


# Late Holocene hydrology of Lake Maharlou, southwest Iran, inferred from high-resolution sedimentological and geochemical analyses

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**Abstract** Sedimentological (dry density, micro-facies analysis on thin-slides, X-ray diffraction, layer counting) and elemental analyses (X-ray fluorescence) of a 3.5-m-long sediment core (MAH-B) from saline Lake Maharlou (SW Iran) were used to infer hydrological changes over the last ~ 3800 years. The sediment consists of thin, alternating beds of evaporites (halite, gypsum), carbonates (calcite, aragonite) and detrital minerals (quartz, muscovite, chlorite). We interpret the data with respect to three main hydrologic conditions: (1) nearly complete desiccation of the lake, marked by frequent halite layers, (2) hypersaline conditions marked by gypsum abundance, (3) wet

conditions, characterized by high river-borne terrigenous sediment input with high potassium content. Distinct flood layers indicate fluvial activity. From about 3800–2000 cal year BP, terrigenous sediment supply was high, with peaks at 3700–3650, 3500–3450, and 3400–3250 cal year BP. Evaporative conditions were not common. From ca. 2000 cal year BP to present, detrital minerals are less abundant, and the increase in gypsum abundance suggests drier climate with enhanced evaporation. Frequent desiccation events occurred from 1100 to 700 cal year BP. The late Holocene hydrology of Lake Maharlou corresponds well with records of dune formation and phases of river alluviation in Iran and the Arabian Peninsula, and with regional speleothem records. In addition to the influence of climate change on the lake hydrology, ancient humans modified water drainage in the catchment. Periods of subterranean tunnel (*qanats*) use correspond to dramatic shrinkage of the water body. We propose that climate changes drove long-term water shortages that were enhanced by anthropogenic activities, leading to more frequent desiccation of the lake during the last millennium.

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## Introduction

Arid and semi-arid regions, which cover one quarter of the continental Earth surface, are particularly vulnerable to droughts because of the scarcity of water resources. In Iran, especially in the southeast and southwest regions, the severity of drought has increased during the last 30 years (Bari Abarghouei et al. 2011). Climate models predict that Iran's wet regions will receive more rainfall, while its dry regions will be generally drier, continuing this trend of protracted droughts (Abbaspour et al. 2009). Droughts superimposed on the overall drying have significant economic and environmental consequences. The severe drought in AD 1999–2000 resulted in a 30% loss of productivity in the agricultural sector of Iran compared with previous years (Salami et al. 2009) and also badly affected livestock, fisheries, and forestry. The combination of drought, the rising demand for water for irrigation and domestic use and poor water management have critically affected all water resources of the country (Motagh et al. 2008).

During the last decade, numerous lakes of the semi-arid Iranian plateau (Urmia, Hamoun, Parishan, Maharlou) have reached critical water levels. Lake Urmia, one of most emblematic lakes in the Near East because of its large size, lost 70% of its surface area between 2002 and 2014, despite its environmentally protected status (Wurtsbaugh et al. 2017). The disappearance of lakes in several regions of Iran reveals the serious problem of long-term sustainability of surface water resources in the region, but also highlights the depleted state of the ground-water resources. The main cause of the recent drop in lake levels is anthropogenic alteration of the water system (Tourian et al. 2015), including dam construction, channel deviation, and groundwater pumping, all of which result in a water deficit.

Paleolimnological studies allow to better understand changes in a lake's hydrological regime over long timescales (Dearing 2013), providing boundary conditions that help establish sustainable management strategies of water resources (Coops et al. 2003). In arid and semi-arid regions of Iran, sediment archives from hydrologically closed lakes have demonstrated their high potential for deciphering paleohydrological changes (Stevens et al. 2012; Jones et al. 2015; Aubert et al. 2017) through a range of biotic and abiotic indicators (ostracods, diatoms, chironomids, oxygen

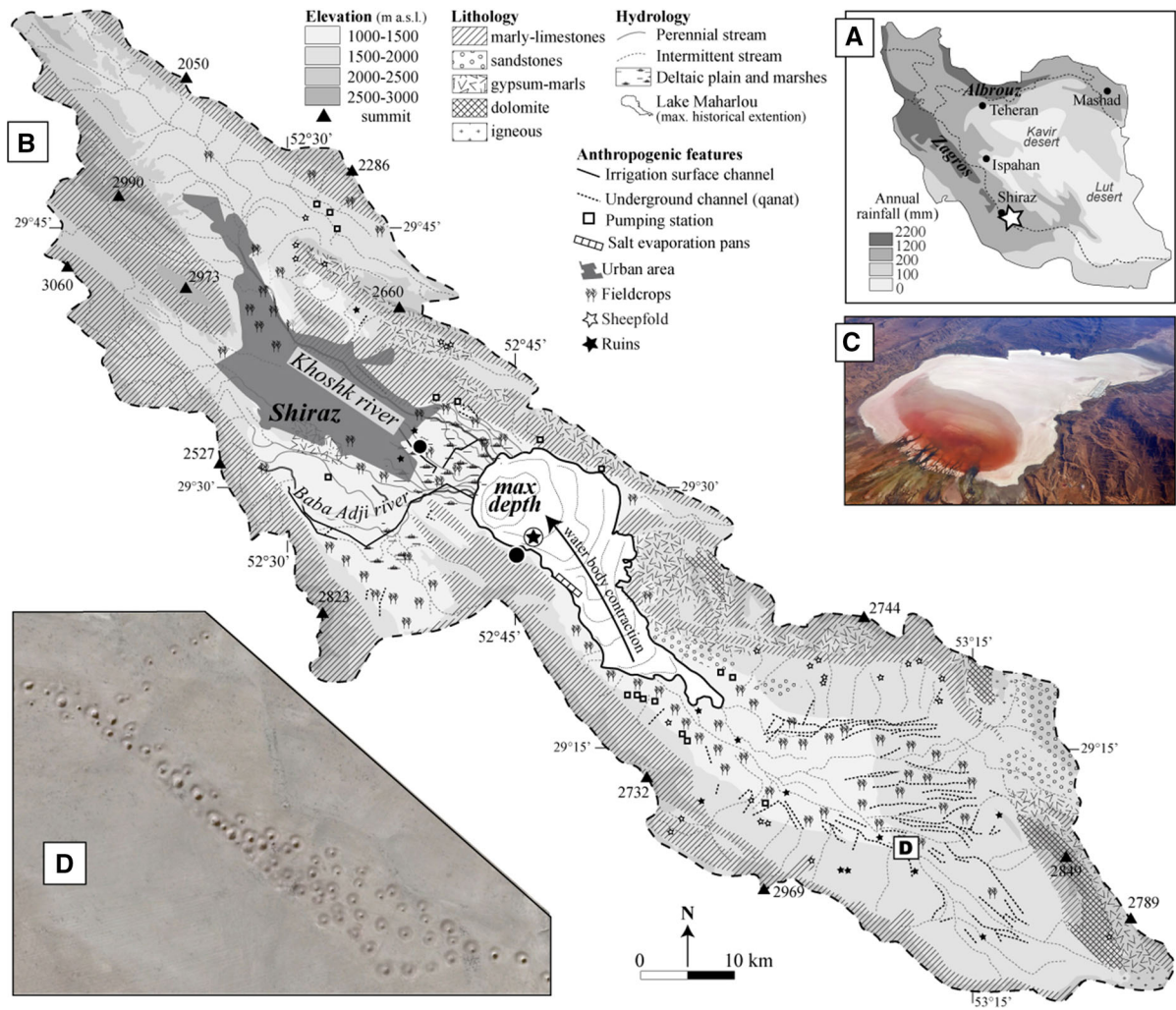
isotopes of endogenic carbonates). Fewer investigations have reconstructed lake paleohydrology in Iran with sedimentology and major element geochemistry, with the exception of Lake Urmia (Kelts and Shahrabi 1986) and the recent studies on Lake Neor (Sharifi et al. 2015). These proxies may provide valuable high-resolution data from which lake desiccation and fluvial activity can be reconstructed (Neugebauer et al. 2016). In this study, we present new high-resolution sedimentological (dry bulk density, micro-facies analysis, X-ray diffraction) and major element (X-ray fluorescence) analyses of MAH-B, a 3.5-m-long sediment core, to investigate centennial-scale variations in the hydrology of Lake Maharlou for the last 3800 years.

## Site description

Lake Maharlou is located 10 km southeast of the city of Shiraz (29°29'N, 52°45'E), in an intermontane depression of the Zagros Fold-Thrust Belt. The lake, with an elevation of 1460 m a.s.l., lies in a vast, 3928-km<sup>2</sup> endorheic watershed surrounded by mountains that reach 2800 m a.s.l. Basement rocks mainly consist of Mesozoic/Cretaceous limestone-dolomites and marls. Several evaporitic outcrops, including gypsum, anhydrite, and halite beds, are exposed in the eastern part of the lake (Faghih et al. 2012; Fig. 1).

Iranian climate is characterised by hot and dry summers associated with northerly winds (i.e. Shamal) from the Mediterranean high to South Asian monsoon low (Tyrlis et al. 2013). Persistence of this synoptic pattern produces great dust storms along the Persian Gulf that do not affect the southern continental slope of the Zagros (Nabavi et al. 2016). Winters are dominated by cold dry air from the Siberian high. Penetration of the most intense depressions from the Atlantic and Mediterranean triggers winter and spring precipitation (Alijani and Harman 1985), especially in the Zagros massif where snow is deposited at high altitude.

Lake Maharlou was a nearly permanent lake in the recent past, but since its near desiccation during summer 2002, seasonal drying of the lake has become a common phenomenon (Dumas et al. 2003; Fayazi et al. 2007). Today, the depth of the lake varies seasonally from a completely dried surface in summer months to about 3.5 m in early spring. The depth of the basin is greater to the northwest (Fig. 1). The area of the lake at a high stand of 3.5 m depth is 275 km<sup>2</sup>



**Fig. 1** **a** Spatial distribution of annual precipitation in Iran and location of the study site near the city of Shiraz (white star). **b** Map of the Lake Maharlou watershed, including topography, geology, hydrology, and anthropogenic features. Note that map details differ between the northern and southern part of the watershed because several sources of information were integrated. The black star indicates the MAH-B coring location. The two black dots indicate the location of the surface sediment

samples. **c** Aerial photograph of Lake Maharlou during the phase of nearly complete drying, with the view toward the southeast. Spatial extension of the halite crust is visible in white and hypersaline lake waters in red. **d** Satellite image (Google Earth) of the southeastern part of the watershed (location in **b**), representing several generations of regular holes of a *qanat* irrigation system. (Color figure online)

(Dumas et al. 2003). A detailed hydrologic study of the lake based on instrumental measurements from 1988 to 1992 and hydroclimate data provided by the Karstic Research Center of Teheran showed that inputs to the lake originate from catchment runoff (85%) and karstic springs (15%) that feed two rivers, Khoshk and Baba Adji (specific discharge of 3.59 and 1.85 L/s/km<sup>2</sup>, respectively). These rivers form a delta in the northern part of the lake (Dumas et al. 2003). Karstic losses, a common phenomenon in Zagros (Rezaei et al.

2013), are considered negligible since the basin floor is covered by a thick impermeable clay deposit. Precipitation is seasonal, occurring mainly from January to April. Annual mean precipitation is estimated at 350 mm/year (Shiraz station). During the monthly-instrumented period of 1988, the negative balance between precipitation and evaporation was compensated for by karstic inputs, preventing complete drying of the lake (Dumas et al. 2003). Altogether, the hydrological regime of Lake Maharlou is driven by

karstic resurgence throughout the year, winter/spring rainfall, and summer evaporation. The watershed of Lake Maharlou is also occupied by a dense network of surface and underground tunnels (*qanats*) and small water pools that are partially abandoned today (Fig. 1b). These infrastructures, essentially located in foothills, provide a freshwater supply for agriculture and pastoral activities on the plain.

## Materials and methods

Four contiguous drives (MAH-B-D1, D2, D3 and D4), each 1 m in length, were taken with a Russian Corer in the southwest quadrant of the northwestern depression of Lake Maharlou (29°27'40.5"N; 52°43'45.6"E) in September 2012. Main sedimentary facies were described using macroscopic visual features (colour, texture and structure, presence of macrofossils). The result is a master core, 355 cm in length, designated MAH-B.

Seven samples were radiocarbon dated at the Poznan Radiocarbon Laboratory. To avoid contamination by fossil organic carbon and the 'hard-water effect,' only botanical macroremains and charcoals were dated. Because of the paucity of terrestrial macroremains, sediment samples were wet sieved at 100 µm, avoiding the use of alcohol. Terrestrial organic particles were then picked one by one under a binocular microscope. Humic acid removal pretreatment (Poznan Radiocarbon Laboratory) was skipped because of the small weight of the sample for preparation of graphite. The age-depth model was developed using a smooth spline (type 4 argument) in the Clam R package (Blaauw 2010), based on probability distributions of the <sup>14</sup>C ages, and the IntCal13 calibration curve (Reimer et al. 2013). Ten thousand iterations were run to calculate the best age probability and 2σ envelope. Salt layers that represent rapid sedimentation were considered instantaneous deposits and were excluded from the age-depth model.

Two sediment segments (110 mm × 20 mm) were sampled for thin sections at 110–120 cm and 279–291 cm. The segments were freeze-dried and impregnated with epoxy resin (MK Factory) under vacuum following the procedure of Brauer and Casanova (2001). Polished slides were observed under a compound microscope at 10×, 50×, and 100× magnification (IMBE laboratory). XRF elemental analysis

was performed on a 1-m-long U-Channel with an ITRAX™ core scanner (Cox Analytical Systems) at the CEREGE laboratory. High-resolution visual scans were acquired with an optical camera before covering the sediment with polyethylene foil to prevent drying during analysis. The radiographic scan was done at 200-µm increments with parameters fixed at 45 kV, 40 mA and a dwell time of 700 ms. The relative abundances of Si, K, Fe, Ti, Ca, S, Sr, Cl were measured by X-ray excitation using a Mo-source tube set at 30 kV and 45 mA for 15 s, at a step of 500 µm. Significant changes in the geochemical composition were defined by cluster analysis, both stratigraphically constrained and unconstrained, using the CONISS algorithm of the Rioja R package (Juggins 2012). The number of significant zones was defined by comparison to the Broken-stick model (Bennett 1996). A visual layer counting was done with the support of a high-resolution optical scan, a radiographic scan, and XRF element variations. The minimal resolution of layer counting was 1 mm.

Contiguous sub-samples, 2-cm thick, were collected on the remaining core. Seventy samples of a constant volume of 7 cm<sup>3</sup> were taken at approximately 10-cm intervals, dried at 60 °C for 48 h, weighed to calculate dry bulk density, and then crushed using an agate mortar. Additionally, two surface sediment samples, one from the Khoshk riverbed, and one from soil on the west bank of the lake near the coring site (Fig. 1), were analysed to characterize detrital mineral species of the watershed. Mineral composition of 20 bulk samples was characterised by X-ray diffraction on powder deposited on zero-background silicon plates (CEREGE laboratory) with a PANalytical X'Pert Pro diffractometer. To verify that drying temperature did not modify mineralogical species, one air-dried sample (< 25 °C) was measured. Samples were scanned from 5° to 70° (2θ), with an increment step of 0.033° and an acquisition time per sample of 3 h. The X-ray emission source, cobalt (λ = 1.79 Å), was set at 40 kV and 40 mA. Peak identification was performed using the automated peak-match software X'Pert High Score Plus (Degen et al. 2014).



**Results**

**Chronology**

Of the seven AMS <sup>14</sup>C ages (Table 1), six were used to create the chronological model. Given its very low carbon weight (0.03 mg C), Poz-83775 (269 cm) was very close to the detection limit of the instrument and was rejected. The remaining ages are stratigraphically coherent if either sample Poz-83773 or Poz-81500 is excluded (Fig. 2). These two samples have low but equivalent weights of carbon (0.1 mg C). Poz-83773 comprises a mixture of numerous half-mm-long particles of charcoal, seeds, and insects, whereas Poz-81500 consists solely of charcoal (~ 3 mm long). The most common cause of age bias in dating terrestrial material from lake sediment is long-term retention in the lake catchment. It is more likely that Poz-83773 was reworked than that Poz-81500 was rejuvenated. Sample Poz-83775 was accepted a posteriori because the 2σ calibration range fits the model uncertainty. Thus, the final age-depth model is based on six <sup>14</sup>C dates. The age-depth model (Fig. 2) indicates that MAH-B spans the last 3800 ± 370 years, with an average 2σ uncertainty of 250 years. The highest sedimentation rates occurred from 3700 to 3250 cal year BP (mean = 1.5 mm/year). The sedimentation rate was lowest (0.5 mm/year) between 3250 and 1000 cal year BP, increased from 2000 to 400 cal year BP (1.1 mm/

year), and decreased again from 400 cal year BP to Present (0.4 mm/year).

**Bulk sediment composition**

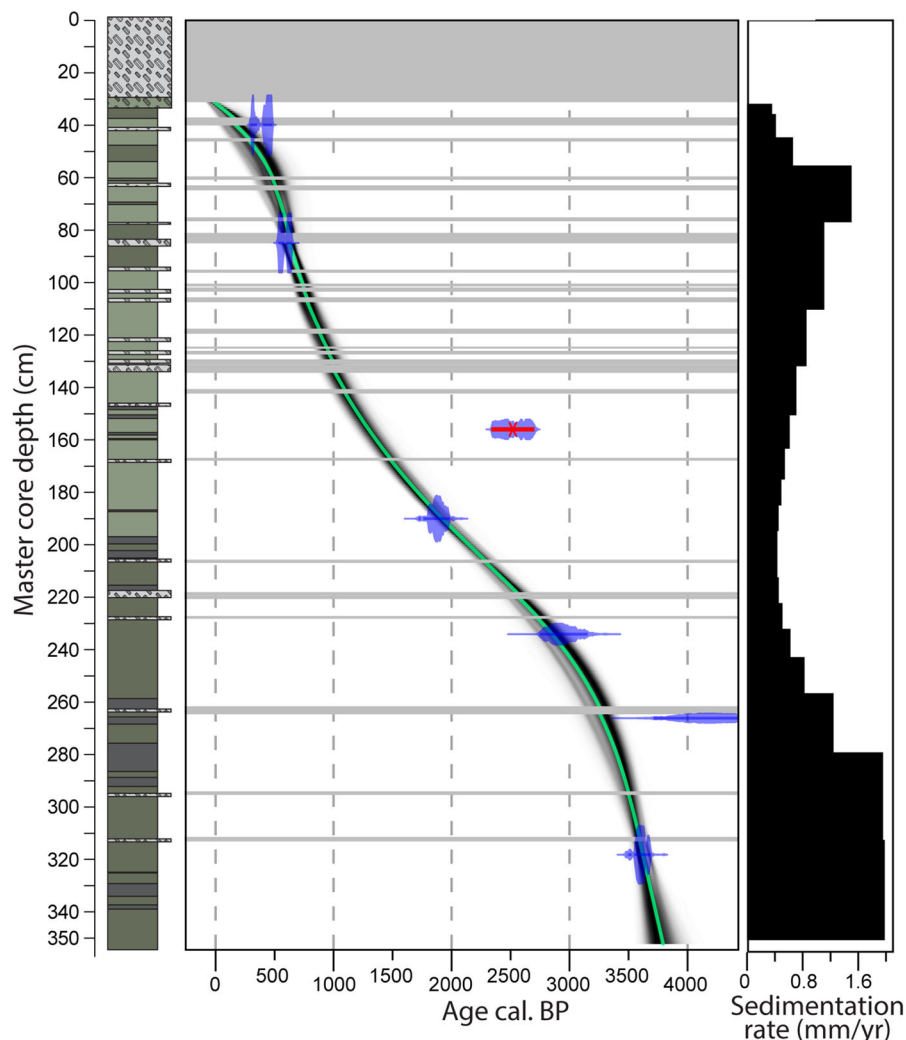
MAH-B is composed of very thinly bedded (1–2-cm thick) to laminated (0.1–1 cm) sediments that display an alternation of sand-size crystals and beige to light brown, silt–clay layers (Fig. 3). Contacts between beds are generally sharp, mainly because of the abrupt changes in particle size. The dry bulk density varies between 1.2 and 2.0 g/cm<sup>3</sup>. The upper 30 cm of the core is exclusively halite, reflecting the presence of the decimetre-thick salt crust encountered during coring. XRD measurements indicate the presence of both chemical and clastic components in all 19 samples. The comparison between the samples dried at different temperatures (25 °C vs 60 °C) indicates that sample desiccation induced a partial dehydration of gypsum to bassanite. Thus, the presence of bassanite is a preparation artifact and is considered to have been gypsum. Ranking minerals from the most dominant to least, the sediment consists of a constant presence of gypsum (CaSO<sub>4</sub>·2H<sub>2</sub>O), halite (NaCl), strontium in aragonite (CaCO<sub>3</sub>-Sr), calcite (CaCO<sub>3</sub>), quartz (SiO<sub>2</sub>), muscovite (KAl<sub>2</sub>(Si,Al)O<sub>10</sub>(OH)<sub>2</sub>) and chlorite ((Mg,Fe)<sub>3</sub>(Si,Al)<sub>4</sub>O<sub>10</sub>). Dolomite (CaMg(CO<sub>3</sub>)<sub>2</sub>) and celestine (SrSO<sub>4</sub>) occur regularly, but XRD peak heights are lower. Occurrences of pyrite (FeS<sub>2</sub>) at 92 and 337 cm, glauberite (Na<sub>2</sub>Ca(SO<sub>4</sub>)<sub>2</sub>) at 32 cm, and

**Table 1** Radiocarbon ages obtained on core MAH-B

Core section	Core section depth (cm)	Master core depth (cm)	Material	Laboratory code	Age <sup>14</sup> C BP	Age cal year BP (± 2σ)	Weight (mg C)
D1	43	43	Charcoals, Coleoptera, seed	Poz-81502	390 ± 30 BP	320–509	0.2
D1	88	88	Plant fragment	Poz-81501	655 ± 35 BP	556–673	0.15
<i>D2</i>	<i>62–70</i>	<i>159</i>	<i>Charcoals, Coleoptera, seed</i>	<i>Poz-83773</i>	<i>2450 ± 50 BP</i>	<i>2359–2707</i>	<i>0.12</i>
D3	14	193	Charcoal	Poz-81500	1950 ± 50 BP	1741–2002	0.1
D3	58–66	237	Charcoals, Coleoptera	Poz-83774	2820 ± 90 BP	2762–3162	0.07
D3	90–96	269	Charcoals, Coleoptera, seed	Poz-83775	3800 ± 180 BP	3717–4803	0.03
D4	66–74	321	Charcoals, Coleoptera, seed	Poz-83776	3370 ± 35 BP	3496–3697	0.16

The rejected age is in italics

**Fig. 2** Age-depth model for core MAH-B and lithostratigraphy and sedimentation rate. The AMS  $^{14}\text{C}$  ages are represented by their probability density functions. The age rejected before modeling is crossed in red. The model calculated with 10,000 iterations is given in the grey scale (higher probability is darker), and the best model is shown as a green line. Grey horizontal band illustrates a layer interpreted as non-continuous sedimentation (coarse crystal beds), which were “removed” before age modeling

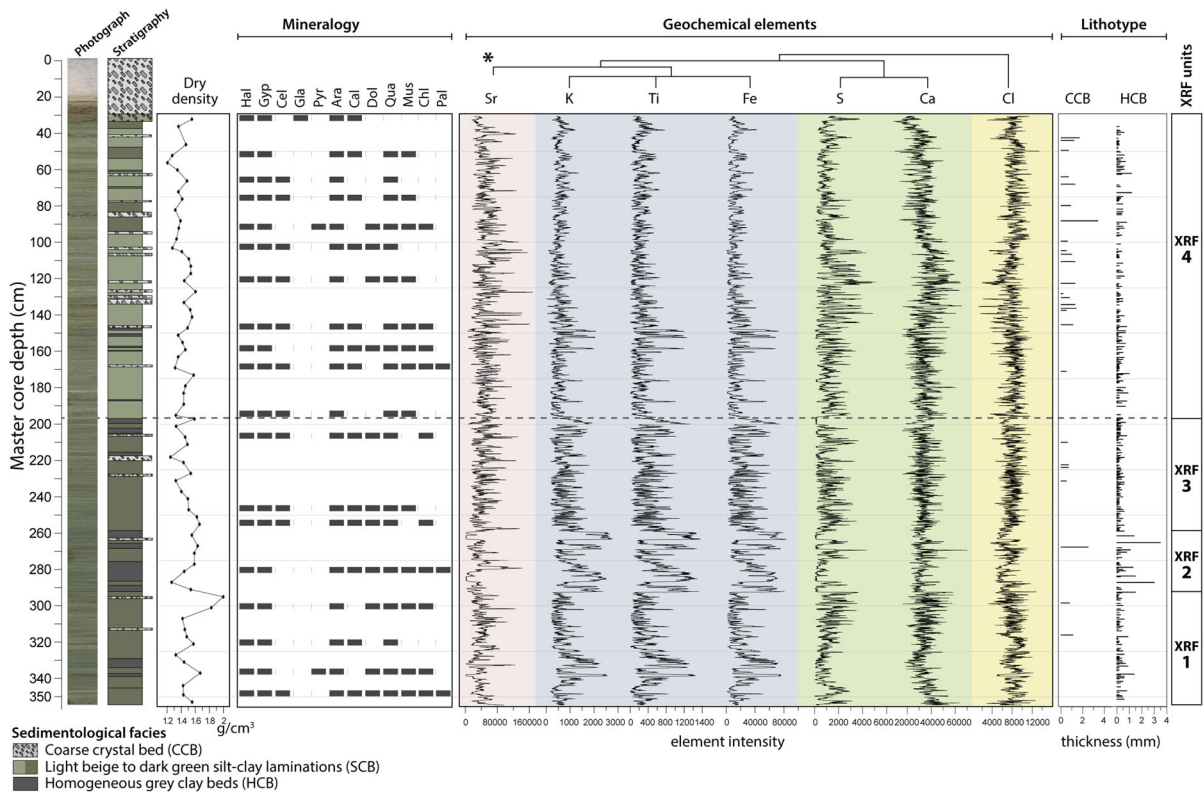


palygorskite ( $(\text{Mg,Al})_2\text{Si}_4\text{O}_{10}(\text{OH})\cdot 4(\text{H}_2\text{O})$ ) at 159, 281 and 349 cm were also noted. Mineralogical analysis of the surface sediment from the west bank of the lake near the coring site, as well as the sample from the Khoshk riverbed, the main tributary river of Lake Maharlou (Fig. 1), showed a similar composition of mostly quartz and calcite, with minor amounts of dolomite.

#### Lithofacies

Microscopic examination of thin sections and sieved fractions greater than  $100\ \mu\text{m}$  provided qualitative information on sediment texture and constituents, enabling the designation of three lithofacies (Table 2).

Coarse Crystal Beds (CCB) are characterised by the nearly exclusive presence of euhedral mm-size halite crystals free of finer matrix (Fig. 4b). CCBs are represented in dark grey on the radiograph (Fig. 4a). Beige to grey Silt–Clay Beds (SCB) consist of non-graded silt particles of mainly palisade gypsum crystals (g1 in Fig. 4c), rounded gypsum crystal (g2 in Fig. 4c) and aragonite (Fig. 4d) with some halite crystals in association with a finer matrix of calcite (Fig. 4e) and clay. Dark SCB (D-SCB) are richer in clay compared to lighter SCB (L-SCB). These beds are represented in medium grey in radiographic images (Fig. 4a). Homogeneous Clay Beds (HCB) consist of non-graded fine-silt to clay with scattered amorphous



**Fig. 3** Lithostratigraphy, mineralogy, dry bulk density, XRF elemental profiles, and lithofacies deposition in core MAH-B. Mineralogical species are abbreviated: *Hal* Halite, *Gyp* Gypsum, *Cel* Celestine, *Gla* Glauberite, *Pyr* Pyrite, *Ara* Aragonite,

*Cal* Calcite, *Dol* Dolomite, *Qua* Quartz, *Mus* Muscovite, *Chl* Chlorite, *Pal* Palygorskite. Elements (Sr, K, Ti, Fe, S, Ca, Cl) are shown following their degree of similarity, given by cluster analysis (asterisk). (Color figure online)

organic particles, charcoals, and quartz. HCBs are represented in light grey in X-radiograph images.

The occurrence of the three lithofacies is repetitive down core. Most of the 26 CCBs (excluding the modern one) were deposited between 145 cm and the core top. CCB thickness ranges between 0.2 and 35.7 cm, with an average of 1.0 cm. L-SCB (224 beds) and D-SCB (196 beds) are relatively homogeneously distributed along the profile, with thicknesses ranging from 0.1 to 2.9 cm and average values of 0.42 and 0.48 cm, respectively. The 385 HCB have thicknesses ranging between 0.1 and 3.5 cm (mean 0.3 cm). HCBs are thicker between 285 and 260 cm (> 1 cm).

**Geochemical variations**

Depth-constrained cluster analysis of the high-resolution XRF profiles revealed four geochemical groups: K–Ti–Fe, Sr, Ca–S, and Cl. K–Ti–Fe are mostly

associated with Sr and to a lesser extent with Ca and S. Cl is an isolated unit (Fig. 3).

*The Cl group*

Cl intensity represents halite abundance in the bulk sediment. The larger Cl fluctuations correspond to CCB, although isolated crystals of halite sometimes occur in lithofacies L-SCB and D-SCB (Fig. 4A-1 and A-2). Cl intensities are minimum in HCB beds characterised by higher detrital mineralogenic elements (Fe, Ti, K).

*The group Ca, S*

Ca can be associated with gypsum, aragonite, calcite, and dolomite, whereas S may originate from gypsum, pyrite, and celestine. Both elements, however, are more concentrated in light SCBs (Fig. 4a) that consist mainly of gypsum and aragonite. Low correlation

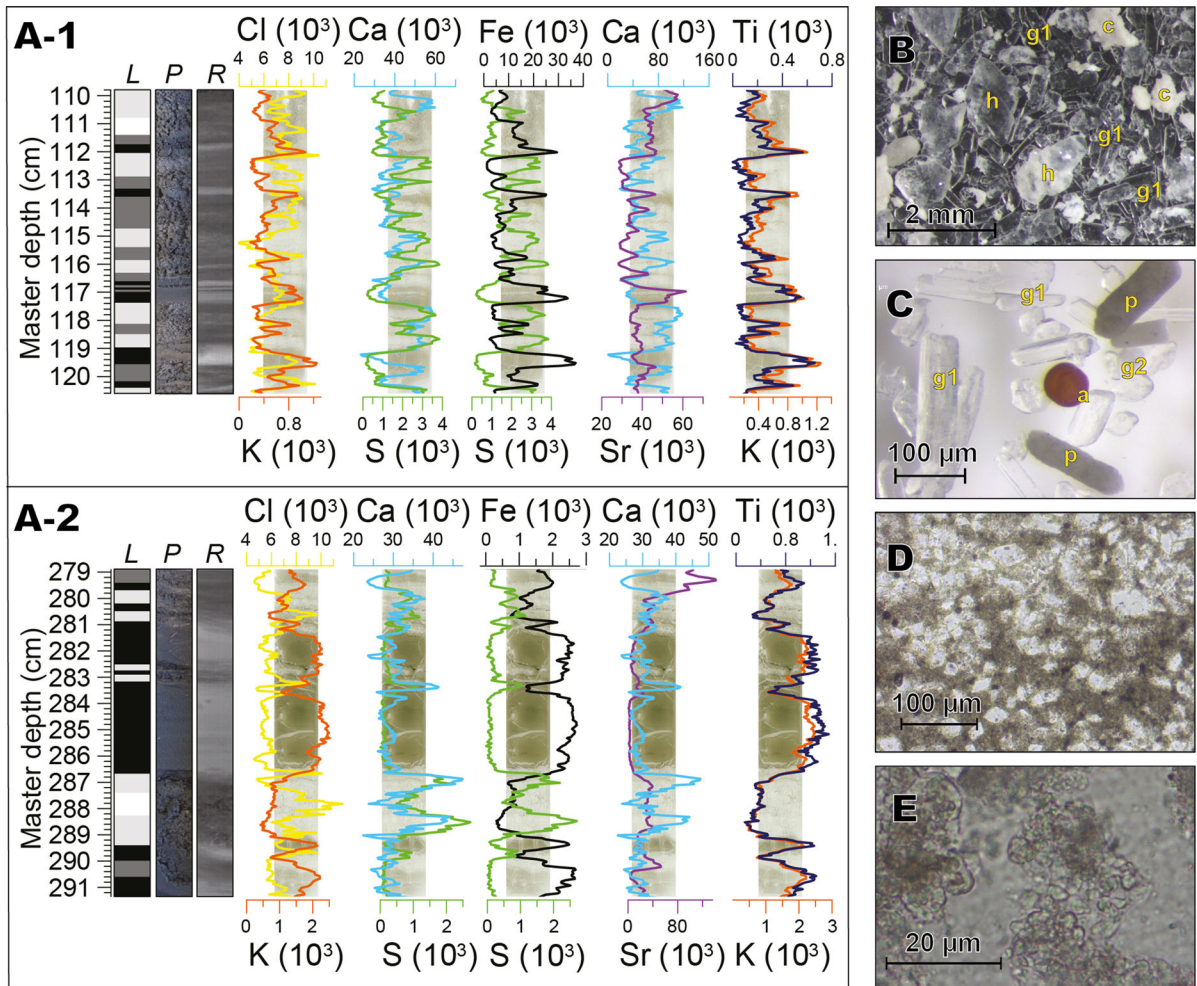
**Table 2** Sedimentological and geochemical composition of the MAH-B lithofacies

Lithofacies (CODE)	Description	Main minerals	XRF element	Layer thickness (mm)	Interpretation
Coarse Crystal Bed (CCB)	Transparent euhedral cubic crystal (mm-size)	Halite	Cl	From 0.2 to 3.4 cm (mean: 1.0 cm), excluding the 35.7 cm of top core	Residual indicator of a summer salt crust indicating a complete drying of the lake surface
Beige to grey Silt–Clay bed (SCB)	Non-graded silt–clay mud made of palisade gypsum crystal, eroded gypsum, aragonite and of a finer matrix comprising slightly rounded rhombs to axially symmetric ellipsoids of calcite, with muscovite and chlorite  The fraction > 100 µm (Fig. 3) is composed of few halite crystals (h), gypsum needles (g), <i>Artemia</i> cysts (a), concretion of <i>Artemia</i> pellets (p), charcoal particles, gastropods and ostracods	Gypsum	S, Ca	L-SCB (Light Silt–Clay Bed):	Sediment accumulation dominated by dilution between three main chemical components (halite, calcium sulphate and calcium carbonate) and of a fourth clastic inorganic (quartz, muscovite and chlorite). Presence of subangular gypsum crystals indicates eolian deflation  Beds are lighter or darker depending on mineral abundance in evaporate/calcium carbonates or in clay. Lighter beds indicate mainly chemical precipitation, and darker ones a more important detrital supply in a context of sedimentation dominated by chemical precipitation
		Celestine	S, Sr	From 0.1 to 2.9 cm (mean: 0.4 cm)	
		Calcite	Ca		
		Aragonite	Sr, Ca	D-SCB (Dark Silt–Clay Bed):	
		Dolomite	Ca	From 0.1 to 2.5 cm (mean: 0.4 cm)	
Homogeneous Clay Bed (HCB)	Non-graded clastic fine-silty to clayey mud. Charcoals and amorphous particles are scattered within the bed	Quartz	Si	From 0.1 to 3.5 cm (mean: 0.3 cm)	Clastic detrital inputs intercalating the biochemical lacustrine sedimentation, indicating phases of abundant river detrital supply
		Muscovite	K, Ti, Fe		
		Chlorite			

between S and Fe, the rare occurrence of pyrite in the entire sediment sequence (Fig. 3), and the low S intensities in the samples where pyrite was detected by XRD, indicate that S does not track pyrite. Poor correlation between S and Sr concentrations, along with the rare presence of celestine (Fig. 4a), also argues that S does not track celestine. The positive linear relationship between S and Ca likely reflects the dominance of gypsum in controlling the S concentration, but the modest correlation coefficient ( $r = 0.5$ ;  $n = 6360$ ) suggests that Ca values are also affected by other Ca-rich minerals (calcite, aragonite, and dolomite). Comparison between the ratio Ca/S (data not

shown) and detrital elements presents a moderate positive correlation (e.g. with K,  $r = 0.33$ ). This relationship suggests that some of the Ca may have a detrital origin. This supposition is supported by the presence of Mesozoic/Cenozoic carbonate rocks in the watershed and the presence of these minerals in riverbed sediments. Together, these results show that Ca probably has multiple mineral associations and origins (i.e. endogenic and allochthonous), whereas S is more securely driven by gypsum abundance only. As gypsum minerals occur mainly as palisade crystals, the main process of gypsum deposition in the lake sediment is as in situ, perpendicular growth relative to





**Fig. 4 a** Macro- and microscopic characterisation of the lithofacies from two sediment segments corresponding to two thin sections **A-1** (110–120 cm) and **A-2** (279–291 cm). Lithofacies (L) distinction (white: Coarse Crystal Beds (CCB); light grey: Light Silt–Clay Beds (L-SCB); dark grey: Dark Silt–Clay Beds (D-SCB); and black: Homogeneous Clay Beds (HCB)), photographic image (P), radiographic scan image (R), and superposed thin-slide optical scans and high-resolution

curves of the XRF intensities of Cl, K, Ca, Fe, S and Ti (corresponding scale to curve is shown by colour). **b, c** > 100- $\mu$ m sieved fraction made of: (h) halite, (g1) palisade gypsum crystal, (g2) subangular gypsum crystal, (c) calcium carbonate, (a) diapauses cysts of *Artemia* and (p) *Artemia* aragonite pellets. **d** Euhedral gypsum crystals included in a calcite-clay matrix. **e** Ellipsoidal coalescent calcite particles

the sedimentation surface, in a hypersaline pool. In lesser abundance, abraded gypsum crystals suggest allochthonous processes that may be associated with mechanical erosion from evaporitic outcrops or reworking by wind on the exposed dry lake surface. Absence of gypsum in riverbed sediments, together with the lack of a relationship between S and detrital element (K, Fe, Ti), refutes the hypothesis of supply by rivers.

*The group K, Ti, Fe*

Intensities of these elements are highest in HCBs and relatively high in D-SCBs. This group is associated with the presence of muscovite, chlorite and palygorskite minerals. Muscovite and chlorite are products of weathering of Mesozoic/Cenozoic rocks and are incorporated into the lake sediments as terrigenous, silt–clay-sized particles that are transported by fluvial or eolian processes. Although origin of palygorskite mineral may be detrital or endogenic, its occurrence in

Lake Maharlou is associated only with the HCB lithofacies, suggesting a strictly detrital origin. Moreover, if the Si XRF signal is low (data not shown), Si shows the same variability as K, Ti, and Fe, indicating the concordance of quartz together with phyllosilicates in HCB. Rare presence of diatoms confirms this interpretation. Two alternative interpretations can be proposed for the detrital element group. K, Ti, Fe may mostly track phyllosilicates that are deposited from turbid river discharge during heavy rainfall and thus be related to relatively high lake level, or phyllosilicates may be of eolian origin signalling arid conditions and thus be related to lake desiccation.

#### *The Sr group*

Peaks of Sr intensities are generally, but not exclusively, associated with L-SCBs and D-SCBs. Higher Sr intensities occur at depths where no celestine was recorded (e.g. 280–340 cm). Besides, Sr is often used to track Sr-rich aragonite (Böning and Bard 2009) that possibly reflects the abundance of aragonite-rich fecal pellets of the brine shrimp, *Artemia* (Fig. 4c). The lack of correlation between Ca and Sr, however, cannot confirm this hypothesis. Dominance of gypsum likely overrides the less abundant aragonite and complicates the Ca/Sr relationship.

## Discussion

### Defining palaeoenvironmental proxies

#### *Halite and gypsum as indicators of lake desiccation*

The sedimentological and geochemical analyses of MAH-B revealed a large range of biochemical and evaporitic products that are characteristic of different stages of concentration of the water (brine), related to retraction of the lake. Along the gradient of brine ionic concentration, low-solubility minerals such as calcite, aragonite and dolomite, are precipitated first, followed by gypsum, and finally halite (Hardie and Eugster 1970). Based on the analysis of the ion concentration of Lake Maharlou brine in summer 2002 and spring 2005, Fayazi et al. (2007) showed that during the driest conditions the salt crust comprises only halite, whereas under wetter conditions, several evaporite minerals, including mainly halite, gypsum, and

calcium, are precipitated. In MAH-B, gypsum is mainly present in the form of elongated prismatic crystals and occurs in SCBs in association with carbonates and scattered halite grains. Gypsum formation requires a pool of water for the growth habit of the crystals.

In addition to variations in brine concentrations, other environmental factors can affect which minerals are deposited in a lake. Heterogeneities of basin morphology and proximity to freshwater inputs induce spatial variability with respect to the saturation point of different minerals (Bowler 1986), which are reached at different stages of water body contraction, depending on the location of the coring site (Rahimpour-Bonab and Abdi 2012). Distribution of cation concentrations in surface waters of Lake Maharlou indicates that the southeast of the lake has higher concentrations of  $\text{Na}^+$  and  $\text{Cl}^-$  (Fayazi et al. 2007) than the northeast deep basin, which is fed by the main tributaries. The MAH-B core site was located in the latter deep basin, but relatively close to the lake margin (Fig. 1). This location is more prone to surface crust formation than the deepest part of the lake, but less so than the shallower southern plateau. Thus, MAH-B is sensitive to an intermediate stage of water body contraction, corresponding to half the maximum current lake surface. Post-depositional mineral dissolution caused by cyclic flooding-desiccation also affects the integrity of hydrologic reconstruction (Last and Schweyen 1983). Dissolution is a differential process, first affecting the most soluble salts. Thus, despite the fact that gypsum can represent the complete record of changes in lake sedimentation, more soluble halite layers probably underestimate the number and intensity of past extreme desiccation events. Thus, gypsum abundance, tracked by S XRF intensity, is interpreted as an indicator of hypersaline conditions in Lake Maharlou, with some water still present. The frequency of indurated halite beds (CCB lithofacies) is interpreted as the residue of summer salt crusts, which indicate desiccation of more than 50% of the lake surface, thus corresponding to very low lake levels (< 1 m) even in the deepest part of the lake.

#### *Detrital fraction as an indicator of river activity*

Detrital minerals (phyllosilicates and quartz) have been identified in HCB and D-SCB lithofacies and characterised by high XRF intensities of K, Ti, and Fe.

These minerals have two possible origins, eolian or alluvial. Dust storms that carry high sediment loads originate from two main atmospheric patterns. Shamal summer winds, oriented northwest–southeast, come from the Mediterranean and produce large dust deposits along the Euphrates Valley to the Persian Gulf. Frontal winds, coming from Saudi Arabia, generate dust storms in the northern part of the Zagros to the Caspian Sea (Nabavi et al. 2016). A retrospective atmospheric analysis, covering the period 1980 to Present (Nabavi et al. 2016), indicated that the Lake Maharlou Basin has not received substantial dust transport, in part because of the protected position of the basin with respect to the dominant summer winds (Rezazadeh et al. 2013). The same situation was observed for a major dust storm in July 2009, which did not trigger dust deposition in the study region (Hamidi et al. 2014). In addition, mineralogical and elemental studies of dust deposited in sediment traps in the Persian Gulf, along the Iranian coast, indicated the dust was dominated by calcite (60%) and quartz (15%) minerals (Ahmady-Birgani et al. 2015), with grain sizes of silt and clay. In Lake Maharlou sediments, although the grain size of dust is similar to the HCBs particles, these beds have low to no calcite content and are depleted in Ca (Fig. 4a). It is thus unlikely that detrital-rich HCBs correspond to dust-storm deposits. In contrast, the hypothesis of terrigenous input from rivers is supported by the occurrence of intense floods along the Khoshk River in the last few decades, which caused economic losses in the city of Shiraz (e.g. in December 1986 and February 1998) (Kowsar et al. 2009; Mardaneh and Keshtkar 2013; Jamali et al. 2015). Historical archives also mention major floods that devastated the town in AD 1630 and 1668. The hydrograph of the Khoshk River system indicates discharges of 200 and 350 m<sup>3</sup>/s with return periods of 25 and 100 years, respectively (Rivaz and Musavi-Jahromi 2012). Considering these two possible origins of the detrital material, i.e. dust storms and river supply, the latter is more likely to cause HCB layers and increases in K intensities.

HCB lithofacies are characterised by absence of graded microfabrics, suggesting that detrital sediment supply entered the lake and that the kinetics of deposition results from uniform suspension in the water column (Siegenthaler and Sturm 1991). The absence of internal, microscopic, graded bedding may also result from deposition distal from the river mouth

(Czymzik et al. 2013). MAH-B is located in the same northern basin as the river delta and in the path of the entering flow. The mouth of Khoshk River forms a vast (8.5-km-wide), relatively flat (3° slope), fluvio-lacustrine delta that encompasses the entire width of the lake. This complex deltaic environment consists of numerous wetlands in which divergent streams, both intermittent and perennial, flow. Taking into account the low potential velocity and relatively low stream discharge, two explanations for the absence of coarse detrital particles in MAH-B are (i) quasi-permanent presence of dense aquatic vegetation (*Juncus*, *Phragmites*) in wetlands developed in the Khoshk River Delta environment, which buffer and trap the coarse sediment fraction, and (ii) sediment supply feeding the lake system and the coring site consists of turbid waters, with only clays/silts in suspension.

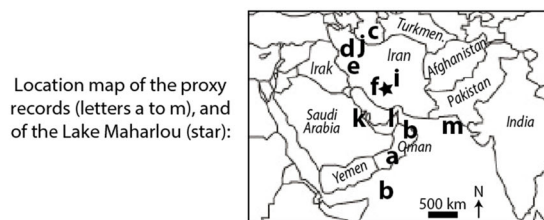
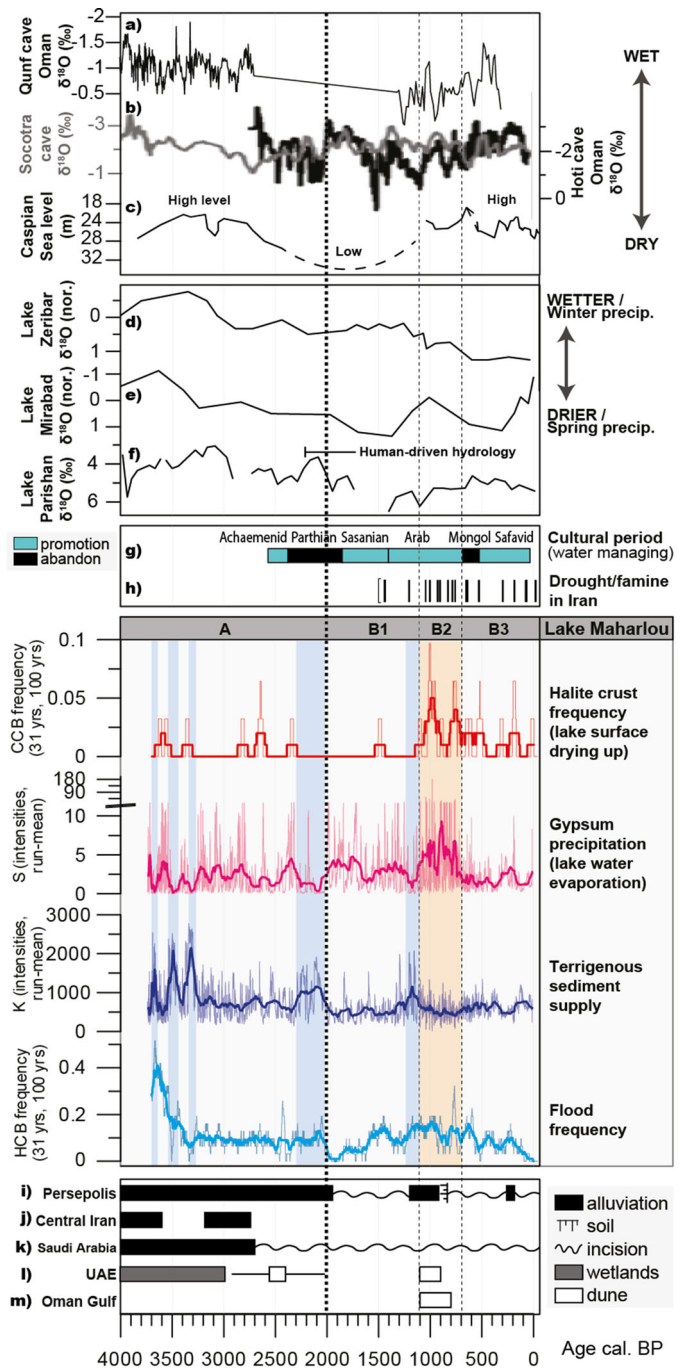
Climate: a forcing mechanism for hydrologic changes in Lake Maharlou?

*3800 to 2000 cal year BP (period A): enhanced detrital supply during wetter climate conditions*

The base of MAH-B, from 3800 to 2000 cal year BP, records peak abundance of terrigenous sediment (high K intensities, Fig. 5). High fluvial activity during this period corresponded with relatively low evaporative concentration of the lake. Rare phases of lake contraction (inferred from S intensities and CCB frequency, Fig. 5), however, did occur at 3700–3650, 3500–3450, and 3400–3250 cal year BP. Thus, this period is interpreted as having an overall positive hydrological balance, corresponding to higher river flow and lower water loss through evaporation. Superimposed on these general conditions are three K peaks and a peak in HCB frequency from 3800 to 3250 cal year BP. Enhanced detrital supply could be caused by rapid snow melt or large spring rains (Masih et al. 2011), possibly related to a highly unstable climate. Alternatively, it could be caused by the river adjusting course, moving its mouth closer to the coring location. From 3250 to 2000 cal year BP, flood frequency and K intensities decreased, suggesting more stable hydrological conditions and fewer pulses of fluvial sediment.

Higher river discharge in the Maharlou catchment fits well with other records that show a period of river alluviation in the Pulvar (Rigot 2010) and Kor (Kehl

**Fig. 5** Palaeohydrology reconstructed from Lake Maharlou proxies compared to records in southwest Asia, as a function of age (cal year BP). **a** Qunf Cave, southern Oman (Fleitmann et al. 2003, 2007). **b** Socotra Cave, Yemen (Van Rempelbergh et al. 2013); Hoti Cave, northern Oman (Fleitmann et al. 2007). **c** Level of the Caspian Sea (Rychagov 1997; Kroonenberg et al. 2007). **d** Lake Zeribar (Stevens et al. 2001; Snyder et al. 2001). **e** Lake Mirabad (Stevens et al. 2006). **f** Lake Parishan, Iran (Jones et al. 2015). **g** Historical cultural periods in the Fars region (Iran) and attention paid to repair and construction of new *qanats*. **h** Historical archives citing drought or famine in Iran (Frye 1984; Melville 1984; Daryae 2014). **i** Geomorphic record of the Pulvar fluvial valley of the Persepolis basin, Iran (Rigot 2010). **j** Geomorphic record of the Central Iranian Plateau (Schmidt et al. 2011). **k** Geomorphic record of the Wadi ad Dawasir in central Saudi Arabia (Matter et al. 2016). **l** Inter-dune lake Awafi (Parker et al. 2006) and Idhn in United Arab Emirates (Goudie et al. 2000). **m** northern Arabian Sea offshore marine record (Lückge et al. 2001)





et al. 2009) River Valleys, which are located 50 km north of Lake Maharlou in the Persepolis Basin (Fig. 5i), alluviation in the northern part of the Iranian Plateau (Schmidt et al. 2011, Fig. 5j) and in the Arabian Peninsula (Matter et al. 2016, Fig. 5k). River terrace formation probably indicates that sediment was available on slopes and that water discharge was sufficient to transport those sediments. More humid conditions, which would be associated with stable river input, also explain the reduction in evaporative concentration of the lake surface. A generally wetter climate with pulsed river activity is consistent with the presence of inter-dune wetlands that indicate short-lived wet phases in the Arabian Peninsula (Parker et al. 2006, Fig. 5l) and in a marine record from the Gulf of Oman (Lückge et al. 2001, Fig. 5m). The interpretation of wetter conditions in Southwest Asia from 3800 to 2000 cal year BP is also consistent with proxies that are more sensitive to effective moisture (Fig. 5a, b, c) than are geomorphic records. Lower values of the  $\delta^{18}\text{O}$  of speleothems from Qunf Cave in Oman (Fleitmann et al. 2003, 2007, Fig. 5a) and in Gejkar Cave in Iraq (Flohr et al. 2017), and higher levels of the Caspian Sea (Rychagov 1997, Fig. 5c), all argue for a regional increase in moisture. In the Iranian Plateau, higher lake levels from 3700 to 1900 cal year BP are documented by dominance of aquatic plants in the pollen diagram from Lake Almalou, with a peak high stand between 3700 and 3400 cal year BP (Djamali et al. 2009a). In Lake Parishan, some 100 km to the west of Lake Maharlou, both geochemical and paleoecological proxies suggest that between 3800 and 2200 cal year BP the lake was higher than periods before and after this interval (Jones et al. 2015).

*2000 to Present (period B): increasing aridification and more lake low stands*

The main change in the catchment/lake hydrology occurred at ca. 2000 cal year BP, with an increase in gypsum abundance attesting to more evaporative conditions in the lake and decreased terrigenous input (Fig. 5). The timing of the transition is consistent with a major environmental change recorded in geomorphic records from the Iranian Plateau and the Arabian Peninsula (Fig. 5i–l): increased river incision and dune development at that time suggest drier climatic conditions. More arid conditions in this region are in

phase with a lower Caspian Sea level (Rychagov 1997) and more positive  $\delta^{18}\text{O}$  values from the Qunf Cave (Fleitmann et al. 2003). A transitory increase in the detrital supply in Lake Maharlou between 1300 and 1100 cal year BP is concomitant with increased alluvial sediment deposition in the Persepolis Basin (Fig. 5i) that may reflect a short phase of more humid conditions (Rigot 2010).

From 1100 to 700 cal year BP (period B2, Fig. 5), Lake Maharlou experienced maximum reduction of its water body and very frequent desiccation events, as revealed by high gypsum abundance and frequent deposition of halite beds (CCB frequency), respectively. Dune formation and dust events in the Arabian Peninsula (Goudie et al. 2000; Lückge et al. 2001) suggest that a more arid climate might be responsible for the extreme evaporation of Lake Maharlou. In the Persepolis Basin, soils developed on previously deposited terraces (Rigot 2010), also suggesting very low fluvial activity. Speleothem records from northern and southern Oman seem to also indicate several short periods of drought (Fleitmann et al. 2007). The most compelling argument, however, is the frequency of drought and famine recorded in historical archives from Iran during this interval (Daryae 2014; Sharifi et al. 2015; Fig. 5h).

Lake Maharlou underwent few desiccation periods and was likely a perennial water body from 700 cal - year BP to Present (period B3, Fig. 5), as indicated by a decrease in halite crust formation and gypsum abundance. Detrital input, however, remained low. These conditions correspond to fewer droughts and famines in Iran (Daryae 2014). Hoti and Qunf Cave (Fleitmann et al. 2003, 2007) display modest decreases in  $\delta^{18}\text{O}$  at that time, suggesting relatively wetter climatic conditions than in period B2. This interval corresponds roughly to the Little Ice Age (600–150 cal year BP), during which cooling associated with lower solar irradiance and higher frequency of stratospheric volcanic eruptions may have reduced evaporation stress (greater effective moisture), resulting in less halite crust formation at Lake Maharlou. Similar results are documented for Lake Almalou (Djamali et al. 2009a) and Caspian Sea level (Naderi Beni et al. 2014). Contrasting dynamics, however, have been documented at other sites, suggesting a decrease in effective moisture (Jones et al. 2006; Flohr et al. 2017) that may have induced lower lake level. Finally, coherence between many proxy records from



the Mediterranean and Southwest Asia, with respect to a general drying trend from the middle to late Holocene, may be explained by a weakening of the winter storm tracks that originate from the Mediterranean and North Atlantic, even if regional-scale experiments suggest a much more complex circulation pattern (Brayshaw et al. 2011) caused by the orography in this region (Zarrin et al. 2011).

Equivocal palaeohydrological interpretations of lake records in the Zagros Massif: how strong was the influence of humans?

Comparisons with geomorphic and speleothem archives suggest that hydrological changes at Lake Maharlou were driven by late Holocene climatic changes that were manifest over much of southwest Asia. Comparison of the record presented here with those from three other lakes in the Zagros Massif [Lake Zeribar (Fig. 5d), Lake Mirabad (Fig. 5e) and Lake Parishan (Fig. 5f)], however, reveals complexities. During the last 4000 years, the lowest  $\delta^{18}\text{O}$  values measured on endogenic calcite at Mirabad and Zeribar occurred during periods of high terrigenous supply at Lake Maharlou (Fig. 5). Then at 2000 cal year BP, increasing  $\delta^{18}\text{O}$  values corresponded to an increase of gypsum abundance in Lake Maharlou. These general trends suggest a common forcing factor affected lake-water balances of the Zagros region. It has been proposed, however, that  $\delta^{18}\text{O}$  values are not controlled entirely by hydrologic balance (Jones and Roberts 2008), and that the seasonality of precipitation should be considered (Snyder et al. 2001; Stevens et al. 2001, 2006). In Lake Parishan, Jones et al. (2015) did not find any common climatic explanation for the increase of the  $\delta^{18}\text{O}$  with changes in ostracod assemblages, and variations of other geochemical and biological indicators occurring at 2000 cal year BP. These authors invoked large-scale anthropogenic disturbances of the water cycle in the watershed to explain variations in the lake hydro-geochemistry.

With respect to Lake Maharlou, anthropogenic activities caused or simply amplified the decrease in both fluvial activity and lake volume. Notably, the HCB frequency and K intensities diverge after 2000 cal year BP, suggesting a modification of the relationship between the flood frequency and the sediment supply. This pattern could be explained by ecological and geomorphic factors in the face of

climate variability, possibly influenced by anthropogenic activities (Brisset et al. 2017). For example, lower detrital supply with relatively frequent flood events may suggest that the erosive effect of precipitation and runoff was buffered by vegetation cover in the lake catchment, a change in the ratio of discharge to sediment load (Brisset et al. 2014), development of riparian vegetation along the river banks and the delta, or by sediment mining for urban development (Rinaldi et al. 2005).

A 4800-year pollen diagram from another core taken in Lake Maharlou (Djamali et al. 2009b, 2010a, b) shows that vegetation cover was generally dominated by *Quercus* and *Pistacia*, scrub species typical of the Zagros Mountains. A marked increase in *Juglans*, *Platanus*, and *Olea* in that core, however, attests to intensification of cultivation practices and urban development, both of which were favoured by advances in irrigation techniques. Temporal comparisons with this core are not possible because radiocarbon ages were obtained on bulk carbonate (Djamali et al. 2009b) and variable reservoir effect prevents reconciliation with dates on macrofossils. Effective management of water resources has been a common way to cope with arid conditions in Southwest Asia (Magee 2005). Archaeological findings indicate that subterranean tunnels (i.e. *qanat* or *kariz* systems) were used to drain groundwater from foothills to irrigate plains since the 1st and possibly 2nd millennium BCE (Al Tikriti 2002; Magee 2005; Fattahi 2015). Because of the strong concentration of precipitation in winter/spring seasons throughout most of the Iranian Plateau, the introduction of *qanat* technology provided access to perennial water sources, modifying people's ability to subsist within this region (Manuel et al. 2017). Estimates of the number of functioning *qanats* in Iran range from 22,000 to 36,000 (Semsar Yazdi and Askarzadeh 2007). The length of some *qanats* may extend over several hundred km, with most ranging from 5 to 10 km (Wilkinson et al. 2012). The *qanat* irrigation network located in the watershed of Lake Maharlou is very dense (Fig. 1b). Even if a large number of those *qanats* are currently abandoned, their high abundance argues for strong disturbances in the drainage basin of the lake through changes in groundwater input to the lake, which in turn would accelerate lake drying. The importance given to the management of the water supply has not been the same throughout the different

empires that ruled in Persia ('Persis') (Fig. 5g). According to Manuel et al. (2017), little attention was paid to water system sustainability during the Seleucid and Parthian Empires (2197–1726 cal year BP), whereas during the reign of the Sasanians (1726–1300 cal year BP), the state supported the development and maintenance of *qanats*, but also surface water networks. Thus, the turning points in Lake Maharlou hydrology coincide with these imperial periods: the first major reduction of the water body (gypsum abundance, period B1 in Fig. 5) occurred at 2160–1980 cal year BP ( $2\sigma$ ). Then, the period of more frequent, nearly complete drying of the lake (CCB bed frequency, period B2 in Fig. 5), occurred during the Islamic period, while *qanats* were still utilised (Manuel et al. 2017). The long-term utilisation of the water resources, over millennia, may thus have been an important factor in Lake Maharlou drying during the past, similar to the situation at neighbouring Lake Parishan (Jones et al. 2015).

Finally, temporal synchronicity of climate proxies and intensity of human activities suggest that both climate and human drivers influenced Lake Maharlou hydrological changes over the last four millennia. Even if it remains difficult to decipher the relative role of each forcing factor on the lake's evolution, a probable and parsimonious interpretation might be that large-scale climate changes may have induced a water deficit that was further aggravated by alteration of the water cycle by anthropogenic activities. Thus, in the face of the global change and an increasing need for fresh water for agriculture and urban use, particular attention should be given to sustainable water resource management. Historical instrumental records, combined with secular to millennial-scale records of hydrological changes, provide helpful information to understand the function of water resources.

## Conclusions

Analyses carried out on core MAH-B from Lake Maharlou reveal four past hydro-environmental states: (1) near or total desiccation of the lake (frequency of halite beds), (2) an intermediate phase of lake water body retraction (gypsum abundance), (3) high terrigenous sediment supply from rivers (potassium abundance), and (4) river flood occurrence (frequency of the individual beds of homogeneous clays). Of these

four states, two major periods were identified over the last 3800 years. From 3800 to 2000 cal year BP, the terrigenous supply was high (notably in the intervals 3700–3650, 3500–3450, and 3400–3250 cal year BP), and evaporative conditions of the lake surface were relatively low. From 2000 cal year BP to Present, an increase in gypsum attests to more evaporative conditions, concomitant with a decrease in fluvial activity. The contraction of the water body reached a maximum from 1100 to 700 cal year BP, during which the lake frequently dried.

The comparison of the Lake Maharlou records to geomorphic activity and to climatically driven regional records suggests that a common climate forcing controlled late Holocene hydrology in southwest Asia. Comparison with the lake records located in the Zagros Massif is more challenging because these records yielded equivocal conclusions about hydrological changes inferred from different climate proxies. Indeed, ancient human impact on local and regional hydrologic balance is likely to have impacted lake systems for millennia. Subterranean tunnels (*qanats*) are an important component of the Maharlou watershed. Periods of construction of new *qanats* and repair (Sassanid and Islamic periods) seem to have corresponded to the dramatic contraction of the water body, suggesting that human disturbances of the water cycle in the lake watershed played a role in the lake water deficit, along with climate change. As both periods of climatic change and human impact correspond to hydrological variability of Lake Maharlou, we suggest that weaker winter storm tracks originating from the Mediterranean may have induced water shortage that was aggravated by human activity, leading to more frequent drying of Lake Maharlou. In turn, scarcity of the water resource has had a severe impact on food security, as famines have coincided with exceptionally frequent lake drying during the last millennium.

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