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# An assessment of radionuclides level, radon and thoron exhalation rate in hill and field soil of Mahendergarh district in Haryana, India

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#### Abstract

To investigate the activity concentrations of naturally occurring radionuclides such as <sup>238</sup>U, <sup>232</sup>Th, and <sup>40</sup>K, as well as the presence of radon (<sup>222</sup>Rn) and thoron (<sup>220</sup>Rn) in the vicinity of the Aravalli Mountain range in Mahendergarh, India, a comprehensive study was conducted. We meticulously examined soil samples obtained from both field and hill areas using NaI (Tl) detector based on gamma spectroscopy. It is noteworthy that concentrations were found lower than the global average values. Notably, the hill soil samples exhibited a higher activity concentration in comparison to the field soil samples. Overall, in terms of radium equivalent activity (<sup>226</sup>Ra), gamma absorbed dose rate, and the internal hazard index, our findings did not reveal any significant radiological risks.

Keywords Naturally occurring radionuclides  $\cdot$  Smart RnDuo radon monitor  $\cdot$  Hill soil  $\cdot$  Exhalation rates  $\cdot$  Aravalli mountain range

# Introduction

Soil is the principal reservoir for all essential life supporting components either directly or indirectly. It contributes significantly to the natural background radiation and these radiations exposed to the surroundings [1, 2]. Soil constitutes the uppermost layer of the earth crust, formed through a series of physiochemical changes including decomposition, water movement, and weathering of solid rock. Within the earth crust, rocks and minerals naturally emit low levels of radiation due to the presence of radioactive isotopes. Soil comprises minerals and rocks that naturally erode, releasing radioactive elements, particularly uranium (<sup>238</sup>U), thorium (<sup>232</sup>Th), and potassium (<sup>40</sup>K), along with their decay

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products, as inherent components of soil. Radiation is an integral aspect of our environment, and human exposure and radiation occurs through routine interactions, such as exposure to sunlight and natural background radiation [3].

Background radiation encompasses cosmic radiation that constantly permeates the atmosphere from space. The average annual natural background radiation exposure for humans is approximately 1.1 mSv, sourced from cosmic radiation (0.35 mSv) and from atmospheric sources its value is 0.05 mSv [4]. The distribution of radionuclides in soil and their radiological impacts significantly influence human health. These radionuclides account for at least 80% of natural radiation exposure [5, 6], with the remaining 20% stemming from human activities. Elevated levels of anthropogenic radiation, originating from <sup>238</sup>U and its decay products in geological materials as well as <sup>232</sup>Th, predominantly found in zircons, igneous rocks, and monazite sands are significant contributors to high background radiation levels. In some regions, the presence of monazite sands has resulted in exceptionally high background radiation levels, increasing radiation exposure in various countries [7–15]. Pegmatite, granite, diorite, and gneiss rock samples of North Pakistan were highly radioactive and should not be used as constructing material [16]. Globally, naturally the high background radiation place is Ramsar City, Iran. In Iran, high background radiation is due to the presence of <sup>226</sup>Ra in local rocks [17]. Mrima Hill of Kenya country, known for high background radiation due to the presence of heavy minerals like carbonatites, and monazites. Here, activity concentration of naturally occurring radioactive elements and dose rate were found above the global value [18]. Also, in India, Kerala has a high level of radiation, and the attribution of radiation is due to monazite sand containing enriched thorium [8]. Prior research has shown that higher levels of radon and thoron in the environment significantly increase the risk of lung cancer even in non smokers [19-24]. Radon (<sup>222</sup>Rn) and thoron (<sup>220</sup>Rn) are released from soil and construction materials into the environment through emanation and exhalation. The exhalation rate is influenced by various factors, including the content of <sup>226</sup>Ra in the soil, rock composition, porosity, permeability, temperature, humidity, and meteorological conditions [25-28]. Thus, this study project also included the measurement of the exhalation rate for soil samples to assess potential health risks.

Although radionuclide levels in soil have been measured in various regions of Haryana over the past few decades [28–32]. No prior investigations have been reported on these radioactive elements in the soils of the Mahendergarh district in Haryana. Given the common association of <sup>238</sup>U and <sup>232</sup>Th with the Aravalli Hills, radiation exposure from this region could be an environmental concern. Consequently, it is essential to conduct a qualitative analysis of these radionuclides in this specific research area. Buildings in India commonly utilize bricks that incorporate approximately 80% soil [20]. Therefore, this study aims to determine whether the soil in this area is suitable for construction without posing risks to human health.

Radionuclide <sup>226</sup>Ra is known to migrate more readily in the environment and its decay product, radon gas escapes from the soil. While various natural radionuclides such as those in the <sup>235</sup>U series, <sup>176</sup>Lu, <sup>87</sup>Rb, and <sup>147</sup>Sm exist in the environment at low levels and their contributions to human radiation exposure are relatively low [33]. As a result, this study focuses on the assessment of radionuclides <sup>238</sup>U, <sup>232</sup>Th, and primordial radionuclide <sup>40</sup>K in the soil, utilizing a gamma ray spectrometer to measure element activity concentrations and calculate the exhalation rates of <sup>222</sup>Rn and <sup>220</sup>Rn for the soil samples by employing the SMART RnDuo portable radon monitor.

# Geological characteristics of the study area

The research was conducted in the vicinity of the Aravalli Mountain range located in Mahendergarh, Haryana, India. This district spans between  $24^{\circ}$  47' to  $28^{\circ}$  26' N latitudes and 75° 56' to 76° 51' E longitudes, covering an area of 1899 km<sup>2</sup>. The region primarily falls within the Indo Gangetic plains geomorphological zone. The predominant soil types in this district include arid soil, blown sand, and alluvium. These soils typically contain subsurface lime nodules and exhibit calcareous characteristics. The geological substrata of the district consist of rocks belonging to the Delhi and Delhi systems, overlain by recent alluvial deposits and blown sand.

The area surrounding the district features prominent hill formations, notably the Madhogarh Hill, Dhosi Hill, and the Tosham Hill range. These hills are situated within the Aravalli Mountain range and primarily consist of metasedimentary rocks [34]. These metasedimentary rocks predominantly comprise quartzite and contain a relatively low number of pegmatites, slate granite, and phyllite. The Tosham Hill range, an essential component of the Aravalli Mountain range, falls under the Archean Bilwala basement rock category and primarily comprises quartz and granite porphyries known for their high thermal conductivity.

Climatically, the study area experiences an annual average rainfall of approximately 500 mm, with uneven distribution across the region. Moreover, the district is situated near the Dohan River, which is currently facing the threat of extinction. The Krishnavati River originates from the Aravalli range, near the Dariba copper mine in the state of Rajasthan.

# **Experimental procedure**

### Sample preparation

In this research, a total of 28 soil samples were initially chosen randomly from various surface areas, including rock formations within the Aravalli range, for the determination of radionuclides. Subsequently, 17 of these soil samples were selected for the measurement of radon and thoron exhalation rates in Mahendergarh, Haryana, India, as illustrated in Fig. 1. Each sample was obtained at a depth of 45 cm and was accurately positioned using GPS coordinates.

The collected samples, each weighing approximately 1 kg, underwent a series of processing steps, including grinding, sieving, and homogenization, resulting in a particle size of 100 mesh, achieved through the use of a crushing machine. Following this, the prepared samples were dried at 110 °C for 12 h to ensure the complete removal of moisture. After the samples were weighed, they were placed into sun pet jars, hermetically sealed, and left undisturbed for over one month. This critical step allowed for the establishment of secular equilibrium, as described in previous studies [35, 36]. By doing so, it ensured that radon gas was contained within the samples, and its decay products remained within the samples for subsequent measurements and analysis.



Fig. 1 Surveyed area map for soil samples in Mahendergarh, district, Haryana India

rnual Singhana

#### Instrumentation and calibration

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For the determination of the concentration of natural radionuclides, a  $\gamma$ -ray spectrometer employing NaI (Tl) scintillation detector with dimensions of  $2'' \times 2''$  was utilized. The detector (MODEL: NETS – ØM) was supplied by electronic enterprises (I) PVT. Ltd PARA Electronics-Mfg. Division Mulund Mumbai. The detector boasts a resolution of (FWHM) 1.85 keV for the 1.33 MeV gamma line of <sup>60</sup>Co [37]. Energy calibration was performed using point sources of <sup>60</sup>Co and <sup>137</sup>Cs.

#### Measurement of radioactivity concentration

The measurement of radioactivity concentration involved the use of  $\gamma$ - rays emitted by specific radioactive isotopes for analysis. Notably, the  $\gamma$ -rays of interest included 186.2 keV for <sup>238</sup>U, 911 keV, 968 keV for <sup>232</sup>Th, and 1460.8 keV for <sup>40</sup>K [38]. The counting period for each sample was set at 80,000 s to ensure robust statistical data. Subsequent analysis of the obtained counts facilitated the calculation of the activity concentration of radioactive elements, specifically <sup>238</sup>U, <sup>232</sup>Th, and <sup>40</sup>K reported in Bq/kg.

76°30.00'E

Haryana Mahendragarh Soil Sampling Locations

Radionuclides Concentration

along with Radon and Thoron Exhalation Rates Radon and Thoron

Exhalation Rates

#### **Theoretical calculations**

Given the non-uniform distribution of natural radionuclides ( $^{238}$ U,  $^{232}$ Th, and  $^{40}$ K) within the soil, an assessment of radiological risks associated with soil usage was performed using a single index incorporating the activity of various radionuclides. The activity concentrations of uranium, potassium, and thorium were calculated employing the following Eq. (1) [39, 40].

Activity concentration(Bq kg<sup>-1</sup>)  
= 
$$\frac{\text{Net count rate}}{\text{efficiency } \times \text{ sample weight } \times \text{ abundance}}$$
 (1)

Radium (Ra<sub>eq</sub>) equivalent activity is used to compute the total radiation exposure caused by radionuclides ( $^{238}$ U,  $^{232}$ Th, and  $^{40}$ K). The calculation is performed by using the following Eq. (2) [38, 41, 42].

$$Ra_{eq}(Bq kg^{-1}) = A_{U}(Bq kg^{-1}) + 1.43 A_{Th}(Bq kg^{-1}) + 0.077 A_{K}(Bq kg^{-1})$$
(2)

where  $A_U$ ,  $A_{Th}$ , and  $A_K$  are the concentrations of <sup>238</sup>U, <sup>232</sup>Th, and <sup>40</sup>K in Bq kg<sup>-1</sup> respectively [7].

#### Hazard index

Internal hazard index  $(H_{in})$  measures the effect of radionuclides on the lungs and other organs. Its value must be less than unity. Using the following Eq. (3), one can determine the risks due to naturally occurring radionuclides [42].

$$H_{\rm in} = \frac{A_{\rm U}}{185} + \frac{A_{\rm Th}}{259} + \frac{A_{\rm K}}{4810} \tag{3}$$

The constant terms are used, it is assumed that the radiation doses for  $^{238}$ U,  $^{232}$ Th, and  $^{40}$ K were 185 Bq kg<sup>-1</sup>, 259 Bq kg<sup>-1</sup>, and 4810 Bq kg<sup>-1</sup> to provide equal gamma radiation dose [43, 44].

#### Absorbed dose rate (AAD)

It can be calculated by using the following Eq. (4) [7] using activity concentrations of  $^{238}$ U,  $^{232}$ Th, and  $^{40}$ K (UNSCEAR 2000).

$$\begin{aligned} AAD(nGyh^{-1}) &= 0.462 A_U(Bq kg^{-1}) \\ &+ 0.604 A_{Th}(Bq kg^{-1}) + 0.0417 A_K(Bq kg^{-1}) \end{aligned} \tag{4}$$

where  $A_U$ ,  $A_{Th}$ , and  $A_K$  are the radioactivity concentrations of natural radionuclides in soil samples [45].

#### Radon exhalation rate measurement

To determine the exhalation rate of <sup>222</sup>Rn in soil samples, a Portable Radon Monitor developed by BARC (Bhabha Atomic Research Centre) was employed, as depicted in

**Fig. 2** Experimental set up used for the assessment of <sup>222</sup>Th exhalation rate in soil samples

Fig. 2. This monitor operates by detecting alpha particles produced by <sup>222</sup>Rn and its progenies, namely <sup>218</sup>Pb and <sup>214</sup>Po.

The procedure involved placing the soil sample within a stainless steel cylindrical container with a known weight (M). The container measured 8.2 cm in height and had an inner diameter of 10 cm. A progeny filter was utilized to selectively collect radon while effectively eliminating the <sup>222</sup>Rn descendants. In addition, a pinhole plate was employed to suppress thoron, which is relatively short lived. This step was essential to account for the diffusion time delay, a phenomenon in which the transmission of <sup>220</sup>Rn takes longer compared to radon transmission due to the shorter lifetime of thoron. Each measurement was carried out for 9 h to ensure a comprehensive assessment of the accuracy and precision of the experimental setup, a quality control measure verified by the creator of the SMART RnDuo [46].

The radon mass exhalation  $(J_m)$  rate is calculated by analysing the radon concentration C(t), and applying Eq. (5) [46–50].

$$C(t) = \frac{J_{\rm m}M}{V\lambda_{\rm e}} \left(1 - e^{-\lambda_{\rm e}t}\right) + C_0 e^{-\lambda_{\rm e}t}$$
<sup>(5)</sup>

where  $J_{\rm m}$  is the radon mass exhalation rate in mBq/kg/h,  $V({\rm m}^3)$  is the sum of the effective chamber volume and the volume of the scintillation cell,  $M({\rm kg})$  is the sample mass,  $\lambda_{\rm e}(h^{-1})$  is the effective decay constant due to decomposition of <sup>222</sup>Rn, back diffusion, leakage rate of chamber,  $C_0$  is the initial radon concentration in the chamber at time t=0.

#### Measurement of thoron surface exhalation rate

To assess the concentration of <sup>220</sup>Th, a flow mode sampler was connected to the inlet of the monitor pump as shown in Fig. 3. This sampler was exclusively used for the quantification of thoron. The quantification process involved a 15 min cycle during which measurements of thoron levels



and background counts were recorded. Following this cycle, a 5 min hold up period was observed to ensure that the <sup>220</sup>Th had nearly completely decayed. Subsequently, a final 5 min count was conducted to determine the number of background counts associated with that specific cycle.

The surface rate  $(J_{st})$  of thoron in Bq/m<sup>2</sup>/s was calculated using Eq. (6), as previously employed and validated in related studies [51–53].

$$J_{\rm st} = C_t \frac{V\lambda}{A} \tag{6}$$

where  $C_t$  is the build up <sup>220</sup>Rn concentration (Bq/m<sup>3</sup>) within the chamber as determined by a portable monitor throughout 15 min cycle. V (m<sup>3</sup>) is the leftover air volume enclosed by the loop. A (m<sup>2</sup>) is the sample surface area.  $\lambda$  is the decay constant of <sup>220</sup>Rn (0.012464 s<sup>-1</sup>).

# **Results and discussion**

# Naturally occurring radionuclides

Using NaI (Tl) detector, the radioactivity concentration in the study area was determined. The activity concentration of  $^{238}$ U,  $^{232}$ Th, and  $^{40}$ K varied within the range of (0.06–1.81) Bq/kg, (0.09–2.37) Bq/kg, and (3.09–10.9) Bq/kg with mean values of 0.84 Bq/kg, 1.16 Bq/kg, and 7.08 Bq/kg, respectively. It is noteworthy that these activity concentrations of  $^{238}$ U,  $^{232}$ Th,  $^{40}$ K, and  $^{226}$ Ra<sub>eq</sub> radionuclides are below the permissible limits of the world average values of 32 Bq/kg, 30 Bq/kg, 400 Bq/kg, and 370 Bq/kg, as outlined by UNSCEAR in 2000.

The coefficient of variability was highest for  $^{238}$ U (49%) and lowest for  $^{40}$ K (23%). Notably, the highest concentrations of uranium, thorium, and potassium were found in Jhhankhadi village, Narnaul Singhana bottom hill, and Narnaul Dhosi hill, while the maximum radium equivalent activity concentration was observed in Narnaul Singhana bottom hill.

Analysis of Fig. 4 reveals that the activity concentration of  $^{40}$ K in all soil samples was greater than that of  $^{238}$ U and  $^{232}$ Th, which is consistent with soil expectations. This variation may be attributed to geological disparities, the application of chemical fertilizers for agricultural purposes, and the presence of the Aravalli Hills, all contributing to the increased radioactivity in the study area.

Furthermore, as illustrated in Fig. 4 the activity concentration of <sup>232</sup>Th surpasses that of <sup>238</sup>U in all soil samples, except for seven locations. This finding suggests the prevalence of thorium rich soil in the study area, corroborating the



Fig. 4 Variations of radioactive nuclides activity content in soil samples



#### Tubes for connections



notion that thorium is more abundant in nature than uranium, in alignment with the World Nuclear Association Report in 2020 on thorium. The contribution of radionuclides to the absorbed dose rate in air depends on their concentration in the soil, with absorbed dose rates varying from 0.46

Table 1 Radon mass exhalation rate at survey site areas

Locations	Radon exhalation rate (mBq/kg/h)	Thoron surface exhalation rate (Bq/ m <sup>2</sup> /h)
Lawan, field soil	1.96	19.2
Majra Kalan, field soil	2.68	13.8
Khayra, field soil	7.05	17.4
Khatod, field soil	61.4	10.2
Madhogarh, mid hill soil	240	34.8
Madhogarh, top hill soil	90	20.4
Barda, field soil	4.38	6.6
Pali, field soil	2.99	31.8
Narnaul Singhana hill soil	119	31.2
Narnaul, Dhosi hill soil	224	25.8
Kaliana, mid hill soil	224	29.4
Kaliana, top hill soil	141	11.4
Mandola, field soil	5.4	11.4
Dholi, field soil	16.4	4.8
Digrota, field soil	3.01	5.4
Nangal Mala, field soil	16.9	9.0
Balana, field soil	14.7	16.2
Mean	51.9	17.4

**Fig. 5** Radon mass and Thoron surface exhalation rate in soil samples of the area under study

to 2.28 nGy/h, averaging 1.38 nGy/h. In comparison to the world average absorbed dose rate of 86 nGy/h (UNSCEAR, 2000), the calculated values in the soil samples are significantly lower, underscoring the region safety in terms of radiation exposure. Additionally, the calculated average value of the internal hazard index ( $H_{\rm in}$ ) was determined to be 0.01, indicating values lower than the safe threshold.

#### **Radon and Thoron exhalation rate**

The mean values of  $^{222}$ Rn mass and  $^{220}$ Rn surface exhalation rate were found to be 51.9 and 17.4 Bq/m<sup>2</sup>/h, respectively as reported in Table 1. The mean value of radon mass exhalation rate is approximately 9% lower than the world average value of 57 mBq/kg/h, while the mean value of thoron surface exhalation rate is nearly 99.5% lower than the world average value of 3600 Bq/m<sup>2</sup>/h, as reported by UNSCEAR in 2000.

In the bar graph presented in Fig. 5 and Table 1, it is evident that the maximum thoron exhalation rate was recorded for Madhogarh Mid Hill (location no. 5), measuring 34.8 Bq/m<sup>2</sup>/h, while the maximum mass exhalation rate of 240 mBq/kg/h was also found in Madhogarh Mid Hill. The figure illustrates high peaks representing rock samples and lower peaks representing field area samples. Specifically, the samples with sample numbers 5, 6, 9, 10, 11, and 12 correspond to hill soil samples, exhibiting higher radon mass and thoron surface exhalation rates compared to the rest of the samples, which are field soil samples.





Fig. 6 Correlation between  $^{222}\text{Rn}$  mass exhalation rate and  $^{238}\text{U}$  activity concentration

# Correlation between radionuclides present in soil samples

Figure 6 illustrates the correlation between  $^{238}$ U and  $^{222}$ Rn, which is weak but positive, with an  $R^2$  value of 0.0044. As demonstrated in Fig. 7, thorium exhibits a weak yet positive correlation (0.0299) with thoron. It is important to note that no strong statistical relationship exists between these radionuclides, indicating that the radioactive content in soil samples is influenced by the diverse nature of these radionuclides. Therefore, the distribution of one radionuclide in

the soil does not depend on the concentration of another radioactive element.

To further contextualize the results, the study area was compared with hilly areas in India, particularly those surrounded by the Aravalli hills, as shown in Table 2. The findings revealed that the radon exhalation rate in the study area was similar to that of Granitic hills in Karnataka [54]. However, the values for radon exhalation rate were lower in comparison to Kamaun Hills in Uttarakhand [55], submountainous regions in Jammu & Kashmir [56], Shivalik Hills in Himalaya [57], the Himalaya foothill region in Uttarakhand [7] and Shivalik Hills in Haryana and Himachal Pradesh [58].

# Differentiation of natural radioactivity in rock soil samples

The natural hills in the study area, such as Madhogarh, Dhosi, Kaliana, and Singhana Hills are formed through various geological processes. These hills can also result from erosion, where rocks, soil, and sediments accumulate in a particular area. The presence of natural radionuclides and the exhalation of <sup>222</sup>Rn and <sup>220</sup>Rn were found to be greater in hill soil samples as compared to field samples. The statistical data plotted for gamma absorbed dose rate values and hazard index is presented in Fig. 8. Here, the maximum hazard



**Fig. 7** Correlation between <sup>220</sup>Th surface exhalation rate and <sup>232</sup>Th activity concentration

**Table 2** Comparative study ofsurveyed area with India hillyareas

Hilly areas	Radon exhalation rate (mBq/ kg/h)		References
	Min	Max	
Kamaun hills region, Uttarakhand	16	54	Semwal et al. [55]
Sub mountainous region, J & K	$15 \pm 0.4$	$38 \pm 0.8$	Kaur et al. [56]
Siwalik hills Himalaya, Jammu & Kashmir	$7 \pm 0.6$	$48 \pm 1.3$	Kaur et al. [57]
Himalaya foothills region, Uttarakhand	16	111	Anamika et al. [7]
Granitic hills region, Karnataka	$76\pm 6$	$269 \pm 19$	Poojitha et al. [54]
Shivalik hills, Haryana &Himachal Pradesh India	$50 \pm 1$	$143 \pm 6$	Chauhan et al. [58]
Aravalli hills, Mahendergarh Haryana	1.96	240	Present study

index and absorbed dose rate were found for Narnaul Singhana Bottom Hill. The distribution of radionuclides in hill soil samples is depicted in Fig. 9, where it can be observed that the activity concentration of <sup>238</sup>U was highest for Dhosi Hill, <sup>232</sup>Th was most prominent in Narnaul Singhana Hill, <sup>40</sup>K concentration was at its peak in Dhosi Hill, and <sup>220</sup>Rn and <sup>222</sup>Rn were elevated in Madhogarh Mid Hill soil samples. Narnaul Dhosi Hill predominantly consists of quartzite, with some pegmatite, slate, granite gneiss, phyllite, schist, and various basic rocks. The sharp contact between pegmatite and quartzite, along with the blurred contact between granite and pegmatite at Narnaul, serve as strong evidence of geological movement within the Aravalli orogenic belt [59]. The high concentration of heat producing elements in granite indicates the presence of radioactive elements like <sup>238</sup>U, <sup>232</sup>Th, and <sup>40</sup>K. Madhogarh Hill, an isolated hill within the Aravalli range, is believed to release significant amounts of radon and thoron gases due to the presence of basic rocks.

Rocks found in the Kaliana Hills predominantly consist of mica, quartz grains, sedimentary quartzite, and flexible sandstone known as Itacolumite. While both flexible and non flexible sandstone can be found in the Kaliana area, non flexible sandstone is more abundant relative to flexible sandstone. Cementing materials such as floor tiles are produced







Fig. 9 Distribution of naturally occurring radionuclides material  $(^{238}U, ^{232}Th, ^{40}K)$  and exhalation rate in different rock samples

using these sandstones. Although the concentration of radionuclides was higher in Kaliana Hill as compared to field soil, it remained within safe limits. This suggests that Kaliana soil samples can be utilized for building materials [60]. Narnaul Singhana Hill is part of parametamorphites belonging to the Delhi Subdivision, and it is characterized by schists, amphibole quartzite, and mineralized shear zones. These features may account for the high concentration of thorium in the area. Additionally, inhomogeneity in the distribution of radioactive elements in different geological layers of the hill was observed, with some layers being more radioactive than others. This variation is attributed to temperature differences at different elevations of the hill, with the top elevation having lower temperatures than the midsection. The distribution of radioactive elements is also influenced by the type of rock in a given layer, emphasizing the need for future researchers to conduct radiometric measurements at specific hill locations.

Certain types of hills, such as those with sedimentary rocks like Kaliana Hill, exhibit lower radioactivity concentrations in comparison to hills containing granites and some metamorphic rocks. As demonstrated in Fig. 9, Kaliana Hill did not have the highest radionuclide concentrations, as it primarily consists of sedimentary rocks. According to the literature survey, sedimentary rocks consist of low radioactivity as compared to igneous rocks [61]. Deposition, soil erosion, and topographical variations affect the morphological, chemical, and physical characteristics of the soil [62]. In hilly areas, the radioactivity is high which is interrelated with dental fluorosis [63]. Mahendergarh area is also a fluorosis endemic red zone alert area, here fluoride distribution in groundwater is due to fluoride bearing rocks. Here dental fluorosis was diagnosed, and found high level of fluoride [64]. The present area is an industries free, pollution free area, and inhabitants of this area are dependent on agriculture. So, industrial aerosols are not responsible for high radioactivity in this area. This is the reason that there is no radioactivity was found in the field soil samples. Also, one can say that the geology of this area, and radiation bearing rocks may be responsible for high activity concentration in hills soil samples.

# Conclusions

In conclusion, the activity concentration values in the study area are influenced by geographical conditions, soil composition, and the presence of the Aravalli hills. These values were found to be lower than the world average, indicating that the radiation levels in the area are within safe limits. Additionally, the prevalence of higher <sup>232</sup>Th concentration over <sup>238</sup>U in all samples suggests that the soil in the study region is thorium enriched. The hills with granite and gneiss rocks exhibited the highest radioactivity concentrations, whereas sandstone rocks had lower concentrations. When compared to the world average absorbed dose rate, the calculated values suggest that the study area is safe in terms of radiation levels. As this study is the first of its kind, it holds importance for future researchers in the field of natural radioactivity mapping and radiometric measurements in hilly areas. In summary, the results of this study indicate that the soil in both hill and field areas is suitable for use in construction materials and does not pose any health risks to residents.

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# Declarations

**Conflict of interest** The authors declare that they have no known conflicting financial interests or personal relationships that could have ap-

peared to influence the findings represented in this paper. The authors declare that there is no conflict of interest in the publication of the paper.

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