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Natural hazards in Australia: floods

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Abstract Floods are caused by a number of interacting factors, making it remarkably difficult to explain changes in flood hazard. This paper reviews the current understanding of historical trends and variability in flood hazard across Australia. Links between flood and rainfall trends cannot be made due to the influence of climate processes over a number of spatial and temporal scales as well as landscape changes that affect the catchment response. There are also still considerable uncertainties in future rainfall projections, particularly for sub-daily extreme rainfall events. This is in addition to the inherent uncertainty in hydrological modelling such as antecedent conditions and feedback mechanisms.

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Research questions are posed based on the current state of knowledge. These include a need for high-resolution climate modelling studies and efforts in compiling and analysing databases of sub-daily rainfall and flood records. Finally there is a need to develop modelling frameworks that can deal with the interaction between climate processes at different spatio-temporal scales, so that historical flood trends can be better explained and future flood behaviour understood.

1 Introduction

Floods are one of the most dangerous natural hazards worldwide, with thousands of people dying and hundreds of millions of dollars damage on average per event (Guha-Sapir et al. 2015). Particularly significant flood events can cost much more; for example the Queensland 2010–2011 floods caused over $AUD2 billion infrastructure damage and even larger indirect costs to the economy. To better manage flood risk in the future, it is necessary to know whether and how flood magnitude and frequency is changing. Although the annualised total cost of floods has been increasing over time in Australia (e.g., Guha-Sapir et al. 2015), it is unclear whether the trends are due to changes in reporting mechanisms, population, land use, infrastructure, and/or in the frequency and magnitude of flood causing mechanisms. The limited assessments that have been conducted directly on trends in Australian flood data have suggested that, if anything, the magnitude of floods has remained unchanged or even decreased in many parts of the country (Ishak et al. 2013). The reasons for these changes are not fully understood.

Explaining changes in flood hazard is challenging because of the interactions between meteorological and catchment conditions. Floods are primarily caused by intense rainfall events on site or upstream, but are also influenced by the location, pattern and duration of the rainfall event, the overall catchment wetness prior to the event, and the hydraulic characteristics of the catchment. Furthermore, most inhabited catchments in Australia have been anthropogenically modified.. To attribute changes in flood behaviour to one or more causes requires a deep understanding of the nature, timing and extent of these various influences. The presence of long-term natural variability in the climate system (Kiem et al. 2003), as well as anthropogenic climate change (CSIRO, Bureau of Meteorology 2015), adds to the difficulties in attributing changes in floods to any one cause or combination of several causes. Despite this, understanding future flood hazard is important to prioritise future investment in infrastructure, floodplain management practices and resources for operational flood forecasting.

This paper is part of the Special Issue on changes to natural hazards in Australia, and describes our present understanding of the causes of floods in Australia, how flood hazard has changed over time and how it is projected to change in the future. The paper also presents a number of open research questions that should be prioritised to better understand historical and future change to Australian flood hazard.

2 Understanding floods

2.1 Defining flood types and hazards

Beyond the broad definition of 'unwanted water covering otherwise dry land', there is no universal taxonomy for floods. A prominent distinction is between coastal floods caused by high tides, storm surge and strong winds, and fluvial floods caused by rainfall and possibly snowmelt (Blöschl et al.

2015). The primary driver of fluvial floods in Australia is heavy rainfall, although the rainfall itself does not need to be extreme to lead to an extreme flood impact (Leonard et al. 2014). Coastal hazards are the subject of a companion paper in this Special Issue (McInnes et al., this issue), whilst the hazards from tropical cyclones and storms are discussed in Walsh et al. (this issue).

Another important distinction in flood classification is their temporal and spatial scale.. Flash floods can occur anywhere in the landscape but are normally limited to relatively small catchments (Blöschl et al. 2015; Hapuarachchi et al. 2011). Widespread floods are generally due to prolonged rainfall or sequential large events and are generally related to larger-scale circulation patterns such as extratropical cyclones. Flooding can also be related to longer timescale variations in atmospheric and oceanic states (Kiem et al. 2003). This paper focuses on fluvial floods and explores their relationship to the climatic and meteorological mechanisms that generate such events.

Flood hazard is generally defined in terms of the magnitude and probability of occurrence. Flood magnitude is often quantified in terms of peak streamflow, but other measures include rate of flow or water level increases, flow velocity, and area or duration of inundation.. In a design context, *Australian Rainfall and Runoff* (Engineers Australia 1987) provides guidance on best-practice methods for flood estimation, which depend on the frequency of the event. In an operational setting in Australia, flood magnitude is reported as a relative flood severity (minor, moderate or major), by comparing the flood height to pre-specified thresholds at forecast locations. The focus of this review is peak streamflow magnitude because this is the most commonly available information.

2.2 Causes and mechanisms of flood hazard

When trying to understand how changes to the climate affect flood hazard, it is necessary to account for processes across a wide range of timescales depicted in Fig. 1. There are differences in the timescales of the rainfall event itself (T_p) compared to the flood hydrograph (T_Q); this difference can be particularly pronounced in the catchments west of the Great Dividing Range in Australia where flood events take the order of months to flow along the Darling River. Geological influences (Gaál et al. 2012) lead to variations in floodplain extent and nature of the flood hazard, and hence the population affected. Landforms and soil affect antecedent conditions; this is further explored in Fig. 2. Figure 1 also shows that geological timescales affect the type and occurrence of rainfall events e.g. strong orographic rainfall around Coffs Harbour on the east coast leads to a high risk of flash flooding. Shown at the top of Fig. 1 are the important global climate patterns (e.g., El Niño/Southern Oscillation (ENSO), Southern Annular Mode (SAM), Interdecadal Pacific Oscillation (IPO)) that have a strong influence on inter-annual, annual and seasonal variations in the occurrence and magnitude of floods (e.g., see Westra et al. this issue). In addition, these large-scale climate patterns can control synoptic-scale meteorological processes as reviewed in Walsh et al. (this issue).

There are significant regional variations in flood-producing meteorological processes around Australia; a review of Australian rainfall variability is provided in Risbey et al. (2009). Other mechanisms include East Coast Lows in the south-east (Callaghan and Power 2014), tropical cyclones in the north or more generally thunderstorms, the latter particularly relevant to flash flooding (McKay 2007). The challenge is how to attribute observed flood trends to either large scale mechanisms or the rainfall processes that can occur locally over very short timescales, or both.

A catchment's runoff generating mechanisms also play a major role in determining how changes to extreme rainfall will be translated into changes in flood hazard. Two important

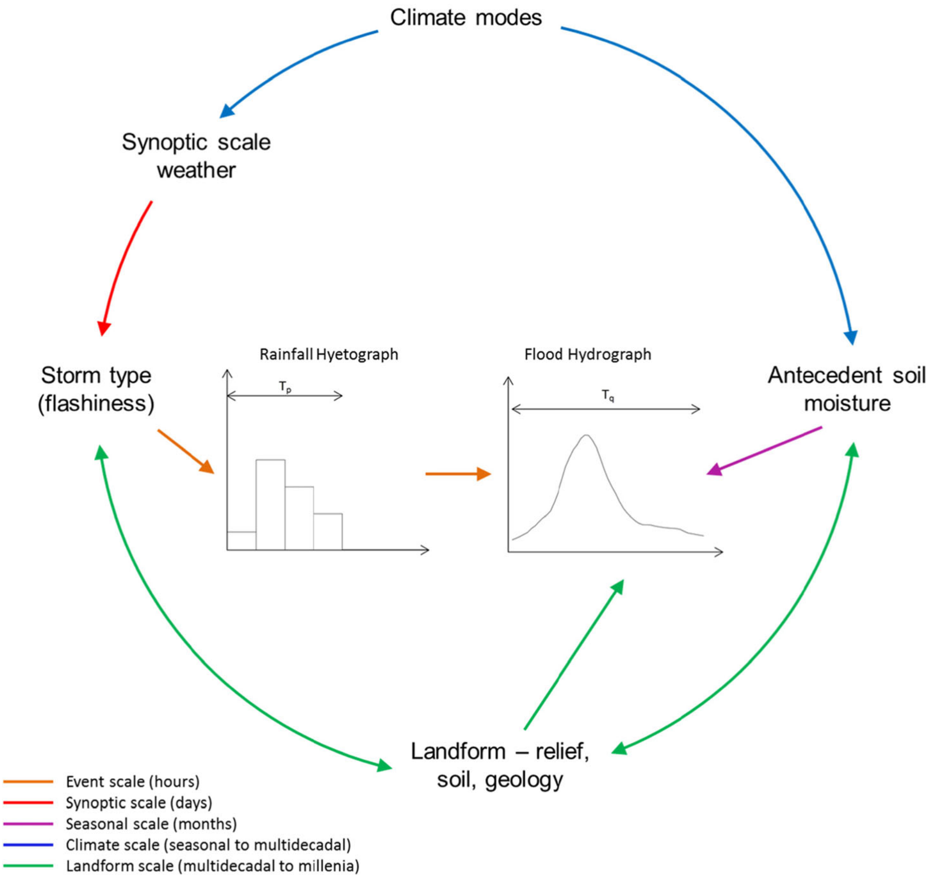

Fig. 1 Major Interactions between flood processes on different timescales (after Gaál et al. 2012). T_p is the period over which the flood producing rainfall occurs depicted here using a rainfall hyetograph and can be significantly shorter than the flood duration (T_q). Both T_p and T_q are affected by the longer timescale processes shown here at synoptic, seasonal, climate and geological scales

mechanisms — infiltration excess and saturation excess runoff — are responsible for a significant portion of flood runoff in Australian catchments (Fig. 2). The saturation excess mechanism is more common in humid catchments with relatively high water tables or wet antecedent conditions (Fig. 2a). The event rainfall is shown in Fig. 2a increases the extent of saturated land. The infiltration-excess mechanism (Fig. 2b) is likely to apply in situations where the instantaneous rainfall intensity is significantly higher than the hydraulic conductivity of the soil (Mirus and Loague 2013). Trancoso et al. (2016) classify 355 catchments in eastern Australia according to climate and runoff mechanisms and such an approach could be extended to consider flood generation as well. Both mechanisms affect the relative importance of extreme rainfall compared to the antecedent moisture conditions. As shown in Fig. 2, in both cases catchment evapotranspiration prior to the flood also has a large impact on antecedent conditions. Potential future changes to evapotranspiration are discussed in Kiem et al. (this issue). Thus, flood hazard can depend both on the longer-term water balance that determines antecedent catchment wetness and groundwater levels, as well as the intensity of the rainfall event.

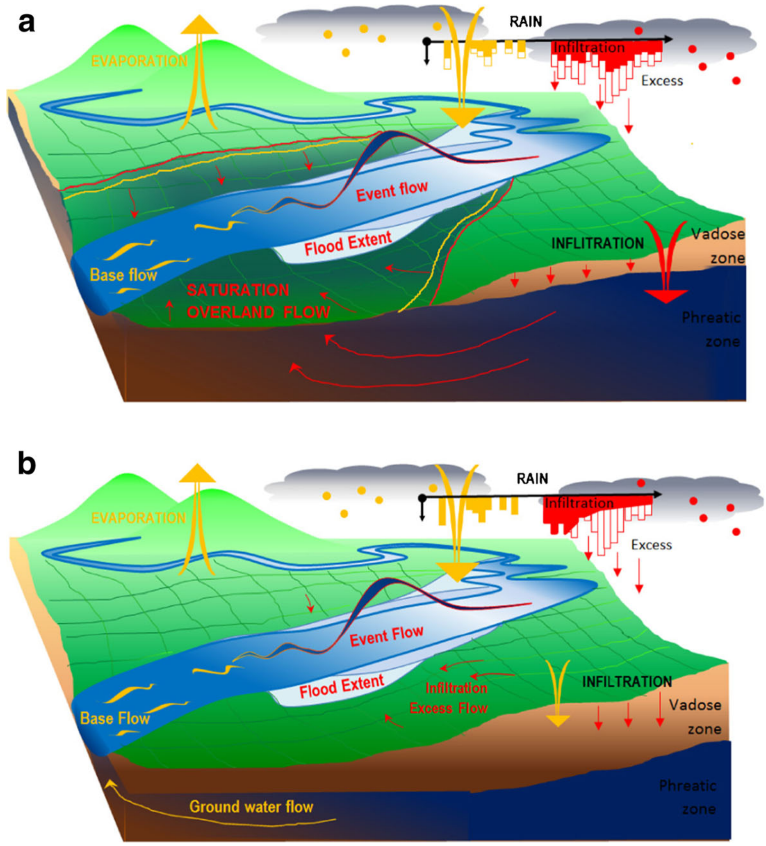

Fig. 2 Conceptual representation of runoff generation processes relating to floods. *Red* highlights water fluxes during the event, and orange highlights antecedent water fluxes over a longer time period, with arrow sizes proportional to flux magnitude. **a** Saturation excess flow where the overland flow is proportional to the area of the catchment fully saturated. A longer term water balance determines the initial degree of catchment saturation and the event determines the growth in the saturated area. **b** Infiltration excess where the initial degree of infiltration depends on antecedent rainfall. The overland flow is proportional to excess rainfall and infiltration diminishes during the event

2.3 Modelling floods

Due to the complexity of flood-generating processes and data limitations it is necessary to make simplifying assumptions when estimating flood hazard Where streamflow data are available, a flood frequency analysis (which links the flood magnitude and its probability of occurrence) is often considered to be preferable to rainfall-driven flood modelling approaches (Engineers Engineers Australia 1987). Unfortunately this method does not explicitly represent the connection between the causative variables and flood hazard, preventing extrapolation under changing catchment and/or climate conditions.

When streamflow observations are not available the flood hazard must be estimated using rainfall-runoff models. A common modelling approach uses design rainfalls (i.e., rainfall estimates from Intensity-Frequency-Duration relationships) in a hydrologic model of the catchment (Engineers Engineers Australia 1987). The most popular design flood models used in Australia tend to be used with a simple loss model similar to the infiltration-excess mechanism, although as shown in Fig. 2 the saturation excess mechanism controls runoff in many catchments. The resulting hydrographs are used directly to estimate flood hazard or in a hydraulic model to estimate velocities and inundation extents. Alternatively, a continuous simulation approach can be adopted using long time series of observed or synthetic rainfall. Antecedent conditions are thus modelled explicitly, in contrast to design events where empirical losses represent catchment wetness. The paucity of sub-daily rainfall observations has hampered widespread implementation of continuous simulation. In both the design event and continuous simulation methods, the flood-producing mechanisms for individual catchments are rarely explicitly taken into account, which is particularly important when modelling future climate settings where the relative importance of individual mechanisms may change. The extent to which the models can be extrapolated under non-stationary climatic forcings therefore remains untested.

3 Historical changes to floods and causative variables

3.1 Historical changes to floods

The most comprehensive study on historical trends in floods in Australia to date (Ishak et al. 2013) considered annual maximum series (AMS) of instantaneous streamflow data from 330 catchments across the country with minimal regulation or land cover change. The directions of trends in the AMS were found to be mixed (i.e., both decreasing and increasing), although only 30 % of stations had significant trends (10 % level). Many more stations had decreasing trends rather than increasing ones. Most of the decreasing trends were found along the eastern and south-eastern coasts, although it is important to note that this is also where the vast majority of stations used in the analyses are located. Ishak et al. (2013) also restricted the analysis period to exclude the multi-year Millennium Drought (2001–2009), which reduced by about half the number of locations with significant decreasing trends. Some of the significant trends were also removed when the statistical tests were conditioned on SAM, ENSO and IPO, highlighting the influence of large-scale climate variability on extremes. However these indices themselves may have some covariation with long-term warming trends.

Given that the average record length used by Ishak et al. (2013) was only 38 years, this demonstrates the difficulty in separating non-stationarity related to anthropogenic climate change from that associated with natural low frequency climate oscillations. Although there was considerable quality control on the AMS database, the influence of non-climatic causes of non-stationarity, such as catchment modification or regulation, cannot be completely ruled out from the Ishak et al. (2013) analyses. In addition to this national study, some regional scale assessments have also been carried out. For example, statistical modelling of rainfall variables and satellite imagery suggests that summer flooding in the large Upper Fortescue basin in north-western Australia has been particularly severe and remarkably sustained between 1999 and 2006 when compared to the previous century (Rouillard et al. 2015).

There has been more research into inter-annual and multi-decadal variability than trend analyses for floods in Australia. A recent global analysis of floods (Ward et al. 2014) shows that flood volumes are significantly higher than average in the north-west and arid inland Australia during La Niña years and lower during El Niño. The variability in streamflow and flooding on interannual timescales related to ENSO is supported by similar observations from eastern Australia (Chiew et al.

1998; Kiem et al. 2003; Power et al. 1999; Verdon et al. 2004), and decadal/multi-decadal timescales related to IPO or low frequency Pacific sea surface temperature variability (Kiem et al. 2003; Kiem and Verdon-Kidd 2013; Micevski et al. 2006; Power et al. 1999).

The high variability in Australia's climate, coupled with the low population density, means that high-quality, long-term streamflow gauge records are relatively rare. The value of maintaining these networks in the face of funding pressure has been demonstrated by Cordery (2003). The inability to separate long-term trends from climate mode variability is directly attributable to the relatively short streamflow records and, despite many advances in the last decade, still quite limited understanding into what role the climate modes play in relation to flood hazard (and also a lack of long-term information about the climate modes). Modelling studies may be useful to supplement this gap given that rainfall records are generally longer and provide better coverage of the continent. Examples include CSIRO's Sustainable Yields projects[1] which have covered large parts of Australia. However, the focus of these projects was on long-term yield rather than floods and the opportunity to use the modelling to examine trends and variability over the historical period was not fully capitalised, pointing at a promising opportunity.

3.2 Historical changes to extreme rainfall

Increases in the proportion of heavy rainfall have been detected across Australia with heavy daily rainfall accounting for an increased proportion of total annual rainfall since the 1970s (CSIRO and Bureau of Meteorology 2015). Regionally, however, significant variability exists. Gallant et al. (2007) found that the proportion of total rainfall stemming from extreme events has increased since the 1950s along the eastern coast and in south-western Australia, and since the 1970s in the south-east.

Observed trends in extreme rainfall across much of Australia have been consistent with mean rainfall trends, although trends in the extremes have typically been greater than the mean trend (Alexander et al. 2007); that is, where average rainfall is increasing, the extremes have tended to increase at a faster rate. The north-west of Australia has been relatively wet during the late 20th century, with increases in rainfall and the frequency of extreme rainfall events over the monsoonal and sub-tropical north-west, particularly during summer (Evans et al. 2014; Shi et al. 2008). Other studies have considered non-stationarity in extreme rainfalls for single sites (Jakob et al. 2011a; Yilmaz and Perera 2013) or small collections of stations (Jakob et al. 2011b; Laz et al. 2014). In general these studies have found that rainfall extremes for short duration events tend to have increasing trends. These increasing trends are not as common for longer duration events (Westra and Sisson 2011).

These trends are generally unexplained. One approach that may be useful is that of Hardwick-Jones et al. (2010) who examined the relationship between observed temperature and extreme precipitation using historical precipitation records. This approach could be extended to project future changes although there are still some uncertainties with respect to the role that moisture availability plays in these relationships.

3.3 Attribution of changes in floods

Attribution is the process of identifying the reasons for significant changes in a climatic or hydrologic variable. This is a relatively new area of interest in hydrology and Merz et al. (2012)

[1] http://www.csiro.au/en/Research/LWF/Areas/Water-resources/Assessing-water-resources/Sustainable-yields

argue that much more rigour is required in attributing flood hazards; most studies to date have focused on detection (identifying the changes in the variable, e.g. Ishak et al. (2013)). An attribution framework needs to show that the changes are consistent with the driver of change and just as importantly that the changes are inconsistent with other alternative drivers. Finally a level of confidence in the attribution is required (Merz et al. 2012). Harrigan et al. (2014) also suggest that rather than focussing on single drivers of change, that a multiple working hypotheses framework should be adopted. Similar attribution studies have not been undertaken in Australia and this is an area of research that should be pursed. The concept of *Fraction of Attributable Risk* (Allen 2003) can be used to investigate the contribution of anthropogenic climate change to that a specific flood event (e.g. Pall et al. 2011). In a recent Australian example, Evans and Boyer-Souchet (2012) used an ensemble of high-resolution model simulations to examine the record rainfall totals in Queensland during the 2010 and 2011 floods and found that higher than average sea surface temperatures to the north of Australia contributed 25 % more precipitation than average La Niña conditions.

One of the working hypotheses that should be considered when attributing observed changes is catchment modification. The main source of non-stationarity in catchment conditions is urbanisation, which is known to increase flooding in various ways (Shuster et al. 2005). The increase of impervious surfaces is well known, but other parts of the urban landscape, such as parks, also have lower infiltration rates. Rapid concentration of the flows leads to a shorter and more intense flood peak as well as increases in flood volume. This is an issue particularly for flash floods in small catchments with urban development.

Links and feedback mechanisms with landscape processes are also not fully understood and to date have only been explored over relatively short timescales (Beringer et al. 2011; Wu et al. 2013). In rural areas, forest clearing, bushfire and afforestation activities have a large role in changing catchment conditions and hence flood hazard. However, although a direct link has often been implied between forest cover and flooding, the impact of such activities primarily depends on the effect on soil infiltration capacity, which does not necessarily have to be directly affected by forest cover change (van Dijk and Keenan 2007). Nonetheless, analysis of changes in streamflow after large-scale tree clearing in inland Queensland did suggest a small increase in catchment flood response (Pena-Arancibia et al. 2012).

4 Future changes to floods and causative variables

4.1 Future changes to floods

Limited research has been undertaken to examine changes to future flood hazard globally (Arnell and Gosling 2014) and this is also the case within Australia. General comments on increasing risk of floods and droughts are very common, particularly in the climate change adaptation literature. However, these are typically based on the assumption that changes to rainfall extremes (see Section 4.2) will translate directly to changes in flood. IPCC (2013) and other assessments (e.g., Kundzewicz et al. 2013) have concluded that there is only *low confidence* in numerical projections of changes to flood frequency or magnitude.

Most research has considered specific catchments with a limited subset of driving climate scenarios. Some global studies have considered either large Australian catchments or used a grid-based approach with a land surface model or river routing model to estimate changes in flood magnitudes at all locations around the world. Hirabayashi et al. (2013) used discharge projections from 11 GCMs for a range of future scenarios within a global river routing model.

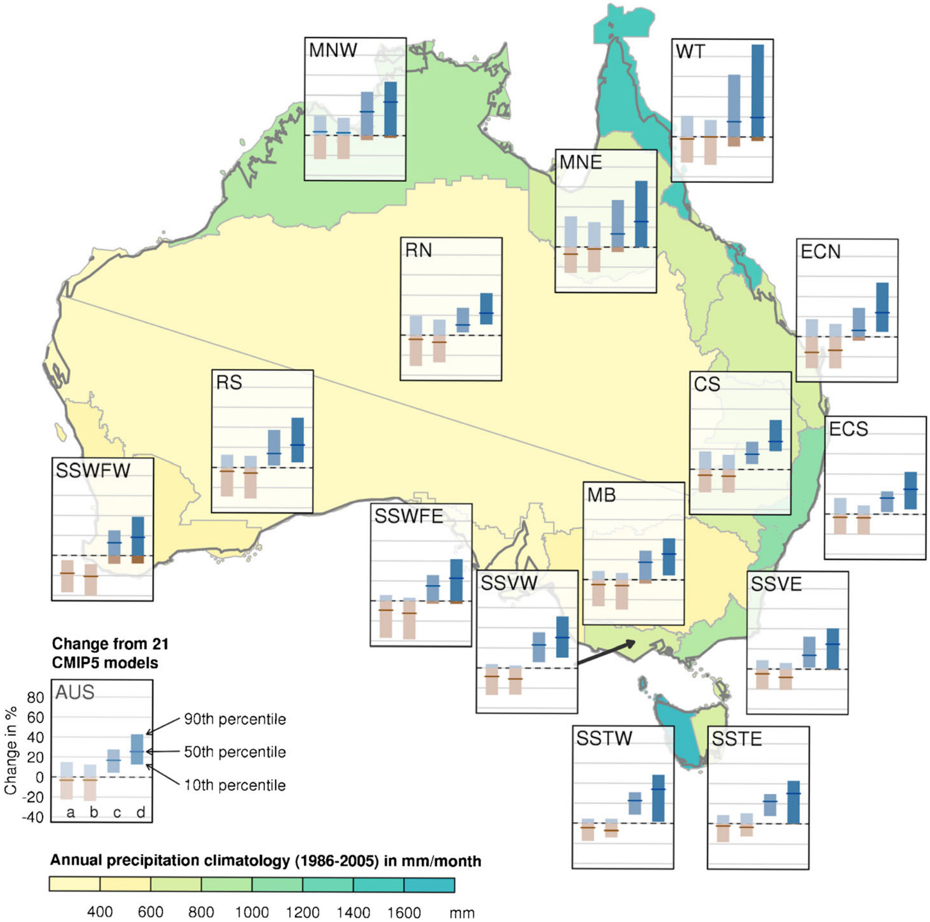

Fig. 3 *Bars* showing median and the 10th to 90th percentile range of projected change in daily rainfall for 2080–2099 relative to 1986–2005 for RCP8.5. Each box shows from *left*: (**a**) annual mean rainfall based on a set of 39 models and from a consistent subset of 21 CMIP5 models the (**b**) annual mean rainfall, (**c**) annual maximum daily rainfall, and (**d**) 20 year return level of the annual wettest day rainfall. *Blue* indicates increase and brown indicates decrease. The Australia average results are shown in the bottom left. Reprint from Figure 7.2.13 in CSIRO and Bureau of Meteorology (2015) (http://www.climatechangeinaustralia.gov.au/media/ccia/2.1.5/cms_page_media/178/TR_Figure7.2.13.png). Reproduced by permission of CSIRO Australia, © CSIRO

It was found that increases in the frequency of floods are likely in northern Australia and along the east coast, whilst flood frequency was projected to decrease in south-western Western Australia. The Murray-Darling basin, which provides over 40 % of Australia's agricultural output and accounts for 70 % of the total irrigated area in Australia (location shown as MB in Fig. 3), was projected to have increased flood magnitudes by 8 out of 11 models but the projected return periods of the current 100-year flood were found to vary from 1 year to approximately 5000 years, the highest variability seen in the 30 basins that were analysed.

Global assessments necessarily make simplifying assumptions in terms of the methods of calculating climate changes. The uncertainty introduced by using simplified hydrologic or hydraulic routing models can be as large as the climate model uncertainty (Dankers et al. 2014). Another issue is that global assessments often focus on percentage changes. However, this does not always translate into impacts of practical significance. For example, a 20 % change to floods in central Australia might be much smaller increase in volume and in the number of people affected than a 5 % change to floods in northern Australia or the heavily populated south-east coast.

As highlighted by Arnell and Gosling (2014), "small-scale flooding from small rivers … is not included, and neither is flash-flooding within urban areas caused by intense rainfall" in these global assessments. In some ways, understanding flash flooding in urban catchments is a simpler task because the effects of antecedent conditions are generally smaller and therefore the change in floods can be more easily related to changes in the extreme rainfalls, while evapotranspiration and pre-event catchment wetness are secondary influences. However, even in this simpler situation, there has been limited catchment-scale research of potential changes to flood risk due primarily to the uncertainties in future rainfall projections discussed in the next section.

4.2 Future changes to causative variables

The focus of this section is on changes to rainfall extremes in Australia, firstly on the basis of projections from GCMs are discussed in the following section followed by results from downscaled climate projections.

4.2.1 Rainfall projections from GCMs

Plausible changes in rainfall extremes have recently been assessed as part of the revised climate change projections for Australia (CSIRO and Bureau of Meteorology 2015). Daily rainfall extremes were defined as the magnitudes associated with a 20-year return period event. The projections are available for eight regions, which provides a useful summary of the potential changes in extremes across the country, however the disadvantage is that for some of the very large regions the 20-year return value is derived as the average of grid cell estimates, obscuring sub-regional variations.

Assessments of the rainfall extremes are based on simulations from 21 GCMs from the Coupled Model Intercomparison Project 5 archive for two Representative Concentration Pathways (RCP4.5 and 8.5) over for two future time horizons: a near future (2020–2039) and a more distant period (2080–2099). Changes are reported relative to the baseline period 1986–2005 used in IPCC (2013). Projections for 2080–2099 in Fig. 3 show very similar results across all regions, indicating that rainfall amounts associated with extreme rainfall events are likely to increase across Australia. Indeed, when comparing projected changes in annual mean rainfall, annual maximum rainfall and the 20-year return value, even in regions where mean annual rainfall is expected to decrease, the annual maximum rainfall is projected to increase. For the 2080–2099 time horizon and RCP 8.5 scenario, median changes typically suggest an increase in annual maximum rainfall of around 25 %.

Potential changes to mean annual rainfall are also relevant to flood hazard as they may affect antecedent conditions in the catchment. For example, returning to the example of the Murray-Darling Basin, the multi-model median increase in rainfall extremes is much larger than the slight decrease in mean annual rainfall, providing support to the projected changes in flood hazard from Hirabayashi et al. (2013). Changes in antecedent conditions are dealt with in more detail in Kiem et al., (this issue). However as highlighted earlier, the interplay between antecedent conditions and extreme rainfall leads to a non-linear response in many catchments and can complicate the interpretation of rainfall changes with respect to flood hazard.

In summary, there is *high confidence* that rainfall extremes will increase across most of Australia. However, there is only *low confidence* in the magnitude of change (CSIRO and Bureau of Meteorology 2015). The reason for this is that many of the processes associated with extreme rainfall, summarised in Westra et al. (this issue) and Walsh et al. (this issue), are poorly resolved in GCMs.

4.2.2 Regional climate projections

While a number of projects have used statistical and dynamical methods to downscale future climate projections over regions of Australia, few of them have explicitly looked at changes in rainfall extremes. Evans and McCabe (2013) examined future changes in moderate precipitation extremes (defined as the annual total of all days above the 99th percentile of daily precipitation) over the Murray-Darling basin and Eastern Seaboard derived from downscaling a single GCM with a single Regional Climate Model (RCM) at 10 km resolution. They show that while the GCM projected decreasing mean precipitation, it consistently projected increasing precipitation at the 99th percentile and above. The RCM projected similar mean changes but did not always project increasing extreme precipitation. White et al. (2013) examined changes in precipitation extremes by downscaling six GCMs using a single RCM (~10 km resolution) over Tasmania and projected increases in maximum 1 and 5-day precipitation intensities, separated by longer dry spells.

In contrast to these studies, Perkins et al. (2014) investigated changes in precipitation extremes (defined as the 20- year return period) produced by an RCM with a lower resolution (60 km) model that covered all of Australia. They found that the changes projected by the RCM differed from the host GCMs and did not support an increase in extreme precipitation across the entire country. These studies suggest caution should be used when relying on projections of precipitation extremes made by models that do not resolve key processes, which is a problem for GCMs and low resolution RCMs. In particular, for sub-daily precipitation extremes it has been suggested that convection-permitting resolution models are required (Kendon et al. 2014; Westra et al. 2014) but few such studies have been performed over Australia (Argüeso et al. 2013) and one should be equally cautious of projections based on single models or very small ensembles.

5 Discussion and recommendations

Flood hazard in Australia is expected to change in the future but at this stage it is not possible to even universally predict the direction of these changes. Trends in rainfall extremes have been explored, but work remains to be done to clarify the relative influences of temperature and moisture availability on extreme precipitation at sub-daily durations. There is some confidence in very high-resolution RCM projections of extremes but these models have not yet been run over the majority of the country. In addition, there are still major gaps in explaining the direction and causes for historical flood trends and variability, and as such projections for future flood behaviour are highly uncertain.

The review of historical analyses on flood hazard shows that there is still extensive research required on how to identify and separate the influences of long-term trends from other climate cycles that occur and interact over a range of timescales. This is an important challenge with respect to precipitation, and the challenge of converting improved understanding of precipitation variability to understanding of the variability of flood hazard adds further complexity. Reliable data sets of flood records are required, similar to the coastal flood database that was recently developed for the period 1860–2012 (Callaghan and Power 2014). The development of the Hydrologic Reference Stations dataset (Bureau of Meteorology 2015) has provided an excellent starting point for a national dataset, although for small catchments the daily resolution data will not capture the peak flood flow. Paleo-flood records, which extend for several hundreds of years to thousands of years, may assist in placing the instrumental records in context for very large catchments, especially those with relatively short and sparse instrumental networks. Examining the temporal variability of these flood databases will provide a useful baseline for understanding the direction of historical changes and, as discussed below, modelling studies will be needed to attribute the trends to their causes and extrapolate to the future.

There is a need for robust future projections of extreme precipitation over daily and longer durations. Such projections need to capture both the large-scale influences of various climate modes, as well as the local to regional scale processes that are the proximal cause of extreme precipitation, including phenomena such as East Coast Lows, bands of thunderstorms, interactions with fronts and topography, and tropical cyclones. A large ensemble of climate projections at resolutions that can capture these phenomena (>3 km) is required to engender more confidence in projections of future floods. The need for high-resolution simulations has also been identified by McInnes et al. (this issue) to improve understanding of storm surge and waves, which also affects coastal flood hazards. As noted in Section 4, for sub-daily precipitation extremes convection-permitting resolution models are required but due to the computational cost of running such models there has been limited work at this scale in Australia, or even globally. Further work is needed to inform engineering practices that depend on these short timescale extremes.

Even if future sub-daily precipitation extremes were completely understood, there remain outstanding questions on the best methods to translate knowledge in future rainfall extremes to future flood hazards. Assessments of changes to the whole spectrum of precipitation from light to heavy events (e.g., Lau et al. 2013) are a valuable starting point for considering future flood hazard but do not provide information on the sequence of flood producing rainfall compared to the pre-event rainfall and thus the catchment antecedent conditions that will affect the severity of the flood. Continuous simulation models for rainfall and runoff are required to address this issue. Developing a common framework for implementing such models to understand changes at the individual catchment scale could be a useful starting point. This framework could cover continuous simulation methods to address climate changes in precipitation, as well as appropriate rainfall runoff models for assessing flood risk. Another useful contribution would be to identify how long continuous sequences of rainfall and runoff data need to be to separate the impacts of natural variability in the climate system from climate change signals with respect to flood hazard in particular.

Further exploration of the relative importance of different runoff-generating mechanisms in Australian catchments, including how these mechanisms vary geographically, will be critical to help better understand the climatic controls on floods, and in particular the relative influence of antecedent moisture on flood magnitude. Insights are likely to be gained through investigation of stores and fluxes of moisture in highly instrumented experimental catchments, combined with numerical experiments such as those conducted by Mirus and Loague (2013) to assess the sensitivity of various assumptions on catchment sensitivity to changes in rainfall extremes.

Our ability to resolve processes across an ever-wider range of spatial and temporal scales, together with the increase in availability of observational data from in-situ and remotely sensed sources, suggests that substantial improvements in our ability to attribute changes to floods are possible. However, in many cases changes in flood hazard will occur because of simultaneous changes to multiple processes (including climatic and land use changes), so that the specific contribution of individual processes will remain difficult to isolate. Furthermore, there remains significant uncertainty in key driving variables such as extreme rainfall. Finally, the 'uniqueness of place' (Beven 2000) of individual catchment processes indicate that changes are likely to be geographically diverse, posing substantial challenges to continental-scale assessments of historical and future changes to flood hazard. As a result, the attribution of changes to flood hazard to specific causes remains a significant challenge.

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