



# Phosphorus availability and arbuscular mycorrhizal fungi limit soil C cycling and influence plant responses to elevated CO<sub>2</sub> conditions

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**Abstract** Soil organic matter (SOM) decomposition and organic phosphorus (P) cycling may help sustain plant productivity under elevated CO<sub>2</sub> (eCO<sub>2</sub>) and low-P conditions. Arbuscular mycorrhizal (AM) fungi and their role in P-acquisition and SOM decomposition may become more relevant in these conditions. Yet, experimental evidence of AM fungi and P availability interactive effects on soil carbon (C) cycling under eCO<sub>2</sub> is scarce with the potential mechanisms of this control being poorly understood. We performed a pot experiment with soil and a grass from a low-P ecosystem where plant biomass and soil C cycling have been mostly unresponsive to eCO<sub>2</sub>. We manipulated AM fungi, P, and CO<sub>2</sub> levels and assessed their impacts on soil C cycling and plant growth using continuous <sup>13</sup>C plant labelling to isolate

and measure short-term changes in total and SOM-derived fractions of respired CO<sub>2</sub>, dissolved organic C (DOC) and microbial biomass (MBC), as relevant components of the soil C cycle. Increases in SOM decomposition and microbial C use were hypothesised to support plant growth under eCO<sub>2</sub> and low-P with AM fungi intensifying this effect. However, we did not detect simultaneous significant impacts of the three experimental factors. We observed instead increased root biomass and nutrient uptake with eCO<sub>2</sub> and AM presence and lower SOM-derived DOC and MBC with low-P, decreasing further with AM inoculation. Taken together, our findings in this model plant-soil system suggest that, AM fungi can support root biomass growth and nutrient uptake under eCO<sub>2</sub> and protect the SOM pool against decomposition even in low-P conditions. Contrary to reports from N-limited ecosystems, our results allow us to conclude that C and P biogeochemical cycles may not become coupled to sustain an eCO<sub>2</sub> fertilisation effect and that the role of AM fungi protecting the SOM pool is likely driven by competitive interactions with saprotrophic communities over nutrients.

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## Introduction

The effect of increased atmospheric CO<sub>2</sub> concentrations on plant productivity and its potential impacts on soil carbon (C) are regulated by nutrient availability (Ainsworth and Long 2004; Finzi et al. 2011; Reich et al. 2014; Terrer et al. 2019). Higher plant productivity and increases in soil C with elevated CO<sub>2</sub> (eCO<sub>2</sub>) have been observed when combined with nutrient addition (Maroco et al. 2002; van Groenigen et al. 2006; De Graaff et al. 2006; Hungate et al. 2009; Dieleman et al. 2010; Piñeiro et al. 2017). In the absence of external nutrient supply, higher soil organic matter (SOM) decomposition under eCO<sub>2</sub> contributes to meet plant nutrient demands allowing for increases in plant biomass (Hoosbeek et al. 2004; Finzi et al. 2006; Reich et al. 2006; Drake et al. 2011; Phillips et al. 2012; Carrillo et al. 2014). In this case, positive responses of plant biomass with eCO<sub>2</sub> are strongly related with decreases in soil C (Terrer et al. 2021). However, it has also been observed that in ecosystems with soil nutrient limitations, eCO<sub>2</sub>-induced increases in plant productivity are constrained (Reich et al. 2006; Reed et al. 2015; Ellsworth et al. 2017) resulting in no effects on soil C. The current understanding about nutrient availability regulating eCO<sub>2</sub> impacts on C cycling has been obtained mainly from northern hemisphere systems, where nitrogen (N) often constrains ecosystem productivity and soil microbial activity. But the impacts of eCO<sub>2</sub> on soil C cycling may be dependent on what the most limiting nutrient is (Dijkstra et al. 2013). Phosphorus-limited tropical and subtropical ecosystems are relevant C sinks that cover a vast area (Soepadmo 1993; Pan et al. 2011; Keenan et al. 2015) and although the role of P availability regulating the impact of eCO<sub>2</sub> on soil C storage is broadly recognised (Vitousek et al. 2010; Goll et al. 2012; Reed et al. 2015; Sun et al. 2017), the mechanisms of this control are poorly understood (Norby et al. 2015; Ellsworth et al. 2017; Terrer et al. 2019).

In contrast to N, soil P cycling is less coupled to soil C dynamics because P mobilisation is not necessarily linked with SOM decomposition. The P cycle can be divided into inorganic and organic sub-cycles (Mullen 2005; Liu and Chen 2008). Mobilisation of inorganic P occurs mainly via physicochemical mechanisms without SOM mineralisation nor CO<sub>2</sub> production, thus C and the inorganic P cycle are not

coupled. On the other hand, organic P is made available via P-hydrolytic enzymes that release phosphates without CO<sub>2</sub> production (Nannipieri et al. 2011) or as a consequence of substrate degradation (Sharma et al. 2013) driven by the nutrient and energy needs of soil organisms (Tate 1984) and involving CO<sub>2</sub> release as a consequence of SOM decomposition. Most P is made available via inorganic P mobilisation and via P-hydrolytic enzymatic activity, particularly with P sufficiency. But in low-P systems, mineral P sources are depleted and P derived from rock weathering is minimal (Reed et al. 2011; Goll et al. 2012). Therefore, the majority of P acquisition occurs via internal recycling of organic P sources (Johnson et al. 2003; Fox et al. 2011; Reed et al. 2011). Most of the research on the effects of eCO<sub>2</sub> on soil P availability are focused on its impacts on inorganic P and organic P cycling that do not involve SOM decomposition. For example, higher microbial P enzymatic activity is reported with eCO<sub>2</sub> (Moorhead and Linkins 1997; Kelley et al. 2011; Hasegawa et al. 2016; Souza et al. 2017) as well as higher production of organic acids and siderophores involved in P mobilisation (Watt and Evans 1999; Tarnawski and Aragno 2006; Fransson and Johansson 2010; Högy et al. 2010). Moreover, increases in fine root production and changes in pH and soil moisture with eCO<sub>2</sub> can indirectly enhance inorganic P cycling in soils (Dijkstra et al. 2012; Jin et al. 2015; Hasegawa et al. 2016). It is likely that organic P cycling involving SOM decomposition is also affected by eCO<sub>2</sub> due to the influence of eCO<sub>2</sub> on ecosystem stoichiometry (Gifford et al. 2000; Loladze 2002, 2014) but experimental evidence for this is scarce. A better understanding about how eCO<sub>2</sub> and P availability influence soil C cycling is important to better predict changes in global soil C due to eCO<sub>2</sub> (Finzi et al. 2011; Goll et al. 2012; Yang et al. 2016).

In low-P conditions, organic P cycling may help sustain plant productivity under eCO<sub>2</sub> at the expense of higher SOM decomposition. As C allocation belowground increases with eCO<sub>2</sub> (Canadell et al. 1995; Cotrufo and Gorissen 1997; Pausch and Kuzyakov 2018), higher labile C availability can lead to increased microbial biomass and activity and higher SOM decomposition may allow for a sustained nutrient release for plant uptake and growth in low-P soils. However, evidence from field experiments exposed to eCO<sub>2</sub> suggests that under P-limiting conditions, changes in soil C cycling are minor or undetected

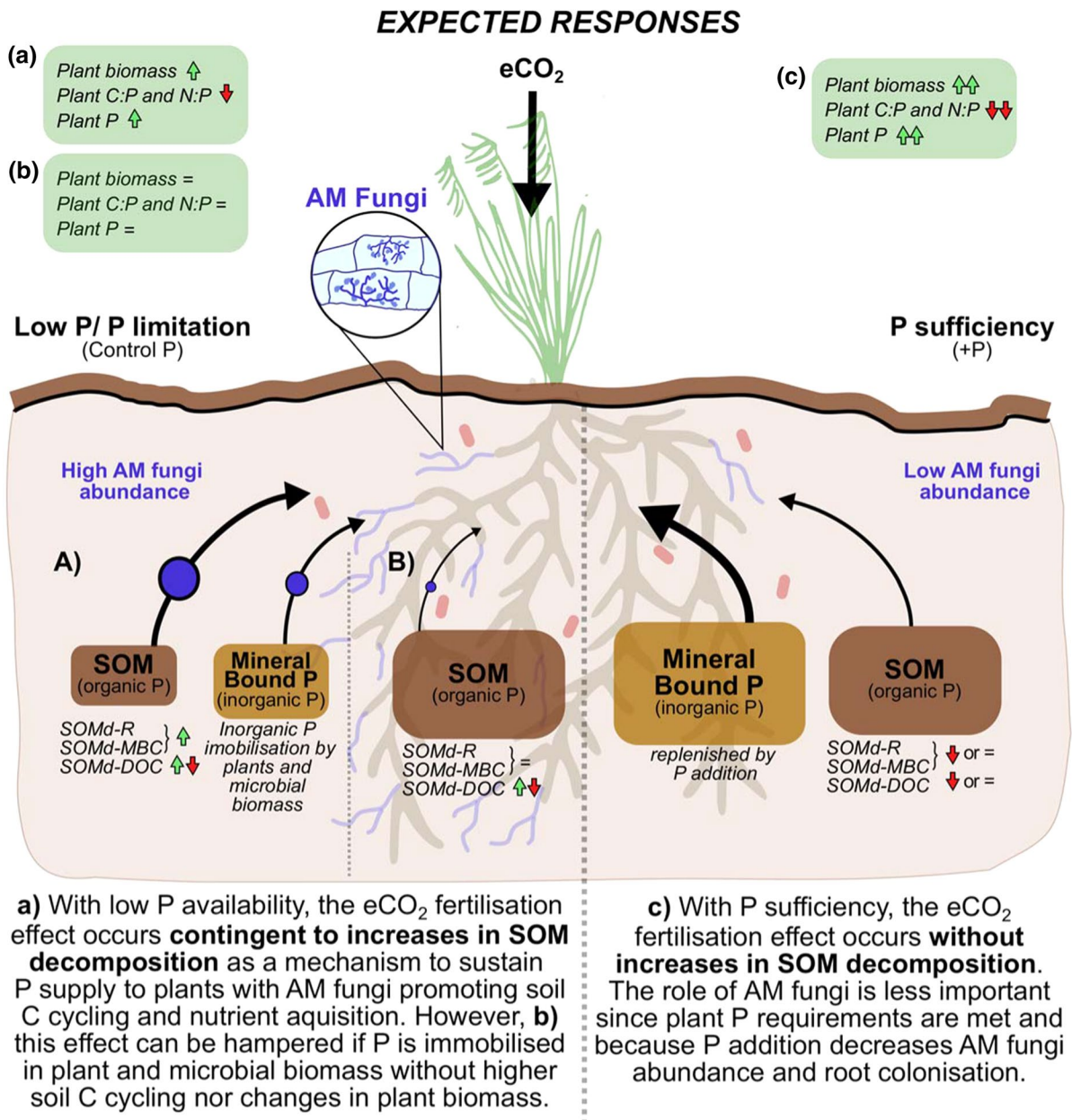
(Dijkstra et al. 2013; Drake et al. 2016; Castañeda-Gómez et al. 2020) and that the CO<sub>2</sub> fertilization effect is reduced or not observed (Conroy et al. 1992; Peñuelas et al. 2012; Goll et al. 2012; Zhang et al. 2014; Deng et al. 2015; Ellsworth et al. 2017). Thus, if P availability becomes limiting, ecosystem productivity and microbial-mediated SOM decomposition may be hampered, preventing further changes in the soil C pool. Soil microbial communities play a key role in SOM decomposition and both organic and inorganic P cycling. Moreover, the responses of SOM decomposition and C stabilisation to P availability have been found to be largely mediated by changes in saprotrophic microbial community composition and activity (Liu et al. 2018; Soong et al. 2018; Fang et al. 2019; Feng and Zhu 2019; Yuan et al. 2021) and by symbiotic fungi such as arbuscular mycorrhizal (AM) fungi (Xu et al. 2018). Yet the role of saprotrophic and mycorrhizal fungi altering SOM decomposition under eCO<sub>2</sub> conditions and low-P availability is poorly understood.

AM fungi are known to greatly contribute to P mobilisation under eCO<sub>2</sub> conditions, especially in ecosystems with low soil P availability (Delucia et al. 1997; Matamala and Schlesinger 2000; Terrer et al. 2016) where AM fungi tend to be more abundant (Treseder and Cross 2006). AM fungi also alleviate P limitation conditions for plants (Willis et al. 2013; Soka and Ritchie 2014) possibly via excretion of P hydrolytic enzymes (Joner et al. 2000; Sato et al. 2015) and certainly via uptake of inorganic nutrients from the soil solution (Javot et al. 2007; Smith et al. 2011). They have also been linked with increased SOM decomposition with climate change (Talbot et al. 2008; Cheng et al. 2012; Carrillo et al. 2015; Wei et al. 2019). Higher C transfer to AM fungi under eCO<sub>2</sub> enhances AM fungi growth and activity (Mohan et al. 2014) and so their contribution to plant P acquisition in P-limited ecosystems may be enhanced with eCO<sub>2</sub> (Cavagnaro et al. 2011). On the other hand, AM fungi can also suppress eCO<sub>2</sub>-mediated increases in plant growth due to increased P immobilisation in AM fungal biomass and increased competition over P with saprotrophic communities (Jin et al. 2015) that can simultaneously decrease AM fungi-mediated SOM decomposition (Talbot et al. 2008). If P availability is low, the role of AM fungi mobilising P and promoting plant growth may be constrained due to P-limitation (Treseder and Allen 2002), which might

further regulate impacts of eCO<sub>2</sub> on ecosystem productivity in P-limited ecosystems. Due to the significant role of AM fungi on P and SOM dynamics, predictions of altered soil C stocks under eCO<sub>2</sub> require an understanding how eCO<sub>2</sub> can change AM fungal activity, their impact on saprotrophic communities and plant responses to eCO<sub>2</sub>, particularly under P limitation.

In this study, we explored the role of P availability and AM fungi mediating responses of soil C cycling and plant productivity to eCO<sub>2</sub> in a model plant-soil system. Our aim was to answer two questions: How does P availability controls soil C cycling responses to eCO<sub>2</sub> and how does AM fungi mediate these responses. We used plant and soil material from the Eucalyptus Free Air CO<sub>2</sub> enrichment (EucFACE site) as a model system to answer these questions. At this site, both plant growth and soil C responses to eCO<sub>2</sub> have been limited and it has been hypothesised that P limitation is preventing stronger ecosystem responses to eCO<sub>2</sub>. We manipulated P availability, AM fungal presence and atmospheric CO<sub>2</sub> levels. An Australian native grass species (*Microlaena stipoides*) with the ability to grow in a wide range of P availability conditions and forming arbuscular mycorrhizal associations (Hill et al. 2010; Clark et al. 2014) was grown in controlled environment chambers. This grass species is also a dominant understory species present at the EucFACE site (Hasegawa et al. 2018). The growth chambers used allowed the continuous isotopic labelling of plant tissues so we could directly quantify short-term changes in soil C cycling by measuring relevant total and SOM-derived fractions of relevant soil C components: respired CO<sub>2</sub> (R; the product of decomposition), dissolved organic C (DOC; considered an active pool of C) and microbial biomass (MBC; the agent of SOM decomposition). This approach allowed for a comprehensive assessment of the impact of eCO<sub>2</sub>, P availability and AM fungal presence on short-term soil C cycling and plant growth in a low-P model system. Changes in saprotrophic communities were also measured (using phospholipid fatty acids) in relation to eCO<sub>2</sub> conditions, P treatment and AM fungal presence. These responses, along with nutrient contents in plant tissues and available soil nutrients, were collected to test our mechanistic hypotheses (Fig. 1).

We hypothesised that under eCO<sub>2</sub> and low P conditions, increases in plant growth will be



**Fig. 1** Conceptual framework and expected responses of measured variables. **(a)** Under eCO<sub>2</sub> and low P conditions, increases in plant growth and nutrient acquisition will be contingent to increases in soil C cycling facilitated by AM fungi presence. However, **(b)** if P conditions are too low, soil microbial activity will be limited, including that of AM fungi, and thus, changes to soil C cycling, plant biomass or plant nutrient contents will be constrained. Increased plant biomass under eCO<sub>2</sub> and P-sufficiency **(c)** will not be necessarily accompanied by changes in soil C cycling, with most of P acquisition occurring from the replenished mineral P pool without the

involvement of AM fungi. Soil C cycling is measured in this study as soil organic matter-derived (SOMd) C components: R, respired CO<sub>2</sub>; MB, microbial biomass; DOC, dissolved organic C. Green arrows: increase; red arrows: decrease; equal sign (=): no effect. Box sizes represent the relative size of each pool. The thickness of the arrows from SOM and mineral bound P pools represent the relative flux of P from these pools to plants. The size of the circles on top of each arrow represent the relative role of AM fungi on P acquisition and SOM decomposition from these pools

contingent to increases in soil C cycling, evidenced as higher SOM-derived respired  $\text{CO}_2$ , MBC and altered SOM-derived DOC (Fig. 1a). In this scenario, both low-P availability and  $\text{eCO}_2$  promote AM fungal colonisation rates and activity and thus, AM fungi presence will promote soil C cycling and enhance organic P cycling and nutrient acquisition detected as higher plant P, and lower C:N and C:P in plant tissues. P uptake by AM fungi from the depleted mineral P pool will be therefore reduced in these conditions (AM role reflected on the circles on top of each flux in Fig. 1a). Alternatively, we hypothesised that no fertilization effect would be observed if P conditions were too low due to a limitation of soil microbial activity, including that of AM fungi, without changes to soil C cycling, plant biomass or plant nutrient contents (Fig. 1b). Contrarily, increased plant biomass under  $\text{eCO}_2$  and P-sufficiency will not be necessarily accompanied by changes in soil C cycling, with most of P acquisition occurring from the replenished mineral P pool (Fig. 1c). In this scenario, AM fungi will not play a relevant role in SOM decomposition nor nutrient acquisition as high P availability decreases AM fungi abundance and activity.

## Materials and methods

We set up a factorial experiment with three factors: **phosphorus (P) treatment**: P addition (+P)/control; **AM fungi**: Active AM inoculum (AM)/Non-mycorrhizal (Sterile AM inoculum, -AM) and  **$\text{CO}_2$** : ambient (400 ppm— $\text{aCO}_2$ ) and elevated (640 ppm— $\text{eCO}_2$ ) for a total of 8 treatment combinations with 4 replicates each (+P/+AM/ $\text{aCO}_2$ ; +P/+AM/ $\text{eCO}_2$ ; +P/-AM/ $\text{aCO}_2$ ; +P/-AM/ $\text{eCO}_2$ ; control/+AM/ $\text{aCO}_2$ ; control/+AM/ $\text{eCO}_2$ ; control/-AM/ $\text{aCO}_2$ ; control/-AM/ $\text{eCO}_2$ ;  $n=32$ , Fig. SI 1a). Plants of *Microlaena stipoides* (Labill.) R.Br., a native Australian  $\text{C}_3$  grass with a broad P range, were grown from seed in 125 mm pots containing around 500 g of dry soil. Plants were grown in growth chambers for 12 weeks and soil water content was kept between 15 and 20% gravimetric content by addition of MilliQ water as needed every 2–3 days.

## Soil characteristics and treatment preparation

Soil from 0 to 15 cm depth was collected from a Cumberland plain natural woodland in Western Sydney (33° 37' 01" S, 150° 44' 26" E, 20 m.a.s.l.) from the EucFACE site. Soil at this site is an aeric podsol, slightly acidic ( $\text{pH } 5.38 \pm 0.02$  at 0–10 cm) (Ross et al. 2020), with a N content of  $677 \text{ mg Kg}^{-1}$ , low total C (1.8%, 0–15 cm), and P content ( $76.28 \text{ mg Kg}^{-1}$ ) (Hasegawa et al. 2016). Previous studies have demonstrated that both the vegetation (Crous et al. 2015) and soil fauna (Nielsen et al. 2015) are limited by P within this ecosystem. Soil was sieved (to 2 mm) and sterilised (gamma irradiated, 50 kGy) to remove viable AM fungal propagules. To reintroduce a homogenous microbial community to all pots, excluding AM fungi, we prepared a microbial inoculum. For this, approx. 3 Kg of freshly collected soil from a grassland nearby the EucFACE site was mixed with water in a 1:3 proportion by volume and the suspension passed through a 20  $\mu\text{m}$  mesh sieve to remove AM fungal spores (Brundett and Australian Centre for International Agricultural Research 1995). Prior to potting, the filtrate was added to the sterile soil at a rate of 50 mL filtrate per Kg soil. The soil was incubated for a week at room temperature, mixing it daily.

The soil was then divided into four subsets based on whether they were to receive additional P and AM fungal inoculum. For the former, triple super phosphate (Richgro, super phosphate fertiliser supplement 9.1% P w/w) was added in a rate of 0.4 g/kg dry soil. The triple super phosphate powder was weighed, diluted and sprayed into the soil while low-P treatments received MilliQ water only. The soil was then thoroughly mixed and left to settle for a day. Next, AM fungi treatments were created by applying AM fungal inoculum in a 1:10 AM fungi inoculum:soil with non-mycorrhizal controls produced by autoclaving the live AM inoculum (121 °C, 2 h) and applying it to the soil in the same manner (See Supplementary Information for AM inoculum production). Once the four types of soils were prepared, soil was added to pots and five surface sterilised *M. stipoides* seeds (30%  $\text{H}_2\text{O}_2$  for 10 min followed washing) were sown per pot. Pots were randomly split and placed in  $\text{aCO}_2$  and  $\text{eCO}_2$  chambers. After 2 weeks of growth, pots were thinned to one plant per pot.

Unplanted pots ( $n=16$ ) with the different P and CO<sub>2</sub> treatments (+P/–AM/aCO<sub>2</sub>; +P/–AM/eCO<sub>2</sub>; control/–AM/aCO<sub>2</sub>; control/–AM/eCO<sub>2</sub>  $n=4$  each), were included to estimate effects on SOM decomposition in the absence of plants and AM fungi (Fig. SI 1a). Unplanted pots were all prepared using soil with sterile AM inoculum and were kept under the same conditions as the planted pots for the duration of the experiment. The dried soil  $\delta^{13}\text{C}$  from these unplanted pots was measured and an averaged value of all unplanted pots was used for the isotopic partitioning. See “[Respired, dissolved organic C \(DOC\) and microbial biomass isotopic composition and partitioning](#)” section.

#### Growth chambers set up

The growth chambers (six in total, three per CO<sub>2</sub> treatment) were modified using the approach of Cheng and Dijkstra (2007) to achieve a continuous <sup>13</sup>C-labelling of plant tissues in both aCO<sub>2</sub> and eCO<sub>2</sub> treatments. The chambers were adapted to take an influx of naturally <sup>13</sup>C-depleted CO<sub>2</sub> ( $\delta^{13}\text{C} = -31.7\text{‰} \pm 1.2$ ) delivered during the photoperiod, combined with a scrubbing system made of a soda lime-filled PVC tube (72 L) that allowed continuous supply of CO<sub>2</sub>-free air (Fig. SI 1b). Chambers were adjusted to a 16h/8h photoperiod with 25 °C/18 °C, 60% relative humidity, and light intensity of 900  $\mu\text{mol}/\text{m}^2\text{s}^{-1}$ . The chamber atmosphere was sampled frequently to confirm depletion in <sup>13</sup>C. Air samples from the chambers were extracted via a pump system into a gas tight tedlar bag (Tedlar® SCV Gas Sampling Bag) and analysed for  $\delta^{13}\text{C}$  in a PICARRO G2201i isotopic CO<sub>2</sub>/CH<sub>4</sub> analyser (Picarro Inc., Santa Clara, CA, USA). Pots were randomly moved among growth chambers with the same CO<sub>2</sub> condition to prevent “chamber effects”. Moreover, to avoid plant-uptake of <sup>13</sup>C from outside the chambers, watering and all other manipulations were performed during the night period aided by green light.

#### Harvest and sample processing

##### *Gas sampling of the plant-soil system*

After 12 weeks of growth, we quantified rates of total respired CO<sub>2</sub> (R) and its C isotopic composition as described by Carrillo et al. (2014, 2015). Briefly, pots

were placed on an elevated platform inside a water filled tray and covered with a PVC chamber (45 cm H x 15 cm D) adapted with an air-tight rubber stopper for air sampling. Free-CO<sub>2</sub> air was circulated for 2 h using an aquarium pump connected to a CO<sub>2</sub> scrubber (50 cm H x 4 cm D PVC tubing filled with soda lime). After 2 h of scrubbing, an air sample was taken to determine baseline CO<sub>2</sub> concentrations using a 7890 A gas chromatograph with a G1888 network headspace sampler (Agilent Technologies, USA). Later, the CO<sub>2</sub> scrubbers were removed and the pump reconnected to the PVC chamber to allow for air circulation while pots were incubated. After 2 to 3 h of incubation, a gas sample was collected in an airtight gas collection bag (Tedlar® PVDF, 1 L) using an aquarium pump system and analysed for its C isotopic composition.

##### *Plant and soil harvest*

One day after gas sampling, pots were destructively harvested. Aboveground biomass was cut and roots separated from the soil and washed. Plant aboveground biomass, roots and a subsample of fresh soil were oven dried at 60 °C for measurements of dry plant biomass, soil gravimetric water content and total nutrients. A fresh soil subsample was stored at –20 °C for microbial community analyses and the remaining fresh soil was used for assessments of dissolved nutrients and microbial biomass.

##### *Plant, soil nutrients and microbial biomass C and N*

Two sub-samples of soil were weighed. Dissolved nutrients in soil were extracted from the first subsample with a 0.05 M K<sub>2</sub>SO<sub>4</sub> solution in a 4:1 solution to soil ratio, shaking at 180 rpm for an hour. Samples were filtered through a Whatman # 42 filter paper and frozen (–20 °C) until analyses. The second subsample was fumigated for 5 days with chloroform and then nutrients were extracted as for unfumigated samples. The fumigated and unfumigated extracts were analysed for total dissolved organic C and N (Shimadzu® TN, TOC-L, Japan) and microbial biomass C and N calculated by subtracting fumigated and unfumigated samples (Vance et al. 1987). The volume left of these K<sub>2</sub>SO<sub>4</sub> extractions was used to obtain the isotopic composition of DOC and microbial biomass (see section below). Phosphates were

extracted following the Bray 1-P method for acidic soils, in a 1:7 solution:soil ratio using a 0.03 M  $\text{NH}_4\text{F}$  solution in 0.025 M hydrochloride (HCl) adjusted pH to  $2.6 \pm 0.05$  with HCl (Rayment et al. 2010). Samples were manually shaken for 60 s and immediately poured over a Whatman # 42 filter paper. Collected extracts were analysed by colorimetry (AQ2 Discrete Analyser, SEAL Analytical, Mequon, WI, USA).

Total P and N were determined from ground oven-dried soil and root samples. Total P concentration was analysed by an X-ray fluorescence spectrometer (PANalytical, epsilon 3. 10Kv, 0.9 mA. Lelyweg, Almelo, Netherlands) while total C and N from soil, roots and soil extracts were analysed along their isotopic composition.

Respired, dissolved organic C (DOC) and microbial biomass isotopic composition and partitioning

Gas samples from plant-soil system incubations were analysed in a PICARRO analyser (G2201i; Picarro, Santa Clara, CA, USA. Precision values below 0.16‰) for the  $\delta^{13}\text{C}$  and  $\text{CO}_2$  concentration of the total respired  $\text{CO}_2$  (R). For the isotopic composition of the DOC and microbial biomass C (MBC), fumigated (*f*) and unfumigated (*uf*) soil  $\text{K}_2\text{SO}_4$  extracts were oven dried at 60°C. The dried extracts were scraped and weighed for analysis on a Thermo GC-C-IRMS system (Trace GC Ultra gas chromatograph, Thermo Electron Corp., Milan, Italy; coupled to a Delta V Advantage isotope ratio mass spectrometer through a GC/C-III); University of California, Davis Campus, USA). The  $\delta^{13}\text{C}$  of unfumigated soil extracts was used as the isotopic composition of the DOC while  $\delta^{13}\text{C}$  of MBC ( $\delta^{13}\text{C}_{\text{MBC}}$ ) were calculated from  $\delta^{13}\text{C}$  values of both fumigated (*f*) and unfumigated (*uf*) extractions (See Supplementary Information for calculations of the isotopic composition of the microbial biomass). Samples of dried soil were also analysed on the Thermo GC-C-IRMS system for C, N and  $\delta^{13}\text{C}$ .

To calculate the fractions of SOM-derived C in the total respired  $\text{CO}_2$ , DOC and MBC, we used isotopic partitioning applying the formula:  $\text{SOM.C}_{R,DOC,MBC} = (\delta^{13}\text{C}_{R,DOC,MBC} - \delta^{13}\text{C}_p) / (\delta^{13}\text{C}_{\text{control}} - \delta^{13}\text{C}_p)$ . Where  $\delta^{13}\text{C}_{R,DOC,MBC}$  were the  $\delta^{13}\text{C}$  of either the  $\text{CO}_2$  measured in the total respired  $\text{CO}_2$  (R), dissolved organic C (DOC) or microbial biomass C (MBC)) from each pot;  $\delta^{13}\text{C}_p$  is the isotopic ratio of the plant source

averaged across P treatments per  $\text{CO}_2$  condition (root biomass at a $\text{CO}_2$   $\delta^{13}\text{C} = -40.01 \pm 0.08$  and e $\text{CO}_2$   $\delta^{13}\text{C} = -43.20 \pm 0.04$ ) and  $\delta^{13}\text{C}_{\text{control}}$  is the average  $\delta^{13}\text{C}$  of the native SOM source obtained from the dried bulk soil from unplanted pots across  $\text{CO}_2$  and P conditions ( $\delta^{13}\text{C} = -25.34 \pm 0.02$ ). The fractions of plant-derived C were obtained by subtracting SOM-derived fractions from the unit. Measured total respired  $\text{CO}_2$ , DOC and MBC were partitioned into SOM-derived and plant-derived, using these fractions to obtain the mass of SOM derived R, MBC and DOC. In this study, we focus only on the SOM-derived C components.

Microbial community analysis: Phospholipid-derived fatty acids (PLFA) and neutral lipid-derived fatty acids (NLFA)

Soil PLFAs were extracted to assess the overall microbial communities while the 16:1 $\omega$ 5c NLFA was used as an indicator of arbuscular mycorrhizal (AM) fungi presence and abundance (Olsson 1999). Freeze-dried soil from planted pots ( $n=32$ ) were extracted following the protocol by Buyer and Sasser (2012) with modifications by Castañeda-Gómez et al. (2020) (See Supplementary Information for PLFA and NLFA analyses). Functional groups were defined as shown in Table SI 1. Fungal to bacterial ratio (F:B) was calculated by dividing fungal PLFAs (not including AM fungi) by the sum of bacterial PLFAs. The sum of individual lipids was used as an indicator of the size of the community ( $\mu\text{g PLFA g}^{-1}$  dry soil).

Statistical analyses

The effect of  $\text{CO}_2$  condition, AM fungi treatment and P addition and their interactions on the response variables was analysed with a linear model fitted with the function “lm” from the stats package in R version 3.3.2 (R Core Team 2019). This approach was used instead of a mixed effects modelling approach since pots were moved among chambers with the same  $\text{CO}_2$  treatment so it was not possible to estimate a random effect associated with each chamber. The normality of the residuals of each model was inspected to check the appropriateness of the fit and transformations were performed when needed. Statistical significance was determined performing an ANOVA (Analyses of variance) with the “Anova” function (“car”

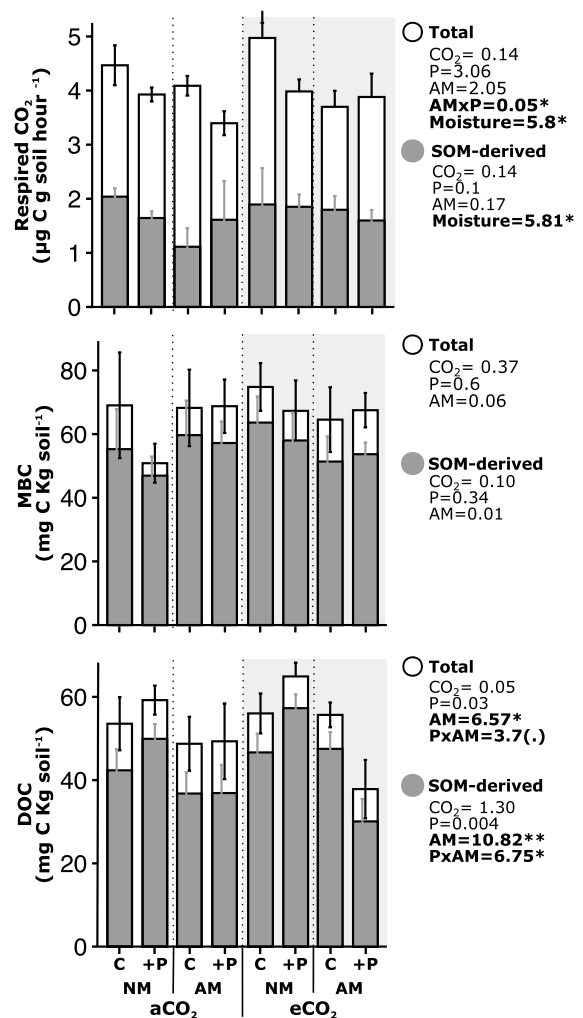
package, Fox et al. 2021). Multiple mean comparisons were performed using the Tukey test using the “glht” function (“multcomp” package, Hothorn et al. 2021). As soil moisture can affect soil respiration measurements, we tested for correlations between soil moisture and the total respired CO<sub>2</sub>, SOM-derived respired CO<sub>2</sub> fraction and total mass of SOM-derived respired CO<sub>2</sub>. For response variables with a significant correlation with moisture, a 3-way ANCOVA was performed with soil moisture as covariate, to account for the variability brought by the slightly different moisture contents at the time of CO<sub>2</sub> sampling. The homogeneity of the regression slopes, normality of residuals and homogeneity of variances was tested when performing the ANCOVA. Significance levels were: ≤ 0.1 (.), ≤ 0.05 (\*), ≤ 0.01 (\*\*), ≤ 0.001 (\*\*\*)

Microbial communities were analysed as the PLFA-based total microbial biomass (sum of all individual lipids, in µg PLFA g<sup>-1</sup> dry soil), as absolute microbial abundance (reflecting the size or biomass of the community, in µg PLFA g<sup>-1</sup> dry soil) and as relative abundance (in percentage, reflecting microbial community composition) of the different microbial groups and as individual lipids. To visualise significant three-way interactions of experimental factors in the abundance of microbial groups from the ANOVA, the “emmip” function from the “emmeans” package (Lenth et al. 2020) was used to show the estimated marginal means from the fitted linear model.

## Results

### Influence of elevated CO<sub>2</sub>, P availability and AM fungi on SOM decomposition and soil C cycle components

Higher SOM-C losses under eCO<sub>2</sub> were expected for control P conditions (low-P), particularly when AM fungi were inoculated (Fig. 1a). However, we found that in general, there was not a significant interactive impact of the three experimental factors on the soil C cycle, represented in the measured SOM-derived fractions of respired CO<sub>2</sub>, MBC and DOC, as initially expected. Instead, we found that the interaction of AM and P treatments determined the fate of the soil C cycle components, regardless of eCO<sub>2</sub>. The total respired CO<sub>2</sub> (Fig. 2) was higher for control P conditions but only for uninoculated AM pots (–AM.



**Fig. 2** Mean values ( $\pm$  standard error,  $n=4$ ) of the total (white) and SOM-derived (grey) respired CO<sub>2</sub>, microbial biomass C (MBC) and dissolved organic C (DOC) per CO<sub>2</sub> condition, AM fungi (with (+AM) and without (–AM) mycorrhizal fungi) and P addition treatment (C: control and +P: P addition). CO<sub>2</sub> conditions are shown on light grey (elevated) and white (ambient) backgrounds.  $F_{(1,24)}$  values of ANCOVAs (for respired CO<sub>2</sub>) and ANOVAs from linear model for response variables displayed on the right. Significance levels: ≤ 0.1 (.), ≤ 0.05 (\*), ≤ 0.01 (\*\*), ≤ 0.001 (\*\*\*)

Only significant interactions of the treatments displayed  $P < 0.05$ , Tukey’s multiple comparison). However, we did not find any significant effects of CO<sub>2</sub> conditions, AM or P treatments on the SOM-derived R (Fig. 2). On the other hand, most of the MBC was derived from SOM (above 80%, Table SI 2) and while neither the total MBC nor the SOM-derived MBC (Fig. 2) were affected by the experimental factors,



the C sources used by the MBC were affected by the interaction between AM and P treatments (Table SI 2). The fraction of MBC derived from SOM was marginally lower by 6% for control P and –AM pots ( $P=0.1$ , Tukey's multiple comparison, Table SI 2), further decreasing for +AM pots. Similar to MBC, most of the DOC was derived from SOM (above 80%, Table SI 2) and both the total DOC and SOM-derived DOC were affected by the interaction between AM and P treatments (Fig. 2) with significantly lower SOM-derived DOC for control P and –AM pots ( $P=0.001$ , Tukey's multiple comparison) and further decreasing for +AM pots. Finally, although not evidenced in the total or SOM-derived R and DOC, the fractions of SOM-derived R and DOC significantly increased with  $e\text{CO}_2$  (Table SI 2).

Soil nutrient responses to elevated  $\text{CO}_2$ , P availability and AM fungi

We expected to see decreases in soil dissolved P and N under  $e\text{CO}_2$ , control P and AM fungi presence due to higher nutrient acquisition and immobilisation by the fungal biomass. As hypothesised, dissolved P

significantly decreased with  $e\text{CO}_2$ , particularly under control P and +AM conditions (Interactive effects of  $\text{CO}_2$  by P and  $\text{CO}_2$  by AM, Table 1). Congruently, we observed significantly higher dissolved C:P in  $e\text{CO}_2$  conditions for control P and for +AM pots (Fig. 3). Similarly, Dissolved N decreased with AM fungi presence and marginally decreased under  $e\text{CO}_2$  conditions in +AM pots (Table 1). Thus, higher dissolved C:N observed for control P conditions for AM pots while higher dissolved C:N was detected for AM pots under  $e\text{CO}_2$  (Fig. 3). Finally, AM fungi significantly decreased total P (–AM:  $175.2 \pm 11.1$  ppm; +AM:  $206.8 \pm 15.1$  ppm.  $AM\_F_{(1,24)}=5.41$ ,  $AM\_P=0.03$ ) while higher total soil P was observed as a result of P addition (control:  $155.1 \pm 8.8$  ppm; +P:  $226.9 \pm 11.7$  ppm.  $P\_F_{(1,24)}=28.01$ ,  $P\_P<0.00$ ) although soil dissolved P did not increase with P addition (Table 1).

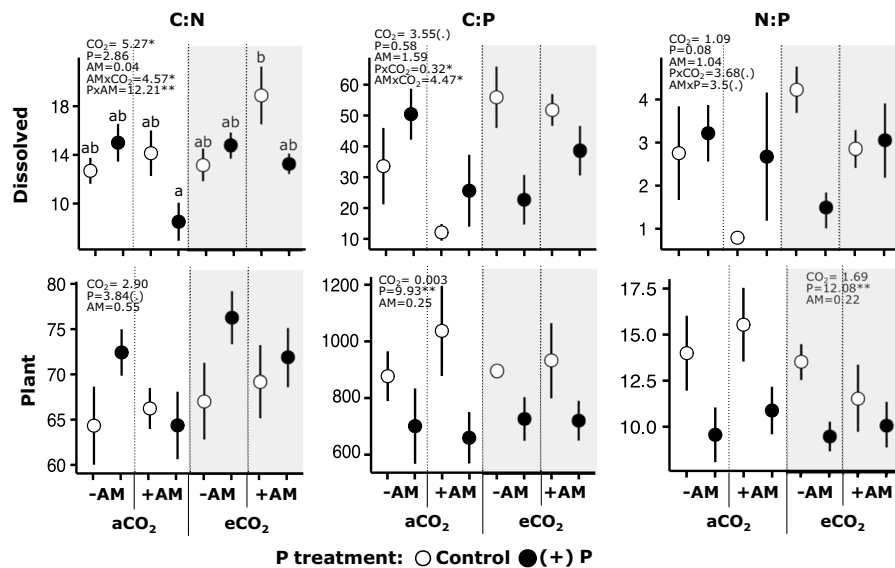
Plant responses to elevated  $\text{CO}_2$ , P availability and AM fungi

We expected that plant biomass would increase with  $e\text{CO}_2$  and low P conditions in AM fungi presence,

**Table 1** Soil dissolved nutrients [ppm. TN—total nitrogen and phosphates ( $\text{PO}_4$ )] per  $\text{CO}_2$  condition, AM fungi and P addition treatment

			TN $\times$ (ppm)	$\text{PO}_4 \times$ (ppm)
a $\text{CO}_2$	–AM	Control	4.2 (0.4)	2.5 (0.8) <sup>ab</sup>
		+P	4.2 (0.5)	1.3 (0.1) <sup>ab</sup>
	+AM	control	3.6 (0.6)	3.9 (0.3) <sup>b</sup>
		+P	4.8 (1.0)	2.8 (0.6) <sup>ab</sup>
e $\text{CO}_2$	–AM	Control	4.4 (0.1)	1.1 (0.1) <sup>a</sup>
		+P	4.5 (0.5)	3.4 (0.8) <sup>ab</sup>
	+AM	Control	3.7 (0.2)	1.1 (0.1) <sup>a</sup>
		+P	3.0 (0.8)	1.0 (0.1) <sup>a</sup>
		$\text{CO}_2$	1.74	<b>8.84**</b>
		P	0.16	0.0001
		AM	<b>6.58*</b>	0.08
		$\text{CO}_2 \times \text{P}$	0.60	<b>7.64*</b>
		$\text{CO}_2 \times \text{AM}$	3.64 (.)	<b>11.45**</b>
		AM $\times$ P	0.15	3.46
	P $\times$ AM $\times$ $\text{CO}_2$	1.22	2.90	

+AM: with and –AM: without mycorrhizal fungi; a $\text{CO}_2$ : ambient  $\text{CO}_2$  and e $\text{CO}_2$ : elevated  $\text{CO}_2$ ; C: control P, +P: P addition. Mean values ( $\pm$  standard error,  $n=4$ ) followed by the same letter are not significantly different ( $p \leq 0.05$ , Tukey multiple comparison test), no letters indicate non-significant effect of the treatments. Below, ANOVA results from linear model (lm) for response variables ( $\times$  natural log transformed).  $F_{(1,24)}$  values displayed with significance levels:  $\leq 0.1$  (.),  $\leq 0.05$  (\*),  $\leq 0.01$  (\*\*),  $\leq 0.001$  (\*\*\*)



**Fig. 3** Mean values ( $\pm$  standard error,  $n=4$ ) of mass nutrient ratios (columns) in plant and soil solution (dissolved) per treatment. Elevated ( $eCO_2$ ) conditions shown as grey areas and P treatment as white (control) or black (+P: P addition) circles. +AM (with) and –AM (without) mycorrhizal fungi.  $F_{(1,24)}$  val-

ues of ANOVAs from linear model for response variables displayed. Significance levels:  $\leq 0.1$  ( $\cdot$ ),  $\leq 0.05$  ( $*$ ),  $\leq 0.01$  ( $**$ ),  $\leq 0.001$  ( $***$ ); only significant interactions of the treatments displayed. Treatments with different letters are significantly different ( $p \leq 0.05$ , Tukey multiple comparison test)

thanks to the role of this mycorrhizal fungi in SOM decomposition and nutrient acquisition (Fig. 1). We detected significant increases in root biomass under  $eCO_2$  and AM fungi presence, but this effect was independent of P conditions (Table 2). In low P and  $eCO_2$  conditions, higher nutrient resorption and immobilization would lead to decreases in plant C-to-nutrient ratios but we did not find a significant impacts of  $eCO_2$  on C-to-nutrient ratios and thus no clear evidence for increased P immobilisation under  $eCO_2$  was found. However, we detected significant decreases in plant N with  $eCO_2$  for AM treatments (Table 2.  $P=0.05$ , Tukey's multiple comparison). Increases in plant biomass with P addition were not observed either, but we detected increases in plant P concentration as well as the plant P pool in both shoot and roots with P addition (Table 2). As a consequence, plant C:P and N:P significantly decreased with P addition (Fig. 3).

#### Microbial communities including AM

AM fungi presence and abundance was assessed with the 16:1w5c neutral lipid. As expected, we detected higher AM fungi presence in +AM pots under control

P conditions and a significant decrease in AM fungal neutral lipid with P addition to a level approaching the uninoculated pots (AMxP:  $F_{(1,24)}=6.28$ ,  $P < 0.001$  Table SI 2). We assessed the biomass and composition of the microbial community as we expected that AM fungi would alter soil saprotrophic communities as a mechanisms to enhance SOM decomposition under low-P conditions. However, we only observed a significant impact of  $CO_2$  conditions on the absolute and relative microbial community abundance (Table SI 3) driven by marginally significant increases in Gram negative bacteria and significant increases in Fungi and Protozoa (Table SI 4).

#### Discussion

Our experiment investigated whether under low-P conditions, increased SOM decomposition and organic P cycling aided by arbuscular mycorrhizal fungi would allow for a sustained plant fertilisation effect with  $eCO_2$ . We expected to observe increases in plant biomass under  $eCO_2$  conditions and low-P availability paired with higher soil C cycling, represented in the measured SOM-derived fractions of relevant C

**Table 2** Plant biomass (shoot and roots), nutrient contents and plant nutrient pools (plant P and N content multiplied by the shoot or root biomass) per CO<sub>2</sub> condition, AM fungi and P addition treatment

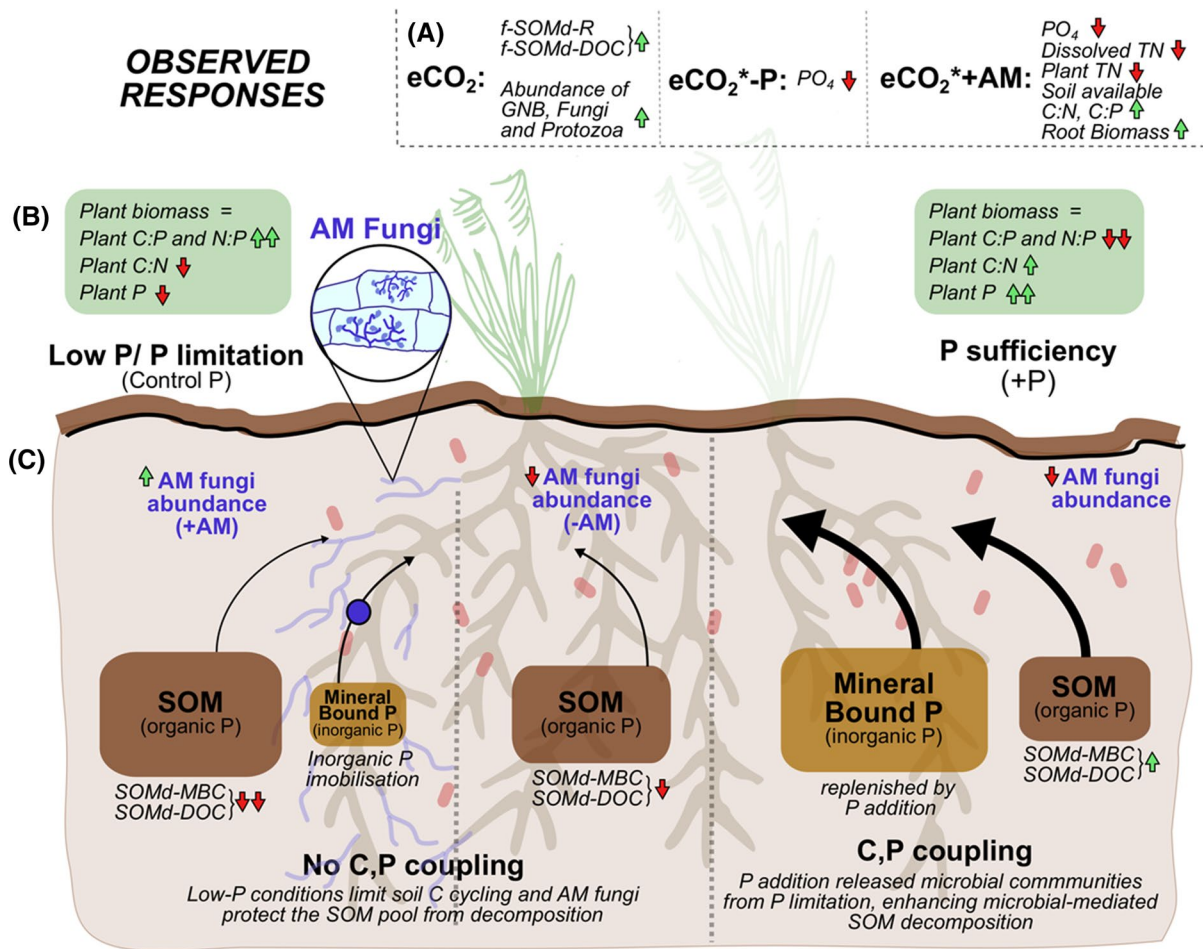
			Biomass		Nutrient Content		P pool		N pool	
			Shoot	Root	Plant N	Plant P	Shoot P	Root P	Shoot N	Root N
			g		mg/Kg		mg			
aCO <sub>2</sub>	–AM	Control	2.7 (0.2)	2.1 (0.1) <sup>ab</sup>	5.5 (0.2) <sup>a</sup>	0.5 (0.07) <sup>a</sup>	1.2 (0.2) <sup>a</sup>	0.9 (0.1) <sup>a</sup>	15.1 (0.9)	11.6 (0.7)
		+P	2.8 (0.3)	2.0 (0.3) <sup>ab</sup>	5.3 (0.1) <sup>a</sup>	0.6 (0.09) <sup>b</sup>	1.6 (0.2) <sup>b</sup>	1.3 (0.4) <sup>b</sup>	14.5 (1.5)	10.5 (1.3)
	+AM	Control	2.8 (0.1)	1.7 (0.3) <sup>a</sup>	6.0 (0.4) <sup>b</sup>	0.4 (0.07) <sup>a</sup>	1.2 (0.2) <sup>a</sup>	0.7 (0.04) <sup>a</sup>	16.8 (1.1)	10.1 (0.1)
		+P	4.3 (1.2)	1.8 (0.3) <sup>a</sup>	6.2 (0.4) <sup>b</sup>	0.6 (0.08) <sup>b</sup>	3.0 (1.1) <sup>b</sup>	1.1 (0.1) <sup>b</sup>	19.4 (2.2)	13.8 (2.2)
eCO <sub>2</sub>	–AM	Control	3.2 (0.5)	1.7 (0.2) <sup>ab</sup>	5.9 (0.3) <sup>ab</sup>	0.4 (0.01) <sup>a</sup>	1.4 (0.2) <sup>a</sup>	0.7 (0.1) <sup>a</sup>	19.3 (3.8)	9.9 (1.4)
		+P	3.2 (0.4)	2.1 (0.2) <sup>ab</sup>	5.1 (0.2) <sup>ab</sup>	0.6 (0.06) <sup>b</sup>	1.8 (0.3) <sup>b</sup>	1.2 (0.2) <sup>b</sup>	16.6 (3.1)	11.3 (1.4)
	+AM	Control	3.4 (0.3)	2.4 (0.4) <sup>b</sup>	5.5 (0.2) <sup>a</sup>	0.4 (0.09) <sup>a</sup>	1.5 (0.3) <sup>a</sup>	0.1 (0.1) <sup>a</sup>	18.1 (26)	13.3 (2.04)
		+P	3.6 (0.3)	2.4 (0.3) <sup>b</sup>	5.3 (0.2) <sup>a</sup>	0.6 (0.09) <sup>b</sup>	2.0 (0.3) <sup>b</sup>	1.2 (0.03) <sup>b</sup>	18.9 (1.2)	12.5 (1.7)
	CO <sub>2</sub>		0.3516	2.1	2.48	0.19	0.03	0.40	1.65	0.03
	P		1.0574	0.2835	2.52	<b>8.87**</b>	<b>5.62*</b>	<b>10.98**</b>	0.08	0.55
	AM		2.1422	0.1943	2.35	0.03	1.50	0.11	2.60	2.19
	CO <sub>2</sub> xP		0.9543	0.4923	1.13	0.54	1.08	0.002	0.32	0.22
	CO <sub>2</sub> xAM		0.4513	<b>5.1081*</b>	<b>5.22*</b>	0.03	0.67	2.25	0.35	0.43
	AMxP		1.241	0.4535	2.81	0.08	1.24	0.28	1.10	0.38
	PxAMxCO <sub>2</sub>		0.758	1.0528	0.04	0.08	0.91	0.49	0.01	2.67

+AM: with and –AM: without mycorrhizal fungi; aCO<sub>2</sub>: ambient CO<sub>2</sub> and eCO<sub>2</sub>: elevated CO<sub>2</sub>, C: control P, +P: P addition. Mean values (± standard error, n=4) followed by the same letter are not significantly different (p≤0.05, Tukey multiple comparison test), no letters indicate non-significant effect of the treatments. Below, results of the ANOVA from linear model (lm) for response variables (log transformed). F<sub>(1,24)</sub> values displayed with significance levels: ≤ 0.1 (.), ≤ 0.05 (\*), ≤ 0.01 (\*\*), ≤ 0.001 (\*\*\*)

cycle components (R, respired CO<sub>2</sub>; MBC, microbial biomass C and DOC, dissolved organic C). Further enhancements in soil C cycling were expected with AM fungi inoculation due to the role of this symbiotic fungi on P acquisition and SOM decomposition, expected to increase further under eCO<sub>2</sub> and low-P conditions. We did not find strong evidence suggesting that increases in soil C cycling with low-P availability and AM fungi presence supported plant growth under eCO<sub>2</sub>. Instead, we found that the impacts of the experimental factors were generally independent of each other. Contrary to our hypotheses (Fig. 1), the detected increases in root biomass with eCO<sub>2</sub> conditions and AM presence along with significant decreases in soil dissolved nutrients occurred without negative impacts on soil C cycling (Fig. 4 A). This suggests that AM fungi presence was supporting the eCO<sub>2</sub> fertilization effect and facilitating soil nutrient uptake without enhancing SOM decomposition. On the other hand, we observed significantly lower SOM-derived MBC and DOC in low-P and –AM fungi conditions, further decreasing with AM inoculation (Fig. 4). This suggests that low-P conditions

were limiting soil C cycling and that AM fungi were protecting SOM against decomposition from saprotrophic communities likely via a more efficient nutrient uptake and immobilization by AM fungi. Finally, soil C cycling (SOM-derived R and DOC) and microbial biomass increased with eCO<sub>2</sub> mainly due to the response of Gram negative bacteria, fungi and protozoa (Fig. 4), which points to an increase in microbial-derived SOM decomposition with eCO<sub>2</sub>. Taken together, our findings in this model plant-soil system demonstrate that C and P biogeochemical cycles may not become strongly coupled to sustain an eCO<sub>2</sub> fertilisation effect under low-P conditions and that AM fungi might not support increases in SOM decomposition to sustain positive plant responses to eCO<sub>2</sub>. Instead, our observations highlight the role of AM fungi protecting the SOM pool against decomposition in eCO<sub>2</sub> conditions and facilitating nutrient acquisition without negative impacts on the soil C pool.

The CO<sub>2</sub> fertilization effect is modulated by the interaction of mycorrhizal associations and nutrient availability (Treseder 2004; Terrer et al. 2016, 2019). Low-P conditions can significantly limit positive



**Fig. 4** Observed responses of relevant measured variables with the impact of (A)  $eCO_2$  alone and when interacting with P and AM fungi treatments. (B) P availability alone on plant variables (green boxes) and (C) AM fungi treatments interacting with P availability on soil organic matter-derived (SOMd) C components: R, respired  $CO_2$ ; MB, microbial biomass; DOC, dissolved organic C. Left side control P, right side +P. Green

arrows: increase; red arrows: decrease; equal sign (=): no effect. Box sizes represent the relative size of each pool. The thickness of the arrows from SOM and mineral bound P pools represent the relative flux of P from these pools to plants. The size of the circles on top of each arrow represent the relative role of AM fungi on P acquisition and SOM decomposition from these pools

responses of root biomass to  $eCO_2$  (Jiang et al. 2020), particularly in AM-associated plants (Terrer et al. 2021). In this study however, the observed increase in root biomass under  $eCO_2$  conditions occurred when AM fungi were present, regardless of P availability (Table 2; Fig. 4A). A recent study by Frew et al. (2021) found that contrary to our observations, root biomass of *Microlaena stipoides* did not increase under  $eCO_2$  and AM fungi presence. Yet, soils in this experiment were relatively high in P ( $254.75 \text{ mg Kg}^{-1} \text{ P}$ ) whereas our soils were low in P (approx.  $76.28 \text{ mg Kg}^{-1} \text{ P}$ ). Despite the lack of significant

impacts of P availability in our study, the differences observed on the impact of AM fungi on root growth under  $eCO_2$  conditions between this and the study by Frew et al. (2021) suggests that the role of AM fungi on the  $eCO_2$  fertilization effect might indeed be dependent on P availability but that the P levels used in this experiment were possibly not high enough to detect this impact. Increases in plant root biomass and root exudation have been linked to enhanced nutrient uptake due to higher plant nutrient demands under  $eCO_2$  (Iversen 2010; Dong et al. 2021) and as a response to nutrient deficiency (Shipley and Meziane

2002). In our study, AM fungi were likely facilitating positive root biomass responses to  $e\text{CO}_2$  as well as enhanced nutrient uptake as supported by the significant decreases in soil dissolved N and P with AM presence under  $e\text{CO}_2$  conditions (Table 1; Fig. 4A). Lower dissolved nutrients in soil for +AM and  $e\text{CO}_2$  without any changes on the measured SOM-derived C pools nor altered soil microbial communities, suggests that nutrient uptake by AM fungi in this plant-soil system was not linked with enhanced SOM decomposition. Finally, AM fungi also marginally decreased plant N contents in the aboveground biomass with  $e\text{CO}_2$  (Table 2; Fig. 4A). This might indicate that the obtained soil dissolved nutrients by this symbiotic fungi are not necessarily being transferred to the aboveground plant biomass. AM fungi can potentially become parasitic and reduce the symbiotic benefits for their plant host if the nutrient exchange efficiency is low, particularly that of P (Cotton 2018) as it might have been the case in our study. Alternatively, greater root biomass growth than the gains in nutrient concentration can be also driving the observed lower plant N concentrations (Dong et al. 2018).

We hypothesised that higher soil C cycling evidenced as increase SOM-derived fractions of soil C components (R, MBC, DOC) would occur with low-P availability as organic P cycling is crucial when mineral P sources are depleted (Reed et al. 2011). We were also expecting that AM fungi would significantly increase SOM decomposition under low-P conditions given their role in both enhanced P acquisition and SOM decomposition (Fig. 1). Studies have shown that AM fungi can increase SOM decomposition under elevated  $\text{CO}_2$  (Cheng et al. 2012; Kowalchuk 2012; Carrillo et al. 2015) while others show that AM fungi enhance nutrient acquisition from organic substrates, particularly under low-P availability or poor nutrient conditions (Xu et al. 2018; Bunn et al. 2019). But contrary to our expectations, we found that under low-P conditions, SOM decomposition decreased as indicated by significant lower SOM-derived MBC and DOC and that soil C cycling decreased even further with AM fungi presence, regardless of  $e\text{CO}_2$  (Figs. 2 and 4). This results demonstrate that low-P conditions as well as AM fungi presence were limiting SOM decomposition. Limitation of SOM decomposition in low-P conditions is further supported by the observed increase in soil C cycling with P

addition, which might have occurred thanks to the release of saprotrophic communities from P-limiting conditions (Cleveland et al. 2006; Mori et al. 2018). The observed lack of significant increases in SOM decomposition with AM fungi presence might be related to enhanced competition between AM fungi and saprotrophic communities over P and nutrient immobilisation in the fungal biomass with the consequent limitation of saprotrophs and their impact on soil C cycling in this plant-soil system (Zhou et al. 2019). This mechanism for SOM protection against decomposition has been attributed to AM fungi (Frey 2019) and the decreases in soil dissolved nutrients with AM fungi presence under  $e\text{CO}_2$  further support the idea that these symbiotic fungi might outcompete saprotrophic microbes in nutrient uptake.

Several mechanisms have been proposed to explain AM fungi contribution to soil C cycling and how this impact will shift with climate change factors such as  $e\text{CO}_2$  Talbot et al. 2008; Wei et al. 2019; Frey 2019; Parihar et al. 2020). Yet, multifactorial experiments assessing how nutrient availability, specifically P, impact AM fungi-mediated SOM decomposition under  $e\text{CO}_2$  are scarce. We aimed to capture tripartite interactions between P availability,  $e\text{CO}_2$  conditions and AM fungi presence, however we could not detect significant simultaneous effects of these experimental factors on the measured variables. First, the observed increases in root biomass under  $e\text{CO}_2$  and AM fungi presence occurred independently of P availability and without impacts to the SOM pool. Second, decreases in soil C cycling for low-P conditions and AM fungi presence occurred independently of  $e\text{CO}_2$  conditions (Fig. 4). The lack of significant interactive impacts of P availability on root biomass responses with AM fungi presence and  $e\text{CO}_2$  might be related with the chosen methodology for P addition. In this study, only one addition of P as triple superphosphate was done at the start of this experiment. Triple superphosphate is highly soluble in water and becomes rapidly available for plant uptake (Mullins et al. 1995) by the initial weeks of application (Ghosal and Chakraborty 2012). Moreover, plant P luxury consumption, where higher P immobilisation is not necessarily related to increases in growth (Brar and Tolleson 1975) has been observed for *Microlaena stipoides* (Robinson et al. 1993; Nie et al. 2009), which explains the observed higher P concentrations in plants with P addition without increases in biomass under any

CO<sub>2</sub> or AM fungi treatment (Fig. 4). Furthermore, the lack of impacts of AM fungi on the SOM pool under eCO<sub>2</sub> conditions, despite enhanced nutrient uptake and positive impacts on root biomass, could potentially be related with the preferential C allocation of this C<sub>3</sub> plant species to increase root biomass rather than to sustain a mutualistic relationship with AM fungi (Frey 2019). Thus, the lack of a more efficient funnelling of labile C to saprotrophic communities by AM fungi likely restricted an impact on the SOM pool in these conditions. Higher DOC is usually used as evidence for increased C allocation belowground and higher rhizodeposition with eCO<sub>2</sub> (Lukac et al. 2003; Freeman et al. 2004; Drake et al. 2011; Phillips et al. 2011) but it can also be the result of higher SOM decomposition (Hagedorn et al. 2002, 2008). Supporting the idea that not enough photosynthate was reaching soil microbial communities, our isotopic analyses allowed us to detect that the observed increases in DOC with eCO<sub>2</sub> were the result of enhanced SOM decomposition given the observed increases in SOM-derived DOC fraction (Table SI 2), rather than the result of higher C allocation to belowground microbial communities. An inefficient C transfer to AM fungi and saprotrophs under low-P conditions might also explain the lack of a significant impact of eCO<sub>2</sub> on soil C cycling in low-P+AM fungi pots. Although this well-controlled experiment was able to provide mechanistic insights into potential effects of P availability on plant and microbe responses to eCO<sub>2</sub>, it was a short-term pot study with a single grass species and relatively low replication number. Longer-term investigations using a variety of species and higher replication at the pot level to enhance statistical power are needed to more fully assess potential future responses of P-limited ecosystems to eCO<sub>2</sub>. Moreover, impacts of P availability and P manipulations are likely to be more relevant if applied at continuous rates in the form of a nutrient solution rather than as a single application.

Current understanding of the nutrient dependency of the impacts of eCO<sub>2</sub> on plant productivity has focused on N-limited ecosystems. Higher SOM decomposition under eCO<sub>2</sub> when N availability is low occurs as a mechanism to sustain nutrient supply and plant growth, with mycorrhizal fungi aiding to deliver the mined nutrients to the plants. For P-limited ecosystems however, low-P availability generally constrains ecosystem responses to CO<sub>2</sub> enrichment

(Reich et al. 2006; Reed et al. 2015; Ellsworth et al. 2017; Jiang et al. 2020) and the role of AM fungi mediating plant and soil C responses to eCO<sub>2</sub> with low-P availability is not fully clear. This pot study was carried out with natural soils and a grass species growing in a P-limited ecosystem where the EucFACE site is located, and thus it serves as a model system that allows us to infer possible responses at the ecosystem level. Previous studies at the EucFACE site have shown that higher soil C losses at this site are not driven by the impact of eCO<sub>2</sub> on soil microbial communities but are rather reliant on seasonal increases in water and nutrient availability (Castañeda-Gómez et al. 2020, 2021). Our observations from this soil-plant system also suggest that the low-P conditions in this Australian native woodland might enhance competition over nutrients with AM fungi that outcompete saprotrophic communities over these limited resources, leading to an overall decrease in SOM decomposition and limitation the eCO<sub>2</sub> fertilization effect unless a more efficient transfer of photosynthate can be allocated to the mycorrhizal symbionts which may potentially allow for enhanced AM fungi mediated soil C cycling.

In this experiment we show that the impacts of CO<sub>2</sub> enrichment and P availability on plant growth and soil C cycling were independent of each other and are not likely to become more coupled with eCO<sub>2</sub> conditions. We also demonstrate that AM fungi presence is contributing to SOM protection, likely via competition over nutrients with saprotrophic communities or maybe due to a reduced transfer of C from the plant host to AM fungi and therefore, to the saprotrophic communities. Our findings highlight that ecosystem responses to eCO<sub>2</sub> with P limitation differ from those reported for N-limited systems in that low-P conditions do not necessarily lead to higher SOM decomposition as a mechanism to sustain plant growth as usually observed for N-limited systems. Thus, inferences of the behaviour of P-limited ecosystems based on current knowledge about N-limited ecosystems are not ideal. In this way, our results also contribute to the current gap in knowledge regarding the impacts of soil C cycling with low-P availability exposed to eCO<sub>2</sub> conditions.

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**Author contributions** LCG conceived the original idea, set up and harvested the experiment, carried out the laboratory and statistical analyses and wrote the manuscripts with input from YC, JRP, and EP.

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**Data availability** Data available at Figshare: <https://doi.org/10.6084/m9.figshare.19964411.v1>.

**Code availability** Sample R code is included in the supplementary material.

## Declarations

**Conflict of interest** The authors declare there are no conflict of interest.

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