1) residue removal, RR₀; (2) single level residue retention, RR₁; and (3) double level residue retention, RR₂. Foliar δ^{13} C and δ^{18} O were determined for samples at ages 2, 4, 6 and 10 years, and the atmospheric ¹³C discrimination (Δ^{13} C), intercellular CO₂ concentration (C_i) and WUE were determined from the δ^{13} C data. Litter needle δ^{13} C and δ^{18} O were also determined over 15 months between ages 9 and 10 years. These parameters or variables where correlated to each other as well as to the periodic mean annual increment of tree diameter at breast height (PAID) across the treatments and over time. Foliar δ^{13} C and δ^{18} O were also related to published data of foliar K concentrations of the same trees.

Results Significant variations of foliar δ^{13} C, and therefore WUE and Δ^{13} C, across treatments were only observed at ages 4 and 10 years old, and foliar δ^{18} O at age 4 years old only. The results showed increasing foliar δ^{13} C, δ^{18} O and WUE, and decreasing Δ^{13} C and Ci, from RR₀ to RR₂ treatments. However, while the WUE was positively related to the PAID and PAIB at age 4 years, it was negatively related to PAID and PAIB at age 10 years old. Litter needle δ^{13} C, indicative of WUE, was also negatively related to the PAID at age 10 years old. At age 4 years, foliar δ^{13} C and δ^{18} O were positively related with a steep slope of 7.70 ‰ across treatments, and that both isotopes were positively related to foliar K concentrations. Similarly, δ^{18} O was negatively related to the Δ^{13} C. No significant relationship can be determined between foliar δ^{13} C, or Δ^{13} C, and δ^{18} O at age 10 years old. In addition, WUE

was increasing (p < 0.001) and Δ^{13} C and C_i decreasing (p < 0.001) with decreasing PAID over time. *Conclusions* The variations at age 4 years in foliar δ^{13} C or Δ^{13} C and δ^{18} O and increasing WUE with increasing growth rate suggest growth induced water-stress with increasing residue-loading rate as a result of the nutritional effect of the harvest residues on tree growth. At age 10 years, the negative relationships between WUE and PAID indicate nutrient limitation has an over-riding effect on δ^{13} C variations rather than g_s. This was due to the lack of a significant relationship between foliar Δ^{13} C and δ^{18} O at this age, as well as over time.

Keywords Exotic pine plantation \cdot Foliar δ^{13} C and δ^{18} O \cdot Harvest residue

management \cdot Potassium \cdot Stomatal conductance \cdot Tree

growth

Introduction

Plant organic matter carbon (C) and oxygen (O) isotopic ratios (δ^{13} C and δ^{18} O, respectively) are widely used to determine the influence of genetic and environmental factors on plant growth (Xu et al. 2000; Warren et al. 2001; Barbour et al. 2002; Prasolova et al. 2005; Keitel et al. 2006; Huang et al. 2008a; Hasselquist et al. 2010). They are regarded as integrative measures of ecophysiological processes over the period in which the plant material was formed, and therefore can better account for tree growth than instantaneous measurements of stomatal conductance, water use efficiencies and net photosynthesis (Zhang and Cregg 1996; Prasolova et al. 2003; Xu et al. 2003; Barbour 2007; Huang et al. 2008b).

The theory behind varying plant δ^{13} C composition as a result of varying soil moisture conditions is well established (Farquhar et al. 1982; Farquhar and Richards 1984; Barbour et al. 2000; Warren et al. 2001). During photosynthesis, the C fixing enzyme, ribulose bisphosphate carboxylase/oxygenase (Rubisco), discriminates against the heavier ¹³CO₂. This discrimination, however, diminishes as the leaf internal CO₂ concentration (C_i) decreases (Farquhar et al. 1982; Powers et al. 2009; Hasselquist et al. 2010), such as when stomatal conductance (g_s) is decreasing in response to low soil moisture. Thus, a significant correlation between plant δ^{13} C and soil moisture availability can be established (Farquhar and Richards 1984; Xu et al. 2000; Warren et al. 2001; Huang et al. 2008c; Ibell et al. 2013).

The C_i , however, can also be influenced by the photosynthetic capacity, therefore confounding the relationship between g_s and $\delta^{13}C$ of plant tissues. This limitation can be overcome by the combined determination of δ^{13} C and δ^{18} O, to separate the effects of g_s and photosynthesis on δ^{13} C variation, since the δ^{18} O signature is not dependent on Rubisco activity (Scheidegger et al. 2000; Xu et al. 2000; Keitel et al. 2003; Barbour 2007). The ¹⁸O fractionation theory had been discussed extensively in Barbour et al. (2000) and Barbour (2007). In brief, it suggests, in general, that plant δ^{18} O is negatively related to g_s , and that a positive relationship between δ^{18} O and δ^{13} C indicates that δ^{13} C is largely driven by g_s (Saurer et al. 1997; Farquhar et al. 1998; Barbour et al. 2000; Keitel et al. 2003; Huang et al. 2008a, c). A recent study, however, showed that relationships between WUE and growth can also be used to differentiate photosynthetic capacity (A_{max}) from changes in gs due to water stress (Silva and Anand 2013).

A number of studies have indicated that tree growth and bulk stem wood or tree ring cellulose δ^{13} C are related (Dupouey et al. 1993; McNulty and Swank 1995; Livingstone and Spittlehouse 1996; Garcia-G et al. 2004; Fernandez et al. 2006; Sun et al. 2010; Gomez-Guerrero et al. 2013). Furthermore, studies of conifer plantation species showed that whole leaf or foliar δ^{13} C could be related to cumulative tree growth (Högberg et al. 1993; Xu et al. 2000). Limited studies, however, have successfully related either whole stem wood or leaf δ^{18} O to tree growth (Xu et al. 2000, 2009). In addition, the use of leaf litter δ^{13} C and δ^{18} O, which are likely to be representative of longer eco-physiological history, has been largely unexplored.

A study at ages 0-6 years (Simpson et al. 2003), and our recent study at age 10 years of the same plantation (Tutua et al. 2008), on the impact of residue management on tree nutrition and growth showed significant variation among treatments of the cumulative tree growth and periodic mean annual increment (PAI) prior to the peaking of growth rate. This variation of tree growth between treatments could not be related to foliar nutrient concentrations, other than foliar potassium (K) concentration, the only nutrient that showed significant residue treatment effects in both the soil and the foliage (Tutua et al. 2008). It was speculated, however, that the variation of foliar K concentration was a response to tree water stress (Prasolova et al. 2005; Fernandez et al. 2006), where larger or fast growing trees in the residue retention treatments were much more water stressed than those in the no-residue treatments (Tutua et al. 2008), consistent with studies showing growth-induced water stress due to the high demand of faster growing trees for soil water when N and P supply are non-limiting (Högberg et al. 1993). Potassium is important for maintaining leaf water and pressure potential, allowing photosynthesis to proceed under water stressed conditions (Ashraf et al. 2001; Pervez et al. 2004). Therefore, foliar K is expected to relate to foliar δ^{18} O and δ^{13} C. On the other hand, the variation in foliar K concentrations could merely be due to a luxury uptake of K under the residue retention treatments.

Studies that have clearly separated the mulching or nutritional effects of logging harvest residues through their influence on gs and/or Amax and therefore foliar δ^{13} C and δ^{18} O and their relationships to tree growth are limited. Therefore, this study aimed to better understand the impact of harvest residues on tree growth, in relation to the soil moisture or nutrition hypotheses, through a retrospective analysis of foliar and litter needle δ^{13} C and δ^{18} O. The main objectives of this study were to: (1) determine the influence of residue management on the foliar $\delta^{13}C$ and $\delta^{18}O$ of archived and current year samples; (2) establish relationships among the δ^{13} C, or WUE and Δ^{13} C, δ^{18} O, foliar K concentrations and tree growth parameters to assess whether gs and Amax or water and nutrient limitations could explain the variation in tree growth under different residue management regimes; and (3) explore the potential of using litter needle $\delta^{13}C$ and $\delta^{18}O$ as a tool for understanding the influence of environmental factors on tree growth in the exotic pine plantation of Subtropical Australia.

Methodology

Site description and experimental design

This study was carried out in an existing harvest residue management experiment established in July 1996. The details of this site and experimental design were as described by Simpson et al. (2003) and Chen and Xu (2005). In brief, the site was located at Toolara State Forest ($26^{\circ}00'$ S, 152° 49'E), southeast Queensland, Australia. It is generally flat, with a deep sandy soil classified as gleyic acrisols (FAO 1974). The climate is humid sub-tropical with a mean annual rainfall of 1,354 mm, with 56 % falling in December to March (Simpson et al. 2003). The July to September period is relatively dry, which may extend to November (Xu et al. 2000). The summers are hot and moist with a midsummer mean daily temperature of 24.9 °C and a relative humidity of 70 %, while the winters are mild, with a mid-winter mean daily temperature of 14.0 °C and a mean relative humidity of 64 %. Exotic pine plantation trees in this region often experience both well-watered and water-limited conditions in a year, even in a wet summer season (Xu et al. 2000; Prasolova et al. 2005; Ibell et al. 2013). In addition, annual rainfall has been below average for the last 10 years.

The current experiment is a randomised complete block design with six treatments and four blocks. Gross plots are 12 rows by 12 trees at 3×3 m spacing and net plots are 8 rows by 8 trees (0.058 ha). While there are 6 treatments (Simpson et al. 2003), this study focussed on the harvest residue loading rates, which included: (1) residue removal+50 kg P ha⁻¹ added; (2) residues retained+50 kg P ha⁻¹ added; and (3) double quantities of residues retained+50 kg P ha⁻¹ added. Moving harvest residues from the residue removal treatments to plots with harvest residues created the double residue retention treatment (Simpson et al. 2003). The three treatments are referred to as RR₀, RR₁ and RR₂. respectively. The plots were planted in mid 1996 with the F1 hybrid between slash pine (Pinus elliottii var. elliottii) and Carribean pine (Pinus caribeae var. *hondurensis*) seedling stocks from 6 different families. Each family was randomly allocated a row, with a double up of 2 families per row in 2 rows to make up the 8 rows.

Measurement of growth and foliar and litter samplings

Detailed tree growth measurements and foliar and litter sampling methods were as described by Tutua et al. (2008). The growth indices, diameter at breast height (DBH), basal area (BA), tree height (HT) and periodic mean annual increments (PAI) over a 2 year period of DBH (PAID), BA (PAIB) and HT (PAIH), were determined at ages 2, 4, 6, 8 and 10 years (Tutua et al. 2008). Periodic annual increments (PAI), rather than annual increments (AI) of growth was used because tree measurements and sampling were only carried out in alternate years.

Foliar samples at age 10 years and archived samples at ages 2, 4 and 6 years (1998, 2000 and 2002, respectively; no foliar sampling in 2004) were all collected from the northward facing (sunny) side of the tree canopy (Xu

et al. 2000; Tutua et al. 2008), and sampling was conducted on the same trees over the 10 years. Fifty fascicles of the most recent, fully expanded needles (approximately 1 year old) were collected from the 4 trees within a plot and bulked as one sample (Simpson et al. 2003). Even though the 4 trees within a plot may represent different families, pooling the samples was expected to minimise any variations in isotopic compositions between the families. In addition, a study of the F1 hybrid trees indicates that environmental effects on δ^{13} C are usually greater than familial differences (Prasolova et al. 2005). Litter needles were collected quarterly from July 2005 to June 2006 by Tutua et al. (2008). All plant materials were oven dried at 60 °C for 5 days before ground to powder in a puck and ring mill.

Chemical analyses

Foliar and litter C isotope compositions (δ^{13} C) were determined by an Eurovector 3000 elemental analyser (Milan, Italy) coupled to a GVI Isoprime mass spectrometer (Manchester, UK). The O isotope compositions (δ^{18} O) were determined by a VARIO EL III elemental analyser (Hanau, Germany) coupled to a Sercon Hydra 20–20 mass spectrometer (Crewe, UK). The foliar and litter δ^{13} C and δ^{18} O were calculated relative to the PDB and IAEA VSMOW standards, respectively (Barbour et al. 2000; Xu et al. 2000; Huang et al. 2008a). Foliar K concentration data for the same trees and year was obtained from Tutua et al. (2008).

Precipitation measurements

Rainfall data for the site was obtained from January 1997 to September 2006, the last foliar sampling date. Total growth period rainfall of 12 months before the foliar sampling was carried out was also determined to correlated with the isotopic variables. This was because this rainfall would have influenced the δ^{13} C and δ^{18} O of the 1 year old foliar needles as they formed.

Δ^{13} C, C_i and WUE calculations

Foliar δ^{13} C is a good indicator of WUE for comparisons across treatments in a given year. However, since we were also interested in the long term temporal variations in isotopic compositions and tree growth parameters, we calculated foliar Δ^{I3} C using the foliar δ^{13} C values to take into account the temporal variations of atmospheric CO_2 and $\delta^{13}C$ (Silva and Horwath 2013). The physiologically relevant $\Delta^{13}C$ is expressed as:

$$\Delta^{13}C = a + (b-a) (C_i/C_a)$$

where a is the discrimination against atmospheric ${}^{13}\text{CO}_2$ during diffusion through the stomata (-4.4‰), b is the net discrimination due to carboxylation (-27‰), C_i is the internal CO₂ concentration or partial pressures and C_a is the atmospheric CO₂ concentration for a given year. The C_a value for each year of sampling was obtained from direct measurements from the South Pole at the website of the Carbon dioxide Information Analysis Centre (Sun et al. 2010). The C_i values were estimated from the following formulae (Silva and Anand 2013) expressed as:

$$C_{i} = C_{a} \big[\delta^{13}C_{plant} - \delta^{13}C_{atmosphere} + a \big] / (a-b) \big]$$

where $\delta^{13}C_{plant}$ and $\delta^{13}C_{atmosphere}$ are the plant and atmospheric isotopic carbon ratios, respectively. The $\delta^{13}C_{atmosphere}$ value for each year was estimated from the exponential function given in Feng (1998). Finally, the WUE was calculated following the derivations of C_i as follows:

WUE =
$$C_a[1-(C_i/C_a) \times 0.625]$$

where WUE, which is defined as the ratio between carboxylation and stomatal conductance, is a function of C_i and C_a , and that the conductance of CO_2 molecules is 0.625*g* for leaf conductance to water vapour (Silva and Anand 2013).

Statistical analyses

An analysis of variance (ANOVA) was carried out on δ^{13} Cand δ^{18} O and variables such as Ci, Δ^{13} C and WUEto detect differences between the treatments at p < 0.05. The least significant difference (LSD) test at p < 0.05 was carried out to determine the degree of the variations between the treatments. Pearson's correlation and linear regression analyses were carried out to determine relationships among tree growth indices and foliar δ^{13} C, Ci, Δ^{13} C, WUE and δ^{18} O and, and . Litter δ^{13} C sampled from July 2005 to July 2006 were regressed with foliar δ^{13} C and δ^{18} O sampled in July 2006 and PAID (2004–2006). The Statistix software (Version 8.0) was used for all the statistical analyses. Since we proposed that the variations in foliar K concentrations was a response to water stress, we

regressed foliar K concentrations data obtained from the same trees in Tutua et al. (2008) with ¹³C and ¹⁸O in this present study to assess the validity of the proposition.

Results

Precipitation measurements

Figure 1 showed the rainfall distribution at the trial site from January 1997 to September 2006, the last sampling date of this study. The rainfalls for each year was below the annual mean for this site, and in general decreased over time, with the lowest 12 months rainfall in 2002. Although year 2000 had the highest rainfall, the last 2– 6 months prior to sampling were relatively drier compared to the other years. In contrast, rainfall was more evenly distributed throughout the 12 months before sampling in 1998, therefore remaining relatively high within 2–6 months prior to sampling. Growth period rainfall, defined as the rainfall over the period in which the pine needles formed (1 year), also declined from 1998 to 2006.

Foliar δ^{13} C, Δ^{13} C, WUE and δ^{18} O and tree growth

Significant treatment effect on foliar δ^{13} C occurred in 2000 and 2006 only (Table 1), where δ^{13} C was lower in the RR₀ treatment compared to the RR₁ and RR₂

treatments in both years. The maximum differences in δ^{13} C variations between the treatments were 0.90 ‰ and 0.87 ‰ in 2000 and 2006, respectively. The δ^{18} O, however, only showed significant treatment effects in 2000 when δ^{18} O was the lowest in the RR₀ treatment (p < 0.05) compared to the RR₁ and RR₂ treatments (Table 1). These 2 years represented the periods of peak growth (2000) and declining growth (2006) according to the PAIB and PAID trends reported in Tutua et al. (2008).

Calculations of C_i , $\Delta^{13}C$ and WUE from the foliar $\delta^{13}C$ data also showed significant variations among the residue management treatments in year 2000 and 2006 (Table 1) with significantly greater WUE, strongly reflecting trends in foliar δ^{13} , of trees in the RR₂ and RR₁ treatments than the RR₀ treatment. The results also showed a concomitant decrease in C_i and therefore $\Delta^{13}C$ of the trees with increasing residue loading rates in both years (Table 1).

Regression analyses of PAID and PAIB with foliar δ^{13} C and WUE in 2000 (age 4 years) showed a positive relationship with both variables (Fig. 2a–d). The results showed that WUE can explain about 36 % and 40 % of the increasing PAID and PAIB, respectively, at this stage. In contrast to year 2000, regressions of PAID and PAIB with foliar δ^{13} C and WUE in year 2006 (age 10 years), showed a negative relationship (Fig. 3a–d). The increasing WUE explained 70 % and 49 % of the declining PAID and PAIB, respectively, with increasing residue loading



Fig. 1 Rainfall distribution at Toolara State Forest from January 1997 to September 2006

regimes (RR_0 , RR_1 and RR_2), measured from 1998 to 2006. No foliar sampling was conducted in 2004

Treatments	Year				
	1998	2000	2002	2006	
Foliar δ^{13} C (‰)					
Residue removal (RR ₀)	$-30.70a^{a}$	-30.63b	-27.88a	-29.25b	
Single residue retention (RR ₁)	-30.80a	-30.30a	-28.00a	-29.00ab	
Double residue retention (RR ₂)	-30.73a	-30.25a	-27.75 <i>a</i>	-28.85a	
Foliar $\delta^{18}O$ (‰)					
Residue removal (RR ₀)	25.70 <i>a</i>	26.33 <i>b</i>	28.98 <i>a</i>	30.93 <i>a</i>	
Single residue retention (RR ₁)	26.53 <i>a</i>	30.48 <i>a</i>	27.63 <i>a</i>	27.20 <i>a</i>	
Double residue retention (RR ₂)	24.80 <i>a</i>	29.85 <i>a</i>	27.55 <i>a</i>	29.35a	
Δ^{13} C (‰)					
Residue removal (RR ₀)	22.64 <i>a</i>	22.50a	19.67 <i>a</i>	20.91 <i>a</i>	
Single residue retention (RR ₁)	22.77 <i>a</i>	22.17b	19.82 <i>a</i>	20.63 <i>ab</i>	
Double residue retention (RR ₂)	22.64 <i>a</i>	22.12b	19.55 <i>a</i>	20.49b	
Plant Ci (ppmv)					
Residue removal (RR ₀)	293.52 <i>a</i>	293.88 <i>a</i>	250.48 <i>a</i>	276.50a	
Single residue retention (RR ₁)	295.56a	288.61 <i>b</i>	252.87a	271.82 <i>ab</i>	
Double residue retention (RR ₂)	293.60 <i>a</i>	287.83 <i>b</i>	248.41 <i>a</i>	269.55b	
WUE (µmol mol ⁻¹)					
Residue removal (RR ₀)	43.86 <i>a</i>	45.72 <i>b</i>	75.11 <i>a</i>	63.75 <i>b</i>	
Single residue retention (RR ₁)	42.59 <i>a</i>	49.01 <i>a</i>	73.62 <i>a</i>	66.68 <i>ab</i>	
Double residue retention (RR ₂)	43.81 <i>a</i>	49.50 <i>a</i>	76.41 <i>a</i>	68.09a	

^a Means with the same letter are not significantly different (p>0.05)





Fig. 2 Relationships between foliar δ^{13} C and periodic annual increments of tree diameter at breast height (PAID) (a) or basal area (PAIB) (b), and between tree WUE and PAID (c) or PAIB (d)

rate and decreasing Δ^{13} C in the order: RR₀>RR₁>RR₂. These observations were also consistent with the decreasing C_i across the treatments (Table 1).

Pearson's correlations of tree growth indices and isotopic variables of the year 2000 data also yielded similar results as the regression analyses. These showed that the isotopic ratios and WUE were positively related to almost all growth indices, while the $\Delta^{13}C$ was negatively related to the growth indices (Table 2). In 2006, foliar δ^{13} C, Δ^{13} C and WUE showed stronger relationships than in year 2000 as indicated by the coefficients with all growth indices, except HT. However, there were no significant correlations between δ^{18} O and the growth indices, or between foliar δ^{13} C and δ^{18} O and between Δ^{13} C and δ^{18} O. This was in contrast to year 2000, where there was also a significant correlation between foliar δ^{13} C and δ^{18} O and between Δ^{13} C and δ^{18} O (Table 2). This was further illustrated by the regression analyses of foliar δ^{18} O with Δ^{13} C and foliar δ^{13} C, indicating that these parameters could explain at least 56 % of the variations in foliar δ ¹⁸O (p < 0.005) across the treatments (Fig. 4b and c), with a slope of 7.70% increase in δ^{18} O per 1.0% increase in δ^{13} C (Fig. 4b). This study also found that foliar δ^{13} C, δ^{18} O and K concentration in 2000 were significantly related (Fig. 4a). Increasing δ^{13} C and δ^{18} O across the treatments coincided with the concomitant increase in foliar K concentration (Fig. 4a). In addition, both foliar K concentration and δ^{18} O together improved the regression model and were able to explain about 67 % of the variation of δ^{13} C (p<0.01) in comparison to the foliar δ^{18} O alone. This result was in contrast to year 2006 where there were no relationship between foliar δ^{18} O and δ^{13} C (Table 2) or between foliar δ^{18} O and foliar K, except for foliar δ^{13} C and K concentrations (r= 0.87; p<0.001).

Long -term variations in isotopic compositions and tree growth

Table 2 also shows that both foliar $\delta^{13}C$ and $\delta^{18}O$ increased over the last 8 years (p < 0.005), with means ranging from -30.80% to -28.85% for $\delta^{13}C$, and from 24.80% to 30.93% for δ^{18} O from 1998 to 2006, respectively. Similarly, C_i and Δ^{13} C were also decreasing (p < 0.05 and p < 0.0005, respectively), while WUE was increasing over this time period (p < 0.0001). The increase in WUE was about 45-55 % from 1998 to 2006. However, Pearson's correlation analyses showed no significant correlation between foliar δ^{18} O and δ^{13} C or Δ^{13} C, which accounted for changes in C_a and $^{13}C_{\text{atmosphere}}$ concentrations over the 8 years (Table 3). Similarly, there were no significant correlations observed between foliar δ^{18} O and the declining PAID or PAIB, and even the foliar needle growth period rainfall. On the other hand, foliar δ^{13} C, and therefore WUE, showed very





Fig. 3 Relationships between foliar δ^{13} C and periodic annual increment of the basal area (PAIB) (a) or periodic annual increment of diameter at breast height (PAID) (b), and between tree WUE and PAID (c) or PAIB (d) at age 10 years of a F1 hybrid

exotic pine plantation under three harvest residue management regimes: (1) residue removal; (2) single residue retention; and (3) double residue retention

Growth indices	Variables			
	Foliar δ^{13} C	Foliar δ^{18} O	WUE	$\Delta^{13}C$
Year 2000 (age 4 years)				
Diameter at breast height (DBH) ^a	0.66 ^{*d}	0.73 **	0.66^{*}	-0.66*
Basal area (BA)	0.64 *	0.73 **	0.65^{*}	-0.65*
Height (HT)	0.63 *	0.61*	0.62^{*}	-0.62*
Volume index (VI) ^b	0.64 *	0.69 *	0.63*	-0.63*
Periodic annual increment of DBH (PAID) ^c	0.59 *	0.56*	0.60^{*}	-0.60*
Periodic annual increment of BA (PAIB)	0.62 *	0.69 *	0.63*	-0.63*
Foliar δ^{18} O	0.76 ***	-	0.75 **	-0.75**
Foliar $\delta^{13}C$	_	0.76 ***	_	-
Year 2006 (age 10 years)				
Diameter at breast height (DBH)	0.78 ***	ns	0.77 **	-77**
Basal area (BA)	0.80 ***	ns	0.80 **	-80**
Height (HT)	ns	ns	ns	ns
Volume index (VI)	0.70 **	ns	0.71**	-0.71**
Periodic annual increment of DBH (PAID) ^c	-0.81 ***	ns	-0.84 ***	0.84 ***
Periodic annual increment of BA (PAIB)	-0.68 *	ns	-0.70^{*}	0.70*
Foliar $\delta^{13}C$	_	ns	1.00***	1.00***
Foliar δ^{18} O	ns	_	ns	ns

Table 2 Pearson's correlation coefficients of tree growth indices with foliar δ^{13} C, δ^{18} O, Δ^{13} C and WUE at ages 4 and 10 years of the F1 hybrid exotic pine plantation under three harvest residue management regimes (RR₀, RR₁ and RR₂)

^a Measured at 1.3 m above ground level

^b Calculated according to Xu et al. (2000), where $VI = \pi DBH^2 H/12$

^c Mean of periodic growth calculated from growth over 2 years

^d Asterisks *, **and *** indicate significance at p < 0.05, 0.01, and 0.001, respectively, and ns not significant (p > 0.05)

strong negative correlations with the declining PAID and PAIB over the 8 years. These parameters were also very strongly correlated to the foliar needle growth period rainfall (Table 3), where the declining rainfall was positively related to the declining PAID or PAIB, and negatively related to the WUE.

Litter needle δ ^{13}C and δ ^{18}O

Litter needle δ^{13} C ranged from -29.53‰ to -29.13‰ throughout the seasons (Table 4). Although the variations in litter needle δ^{13} C between the treatments at each sampling were relatively small (0.27–0.33‰) compared to the foliar δ^{13} C, they were significant (*p* < 0.01) in three of the four seasons assessed. In each case, δ^{13} C increased in the following order: RR₀<RR₁<RR₂, a similar trend to that seen in the foliage. Significant differences in litter needle δ^{18} O between the treatments were observed in the January–March 2006 season only,

where higher δ^{18} O occurred in the RR₀ treatment compared to the RR₁ treatment (Table 4). However, this observation could not be related to litter needle δ^{13} C.

This study also showed that litter needle $\delta^{13}C$ and foliar δ^{13} C were related (Fig. 5a–b). However, only litter needle δ^{13} C from the July – September 2005 season was significantly related ($r^2=0.47$; p<0.05) to the PAID (2004–2006 period) (Fig. 5c). Regressing PAID with litter needle $\delta^{13}C$ for all seasons that had significant $\delta^{13}C$ variations showed that litter needle δ^{13} C, in general, could be related to the PAID ($r^2=0.26$; p<0.005) (Fig. 5d). Furthermore, Table 4 shows a temporal trend in the variation of litter needle $\delta^{13}C$ between the treatments. The occurrence or absence of significant variations in litter needle δ^{13} C between the treatments for each period appeared to closely follow the seasonal changes in the rainfall and temperature, where a greater variation between the treatments occurred in the periods with the lowest rainfall or highest daily temperature (Fig. 6).

22 24

-31.0

0

-30.6

Foliar δ^{13} C (‰)

0

0.7

0.6

0.4

0.3

34

32

30

28

26

24

22

Foliar K (%) 0.5

Foliar $\delta^{18} O(\%)$



0

-30.2

0

-30.4

0

(b)

-30.0



Fig. 4 Relationships between foliar K concentration and δ^{18} O or δ^{13} C (a), and between foliar δ^{13} C and δ^{18} O (b), and Δ^{13} C and δ^{18} O at age 4 years (year 2000) of a F1 hybrid exotic pine

-31.0

-30.8

Discussion

Foliar δ^{13} C and δ^{18} O variations

Harvest residue management clearly influenced the variations in foliar δ^{13} C and δ^{18} O in 2000, and δ^{13} C in 2006 (Table 1), as indicated by the consistent increase in foliar δ^{13} C in the order RR₀<RR₁<RR₂ in both years. Whilst the maximum variations in foliar δ^{13} C across the

plantation under three harvest residue management regimes: (1) residue removal; (2) single residue retention; and (3) double residue retention

treatments were only 0.90% and 0.87% in 2000 and 2006, respectively, they were consistent with those studies having small variations of 1.0% or less, yet were useful in explaining the changes in gs, Amax and tree growth (Olbrich et al. 1993; Fischer et al. 1998; Barbour et al. 2000; Xu et al. 2000). Thus the correlation of foliar δ^{13} C with the growth indices of the F1 hybrid exotic pine was consistent with other studies on conifer species, using foliar δ^{13} C (Högberg et al. 1993) or Δ^{13} C (Xu et al.

Table 3 Pearson's correlation coefficients of the relationships between long-term changes in foliar δ^{13} C and δ^{18} O, Δ^{13} C, WUE and periodic mean annual increment of tree diameter at

breast height (PAID) and tree basal area (PAIB) over a 10-year period in the F1 hybrid exotic pine plantation in subtropical Australia

Variables PAID ^a PAIB Foliar δ^{18} O Foliar δ^{13} C Δ^{13} C	WUE
Foliar δ^{18} O ns ns	
Foliar δ^{13} C -0.70^{***b} -0.62^{***} ns	
Δ^{13} C 0.83*** 0.66** ns -1.00***	
WUE -0.86*** -0.70*** ns 0.99*** -1.00***	
Rainfall 0.93 *** 0.86 *** ns -0.86*** 0.90***	-0.92***

^a PAID and PAIB were means of periodic annual increment of growth calculated over 2-year intervals

^b Asterisks *, **and *** indicate significance at p < 0.05, 0.01, and 0.001, respectively and ns not significant (p > 0.05)

Treatment	Sampling Periods				
	Jul-Sept05	Oct–Dec05	Jan-Mar06	Apr-Jun06	
Litter δ^{13} C (‰)					
Residue removal (RR ₀)	$-29.46b^{a}$	-29.41b	-29.50b	-29.53a	
Single residue retention (RR ₁)	-29.35b	-29.37b	-29.51 <i>b</i>	-29.26 <i>a</i>	
Double residue retention (RR ₂)	-29.19 <i>a</i>	-29.13 <i>a</i>	-29.17 <i>a</i>	-29.42a	
Litter δ^{18} O (‰)					
Residue removal (RR ₀)	21.48 <i>a</i>	21.65 <i>a</i>	21.88 <i>a</i>	21.18 <i>a</i>	
Single residue retention (RR ₁)	21.20 <i>a</i>	21.65 <i>a</i>	21.08 <i>b</i>	20.80 <i>a</i>	
Double residue retention (RR ₂)	21.05 <i>a</i>	21.33 <i>a</i>	21.35 <i>ab</i>	21.15 <i>a</i>	

Table 4 Litter needle δ^{13} C and δ^{18} O of a 10-year-old F1 hybrid exotic pine plantation under three harvest residue management regimes in subtropical Australia

^a Means assigned with the same letter in each column were not significantly different (p>0.05)

2000) and stem or wood cellulose δ^{13} C (Dupouey et al. 1993; McNulty and Swank 1995). Relating foliar δ^{13} C and WUE or Δ^{13} C to growth rate indices such as PAID in this study or basal area annual increments (BAI) has also been widely used recently (Sun et al. 2010; Gomez-Guerrero et al. 2013; Silva and Anand 2013). This was because relating isotopic variations to growth rates are

useful in explaining the dynamic changes in tree growth over time.

The absence of significant variations in foliar δ^{13} C and δ^{18} O among the treatments in some years could be due to high water availability (Keitel et al. 2003) in the 6–12 months leading to each sampling date, as reflected by the rainfall distribution (Fig. 1). In year 2000, the



Fig. 5 Relationships between foliar δ^{13} C sampled in 2006 and litter needle δ^{13} C sampled from the July to September 2005 period (a) and all periods that have significant δ^{13} C variations (Table 4) (b). Litter needle δ^{13} C was also significantly related to the periodic annual increment of diameter at breast height (PAID) of the 2004–

2006 period (age 8–10 years) as shown by the negative relationship between PAI.D and litter needle $\delta^{13}C$ sampled in the July–September 2005 period (c) and periods that have significant litter needle $\delta^{13}C$ variations (d)



Fig. 6 The occurrence of litter needle δ^{13} C variation between the treatments for each season or quarter from the 2005 to 2006 litter collection period was closely related to the rainfall (**a**) and temperature (**b**) patterns of the same periods in the previous year (2004–2005). The vertical bars are least significant differences (LSD_{0.05}), and the sampling periods were: (1) Jul–Sept 2005, (2) Oct–Dec 2005, (3) Jan–Mar 2006, (4) Apr–Jun 2006, and (5) Jul–Sept 2006

relatively drier 4 to 6 months prior to sampling was probably critical in the variation across the treatments, especially for δ^{18} O, which often occur in concurrence with the variations in δ^{18} C (Barbour et al. 2000). Rainfall distribution (Garcia-G et al. 2004), duration of drier conditions (Fernandez et al. 2006), N nutrition (Högberg et al. 1993; Huang et al. 2008b), ambient temperature and vapour pressure deficit (VPD) (Korol et al. 1999; Barbour et al. 2002) have been reported to influence tree foliar δ^{13} C and δ^{18} O.

Foliar isotopic compositions, WUE and K concentration and tree growth

The negative relationship between foliar δ^{13} C or WUE and PAID or PAIB at age 10 years (2006) (Fig. 3) was similar to a study of *Pinus greggii* using wood cellulose δ^{13} C (Garcia-G et al. 2004). Similarly, negative relationships between WUE, calculated from wood ring δ^{13} C, and BAI have also been reported (Gomez-Guerrero et al. 2013). This growth response indicates an environmental stressor affecting tree growth (Silva and Anand 2013), and the effect on the growth rate is more severe on the larger trees in the RR_2 (Tutua et al. 2008). The growth response was consistent with decreasing g_s , which was indicated from the decreasing C_i across the treatments in this study (Table 1). Reductions in C_i is more likely due to closing stomata in response to decreasing g_s , thus affecting the photosynthetic rate and therefore growth (Farquhar et al. 1982; Fischer et al. 1998; Barbour et al. 2000; Garcia-G et al. 2004). The greater WUE of trees in the RR₂ treatment (with least growth rates) in year 2006 indicated that gs was probably more affected rather than C assimilation, given the decreasing $\Delta^{13}C$ as C_i decreases (Table 1). The suggested decrease in gs with increasing residueloading rates, however, was unexpected against the background of published work, which suggests that the residues would conserve soil moisture (Proe et al. 1999; Scott et al. 2005). In addition, the lack of a significant relationship between foliar δ^{13} C or Δ^{13} C and foliar δ^{18} O in concurrence with a negative WUE/growth relationship indicated that other environmental stressors, other than reduced gs, could be more important drivers of foliar δ^{13} C variations (Silva and Anand 2013; Gomez-Guerrero et al. 2013). Studies have shown that nutrient limitations could cause such a growth response when the photosynthetic capacity is limited by leaf nutrient content and soil nutrient availability (Silva and Anand 2013). This interpretation is consistent with the nutrient analyses of the same trees in 2006, which showed that both foliar N & P concentrations were either marginal or below critical levels (Tutua et al. 2008). Furthermore soil nutrient analyses in year 2006 showed that although total N increased and were significantly greater in the RR1 and RR₂ treatments, mineral N pools (NH₄⁺ and NO₃⁻) only ranged from 9.8–10.9 mg kg⁻¹ to 0.26–0.36 mg kg⁻¹, respectively, which were much lower than those reported at age 6 of the same plantation (Chen and Xu 2005; Tutua 2009). Similarly, labile P pools at age 10 years had decreased compared to those at age 2 years old of the same plantation (Mathers and Xu 2003; Tutua 2009). The low availability of nutrients for plant uptake, however, could be due to low mineralisation rates, which may result from soil water deficit (Högberg et al. 1993). The influence of soil water deficit could not be ruled out given the decreasing growth period total rainfall and that the total annual rainfall at this site had been below

average over the 10 years of the plantation (Fig. 1). Nonetheless, nutrient limitations seemed to have an over-riding effect given the lack of a significant relationship between Δ^{13} C and foliar δ^{18} O (Gomez-Guerrero et al. 2013).

On the other hand, the positive relationship of foliar δ^{13} C or WUE with tree growth (Fig. 2a–d) across increasing residue loading rate in 2000 indicates a stimulus to growth (Högberg et al. 1993; Silva and Anand 2013). This stimulus is most likely to be the unlimited nutrient availability from the decomposing residues during the early growth stages, thus enhancing the photosynthetic capacity and WUE of the trees. This proposition was consistent with our other parallel residue management study adjacent to this plantation, which showed greater levels of available or labile N and P after 18 months where harvest residues were retained (Tutua et al. 2013).

However, the decreasing C_i and $\Delta^{13}C$ from RR₀ to RR₂, and the positive relationship between δ^{13} C or the negative Δ^{13} C and δ^{18} O relationship (Fig. 4) indicated that g_s was the main driver of the δ^{13} C variations across the treatments in 2000 (Fischer et al. 1998; Barbour et al. 2000; Scheidegger et al. 2000). This would be inconsistent with the gs theory as a mechanism for growth, where decreasing gs would lead to reduced photosynthesis and growth (Fischer et al. 1998; Barbour et al. 2000). In addition, higher g_s expected under residue retention treatments was supposed to cool the leaves, and therefore the canopy, resulting in increased internal CO₂ concentration (increased Δ^{13} C), and higher light-saturated photosynthetic rate of plants (Fischer et al. 1998). Therefore, we suspected that the increasing δ^{13} C and δ^{18} O with increasing tree growth rates (Fig. 4), was the result of higher growth rates in the RR1 and RR2 treatments, causing periodic water stress to which foliar δ^{18} O was sensitive. Growth-induced water stress had been suggested by some published work (Högberg et al. 1993; Xu et al. 2000; Garcia-G et al. 2004). It has been shown that the addition of an N source to an N-limited soil would increase foliage biomass and therefore growth rate, which in turn could lead to water stress as a result of growth demands for water (Högberg et al. 1993). This effect might explain the very steep slope of the $\delta^{18}O/\delta^{13}C$ relationship (7.70 % increase in δ^{18} O per 1% increase in δ^{13} C) in this study (Fig. 4b) compared to those reported by Barbour et al. (2000) (2.9 ‰) and Keitel et al. (2003) (1.11‰), where the higher growth demands for water established

a steep gradient of g_s across the treatments. We would speculate that growth demands for soil water might have increased soil water deficit (SWD) over time in the RR₁ and RR₂ treatments as suggested for in 2006, due to the below-average rainfalls over the last 10 years, which were probably not enough to pass through the thick organic/residue layer to reach the mineral soil as observed in other studies (Huang et al. 2008b). Nonetheless, these results confirmed the nutritional value of the residues during early growth, which could not be confirmed from foliar nutrient analyses across treatments (Simpson et al. 2003; Tutua et al. 2008) possibly due to dilution effects of the trees.

The significant $\delta^{13}C/\delta^{18}O/K$ relationships (Fig. 4a) in 2000 supported our proposition that significant K variation across treatments was probably a response to water stress (Xu et al. 2000; Tutua et al. 2008). This was due to the fact that foliar K concentrations were nonlimiting in all the treatments and therefore cannot be a driver of growth variations (Tutua et al. 2008). The result, however, is consistent with studies showing enhanced K supply under decreased gs (Fernandez et al. 2006). Potassium functioned through its maintenance of leaf water and pressure potential (Ashraf et al. 2001; Pervez et al. 2004), allowing photosynthesis to proceed under water stressed conditions. The weak relationship between foliar K concentration and growth indices in 1998 when rainfall was well distributed as reported in Tutua et al. (2008) is also consistent with the proposition that the foliar K variations were a response to water stress. Although lack of direct measurements of soil moisture content is a limitation of this study, the observations parallel a separate study of the F1 hybrid exotic pine clones, showing a correlation between foliar mineral, including K, concentrations and $\delta^{13}C$ in water stressed trees (Prasolova et al. 2005).

Long-term variations of foliar isotopic composition and WUE and declining tree growth

The long-term changes in foliar isotopic compositions and WUE over the 8 years also revealed some interesting trends relevant to long-term tree growth and climate change. The increasing δ^{13} C, WUE and δ^{18} O, or decreasing Δ^{13} C, over time were consistent with the declining growth rates and rainfall over the 10 years (Table 3). Again as shown in 2006, the increasing WUE and declining PAID over time (negative coefficients) (Table 3) was consistent with reduced g_s induced by SWD (Farquhar and Richards 1984; Barbour et al. 2000). This was supported by the decreasing C_i and rainfall with the declining PAID over time (Table 3). However, the lack of a significant correlation between Δ^{13} C and δ^{18} O over time indicated that the decrease in C_i and $\Delta^{13}C$ or increase in WUE could not be driven solely by gs. As indicated earlier, this growth response to increasing WUE could also be due to nutrient limitation on photosynthetic rate, increasing ambient CO₂ and warming associated with climate change (Sun et al. 2010; Gomez-Guerrero et al. 2013; Keenan et al. 2013). These are all possible causes given the increase in ambient CO₂ by 15 ppm from 1998 to 2006 and the apparent changing climate indicated by the below average rainfalls. The nutrient limitation proposition, however, was supported by our earlier report that the decreasing foliar N and P concentrations were significantly related to the decreasing PAID over time (Tutua et al. 2008). A number of studies reported that nutrient limitation is one of the causes of declining forest growth (Gower et al. 1996; Jokela and Martin 2000; Gomez-Guerrero et al. 2013; Silva and Anand 2013). Soil nutrient and moisture interactions are also important to consider given that soil moisture could influence nutrient mineralisation, mass movement and therefore uptake by the trees (Högberg et al. 1993). Regardless, these results have significant implications for climate change adaptations. They demonstrated that increasing ambient CO₂ from anthropogenic sources, warming and therefore increases in WUE do not always translate to tree growth when there are other limiting factors (Peňuelas et al. 2010; Silva and Anand 2013).

Litter δ ^{13}C and δ ^{18}O

The significant relationships between foliar and litter δ^{13} C across the treatments (Fig. 5a–b), demonstrates that a significant quantity of foliar δ^{13} C was part of leaf structural C, which remained with the litter after senescence. The apparent depletion in litter δ^{13} C in all seasons relative to foliar δ^{13} C in 2006, and the narrower range (0.27–0.33‰) across the treatments (Table 4), was reflective of C mobilisation or re-translocation during senescence. Regardless of this, litter needle δ^{13} C was able to show significant variations across the treatments in at least two seasons, with the highest δ^{13} C in the RR₂ treatment, consistent with the trend of foliar δ^{13} C in 2006. This consistency was further demonstrated in the significant relationship between litter δ^{13} C and PAID in

2006 (Fig. 5d). The negative relationship between litter needle δ^{13} C and tree growth in 2006 was reflective of the inverse relationship between foliar δ^{13} C and PAID or PAIB. While this might suggest moisture stress, and therefore g_s, as the controlling factor for the variation in foliar δ^{13} C and tree growth (Farquhar and Richards 1984; Zhang and Cregg 1996; Barbour et al. 2000), this proposition was weakened by the lack of a significant relationship between litter δ ^{13}C and δ ^{18}O across the treatments. It is more likely due to nutrient limitation as discussed above or the interactions of both stressors. This is because the occurrence or absence of litter needle $\delta^{13}C$ variations between the treatments across the seasons reflects the temperature and rainfall patterns of this site, whereby seasons with higher mean maximum temperatures and relatively low rainfalls of the previous year resulted in the significant variations in litter δ^{13} C of those seasons in the present year (Fig. 6). Although the litter δ^{13} C measurements were carried out for 15 months only, the results were consistent with studies of wood cellulose δ^{13} C and δ^{18} O variations, where increasing temperature and low rainfall resulted in increasing wood cellulose δ^{13} C and δ^{18} O (Korol et al. 1999; Barbour et al. 2002), similar to that with litter δ^{13} C in this study. Thus, litter δ^{13} C may be a potential indicator of environmental stressors and that its use could substitute for foliar sampling of taller trees.

This study has been able to elucidate the influence of environmental stressors on tree growth using foliar and litter isotopic compositions, including WUE, and their relationship to growth rates. However, perhaps future studies in this plantation need to carry out direct measurements of soil moisture and gas exchange to backup isotopic variations relating to environmental stressors. In addition there is a need to look into soil moisture and nutrient interactions, especially at mineralisation rates, microbial activity and understory vegetation competition for water and nutrients for a better understanding of long term growth patterns. Future work may also need to increase the sample size from 4 trees per plot in this study as well as conduct the study in more than one site for a broader understanding in subtropical conditions.

Conclusion

This study demonstrated that harvest residues can have a significant influence on the variations in foliar $\delta^{13}C$ and

 δ^{18} O of the F1 hybrid exotic pine trees. The foliar δ^{13} C or WUE variations and their relationships with periodic tree growth rate have been helpful in understanding the mechanism that regulates tree growth and especially plantation productivity. The significant relationships among foliar δ^{13} C. δ^{18} O and K during early growth clearly indicated growthinduced water stress of trees in the RR1 and RR2 treatments, a situation due to greater availability of N and P where residues were retained. This suggests that harvest residue retention in these plantations had a significant nutritional effect on tree growth gain observed over the 10 years. The negative relationship between foliar and litter $\delta^{13}C$ or WUE and growth rates at age 10 years indicated that nutrient limitation has an over-riding effect on the photosynthetic capacity of the trees. Similarly, the long term foliar isotopic compositions and WUE trends over the 10 years also indicated nutrient limitations as the cause of the decline in the growth rates. Litter needle δ^{13} C has the potential for studying environmental conditions in relation to tree growth as demonstrated in this study.

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