## **ORIGINAL PAPER**



# **An improved nonlinear creep damage model of slates considering freeze–thaw damage and bedding damage**

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#### **Abstract**

To study the creep behavior of bedding slates after freezing and thawing, the RLW-2000 rock creep triaxial instrument was used to carry out the creep test, and the creep curves of slate under diferent bedding angles and diferent freeze–thaw cycles were obtained. Then, based on the slate creep test results and the fractional-order theory, a new nonlinear creep damage model considering freeze–thaw damage and bedding damage was proposed. The creep damage model can not only describe the changes in the three creep stages (primary creep, steady-state creep, and accelerated creep) but also refect the infuence of freeze–thaw and bedding coupling damage on creep characteristics. The mathematical optimization analysis software 1stOpt was used to identify the parameters of the nonlinear creep damage model. Finally, the infuence of stress level and bedding angle on creep parameters was analyzed, and the creep characteristic parameters in the creep damage model were sensitively analyzed. The consistency between the test data and the predicted results showed that the nonlinear creep damage model proposed in this study can accurately refect the creep behavior of rock with freeze–thaw and bedding-damaged.

**Keywords** Bedding slate · Creep characteristics · Freeze–thaw (F-T) cycles · Fractional derivative · Creep damage model

# **Introduction**

The freezing and thawing of rock and soil is a natural phenomenon in cold regions. The repeated freeze–thaw action gradually leads to the deterioration of rock mechanical properties, which causes great harm to the construction of geotechnical engineerings, such as slipping of the open-pit mining slopes, cracking and collapsing of tunnel linings, frost heave deformation of railway subgrades, and deterioration of geotechnical materials (Ghobadi and Babazadeh [2015;](#page-15-0) Jin [2010](#page-16-0); Kawamura and Miura [2013;](#page-16-1) Knutsson et al. [2018\)](#page-16-2). In addition, the existence of weak planes (i.e., bedding, jointing, layering, faw, and fssuring et al.) in the rock mass is quite common in nature (Lee and Pietruszczak [2008](#page-16-3); Zhang et al. [2019b](#page-16-4)), and the strength characteristics of anisotropic rocks are also related to the bedding planes orientation (Fu et al. [2018\)](#page-15-1). The existence of bedding planes will aggravate

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the development of rock freeze–thaw damage (Zhang et al. [2019a\)](#page-16-5). One of the most important mechanical properties of rocks creep properties is a key factor afecting the design, working life, safety, and stability of underground caverns in bedding slate (Cornet and Dabrowski [2018](#page-15-2)). Therefore, it is of great signifcance to study the creep characteristics of bedding slates after freezing and thawing to improve the safety of rock engineering projects in cold regions.

The effect of freezing and thawing on the mechanical properties of rocks is a topic of particular concern to researchers (Hale and Shakoor [2003;](#page-15-3) Jia et al. [2019;](#page-16-6) Khanlari et al. [2015](#page-16-7); Sun et al. [2020](#page-16-8); Yang et al. [2010](#page-16-9); Yavuz [2011](#page-16-10)). Wang et al. ([2016\)](#page-16-11) studied the infuence of freeze–thaw cycles on static compression and dynamic impact experiments of red sandstone and used a decay model to describe the reduction of rock strength with the increasing number of freeze–thaw weathering cycles. The efects of freeze–thaw cycles on mechanical characteristics of the Gödene stone in Konya (Turkey) were investigated by Gökçe et al. ([2016\)](#page-15-4). The greatest factor during the freeze–thaw process was the existence of water. The water contained in the body (in discontinuities and pores) of a building stone freezes along with the fall of temperature below 0 °C. Besides, the efect of freeze–thaw action on the durability of rocks was

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more severe in sodium sulfate solution as compared with freshwater (Jamshidi et al. [2016\)](#page-16-12). Yang et al. [\(2019\)](#page-16-13) studied the time-dependent behavior of quartz sandstone and quartzite in four diferent chemical solutions (acidic HCl solution, neutral NaCl solution, alkaline NaOH solution, and distilled water) after freezing and thawing. The elastic modulus of quartzite is the most damaged after HCl solution corrosion, followed by NaOH, NaCl, and pure water turn. Wang et al. [\(2021](#page-16-14)) conducted a shear creep test on red sandstone in a freeze–thaw environment, analyzed the shear creep characteristics of rocks in diferent temperature ranges, and established the Burgers model considering freeze–thaw cycle and damage. Kodama et al. [\(2019](#page-16-15)) conducted a series of uniaxial compression and creep tests on dry and wet samples of tuff under low-temperature conditions  $(-20^{\circ}C)$  and studied the time dependence of the mechanical behavior of frozen rocks. Yang et al. ([2021\)](#page-16-16) analyzed and discussed the infuence of freeze–thaw cycles on the creep characteristics of saturated gneiss and, based on the test results, proposed a variable parameter model that considers the creep damage of the freeze–thaw cycles on the gneiss.

In recent years, many scholars have studied the creep characteristics of rocks (Bai et al. [2021](#page-15-5); Feng et al. [2020;](#page-15-6) Hou et al. [2018;](#page-16-17) Liu et al. [2020](#page-16-18); Wang et al. [2019](#page-16-19)). Multi-stage loading triaxial creep test was performed on the cataclastic rock, and the creep mechanical behavior of the rock was studied (Zhang et al. [2016](#page-16-20)). Rassouli and Zoback ([2018](#page-16-21)) conducted a series of long-term creep experiments on horizontal and vertical shale samples rich in clay and carbonate. Hu et al. [\(2019](#page-16-22)) performed a cyclic loading and unloading creep test on the artifcially layered cemented samples and defned time-independent (instantaneous elasticity and instantaneous plasticity) and time-dependent (viscoelastic and viscoplastic) deformations. Based on the test results, many scholars believe that there is a signifcant coupling characteristic between rock damage and creep and proposed a corresponding creep damage model (Hou et al. [2018](#page-16-17); Ma et al. [2017](#page-16-23); Wang et al. [2019](#page-16-19); Zhang et al. [2009\)](#page-16-24). Zhou et al. ([2011](#page-16-25)) replaced the Newtonian damper in the classical Nishihara model with the Abel damper and proposed a creep constitutive model based on the fractional derivative. Bai et al. [\(2021](#page-15-5)) established a nonlinear viscoelastic-plastic constitutive model of red sandstone with a single ice crack based on the fractional calculus theory and Kachanov damage theory. Feng et al. ([2020\)](#page-15-6) improved the Nishihara model and established a nonlinear creep damage model considering initial damage and damage evolution. Wang et al. [\(2019\)](#page-16-19) combined the theory of continuous damage mechanics and used elastic modulus to characterize rock damage, and established a constitutive model of salt rock creep damage. Based on the multi-stage loading creep test of sandstone with diferent initial damage degrees, Hou et al. [\(2018](#page-16-17)) proposed a new non-linear creep damage model of rock. Cong and Hu [\(2017\)](#page-15-7) conducted a triaxial creep test on sandstone under low confning pressure and established an improved Burgers model to describe initial creep and steady creep. Zhang et al. [\(2020a](#page-16-26)) analyzed the infuence of freeze–thaw cycles on the mechanical properties of water-saturated rocks, combined with rheological theory, and proposed an elastic-visco-plastic model based on the stress function. Kuhn and Mitchell ([1992\)](#page-16-27) proposed the rate process theory (RPT) and used the discrete element method (DEM) to simulate the soil creep process. Gutiérrez-Ch et al. ([2021\)](#page-15-8) used the DEM combined with RPT to simulate the rock response in the creep stage of slate.

The above research results well illustrate the rock creep mechanism, but mainly focus on the creep behavior of homogeneity rock mass under indoor temperature, lacking the research on the creep behavior of rock mass under freeze–thaw cycles, especially the freeze–thaw bedding rock. Although the creep model of rock materials has been improved in recent years, most general models cannot fully describe the creep characteristics of bedding slates after freezing and thawing. Therefore, this study conducted triaxial compression creep tests on bedding slate samples with diferent freeze–thaw cycles. Then, based on freeze–thaw and bedding coupling damage and fractional calculus theory, a creep constitutive model considering freeze–thaw damage and bedding damage was established. Then, the infuence of stress level and bedding angle on creep parameters was analyzed. Through the sensitivity analysis of the creep characteristic parameters, the rationality of the creep damage model proposed in this study was verifed. This study can be used to evaluate the long-term stability of bedding tunnels and bedding slopes in cold areas and provide a theoretical basis for the support design of tunnels and slopes.

## **Experimental procedures**

## **Experimental material**

This study takes the bedding slate as the research object. The slate samples are taken from a tunnel project in a seasonal frozen area in northern China. X-ray difraction (XRD) analysis (Fig. [1](#page-2-0)) indicated that the slate used in this study was composed of quartz, mica, calcite, and other clay minerals. According to the test specifcation recommended by the International Society for Rock Mechanics (ISRM), the standard cylindrical sample with a diameter of 50 mm and a height of 100 mm was prepared. The non-parallelism of both ends of the sample was controlled within  $\pm 0.02$  mm, as shown in Fig. [2](#page-2-1)a. The basic physical and mechanical parameters of the slate sample are presented in Table [1,](#page-2-2) and the curves of the uniaxial compressive strength tests are in Fig. [3](#page-3-0). The dry density of the slate is 2.71  $g/cm<sup>3</sup>$ , the natural water content is 0.45%, the saturated water content is 1.69%,



<span id="page-2-0"></span>**Fig. 1** X-ray difraction (XRD) analysis

and the porosity is 3.25%. For each bedding orientation  $(0^{\circ})$ , 30°, 45°, 60°, and 90°), 10 samples were prepared. In total, 50 samples were prepared, and each set of bedding orientation samples was divided into 5 groups, with 2 samples per group. Each group was subjected to a diferent number of F-T cycles, i.e., 0, 20, 40, 60, and 80 cycles.

#### **Experimental equipment and program**

1. Freeze–thaw cycle tests

The appearance, control panel, and internal structure of the freeze–thaw equipment are shown in Fig. [2](#page-2-1)b. The device has a sinusoidal and linear (including constant temperature)

<span id="page-2-2"></span>**Table 1** The physical and mechanical parameters of the slate

Bedding angle $(°)$	Mean P-wave velocity $(m\bullet s^{-1})$ 5625	Uniaxial compressive strength (MPa)	Elastic modulus (GPa)	Poisson ratio	
$\mathbf{0}$		161.70	60.83	0.27	
30	5431	68.90	37.26	0.22	
45	5394	79.14	45.09	0.21	
60	5440	77.41	42.74	0.24	
90	5728	170.25	73.19	0.31	

regular compound programming capability, which is suitable for F-T cycles ranging from−70 °C to 150 °C. The digital screen displays the set temperature and actual temperature. The temperature uniformity is less than 2.0 °C, whereas the temperature deviation is within  $\pm 2.0$  °C.

This study considers the freeze–thaw cycle (−20 °C  $-20$  °C) as the external environmental conditions of the sample. Each F-T cycle is set as  $12 h (-20 \degree C)$  for 6 h, 20 ℃ for 6 h), and the number of F-T cycles is set as 0, 20, 40, 60, and 80 times, respectively. Before the freeze–thaw cycle tests, the prepared sample is saturated in the saturator for 120 h; then the sample is taken out and put into the water container and then put into the freeze–thaw equipment for the freeze–thaw cycle. As the critical degree of saturation has an important infuence on the freeze–thaw damage of the rock (Al-Omari et al. [2015\)](#page-15-9), the samples always maintain a saturated state during the freeze–thaw process.

2. Triaxial compression creep test



<span id="page-2-1"></span>



<span id="page-3-0"></span>**Fig. 3** Curves of slate uniaxial compressive strength under diferent bedding angles

 The multi-functional RLW-2000 rock triaxial apparatus developed by Dalian Maritime University is used for the uniaxial compression test. The test equipment is shown in Fig. [2](#page-2-1)c. The equipment can be used for a full stress–strain test and rheological test under high and low temperature, high pore pressure, and permeability environment. The full digital servo control, ball screw, and hydraulic technology produced by the German DOLI company can control osmotic pressure, confning pressure, and axial pressure stably. The maximum confning pressure is 80 MPa, and the maximum osmotic pressure is 60 MPa. The control accuracy is within  $\pm 0.01\%$ .

According to the survey data, the ground stress of the tunnel is that the maximum horizontal principal stress is 4.83 MPa—5.74 MPa, and the minimum horizontal principal stress is 4.29 MPa—4.87 MPa. This test only studied the mechanical properties of the bedding slate under the confning pressure of 5 MPa in the above stress range. Table [2](#page-3-1) shows the triaxial compression creep test program for bedding slate samples. The axial load was applied at a constant rate of 0.5 MPa/ min until the loading was stopped at predetermined axial stress (Chen et al. [2018](#page-15-10)). Due to the large dispersion of strength of slate with diferent bedding angles, the initial loading axial stress of the sample is diferent. During a

<span id="page-3-1"></span>**Table 2** Testing program of loading creep

Property	Slate					
Bedding angle $(°)$		30	45	60	90	
Initial value $\sigma_1$ (MPa)	10	10	10	10	10	
$\Delta\sigma$ (MPa)	20	5	10	10	20	
Confining pressure $\sigma_3$ (MPa)	5					

multi-loading triaxial creep test, four to seven levels of axial stress were applied. The minimum holding time for each loading step was 24 h, and the creep strain rate of the sample was less than 0.001 mm/h.

## **Experimental results and analysis**

## **Creep deformation**

The axial strain–time curve of the slate sample under diferent bending angles and diferent freeze–thaw cycles is shown in Fig. [4](#page-4-0). It can be seen from Fig. [4](#page-4-0) that the creep characteristics of the sample are related to the freeze–thaw cycles, bedding angles, and axial stress. Taking the sample with  $\beta = 0^{\circ}$  as an example, it can be seen from Fig. [4a](#page-4-0) that when the axial stress is 10 MPa, the instantaneous elastic strains are 0.0094%, 0.0155%, 0.0209%, 0.0283%, and 0.0357%, respectively, for the samples after freeze–thaw 0, 20, 40, 60 and 80 cycles, and the creep strains are 0.0288%, 0.0381%, 0.0495%, 0.0579%, and 0.0692%, respectively. The samples after freeze–thaw 0, 20, 40, 60, and 80 cycles experienced creep failure at 110 MPa, 110 MPa, 90 MPa, 90 MPa, and 90 MPa, respectively. It can be seen from the above that, under a certain stress level, the instantaneous elastic strain and creep strain of the sample both increase with the increase of the freeze–thaw cycles. However, the creep failure stress of the samples decreases with the increase of the freeze–thaw cycles.

The slate sample produces instantaneous strain at the moment when the stress is applied, and creep strain occurs under long-term stress. Under the action of the frst-level loading stress, the creep deformation of the sample is obvious. The slope of the creep curve increases with increasing stress. The creep curve shows a deceleration creep stage and steady-state creep stage. The creep behavior of the slate sample becomes more and more obvious when the stress continues to be applied. When loaded to the last level of stress, the axial strain of the sample gradually increases with time, showing the typical three stages of rock creep behavior, that is, the primary creep stage, steady-state creep stage, and accelerated creep stage. And the sample failure occurs in the accelerated creep stage. Under certain stress levels, with the increase of freeze–thaw cycles, the instantaneous strain and creep strain of bedding slate increase, and the creep curve changes gradually to steep.

In addition, the axial strain–time curve of the 80 F-T cycles slate under different bedding angles is shown in Fig. [5.](#page-4-1) Under certain freeze–thaw cycles, the creep characteristics of the sample are related to the bedding angles and axial stress. It can be seen from Fig. [5](#page-4-1) that the failure stress of samples with  $\beta = 0^{\circ}$ , 30°, 45°, 60° and 90° are 90, 30, 50,

<span id="page-4-0"></span>



50, and 90 MPa, respectively. With the increase of the bedding angle, the failure stress of the sample frst decreased and then increased, which is similar to the law of uniaxial compressive strength.



<span id="page-4-1"></span>**Fig. 5** Curves of slate creep test under diferent bedding angles (80 F-T cycles)

#### **Creep strain rate**

The creep rate is an important indicator refecting the performance of the creep curve. Under the last stage load, the slate samples with diferent freeze–thaw cycles and diferent bedding angles experienced creep failure. The relationship curve between the axial strain rate and time in the creep test is determined by calculating the slope of the creep curve. Figure [6](#page-5-0) shows the curves of axial strain and strain rate with the time of slate sample with diferent bedding angles under 80 F-T cycles.

It can be seen from Fig. [6](#page-5-0) that the slate sample develops from steady-state creep to acceleration creep after deceleration creep, and then failure occurs in the acceleration creep stage. The duration of the deceleration creep stage and acceleration creep stage is shorter than that of the steady-state creep stage.

Taking the sample with  $\beta = 45^\circ$  as an example, it can be seen from Fig. [6](#page-5-0)c that the strain rate of the sample decreased from  $0.733 \times 10^{-2}$  /h to  $0.017 \times 10^{-2}$  /h and then stabilized in the *AB* region, which is the primary creep stage. In the *BC* region of the creep curve, when the strain rate is  $0.017 \times 10^{-2}$ /h, the strain rate remains unchanged and the axial strain



<span id="page-5-0"></span>**Fig. 6** The axial strain rate of the slate sample under 80 F-T cycles versus time

increases slightly, which is considered the steady-state creep stage. In the *CD* region, the strain rate increases sharply until the rock failure, which can be regarded as the accelerated creep stage. Therefore, the creep characteristics of the bedding slate have obvious time-dependent characteristics. The creep process includes three classical stages: initial creep stage (*AB*), steady-state creep stage (*BC*), and accelerated creep stage (*CD*) (Boukharov et al. [1995](#page-15-11); Main [2010](#page-16-28); Ngwenya et al. [2001\)](#page-16-29).

In Fig. [6c](#page-5-0),  $T_1$  (5.89 h) is used as the primary creep limit, and  $T<sub>2</sub>$  (22.3 h) and  $T<sub>3</sub>$  (23.6 h) are used as the start time and end time of accelerated creep. The primary creep, steadystate creep, and accelerated creep times are 5.89 h, 19.28 h, and 1.3 h, respectively. In the long-term steady-state creep stage, the cumulative creep strain increases with time. When the cumulative creep strain is greater than the steady-state creep limit strain, the sample will enter the accelerated creep stage. During the accelerated creep stage, the axial strain and the strain rate suddenly increase, which indicates that as the microcracks in the specimen expand and propagate, creep failure eventually occurs.

The failure stress, creep duration, and creep rate of the slate with  $\beta = 45^\circ$  under different freeze–thaw cycles were compared, as shown in Table [3.](#page-6-0) The creep rate of the sample increases with the increase of the number of freeze–thaw cycles. For the sample after freeze–thaw 0 cycles, the initial creep rate, steadystate creep rate, and accelerated creep rate are  $0.408 \times 10^{-2}$ ·h<sup>-1</sup>,  $0.0106 \times 10^{-2}$ ·h<sup>-1</sup>, and  $1.073 \times 10^{-2}$ ·h<sup>-1</sup>, respectively, while for the sample after 80 F-T cycles, they are  $0.733 \times 10^{-2}$ ·h<sup>-1</sup>,  $0.0170 \times 10^{-2}$ ·h<sup>-1</sup>, and  $1.597 \times 10^{-2}$ ·h<sup>-1</sup>, respectively. With the increase in freeze–thaw cycles, the failure stress and creep duration of the sample decrease gradually. Creep failure occurred at 70 MPa for the sample after freeze–thaw 0 cycles, and the creep duration is 143.80 h. With the increase in freeze–thaw cycles, the failure stresses of the samples after freeze–thaw 20, 40, 60, and 80 cycles are 60 MPa, 60 MPa, 60 MPa, and 50 MPa, respectively, and the creep duration is 121.91 h, 120.51 h, 119.31 h, and 96.45 h, respectively.

In addition, the creep rate curve of slate under diferent freeze–thaw cycles and diferent bedding angles is shown in Fig. [7](#page-7-0). Taking the sample with  $\beta = 45^\circ$  as an example, it can be seen from Fig. [7](#page-7-0)a that the initial creep rate and accelerated creep rate increase with the increase of freeze–thaw cycles, but the steady creep rate does not change much. The steady creep rates under diferent freeze–thaw cycles are all close to 0. Taking the sample with F-T 80 cycles as an example, as the bedding angles increases, the initial creep rate and accelerated creep rate frst increase and then decrease, but the steady creep rate does not change much, as shown in Fig. [7b](#page-7-0). The variation law of the initial creep rate and the accelerated creep rate of the sample may be related to the bedding structure of the rock.

## **Creep damage model**

It can be seen from the above that the damage caused by the freeze–thaw cycles and the bedding angles afects the creep mechanical properties and creep parameters of the slate. Therefore, it is necessary to carry out research using a creep damage constitutive model that considers the freeze–thaw cycles and bedding angle.

Fractional calculus is a branch of mathematics that studies the possibility of considering the power law of real or complex diferential and integral operators. Fractional calculus can solve the problem in mathematical modeling. Compared with integer calculus, the theoretical model of fractional calculus is in good agreement with experimental results. Fractional calculus uses fewer parameters and more concise expression, so it has been applied in many felds. In recent years, fractional calculus has been gradually developed in the feld of rock rheology, and fractional calculus is often defned by Riemann–Liouville's theory. Riemann–Liouville's fractional derivative takes the derivation of a constant, and the result is not zero (Kilbas et al. [2006\)](#page-16-30). In addition, in the feld of rock creep constitutive models, the Riemann–Liouville fractional calculus operator theory is the most widely used (Liu et al. [2020\)](#page-16-18). For the *m*-order Riemann–Liouville integral of function *f* (*t*), as shown in Eq.  $(1)$  $(1)$ :

<span id="page-6-0"></span>





<span id="page-7-0"></span>**Fig. 7** Creep rate of slate under diferent freeze–thaw cycles and diferent bedding angles

$$
\frac{d^{-m}[f(t)]}{dt^{-m}} = D_t^{-m}f(t) = \frac{1}{\Gamma(m)} \int_0^t (t - \xi)^{m-1} f(\xi) d\xi \tag{1}
$$

The fractional diferential is shown in Eq. [\(2](#page-7-2)):

$$
\frac{d^{m}f(t)}{dt^{m}} = D_{t}^{m}f(t) = \frac{d^{k}}{dt^{k}} \left[ D_{t}^{-(k-m)}f(t) \right]
$$
\n(2)

where  $\Gamma(m)$  is the Gamma function,  $\Gamma(m) = \int_0^\infty t^{m-1} e^{-t} dt$ , *k* is a positive integer, *m* is the fractional order and *m* lies between  $0 \le m \le 1$ .

The fractional derivative is applied to the rheological model (Kelvin and Maxwell) by replacing the damper with the spring. The ideal solid (spring) follows Hooke's law, and the Newtonian fuid (damper) follows Newton's viscosity law (Kabwe et al. [2020](#page-16-31); Yin et al. [2013\)](#page-16-32).

## **Soft‑matter element**

The stress–strain relationship of the ideal solid satisfes Hooke's law, and the ideal fuid satisfes Newton's viscosity law. Therefore, the stress–strain characteristics of geotechnical materials between ideal solid and ideal fuid can be expressed by soft-matter element (Abel dashpot), as shown in Fig. [8](#page-7-3) (Zhou et al. [2011](#page-16-25)).

The stress versus strain relationship of the soft-matter element can be calculated according to Eq. ([3](#page-7-4)):

$$
\sigma(t) = \eta \frac{d^m \epsilon(t)}{dt^m}
$$
 (3)

<span id="page-7-1"></span>where  $\eta$  is the viscosity coefficient of the soft-matter element. *m* is the fractional order and *m* lies between  $0 \le m \le 1$ . Equation [\(3](#page-7-4)) can be reduced to Hooke's elastic law when *m* is 0 and similarly to Newton's fuid law when *m* is 1. The soft-matter element can describe stress versus strain characteristics between ideal elastic and fuid materials when *m* is between 0 and 1. The soft-matter element can describe the accelerated creep state of the rock when *m* is greater than 1.

<span id="page-7-2"></span>The creep constitutive equation of the soft-matter element can be obtained using the Riemann–Liouville fractional diferential operator theory, as shown in Eq. ([4\)](#page-7-5):

<span id="page-7-5"></span>
$$
\varepsilon(t) = \frac{\sigma}{\eta} \frac{t^m}{\Gamma(1+m)}\tag{4}
$$

The creep rate of the Abel dashpot element can be obtained from the derivation of Eq. ([4](#page-7-5)), as shown in Eq.  $(5)$  $(5)$ :

$$
\dot{\varepsilon}(t) = \frac{\sigma}{\eta} \frac{m t^{m-1}}{\Gamma(1+m)} dt
$$
\n(5)

Equation ([5\)](#page-7-6) is a decreasing function when *m* is between 0 and 1. That is, the strain rate decreases continuously in the creep process, which can describe the deceleration creep

<span id="page-7-6"></span>

<span id="page-7-4"></span><span id="page-7-3"></span>**Fig. 8** Soft-matter element

characteristics of rock materials. Equation ([5\)](#page-7-6) is an increasing function when *m* is greater than 1, which describes the accelerated creep characteristics. Equation [\(5](#page-7-6)) can be transformed into Eq. [\(6](#page-8-0)) when *m* is 1:

$$
\dot{\varepsilon}(t) = \frac{\sigma}{\eta} \tag{6}
$$

Equation ([4](#page-7-5)) shows that when *m* is 1, the strain rate  $\dot{\epsilon}(t)$ is constant, which can describe the steady-state creep characteristics of rock materials.

Equation ([4\)](#page-7-5) denotes the creep strain characterized by the Abel dashpot. Substituting  $\sigma$  = 30 MPa and  $\eta$  = 3 GPa·*h* into Eq. [\(4](#page-7-5)), one fnds a series of creep curves under the cases of diferent derivative orders *m* (Fig. [9\)](#page-8-1), showing the dependence of creep strain on the derivative order.

Figure [9a](#page-8-1) shows that for an *m* value between 0 and 1, the strain versus time curves gradually increase, although the growth rate slows gradually, which describes the steadystate creep characteristics of the geomaterials. In addition, the growth rate of the creep curves increases signifcantly when *m* is greater than 1 (Fig. [9b](#page-8-1)), which in turn describes the accelerated creep characteristics.

#### **Bedding, freeze–thaw, and load coupling damage**

According to the creep test results of slate with diferent bedding angles and diferent freeze–thaw cycles, the bedding, and freeze–thaw coupling damage was introduced. A nonlinear viscoelastic plastic creep model considering the coupling damage of freeze–thaw and bedding was established to describe the creep mechanical properties of slate.

According to Lemaitre's strain equivalence principle, the strain caused by the action of total stress  $\sigma$  on the damaged material is equivalent to that caused by the action of effective stress  $\sigma'$  on the non-damaged material.

<span id="page-8-3"></span><span id="page-8-2"></span>

$$
\varepsilon = \frac{\sigma}{E'} = \frac{\sigma'}{E} \tag{7}
$$

where  $E$  and  $E'$  are the elasticity modulus of non-damaged rock and damaged rock, respectively.

<span id="page-8-0"></span>At present, there is no mature method to refect the infuence of bedding angle in the creep model of bedding rock. Some scholars regard bedding as a kind of damage in the rock mechanics model (Shi et al. [2020;](#page-16-33) Zhang et al. [2021](#page-16-34)). Therefore, this study tries to introduce rock bedding into the creep model in the form of a damage variable.

According to Eq. ([7](#page-8-2)), the bedding angles damage variable  $D_\beta$  of slate is defined as

$$
D_{\beta} = 1 - \frac{E_{\beta}}{E_0} \tag{8}
$$

where  $\beta$  is the bedding angle,  $E_{\beta}$  is the elastic modulus of the slate sample with bedding angle  $\beta$ , and  $E_0$  is the elastic modulus of the slate sample with bedding angle  $\beta = 0^\circ$ .

The freeze–thaw damage variable  $D_n$  of slate is defined as (Zhang et al. [2020b\)](#page-16-35)

<span id="page-8-4"></span>
$$
D_n = 1 - \frac{E_n}{E} \tag{9}
$$

where *n* is the freeze–thaw cycles,  $E_n$  is the elastic modulus of the slate sample at certain bedding angles after freeze–thaw *n* cycles, *E* is the elastic modulus of an unfrozen and thawed slate sample at certain bedding angles.

The total damage variable of rock with bedding angles under freeze–thaw cycles was obtained from Eqs. ([8\)](#page-8-3) and ([9\)](#page-8-4).

<span id="page-8-5"></span>
$$
D_{\beta n} = D_{\beta} + D_n - D_{\beta} D_n \tag{10}
$$

Equation  $(10)$  $(10)$  characterizes the nonlinear relationship between the two types of damage caused by bedding angles and freeze–thaw cycles and total damage. The

0 5 10 15 20 25 30 0.00 0.05 0.10 0.15 0.20 0.25 0.30  $Strain\%$  $\frac{15}{\text{Time/h}}$ *m*=0.1 *m*=0.2 *m*=0.3 *m*=0.5 *m*=0.8  $m=0.9$  $=1.0$ 0 10 20 30 40 50 0.00 0.05 0.10 0.15 0.20 0.25 0.30  $Strain\%$ Time/h *m*=1 *m*=1.5 *m*=2 *m*=5 *m*=10

**(a)**  $0 \le m \le 1$  **(b)**  $m \ge 1$ 

<span id="page-8-1"></span>

ing efect on total damage. In addition, when the applied load reaches or exceeds certain axial stress, load damage will occur inside the rock, so the infuence of load damage on creep parameters should also be considered (Yang et al. [2021\)](#page-16-16). Since the load damage in the two stages of deceleration creep and steady-state creep is relatively small, this study only considers the damage caused by the stress in the accelerated creep. The load damage  $D<sub>s</sub>$  is:

$$
D_s = 1 - e^{-\alpha t} \tag{11}
$$

According to Eqs.  $(8)$  $(8)$  $(8)$ ,  $(9)$ , and  $(11)$  $(11)$  $(11)$ , the coupling damage of bedding damage, freeze–thaw damage, and load damage at the accelerated stage can be expressed as

$$
D_{\beta ns} = D_{\beta} + D_{n} + D_{s} - D_{\beta}D_{n} - D_{\beta}D_{s} - D_{n}D_{s} + D_{\beta}D_{n}D_{s}
$$
\n(12)

Figure [10](#page-9-1) shows the creep curves of rock at diferent stress levels, and Fig. [11](#page-9-2) shows the nonlinear creep model considering bedding, freeze–thaw, and load coupling damage. Combined with the classical creep model (Maxwell, Kelvin, and Binham model) and the above-improved creep components, a nonlinear viscoelastic plastic creep model considering the bedding and freeze–thaw coupling damage was established, as shown in Fig. [11](#page-9-2).

#### **Instantaneous elastic deformation component**

In the rock creep test, when the stress level is less than the long-term strength of the rock, the sample will produce instantaneous strain during the loading process. Since the loading time is shorter than the later creep time, the sample can be considered that the elastic strain is instantaneously completed. The Hooke body is used to describe



<span id="page-9-1"></span>**Fig. 10** Creep curve of rock under diferent stress levels



<span id="page-9-2"></span>**Fig. 11** Nonlinear creep model considering bedding, freeze–thaw, and load coupling damage

<span id="page-9-0"></span>the instantaneous elastic deformation of the rock, and the constitutive relation is given by

$$
\varepsilon_{\rm M} = \frac{\sigma'}{E_{\rm M}}\tag{13}
$$

where  $\varepsilon_M$ ,  $E_M$  are the strain and elasticity modulus of the Hooke body, respectively.  $\sigma'$  is deviatoric stress.

### **New nonlinear visco‑elasticity component**

The strain rate of rock gradually decreases and approaches 0 at last. Therefore, the strain curve in the decay creep stage has obvious nonlinear characteristics. The main reason for the model is that the parameters of the model are constant and do not change with time. In this paper, the modifed Kelvin model is used to express the nonlinear characteristics of the rock creep curve, that is, the software element is parallel with the Hooke body to describe the decay creep deformation of rock. According to the series–parallel relationship, it can be obtained that:

$$
\begin{cases} \varepsilon_{\rm K} = \varepsilon_{\rm H} = \varepsilon_{\rm N} \\ \sigma' = E_{\rm K} \varepsilon_{\rm K} + \eta_{\rm K} \frac{d^{m_1} \varepsilon_{\rm K}(t)}{dt^{m_1}} \end{cases}
$$
(14)

where  $\varepsilon_{\rm H}$ ,  $\varepsilon_{\rm N}$  are the strain of the Hooke body and the strain of the soft-matter element, respectively.  $E_K$ ,  $\varepsilon_K$  are the elasticity modulus and strain of the Kelvin body, respectively.  $\sigma'$ is deviatoric stress,  $\eta_K$  is the viscosity coefficient of the softmatter element, *m* is the fractional order and *m* lies between 0≤*m*≤1.

<span id="page-9-3"></span>Equation  $(6)$  $(6)$  can be rewritten as

$$
\frac{\sigma'}{\eta_K} = \frac{E_K}{\eta_K} \varepsilon_K + \frac{d^{m_1} \varepsilon_K(t)}{dt^{m_1}}
$$
(15)

Considering an initial condition  $\varepsilon_K(t) = 0$  when time  $t = 0$ , and suppose  $a = E_K / \eta_K$ ,  $b = \sigma' / \eta_K$ , Eq. ([15](#page-9-3)) can be rewritten as

<span id="page-9-4"></span>
$$
b = a\varepsilon_{\mathcal{K}} + \frac{d^{m_1}\varepsilon_{\mathcal{K}}(t)}{dt^{m_1}}\tag{16}
$$

Based on the theory of fractional calculus (Kilbas et al. [2006](#page-16-30)), the relationship between the Riemann–Liouville fractional derivative and the Caputo fractional derivative  ${}^CD_t^m f(t)$  is given by

$$
D_t^m f(t) = C D_t^m f(t) + \sum_{k=0}^{n-1} \frac{t^{k-m} f^{(k)}(0)}{\Gamma(1+k-m)} (0 < m < n)
$$
 (17)

Considering an initial condition  $\varepsilon_K(0) = 0$ , one gets  $D^m\big[\varepsilon_K(t)\big] = C D^m\big[\varepsilon_K(t)\big]$ , then Eq. [\(16](#page-9-4)) can be rewritten as

$$
b = a\varepsilon_{\mathcal{K}} + \,^C D^m \big[ \varepsilon_{\mathcal{K}}(t) \big] \tag{18}
$$

Taking a Laplace transform on both sides of Eq. ([18](#page-10-0)), we get

$$
\frac{b}{s} = aE(s) + s^m E(s) \tag{19}
$$

i.e.,

$$
E(s) = \frac{b}{s(s^m + a)}\tag{20}
$$

Taking Laplace inverse transform on Eq.  $(20)$  $(20)$ , we get

$$
\varepsilon_{\mathcal{K}}(t) = b \int_{0}^{t} (t - s)^{m-1} E_{m,m} \left[ -a(t - s)^{m} \right] ds \tag{21}
$$

where

$$
E_{m,m}(x) = \sum_{k=0}^{\infty} \frac{x^k}{\Gamma(km+m)}
$$
\n(22)

It can be further given by

$$
\varepsilon_{\mathcal{K}} = b \sum_{k=0}^{\infty} \frac{(-a)^k t^{m_1(1+k)}}{m_1(1+k)\Gamma[m_1(1+k)]}
$$
(23)

Substituting  $a = E_K / \eta_K$  and  $b = \sigma / \eta_K$  into Eq. ([21](#page-10-2)), we obtain the general solution:

$$
\varepsilon_{\rm K} = \frac{\sigma}{\eta_{\rm K}} \sum_{k=0}^{\infty} \frac{(- (E_{\rm K}/\eta_{\rm K}))^k t^{m_1(1+k)}}{m_1(1+k)\Gamma[m_1(1+k)]}
$$
(24)

where  $k$  is a positive integer,  $m_1$  is the fractional order and  $m_1$  lies between  $0 \le m_1 \le 1$ .

#### **Viscous rheological component**

When the stress level is close to the long-term strength of the rock, the strain of the rock is linear with time after the deceleration creep stage, and the stress–strain relationship in this process is described by the viscous body. Newton's

body is used to describe the steady-state creep deformation of rock, and the constitutive relation is given by

$$
\sigma' = \dot{\varepsilon}_{\rm v} \eta_{\rm M} \tag{25}
$$

where  $\dot{\epsilon}_v$  is the strain rate of the Newton body,  $\sigma'$  is deviatoric stress,  $\eta_M$  is the viscosity coefficient of the Newton body.

#### **Nonlinear visco‑plasticity component**

<span id="page-10-0"></span>When the stress level exceeds the long-term strength of the rock, the rock passes through the deceleration creep stage and the steady-state creep stage and quickly enters the acceleration creep stage. In the acceleration creep stage, the strain and strain rate increase rapidly, and the nonlinear characteristics of the strain curve are obvious. When *m* is greater than 1, the strain of the soft-matter element increases signifcantly with time, showing the characteristics of accelerated creep, and the creep characteristics gradually increase with the increase of the *m* value. Therefore, the viscoplastic element with soft-matter element and plastic slider in parallel can be used to describe the viscoplastic deformation of rock, to better refect the accelerated creep stage of rock.

<span id="page-10-2"></span><span id="page-10-1"></span>According to the principle of parallel connection of elements in Fig. [11,](#page-9-2) the stress–strain relationship of the Binham body can be obtained as follows

$$
\varepsilon_{\rm B} = \begin{cases}\n0 & , \sigma < \sigma_{\rm s} \\
\frac{\sigma' - \sigma_{\rm s}}{\eta_{\rm B}} \frac{t^{m_2}}{\Gamma(1 + m_2)}, \sigma \ge \sigma_{\rm s}\n\end{cases} \tag{26}
$$

where  $\eta_B$  is the viscosity coefficient of the Binham body,  $m_2$ is the fractional order and  $m_2$  lies between  $0 \le m_2 \le 1$ .  $\sigma_s$  is the long-term strength of the slate, which can be obtained experimentally.

#### **Nonlinear creep damage model**

 $\overline{a}$ 

According to Fig. [11](#page-9-2), the components of the nonlinear creep damage model satisfy the following equations

$$
\sigma = \sigma_{\rm M} = \sigma_{\rm K} = \sigma_{\rm B}
$$
\n
$$
\varepsilon = \varepsilon_{\rm M} + \varepsilon_{\rm K} + \varepsilon_{\rm B}
$$
\n
$$
\sigma_{\rm M} = E_{\rm M} (1 - D_{\beta n}) \varepsilon_{\rm e} = \eta_{\rm M} (1 - D_{\beta n}) \dot{\varepsilon}_{\rm v}
$$
\n
$$
\sigma_{\rm K} = E_{\rm K} (1 - D_{\beta n}) \varepsilon_{\rm K} + \eta_{\rm K} (1 - D_{\beta n}) \frac{d^{m_1} \varepsilon(t)}{dt^{m_1}}
$$
\n
$$
\sigma_{\rm B} = \eta_{\rm B} (1 - D_{\beta n s}) \frac{d^{m_2} \varepsilon(t)}{dt^{m_2}} + \sigma_{\rm s} (1 - D_{\beta n s})
$$
\n(27)

where  $\sigma$ ,  $\varepsilon$  are the total stress and total strain, respectively.  $\sigma_{\rm M}$ ,  $\sigma_{\rm K}$ ,  $\sigma_{\rm B}$  are the stress on Maxwell's body, Kelvin's body, and Binham's body, respectively.  $\epsilon_M$ ,  $\epsilon_K$ ,  $\epsilon_B$  are the strain of Maxwell's body, Kelvin's body, and Binham's body, respectively.

The creep model considering freeze–thaw and bedding coupling damage shown in Fig. [11](#page-9-2) is given by

 $\beta$ =0° as an example, when the stress levels are 10, 30, and 50 MPa, the test curve is characterized by the deceleration creep stage, and the formula under the conditions  $\sigma' < \sigma_s$ in Eq. ([28\)](#page-11-0) is used for fitting. When  $\sigma' \geq \sigma_s$ , the rock enters the accelerated creep stage, the formula under the condition  $\sigma' \geq \sigma_s$  of Eq. [\(28\)](#page-11-0) is used for fitting. In addition, the long-term strength  $(\sigma_s)$  of rock mass is an important indica-

$$
\varepsilon = \begin{cases}\n\frac{\sigma}{E_{\text{M}}(1 - D_{\beta n})} + \frac{\sigma}{\eta_{\text{M}}(1 - D_{\beta n})}t + \frac{\sigma}{\eta_{\text{K}}(1 - D_{\beta n})} \sum_{k=0}^{\infty} \frac{(- (E_{\text{K}}/\eta_{\text{K}}))^{k} t^{m_{1}(1+k)}}{m_{1}(1+k) \Gamma[m_{1}(1+k)]}, \sigma < \sigma_{\text{s}} \\
\frac{\sigma}{E_{\text{M}}(1 - D_{\beta n})} + \frac{\sigma}{\eta_{\text{M}}(1 - D_{\beta n})}t + \frac{\sigma}{\eta_{\text{K}}(1 - D_{\beta n})} \sum_{k=0}^{\infty} \frac{(- (E_{\text{K}}/\eta_{\text{K}}))^{k} t^{m_{1}(1+k)}}{m_{1}(1+k) \Gamma[m_{1}(1+k)]} + \frac{\sigma' - \sigma_{\text{s}}(1 - D_{\beta n \text{s}})}{\eta_{\text{B}}(1 - D_{\beta n \text{s}})} \frac{t^{m_{2}}}{\Gamma(1 + m_{2})}, \sigma \ge \sigma_{\text{s}}\n\end{cases}
$$
\n(28)

# **Model verification**

# **Parameter identification and validation of test results**

Adopt the Boltzmann superposition principle (Landel and Nielsen [1993\)](#page-16-36) to process the creep curve. Based on the experimental results, the universal global optimization algorithm in the mathematical optimization analysis software 1stOpt is used to carry out a regression analysis of the model in this paper, and the optimal model parameter values are determined. Taking the creep test of slate with <span id="page-11-0"></span>tor of long-term stability and safety considering the timedependent behavior (Damjanac and Fairhurst [2010](#page-15-12)). The steady-state creep rate method was used for data analysis in this study. The parameter  $\sigma_{\rm s}$  and the creep parameter identifcation results are shown in Table [4.](#page-11-1) The ftting curve of the triaxial creep test of slate samples with diferent bedding angles after undergoing 80 freeze–thaw cycles is shown in Fig. [12](#page-12-0).

Figure [12](#page-12-0) shows the comparison between the creep test curve and ftting curve of the theoretical model of slate samples with diferent bedding angles after undergoing 80 freeze–thaw cycles. Table [4](#page-11-1) shows that the theoretical



<span id="page-11-1"></span>**Table 4** Creep damage model parameters under diferent conditions ( $F-T$  cycles $=80$ )



<span id="page-12-0"></span>**Fig. 12** Creep data of slate with diferent bedding angles after 80 F-T cycles and ftting results of the creep damage model: (**a**) 0°, (**b**)30°, (**c**) 45°, (**d**) 60°, (**e**) 90°

<span id="page-13-0"></span>**Table 5** Determination results of creep parameters of slate





curve and the test data fit well, and the fitting coefficients are both greater than 0.9. The model ftting curve refects the characteristics of decelerating creep, steady-state creep, and accelerated creep stage under diferent bedding angles and diferent stress. This verifes the accuracy and applicability of the freeze–thaw and bedding coupling damaged creep constitutive model established in this study.

## **Prediction of test results for slate**

To further illustrate the rationality of the model proposed in this study, the slate creep test results provided by Mao et al. [\(2006\)](#page-16-37) are used to further verify the model. The determination results of the model parameters are shown in Table [5.](#page-13-0) Figure [13](#page-13-1) shows the comparison between the theoretical curve of the model and the experimental results. It can be seen that the theoretical curve of the model is in good agreement with the experimental results, indicating that the model can well characterize the creep mechanical properties of slate, which also proves the rationality of the model in this study.



<span id="page-13-1"></span>**Fig. 13** Comparison between the theoretical curve and test data of slate

## **Discussion**

## **Influence of the stress level on model parameters**

Figure [14](#page-14-0) shows the effect of stress level on creep parameters of  $E_M$ ,  $\eta_M$ ,  $E_K$ ,  $\eta_K$  and  $m_1$ , respectively. It can be seen that parameters  $E_M$  and  $E_K$  increase with the increase in stress level. However, parameters  $\eta_M$  and  $\eta_K$  decrease with the increase in stress level. The value of fractional order  $m<sub>1</sub>$  increases with the increase in stress level. That is, the creep rate of the creep curve increases with the increase of stress. Under the action of axial stress, the stress damage of slate increases with the increase of creep time. Under the action of the last stage stress, the creep rate of the slate sample is the largest, so the fractional-order  $m_1$  is the largest. When the sample fails, the creep rate increases signifcantly, and the creep curve enters the accelerated creep stage.

It can be seen from the above that parameters  $E_M$ and  $E<sub>K</sub>$  increase with the increase of stress level, while parameters  $\eta_M$  and  $\eta_K$  decrease with the increase of stress level. The fractional-order  $m_1$  reflects the creep rate of rock in the creep process. At the maximum stress level, when  $\beta$  is 60°, the value of  $m_1$  is the largest. That is to say, before the accelerated creep stage, the creep characteristics of rock samples with  $\beta = 60^{\circ}$  are the most obvious.

# **Influence of viscosity coefficient**  $\eta_B$ **and the fractional derivative** *m***<sup>2</sup>**

Figure [15](#page-15-13) shows the infuence of bedding angles on the viscosity coefficient  $\eta_B$  and the fractional derivative  $m_2$ . It can be seen from Fig. [15](#page-15-13) that with the increase of bedding angle, the changing trend of model parameters  $\eta_B$  and  $m_2$ first decreases and then increases. When  $\beta$  is 60°, these two parameters are the smallest. That is, the corresponding cohesive force *c* and internal friction angle *φ* are also the smallest, and the shear strength of the sample is also the smallest.



<span id="page-14-0"></span>**Fig. 14** Influence of the stress level on model parameters of (**a**)  $E_M$ , (**b**)  $\eta_M$ , (**c**)  $E_K$ , (**d**)  $\eta_K$ , (**e**)  $m_1$ 



<span id="page-15-13"></span>**Fig. 15** Influence of bedding angles on model parameters: $\eta_B$  and  $m_2$ 

# **Conclusion**

To study the creep characteristics of bedding slates under diferent freeze–thaw cycles, a series of triaxial creep tests were carried out. Then, based on freeze–thaw and bedding coupling damage and fractional calculus theory, a creep constitutive model considering freeze–thaw damage and bedding damage was established. The main conclusions are as follows:

- 1. The test results show that the number of freeze-thaw cycles and the bedding angle have a signifcant efect on the creep behavior of the slate. The creep curve gradually changed from two-stage steady-state creep to three-stage unsteady-state creep, and fnally accelerated creep. Under certain bedding angles, the instantaneous deformation, creep deformation, axial initial creep rate, and the steady-state creep rate of slate increase with the increase of freeze-thaw cycles and axial stress.
- 2. According to the nonlinear rheological theory and fractional calculus theory, considering the freeze-thaw damage and bedding damage, a creep damage model which can simultaneously describe the instantaneous elastic strain, vicious strain, nonlinear viscoelastic strain, and nonlinear viscoplastic strain of freeze-thaw bedding slate is established in this study. Then, the one-dimensional fractional diferential creep damage model is derived to three-dimensional. Finally, a simple and feasible method of model parameter identifcation is given. The experimental verifcation shows that the model can efectively express the infuence of the bedding angles and the freeze-thaw cycles on the creep characteristics of slate, and it can well refect the nonlinear creep characteristics of the acceleration stage of slate.

3. Through the sensitive analysis of the creep characteristic parameters in the creep damage model, it is found that as the stress level increases,  $E_M$  and  $E_K$  gradually increases, while  $\eta_M$  and  $\eta_K$  gradually decrease. The faster the axial strain increases after the rock enters the accelerated creep stage, the shorter the time of the tendency to fail. In addition, with the increase of the bedding angle, the changing trend of the model parameters  $\eta_B$  and  $m_2$  first decreases and then increases.

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