TECHNICAL ARTICLE

Study on the Height of the Mining‑Induced Water‑Conducting Fracture Zone Under the Q₂ Loess Cover of the Jurassic Coal Seam in Northern Shaanxi, China

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Abstract

A method to calculate the height of a water-conducting fractured zone (HWCFZ) was developed based on the plate and shell theory, and the development of the HWCFZ in bedrock and $Q₂₁$ loess strata is discussed in detail. First, the subsidencedefection curve equation of the overlying stratum is theoretically derived, and then the ultimate defection and free space height of rock strata are calculated. Moreover, the strata tensile strain is calculated by using integral calculus. In addition, the failure state of the rock is analyzed by comparing the theoretically calculated tensile strain with the experimentally measured yield tensile strain, allowing one to attain the maximum value of HWCFZ. This approach was tested at the Jinjitan coal mine; the theoretically predicted, experimentally measured, and numerically computed maximum HWCFZ values were 189.5, 187.3, and 188.5 m, respectively, demonstrating the accuracy of the proposed method. These results are highly signifcant for safe and environment-friendly coal mining in northwest Shaanxi, China.

Keywords Plate and shell theory · Yield tensile strain · Field measurements · Numerical simulations

Introduction

Coal is China's major source of energy in China. In 2016, coal production and consumption supplied more than 70% of its energy (Liu et al. [2018;](#page-10-0) Zhang et al. [2009](#page-10-1)) and despite the continuous growth of renewable energy technologies, the total production and consumption of coal is expected to still exceed 50% of China's total energy resources in 2030 (Comprehensive Group on Consultation and Research in the Field of Energy [2015](#page-9-0)). The Jurassic coalfeld in northern Shaanxi is China's largest coalfeld, containing a third of China's total coal reserves (Yi [2008\)](#page-10-2), and the coal seam is thick and shallowly buried. This means that large-scale and high-intensity mining is liable deform and destroy overlying strata, decreasing their capacity to supply water to the area.

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This has resulted in several issues, such as water inrush accidents in mines, industrial and agriculture problems, reduced vegetation, intensifed desertifcation, and environmental concerns (Adam and Paul [2000;](#page-9-1) Andreas and Nikola [2011](#page-9-2); Hu et al. [2019](#page-10-3); Shao et al. [2014](#page-10-4); Wang and Hang [2010\)](#page-10-5).

Based on the deformation and failure characteristics of the overlying strata, there are three distinct zones of disturbance in overburden strata, i.e. the caved, fractured, and continuous deformation zones (Palchik [2003](#page-10-6)). The height of the water-conducting fractured zone (HWCFZ) is the sum of the heights of the fractured and caved zones (Adhikary and Guo [2014\)](#page-9-3). Therefore, a method to accurately predict the maximum value of the HWCFZ is of great theoretical and practical signifcance.

Scholars have proposed different methods, including empirical formulas, statistical analysis, theoretical calculations, feld measurements, and numerical simulations, to predict the maximum value of the HWCFZ. For instance, Liu [\(1995\)](#page-10-7) carried out regression analysis of experimental data of diferent coal mines in north China and obtained an empirical formula. However, the geological conditions and mining technology are signifcantly diferent in the eastern and western parts of China. Therefore, the derived empirical formulas are not applicable to coal mines in northern

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Shaanxi (Fan and Zhang [2015\)](#page-9-4). Hu et al. [\(2012](#page-9-5)) summarized the non-linear statistical relationship between the HWCFZ and the mining height, hard rock lithology ratio, coalface length, mining depth, and other mining factors, and obtained a formula to predict the HWCFZ based on diferent mining factors. Zhao and Wu ([2018\)](#page-10-8) predicted coal seam roof HWCFZ by random forest regression analysis. However, the selection of infuencing factors, i.e. the determination of weight index and weight analysis, is based on the judgment of decision-makers, which can be infuenced by several factors. Liu et al. [\(2015\)](#page-10-9) analyzed the movement characteristics of overlying strata and e obtained a theoretical formula to predict the HWCFZ based on diferent mining factors. Majdi et al. [\(2012](#page-10-10)) considered the upper-pressure relief zone after coal seam mining and derived fve mathematical methods to estimate the maximum value of the HWCFZ. Based on the failure characteristics of diferent lithologic rocks, Wang et al. ([2012\)](#page-10-11) established a computational mechanics model to predict the HWCFZ using the strain and stress as the indices of water fow in soft and hard rock strata. Che et al. ([2016](#page-9-6)) proposed a novel approach to determine by constructing a damage range model of overlying strata, based on a theory of rock fragmentation and expansion and the basic principles of stress balance in underground space after mining. Zhang et al. [\(2017](#page-10-12)) applied system theory, and explained the process of overburden failure and generation of water-conducting fssure zone from an energy conservation viewpoint. This allowed the approximate value of the HWCFZ to be obtained by a system input–output method.

However, these theoretical methods are based on the development of the HWCFZ in bedrock, and does not consider the development of the HWCFZ in overburden soil. Sui et al. ([2015](#page-10-13)) studied the changes in flushing fluid consumption, water level, drilling speed, air suction, and gas emission during drilling and used these to establish a rock quality designation (RQD) and systematically analyze the fracture development and the HWCFZ. However, in the case of large drillholes and well leakage, their method is unlikely to work. Lv et al. [\(2016\)](#page-10-14) used borehole imaging to collect image information of borehole interiors and determined the development characteristics, connectivity and fragmentation of mining rock mass and obtained a HWCFZ value however, the proposed method can only be used in boreholes without water holes and casings. Zhang et al. (2018) (2018) targeted the recovery and exploitation of coal resources retained after fully mechanized top-coal caving mining, and used UDEC numerical simulation software and a multivariate non-linear regression analysis method to obtain a mathematical relationship for calculating the HWCFZ for block coal mining. And, based on the geological conditions and flling methods of a copper-lead–zinc mine and a coal mine, Liu et al. [\(2009\)](#page-10-16) and Cheng et al. ([2017](#page-9-7)) obtained HWCFZ values by fuid–solid coupling calculations in FLAC3D.

 Q_{21} loess is widely distributed in northwest China. It has a small elastic yield limit, obvious structural strength and collapsible deformation, and its mechanical properties are strongly afected by these characteristics. The development of the HWCFZ under the cover of the Q_{21} loess is the root cause of typical water disasters in China's western mining area. Therefore, we developed an HWCFZ prediction method based on the plate and shell theory (Huang [1987\)](#page-10-17) in theoretical analysis. The Jinjitan coal mine was used as a case study; the development of the HWCFZ was developed in overlying bedrock and the Q_{21} loess was investigated using theoretical calculations, feld measurements, and numerical simulations. The feasibility and accuracy of the theoretical calculation method was then verifed by comparing the theoretically predicted results with feld measurements and numerically computed results.

Study Region

The Jinjitan coal mine is located in Yulin city, Shaanxi province (Fig. [1](#page-2-0)a). The main landforms in the study area are wind-blown sand with scattered loess ridges and valley terrace landforms. The average annual rainfall in Yulin (2005–2015) can be estimated as 406.18 mm, and the rainfall in June, July, August and September accounts for about two-thirds of the total annual rainfall. Moreover, the average annual evaporation can be estimated as 1775 mm; thus, the region is arid-semiarid and its ecological environment is already fragile (Liu et al. [2019](#page-10-18)). The Quaternary phreatic aquifer mainly receives infltration recharge from atmospheric precipitation, which is restricted by rainfall conditions, topography, and other factors. Given the high evaporation rate and coal mining, water resources are short and desertifcation is intensifed. Therefore, the study area has become a key area for soil and water conservation and desertifcation control.

The strike and inclined length of the J coalface is 5.1 km and 300 m, respectively, and the mining area is 1.53 km² (Fig. [1](#page-2-0)b). The coal seam dip is sub-horizontal with a mining thickness of 8.6 to 11.2 m, at a depth of 259 to 284 m. The stratigraphic sequence of the study area is shown in Table [1](#page-2-1).

Due to the shallow burial depth of the coal seam in the study area, the HWCFZ can penetrate through the bedrock into the Q_{21} Q_{21} Q_{21} loess layer (Fig. 2). The Q_{21} loess layer is the main relative aquifuge in the study area, and controls the leakage of phreatic water. The HWCFZ development decreases the waterproof performance and enhances water leakage of the phreatic aquifer in the sand layer of the Quaternary Salawusu formation.

Fig. 1 Location of study area: **a** Jinjitan colliery in Yulin city, **b** No. J coalface of Jinjitan

Theoretical Analysis

Before mining, the surrounding rock is in a state of stress balance. However, the stress of the surrounding rock is redistributed by mining. Based on the redistributed stress, the overlying strata can be divided into decompression, pressurization, and pressure stabilization zones (Qian and Shi [2003\)](#page-10-19). Herein, the strata in the decompression area above the coalface after mining was the research focus. According to the J coalface borehole data, the ratio of the thickness of overlying stratum to length (or width) of the coalface is far less than 1/5. Therefore, the overlying bedrock can be considered a thin plate with four fxed edges, and the fracture mechanism of the bedrock can be analyzed using the plate and shell theory (Hang et al. 1987). If the direction along the coalface is considered as the *x*-axis, a thin plate mechanical model of overlying bedrock strata,

Fig. 2 HWCFZ in bedrock and Q_{21} loess

Fig. 3 Mechanical model of the stratum: **a** Top view, **b** A-A section

(b)

which is inclined to the *y*-axis and makes an intersection point at *o*, can be established, as shown in Fig. [3.](#page-3-1)

The equations and boundary conditions of the proposed mechanical model can be given as:

$$
\frac{\partial^4 w}{\partial x^4} + 2 \frac{\partial^4 w}{\partial x^2 \partial y^2} + \frac{\partial^4 w}{\partial y^4} = \frac{q(x, y)}{D} \tag{1}
$$

$$
(w)_{x=0} = (w)_{x=a} = (w)_{y=a} = (w)_{y=b} = 0
$$
\n(2)

$$
\left(\frac{\partial w}{\partial x}\right)_{x=0} = \left(\frac{\partial w}{\partial x}\right)_{x=a} = \left(\frac{\partial w}{\partial y}\right)_{y=0} = \left(\frac{\partial w}{\partial y}\right)_{y=b} = 0 \tag{3}
$$

where $q_{(x,y)}$ represents the load on each overlying bedrock and *D* refers to the fexural stifness of the plate. The value of *D* can be given as: D = $\frac{Eh_i^3}{12(1-v^2)}$. To satisfy the differential equation of elastic surface of the thin plate, the defection function (w) can be represented as a triple series $(Xu\ 2006)$ $(Xu\ 2006)$ $(Xu\ 2006)$.

$$
w = \sum_{m=1}^{\infty} \sum_{n=1}^{\infty} A_{mn} \sin \frac{m\pi x}{a} \sin \frac{n\pi y}{b}
$$
(4)

where *m* and *n* represent positive integers (1, 2, 3…) and *Amn* can be determined from Eq. [\(1](#page-3-2)). Therefore, $q_{(x,y)}$ can also be expanded into double trigonometric series at intervals [0, a] and [0, b], as given below:

$$
q_{(x,y)} = \sum_{m=1}^{\infty} \sum_{n=1}^{\infty} a_{mn} \sin \frac{m\pi x}{a} \sin \frac{n\pi y}{b}
$$
 (5)

Herein, a_{mn} can be obtained from the orthogonality of the trigonometric function, $\sin \frac{m\pi x}{q}$ on [0, a]. Furthermore, by mul-tiplying Eq. ([5\)](#page-3-3) by $\sin \frac{m\pi x}{a}$, $\sin \frac{n\pi y}{b}$ and integrating *X* and *Y* in [0, a] and [0, b] intervals, we can obtain a relationship for a_{mn} :

$$
a_{mn} = \frac{4}{ab} \int_{0}^{a} \int_{0}^{b} q_{(x,y)} \sin \frac{m\pi x}{a} \sin \frac{n\pi y}{b} d_x d_y
$$
 (6)

By substituting Eq. (4) (4) and Eq. (5) (5) into Eq. (1) (1) , a simplified version of Eq. (6) (6) can be obtained:

$$
A_{mn} = \frac{a_{mn}}{\pi^4 D \left(\frac{m^2}{a^2} + \frac{n^2}{b^2}\right)^2}
$$
 (7)

Furthermore, by substituting *Amn* into Eq. ([4\)](#page-3-4), a generalized solution of the defection of four-sided clamped plates can be obtained:

$$
w = \frac{1}{\pi^4 D} \sum_{m=1}^{\infty} \sum_{n=1}^{\infty} \frac{a_{mn}}{\left(\frac{m^2}{a^2} + \frac{n^2}{b^2}\right)^2} \sin \frac{m\pi x}{a} \sin \frac{n\pi y}{b}
$$
(8)

A mathematical expression for q_i can be obtained from the key stratum theory (Qian [2003\)](#page-10-21) by assuming that the overlying bedrock in the coal seam is subjected to a uniformly distributed load (q_i) , as given below:

$$
q_i = \frac{E_i h_i^3 (\gamma_1 h_1 + \gamma_2 h_2 + \dots + \gamma_n h_n)}{E_1 h_1^3 + E_2 h_2^3 + \dots + E_n h_n^3}
$$
(9)

Moreover, Eqs. $(6, 7)$ $(6, 7)$ $(6, 7)$ can be used to obtain the deflection function of the overlying bedrock slab under a uniform load:

$$
w = \frac{16q_i}{\pi^6 D} \sum_{m=1}^{\infty} \sum_{n=1}^{\infty} \frac{1}{mn \left(\frac{m^2}{a^2} + \frac{n^2}{b^2}\right)^2} \sin \frac{m\pi x}{a} \sin \frac{n\pi y}{b}
$$
(10)

One should note that the given series rapidly converges and only the 1st term needs to be satisfed (Huang [1987](#page-10-17)). Therefore, when $m = n = 1$, the deflection function of the thin plate can be given as:

$$
w = \frac{16q_i}{\pi^6 D \left(\frac{1}{a^2} + \frac{1}{b^2}\right)^2} \sin \frac{\pi x}{a} \sin \frac{\pi y}{b}
$$
(11)

However, at the central positions, i.e. $x = \frac{a}{2}$ and $y = \frac{b}{2}$, of the thin plate, the defection of the thin plate reached its maximum value: $w_{max} = \frac{16q_i}{\pi^6 D(\frac{1}{a^2} + \frac{1}{b^2})^2}$.

The ultimate span a_m and b_m of the thin plate depends on the physical and mechanical properties, such as thickness and load of rock stratum. The failure condition of the thin plate can be given as: $\sigma_t > [\sigma_t]$. The a_m and b_m can be derived from the tensile strength of the plate (Liu et al. [2018\)](#page-10-0):

$$
a_m = \frac{ab_m}{b} \tag{12}
$$

$$
b_m = \sqrt{\frac{\sigma_t h^2}{6kq}}
$$
\n(13)

Herein, K refers to the shape coefficient of the thin plate, as given below:

$$
k = 0.00302 \left(\frac{a}{b}\right)^3 - 0.03567 \left(\frac{a}{b}\right)^2 + 0.13953 \left(\frac{a}{b}\right) - 0.05859
$$
\n(14)

Thus, the ultimate defection function of bedrock thin plate in layer *i* can be given as:

$$
w_{i-max} = \frac{16q_i}{\pi^6 D \left(\frac{1}{a_m^2} + \frac{1}{b_m^2}\right)^2} \sin \frac{\pi x}{a_m} \sin \frac{\pi y}{b_m}
$$
(15)

HWCFZ Analysis

Compared with the horizontal state, the bent and subsided stratum is elongated after bending. Herein, the interlayer tensile deformation of strata is used to express this phenomenon. There is a direct relationship between the water conductivity and tensile deformation of strata (Gao et al. [2012](#page-9-8)). With the increase of strata elevation, the subsidence curve of the strata became gentler, and both the tensile deformation and fracture

width of the strata gradually decreased. When the strata location reached a certain height, the generated micro-cracks were no longer conductive between the upper and lower layers due to the small amount of tensile deformation of rock strata. Hence, the development of water-conductive cracks is stopped, and the upper limit of the HWCFZ is attained.

In the process of overburden failure, only the strata within the range of the HWCFZ breaks up and produces fragmentation. The HWCFZ attained its maximum value when the coefficient of fragmentation of overburden collapse approaches the residual coefficient of fragmentation. The free space height below each caving stratum in the HWCFZ range (Qian [2003\)](#page-10-21) can be given as:

$$
\Delta_i = M - \sum_{j=1}^{i-1} h_i (k_j - 1) \tag{16}
$$

where *M* refers to the mining thickness of coal seam and k_j represents the residual dilatancy coefficient of stratum j , which is taken as a reference (Liu [2018](#page-10-22)).

When $y = b_m/2$, the arc length of the deflection curve of the clamped thin plate in $[0, a_m]$ interval can be calculated from the curve integral method (Tongji University [2014\)](#page-10-23)

$$
S_{(x)} = 2\int_{0}^{\frac{a_m}{2}} \sqrt{1 + w(x)^2} d_x \tag{17}
$$

Herein, the tension ratio of the bedrock layer can be given as:

$$
\epsilon = \frac{S_{(0 < x < a_m)} - a_m}{a_m} \tag{18}
$$

The tensile strain (ε) of each bedrock stratum was calculated and compared with the yield tensile strain $\varepsilon_{[t]}$ of the corresponding bedrock. If $\Delta_i > w_{i-max}$ and $\epsilon > \epsilon_{[t]}$ are tenable at the same time, the fssures between the bedrock conduct water and HWCFZ should continue to increase. On the other hand, if the HWCFZ steadily develops and does not continue to develop upward, this means that the HWCFZ develops to the maximum position. Hence, the maximum value of the HWCFZ can be given as: $H_{max} = h_b$ (Fig. [2](#page-3-0)).

When the coal seam is mined along the coalface to a certain length, the HWCFZ may penetrate the overlying bedrock and develop into the Q_{21} loess layer. Herein, the Q_{21} loess layer is considered as a weak rock layer, similar to the bedrock. Following the abovementioned steps, the free space height and actual tensile strain under the Q_{21} loess layer are calculated, the top interface of the HWCFZ development in the Q_{21} loess layer is analyzed, and the maximum value of HWCFZ is again given by: $H_{max} = h_b + h_L$ (Fig. [2\)](#page-3-0).

Mechanical Theory Calculations

Determination of the Yield Tensile Strain of a Rock Stratum

In the current study, the direct uniaxial tensile tests of sandstone, sandy mudstone, mudstone, and Q_{21} loess were carried out using an MTS810 material testing machine and XL2101B5 strain tester (Fig. [4](#page-5-0)). Then, the experimental results were statistical analyzed to obtain the tensile strength and yield tensile strain of each stratum overlying

Fig. 4 Laboratory test: **a** MTS810 material testing machine and XL2101B5 strain tester; **b** sandstone after tensile failure

the J coalface (Table [2](#page-5-1)). The experimentally measured yield tensile strain of the various rock strata revealed that the yield tensile strain of hard, medium-hard, soft, and weak strata were $\approx 0.05\%, \approx 0.10\%, \approx 0.20\%, \text{ and } > 0.40\%, \text{ respectively.}$ These results are consistent with the previously reported yield tensile strain data of the strata (Gao et al. [2012\)](#page-9-8).

Case Calculation and Analysis

Based on the drilling core data and indoor direct tension test, the overburden characteristics and physical and mechanical

parameters of the J coalface are given in supplemental table S-1. Taking the fourth layer as an example, the theoretical calculations are carried out according to the abovementioned method.

First, the height of the fourth layer stratum free space is calculated using Eq. (16) (16) :

$$
S_4 = 8.6 - 11 \times (1.05 - 1) - 0.2 \times (1.035 - 1)
$$

$$
- 0.25 \times (1.025 - 1) = 8.03
$$
 m

Second, the maximum defection of the fourth layer stratum is calculated using Eq. [\(11\)](#page-4-1):

$$
w_{4-\text{max}} = 4 \times 10400.3 \times 2.045^4 \div \pi^6 \div 154.98 = 4.91 \text{ m}
$$

Third, the actual tensile strength of the fourth layer stratum is calculated using Eq. (18) :

$$
\varepsilon_4 = (2.048 - 2.045) \div 20.45 = 0.0016
$$

Finally, the maximum deflection, free space height, actual tensile strain, and yield tensile strain of the same layer were compared to determine whether the stratum was broken or not. Because $S_4 = 8.03 > w_{4-max} = 4.91$ and $\epsilon_4 = 0.0016 > \epsilon_{[4]} = 0.1\%$, the fourth layer stratum be regarded as broken. The failure state of each bedrock layer is obtained by the same calculation steps (supplemental table S-1).

To accurately determine the top interface of the HWCFZ during its passage through the overburden bedrock strata to the Q_{21} loess, the 39.1 m thick Q_{21} loess was subdivided into eight sublayers, with thicknesses of 3.6, 6.2, 3.8, 5.2, 4.6, 6.7, 5.4, and 3.6 m, based on the lithological description in the borehole column.

Following the above theoretical calculation steps, the free space height, maximum defection and actual tensile strain of each layer were calculated. The results are compared and summarized in supplemental table S-1. Thus, the top interface of the HWCFZ in the Q_{21} loess can be accurately determined. Based on these theoretical calculations, it can be concluded that the HWCFZ penetrates the overburden bedrock stratum and reaches the first small layer of the Q_{21} loess, where it stabilizes and does not continue to develop further. Hence, the HWCFZ attained its maximum value, as given by: $H_{max} = h_b + h_L = 185.92 + 3.58 = 189.5$ m.

Field Measurements

Field monitoring of the dynamic deformation and failure of overlying strata was carried out using Brillouin optical time-domain refectometry (BOTDR) distributed optical fber sensing technology. The sensors were pushed into the monitoring hole before and after the mining of the J coalface, with a range of 400 m. The data were collected

1–2 times per day. The data were processed by an AV6419 Analyzer and Origin 2016 software. Figure [5](#page-6-0) shows the strain distribution of the metal kieso sensing (MKS) optical cable in the JK2 borehole, where the *x*-axis represents strain and *y*-axis refers to the buried depth of the MKS (Liu [2018](#page-10-22)). Herein, the positive and negative strains indicate the tensile and compressive loads at the cable, respectively. The strain was measured after every 0.1 m. The curve shows that the peak value of tensile strain was $+12,200$ με. However, after pushing through the borehole, the coalface was compressed again in the range of 77 to 200 m. When the overlying strata were relatively stable, the tensile strain and compressive strain of the cable were found to be $+ 1600 \mu\varepsilon$ and − 2750 με, respectively. With continuous advancement of the coalface, the overlying strata was gradually destroyed from the bottom to top.

To obtain the dynamic height of the overlying stratum, the deformation state of the overlying stratum was analyzed by comparing the strain of the optical fber with the yield strain of the rock mechanics experiment. Figure [6](#page-7-0) shows the distribution characteristics of deformation and failure zones of the overlying stratum when the distance of the panel and cable was – 52, 6, 46, 98, and 152 m. Under the infuence of mining stress and deformation of the rock surrounding the roadway, the strata mainly renders tensile failure. Moreover, the area of tensile-strain-induced failure increased with the advancement of the coalface.

Based on the theoretical discussion and the position of the fber optic tension–compression transition, it can be concluded that the rock under the HWCFZ experiences tensile load, whereas the rock above the HECFZ is mainly under compressive stress. Moreover, the experimental results of the MKS optical cable in JK2 borehole demonstrated that the buried depth of the maximum development position of the HWCFZ was 76.5 m, whereas the coal seam thickness was 11 m and the depth of the coal seam floor was 274.8 m.

Fig. 5 The strain distribution of MKS cable

Therefore, the maximum value of the HWCFZ can be given as: $H_{max} = 274.8 - 76.5 - 11 = 187.3$ m.

To explore the change of permeability in the fractured rock mass caused by mining, a feld water pressure test was carried out in the afected and unafected areas in working face J using ground drilling. This method is generally used to evaluate the permeability of the rock and soil with poor permeability. In rock with good permeability (such as the rock in the fracture zone), the fow can suddenly increase several times and the pressure cannot be increased. Thus, the permeability of the main lithology (argillaceous sandstone, fne sandstone, and siltstone) in the upper, middle, and lower parts of the area not affected by mining was measured (see supplemental table S-2).

Numerical Simulations

Based on the physical and mechanical parameters of the overlying strata (supplemental table S-1), a soil-rock model was established according to the original roof strata and the development of the HWCFZ under the Q_{21} loess. According to the mining plan of the coalface in the study area, the inclined length of the simulated mining face was 300 m and the planned simulated mining distance was 260 m. To eliminate the infuence of boundary efects, 120 m protective coal pillars were placed on both sides of the model trend and 50 m protective coal pillars were placed on both sides of the model transverse. Thus, the model was set to be 500 m long and 400 m wide. To enhance stabilization of the proposed model, the Aeolian sand in the upper layer of the actual stratum was simplifed to 0.4 MPa force, which acted on the top surface. The model height was 251 m. The boundary conditions were applied on four sides and the bottom of the model. The horizontal displacement in the lateral direction was fixed, whereas both vertical and horizontal displacements of the bottom boundary were fxed. The top of the model was allowed to move freely. The downward gravity stress and lateral gradient stress enhanced the consistency of the proposed model with actual mining conditions.

In general, the failure range of overburden is determined by distinguishing its plastic zone according to its strength and rock mechanics parameters. The range of the plastic zone gradually changes with the mining of coal seam and can be used as the overburden's failure zone (Tu and Yu [2004](#page-10-24); Wang et al. [2005](#page-10-25)). The dynamic change in the plastic zone of the overlying stratum as the coalface advances is presented in Fig. [7](#page-8-0). The failure characteristics of the overlying stratum is summarized below:

- 1. In the early stage of mining, tensile damage occurred above the goaf, and the height of the overburden failure remained low. The failure height of the overburden gradually increased as mining advanced. As the bearing limit of the rock beam above the goaf was not reached, the rock stratum does not collapse, the vertical displacement does not become obvious, and the rate of overburden failure remains slow. However, stress concentration occurred at both ends of the coalface and shear failure zones gradually appeared in the coal seam and surrounding rock in front of the opening eye of the coalface.
- 2. When the coalface advanced to ≈ 60 m, a small number of strata collapsed, shear failure occurred at both ends of the roof strata, and tensile failure occurred in the middle of strata. Furthermore, the rate of overburden failure obviously accelerated with continued mining. A few strata broke as the coalface advanced to ≈ 110 m.
- 3. As the coalface advanced to \approx 140 m, tensile failure points gradually appeared in the strata. Moreover, the height of overburden failure gradually increased and the

Fig. 7 The dynamic change of the plastic zone of the overlying strata in the model, propulsion distance of coalface: **a** 20 m, **b** 70 m, **c** 120 m, **d** 170 m, **e** 220 m, **f** 260 m

failure points become connected. When the coalface further advanced to \approx 170 m, the rock beam and overburden rock were destroyed. As the coalface advanced from 140 to 210 m, the change in the plastic zone showed that overburden failure was not developing in a layer-bylayer scheme, as occurred initially, followed by a gradual increase and coalescence of local failure points.

4. When the coalface advanced to \approx 220 m, the HWCFZ developed into the Q_{21} loess and the bending and subsidence of the upper stratum compacted the cracks in the separated loess layer. As a result, the HWCFZ did not continue to develop upwards and remained stable or exhibited a relatively downward trend.

Based on the dynamic changes of the plastic zone, the HWCFZ-associated dynamics changes are presented in Fig. [8.](#page-9-9) The maximum value of HWCFZ was found to be 188.5 m.

Results and Discussion

The thick and shallow Jurassic coal seam in northern Shaanxi, China is prone to the formation and penetration of a HWCFZ through the overburden bedrock due to the large-scale and high-intensity mining, which can further develop into the Q_{21} loess layer. As the empirical formulae of HWCFZ, summarized by Liu ([1995](#page-10-7)), are not applicable to Northern Shaanxi (Fan and Zhang [2015\)](#page-9-4), diferent reports on HWCFZ do not consider overlying soil (Che et al. [2016](#page-9-6); Liu et al. [2015](#page-10-9); Majdi et al. [2012](#page-10-10); Wang et al. [2012;](#page-10-11) Zhao and Wu [2018\)](#page-10-8). Therefore, a theoretical method was developed based on the plate and shell theory to predict HWCFZ values and the development of the HWCFZ in the bedrock and Q_{21} loess. Field measurements and numerical simulations were carried out to demonstrate the accuracy and reliability of the proposed theoretical model.

Fig. 8 The dynamic change of the HWCFZ in the model

One should note that the mechanical model of a thin plate, which was proposed by Tie et al. ([1977\)](#page-10-26), is based on the plate and shell theory. However, compared with the mechanical model of a masonry beam (Miao et al. [2011\)](#page-10-27) and plastic hinge theory (Zhou et al. [2018\)](#page-10-28), the mechanical model rendered comparable results to the actual situation of the overlying stratum. This was the frst time that the mechanical thin plate model was used to analyze the development of the HWCFZ in a soil layer. Herein, $S_i > w_{i-m\alpha x}$ and $\epsilon > \epsilon_{[t]}$ were used simultaneously as a criterion of rock breakage. By using the J working face of the Jinjitan mine as a case study, the presented theoretical method predicted a maximum HWCFZ value of 189.5 ± 2.2 m.

Furthermore, feld monitoring of dynamic deformation and failure of overlying strata based on BOTDR distributed optical fber sensing technology, reveal that the experimentally measured maximum HWCFZ value was 187.3 m. One should note that optical fber monitoring is a dynamic process. With the advancement of the borehole in the coalface, the stress recovery of the overlying stratum was slightly reduced. Hence, the experimentally measured HWCFZ value was slightly less than the theoretically predicted HWCFZ value. The numerical simulation results rendered a similar maximum HWCFZ value of 188.5 ± 1.2 m.

Conclusions

In summary, based on theoretical calculations, a novel calculation method was developed that accurately predicts HWCFZ values. The main research fndings are summarized below:

- 1. The main steps of the calculation method are as follows: First, the subsidence-defection curve equation of the overlying stratum was theoretically derived, and then, the ultimate defection and free space height of the rock strata were calculated. Then, the strata tensile strain was calculated using integral calculus. Finally, the failure state was analyzed by comparing the theoretically calculated tensile strain with the experimentally measured yield tensile strain of each stratum, thus attaining the maximum HWCFZ value.
- 2. The theoretically predicted HWCFZ value was comparable to the experimentally measured and computed HWCFZ values, which demonstrated the reliability and accuracy of the proposed calculation method. These results are highly signifcant for safe, eco-friendly coal mining in the western mining areas of China, where conservation of groundwater resources is critical.
- 3. The proposed theoretical calculation method provides the basis for HWCFZ prediction, water hazard control and coal mining under water-containing in Jurassic coalfeld of Northern Shaanxi, China.

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