ORIGINAL PAPER

Prediction method and research on characteristics of surface subsidence due to mining deeply buried Jurassic coal seams

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Received: 20 May 2022 / Accepted: 17 September 2022 / Published online: 27 September 2022 © Springer-Verlag GmbH Germany, part of Springer Nature 2022

Abstract

Jurassic coal seams are a common geological feature in the mining areas of western China. Deeply buried Jurassic coal seams were mined in the Hujierte mining area in western China. The Hujierte mining area has a mining depth of more than 600 m, and the alluvium in the mining area is primarily composed of eolian sand. Understanding the characteristics of mining subsidence is crucial for preserving the ecological environment of the mining area. Surface subsidence caused by coal mining has different characteristics depending on geological conditions. However, the characteristics of surface subsidence caused by mining deeply buried Jurassic coal seams have rarely been studied. Therefore, using a coal mine in the Hujierte mining area as an example, the observed data of surface movements in the mining area are analyzed in this study. The results reveal that (1) when the ratio of mining width to depth are 0.44, 0.88, and 1.31, the subsidence factors are 0.129, 0.427, and 0.429, respectively, and (2) although the mining degree of the study area is larger than that of some coal mines in the eastern and western regions of China, the subsidence factor of the study area is still lower than that of these coal mines. Moreover, this study proposes a method for predicting the surface subsidence based on the T-distribution. At the edge of a subsidence basin, the fitting accuracy of the model is better than that of the probability integration method.

Keywords Jurassic coal seams · Eolian sand · Characteristics · Surface subsidence · Prediction method

Introduction

The western mining area of China has large coal reserves (Liu et al. [2019](#page-14-0)). As coal resources in the eastern region are depleted, the western region will be China's main coal production area in the future (Chen et al. [2019\)](#page-14-1). The Shendong and Shanbei coal bases are important coal resource development bases in the mining area of western China (Xu et al. [2021](#page-15-0)). They are located in the border areas of China's Inner Mongolia and Shaanxi provinces. Originally, China primarily mined shallow Jurassic coal seams in these two

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coal bases (mining depth of approximately 50–420 m). In view of the growing energy demand, the development of shallow-buried Jurassic coalfields is incapable of meeting the society's demands. Therefore, China planned a Hujierte mining area in the Shendong coal base. The Hujierte mining area is located in the Mu Us desert, and the alluvium is primarily composed of eolian sand. Deep Jurassic coal seams are mined in the mining area (mining depth over 600 m); the strata of the mining area consist primarily of Jurassic–Cretaceous layers.

Coal mining can cause severe ground subsidence. Mininginduced subsidence can lead to the destruction of cultivated land, buildings and structures, and groundwater resources (Finkelman and Tian [2018;](#page-14-2) Fernandez et al. [2020;](#page-14-3) Ghabraie et al. [2017;](#page-14-4) Kwinta and Gradka [2020](#page-14-5); Salmi et al. [2017](#page-14-6); Tajdus et al. [2018](#page-14-7); Wang et al. [2021](#page-15-1)). Studying the characteristics of mining subsidence is crucial for protecting the ecological environment of a mining area. To understand the characteristics of mining subsidence, China has already established numerous surface movement observation stations in the central and eastern mining areas (Chi [2021;](#page-14-8) Hu et al. [2017](#page-14-9); Li [2012](#page-14-10); Wu [2015;](#page-14-11) Xie [2015](#page-15-2)). Based on the

measured data, the subsidence factor, boundary angle, and other parameters describing the characteristics of mining subsidence are determined. These parameters can be used to guide the rational mining of coal to avoid surface damage caused by mining. In addition, mining subsidence prediction methods have been proposed by scientists to evaluate mining damage. Surface subsidence prediction methods include the influence function method (Li et al. [2014;](#page-14-12) Zhang et al. [2020](#page-15-3)), mechanical methods (Guo et al. [2016](#page-14-13), [2020](#page-14-14)), and numerical simulation methods (Li et al. [2019](#page-14-15); Gong and Guo [2019\)](#page-14-16). In summary, mastering the characteristics of mining subsidence and building a subsidence prediction model are key to the scientific management of coal mining.

Studies have reported that the characteristics of mining subsidence change significantly when geological and mining conditions change significantly (Deng et al. [2014](#page-14-17)). The mining depth, mechanical strength of the overlying strata, and other factors are the most important factors influencing mining subsidence (Deng et al. [2014](#page-14-17)). For example, the greater the mining depth, the lower the surface subsidence under the same conditions (Deng et al. [2014\)](#page-14-17); additionally, the harder the overburden, the lower the surface subsidence and the greater the height of the fractured water-conducting zone (Deng et al. [2014\)](#page-14-17). The geological conditions of the Hujierte mining area differ significantly from those of the numerous mining areas in western China. For example, the mining depth of the Yushen and Shendong mining areas in western China is approximately 50–420 m. However, the mining depth of the Hujierte mining area is more than 600 m. In addition, the geological conditions of the Hujierte mining area are significantly different from those of the mining area in eastern China. For example, the alluvium in the Hujierte mining area is composed primarily of eolian sand, whereas the alluvium in the Huainan mining area in eastern China is composed primarily of clay. Therefore, the characteristics of mining subsidence in the Hujierte mining area may not be the same as those in the other mining areas. In the western region, scholars have studied the regularity of overlying rock mass and surface movement due to the mining of shallow Jurassic coalfields (Chen et al. [2019](#page-14-1); Liu et al. [2019;](#page-14-0) Li [2017,](#page-14-18) [2020](#page-14-19); Peng et al. [2015;](#page-14-20) Shen and Zhu [2019;](#page-14-21) Xu et al. [2021;](#page-15-0) Yan et al. [2018;](#page-15-4) Zhang and Liu [2015](#page-15-5); Zhu [2019](#page-15-6)). However, the prediction methods and the characteristics of surface subsidence due to the mining of deeply buried Jurassic coal seams have rarely been studied. Accurate assessment of the extent of mining damage in the Hujierte mining area is challenging. Therefore, the prediction method and characteristics of surface subsidence caused by mining deeply buried Jurassic coal seams must be studied.

To address the abovementioned issues, herein, we conducted a case study in a coal mine in the Hujierte mining area. By processing the observed data of surface subsidence in the study area, the characteristics of surface subsidence caused by mining deeply buried Jurassic coal seams were analyzed. Accordingly, a mathematical model suitable for predicting surface subsidence in the mining area was proposed. This study is organized as follows: ["Overview of the](#page-1-0) [study area"](#page-1-0) describes the geological and mining conditions in the study area, as well as the layout of the surface movement observation stations. "[Data processing methods](#page-2-0)" describes the methods used to process subsidence data. ["Characteristics](#page-4-0) [of mining subsidence](#page-4-0)" describes the characteristics of mining subsidence in the study area and compares them with other mining areas. In addition, this section describes the shortcomings of the conventional subsidence prediction model for fitting subsidence data. "A model for predicting surface subsidence based on the T-distribution" describes a new ground subsidence prediction model. At the edge of the subsidence basin, the fitting accuracy of the model was found to be better than that of the conventional method. ["Conclusions](#page-13-0)" provides the main results and limitations of this study.

Overview of the study area

Geological and mining conditions

As shown in Fig. [1,](#page-2-1) the Hujierte mining area is located in the Wushen County, China. The mining area covers an area of 3331.7 km^2 , and the coal reserves in the exploration area are 18,004.10 Mt. The Hujierte mining area was planned by the Chinese government in 2013. Only five coal mines have begun operation in the mining area, and the geological and mining conditions of these coal mines are similar. One of these coal mines was selected for this study.

The designed annual output of this coal mine is 4.0 Mt/ year. In chronological order from oldest to youngest, the strata in the mining area are the Jurassic Yanan formation, Jurassic Zhiluo formation, Jurassic Anding formation, Cretaceous Zhidan formation, and Quaternary. The stratigraphic structure is simple and contains no igneous rock. The coalbearing rock series is from the Jurassic Yanan formation. The 3–1 coal seam is the first coal seam to be mined; it is nearly horizontal, with an average thickness of 5.51 m and depth of the seam more than 600 m. District no. 11 was the first district to be excavated.

As can be seen in Fig. [2,](#page-3-0) the coal mine has arranged longwall panels 101, 102, and 103, which are all adjacent, to the east of district no. 11. The width of the coal pillars between the two panels is 30 m, and their mining sequences are panels 101, 102, and 103. A generalized stratigraphic column exposed by a borehole is shown in Fig. [2.](#page-3-0) The geological and mining conditions of the panels are listed in Table [1](#page-3-1).

As listed in Table [1](#page-3-1), D0, L0, *M*, *α*, *H*, *h*, and *C* are the panel width, panel length, mining height, dip angle of the coal seam,

the study area

average mining depth, average thickness of alluvium, and the panel advance rate, respectively. In Table [2](#page-4-1), *N* is the number of particles; and c_1 and c_2 are the acceleration constants.

The geological and mining conditions in the mining area have the following characteristics, as presented in Table [1](#page-3-1) and Fig. [2.](#page-3-0) (1) The coal seam is buried deep and is nearly horizontal. (2) The alluvium is composed primarily of eolian sand, with an average thickness of 128 m. (3) The rock mechanical strength of the Cretaceous strata is generally soft, whereas that of the Jurassic strata, which is composed predominantly of sandstone with medium-hard strength, is typically greater than 30 MPa. In addition, the fully mechanized full-seam coal mining technology is used for mining, whereas a caving method is used to treat the roof.

The surface movement observation station

Two observation lines (lines A and B) were established at locations approximately 714 m and 1531 m from the open-off cut of panel 102, as shown in Fig. [2.](#page-3-0) The length of observation line A was 1275 m, and it contained 41 observation points with an average spacing of 31 m, marked as $A_1, A_2, \ldots A_{41}$. The length of observation line B was 2184 m, and it contained 66 observation points with an average spacing of 33 m, marked as B_1 , B_2 , \ldots B₆₆. A level was used to observe the elevation data at each point in the coal mine. These points were first observed in the coal mine before panel 101 was advanced. During the observation period, some observation sites were damaged by residents. Therefore, some observation data were missing.

Data processing methods

First, the subsidence values were calculated for each point on the observation line. Second, the boundary angle of the observation line was calculated. Note that boundary angle is an important parameter for measuring the scope of a subsidence basin.

As shown in Fig. [3,](#page-4-2) this study defines the boundary angle as the angle between the line connecting the point with 10 mm of subsidence on the observation line to the boundary of the gob, and the horizontal line on the side of the coal pillar. Herein, *γ* is the boundary angle on the rise and *β* is the boundary angle on the dip.

Finally, the parameters of the probability integration method were inverted.

Probability integration method (PIM) is a prediction method for surface subsidence; it is widely used in China. A detailed introduction to the principle of PIM can be found in Deng et al. [\(2014](#page-14-17)). The center of a rectangular longwall panel serves as the origin, and the *X*- and *Y*-axes point to the strike and dip of the coal seam, respectively. Thus, the formula $W_{\text{pim}}(x, y)$ to calculate the subsidence of any point (x, y) on the surface based on the PIM can be expressed by Eq. ([1\)](#page-2-2) (Deng et al. [2014\)](#page-14-17):

$$
W_{pim}(x, y) = \frac{1}{W_{max}} \Big[W_{PIM} \Big(x + \frac{l}{2} - s_{0y} \Big) - W_{PIM}(x - \frac{l}{2} + s_{0z}) \Big] \Big[W_{PIM} \Big(y - H \cot \theta_0 + \frac{L}{2} - DS_{02} \Big) - W_{PIM} \Big(y - H \cot \theta_0 - \frac{L}{2} + DS_{01} \Big) \Big], \tag{1}
$$

Fig. 2 Relative positions of the three panels and stratigraphic features

where $W_{\text{PIM}}(x)$ and $W_{\text{PIM}}(y)$ indicate the prediction equation for the surface subsidence when the scope of the excavation area is semi-infinite along the strike and dip of the coal seam, respectively (Deng et al. [2014](#page-14-17)); *l* and *L* indicate the computational length of the panel along the strike and dip of the coal seam (m), respectively; and s_{0z} and s_{0y} denote the deviation of the inflection point along the left and right sides of the coal seam (m), respectively, and are parameters of PIM.

Note that $W_{pim}(x, y)$ consists of a W_{PIM} , which is described as follows:

$$
W_{PIM} = \frac{W_{max}}{2} \left(1 + erf\left(\frac{x\sqrt{\pi}}{r_p}\right) \right),
$$
 (2)

where r_p is the main influence radius (*m*); and W_{max} indicates the maximum subsidence value (mm), whose calculation method is as follows:

$$
W_{max} = Mq_0 \cos \alpha, \tag{3}
$$

where q_0 indicates the subsidence factor and is an essential parameter of PIM; M is the mining height (mm); and α is the dip angle of the coal seam (°).

In addition, the calculation for r_p in Eq. [\(2](#page-3-2)) is as follows:

$$
r_p = \frac{H}{\tan \beta 0},\tag{4}
$$

where $tan\beta_0$, which is a parameter of PIM, denotes the tangent of main influence angle.

The relationship between the computational length *L* and the actual length $D(m)$ of the panel along the dip of the coal seam is as follows:

$$
L = \frac{D\sin(\alpha + \theta_0)}{\sin\theta_0},\tag{5}
$$

where θ_0 , which is another parameter of PIM, denotes the influence transference angle.

In Eq. ([1\)](#page-2-2), the expressions for DS_{01} and DS_{02} are as follows:

$$
DS_{01} = \frac{s_{01}\sin(\alpha + \theta_0)}{\sin\theta_0},
$$
\n⁽⁶⁾

$$
DS_{02} = \frac{s_{02}\sin(\alpha + \theta_0)}{\sin\theta_0},\tag{7}
$$

where s_{01} and s_{02} indicate the deviation of the inflection point along the dip and rise sides of the coal seam (m), respectively, and are parameters of the PIM.

The PIM parameter reflects the surface subsidence characteristics of the mining area. For example, the subsidence factor and the tangent of main influence angle can reflect the degree of surface subsidence and the morphological characteristics of the subsidence curve in the mining area, respectively. In general, the subsidence factor decreases as surface subsidence decreases. Essentially, the flatter the subsidence curve, the smaller the tangent of main influence angle. Therefore, in this study, the PIM parameters were used to describe the characteristics of the mining subsidence in the study area.

Particle swarm optimization (PSO) is a common method for inverting the PIM parameters. Detailed descriptions on PSO can be found in the literature (Wen [2015\)](#page-14-22). Owing to space limitations, the method is not discussed in detail in this study. PSO uses a fixed inertia weight, which is inconvenient for determining an optimal solution (Cai et al. [2022](#page-14-23)). Therefore, an adaptive particle swarm algorithm (APSO) was used in this study to invert the PIM parameters. The APSO can perform nonlinear adaptive adjustments to the inertial weight according to Eq. [\(8](#page-4-3)) (Cai et al. [2022](#page-14-23)).

$$
\omega = \omega_{max} - \left(\omega_{max} - \omega_{min}\right) \left(\frac{pso}{iter}/G\right)^2,\tag{8}
$$

where ω_{max} and ω_{min} are the maximum and minimum values of the inertia weight, respectively; *pso_{iter}* is the current number of iterations; and *G* is the maximum number of iterations of the APSO.

Characteristics of mining subsidence

Regularity of the surface subsidence

The ratio between the excavation width and depth in district no. 11 is defined as *R*. As shown in Fig. [2](#page-3-0), when panel 101 was advanced, $R = 0.44$. When panel 102 was advanced, $R = 0.88$. When panel 103 was advanced, $R = 1.31$. The terminal mining line of the longwall panel was far from the observation line (more than 1100 m away). When the advancing position of each panel was close to the terminal mining line, the propagation of the mining subsidence to the observation line was difficult. When each panel completed its advance, the maximum subsidence velocity of the surface was less than 0.9 mm/day. The subsidence of the observation point essentially came to a halt. The surface subsidence curves with different values of *R* are shown in Fig. [4](#page-5-0).

As can be seen in Fig. [4,](#page-5-0) the surface subsidence curve became flatter when the first panel finished advancing $(R=0.44)$. The surface subsidence values were small, and the maximum subsidence values of lines A and B were−188 and − 201 mm, respectively. When panel 102 finished advancing $(R=0.88)$, the subsidence values of lines A and B

Fig. 4 Surface subsidence curves with different *R* values of **a** $R = 0.44$, **b** $R = 0.88$, and **c** $R = 1.31$, **d** maximum subsidence

increased significantly. The maximum subsidence values of the two lines were−1589 and−1557 mm, corresponding to an increase of 8.46 and 7.75 times, respectively. When panel 103 finished advancing $(R=1.31)$, the maximum subsidence values of the two lines were−2064 and−2074 mm, corresponding to increases of 29.9% and 33.3%, respectively. At this point, the magnitude of the increase in subsidence decreased.

Boundary angle of the subsidence basin

Next, the characteristics of the boundary angles were analyzed. The boundary angles of each observation line for the different values of *R* are listed in Table [3](#page-5-1).

As presented in Table [3](#page-5-1), the boundary angle decreased with increasing *R*. When *R* increased from 0.44 to 0.88, the boundary angle β of line B decreased from 56.2 to 46.4°, i.e., by 9.8°. As *R* increased from 0.88 to 1.31, the boundary angle *β* of line B decreased by 1.4°. Thus, when panel 102 completed its advance, the scope of the subsidence basin increased significantly. When panel 103 completed its advance, the extent of the subsidence basin enlargement decreased.

Fitting effect of surface subsidence based on PIM

A comparison between the predicted subsidence values based on the PIM and the measured values is shown in Fig. [5.](#page-6-0) In mining subsidence engineering, the relative error *RE* is

Table 3 Boundary angle (°)

typically used to evaluate the fitting effect (%). $RE = m_w/W_{\text{max}}$, where m_w is the root mean square error (mm) and W_{max} is the maximum subsidence (mm). According to the calculation results, when $R = 0.44$, 0.88, and 1.31, the RE values were 4.6, 5.2, and 4.6%, respectively.

In the mining area, the fitting accuracy of the PIM was less than 6% and the fitting effect was good. However, PIM has certain shortcomings. As shown in Fig. [5](#page-6-0)c–f, the extent of the subsidence basin was large when *R* was greater than or equal to 0.88. The surface subsidence curve had the characteristics of a long-tailed distribution. At this point, the fitting accuracy of PIM at the edge of the subsidence basin was poor. Chinese engineers commonly use PIM to predict the boundary of the subsidence basin. Therefore, accurately

Fig. 5 Fitting effect of PIM when **a** $R = 0.44$ for line A, **b** $R = 0.44$ for line B, **c** $R = 0.88$ for line A, **d** $R = 0.88$ for line B, **e** $R = 1.31$ for line A, and $f R = 1.31$ for line B

Table 4 Parameters of PIM for

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predicting the boundaries of mining-induced damage in the Hujierte mining area using PIM is not possible.

Characteristics of the parameters of PIM

The parameters obtained by inversion are listed in Table [4.](#page-7-0)

As presented in Table [4,](#page-7-0) when panel 101 finished advancing (when $R=0.44$), q_0 was very small (less than 0.13). q_0 increased significantly after panel 102 finished advancing $(R=0.88)$; at this point, tan β_0 also increased.

Chinese engineers classify the overlying strata into three types (Hu et al. [2017\)](#page-14-9): hard (uniaxial compressive strength of more than 60 MPa), medium hard (uniaxial compressive strength of 30–60 MPa), and soft (uniaxial compressive strength of less than 30 MPa). The greater the mechanical strength of the overlying strata, the smaller the subsidence factor and the tangent of main influence angle. According to the literature (Hu et al. [2017\)](#page-14-9), q_0 and tan₀ are generally less than 0.55 and 1.92, respectively, during coal mining under conditions of hard overlying strata. Therefore, when *R* is less than or equal to 1.31, the regularity of surface subsidence in the study area exhibits partial characteristics of surface subsidence during mining under hard overlying strata

conditions. After panel 101 completed its advance, the ratio between the height of the fractured water-conducting zone and the mining height was 23.4 (Yang et al. [2019](#page-15-7)). According to the literature, this ratio is typically greater than 18 for coal mining under hard overlying strata (Hu et al. [2017](#page-14-9)). Thus, the obtained results further indicate that the regularity of surface subsidence and overlying strata movement in the Hujierte mining area exhibits partial characteristics of surface subsidence during mining under hard overlying strata conditions.

Deep Jurassic coal seam is mined in the Hujierte mining area. However, the western region contains numerous mining areas where Jurassic coal seams are mined at shallow depths (e.g., the Huojitu and Bulianta coal mines in the Shendong mining area, the Chahasu coal mine in the Xinjie mining area, and the Yushuwan and Longde coal mines in the Yushen mining area). The mining depth of these areas is approximately 70–418 m, and the mechanical strength of the overlying strata is medium-hard. As shown in Fig. [6](#page-7-1), these mining areas are adjacent to the Hujierte mining area. The geological and mining conditions of some panels of these coal mines are listed in Table [5](#page-8-0).

Fig. 6 Geographic location of coal mines in the mining area in western China

Table 5 Geological and mining conditions of some panels of other coal mines (Chen [2015](#page-14-24); Chi [2021](#page-14-8); Hu et al. [2017](#page-14-9); Jia et al. [2019;](#page-14-25) Li [2012,](#page-14-10) [2017](#page-14-18); Peng et al. [2015;](#page-14-20) Shen and Zhu [2019](#page-14-21); Tao and Jia [2021;](#page-14-26) Wu [2015;](#page-14-11) Xie [2015](#page-15-2); Zhang and Liu [2015;](#page-15-5) Zhu [2019](#page-15-6))

A deep coal seam is mined in the Hujierte mining area. In central and eastern China, several mining areas (e.g., Xuehu coal mine in the Yongxia mining area, Panyi and Zhujidong coal mines in the Huainan mining area, and Tangkou coal mine in the Jining mining area) also mine deeply buried coal seams. These mining areas excavate Carboniferous–Permian coal seams at depths of more than 600 m. The overlying strata have a medium-hard mechanical strength. The geographical location of these mines is shown in Fig. [7](#page-8-1), and the geological and mining conditions of some panels of these coal mines are listed in Table [5](#page-8-0).

In Table [5,](#page-8-0) A-L indicates that the alluvium is composed primarily of eolian sand and loess, whereas C-S indicates that the alluvium is composed primarily of clay and sand layers.

A fully mechanized coal mining technology was used to excavate all the panels listed in Table [5](#page-8-0). A caving method was used to treat the roofs. The data presented in Tables [1](#page-3-1) and [5](#page-8-0) reveal that the excavation size and advance rate in the mining area in western China were much larger than those in eastern China. Essentially, the mining intensity in the western China mining area is greater than in the eastern China mining area. Moreover, the alluvium in the western mining area contains less clay and consists primarily of eolian sand and loess.

To determine whether the subsidence characteristics of these mining areas are similar to those of the Hujierte mining area, the PIM parameters were compared with those of the study area. The PIM parameters for each panel are listed in Table [6](#page-9-0).

As presented in Tables [4](#page-7-0) and [6,](#page-9-0) although the mining degree of the study area was greater than that of some coal mines in the eastern and western regions of China, the q_0 of the study area was still less than that of the coal mines. The Hujierte mining area and some coal mines in eastern China excavate deeply buried coal deposits, and the mining intensity in the Hujierte mining area is higher. However, the q_0 of the Hujierte mining area was found to be much lower than that of some coal mines in eastern China. Thus, compared with the mining of shallow Jurassic coal seams in western China and deep coal seams in eastern China, the subsidence in the Hujierte mining area is small. In addition, the tangent of main influence angle in the study area was smaller than that in other mining areas, which indicates that the surface subsidence curve in the Hujierte mining area is flatter in comparison.

Fig. 7 Geographical location of coal mines in the mining area in central and eastern China

Table 6 Subsidence factor and the tangent of main infuence angle in other mining areas (data in the table are from the literature in Table [5](#page-8-0))

Location	Panel	R	q_0	$tan\beta_0$
West	Huojitu 12205	3.24	0.73	1.98
West	Bulianta 31401	1.11	0.54	4.90
West	Yushuwan 20101	0.93	0.50	3.35
West	Chahasu 3101	0.72	0.80	4.67
West	Longde 205	1.38	0.60	2.00
Central	Xuehu 2102	0.22	0.87	1.78
East	Panyi 1252	0.33	0.75	1.80
East	Zhujidong $111(1)$	0.25	0.46	1.70
East	Tangkou 1301	0.22	0.52	2.15

Preliminary analysis of subsidence mechanism

The results presented in the previous section indicate that the surface subsidence in the Hujierte mining area is small. In this section, the mechanism of small surface subsidence in the mining area is analyzed.

A previous study reported that the lithology and thickness of the rock mass are the factors that influence the mining subsidence characteristics of the mining areas (Hu et al. [2017\)](#page-14-9). To reveal the mechanism of the small surface subsidence in the study area, the characteristics of rock masses in the Hujierte mining area were investigated in this section. The characteristics of rock masses in the two types of typical mining areas were compared with those of the study area. One type of mining area is located in western China; herein, shallowly buried Jurassic coal seams are mined (e.g., the Bulianta coal mine in the Yushen mining area), whereas the other type of mining area is located in central

and eastern China, and mines deeply buried coal seams (e.g., the Xuehu coal mine in the Yongxia mining area).

Taking the study area, and the Bulianta and Xuehu coal mines as examples, the ratio of the total thickness of the rock mass with different lithologies to the height of the overlying rock mass was calculated based on their borehole columnar sections (Li [2012;](#page-14-10) Xu [2016\)](#page-15-8), as shown in Fig. [8](#page-9-1)a. Furthermore, the ratio of the number of rock masses in different thickness ranges to the total number was calculated (see Fig. [8](#page-9-1)b).

As shown in Fig. [8,](#page-9-1) the lithology of the overlying strata in the study area and Bulianta coal mine was dominated by sandstone, which accounted for 60.3% and 60.1%, respectively. However, the percentage of mudstone was very low; in the study area, it was only 6.6%. In contrast, the proportion of mudstone in the Xuehu coal mine in the central regions was relatively large, accounting for 4.6 times that of the study area. The overlying strata of the study area contained thicker rock masses compared with the Bulianta and Xuehu coal mines; in the study area, rock masses thicker than 10 m accounted for 63.6%, which is 2.5 times the thickness of the Bulianta coal mine. In contrast, approximately 89.1% of the rock masses in the Xuehu coal mine had thickness in the range of 0–10 m. The Xuehu coal mine did not contain any rock masses with a thickness of more than 20 m.

In summary, the geological conditions of the Hujierte mining area have some distinct features compared with those of other mining areas: (1) The overlying strata are primarily sandstone; and (2) the overlying strata contain multiple layers of rock mass with large thicknesses and medium-hard mechanical strength.

Compared with the Shendong and Yushen mining areas, where the Jurassic coal seams are shallowly buried, the overlying layers of the Hujierte mining area primarily comprise multi-layer thick and medium-hard rock masses, as shown in

Fig. 8 Stratigraphic information in the study area, Bulianta coal mine, and Xuehu coal mine. **a** Proportion of the diferent lithologies; **b** proportion of the thickness of the rock mass

Fig. [8](#page-9-1). They can be considered a key stratum controlling the movement of the strata. The subsidence factor and tangent of main influence angle in the Hujierte mining area are relatively small because the overlying strata are controlled by a multilayer of thick and medium-hard rock masses.

When the first panel of the Hujierte mining area had finished advancing, the height of the fractured water-conducting zone was 23.6 times the mining height. If the primary key stratum is located within the fractured water-conducting zone, the fissure of the primary key stratum passes through. According to the key stratum theory $(X_u 2016)$ $(X_u 2016)$, when the fissure of the primary key stratum passes through it, the fractured water-conducting zone develops toward the top of the overlying strata. However, in the Hujierte mining area, no fractured water-conducting zone develops on the top of the overlying strata, and the overlying strata of the mining area still contain a sagging zone. Therefore, the primary key stratum is not in the fractured water-conducting zone, and the distance (126 m) between its position and the coal seam is greater than 23.6 times the mining height. However, in the Shendong and Yushen mining areas, the distance between the primary key stratum and the coal seam is often less than 7–10 times the mining height (Xu [2016\)](#page-15-8). Therefore, the distance between the primary key stratum and the coal seam is typically greater in the Hujierte mining area than that in the Shendong and Yushen mining areas. Under the same conditions, the farther the primary key stratum is from the coal seam, the smaller the surface subsidence. This is one reason why the subsidence factor of the Hujierte mining area is smaller than that in the Shendong and Yushen mining areas.

Compared with some coal mines in the mining areas in eastern China, where deeply buried coal deposits are extracted, the overlying strata in the Hujierte mining area contain more sandstone and less mudstone. In addition, each layer of rock mass is generally thicker in the Hujierte mining area. In contrast, in the eastern China mining area, the overlying strata contain more mudstones, and the thickness of the rock mass is small (typically 0–10 m). Compared with mudstone, sandstone is less susceptible to plastic deformation when subjected to an external force. Compared with the thin rock masses, the subsidence caused by mining is less in the thick rock masses. In addition, the alluvium in the mining area of eastern China contains numerous clay components. Therefore, the alluvium in the mining area in eastern China is susceptible to compression after mining disturbances. In the Hujierte mining area, the alluvium is composed primarily of eolian sand. Compared with clay, eolian sand is less susceptible to compression by external forces. Therefore, the subsidence factor in the Hujierte mining area is relatively low compared to that of some coal mines in eastern China.

A model for predicting surface subsidence based on the T‑distribution

As shown in Fig. [5c](#page-6-0)–f, the fitting accuracy of PIM at the edge position of the subsidence basin was poor when *R*≥0.88. The improvement in the fitting accuracy of the edge position of the subsidence basin is conducive to the accurate prediction of the boundary of the mining-induced damage. Therefore, a new subsidence prediction model must be developed to improve the fitting accuracy of the edge position of the subsidence basin.

Prediction model

To further improve the accuracy of subsidence prediction, a prediction model of surface subsidence based on the T-distribution (ITM) was proposed in this study.

The probability density function (PDF) and cumulative distribution function (CDF) of the T-distribution are shown in Eqs. [\(9\)](#page-10-0) and [\(10\)](#page-10-1), respectively (Chen [2020;](#page-14-27) Moore et al. [2012\)](#page-14-28).

$$
f_T = \frac{\Gamma\left(\frac{n+1}{2}\right)}{\sigma \Gamma\left(\frac{n}{2}\right) \sqrt{n\pi}} \left[\frac{n + \left(\frac{x-\mu}{\sigma}\right)^2}{n}\right]^{-\left(\frac{n+1}{2}\right)} \sim T\left(\mu, \sigma^2, n\right), \tag{9}
$$

$$
F_T = \int_{-\infty}^{x} f_T dx = \frac{1}{2} + \frac{\left(\frac{x-\mu}{\sigma}\right) \Gamma\left(\frac{n+1}{2}\right) \binom{n+1}{2} F_1\left(\frac{1}{2}, \frac{n+1}{2}; \frac{3}{2}; -\frac{(x-\mu)^2}{n\sigma^2}\right)}{\Gamma\left(\frac{n}{2}\right) \sqrt{n\pi}},\tag{10}
$$

where $\Gamma(\cdot)$ is the gamma function; ${}_2F_1(\cdot)$ is the hypergeometric function; *n* is the degree of freedom, and $n > 0$; *u* is a location parameter; and σ is the scale parameter.

When $u=0$, according to the properties of the continuous distribution function:

$$
\begin{cases}\n\lim_{x \to -\infty} F_T = 0 \\
\lim_{x \to +\infty} F_T = 1 \\
F_T(x = 0) = \frac{1}{2}\n\end{cases}
$$
\n(11)

In fact, the function curves of F_T and W_{PIM} are both S-shaped. The derivation to relate Eq. (10) (10) to Eq. (2) (2) is as follows:

When the location parameter of the normal distribution is u_n =0, its PDF and CDF are as follows (Fang and Xu [2016\)](#page-14-29):

$$
f_{PDF} = \frac{1}{\sqrt{2\pi}\sigma_n} exp\left(-\frac{x^2}{2\sigma_n^2}\right),
$$
\n(12)

$$
f_{CDF} = \frac{1}{2} \left(1 + erf \left(\frac{x}{\sigma_n \sqrt{2}} \right) \right).
$$
 (13)

The relationship between the influence function of the PIM and PDF of the normal distribution is:

$$
w_{PIM} = \frac{1}{r_p} exp\left(-\pi \left(\frac{x}{r_p}\right)^2\right) \sim N \left(0, \left(\frac{r_p}{\sqrt{2\pi}}\right)^2\right).
$$
 (14)

The scale parameter σ_n of the normal distribution was transformed as follows:

$$
\sigma_n = r_p / \sqrt{2\pi} \tag{15}
$$

Substituting Eq. (15) (15) (15) into Eq. (13) (13) (13) , the CDF of the influence function of the PIM can be obtained as follows:

$$
F_{PIM} = \int_{-\infty}^{x} w_{PIM} dx = \frac{1}{2} \left(1 + \text{erf}\left(\frac{x\sqrt{\pi}}{r_p}\right) \right). \tag{16}
$$

The relationship between the normal distribution and the T-distribution is as follows (Fang and Xu 2016):

$$
T(0, \sigma_n^2, n = \infty) \sim N(0, \sigma_n^2).
$$
 (17)

Substituting Eq. (15) (15) (15) into Eqs. (9) (9) and (10) (10) (10) , Eqs. (18) and (19) (19) (19) can be obtained. Furthermore, we set $u=0$.

$$
f_{ITM} = \frac{\sqrt{2}\Gamma\left(\frac{n+1}{2}\right)}{r_p \Gamma\left(\frac{n}{2}\right) \sqrt{n}} \left[1 + \frac{2\pi x^2}{n r_p^2}\right]^{-\left(\frac{n+1}{2}\right)},\tag{18}
$$

$$
F_{ITM} = \frac{1}{2} + \frac{\sqrt{2}x\Gamma\left(\frac{n+1}{2}\right)_{2}F_{1}\left(\frac{1}{2}, \frac{n+1}{2}; \frac{3}{2}; -\frac{2\pi x^{2}}{nr_{p}^{2}}\right)}{r_{p}\Gamma\left(\frac{n}{2}\right)\sqrt{n}}.
$$
 (19)

The following can be derived from Eq. (17) (17) (17) :

$$
f_{ITM}(n = \infty) \sim N \left(0, \left(\frac{r_p}{\sqrt{2\pi}} \right)^2 \right).
$$
 (20)

The following can be derived from Eq. (14) (14) (14) :

$$
f_{ITM}(n = \infty) = \frac{1}{r_p} exp\left(-\pi \left(\frac{x}{r_p}\right)^2\right).
$$
 (21)

Thus, when *n* is positive infinity, f_{ITM} is equivalent to the influence function w_{PIM} . Therefore, their CDFs are equivalent.

$$
\lim_{n \to +\infty} F_{ITM} = F_{PIM}.\tag{22}
$$

When the scope of the excavation area along the strike or dip of the coal seam is semi-infinite, the prediction equation for the surface subsidence based on the PIM can be expressed as follows:

$$
W_{PIM} = \frac{W_{max}}{2} \left(1 + erf \left(\frac{x\sqrt{\pi}}{r_p} \right) \right) = W_{max} F_{PIM} = W_{max} \lim_{n \to +\infty} F_{ITM},
$$
\n(23)

where W_{max} is the maximum subsidence value.

An influence function is typically the PDF of a continuous distribution function. Therefore, the influence function of the prediction model for the ITM-based surface subsidence can be expressed using Eq. [\(24](#page-11-4)):

Fig. 9 Subsidence curve based on Eq. ([25](#page-11-5))

$$
w_{ITM} = \frac{\sqrt{2}\Gamma\left(\frac{n+1}{2}\right)}{r\Gamma\left(\frac{n}{2}\right)\sqrt{n}} \left[1 + \frac{2\pi x^2}{nr^2}\right]^{-\left(\frac{n+1}{2}\right)},\tag{24}
$$

where *r* is the main influence radius.

Therefore, when the scope of the excavation area along the strike or dip of the coal seam is semi-infinite, the prediction equation for surface subsidence based on the ITM can be derived from Eq. ([23\)](#page-11-6), as follows:

$$
W_{ITM} = W_{max}F_{ITM} = W_{max}\left(\frac{1}{2} + \frac{\sqrt{2}x\Gamma\left(\frac{n+1}{2}\right)_{2}F_{1}\left(\frac{1}{2}, \frac{n+1}{2}; \frac{3}{2}; -\frac{2\pi x^{2}}{n^{2}}\right)}{r\Gamma\left(\frac{n}{2}\right)\sqrt{n}}\right).
$$
\n(25)

The subsidence curve for W_{max} = −1 m based on Eq. ([25](#page-11-5)) is shown in Fig. [9](#page-11-7) (calculated with MATLAB and plotted with Origin).

As shown in Fig. [9,](#page-11-7) the thickness of the tail of the curve gradually increases when *r* increases or *n* decreases. Therefore, the ITM is co-regulated by *r* and *n* to change the morphology of the curve. When the main influence radii of the ITM and PIM were the same, and when *n* was positive infinity, the curves of the ITM and PIM coincided, thus proving the correctness of the derivation in this study.

As shown in Eq. [\(1\)](#page-2-2), $W_{\text{pim}}(x, y)$ consists of the W_{PIM} . Replacing W_{PIM} with W_{ITM} in Eq. [\(1](#page-2-2)), the calculation equation for the subsidence of any point on the surface based on ITM is obtained analogously, as follows:

$$
W(x, y) = \frac{1}{W_{max}} \Big[W_{ITM} \Big(x + \frac{l}{2} - s_y \Big) - W_{ITM} (x - \frac{l}{2} + s_z) \Big] \Big[W_{ITM} \Big(y - H \cot \theta + \frac{l}{2} - DS_2 \Big) - W_{ITM} \Big(y - H \cot \theta - \frac{l}{2} + DS_1 \Big) \Big],
$$
 (26)

$$
W_{max} = Mq\cos\alpha, \tag{27}
$$

$$
L = \frac{D\sin(\alpha + \theta)}{\sin\theta},\tag{28}
$$

$$
DS_1 = \frac{s_1 \sin(\alpha + \theta)}{\sin \theta},\tag{29}
$$

$$
DS_2 = \frac{s_2 \sin(\alpha + \theta)}{\sin \theta},\tag{30}
$$

$$
r = \frac{H}{\tan \beta},\tag{31}
$$

where *q*, tan*β*, and *θ* denote the subsidence factor, tangent of main influence angle, and the influence transference angle, respectively; s_z and s_y denote the deviations of the inflection point on the left and right sides of the coal seam, respectively; and s_1 and s_2 indicate the deviations of the inflection point on the dip and rise sides of the coal seam, respectively.

In summary, there were eight ITM parameters. The physical meanings of *q*, tan β , θ , s_z , s_y , s_1 , and s_2 are the same as those of the PIM parameters. In addition, the ITM has one additionally parameter than the PIM. This parameter is the degree of freedom *n*. The smaller the value of *n*, the thicker the tail of the function curve. Therefore, the ITM is theoretically suitable for predicting surface subsidence with a "long-tailed" distribution in the Hujierte mining area. Equation (23) (23) shows that the ITM is equivalent to PIM when the other parameters are equal and *n* is positive infinity.

Effect of the fitting

The ITM was used to fit the subsidence data for the Hujierte mining area when the *R* value was 0.88 and 1.31. The fitting method used was APSO, which is the same as that of PIM. The parameters and accuracies obtained from the fit are listed in Tables [7](#page-12-0) and [8](#page-12-1), respectively. The effects of the fitting are shown in Fig. [10.](#page-13-1)

As shown in Fig. [10](#page-13-1), the predicted value of the ITM was larger at the edge position of the subsidence basin compared with the PIM. Therefore, the ITM is more suitable for predicting surface subsidence with a "long-tailed" distribution in the Hujierte mining area. Moreover, the data presented in Table [8](#page-12-1) revealed that the total fitting error of the ITM and the fitting error of the edge position of a subsidence basin were smaller than those of PIM. Essentially, when *R* was 0.88 and 1.31, the prediction accuracy of the ITM was better than that of PIM. Therefore, compared with PIM, ITM

Table 8 Comparison between the ftting error of PIM and that of ITM

R	ITM		PIM	
	$m_{\rm t}$ (mm)	m_e (mm)	$m_{\rm t}$ (mm)	m_e (mm)
0.88	69.9	61.1	82.6	78.1
1.31	92.9	79.4	93.7	84.3

Here, m_t is the total fitting error, and m_e is the fitting error for the edge position of the subsidence basin

Fig. 10 Comparison between the fitting effect of PIM and that of ITM, when $a \, R = 0.88$ for line A, $b \, R = 0.88$ for line B, $c \, R = 1.31$ for line A, and $dR = 1.31$ for line B

can more accurately determine the boundaries of mininginduced damage, which is conducive to the protection of important buildings and structures in the mining area.

Conclusions

(1) With an *R* of 0.44, 0.88, and 1.31, the subsidence factors were 0.129, 0.427, and 0.429, and the boundary angles were 55.4°, 46.4°, and 45.0°, respectively. Although the mining degree of the study area was higher than that of some coal mines in the eastern and western regions of China, the subsidence factor of the study area was still lower than that of the coal mines. Owing to several layers of thick and medium-hard rock masses in the overlying strata, the surface subsidence in the study area was low. In addition, a prediction model for surface subsidence based on ITM was proposed. At the edge position of a subsidence basin, the fitting accuracy of the ITM was better than that of PIM, which contributes to a more accurate prediction of the boundary of mining-induced damage.

(2) In this study, the law of surface subsidence in the mining area with the advancement of a longwall panel was not investigated. In the future, we will continue to study the dynamic subsidence law in mining areas. In addition, we will continue to develop new subsidence prediction models to significantly improve the accuracy of subsidence predictions.

Acknowledgements This research was supported by the National Natural Science Foundation of China, Graduate Innovation Program of China University of Mining and Technology, and the Postgraduate Research & Practice Innovation Program of Jiangsu Province. We thank the Wushen Country Coal Bureau of the Inner Mongolia Autonomous Region of China and its affiliated coal mine enterprises and the Mining Damage and Protection Research Institute of China University of Mining and Technology for providing information on conditions of geology and mining, and key measured data of the Hujierte mining area.

Funding Funded by the National Natural Science Foundation of China (grant number 51974292); funded by the Graduate Innovation Program of China University of Mining and Technology (grant number 2022WLKXJ105); funded by the Postgraduate Research & Practice Innovation Program of Jiangsu Province (grant number KYCX22_2588).

Declarations

Competing interests The authors declare no competing interests.

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