## ORIGINAL PAPER

# Effect of grazing intensities on the activity and community structure of methane-oxidizing bacteria of grassland soil in Inner Mongolia

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Abstract The effects of different grazing intensities on in situ methane flux and the structure and diversity of the methanotrophic community are measured in the typical grassland of Inner Mongolia. Four grazing intensity sites founded in 1989, control (CK), lowintensity grazing (LG), middle-intensity grazing (MG) and heavy-intensity grazing (HG), were selected. Group-specific PCR-DGGE (polymerase chain reaction-denaturing gradient gel electrophoresis) of 16S rRNA genes for the type I and type II methanotrophs was used to characterize the composition of the methanotrophic community. DGGE patterns were further analyzed using the method of the Shannonwiener index H and non-metric multi-dimensional scaling (MDS). The results showed that there were no significant differences in methane flux among different sites, yet methanotrophic communities showed significant differences. MDS analysis showed that type I

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Y.-F. Wang (⊠) College of Resources and Environment, Graduate University of Chinese Academy of Sciences, Beijing 100049, China e-mail: yfwang@gucas.ac.cn methanotroph community composition at the CK site were significantly different from the three other sites. For type II methanotrophic community composition, it was similar between CK and HG site, and between LG and MG site, while that at the former two sites were significantly different from latter two ones. Additionally Shannon indices of type II methanotrophs were higher at the LG and MG sites than two other sites. Though grazing intensities had an impact on the structure of the methanotrophic community, management-induced changes in the structure of methanotrophic community did not reflect methane consumption capacity across sites. These results suggest that methane consumption is a complex process in soil, and we should be cautious when speculating on the change of methane consumption rates based on a change of methanotrophic community structure.

## Introduction

Atmospheric methane is one of the most important greenhouse gases and contributes approximately 20% of global warming (IPCC 2001). The most important natural source of methane is natural wetlands where methanogenesis occurs (Wuebbles and Hayhoe 2002). Much of this methane (700 Tg year<sup>-1</sup>) is oxidized by aerobic methane-oxidizing bacteria (methanotrophs) in the oxidic cover layer of these anaerobic soils

before it is released into the atmosphere. Aerobic soils also oxidize a significant amount (20–60 Tg year<sup>-1</sup>) of the atmospheric methane (Hütsch 2001).

Methanotrophs are obligately aerobic, gram-negative bacteria that utilize methane as their sole sources of carbon and energy. They can be divided into two distinct physiological group, type I and type II, based on their phylogenetic affiliation, intracellular membrane arrangement, carbon assimilation pathway and phospholipid fatty acid (PLFA) composition (Hanson and Hanson 1996; Bowman 2000). Type I is a kind of low-affinity methanotroph that can only grow in the high methane concentration and low oxygen concentration. Type II methanotrophs are high-affinity organisms that can grow in the low concentration of atmospheric methane (Bender and Conrad 1992). Notably, cultivated isolates of methanotroph were not able to grow on the low mixing ratio of atmospheric methane (Benstead et al. 1998). A new independent culture technique, however, proved that a kind of methanotroph exists in soil which could be the predominant methanotrophs consuming atmospheric methane (Roslev and Iversen 1999; Bull et al. 2000; Crossman et al. 2005).

The typical grassland of Inner Mongolia is one of the representative parts of Eurasia grasslands and is regarded as an important sink for atmospheric methane  $(1.8-2.3 \text{ kg C ha}^{-1} \text{ year}^{-1})$  (Wang et al. 2005). There are about  $3.13 \times 10^8$  km<sup>2</sup> of grassland in China. However, these grasslands have undergone considerable change over the last decades because of grassland management. This has led to more intensive grazing regimes that may eventually result in grassland degradation (Chen and Wang 2000). Effects of grazing intensities on the microbial communities have been conventionally studied by techniques based on the cultivation of the soil microorganisms or by measuring their metabolic activities, e.g., by studying the microbial biomass, soil respiration, dehydrogenase, urease, proteinase, and transferase activities (Liu and Liao 1997). In the past decade a new strategy for the PCR-DGGE method has been widely used to detect the temporal and spatial variation of general bacteria and functional groups. The method has successfully described the structure and diversity of microbial community in waster water sludge (Boon et al. 2002), grassland soil (Clegg 2006), and agricultural soils (Seghers et al. 2003, 2004).

The goal of this study was to evaluate the effects of different grazing intensities on the behavior and structure of methanotrophic community with specificprimers of 16S ribosomal RNA (rRNA) and denaturing gel gradient electrophoresis (DGGE). We combined a kinetic description of the methane consumption capacity of these experimental treatments with a molecular analysis of the methanotrophic populations involved. We hypothesize that management may induce a change in the structure and diversity of the methanotrophic community, which may in turn lead to a change in methane consumption capacity of soils.

#### Materials and methods

Site description and soil sampling

The experimental site,  $600 \text{ m} \times 300 \text{ m}$ , is located in the Xilin River Basin (43°32' N, 116°40' E, 1200 m a.s.l) with temperate terrestrial monsoon climate. The site is 300 m north of the bank of the Xilin River, and the soil type is a sandy calcareous chestnut. The plant community is dominated by Artemisia frigida, Cleistogenes squarrosa and Carex spp. This plant population is regarded as one of degradation stages of the Leymus chinensis and Stipa grandis community due to long-term over-grazing. Soil organic C in the surface layer is 0.5%. According to the long-term records at the Inner Mongolia Grassland Ecosystem Research Station, the temperate semi-arid steppe climate shows a dry spring and a moist summer. Average annual air temperature is -0.4°C, and average air temperature in January and July are -23°C and 17.9°C, respectively. Precipitation is ca. 350 mm/year, which is mainly falling in June, July and August.

Grazing management experimental plots were established in 1989 (Chen and Wang 2000). There are six stocking rates (0.00, 1.33, 2.33, 4.00, 5.33, 6.67 sheep ha<sup>-1</sup>) with three different replicates per treatment except for the 0.00 sheep ha<sup>-1</sup> grazing intensity (that is, control site). Every site area is 100 m × 100 m, and all 16 experimental plots were set up by a randomized block design. Four different stocking rates (0.00, 1.33, 4.00, 6.67 sheep ha<sup>-1</sup>) were selected, which were designated as control (CK), low-intensity grazing (LG), middle-intensity grazing (MG) and heavy-intensity grazing (HG).

Every year, grazing began on May 20 and terminated on October 5. Each plot was rotationally grazed three times per year, each time for 15 days with a rotation interval of 30 days, yielding a total grazing period of 45 days per plot. This rotational grazing facilitated the restoration succession of the degraded *A. frigida* grassland. Soil characteristics under different grazing intensities are listed in Table 1 (Chen and Wang 2000).

Soil samples were taken from a depth of 0 to 5 cm with a 3-cm diameter auger in August 2004. From each study plot, 20 soil cores were collected to form one bulk sample, and soil physical properties were measured at the same time. Other samples were transported to the laboratory and stored at  $-20^{\circ}$ C for molecular analysis after sieving (mesh size <2 mm) to remove plant material.

## CH<sub>4</sub> consumption analyses

Methane consumption rates were measured in situ in 1998 and 1999 as well as in August 2004 and in July 2005 using a closed static chamber method (Wang and Wang 2003). The chambers were equipped with a battery-driven ventilator, and a small vent for pressure equilibration was used for gas flux measurements. Three replicate chambers,  $0.5 \text{ m} \times 0.5 \text{ m} \times 0.5 \text{ m}$ , were used in every plot. During measurement, the chambers were 10-15 m from each other. Gas samples were taken from the chambers with 10-min intervals in 40 min after closure using 10-ml gas tight syringes. All gas samples were analyzed within 6 h with gas chromatograph (HP 5890 II) in the laboratory. To calculate the CH<sub>4</sub> flux rate, liner regressions were performed on the four gas samples collected from each flux chamber. Finally, the CH<sub>4</sub> flux rate was adjusted using air pressure and air temperate within the chamber headspace (Wang et al. 2005).

Extraction and PCR amplification of total DNA from soil

Nucleic acids were extracted from 0.5 g of each sample by using the procedures described by Zhou et al. (2007). In brief, soil was placed in a 2-ml screw-cap tube containing a mixture of ceramic and silica particles (Bio101, Carlsbad, Calif.); the mixture was homogenized for 30 s in a FastPrep bead beater cell disrupter (Bio101). After the nucleic acids had been precipitated and washed twice in 75% (vol/vol) ethanol, the final DNA was re-suspended in 100  $\mu$ l TE buffer. Furthermore, the crude extract was purified with Qiagen gel extraction Kit (Qiagen, Inc.). DNA concentrations were measured by Biospec-mimi (Shimadzu, Kyoto, Japan).

All the primers and PCR conditions are listed in Table 2. For more information about these primers and PCR conditions, we refer readers to the original papers. In order to increase the sensitivity and to facilitate the DGGE by analyzing fragments of the same length, a nested PCR technique was applied. In the first round different group-specific primers and one set of bacterial primers were used, each with their own corresponding PCR protocol. During the second PCR round, the obtained fragments were re-amplified by using the bacterial primers P338F and P518r in the same PCR protocol (Table 2).

The final concentration of different components in the mixture was as follows: 0.4  $\mu$ M of each primer, 200  $\mu$ M of each deoxynucleoside triphosphate, 1.5 mM MgCl<sub>2</sub>, 1× thermophilic DNA polymerase 10× reaction buffer (MgCl<sub>2</sub>-free), 1.25 U per 50  $\mu$ l of

Table 1 Soil characteristics in different stocking rates

|      |                 |                           | e                           |                            |                 |                                  |  |
|------|-----------------|---------------------------|-----------------------------|----------------------------|-----------------|----------------------------------|--|
| Site | pН              | TOC (g kg <sup>-1</sup> ) | $NH_4^+$ – $N (mg kg^{-1})$ | $NO_{3}^{-}N (mg kg^{-1})$ | Total P (%)     | Soil bulk density (g $cm^{-3}$ ) |  |
| СК   | $6.92 \pm 0.02$ | $1.79 \pm 0.07a$          | $4.09 \pm 0.72a$            | $3.06 \pm 0.43a$           | $0.14 \pm 0.02$ | $1.16 \pm 0.03b$                 |  |
| LG   | $7.14 \pm 0.01$ | $1.47 \pm 0.08b$          | $3.00 \pm 0.43a$            | $1.39 \pm 0.37b$           | $0.13 \pm 0.02$ | $1.31 \pm 0.03a$                 |  |
| MG   | $7.07 \pm 0.02$ | $1.47 \pm 0.05b$          | $3.47 \pm 0.71a$            | $1.25 \pm 0.18b$           | $0.13 \pm 0.03$ | $1.31 \pm 0.03a$                 |  |
| HG   | $7.13 \pm 0.02$ | $1.56 \pm 0.07b$          | $3.44 \pm 0.65a$            | $1.00 \pm 0.14b$           | $0.1 \pm 0.03$  | $1.33 \pm 0.03a$                 |  |

CK, control; LG, low-intensity grazing; MG, middle-intensity grazing; HG, heavy-intensity grazing Different letters showed significant differences (P < 0.05)

| Target                | Primers                    | PCR conditions    |              |     |           |     |            | References |                                  |
|-----------------------|----------------------------|-------------------|--------------|-----|-----------|-----|------------|------------|----------------------------------|
|                       |                            | Numbers of cycles | Denaturation |     | Annealing |     | Elongation |            |                                  |
|                       |                            |                   | °C           | min | °C        | min | °C         | min        |                                  |
| First PCR round       |                            |                   |              |     |           |     |            |            |                                  |
| Type I methantrophs   | methT1dF, methT1bR         | 30                | 95           | 1   | 55        | 1   | 72         | 2          | (Wise et al. 1999)               |
| Type II methanotrophs | 27f, methT2R               | 30                | 95           | 1   | 55        | 1   | 72         | 2          | (Lane 1991;<br>Wise et al. 1999) |
| Second PCR            |                            |                   |              |     |           |     |            |            |                                  |
| Bacteria              | P338F <sup>a</sup> , P518r | 30                | 94           | 1   | 56        | 1   | 72         | 1          |                                  |

 Table 2
 PCR primers used in this study

Before the run of cycles, the temperature was held at 95°C for 7 min, and after each run the temperature was kept 72°C for 10 min for final template elongation

<sup>a</sup> A 5'GC-clamp was added for DGGE analysis (Muyzer et al. 1993)

Taq DNA polymerase (Promega, Madison, WI), and DNAse and RNAse free filter sterilized water (LiYuan Apparatus Com., Beijing, China). All reactions were performed in a PTC-200 thermal cycler (MJ Research Co., USA). During the first PCR round, ~50 ng of purified DNA was added to 24  $\mu$ l of PCR master mix, and in the second PCR round, 1  $\mu$ l of amplified product from the first round was added to 24  $\mu$ l of PCR mixture. After each PCR, the size of the amplification product was verified on a 1.2% agarose gel.

## DGGE

DGGE analysis was performed using a Dcode system (Bio-Rad, Hercules, CA) as described previously (Muyzer et al. 1993). For the generation of bacterial community profiles, polyacrylamide gels (8% acrylamide, 0.5× TAE (20 mM Tris, 20 mM acetate, 1 mM EDTA [pH 8.0]), 37:1 acrylamide:bisacrylamide) were cast using 40-60% denaturant [100% denaturant was defined as 7 M urea with 40% (vol/ vol) formamide]. Electrophoresis was run overnight for 12 h at a constant temperature 60°C at 110 V. After electrophoresis, the gels were soaked for 30 min in SYBR green II nucleic acid gel stain (1:10,000 dilution; FMC Bioproducts, Rockland, ME). The stained gels were immediately photographed on UV transillumination table with a Hewlett Packard Scanjet 5370C. The DGGE patterns were determined using the Labworks software (Labworks TM software version 4.0. UVP, UK).

Cloning and sequencing analysis

The specific methanotrophic gene fragments of the first PCR round were cloned into the pGEM-T Easy vector (Promega) according to the manufacture's instructions. Plasmid DNAs were isolated from randomly selected clones and screened for inserts of the expected sizes. In addition, 1 band from the type II methanotrophs pattern was excised from the gel and reamplified with the same reaction procedure. The PCR products were purified and then were cloned into plasmids with a pGEM-T easy vector kit (Promega) as well. The plasmid DNA was sequenced with primers (T7 reverse/forward) on an ABI PRISM 3700 DNA analyzer (ABI Biosystems, Inc.). The nucleotide sequence obtained in this study was sent to the NCBI database. The accession numbers is DQ420621.

# Statistics

The means and standard errors of the mean (SEM) of the methanotroph analytical results are presented per dry matter weight (dry wt) of soil. Analysis of variance (ANOVA) followed by Tukey's test was performed to test for within-treatment differences in methane consumption rates in different grazing intensities. Significant differences are reported at the 0.05 level. The DGGE gel photographs were screened for their presence (1) or absence (0) of bands in the DGGE patterns. A similarity matrix was generated with the binary data using the Jaccard coefficient and analyzed with non-metric multidimensional scaling (MDS) as described by Jaatinen et al. (2004).

The DGGE patterns were analyzed with the labworks software (Labworks TM software version 4.0. UVP, UK). Band intensity of each lane in the densitometric curve was read. The diversity of the microbial community was examined by the Shannon index of general diversity H (Sun et al. 1993). H was calculated on the basis of the band intensity on the gel tracks, which was reflected as peak heights in the densitometric curve. The equation for the Shannon index is:

$$H = -\sum (n_i/N)\log(n_i/N)$$

where  $n_i$  is the height of the peak and N the sum of all peak heights of the densitometric curve.

## Results

Methane consumption rates

Methane fluxes at different stocking rates are shown in Fig. 1. The methane consumption rates at all sites were negative, which showed Inner Mongolia grassland as a sink for atmospheric methane, consistent with previous studies in this area (Wang et al. 2005). The large variations of methane uptake flux observed at one site at the different sampling times confirm flux rates observed by Wang et al. (2005). In total, methane uptake rates were slightly higher on 9/8/98, 7/1/99, 8/1/99 and 9/1/99 than at the other sampling times (Fig. 1). The mean annual precipitation in 1998, 1999, 2004 and 2005 was 357, 214, 364 and 162 mm, respectively. Methane uptake flux at one site at one sampling time also showed large variations (error bars, Fig. 1). When the data of different sampling time at the same site were averaged, however, no significant differences were observed among different grazing sites. Statistical analysis of all methane fluxes at the different sampling times showed that the averages of methane flux in CK, LG, MG and HG sites were 75, 83, 85 and 75  $\mu$ g C m<sup>-2</sup> h<sup>-1</sup>, respectively.



Fig. 1 The methane consumption rates in the different grazing intensity plots. CK, control; LG, low-intensity grazing; MG, middle-intensity grazing; HG, heavy-intensity grazing, different letters showed significant differences (P < 0.05)

### DGGE analysis

Before PCR-DGGE analysis, DNA yields after extraction were examined, and no significant differences were found between the different samples. Figure 2 shows DGGE patterns of the type I methanotrophs and type II methanotrophs. After further MDS analysis of these patterns, the structure of the type I methanotroph community in the grazing treatments was significantly different from site CK (Fig. 3a). No significant differences were observed between the other grazing intensity treatments. Grazing management greatly affected the structure of the type II methanotrophs community as well. Type II methanotroph communities at the LG and MG sites were significantly different from the other two sites (Fig. 3b). Additionally, the Shannon indices (Table 3) indicated that type I and type II methanotrophs communities at the LG and MG sites were higher than at the CK and HG sites.

Each band in the DGGE pattern was assumed to represent one unique ribotype. In the pattern of type II methanotrophs, a phylotype (band) was only detected in LG and MG soil. To determine more information, the band was further cloned and sequenced. After blasting the sequence in the Gen-Bank, its sequence exhibited 100% similarity to an *uncultured methylobacteriaceae* bacterium clone GASP-WA2W2\_D12 (EF072931), which was recently detected in America (unpublished).



**Fig. 2** DGGE patterns of type I methantrophs (**a**) and type II methanotrophs (**b**) in the different grazing intensity treatments in three parallel replicates. CK, control; LG, low-intensity grazing; MG, middle-intensity grazing; HG, heavy-intensity grazing

## Discussion

In this study, we examined the effects of different grazing intensities on methane consumption capacity and methanotrophic community structure. The soilatmosphere exchange of methane comprises two parts: atmospheric methane oxidation by soil and methane produced by the soil methanogenesis (Wager et al. 1996). Methane production in the Inner Mongolia steppe can be negligible for the amount of methane produced by methanogenesis is very little in the aerated soil in semi-arid grassland. The mean methane uptake flux observed agrees with previous studies in this area (Wang et al. 2005). These uptake rates are slightly higher than those reported in the Colorado shortgrass steppe by Mosier et al.



**Fig. 3** Non-metric multi-dimensional scaling (MDS) analysis DGGE patterns of type I methanotrophs (**a**) and type II methanotrophs (**b**). CK, control; LG, low-intensity grazing; MG, middle-intensity grazing; HG, heavy-intensity grazing

(25–45  $\mu$ g C m<sup>-2</sup> h<sup>-1</sup>) (1991, 1996) and in an English grassland (10–19  $\mu$ g C m<sup>-2</sup> h<sup>-1</sup>) (Kammann et al. 2001).

Compared with type I methanotrophs, the type II methanotrophs are likely to be the main group in the grassland soil in Inner Mongolia. These organisms oxidize atmospheric methane at low concentration.

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Table 3 Shannon diversity (H) and evenness (E) of different treatments

|                      | Treatment CK     | Treatment LG     | Treatment MG     | Treatment HG     |
|----------------------|------------------|------------------|------------------|------------------|
| Type I methanotroph  | $0.98 \pm 0.05b$ | $1.03 \pm 0.02a$ | $1.06 \pm 0.09a$ | $0.97 \pm 0.03b$ |
| Type II methanotroph | $1.13 \pm 0.04b$ | $1.26 \pm 0.02a$ | $1.21 \pm 0.03a$ | $1.09 \pm 0.02b$ |

CK, control; LG, low-intensity grazing; MG, middle-intensity grazing; HG, heavy-intensity grazing Different letters showed significant differences (P < 0.05)

Compared with the CK and HG sites, methane uptake at the LG and MG sites was slightly higher. The type II methanotrophic phylotype was obvious at the LG and MG sites and disappeared at the CK and HG sites. This implies that the phylotype seems to play an important role in oxidizing atmospheric methane in this area. The further sequence analysis indicated that it is a kind of *uncultured Methylobacteriaceae* bacterium.

Previous studies indicated that long-term agricultural practices may shift the structure of the methanotrophic community, which results in the change of methane oxidation (Adamsen and King 1993; Seghers et al. 2003). Are the changes in methane oxidation related with the shifting of structure and diversity of methanotrophic community induced by management practices? The research of Jaatinen et al. (2004) indicated that fire significantly increased the methane oxidation rate, but the community structure did not change through comparing pmoA recovery products. It is possible that shifts occurred on a fine phylogenetic scale not distinguishable by pmoA products. However, Seghers et al. (2003) found long-term use of herbicides atrazine and metolachlor resulted in shifting of methanotrophic community structure, but no change in abundance and activity of methanotrophs. In this study, comparing the change of methane uptake flux and the change of methantroph communities structure at different stocking rates sites, we found that although grazing intensity had an impact on the methanotrophic community structure, it did not greatly impact the methanotrophic activity. As one of the primary land uses in Inner Mongolia, the practice of grazing sheep is a complex event. A range of specific factors, such as fecal and urine deposition, shifts in plant rhizosphere exudation, shifts in plant composition and the change in soil structure and aerobicity, changes soil characteristics (Clegg 2006). Methanotroph acitivity in soil is affected by many factors, one of which is ammonium concentration in soil (Bodelier and Laanbroek 2004). The NH<sup>+</sup><sub>4</sub>–N concentrations at the different sites were not significantly different, which may have helped to account for no difference in methane consumption rates across sites (Table 1).

Recently more and more functional groups that participate in consuming atmospheric CH<sub>4</sub> have been found in terrestrial ecosystems (Bull et al. 2000; Henckel et al. 2000; Knief et al. 2003). Recent studies have indicated that atmospheric methane oxidization can also be mediated by a kind of unknown methanotroph, which is similar to type II methanotrophs in the phylogenetic tree (Bull et al. 2000; Henckel et al. 2000; Crossman et al. 2005). The type II methanotroph primers used in this study, which targeted all of the methylosinus/methylocystis groups, are degenerate primers and detect some unknown methanotrophs (Wise et al. 1999). We detected a kind of unknown type II methanotroph in the Inner Mongolian grassland soil and its role in consuming atmospheric methane needs to be further studied.

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