



Dynamics in plant diversity and composition on Australian alpine summits over time

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Abstract

High mountain environments are often characterised by low temperatures and short growing seasons, yet support high plant endemism and biodiversity. While such ecosystems are considered among the most vulnerable to climatic warming, the impacts of climate change on diversity and composition can be complex. To develop a deeper understanding of these dynamics, changes in vegetation over time along five alpine summits that are part of the Global Observation Research Initiative in Alpine Environment (GLORIA), were assessed including species richness, α -diversity, β -diversity, vegetation and growth form cover as well as composition. The five summits of Mount Clarke in the Australian Alps were surveyed in 2004, 2011 and 2019. Despite increases in species richness over time, there was an overall decline in diversity through biotic homogenisation across the summits. Near complete vegetation cover was recorded in 2004 but increased over the 15 years via in-filling and densification, driven by increasing cover of graminoids and shrubs. Consequently, there was also a shift in species composition which was greatest at higher elevations. The results indicate that there has been a shift towards more competitive and thermophilic composition, which may have implications for flammability in a warming and drying climate. Further assessments will be required to more fully explore the effect of climate variation from climate change, and implications for conservation of this and other alpine floras globally.

Keywords Climate change biology · GLORIA · Vegetation dynamics · Alpha and beta diversity · Biotic homogenisation · Densification

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Introduction

High mountain ecosystems are primarily determined by abiotic stressors such as low temperatures and often have short growing seasons (Körner 2003; Nagy and Grabherr 2009). They also often have complex topography that creates environmental heterogeneity and microhabitat differentiation, resulting in high endemism and biodiversity (Körner 2003, 2004; Elsen and Tingley 2015). Above the bioclimatic treeline, most alpine plants are highly specialised, slow growing and long-lived perennials that are resilient to short-term climatic oscillations, but respond to sustained climatic changes (Körner 2003; Nagy and Grabherr 2009; de Witte and Stöcklin 2010). Comparatively, the chemical and biological processes of alpine ecosystems are particularly sensitive to temperature (Körner 2003; Davidson and Janssens 2006; Lütz 2013). With faster rates of warming at higher elevations (Pepin et al. 2015; Wang et al. 2016), alpine ecosystems are considered amongst the most vulnerable and responsive to climate change (Walther et al. 2005; Seddon et al. 2016; Guisan et al. 2019). As a result, there are changes in vegetation due to overall warming but also increasing extreme climatic events such as heatwaves, droughts and fire (Gottfried et al. 1998; Grabherr et al. 2010; Dirnböck et al. 2011; De Boeck et al. 2016; Hock et al. 2019).

With increasing temperatures and longer growing seasons alpine plants can experience direct abiotic stress affecting phenology and physiology (Inouye 2008; Ernakovich et al. 2014; Fu et al. 2015). They are also experiencing increased competition as a result of the expansion of thermophilic species previously restricted to lower elevations (Lamprecht et al. 2018; Steinbauer et al. 2018). In response, there has been an increase in species richness on many alpine summits (Gottfried et al. 2012; Pauli et al. 2012). However, there is a growing ‘extinction debt’ among long-lived alpine plants as climatic thresholds are eclipsed and suitable habitat disappears, particularly for cryophilic species (Dirnböck et al. 2011; Dullinger et al. 2012; Rumpf et al. 2018, 2019). Dispersal lags and fluctuations in biotic interactions are beginning to influence the distribution, diversity and composition of alpine plant communities. Therefore, transient dynamics are likely to dominate the range of responses of alpine plants to climate change in coming decades (Gilman et al. 2010; Anthelme et al. 2014; Alexander et al. 2015, 2018).

Alpine ecosystems are ideal for monitoring changes in vegetation in response to a warming climate. With increasing elevation, the relative importance of biotic interactions such as competition and impacts from anthropogenic land use often decrease (Choler et al. 2001; Körner and Ohsawa 2005; Mark et al. 2015). They also have a global distribution while providing ‘space-for-time natural experiments’ along environmental gradients, and are comparatively simple in their biotic components (Körner 2003; Körner and Paulsen 2004; Grabherr et al. 2010; Körner et al. 2011; Pauli et al. 2015). However, short-term dynamics and natural ecological variability are often difficult to discern from longer-term trends, such as those induced by climate change (Müller et al. 2010). Therefore, multidecadal ecological monitoring using a consistent approach can assist in discerning natural variability from ecosystem dynamics due to a warming climate (Stöckli et al. 2011; Gitzen et al. 2012). Global long-term monitoring initiatives, such as the Global Observation Research Initiative in Alpine Environments (GLORIA—see <https://www.gloria.ac.at>), provide a standardised and comparable approach to assess vegetation dynamics in response to climate change by monitoring permanent vegetation plots on alpine summits (Vittoz and Guisan 2007; Pauli et al. 2015).

With over 100 study areas and 20 years of monitoring using the same protocols (Pauli et al. 2015), changes in alpine summit vegetation have been reported in Europe (Erschbamer et al. 2011; Michelsen et al. 2011; Gottfried et al. 2012; Pauli et al. 2012; Fernández Calzado and Molero Mesa 2013; Stanisci et al. 2016; Artemov 2018; Lamprecht et al. 2018; Steinbauer et al. 2018; Porro et al. 2019), Asia (Gigauri et al. 2016; Chandra et al. 2018; Salick et al. 2019; Hamid et al. 2020), North America (Malanson and Fagre 2013; Smithers et al. 2020), South America (Carilla et al. 2018) and Oceania (Venn et al. 2012, 2014). With the majority of published GLORIA research from Europe, regionally specific dynamics in diversity and edge expansions or contractions are emerging among those summits. In temperate and boreal parts of Europe, upslope migration has resulted in species enrichment and thermophilisation of assemblages (Gottfried et al. 2012; Pauli et al. 2012; Lamprecht et al. 2018), while in the drier alpine areas of Europe species richness on summits is stagnating or declining (Pauli et al. 2012; Fernández Calzado and Molero Mesa 2013).

Results from GLORIA summits outside Europe are becoming available and have provided new geographical and ecological insights about the processes driving alpine vegetation dynamics (Pauli et al. 2012; Rumpf et al. 2019). In Asia, species enrichment and increasing vegetation cover through time was more pronounced on solar aspects and higher elevation summits (Salick et al. 2019), and there is also evidence of thermophilisation (Gigauri et al. 2016). In North America, there has been increasing dissimilarity in assemblages through time, suggesting considerable species turnover on summits (Malanson and Fagre 2013). However, research assessing climate change impacts on alpine vegetation is sparser for the Southern Hemisphere whether using the GLORIA protocol or other methods (Verrall and Pickering 2020). Using the GLORIA protocols, species enrichment and increasing shrub and graminoid cover have been recorded in South America (Carilla et al. 2018) and Australia (Venn et al. 2012, 2014).

Unlike other alpine environments, many alpine summits in the Australian Alps are characterised by aeolian, deep and well-formed humus soils supporting near complete vegetation cover. The alpine climate is relatively mild with long growing seasons and warm summer temperatures (Costin 1967; Costin et al. 2000; Körner and Paulsen 2004; Green and Osborne 2012). Thus, assessing vegetation dynamics in the Australian Alps provides additional perspectives when understanding climate change impacts on alpine summits. Similar to many alpine environments globally, in the Australian Alps temperatures are increasing, there is declining snow cover, increasing variability in precipitation and changing fire regimes (Williams et al. 2008, 2014; Davis 2013; Sánchez-Bayo and Green 2013; Bradstock et al. 2014; McGowan et al. 2018; Holgate et al. 2020).

Monitoring using the GLORIA protocol across five summits in the Australian Alps, found that elevation and climatic factors, such as soil temperatures and growing season, accounted for some differences in vegetation (Pickering et al. 2008; Pickering and Green 2009). A re-survey in 2011 found that mean species richness at the whole summit scale had increased by approximately 12%, and this was largely driven by new graminoid and shrub species (Venn et al. 2012). Additionally, there were increases in overlapping vegetation cover, but the changes were most pronounced at the lowest elevation summit where graminoids and forbs more than doubled in cover. These compositional and growth form dynamics drove shifts in community trait weighted means, favouring species with more competitive traits, with the most pronounced differences at the lowest elevation summit. Across all summits, there was an increase in the cover of species with higher leaf areas (mm^2) and specific leaf area ($\text{mm}^2 \text{mg}^{-1}$) but decreases in height (mm) and leaf dry matter content (mg g^{-1}) (Venn et al. 2014).

Re-surveying in 2019 provides the opportunity to assess the response of vegetation to further changes in climate by assessing diversity and composition dynamics over more time, building upon a multicomponent approach proposed by Porro et al. (2019). Previous assessments primarily focused on species richness (Venn et al. 2012), but by also assessing changes in α -diversity and β -diversity it may be possible to identify other dynamics such as biotic homogenisation through time (Britton et al. 2009; Liberati et al. 2019). Following on from the previous analysis, assessing dynamics in vegetation and growth form cover on summits with near complete vegetation cover can also provide insights into processes such as densification and gap in-filling (Cannone and Pignatti 2014), and hence potential increases in competitive interactions as well as influential species and growth forms (Pickering and Green 2009; Venn et al. 2014).

The aim of this study incorporating data from the 2019 survey, with those from 2004 and 2011, is to extend the assessment of vegetation dynamics in response to climatic warming for the five alpine mountain summits in the Australian Alps. Specifically, the following questions will be addressed: (1) Has species richness continued to increase in line with previous assessments? (2) Has α -diversity and β -diversity fluctuated across surveys, and do they indicate vegetation homogenisation and overall biodiversity loss? (3) Are there increases in the overlapping cover of vegetation and growth forms via in-filling and densification? (4) Has composition shifted over time, and if so, what species or growth forms are influencing these changes? (5) What are the regional implications of climate drivers on changes in vegetation?

Methods

Study area

The only GLORIA study area in Australia (AU-KNP) was established in 2004 in the highest and most diverse alpine environment in Australia centred on Mt Kosciuszko (2228 m a.s.l.), the highest peak in the Australian Alps, and continental Australia (Costin 1967; Costin et al. 2000). Five ‘summits’ along a continuous ridgeline from the highest summit of Mount Clarke (2114 m a.s.l.) to near the bottom of the valley by the Snowy River (see Pickering et al. 2008; Pickering and Green 2009) (Fig. 1a). Although all five summits are above the bioclimatic treeline, which occurs around 1800 m a.s.l. (Costin 1967; Costin et al. 2000; Körner and Paulsen 2004), they vary in biotic and abiotic conditions as they range over 301 m of elevation and a horizontal distance of 2240 m (Table 1). Notably, they differ in climatic conditions including the length of the growing season (Pickering and Green 2009; Venn et al. 2014). As for vegetation structure, the two lowest summits (CL4, CL5) are dominated by shrubs while the three highest summits (CL1, CL2, CL3) are dominated by graminoids. The highest summit (CL1) also includes small tracts of windswept feldmark on the exposed western aspect (CSIRO 1972; Pickering and Green 2009).

Summits in the Australian Alps generally have little topographical relief, with many summits more dome-shaped rather than conical peaks. There is also near complete vegetation cover and high plant biomass compared to many other alpine ecosystems (Costin 1967; Costin et al. 2000; Pickering and Green 2009). Disturbance from summer cattle grazing in the area from the late 1830s to 1958 was limited as stock routes at the time avoided Mount Clarke (Costin et al. 2000; Slattery and Worboys 2020). Mount Clarke is also away from any walking tracks and there is very limited off-track walking in this area

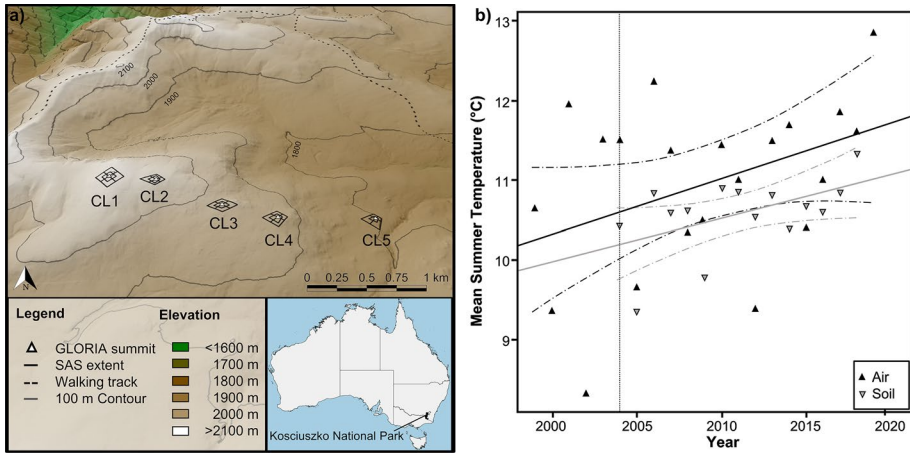


Fig. 1 a Location of the study area within Kosciuszko National Park in Australia and GLORIA summits on Mount Clarke 1–5 (CL1, CL2, CL3, CL4, CL5). b Mean summer temperatures (December to February—December from preceding year) recorded at the Bureau of Meteorology Station 071032 Thredbo AWS (1999–2019, black line; slope=0.0699, $R^2=0.157$, $t=1.879$, $F_{1-19}=3.530$, $p=0.076$) and soil temperature from Mount Clarke summits (2004–2018, grey line; slope=0.0555, $R^2=0.272$, $t=2.206$, $F_{1-13}=4.865$; $p=0.046$). Air and soil temperature are positively correlated (Pearson correlation=0.571, $p=0.026$). First year of vegetation monitoring highlighted by vertical dotted line

Table 1 Characteristics of five summits of Mount Clarke (CL1–CL5)

Characteristic	CL1	CL2	CL3	CL4	CL5
Elevation (m a.s.l.)	2114	2079	1992	1948	1813
Latitude (S)	36.4328	36.4328	36.4356	36.4356	36.4356
Longitude (E)	148.2875	148.2911	148.2961	148.3000	148.3078
5 m SAS area (m ²)	4722	3373	2860	3435	2212
10 m SAS area (m ²)	7581	8196	8551	8920	8664
Mean annual temperature (°C)	4.60	4.62	4.71	5.20	5.21
Mean annual precipitation (mm)	1969	1967	1937	1887	1885
Vegetation belt	Upper alpine	Upper alpine	Lower alpine	Lower alpine	Treeline
Plant community	FMa, TAHa	TAHa	TAHa	TAHa	TAHa
Dominant growth form	Graminoids	Graminoids	Graminoids	Shrubs	Shrubs

Long-term climatic features of the five summits were extracted from WorldClim 2 (reference period 1970–2000) at 30 s spatial-resolution (Fick and Hijmans 2017). Plant community and dominant growth form extracted from vegetation mapping (CSIRO 1972) and prior assessments (Pickering and Green 2009; Pickering et al. 2008), where ‘FMa’ refers to Feldmark (Epacris—Chionohebe alliance) and ‘TAHa’ to Tall Alpine Herbfield (Celmisia—Poa alliance) (Costin et al. 2000)

(Fig. 1a), and very limited evidence of disturbance from the small native mammals found in the region (Green and Osborne 2012; Good and Johnston 2019). Although the population of feral ungulates, such as deer and horses, has been increasing in the area (Driscoll et al. 2019; Ward-Jones et al. 2019), their impacts on these summits have been marginal with little evidence of grazing and the occasional ungulate scat (Pickering and Green 2009; Venn et al. 2012).

While there are no reliable and consistent long-term climatic data for this alpine area, or the Australian Alps in general, air and soil temperature data has more recently become available. For example, hourly air temperature data from 1999 to present are available from an all-weather station maintained by the Australian Bureau of Meteorology approximately 6.3 km to the south of the highest summit (CL1). Soil temperatures have been recorded at 2-h intervals using data loggers (Tinytag Plus—Gemini Data Loggers, Chichester England) at a depth of 10 cm on each aspect of each summit (20 data loggers) from 2004 following the GLORIA protocol (Pauli et al. 2015). These climatic data indicate an overall increase in mean summer air temperatures ($R^2=0.157$, $t=1.879$, $F_{1-19}=3.530$, $p=0.076$) and soil temperatures ($R^2=0.272$, $t=2.206$, $F_{1-13}=4.865$; $p=0.046$) by 0.070°C and 0.056°C per year respectively, beyond the annual variability (Fig. 1b).

Vegetation sampling design

Sampling followed the GLORIA protocol outlined in Pauli et al. (2015) and previous surveys of the five summits (Pickering et al. 2008; Pickering and Green 2009; Venn et al. 2012, 2014). Each summit was split into eight Summit Area Sections (hereafter referred to as “SAS”) by establishing the upper (5 m isoline) and lower (10 m isoline) summit area sections (hereafter referred to as “upper sections” and “lower sections”, respectively) from the highest point, divided by the four cardinal bearings (Fig. S1). For summits with minimal topographical relief, the terminus of the SAS extended 50 m for each section, resulting in variations in the area covered by each SAS (Fig. 1a; Table 1).

Vegetation surveys were conducted in January/February in 2004, 2011 and 2019. This involved surveying vascular plant species composition, bryophyte and lichen (non-vascular) cover, and substrate (solid rock fixed in the ground, loose rock scree, bare soil and organic matter/litter) cover in each SAS. Since this study area has near complete vegetation cover (Costin 1967; Pickering et al. 2008; Pickering and Green 2009), randomised point sampling was used to record the presence and cover of plant species in each SAS supplementing the GLORIA protocol. This involved recording all species intersected by a 1 m length of metal wire (1 mm diameter) lowered vertically toward the ground at 200 randomly located points in each SAS (Evans and Love 1957). After sampling each SAS, a thorough search for rare species that were not recorded during random point sampling was conducted, and any additional species marked as present with a nominal low cover value of 0.1%. From the random point sampling, a species list was generated for all vascular and non-vascular plants for each of the eight SAS at each of the five summits (40 SAS per survey) (Table S1). Vascular plants were characterised into growth forms following Costin et al. (2000) as graminoids (including tussock and non-tussock grasses, sedges and rushes), forbs (including erect, prostrate, creeping and rosette forbs) and shrubs (including subshrubs). Sampling in the 2011 and 2019 re-surveys were conducted ‘blind’, without referring to previous data and species lists (Vittoz et al. 2010a). Species names recorded in the field followed Costin et al. (2000) to be consistent across surveys and the main flora for the region, but have been updated here to match current taxonomy, as outlined by New South Wales Flora Online (<https://plantnet.rbg Syd.nsw.gov.au/>). Species lists were meticulously checked following surveys and where species could not be accurately and repeatedly identified in the field, genus level data were recorded.

Data processing and analysis

Prior to analysis, five species that were only recorded in one SAS and one year were removed from the species list (Table S2), Vittoz et al. (2010a), Pauli et al. (2012) and Porro et al. (2019). Then, the randomised point sampling data were used to calculate overlapping cover for all taxa (referred to as species here after, but includes genera for some as indicated above) (hits per species/200 points \times 100). The overlapping cover of growth forms (graminoids, forbs and shrubs) was calculated as the sum of all species with that growth form. To account for the varying size of SAS areas, overlapping cover of species was standardised by a corrective coefficient determined by the ratio between the SAS area where the species was recorded and the cumulative SAS area (Porro et al. 2019). Prior to analysis, the standardised overlapping vegetation and growth form cover data were log-transformed to satisfy the assumptions of parametric statistical modelling.

Before assessing the dynamics of species richness, the influence of the species-area relationship was corrected by calculating the standardised residuals of a linear mixed model (LMM) fitted on the 40 SAS species richness values in 2004, 2011 and 2019, with the area of each SAS included as a fixed effect. Due to the repeated measures and spatial structure of the sampling design, each SAS was assigned a unique identifying code ('SAS Code'—e.g. CL1N05) which was nested within the respective summits ('Summit Code'—e.g. CL1), with these nested factors included as a random effect in the model. Then, to remove the effect of SAS area on species richness, a linear model (LM) was fitted to the obtained LMM standardised residuals, following the protocol proposed in Vittoz et al. (2010b) and repeated in Porro et al. (2019). To assess species richness dynamics through time for the study area, and determine if summit or section were predictors for species richness, a LMM was conducted with 'Year' (2004, 2011 and 2019), 'Summit' (CL1, CL2, CL3, CL4 and CL5) and 'Section' (upper and lower section) as fixed effects. To test if species richness dynamics were consistent over the study area, interactions of Year, Summit and Section were also included in the model. Then, *p*-values were attained for fixed effects using the package "lmerTest" v3.1.3 (Kuznetsova et al. 2017) and differences between significant levels of fixed effects and interactions were obtained via the package "emmeans" v1.5.2-1 (Lenth et al. 2020) using Tukey's HSD adjustment for multiple comparisons.

To further investigate dynamics in α -diversity, diversity profiles for 2004, 2011 and 2019 were calculated and plotted. These profiles were generated from the standardised vegetation cover for the study area and each summit by calculating the exponential of the Rényi index (Rényi 1961; Hill 1973), which provides a series of linked diversity indices, which depend upon the parameter α and the frequencies of each species, defined as:

$$\exp(H_\alpha) = \exp\left(\frac{1}{1-\alpha}\right) \ln \sum_{i=1}^S p_i^\alpha$$

where p_i is the relative frequency of the i th species and S is the total number of species (Tóthmérész 1995). For $\alpha=0$, this function produces the total species richness, while $\alpha=1$ and $\alpha=2$ represent the Shannon index (Shannon 1948) and Simpson index (Simpson 1949), respectively. Generalised diversity profiles have seen a recent revival in ecological analyses (Porro et al. 2019), as different diversity indices have varying sensitivity to the presence of rare and abundant species, and lone indices cannot effectively assess biodiversity and community structure (Magurran 1988). However, diversity profiles have the ability to demonstrate varying measures of diversity simultaneously and permit ranking of the overall diversity of different groupings (in this case, different years) from high to

low. If a groups profile runs entirely above another, it is unequivocally more diverse. However, groups cannot be unequivocally ranked if the profiles intersect, as one group contains greater diversity of rare species whereas the other contains greater diversity of common species (Tóthmérész 1995, 1998). Thus, to rank such groupings where profiles intersect, the underlying area of the diversity profile function (hereafter referred to as “diversity area”) was calculated using the trapezoid method (Rohde et al. 2012), as proposed by Di Battista et al. (2017). Diversity area is calculated considering the entire domain and thus, does not preference either species richness nor evenness, with a higher value denoting greater biodiversity. Secondly, to understand the dynamics of varying diversity metrics through time, the following indices were calculated for the study area and each summit: Shannon index (Shannon 1948), Sheldon evenness (Sheldon 1969) and 1—Simpsons Index (Simpson 1949), which represents a Dominance index. All calculations and plotting of α -diversity were executed with the software PAST 4.03 (Hammer et al. 2001).

Following on from assessments of α -diversity, β -diversity was also calculated for 2004, 2011 and 2019 to assess if biotic homogenisation had occurred through time. Firstly, the pairwise Simpson dissimilarity (Baselga 2010) was calculated between all the SAS of the study area (40 SAS) as well as the upper and lower section (20 SAS each) for each survey. To assess differences between 2004, 2011 and 2019, the Kruskal–Wallis test with Dunn’s multiple comparison post-hoc test was used to indicate the spatial patterns of diversity without the influence of richness gradients to highlight any differences in β -diversity due to species colonisations and disappearances. Secondly, the multiple-site dissimilarity measures proposed in Baselga (2010) were calculated for the study area, which include β_{NES} (nestedness-resultant multiple-site dissimilarity), β_{SOR} (Sørensen multiple-site dissimilarity) and β_{SIM} (Simpson multiple-site dissimilarity) defined as:

$$\beta_{\text{NES}} = \beta_{\text{SOR}} - \beta_{\text{SIM}}$$

$$= \frac{\left[\sum_{i < j} \max(b_{ij}, b_{ji}) \right] - \left[\sum_{i < j} \min(b_{ij}, b_{ji}) \right]}{2 \left[\sum_i S_i - S_T \right] + \left[\sum_{i < j} \min(b_{ij}, b_{ji}) \right] + \left[\sum_{i < j} \max(b_{ij}, b_{ji}) \right]} - \frac{\left[\sum_{i < j} \min(b_{ij}, b_{ji}) \right]}{\left[\sum_i S_i - S_T \right] + \left[\sum_{i < j} \min(b_{ij}, b_{ji}) \right]}$$

where S_i is the total species richness in the SAS i , S_T is the total species richness in all the SAS and b_{ij} , b_{ji} are the number of species exclusive to SAS i and j respectively. These multiple-site measures best represent varying patterns of β -diversity dependent upon the proportion of species shared between different communities (β_{SOR}), species turnover (β_{SIM}) and spatial nestedness (β_{NES}) (Baselga 2010). Analyses of β -diversity were executed in R by means of “betapart” v1.5.2 (Baselga et al. 2020) and the Kruskal–Wallis test was conducted in the baseline package “Stats” v4.0.3 (R CoreTeam 2020).

To assess overlapping cover dynamics through time, the cover of all species was combined for ‘vegetation’ while species were also grouped by ‘growth forms’: graminoids (including rushes, sedges and grasses), forbs and shrubs (including subshrubs and dwarf shrubs) while ferns were excluded due to their rarity. Growth forms represent broad plant functional types and this type of analysis has commonly been used to assess the responses of alpine vegetation to climate change (Pickering and Green 2009; Venn et al. 2012; Porro et al. 2019). First, the standardised sum of overlapping cover was calculated for vegetation and growth forms for each of the 40 SAS for 2004, 2011 and 2019. To assess the dynamics of vegetation and growth form cover through time, LMMs were fitted to log-transformed data via the “lme4” package v1.1-23 (Bates et al. 2015), with the fixed effect ‘Year’ and ‘SAS Code’ nested within ‘Summit Code’ as the random effect to account for the spatiotemporal influence of the sampling design. F-statistics and p -values were calculated in

R via the “car” v3.0-10 package (Fox and Weisberg 2019). Where significant variation through time had occurred, subsequent LMMs were executed for each summit (8 SAS) where ‘Year’ was the fixed effect while ‘SAS Code’ was the random effect. Differences in standardised vegetation cover between years were determined via the package “emmeans” v1.5.2-1 (Lenth et al. 2020) using Tukey’s HSD adjustment for multiple comparisons.

To assess species and growth form compositional similarity among years, non-metric dimensional scaling ordinations (NMDS) were used to visualise differences in species and growth form composition for the study area. Firstly, these ordinations were generated from Bray–Curtis distance matrices of square-root transformed standardised species and growth form cover for each SAS in each year (Bray and Curtis 1957; Quinn and Keough 2003; Taguchi and Oono 2005). Secondly, influential species and growth forms driving variances between years were determined by the similarity percentage analysis (SIMPER) using a Pearson correlation co-efficient of 0.75 (Clarke 1993), and were overlaid on these ordinations as proportional vectors. Thirdly, to determine if there had been a compositional shift in species and growth forms through time, one-way analysis of similarity (ANOSIM) with ‘Year’ as the fixed effect was executed using 999 random permutations (Clarke 1993). Pair-wise comparisons using the Bonferroni adjustment were then used to determine differences between years. Lastly, ordinations were generated for each summit with subsequent ANOSIM tests to assess temporal changes at the summit level. All calculations and plotting of α -diversity were executed with the software PAST v4.03 (Hammer et al. 2001).

Results

Diversity

A total of 91 plant species from 34 families were recorded across the three surveys in the study area (Table S1). Species richness varied across summits ($p < 0.001$) (Table 2a), with 51 species recorded cumulatively at the highest summit CL1, 63 species at CL2, CL3 and CL4, and 67 species at the lowest summit, CL5. Species richness also varied through time with 73 species recorded in 2004, 76 in 2011 and 87 in 2019 ($p < 0.001$) (Table 2a). Specifically, species richness increased from 2004 to 2011 (mean difference 2.7 ± 1.0 , $p = 0.016$) and from 2011 to 2019 (mean difference 5.5 ± 1.0 , $p < 0.001$). Eleven species were lost between 2004 and 2011 and three between 2011 and 2019 (Table S3). There were 14 new species recorded in 2011 and an additional 14 new species recorded between 2011 and 2019. Between 2004 and 2019, there were 16 new species but only two disappearances.

Dynamics in species richness through time were consistent for all summits except the lowest summit, CL5, as demonstrated by the interaction between Year \times Summit ($p = 0.158$) (Table 2a) and Year for each summit individually (CL1, $p = 0.002$; CL2, $p = 0.005$; CL3, $p = 0.007$; CL4, $p < 0.001$; CL5, $p = 0.262$) (Table 2b). The number of new species was greater than disappearances between 2004 and 2019 for all summits, but was most pronounced for CL2 and CL4, with a net increase of 19 and 17 species, respectively. Sections also had a significant effect on species richness with the lower sections containing significantly more species than the upper sections (mean difference: 2.3 ± 0.6 , $p = 0.004$) (Table 2a), but variation in species richness over time was not consistent between upper and lower sections ($p = 0.004$). In particular, species richness was marginally higher in upper sections in 2004 (mean difference: 1.4 ± 1.4 , $p = 0.322$) but in 2011 (mean difference: 3.4 ± 1.4 , $p = 0.015$) and 2019 (mean difference: 5.0 ± 1.4 , $p < 0.001$), species richness was

Table 2 Results from Linear Mixed Models executed on the residuals of species richness against SAS area, with respect to Year, Summit and Section for: (a) the study area and (b) Summits (CL1, CL2, CL3, CL4, CL5)

Scale	Factor	Df	F	<i>p</i>
(a) Study area				
	Summit	4	20.206	< 0.001
	Year	2	38.038	< 0.001
	Section	1	8.768	0.004
	Summit × Year	8	1.531	0.158
	Summit × Section	4	1.290	0.280
	Year × Section	2	5.839	0.004
	Year × Summit × Section	8	0.554	0.813
(b) Summits				
CL1	Year	2	9.463	0.002
	Section	1	6.139	0.023
	Year × Section	2	3.017	0.074
CL2	Year	2	7.288	0.005
	Section	1	5.072	0.037
	Year × Section	2	0.263	0.072
CL3	Year	2	6.711	0.007
	Section	1	0.197	0.662
	Year × Section	2	0.285	0.755
CL4	Year	2	38.259	< 0.001
	Section	1	0.023	0.882
	Year × Section	2	6.662	0.007
CL5	Year	2	1.443	0.262
	Section	1	1.346	0.261
	Year × Section	2	2.919	0.080

Degrees of freedom (Df), F statistic (F) and significance level (*p*) are displayed for each model, with significant *p*-values (< 0.05) italicised

greater in the lower sections. At the summit level, species richness was greater in the lower sections of the two highest summits (CL1, *p* = 0.023; CL2, *p* = 0.037).

There was a net decrease in α -diversity through time across the study area with declines in Shannon Index and Sheldon Evenness but an increase in the Dominance Index (Table 3). The Dominance Index (0.19) was highest in 2011 which corresponds to the lowest value for the Shannon index (2.45). However, divergent results were recorded across the summits. There was a net increase in diversity for CL2 as demonstrated by the Shannon Index, while there was a decrease in diversity for all other summits. Considering Sheldon Evenness, there was a net decrease for all summits but there was a net decrease in Dominance Index for CL2 and CL5. These results are supported by the comparable diversity profiles for 2004, 2011 and 2019 for the study area (Fig. 2), where a marked decline in profiles demonstrates the asymmetrical relative abundances of plant species, specifically in 2011. The profiles for the study area could not be unequivocally ranked for α -diversity as they intersected each other, yet the diversity area suggests that diversity was marginally higher in 2019 than 2004, which was largely driven by species enrichment through time. From 2004 to 2019, diversity metrics were variable among summits with values for the Shannon Index decreasing across all summits except CL2, while values for the Dominance Index increased only at CL1, CL2 and CL4. However, there was a consistent decline in Sheldon

Table 3 Summary of α -diversity metrics (Shannon Index, Sheldon Evenness and Dominance Index) and Diversity Area that represents the area under the diversity profiles (see Fig. 2) calculated for 2004, 2011 and 2019

Scale	Shannon index			Sheldon evenness			Dominance index			Diversity area		
	2004	2011	2019	2004	2011	2019	2004	2011	2019	2004	2011	2019
Study area	2.60	2.45	2.54	0.18	0.15	0.14	0.14	0.19	0.16	51.98	47.08	<i>52.11</i>
CL1	1.97	1.86	1.79	0.22	0.19	0.18	0.24	0.30	0.31	26.95	24.74	24.48
CL2	1.88	1.80	2.03	0.22	0.18	0.17	0.24	0.27	0.23	26.10	25.79	<i>31.50</i>
CL3	2.40	1.91	2.25	0.25	0.16	0.19	0.20	0.33	0.24	<i>37.65</i>	27.57	35.91
CL4	2.15	2.13	2.11	0.23	0.19	0.17	0.17	0.20	0.20	34.39	34.17	<i>35.04</i>
CL5	2.05	2.10	2.01	0.18	0.19	0.16	0.23	0.21	0.20	31.25	<i>33.67</i>	33.04

The greatest Diversity Area for each scale is italicised

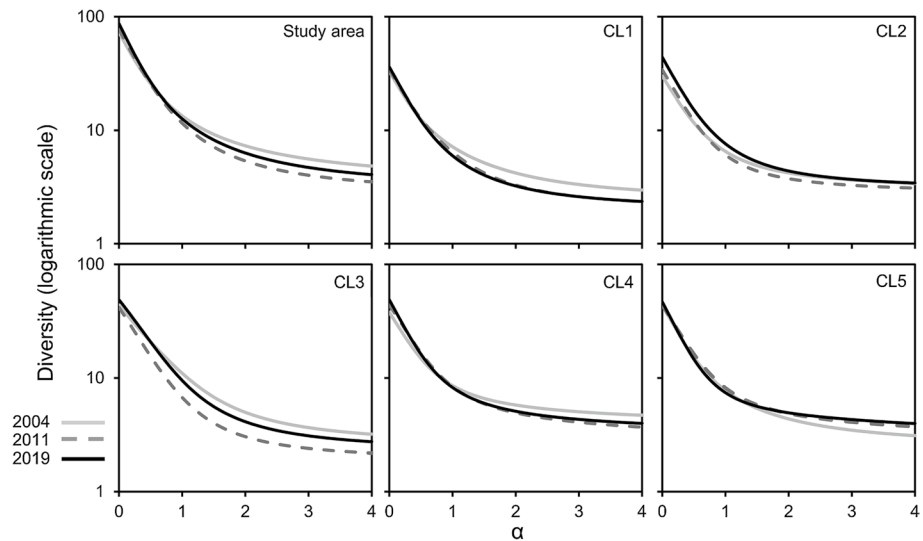


Fig. 2 Diversity profiles for the study area and for each summit for 2004, 2011 and 2019 produced by calculating the exponential of the Renyi index, where metrics of diversity depends upon a parameter α

Evenness across all summits between the first and last survey. Similarly, there were varied patterns in diversity profiles for the five summits (Fig. 2), but a marked increase in species richness for all summits. Since there were no clear differences for any of the profiles at any summit, the diversity area indicates diversity was greatest in 2004 for CL1 and CL3, in 2011 for CL5 and in 2019 for CL2 and CL4.

Considering the dynamics of the ‘pairwise β_{SIM} ’, the overall dissimilarity in the study area decreased significantly between 2004 and 2019 ($p < 0.001$) with the majority of change occurring between 2004 and 2011 ($p < 0.001$) (Fig. 3). Dynamics at the section level were also significant, with a decrease in dissimilarity between 2004 and 2011 for lower sections ($p = 0.026$), and between 2004 and 2019 ($p = 0.032$) (Table S4). However, there was no difference for the upper section between 2004 and 2011 nor 2011 and 2019, but a marked decrease between the first and last survey ($p = 0.006$). Conversely, dynamics

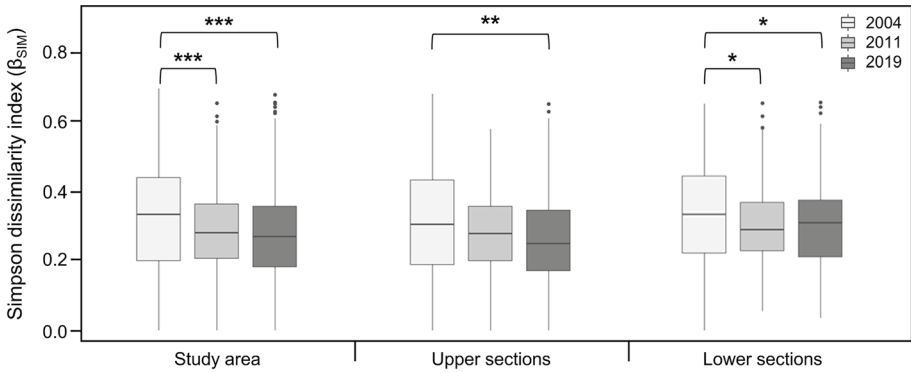


Fig. 3 Pairwise Simpson dissimilarity indices (‘pairwise β_{SIM} ’ as proposed by Baselga (2010)) for the study area upper and lower sections in 2004, 2011 and 2019. Asterisks indicate significant changes between years ($*p < 0.05$; $**p < 0.01$; $***p < 0.001$)

of the ‘multiple-site’ β -diversity metrics were marginal (Table 4). Between 2004 and 2019 for the study area, values for ‘multiple-site B_{SOR} ’ decreased while ‘multiple-site B_{SIM} ’ increased, indicating that β -diversity dynamics was influenced less by shared species and more by species turnover through time. The ‘multiple-site B_{SOR} ’ value was greater for the lower section only in 2004, with the upper sections displaying greater dissimilarity in 2011 and 2019. The ‘multiple-site B_{SIM} ’ values across all years were higher in the lower sections compared to the upper sections yet the lower sections have increased in dissimilarity between 2004 and 2019 while the upper sections have decreased in dissimilarity in the same period.

Cover and composition

In 2004 nearly all of the summit areas were covered by vegetation. However, overlapping cover of vegetation, graminoids and shrubs increased in the study area from 2004 to 2019 ($p < 0.001$) (Table 5a), although there were differential responses among summits (Table 5b), particularly when changes occurred. Graminoid cover increased markedly over the 15 years ($p < 0.001$) but most of the change occurred between 2004 and 2011 (mean difference: 7.8, $p < 0.001$) rather than between 2011 and 2019 (mean difference: 3.5, $p = 0.002$) (Table S5). Shrub cover also increased over the 15 years ($p < 0.001$) but unlike

Table 4 β -diversity multiple-site dissimilarity indices, as proposed by Baselga (2010), for the study area and for the upper and lower sections in 2004, 2011 and 2019

	Multiple-site β -diversity index								
	B_{NES}			B_{SOR}			B_{SIM}		
	2004	2011	2019	2004	2011	2019	2004	2011	2019
Study area	0.1136	0.1199	0.1050	0.9706	0.9691	0.9660	0.8570	0.8492	0.8610
Upper section	0.1316	0.1402	0.1332	0.9702	0.9700	0.9674	0.8386	0.8298	0.8343
Lower section	0.1104	0.1167	0.1007	0.9710	0.9683	0.9647	0.8606	0.8516	0.8640

Table 5 Results from Linear Mixed Models executed on the standardised sum of overlapping vegetation and growth form cover, with respect to Year for: (a) the study area and (b) Summits (CL1, CL2, CL3, CL4, CL5)

Scale	Growth form	Factor	Df	F	<i>p</i>
(a) Study area	Vegetation	Time	1–39	143.919	<0.001
	Graminoids	Time	1–39	111.974	<0.001
	Forbs	Time	1–39	2.045	0.136
	Shrubs	Time	1–39	48.225	<0.001
(b) Summits					
CL1	Vegetation	Time	1–7	20.766	<0.001
	Graminoids	Time	1–7	18.205	<0.001
	Forbs	Time	1–7	3.871	0.046
	Shrubs	Time	1–7	6.711	0.009
CL2	Vegetation	Time	1–7	48.670	<0.001
	Graminoids	Time	1–7	53.941	<0.001
	Forbs	Time	1–7	4.080	0.040
	Shrubs	Time	1–7	7.990	0.005
CL3	Vegetation	Time	1–7	35.405	<0.001
	Graminoids	Time	1–7	42.455	<0.001
	Forbs	Time	1–7	3.099	0.077
	Shrubs	Time	1–7	21.840	<0.001
CL4	Vegetation	Time	1–7	87.291	<0.001
	Graminoids	Time	1–7	40.989	<0.001
	Forbs	Time	1–7	2.474	0.120
	Shrubs	Time	1–7	16.925	<0.001
CL5	Vegetation	Time	1–7	74.413	<0.001
	Graminoids	Time	1–7	107.044	<0.001
	Forbs	Time	1–7	6.080	0.013
	Shrubs	Time	1–7	9.556	0.002

Degrees of freedom (Df), F statistic (F) and significance level (*p*) are displayed for each model, with significant *p*-values (<0.05) italicised

graminoid cover, it was mostly between 2011 and 2019 (mean difference: 3.3, $p < 0.001$) with less variation between 2004 and 2011 (mean difference: 1.9, $p = 0.025$) (Table S5). In contrast to graminoids and shrubs, forb cover was low with no consistent change over time (Fig. 4).

There were clear patterns in cover among individual summits over time (Fig. 5). Overlapping vegetation cover increased across all summits between 2004 and 2019, but for the three lower elevation summits (CL3, $p < 0.001$; CL4, $p < 0.001$; CL5, $p < 0.001$), most of the change occurred between 2004 and 2011, while the majority of change for the two higher summits (CL1, $p = 0.005$; CL2, $p < 0.001$) occurred between 2011 and 2019. There was a similar pattern in graminoid cover, with increases across all summits over the 15 years, but no increases for lowest three summits between 2011 and 2019, while for the two higher summits there were increases between 2011 and 2019. Shrub cover also increased across all summits over the 15 years, but it was only significant between 2011 and 2019 for some summits (CL1, $p = 0.049$; CL3, $p < 0.001$; CL4, $p = 0.012$).

There was a significant shift in species ($R = 0.1134$, $p < 0.001$) and growth form composition ($R = 0.0422$, $p = 0.004$) through time for the study area (Fig. 6). Species composition differed between 2004 and 2011 ($R = 0.1451$, $p < 0.001$) but not between 2011 and

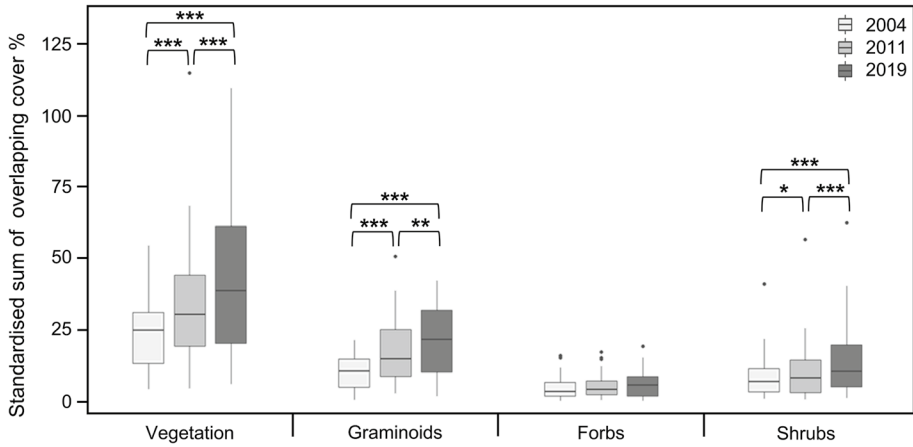


Fig. 4 Pairwise comparisons of the standardised sum of overlapping cover of vegetation and growth forms the study area. Asterisks indicate significant changes between years (* $p < 0.05$; ** $p < 0.01$; *** $p < 0.001$)

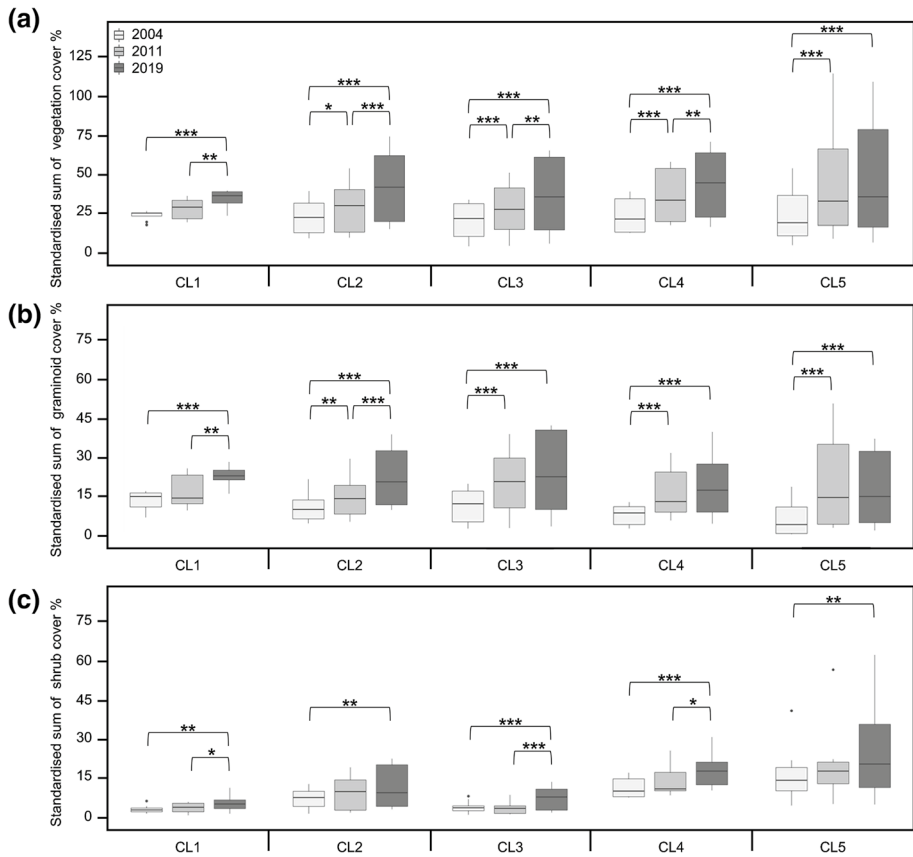


Fig. 5 Pairwise comparisons of the standardised sum of overlapping cover of vegetation and growth forms for the five summits. Asterisks indicate significant changes between years (* $p < 0.05$; ** $p < 0.01$; *** $p < 0.001$)

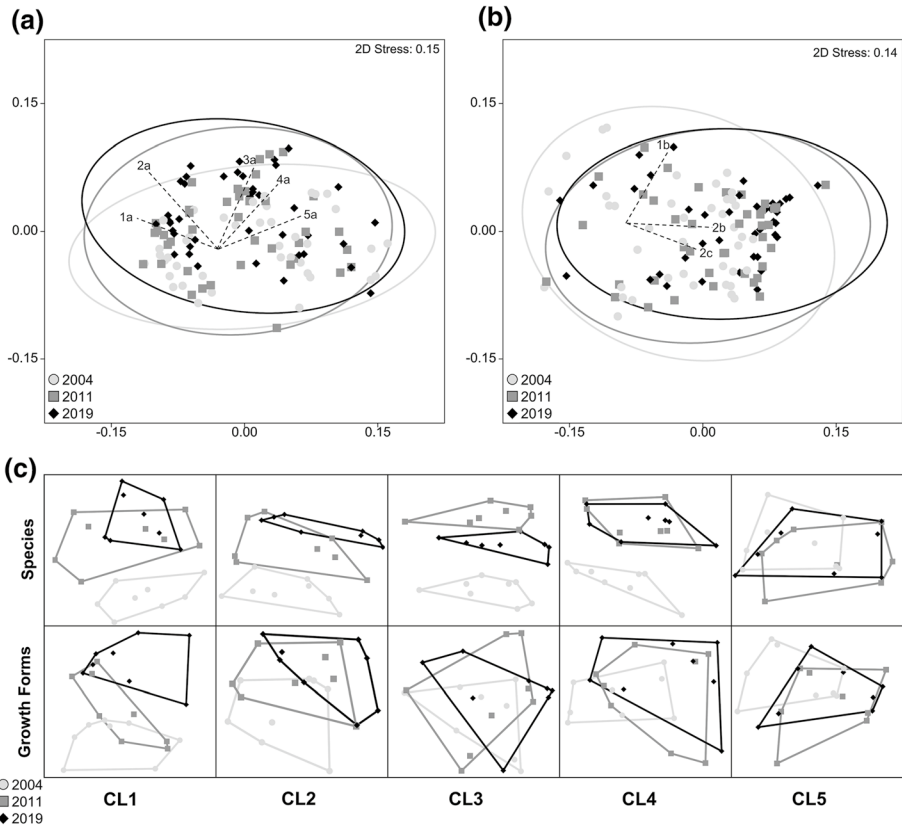


Fig. 6 Non-metric dimensional scaling ordinations using Bray–Curtis similarity index on square-root transformed standardized overlapping cover data with 95% confidence ellipses of **a** species ($R=0.1134$, $p<0.001$), **b** growth form ($R=0.0422$, $p=0.004$) composition shifts across surveys for the study area. Vectors demonstrate Pearson correlations (>0.75) of the relative influence of species (1a—*Trisetum spicatum* subsp. *australiense*, 2a—*Poa* Sp., 3a—*Empodisma minus*, 4a—*Nematolepis ovatifolium*, 5a—*Kunzea muelleri*) and growth forms (1b—Shrubs, 2b—Graminoids, 3b—Forbs) of compositional shift through time. Shifts in **c** species and growth form composition with maximum extent convex hulls for each summit

2019 (Table S6). As for growth forms, composition was similar between 2004 and 2011 and between 2011 and 2019, but there were differences over the 15 years ($R=0.0100$, $p<0.001$). There was a shift in species composition for the four higher summits (CL1, $p<0.001$; CL2, $p=0.029$; CL3, $p=0.005$; CL4, $p=0.011$), but changes in growth form composition was only significant for the highest summit (CL1, $p=0.047$). The majority of the variation in species composition occurred between 2004 and 2011, with shifts for CL1 ($p=0.003$), CL3 ($p=0.002$) and CL4 ($p=0.032$). However, for growth form composition, the only difference was for CL1 between 2011 and 2019 ($p=0.012$).

Discussion

Discerning diversity dynamics

Species enrichment has continued on the summits of Mount Clarke in the Australian Alps consistent with previous assessments (Venn et al. 2012), and with summits in the Central European Alps also using the GLORIA protocol (Erschbamer et al. 2011; Lamprecht et al. 2018; Porro et al. 2019). This contrasts with results from drier alpine environments, such as in the Mediterranean (Pauli et al. 2012) and the Andes (Carilla et al. 2018), where there were declines in species richness. At a summit scale, the maximum increase in species richness for Mount Clarke in the Australian Alps was on the second highest summit (47 to 59 species recorded on CL2 from 2011 to 2019) but was less than the prior assessments (41 to 56 species on CL4 from 2004 to 2011) (Venn et al. 2012). Species richness increased in all but the lowest summit (CL5) over the three surveys, suggesting that species enrichment may be stagnating at lower elevations, which appears to be also occurring in the Altai mountains of Russia (Artemov 2018). Furthermore, the increase in mean species richness for Mount Clarke summits suggests a longer-term trend rather than a short-term fluctuation (Erschbamer et al. 2011; Venn et al. 2012). Between the first and last survey, there were 16 new species and only two of these, *Epilobium gunnianum* and *Oreomyrrhis brevipes*, are restricted to this alpine zone (Wright et al. 2010). Therefore, the increases in species richness were predominately driven by species with broader environmental tolerance including subalpine and montane zones. This is consistent with monitoring at a continental scale from Europe, where species richness and abundance increases were driven by thermophilic species from lower elevations responding to climate warming (Gottfried et al. 2012). Climate induced changes in summit vegetation have also been suggested by Lamprecht et al. (2018), Porro et al. (2019) and Steinbauer et al. (2018), among others. These changes on Mount Clarke summits may also be climate-induced, as conditions in the area have warmed, growing seasons lengthened, and snow cover reduced, while the area has experienced limited other types of disturbance such as from trampling or grazing for nearly 60 years (Slattery and Worboys 2020). Furthermore, the disappearance and re-appearance of species between surveys may be explained by observer error including misidentifications (Vittoz and Guisan 2007; Vittoz et al. 2010a; Futschik et al. 2020), as well as drought legacy effects where certain species may have experience drought-induced mortality but following an increase in precipitation in 2010 (Holgate et al. 2020), competitive species may have established themselves more readily (Venn and Morgan 2009).

There were differing patterns in α -diversity for the Mount Clarke study area. From 2004 to 2019, evenness decreased but there were fluctuations in dominance and diversity (Shannon index). However, disentangling dynamics of species richness and α -diversity has proven difficult and led to mixed conclusions (Britton et al. 2009; Ross et al. 2012). In response, diversity profiles were used here as they embody various α -diversity metrics and dynamics through time. These profiles more effectively encapsulate multifaceted ecological processes driving vegetation dynamics in response to climate change, such as upslope migration, leading or trailing edge range dynamics, localised extinctions, competition and densification (Rumpf et al. 2018, 2019; Porro et al. 2019). Across all summits, diversity was greatest in 2019, but this can largely be attributed to increasing species richness, as denoted by the larger area under the diversity profiles at low α values. Conversely, α -diversity was lowest in 2011, which may due to the dominance of competitive species following a pulse of recruitment at the end of the millennial drought (Venn and Morgan

2009). The second highest summit (CL2) displayed the greatest increase in α -diversity, which was largely driven by 19 new species from 2004 to 2019. Excluding this summit, diversity dynamics seem to follow a trend with elevation, similar to the findings of Porro et al. (2019). The greatest loss of α -diversity over the 15 years occurred at the highest summit (CL1), followed by the third highest summit (CL3), while the two lower summits became more diverse. The loss of diversity over time at higher elevations may be explained by the combination of upslope migration of more thermally-tolerant plants (Artemov 2018; Steinbauer et al. 2018) and the in-filling and expansion of competitive species resulting in densification (Cannone and Pignatti 2014; Rumpf et al. 2018).

Considering β -diversity, the decline in the values of ‘pairwise β_{SIM} ’ through time indicates biotic homogenisation across the study area. These results are similar to those for northern Apennines summits in Italy (Porro et al. 2019), but β -diversity is yet to be investigated for other GLORIA summits. Biotic homogenisation is ongoing for the Mount Clarke study area and was first detected in the lower sections of summits in 2011, but appears to have now progressed upslope to the upper sections potentially demonstrating the influence of elevation on climate-induced vegetation dynamics (Artemov 2018; Rumpf et al. 2018; Smithers et al. 2020). The upper sections were less affected by upslope migration and densification, but over the 15 years there was a strong homogenisation signal, demonstrating the importance of longer-term monitoring (Walther et al. 2005; Alexander et al. 2018; Futschik et al. 2020). The dynamics in β -diversity are mostly attributed to species turnover rather than shared species or nestedness, as indicated by the greater values of ‘multiple-site β_{SIM} ’ (Baselga 2010). Several species were replaced but few were lost from the study area, which is similar to changes documented on Italian summits (Porro et al. 2019) and is consistent with previous assessments of the same summits (Venn et al. 2012).

Densification of vegetation driving compositional dynamics

Unlike many other alpine areas globally, there is near complete vegetation cover across the Mount Clarke summits, as is the case for many other alpine summits in the Australian Alps (Costin 1967; Costin et al. 2000; Pickering and Green 2009; Venn et al. 2017). Therefore, dynamics in vegetation cover were the result of in-filling or densification with more species recorded in areas where vegetation cover is complete but less dense, such as inter-tussock spaces and under shrub canopies rather than colonisation of bare ground (Cannone and Pignatti 2014). There was a considerable increase in graminoid and shrub cover through time, as found in the previous assessment of the summits (Venn et al. 2014) as well as for other alpine and tundra environments (Björk and Molau 2007; Cannone et al. 2007; Wilson and Nilsson 2009; Ababneh and Woolfenden 2010; Myers-Smith et al. 2011; Elmendorf et al. 2012; Gao et al. 2016; Amagai et al. 2018). More specifically, the increasing presence of thermophilic graminoids and shrubs is well documented throughout the GLORIA network (Erschbamer et al. 2009, 2011; Grabherr et al. 2010; Gottfried et al. 2012; Fernández Calzado and Molero Mesa 2013; Gigauri et al. 2016; Stanisci et al. 2016; Lamprecht et al. 2018; Hamid et al. 2020). The process of ‘grassification’ of alpine environments can alter biotic interactions and increase intraspecific and interspecific competition (Choler et al. 2001; Callaway et al. 2002; Alexander et al. 2015). The expansion of graminoids on alpine summits in the Australian Alps may reflect their capacity to out-compete forbs and recruit within the canopy of senescing shrubs following considerable precipitation (Williams and Ashton 1988; Jarrad et al. 2008). Similarly, shrubification has implications for ecosystem processes such as localised snow cover and associated hydrological dynamics

(Körner 2003; Myers-Smith et al. 2011). With an increase in shrub cover and height, they can outcompete graminoids and forbs by forming dense thickets with closed canopies that transform microclimatic conditions and biotic interactions in tundra and alpine environments (Myers-Smith et al. 2011; Ballantyne and Pickering 2015; Cáceres et al. 2015). Shrub cover also affects the duration of snow cover and hence affects hydrological dynamics (Costin et al. 2000; Körner 2003; Myers-Smith et al. 2011; Venn and Green 2018), as well as the size and extent of subnivean spaces that provide critical overwintering habitat for many alpine animals (Green and Osborne 2012).

The graminoid and shrub driven densification of vegetation across the Mount Clarke summits also resulted in shifts in composition. Although many new species were forbs, they remained scarce and their contribution to assemblages was minimal. Earlier assessments of this study area demonstrated a similar pattern of graminoid and shrub driven compositional dynamics (Venn et al. 2012, 2014). Across a wide range of alpine summits monitored using the GLORIA protocol, shifts in composition have also been detected in Europe (Erschbamer et al. 2009; Fernández Calzado and Molero Mesa 2013; Malanson and Fagre 2013; Lamprecht et al. 2018), Asia (Gigauri et al. 2013; Hamid et al. 2020) and South America (Carilla et al. 2018). The majority of changes across the Mount Clarke summits occurred between the first two surveys. Such changes may have been due to lagged responses to climate change and increased competitive interactions as suggested for other alpine summits (Erschbamer et al. 2009; Alexander et al. 2018), but also may have occurred in response to end of a major drought in the region (Venn and Morgan 2009; Holgate et al. 2020) and/or as a consequence of observer error compounded through time (Vittoz and Guisan 2007; Vittoz et al. 2010a; Futschik et al. 2020). There were shifts in species composition in all but the lowest elevation summit, indicating potential stabilisation of assemblages. Since this summit (CL5) was already dominated by dense thickets of thermophilic shrubs with closed canopies found at lower elevations in 2004 (i.e. *Kunzea mullerei* and *Prosanthera cuneata*) (Wright et al. 2010), colonisation by other species and growth forms may have been restricted. Similarly, the divergent responses in diversity, cover and composition of each summit may be explained by varying thresholds for change as determined by the underlying alpine plant communities and differing abiotic factors (Pickering and Green 2009). The exclusion of other species amidst increasing and densifying shrub cover has been established in other alpine ecosystems in Australia, with a positive feedback loop between warming and shrub proliferation (Scherrer and Pickering 2005; Wahren et al. 2013; Camac et al. 2017). Although there are many benefits from assessing growth form or functional type dynamics, Australian alpine plants have shown species-specific varied responses to warming (Hoffmann et al. 2010). Therefore, continued long-term monitoring and more rigorous phenological and experimental warming research is required to more fully understand species-specific responses to climate change in these environments.

Influence of climatic drivers and regional implications

Long-term dynamics in vegetation are often masked by short-term responses and variability to cyclical climate influences (Greenland 1999; Müller et al. 2010). The main climate drivers that cause variability in the study area are the El-Niño-Southern Oscillation (ENSO), Indian Ocean Dipole (IOD) and Southern Annular Mode (SAM). The interaction of these largely independent drivers leads to considerable climatic variability, where warmer temperatures and drought conditions are often the result of a negative IOD, positive SAM during summer and El-Niño (Holgate et al. 2020). Thus, interpreting long-term

changes in vegetation requires consideration of the influence of shorter-term climatic variation. Specifically, the baseline data for the Mount Clarke summits in 2004 was collected during relatively cool conditions, but also in the middle of the ‘Millennial Drought’ (Van Dijk et al. 2013), which persisted in the area until 2010 (Holgate et al. 2020), while the second survey occurred a year after drought-breaking precipitation. This shorter-term climatic variation may explain some of the vegetation dynamics seen, such as an increase in dominance and species richness, as herbaceous vegetation can rapidly respond to an influx of precipitation after the end of drought conditions (Scherrer and Pickering 2005; Venn et al. 2014). The third survey was during a period of below average precipitation but higher temperatures. Thus, difficulties arise from attempting to deduce the influence of climate drivers versus climate change on ecological processes but continued long-term monitoring aims to address this ambiguity (Walther et al. 2002; Müller et al. 2010).

The results for Mount Clarke summits and other studies of Australian alpine areas (e.g. Gallagher et al. 2009; Pickering et al. 2014; Annandale and Kirkpatrick 2017; Kirchoff 2020) indicate that the Australian alpine vegetation is changing, and that it is likely that at least some of the change is associated with warming conditions. Such changes are of particular concern as the area is considered to be one of the most at risk from climate change in Australia (Hughes 2003; Laurance et al. 2011; Hoffmann et al. 2019), and contains many endemic animal and plant species (Costin et al. 2000; Green and Osborne 2012). Australian alpine environments occur as isolated fragments with restricted elevation gradients as treelines are close to summits (Costin 1967; Costin et al. 2000). Thus, there is limited potential for upslope migration, particularly for cryophilic and endemic species as conditions warm, increased variation in precipitation and declining snow cover (Hennessy et al. 2003; Sánchez-Bayo and Green 2013; Holgate et al. 2020).

Of particular issue for Australian alpine regions, is increasing fire severity and frequency with climate change (Hennessy et al. 2005; Zylstra 2018), as these regions have historically experienced few fires (Dodson et al. 1994; Williams et al. 2008; Kirkpatrick and Bridle 2013). While some alpine plant communities have recovered from fire in the past, specialist communities and cryophilic species had limited capacity to recover post fire (Kirkpatrick et al. 2010; Venn et al. 2016). In the past decades, fires in the Australian Alps are becoming hotter, more frequent and covering larger areas with fires burning into alpine areas previously thought to have low flammability. In some cases, there has already been repeat burning of alpine vegetation within a few years (Williams et al. 2008; Camac et al. 2013, 2017; Verrall 2018; Zylstra 2018). It is possible that the densification and infilling of vegetation driven by graminoids and shrubs across the Mount Clarke summits over 15 years, if more widespread, may also increase connectivity, landscape flammability and increased fuel loads as conditions become warmer and drier (Williams et al. 2006; Fraser et al. 2016).

Conclusion

This study found dynamics in alpine vegetation diversity, cover and composition over a 15-year period using a global, standardised monitoring protocol. Across the Mount Clarke summits, air and soil temperatures have increased steadily since the beginning of the millennium. In the same period, α -diversity decreased at higher elevations despite considerable species enrichment. There were no observable patterns in species enrichment along the elevation gradient summits, yet due to upslope migration, lower summit area sections

displayed greater shifts in diversity in comparison to upper summit area sections. More broadly, there was evidence of ongoing biodiversity loss via biotic homogenisation with reductions in β -diversity. There were also marked increases in vegetation cover via in-filling and densification, driven primarily by graminoids and shrubs. Although new species contributed little to overall vegetation cover and composition, there were shifts in species composition. These compositional shifts were more pronounced at higher elevations but the majority of changes occurred between the first and second survey (2004–2011). While this study provides some support for climate-induced dynamics in alpine vegetation, further monitoring is required to disentangle responses to climatic variability versus climate change. However, with clear differences across elevation gradients in Australia's highest and most diverse alpine environment, there is cause for concern. Consequently, as the climate continues to warm, it is likely that there will be further transitions in vegetation assemblages towards more thermophilic species currently characteristic of lower elevations with significant implications for conservation and ecological process in the region including fires.

Supplementary Information The online version contains supplementary material available at <https://doi.org/10.1007/s10531-021-02171-1>.

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Author contributions BV contributed to data collection in 2019, data analysis and writing. KG contributed to project design and data collection in 2004, 2011 and 2019. CMP contributed to project design, data collection in 2004, 2011 and 2019 and revised the manuscript.

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Data availability Data is available at <https://dataverse.harvard.edu/dataset.xhtml?persistentId=doi:10.7910/DVN/VWBZHN>.

Declarations

Conflict of interest All authors declare that they have no conflict of interest.

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