#### **ORIGINAL RESEARCH PAPER**



# **Seismic based characterization of total organic content from the marine Sembar shale, Lower Indus Basin, Pakistan**

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#### **Abstract**

The exploration and production of unconventional resources has increased significantly over the past few years around the globe to fulfill growing energy demands. Hydrocarbon potential of these unconventional petroleum systems depends on the presence of significant organic matter; their thermal maturity and the quality of present hydrocarbons i.e. gas or oil shale. In this work, we present a workflow for estimating Total Organic Content (TOC) from seismic reflection data. To achieve the objective of this study, we have chosen a classic potential candidate for exploration of unconventional reserves, the shale of the Sembar Formation, Lower Indus Basin, Pakistan. Our method includes the estimation of TOC from the well data using the Passey's ΔlogR and Schwarzkofp's methods. From seismic data, maps of Relative Acoustic Impedance (RAI) are extracted at maximum and minimum TOC zones within the Sembar Formation. A geostatistical trend with good correlation coefficient  $(R<sup>2</sup>)$  for cross-plots between TOC and RAI at well locations is used for estimation of seismic based TOC at the reservoir scale. Our results suggest a good calibration of TOC values from seismic at well locations. The estimated TOC values range from 1 to 4% showing that the shale of the Sembar Formation lies in the range of good to excellent unconventional oil/gas play within the context of TOC. This methodology of source rock evaluation provides a spatial distribution of TOC at the reservoir scale as compared to the conventional distribution generated from samples collected over sparse wells. The approach presented in this work has wider applications for source rock evaluation in other similar petroliferous basins worldwide.

**Keywords** Shale gas · TOC · Passey's ΔlogR method · Schwarzkopf's method · Relative acoustic impedance · Correlation coefficient

## **Introduction**

The ever-increasing demand of energy requires a deep insight into all possible ways of fulfilling it. One of the major components of the global energy mix comes from the exploitation of conventional resources (Bocora [2012](#page-16-0); Zou et al. [2016\)](#page-17-0). The increasing rate of the hydrocarbon production from conventional resources show continuous depleting trend for over 50 years (Miller and Sorrell [2014](#page-17-1)). Hence, there is a strong impetus to explore the hydrocarbon potential of unconventional resources i.e. oil/gas shales. Furthermore, the rise of unconventional resource exploitation

 $\boxtimes$  Aamir Ali aamirali.geo@gmail.com could present a new era in the overall global energy mix (Kefferpütz [2010;](#page-17-2) Sunjay [2011;](#page-17-3) Bocora [2012](#page-16-0); Rezaee [2015](#page-17-4)).

Conventionally, shale is a source rock for petroleum systems, showing low permeability (Gluyas and Swarbrick [2009;](#page-17-5) Liu et al. [2012;](#page-17-6) Rezaee [2015](#page-17-4)). Contemporary geophysical and geological techniques have been focused on extracting hydrocarbons from clastics and carbonates due to their favorable permeability. With advances in hydrocarbon exploration techniques in the last few decades, it is now possible to extract hydrocarbons directly from source rocks (Kefferpütz [2010](#page-17-2); Bocora [2012](#page-16-0); Rezaee [2015](#page-17-4)).

The Total Organic Content (TOC) in shale, its thermal maturity and good porosity are key parameters for the generation and accumulation of hydrocarbons (Rybach [1986](#page-17-7); Wood [1988](#page-17-8); Wandrey et al. [2004](#page-17-9); Gluyas and Swarbrick [2009;](#page-17-5) Rezaee [2015](#page-17-4)). Additionally, the vertical and lateral extent of source rocks can be identified and mapped by using seismic and well data (Løseth et al. [2011;](#page-17-10) Rezaee [2015\)](#page-17-4). In fact, TOC content can be precisely determined at boreholes

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using different techniques such as Passey et al. ([1990](#page-17-11)), Schwarzkopf [\(1992\)](#page-17-12) and Bowman [\(2010\)](#page-16-1) etc., and via the analysis of rock samples. This TOC distribution can be further mapped and extended with confidence away from control points on the reservoir scale using seismic data (Løseth et al. [2011](#page-17-10); Rezaee [2015](#page-17-4)).

The objective of the paper is to evaluate the hydrocarbon potential of Sembar Formation in the Badin area of the Lower Indus Basin, Pakistan (Fig. [1\)](#page-1-0), by mapping the TOC content using seismic and well log data within the context of unconventional resource production. Pakistan is endowed with 51 Trillion Cubic Feet (TCF) of recoverable gas shale reserves (Rezaee [2015](#page-17-4)). Shale of Sembar Formation is considered to be the most potential candidate in the region, with favorable characteristics, in terms of thickness, TOC values, and thermal maturity (Wandrey et al. [2004\)](#page-17-9). The workflow followed in the present study is given in Fig. [2.](#page-2-0) The data used in this study comprises of 3-D seismic data  $(13 \times 13 \text{ km}^2)$  and well log/vertical seismic profiling (VSP) data recorded in three wells (A, B and C) (Fig. [3\)](#page-3-0).

## **Geological setting**

The study area is located about 160 km East of Karachi in the Lower Indus Basin, between Latitude 24°5′N–25°25′N and longitude 68°21′E–69°20′E (Fig. [1;](#page-1-0) Kadri [1995](#page-17-13); Alam et al. [2002](#page-16-2); Ahmad et al. [2011;](#page-16-3) Munir et al. [2014](#page-17-14)). The evolution of the Lower Indus Basin began in the Cretaceous, when the tectonic activity occurred in the Indus Basin forming a rift zone (Farah et al. [1984](#page-16-4)). This rift zone exists near the Thar Platform in the form of Sargodha High dividing the Indus Basin into the Upper and the Lower Indus Basins (Kadri [1995](#page-17-13)). As a consequence of this rifting and doming, mainly horst and graben structures are present below the Paleogene unconformity including Cretaceous and older strata (Alam et al. [2002](#page-16-2)). The effect of rifting and doming on our study area (eastern side of Lower Indus Basin) is relatively small, as it lies away from the main deformational zone. The extensional deformational rate increases from east to west and a large area of the eastern Lower Indus Sub-Basin (Badin) is dominated by tilted normal faults (Farah et al. [1984;](#page-16-4) Kadri [1995](#page-17-13)). These faults can be seen and mapped on seismic data



<span id="page-1-0"></span>**Fig. 1** Map showing tectonics and regional geological setting of Pakistan including our study area (highlighted) situated in the Lower Indus Basin (Kazmi and Snee [1989\)](#page-17-15)

<span id="page-2-0"></span>

which have deformed the Cretaceous and older strata (Kadri [1995](#page-17-13)).

Stratigraphically, the Indus Basin ranges from Cambrian to Holocene in age with erosional and non-depositional intervals, while in study area it ranges from Jurassic to Paleogene as shown in Fig. [4](#page-4-0) (Alam et al. [2002\)](#page-16-2). The depositional environment of the Cretaceous Sembar and Goru Formations is shallow marine in nature (Iqbal and Shah [1980](#page-17-16)). These formations are deposited over the regional erosive surface of Chiltan limestone (Jurassic age) along passive continental margins (Iqbal and Shah [1980\)](#page-17-16). The Sembar Formation is a clastic rock with a large quantity of shale and a minor amount of siltstone and sandstone while Goru formation is a mixture of interbedded shale and sandstone (Fig. [4;](#page-4-0) Iqbal and Shah [1980;](#page-17-16) Raza et al. [1990\)](#page-17-17). The shelf to shallow-marine depositional environment persisted till Late Cretaceous time, represented by the carbonates of the Parh, Moghal Kot and Fort Munro Formations. The Pab Formation hosts regressive sand, thus representing a change to a nearshore environment (Wandrey et al. [2004](#page-17-9)).

The Sembar, Goru and Ranikot Formations are major potential conventional/unconventional sources of the Lower Indus Basin (Ahmad et al. [2011;](#page-16-3) Abbasi et al. [2014](#page-16-5)). The hydrocarbon potential of the Sembar Formation in the northern Kirther range is low but improves towards the south and southeast (Eckhoff and Alam [1991\)](#page-16-6). Lithologic composition of the Sembar Formation is variable in different areas (Ahmad et al. [2011](#page-16-3)). The average composition includes Quartz 42%, Clay 47%, Calcite 10% and Pyrite 1% (Ahmad et al. [2011\)](#page-16-3). The gas content is also variable in the basin with  $1\% < CO_2 < 70\%$ ,  $1\% < N < 80\%$  and  $0.1\% < H_2S < 13\%$ . The Sembar Formation extends towards the southern parts of Pakistan and into the offshore area (Eickhoff and Alam [1991](#page-16-6);

Shahzad et al. [2017\)](#page-17-18). In the western side of the Lower Indus Basin, units of the same age (Kzhdumi, Garau and Gadvan Formations) continue into Iran in the Zagros fold-belt and the Persian Gulf Basin (Opera et al. [2013\)](#page-17-19). These Formations possess potential source character in their respective basins within the context of unconventional resource exploitation (Opera et al. [2013](#page-17-19)). On the eastern side of the Lower Indus Basin, the Indian Rajasthan Basin comprises of lithological units, owing to the same depositional and tectonic history (Biswas [1987;](#page-16-7) Klett et al. [2011\)](#page-17-20).

## **Data and approach**

In this research, we have devised a methodology for TOC estimation in the Sembar Formation in Badin area, Lower Indus Basin, Pakistan using 3-D seismic data together with well log data of the three wells (A, B and C). The 3-D seismic cube is acquired and processed using bin size of  $25 \times 25$  m<sup>2</sup> and frequency range of 8–45 Hz, however this data can best resolve the target depth at frequency range of 10–35 Hz. The well data used, comprises of gamma ray (GR), spontaneous potential (SP), caliper (CALI), resistivity, density (RHOB), sonic (DT) and neutron porosity logs.

An indirect method is used for the estimation of TOC using well log data, which is an integral part for our proposed methodology. There exist many techniques for estimation of TOC from well log data such as Ferlt and Rieke ([1980](#page-17-21)), Schmoker ([1981](#page-17-22)), Meyer and Nederlof ([1984](#page-17-23)), Fertl and Chilinger ([1988\)](#page-17-24), Passey et al. ([1990](#page-17-11)), Myers and Jenkyns ([1992\)](#page-17-25), Schwarzkopf ([1992\)](#page-17-12) and Guidry and Walsh ([1993\)](#page-17-26). The contrast in geophysical response of a mature source rock, as compared to a non-source rock,



<span id="page-3-0"></span>**Fig. 3** Base map of the Badin area with in-lines (gray) and cross-lines (green) including three wells (Well A, Well B and Well C)

provides the basis for these indirect techniques (Meyer and Nederlof [1984](#page-17-23); Passey et al. [1990;](#page-17-11) Myers and Jenkyns [1992;](#page-17-25) Schwarzkopf [1992\)](#page-17-12).

A substantial amount of organic matter in a mature source rock is converted to gaseous or liquid hydrocarbon, displacing formation water from the pores, and later revealing anomalous behavior on well logs (Philippi [1968](#page-17-27); Nixon [1973;](#page-17-28) Meissner [1978](#page-17-29)). For example, resistivity, DT, neutron and GR logs show relative increase in the interval of mature source rocks, while density log response shows a decrease in mature source rock intervals, as described in Table [1](#page-4-1) (Passey et al. [1990](#page-17-11)). In this study we have followed the approaches presented by Passey et al. ([1990\)](#page-17-11) and Schwarzkopf ([1992\)](#page-17-12) for the estimation of TOC from well log data.

Reflection seismology (seismic methods) is one of the most important methods for the delineation of subsurface rock properties over larger extents despite its comparatively low resolution. On the other hand, well log/core data with relatively high resolution provides reliable information, but only for very small areal extents i.e. near the boreholes. The uncertainty in the rock property modeling increases as the distance from borehole increases. That is the point where seismic measurements in conjunction with rock physics modeling are very useful and its calibration with available well data can provide us a reliable model at the reservoir scale. The workflow followed in the current study is presented in Fig. [2](#page-2-0) and detailed below.

<span id="page-4-0"></span>**Fig. 4** Generalized stratigraphic chart of the study area (Alam et al. [2002;](#page-16-2) Abbasi et al. [2014,](#page-16-5) [2016](#page-16-8))

AGE (Ma)	<b>PERIODI EPOCH</b>		<b>FORM</b>	LITHOLOGY	<b>NOMANCLATURE</b>	<b>HC SIGNIFICANCE</b>												
						RES.	<b>SOURCE</b>	<b>SEAL</b>										
85			<b>PARH</b>		<b>CHALAK &amp;</b> <b>LIMESTONE</b>													
	CRETACEOUS	<b>UPPER</b>			<b>UPPER GORU</b>													
94			UPPER GORL	<b>LEAST</b> m. List over List over	<b>MARL &amp; SHALE</b>													
		LOWER	LOWER GORU	t de de la collage de la collage de la col	<b>UPPER SAND</b>													
				فانتز الأزمان الأراد أمارات الأراد أمار الأراد أما La de la Carlo de La Carlo 1940 - La Carlo de La Carlo 1940 - La La La Carlo de La	<b>MIDDLE SAND</b>													
				<u> 454 William</u>	<b>LOWER SHALE</b>													
134				自己民民自由的重要原因 <u>international de la</u> <u>Eugurbac Cuandore</u>	<b>BASAL SAND</b>													
															<b>SEMBAR</b>	<u>is de la propincia</u> <b>THEFT</b>	<b>SEMBAR</b> <b>SHALE &amp; SAND</b>	
164	<b>JURASSIC</b>	<b>UPPER</b>	CHILTAN		CHILTAN <b>LIMESTONE</b>													
فللما أأأأأأ لنملأ <b>SOURCE</b> <b>MARL &amp; SHALE</b> <b>SHALE</b> SANDSTONE <b>LIMESTONE</b> <b>SEAL</b> <b>RESERVOIR</b>																		

<span id="page-4-1"></span>**Table 1** Responses of five conventional logs in the organic matter of source rock, in which only density response is low and other four logs (GR, DT, neutron and resistivity) have high response (Fertl and Chilingar [1988;](#page-17-24) Passey et al. [1990\)](#page-17-11)



## **Demarcation of Sembar Formation**

Seismic reflection data is used to understand the regional distribution of the Sembar Formation in combination with well log data (Fig. [3](#page-3-0)). The top and base of the Sembar Formation are identified/marked on seismic data using the available well tops and is confirmed through synthetic seismograms (well-to-seismic tie) generated from well log and VSP data. Time and depth contour maps are prepared from the top and bottom of the Sembar Formation to identify the lateral and vertical trend in the study area. Horizon based Relative Acoustic Impedance (RAI) attribute for the Sembar Formation is applied to trace any acoustic impedance variations and to map TOC at reservoir scale.

## <span id="page-5-1"></span>**Passey overlay technique**

In the ΔLogR method, resistivity logs are superimposed to one of three porosity logs (neutron, DT and RHOB) (Passey et al. [1990;](#page-17-11) Liu et al. [2012\)](#page-17-6). Resistivity logs are plotted on logarithmic scale and other logs (neutron, RHOB and DT logs) are displayed on linear scales (Rider [2002](#page-17-30)). Any combination of two logs can be taken i.e. resistivity/DT, resistivity/neutron or resistivity/RHOB. In this method, the neutron and RHOB logs are mostly affected by borehole wall rugosity, and it is better to use a combination of resistivity/DT for more precise results (Liu et al. [2012](#page-17-6)). Resistivity and DT curves are relatively scaled so that 100 µs/ft of sonic data is equal to two logarithmic cycles of resistivity data (i.e. 100  $\mu s/ft = 2 \Omega m$  or 50  $\mu s/ft = 1 \Omega m$ ). Jianliang et al. ([2012](#page-17-31)) also used GR logs by keeping similar scale factors for sonic data (50 API = 1  $\Omega$ m).

The interval in which both curves are superimposed is known as a baseline. The baseline is mostly established in a fined grained non-source rock where both curves are on the top of each other with small or zero separation. The separation between the curves in the source rock is called ΔlogR which can be calculated mathematically as:

$$
\Delta LogR_{DT} = Log_{10} \left(\frac{R}{R_{Baseline}}\right) - 0.02 \times \left(\Delta t + \Delta t_{Baseline}\right).
$$
\n(1)

here  $\Delta$ Log $R_{DT}$  reflects the separation between DT and resistivity logs, R is the true resistivity from LLD or ILD resistivity logs,  $R_{\text{Baseline}}$  is the value of resistivity corresponding to sonic transit time at baseline,  $\Delta t$  is the sonic transit time and  $\Delta t_{\text{Baseline}}$  is the value of sonic transit time corresponding to resistivity at the baseline. The constant value 0.02 is the scaling factor. Using  $\Delta LogR_{DT}$ , TOC can be estimated as:

$$
TOC_{DT} = \Delta LogR_{DT} \times 10^{(2.297 - 0.1688 \times LOM)}.
$$
 (2)

The  $TOC<sub>DT</sub>$  is the TOC calculated using DT log, and LOM is the level of organic maturity. LOM values can be best obtained from core data. Whenever core data are not available, then conodont alteration indexes and Hood et al. [\(1975\)](#page-17-32) charts can be used for LOM calculation, provided that vitrinite reflectance data is available (Passey et al. [1990\)](#page-17-11).



<span id="page-5-0"></span>**Fig. 5** Superimposing synthetic seismogram over the seismic section of in-line 5821 along with GR log of well A. GR log is used here for differentiating/demarcating shaley intervals over seismic section. The

AI from well A is also displayed for calibration purpose. *F1*, *F2* and *F3* represent the faults having normal sense of shear

It is important to remember that TOC content and the maturity of the source rocks are directly proportional to the ΔlogR separation. A large separation indicates large amounts of organic matter and high maturity for source rocks (Passey et al. [1990\)](#page-17-11).

## <span id="page-6-1"></span>**Schwarzkopf's technique**

For application of Schwarzkopf's technique, we have followed the approach presented by Myers and Jenkyns [\(1992](#page-17-25)). This approach requires a few assumptions before the application of its mathematical relations. The assumptions are: (1) the source rock comprises mud rock as the matrix with density of 2.70  $g/cm<sup>3</sup>$ , water filled density of 1.05  $g/cm<sup>3</sup>$  and kerogen density of 1.1 or 1.2  $g/cm^3$ , (2) porosity and density of the source and non-source interval are same. The porosity for water filled pores ( $\phi_{\text{fl}}$ ) and kerogen ( $\phi_{\text{ker}}$ ) are estimated using the relations given below:

$$
\Phi_{\rm fl} = \frac{\rho_{\rm ns} - \rho_{\rm ma}}{\rho_{\rm fl} - \rho_{\rm ma}},\tag{3}
$$

$$
\phi_{\text{ker}} = \frac{\rho_{\text{s}} - \rho_{\text{ma}}}{\rho_{\text{ker}} - \rho_{\text{ma}}}.
$$
\n(4)

The  $\rho_{\text{ns}}$  is the density of non-source interval (average from log) in  $g/cm^3$ ,  $\rho_s$  is the density of source interval (from log) in g/cm<sup>3</sup>,  $\rho_{\text{ma}}$  is the assumed mud rock density,  $\rho_{\text{fl}}$  is the assumed water filled density and  $\rho_{\text{ker}}$  is the assumed kerogen density. Finally, TOC in wt% can be estimated using the following relation:

$$
TOC = \frac{0.85 \times \rho_{\text{ker}} - \phi_{\text{ker}}}{\rho_{\text{ker}} \left(\phi_{\text{ker}}\right) + \rho_{\text{ma}} \left(1 - \phi_{\text{fl}} - \phi_{\text{ker}}\right)}.
$$
(5)

#### <span id="page-6-2"></span>**Seismic based TOC estimation**

Since acoustic impedance depicts elastic behavior of a material, it therefore should provide information about variations in lithology, mineralogy, and fluid constituents etc. across



<span id="page-6-0"></span>**Fig. 6** Time contour map from the top of Sembar Formation with contour interval of 0.01 s. *F1*, *F2*, *F3* and *F4* represents faults traced over the base map. The color legend provides the information about shallowest and deepest parts of the Sembar Formation

the rock unit (Løseth et al. [2011](#page-17-10)). Moreover, the organic rich source rocks have low densities and acoustic impedances due to the high percentage of TOC compared to the non-source rock with a similar composition (Løseth et al. [2011\)](#page-17-10). To evaluate the source rock potential of the Sembar Formation in Badin area, the zero phase band limited 3-D seismic data of the area is converted from amplitude to RAI using the relation given by Nissen et al. ([2009\)](#page-17-33)

$$
\ln(\rho V) = 2 \int_{t=0}^{t=T} f(t)dt,
$$
\n(6)

where  $f(t)$  is the band limited reflectivity series,  $\rho$  is density and *V* is the compressional wave velocity within the rock layer. The curves of estimated TOC and acoustic impedance (AI) from wells are calibrated with the RAI section. Subsequently, horizon/reflection having low RAI and high TOC is identified throughout the seismic cube within the Sembar Formation with the help of calibrated AI and TOC curves derived from the well log data. After that, RAI is extracted along that interpreted horizons on a grid interval of  $25 \times 25$  m. The choice of this grid interval is based on the fact that the bin size used in the processing of the 3-D seismic volume is  $25 \times 25$  m and this grid interval is able to capture all the finer variations (in terms of RAI) within the data. Furthermore, RAI extracted in this way comes only from the rich TOC zone having low acoustic impedance within the Sembar Formation.

<span id="page-7-1"></span>Similarly, for the sake of comparison, a particular horizon/reflection within the Sembar Formation having high acoustic impedance and low TOC is identified throughout the seismic cube with the help of AI and TOC from well log data, and RAI is extracted and gridded  $(25 \times 25 \text{ m})$ along this horizon as discussed above.

In the next step, for the estimation of TOC from seismic data, a relationship describing the general trend between TOC (well derived property) and RAI (seismic derived property) is required. A relationship is established via a cross-plot between TOC and extracted RAI at the location of the wells for the high/low zones of TOC and RAI within



<span id="page-7-0"></span>**Fig. 7** Time contour map from the bottom of the Sembar Formation with contour interval of 0.05 s. *F1*, *F2*, *F3*, *F4* and *F5* represents faults traced over the base map. The color legend provides the information about shallowest and deepest parts of the Sembar Formation

the Sembar Formation, which is then used for the conversion of RAI at reservoir scale to TOC.

## **Results**

In this research, we have devised a methodology for TOC estimation in the Sembar Formation in Badin area, Lower Indus Basin using 3-D seismic data together with well log data. The results are presented in the relevant sub headings below.

#### **Spatial extent of the Sembar Formation**

A synthetic seismogram is generated using DT and RHOB logs of Well A, and is correlated with the 3-D seismic data of Badin area with a correlation coefficient of  $R^2 = 0.93$  (Fig. [5\)](#page-5-0). The top (1.6 s TWT) and bottom (1.94 s TWT) of the Sembar Formation is marked (Fig. [5\)](#page-5-0) and the time structural maps from the top/bottom of the Sembar Formation are prepared using contour intervals of 0.01 and 0.05 s respectively (Figs. [6,](#page-6-0) [7\)](#page-7-0). As the generated contour maps from both top/bottom of the Sembar Formation reflect the same deformational style (Figs. [6,](#page-6-0) [7](#page-7-0)), therefore, only top and bottom of the Sembar Formation is mapped. However, it's not a usual case to map the top/bottom of the target zone only, if the structural complexity exists within the target zone (Bacon et al. [2007](#page-16-9)). For delineation of exploitation depth, the time structural map generated from the top of the Sembar Formation is converted using velocity derived from seismic and well data (Fig. [8\)](#page-8-0). Furthermore, the time and depth contour maps show the spatial variation of Sembar Formation along with its deformational style. The thickness of Sembar Formation is interpreted to be  $\approx$  366–546 m based on the isopach grid map (Fig. [9](#page-9-0)), starting from the depth of around 2200 m with reference to the datum surface in the study area (Fig. [8](#page-8-0)). It is further inferred from the contour maps that the depth of Sembar Formation increases from east to west with normal faulting resulting in horst and



<span id="page-8-0"></span>**Fig. 8** Depth contour map from the top of Sembar Formation with contour interval of 15 m. *F1*, *F2*, *F3* and *F4* represent the faults traced over the base map. The color legend provides the information about shallowest and deepest parts with of the Sembar Formation

graben structural patterns. Due to the depth and thickness of the Sembar Formation, it has high temperature and overburden pressure, which makes it an ideal source rock within the context of shale gas exploration.

#### **Logs based TOC estimation**

The well log data of three wells (A, B and C) are used for TOC (wt%) estimation of Sembar Formation by using Passey ΔlogR and Schwarzkofp methods as discussed in Sections ["Passey overlay technique](#page-5-1) and [Schwarzkopf's](#page-6-1) [Technique"](#page-6-1) (Table [2](#page-9-1); Figs. [10](#page-10-0), [11](#page-10-1) and [12\)](#page-11-0). In the source rock interval, only the density log shows a decreasing trend, while the response of other logs such as GR, resistivity, DT and neutron shows an increasing trend, marking a rich TOC interval (Table [1;](#page-4-1) Passey et al. [1990](#page-17-11), [2010](#page-17-34)). A baseline is established in all the three wells based on small or no separation between DT and resistivity logs. There is a 250–300 m interval at the center of the Sembar Formation having TOC values in the range of 1–4 wt% (Figs. [10,](#page-10-0) [11](#page-10-1) and [12](#page-11-0)). Based on the estimated TOC within the Sembar Formation, the source rock interval can be designated as having TOC>1 and the non-source interval as having TOC<1.



<span id="page-9-0"></span>**Fig. 9** Isopach map of the Sembar Formation showing its thickness variation. *F1*, *F2*, *F3* and *F4* represent the faults traced over the base map. The color legend provides the information about lateral variation in thickness of the Sembar Formation

<span id="page-9-1"></span>

<b>Correlation</b> Depth		<b>Resistivity</b>	Crossover		Shale Volume W. Saturation	Porosity	Delta Log R	Passey	Schwarzkopf
$_{GR}$	MD	LLS	<b>NPHI</b>	Vshale	$S_{W}$	PHIA	<b>LLD</b>	TOC	<b>TOC</b>
(API) 150 Ю	(m)	$0.2$ (ohm.m) $2000 0.45$	$-0.15 0.3$	0.7 0		$ 0\rangle$	0.4 0.1 $ohm.m$	$415$ -1 (wt%) 5	$-2$ (wt%) 10
SP		<b>LLD</b>	<b>RHOB</b>			PHID			
$-100$ (mv) $150$		[0.2 (ohm.m) 2000]02 (gm/cm3)03		$\exists$ Shale $\equiv$	Water	0 0.4	Delta LogR	TOC (wt%) > 1	TOC (wt%) >1
CL			<b>CrossOver</b>	Sandstone <sup>®</sup>		PHIDT		$TOC (wt$ %) <1	
06 (Inch) 24		LLD/LLS					0.4 150 (us/ft)		$TOC (wt$ %) <1
	iembal								
	2400								
	2500								
	2600								
	2700						asel		
	2800	C.L.C.					≃		

<span id="page-10-0"></span>**Fig. 10** TOC values (wt%) estimated from Passey's ΔlogR and Schwarzkopf's methods in Well A (track 8 and 9). Track 1 is a lithological track, track 2 is a resistivity track and track 3 is a porosity track. The rest of the tracks represent the derived reservoir properties

(shale volume, water saturation, porosities, and Delta LogR) that are used to calculate TOC (wt%) curves from both methods (tracks 8 and 9)



<span id="page-10-1"></span>**Fig. 11** TOC values (wt%) estimated from Passey's ΔlogR and Schwarzkopf's methods in Well B (track 8 and 9). Track 1 is a lithological track, track 2 is a resistivity track and track 3 is a porosity track. The rest of the tracks represent the derived reservoir properties

(shale volume, water saturation, porosities, and Delta LogR) that are used to calculate TOC (wt%) curves from both methods (tracks 8 and 9)

Correlation	Depth		<b>Resistivity</b>	<b>Porosity Logs</b>	Shale Volume W. Saturation		Porosity	Delta Log R	Passey	Schwarzkopf
$_{GR}$	MD		<b>MSFL</b>	NPHI	Vshale	$S_{W}$	PHIA	<b>LLD</b>	TOC	TOC
(APD) 150	(m)	10.2	(ohm.m) 200000.45	$-0.15 0$		1 0 $\bf{0}$		0.3 0.1 (ohm.m)	$(wt\%)$ $500 - 1$	$(wt\%)$ $\mathbf{L}$
SP			LLS	<b>RHOB</b>		Water	PHID	DT	TOC (wt%)>1	TOC (wt%)>1
$-150$ (mv) 150		10.2		(ohm.m) $2000[02$ (gm/cm3) 03E				0.31150 (us/ft) $\bf{0}$		
CAL			LLD	DT			PHIDT			
(Inch) 06	24	10.2	(ohm.m) $2000(140$ (us/ft)	40 <sub>h</sub>	Sandstone		0.3	Delta LogR	TOC (wt%)<1	TOC (wt%)<1
	2500 2600 2700 2800 2900							B		

<span id="page-11-0"></span>**Fig. 12** TOC values (wt%) estimated from Passey's ΔlogR and Schwarzkopf's methods in Well C (track 8 and 9). Track 1 is a lithological track, track 2 is a resistivity track and track 3 is a porosity track. The rest of the tracks represent the derived reservoir properties

(shale volume, water saturation, porosities, and Delta LogR) that are used to calculate TOC (wt%) curves from both methods (tracks 8 and 9)



<span id="page-11-1"></span>**Fig. 13** Interpreted seismic section of in-line 5941 with a display of GR (green), AI (red) and TOC (black) curves at the location of well B and C for calibration purpose. The red horizon is picked at maximum TOC (low RAI) interval, while the blue horizon is picked minimum TOC (high RAI) interval within the Sembar Formation

#### **Analysis of TOC from seismic data**

To analyze TOC from seismic data, the amplitude information in the seismic cube is converted into RAI values using Eq. [\(6](#page-7-1)) and calibrated with the AI and TOC curves calculated from well log data (Fig. [13\)](#page-11-1). The identified low and high TOC zones with the Sembar Formation are gridded in terms of RAI as discussed in Section ["Seismic based TOC](#page-6-2) [estimation](#page-6-2)". From the RAI maps, one can clearly see the distribution of low and high acoustic impedances at well locations in Figs. [14](#page-12-0), [15](#page-13-0), which can be related to high and low TOC. The estimation of TOC from the RAI converted maps needs a relationship in the form of cross-plot between logs derived TOC and seismic derived RAI. The cross-plot of TOC and RAI gives inverse relation with a correlation coefficient  $R^2$ =0.72 (Fig. [16](#page-14-0)). In particular, the data in this cross-plot comes from both high/low TOC and RAI zones and the resulting linear relationship is used for the estimation of both high and low TOC zones within the Sembar Formation. In addition to that, we have used the estimated values of TOC from the Passey's method in cross-plot analysis.

The Passey's method is used, as the estimated TOC values involve the incorporation of laboratory/core data along with well log data making it more reliable than Schwarzkopf's method, which depends only on the well log data i.e. RHOB log. The Equation of the best fit line representing the general trend between TOC and RAI is given as

$$
TOC = -10^{-5} RAI + 1.7356.
$$
 (7)

<span id="page-12-1"></span>Finally, TOC maps are generated from RAI maps using Eq. [\(7](#page-12-1)) (Fig. [17](#page-15-0)). The estimated TOC values from seismic data show good calibration with well log data and suggest that Sembar Formation has suitable potential within the shale gas perspective (Table [2](#page-9-1); Fig. [17\)](#page-15-0). Moreover, we have also shown an estimation of TOC using RAI extracted at a low TOC and high acoustic impedance zone (Fig. [18\)](#page-16-10). By examining the behavior of RAI and TOC at well locations in Figs. [14](#page-12-0), [15,](#page-13-0) [17](#page-15-0) and [18,](#page-16-10) it can be observed that high and low TOC indeed manifests itself in terms of RAI anomalies. In particular, high TOC at well locations in Fig. [17](#page-15-0) show low RAI in Fig. [14](#page-12-0) and low TOC at well locations in Fig. [18](#page-16-10) show high RAI in Fig. [15.](#page-13-0)



<span id="page-12-0"></span>**Fig. 14** RAI extracted at a 25×25 m grid along the maximum TOC zone of Sembar Formation. The map shows low RAI values at the location of wells A, B and C



<span id="page-13-0"></span>**Fig. 15** RAI extracted at a 25×25 m grid along the minimum TOC zone of Sembar Formation. The map shows high RAI values at the location of wells A, B and C

## **Discussions**

Overall, the estimated TOC values for the Sembar Formation in this study ranges from 1 to 4 wt% showing that it lies in the range of good to excellent source rock (Peters et al. [1986\)](#page-17-35). The values of TOC of the Sembar Formation analyzed by Wandrey et al. ([2004](#page-17-9)) from laboratory samples range from 0.5 to 3.5 wt%. So the result of the TOC obtained in this work shows a good agreement with study conducted by Wandrey et al. ([2004\)](#page-17-9). It is further inferred that Kerogen of the Sembar Formation is mostly of Type-II and III in the Lower and Middle Indus Basin (Wandrey et al. [2004](#page-17-9); Ahmad et al. [2011](#page-16-3)). The TOC values calculated from seismic data for the Sembar Formation increases from eastern to western side (Fig. [17](#page-15-0)). Thermally, the rock is much more mature to the western part of the shelf due to deep burial depth and less mature to the eastern part due to shallow burial depth (Fig. [8;](#page-8-0) Wandrey et al. [2004\)](#page-17-9).

Our analysis shows that the Sembar Formation can act as a good potential unconventional resource because it is much mature towards higher concentration of TOC

(Fig. [17\)](#page-15-0). The reason for the high concentration of TOC is the paleo-geographic location of the Sembar Formation during its depositional time. The age and depositional environment as discussed in Geologic settings suggest that during the late Jurrasic and Lower Cretaceous, the Indian plate drifted northward, entering to warmer latitude (Scotese et al. [1988\)](#page-17-36). On its lower shelf, marine shales of the Lower Cretaceous Sembar Formation were deposited over Jurrasic age Chiltan Limestone. During this age, the paleogeographic latitude of the Indian plate was at location of nearly Latitude 20°S and Longitude 68°E (Scotese et al. [1988\)](#page-17-36). The higher concentration of TOC suggests that the existed paleo-environment at that time was suitable for high productivity rate of organic matter. Furthermore, it can also be further inferred that, it was the time when the decomposition rate was slow (means an anoxic environment suitable for preservation of the organic matter existed) as compared to the productivity rate of the organic matter resulting in such high concentration of TOC. The rich TOC potential of the Sembar Formation makes it a



<span id="page-14-0"></span>**Fig. 16** A linear relationship established via a cross-plot between TOC (estimated from well data) and extracted RAI (from seismic) at the location of all the three wells. In particular, both high/low TOC

and RAI zones within the Sembar Formation are used. The linear relationship established in the cross-plot is then used for the estimation of TOC from seismic based extracted RAI values

potential source rock and as well as a potential unconventional play.

For further evaluation of the unconventional potential of a particular unit, the depth of exploitation and its thickness is an essential target. In platform areas, such as the Lower and Middle Indus Basins, the general economic exploitation depths for shale gas are about 3500 m, while that in tectonically unstable fold-belts, it ranges from 1000 to 3000 m (Ahmad et al. [2011\)](#page-16-3). The exploitation depth of Sembar Formation obtained for our study starts from around 2200 m (Fig. [8](#page-8-0)) having considerable thickness (Fig. [9\)](#page-9-0), making it a feasible and suitable candidate for unconventional resource exploitation with higher concentration of TOC (Fig. [17](#page-15-0)).

The analogues of Sembar Formation are present in the Rajasthan Basin (eastern side of the study area) of the Indian subcontinent and in the Zagros fold-belt of Iran (western side of the study area) (Klett et al. [2011](#page-17-20); Opera et al. [2013](#page-17-19)). These analogues units are the most productive source rocks of their respective basins within the context of shale gas exploitation (Klett et al. [2011](#page-17-20); Opera et al. [2013](#page-17-19)). The values of the TOC in the Gurpi and Kazdumi Formation of Zagros fold-belt of Iran ranges from 1.7 to 2.4 wt% and 1.5 to 3.92 wt% respectively (Opera et al. [2013](#page-17-19)). The source potential of the Ghaggar–Hakra Formation of Rajasthan Basin is poorly known due to limited number of well penetrations (Farrimond et al. [2015](#page-16-11)). We expect that this methodology may be helpful in exploration of unconventional resources for such basins worldwide.

# **Conclusions**

The Sembar Formation is an ideal shale gas prospect of the Lower Indus Basin in terms of depositional environment, thickness, organic geochemistry, thermal maturity, mineralogy and porosity. One of the key features for successful shale gas plays requires high TOC. In this study, we have devised a methodology for estimation of TOC from seismic and well log data for the Sembar Formation in Badin area, Lower Indus Basin, Pakistan. The extracted maps of RAI and estimated TOC (1–4%) within the Sembar Formation are evaluated on the reservoir scale with a good correlation coefficient. The estimated TOC values for Sembar Formation show that it lies in the range of good to excellent potential shale oil/gas reservoir. Our results suggest that high and low TOC values indeed manifest themselves in terms of RAI anomalies. Moreover, the results also suggest an excellent calibration of TOC values from seismic at well locations. We anticipate that this methodology can be helpful in the initial assessment and identification of optimum sweet spots (high TOC) for shale gas plays worldwide.



<span id="page-15-0"></span>**Fig. 17** Seismic based TOC using Passey's method for RAI extracted along high TOC and low acoustic impedance zone within the Sembar Formation. The map shows high TOC values at the locations of wells A, B and C



<span id="page-16-10"></span>**Fig. 18** Seismic based TOC using Passey's method for RAI extracted along low TOC and high acoustic impedance zone within the Sembar Formation. The map shows low TOC values at the locations of wells A, B and C

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