

---

# Non-Darcian flow to a well in a leaky aquifer using the Forchheimer equation

Zhang Wen · Guanhua Huang · Hongbin Zhan

**Abstract** Non-Darcian flow to a well in a leaky aquifer was investigated using a finite difference method. Flow in the leaky aquifer is assumed to be non-Darcian and horizontal, while flow in the aquitard is assumed to be Darcian and vertical. The Forchheimer equation was employed to describe the non-Darcian flow in the aquifer. The finite difference solution was compared with the solution of Birpinar and Sen (2004). The latter overestimates the drawdown at early times and underestimates the drawdown at late times; also, the impact of  $\beta_D$  on the drawdown depends on the value of  $B_D$ , where  $\beta_D$  is a dimensionless turbulent factor in the Forchheimer equation and  $B_D$  is the dimensionless leakage parameter. The impact of leakage on drawdown is similar to that of

Darcian flow. A sensitivity analysis indicated that the drawdown is very sensitive to the change in the dimensionless well radius  $r_{cD}$  when  $B_D$  is relatively large, while it is sensitive to the change in  $B_D$  when  $B_D$  is relatively small. The numerical solution has been applied to analyze the pumping test data in Chaj-Doab area of Pakistan. Birpinar ME, Sen Z (2004) Forchheimer groundwater flow law type curves for leaky aquifers. *J Hydrol Eng* 9(1):51–59

**Keywords** Groundwater flow · Forchheimer equation · Finite difference method · Aquitard · Pakistan

## Introduction

Darcy's law has been used for over one and a half centuries for solving groundwater problems (Wen et al. 2009). However, flow can be non-Darcian under certain conditions as long as the flow velocity is relatively high or low (Forchheimer 1901; Dudgeon 1966; Basak 1977; Sen 1990; Choi et al. 1997; Bordier and Zimmer 2000). The limitations of Darcy's law for solving flow problems have long been recognized and even Darcy himself realized that the linear relationship only worked for a certain range of grain size under a certain range of hydraulic gradient (Darcy 1856).

Many scientists have investigated the relationship between the hydraulic gradient and specific discharge for non-Darcian flow (e.g., Forchheimer 1901; Rose 1951; Polubarinova-Kochina 1962; Muskat 1937; Harr 1962; Izbash 1931; Escande 1953; Wilkinson 1956; Slepicka 1961), as summarized by Basak (1978). As can be seen from Table 1 in the paper of Basak (1978), all the relationships can be classified into two types: i.e., polynomial and power functions. Among these empirical (or theoretical) equations, two were commonly used. The first is the Forchheimer equation, which states that the hydraulic gradient is a second-order polynomial function of the specific discharge. The second is the Izbash equation which states that the hydraulic gradient is a power function of the specific discharge. Both equations have advantages and disadvantages. The Forchheimer equation has two terms, the first term represents the viscous term and the second term represents the inertial term. The physical meaning of the Forchheimer equation

---

Received: 12 July 2010 / Accepted: 30 January 2011  
Published online: 19 February 2011

© Springer-Verlag 2011

---

Z. Wen (✉)  
School of Environmental Studies,  
China University of Geosciences,  
388 Lumo Rd., Wuhan, Hubei 430074,  
People's Republic of China  
e-mail: wenzhangcau@gmail.com

Z. Wen · G. Huang  
Department of Irrigation and Drainage,  
College of Water Conservancy and Civil Engineering,  
China Agricultural University, Beijing, 100083,  
People's Republic of China

G. Huang  
Chinese-Israeli International Center for  
Research and Training in Agriculture,  
China Agricultural University, Beijing, 100083,  
People's Republic of China  
e-mail: ghuang@cau.edu.cn

H. Zhan  
Department of Geology and Geophysics,  
Texas A&M University, College Station, TX 77843-3115, USA  
e-mail: zhan@geo.tamu.edu

H. Zhan  
Faculty of Engineering and School of Environmental Studies,  
China University of Geosciences,  
388 Lumo Rd., Wuhan, Hubei 430074,  
People's Republic of China

**Table 1** Definition of the dimensionless variables

$r_D = \frac{r}{m}$	$r_{cD} = \frac{r_c}{m}$	$r_{wD} = \frac{r_w}{m}$	$q_D = -\frac{4\pi m^2}{Q} q$	$B_D = \frac{Bk}{m^2}$
$t_D = \frac{kt}{Sm}$	$s_D = \frac{4\pi km}{Q} s$	$s_{wD} = \frac{4\pi km}{Q} s_w$	$\beta_D = \frac{\beta Q}{4\pi m^2}$	

is evident and it has been validated theoretically by several scientists (Irmay 1958; Whitaker 1996; Giorgi 1997; Sorek et al. 2005). When the velocity is relatively low, the second term, as opposed to the first term, can be ignored. In this case, the Forchheimer equation becomes Darcy's law. The Izbash equation is a fully empirical equation based on numerous experimental data. However, because of the mathematical convenience, it also has been commonly used (e.g., Wen et al. 2006, 2008a, 2008b). In many cases, these two equations can describe non-Darcian flow equivalently well (Bordier and Zimmer 2000; Li et al. 2008).

Because of the high velocities, non-Darcian flow is likely to occur near pumping wells (Wu 2002a, b; Sen 1987, 1988, 1989, 1990; Wen et al. 2006, 2008a, b, c, 2009). Based on the Forchheimer or Izbash equations, many studies have been carried out to investigate the non-Darcian flow near the pumping well. A careful review of existing publications about the non-Darcian flow to a pumping well indicates that three methods have commonly been used to solve this type of non-linear problem: the Boltzmann transform, the linearization method and numerical modeling. Sen (1987, 1988, 1989, 1990) has done extensive studies on non-Darcian flow to a pumping well by using the Boltzmann transform method. Recently, Wen et al. (2008a, b) have proposed a linearization procedure for solving non-Darcian flow to a well based on the assumption that the flow can be described by the Izbash equation. Meanwhile, numerical methods have also been used to solve non-Darcian flow problems (e.g., Ewing et al. 1999; Ewing and Lin 2001; Wu 2002a, b; Mathias et al. 2008; Wen et al. 2009). As commented by many scientists (e.g., Camacho-V and Vasquez-C 1992; Mathias et al. 2008), the Boltzmann transform cannot be used to solve such non-Darcian flow problems in a rigorous mathematical sense. The linearization procedure also has some limitations such as the discrepancies associated with early-time solutions. From this viewpoint, it seems that the present analytical methods can not solve such non-Darcian problems very well and numerical modeling might be a good approach in this area.

Up to now, most models on non-Darcian flow to a well are for confined aquifers such as, e.g., Sen (1987, 1990) and Wen et al. (2006, 2008a). Research on non-Darcian flow in leaky aquifers is too limited. Sen (2000) used a volumetric approach to investigate non-Darcian flow in leaky aquifers with the Izbash equation, and that study has been extended by Birpinar and Sen (2004) to investigate a similar problem with the Forchheimer equation. These two studies were based on the solutions obtained by the Boltzmann transform which was found to be problematic in a rigorous mathematical sense (Camacho-V and Vasquez-C 1992; Mathias et al. 2008). Therefore, the

solutions in these two studies might be questionable. Recently, Wen et al. (2008c) used a linearization procedure and finite difference method to solve non-Darcian flow in a leaky aquifer with the Izbash equation. As mentioned before, the Izbash equation has some disadvantages for describing radial non-Darcian flow. To the authors' knowledge, the investigations about non-Darcian flow in leaky aquifers with the Forchheimer equation are still very limited.

In this paper, a finite difference method will be used to investigate the non-Darcian flow to a pumping well in a leaky aquifer. The wellbore storage is also considered. The flow in the aquifer is assumed to be horizontal and non-Darcian, and the flow in the aquitard is assumed to be vertical and Darcian. The Forchheimer equation will be used to describe the non-Darcian flow in the aquifer. Sensitivity analysis for different parameters under non-Darcian flow conditions will be done in this study and the numerical solution will be applied to analyze the pumping test data in Chaj-Doab area of Pakistan. As Wen et al. (2008b) have investigated non-Darcian flow to a well in an aquifer-aquitard system with the Izbash equation, this study can be regarded as an extended work of Wen et al. (2008b).

## Problem statement

### Governing equations

The system discussed here is similar to that of Hantush and Jacob (1955), as shown in Fig. 1 (Wen et al. 2008b). The following assumptions have been used in this study to make the problem mathematically tractable. First, the aquifer and the bounded upper aquitard are assumed to be homogeneous, isotropic, and horizontally infinite. Second, the flow in the aquifer is assumed to be non-Darcian while the flow in the upper aquitard is assumed to be Darcian, and the flow direction in the aquifer is horizontal while the flow direction in the aquitard is vertical. Third, the aquitard storage is not considered and, fourth, the well fully penetrates the aquifer and the pumping rate is assumed to be constant. Under these assumptions, the mathematical model can be generated as follows:

$$\frac{1}{r} \frac{\partial [rq(r, t)]}{\partial r} - \frac{s(r, t)}{B} = \frac{S}{m} \frac{\partial s(r, t)}{\partial t}, \quad (1)$$

$$s(r, 0) = 0, \quad (2)$$

$$s(\infty, t) = 0, \quad (3)$$

$$2\pi rmq(r, t)|_{r \rightarrow 0} = -Q, \quad (4)$$

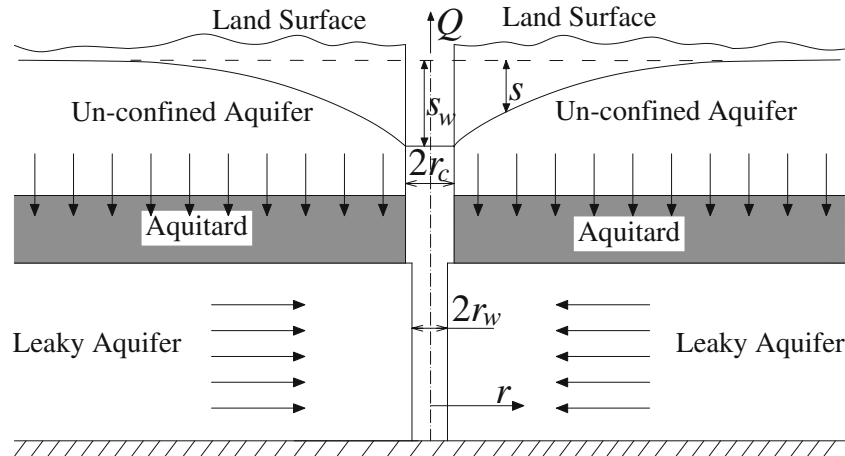


Fig. 1 A schematic diagram of the leaky confined aquifer system (Wen et al. 2008b, with permission from Elsevier)

in which  $r$  is radial distance from the center of the pumping well [L];  $t$  is pumping time [T];  $q(r,t)$  is specific discharge [L/T];  $s(r,t)$  is drawdown [L];  $S$  is storage coefficient of the aquifer;  $m$  is aquifer thickness [L];  $Q$  is pumping rate, which is constant [L<sup>3</sup>/T];  $B$  is the leakage parameter defined as  $m \times m_1/k_1$ , [LT], where  $m_1$  and  $k_1$  are thickness and hydraulic conductivity of the aquitard, respectively. It is notable that a greater value of  $B$ , meaning that a larger aquitard thickness  $m_1$  or a smaller aquitard hydraulic conductivity  $k_1$ , indicates smaller leakage effect. If the hydraulic conductivity of the aquitard  $k_1$  goes to zero, the leakage parameter  $B$  goes to infinity, then the problem investigated here is similar to the Theis model for confined aquifers.

If the wellbore storage cannot be ignored, the boundary condition Eq. (4) should be replaced by

$$2\pi r_w m q(r,t)|_{r=r_w} - \pi r_c^2 \frac{ds_w(t)}{dt} = -Q, \quad (5)$$

where  $r_w$  is the radius of the well screen [L],  $r_c$  is the radius of the well casing [L]. In most cases,  $r_c$  is larger than, instead of equal to  $r_w$ .  $s_w(t)$  is drawdown inside the well [L], which is dependent of the pumping time  $t$ .

The Forchheimer equation will be used to describe the non-Darcian flow in the aquifer. The Forchheimer equation can be expressed as:

$$q + \beta q|q| = k \frac{\partial s}{\partial r}, \quad (6)$$

in which  $\beta$  [T/L] is a non-Darcian factor representing the turbulence of the non-Darcian flow. If  $\beta$  is equal to zero, Eq. (6) becomes the well-known Darcy's law,  $k$  [L/T] can be regarded as the apparent hydraulic conductivity of the aquifer.

### Dimensionless transform

Similar to the analysis of the Theis type curves for Darcian flow, one can define the following dimensionless variables, as listed in Table 1. Notice that a minus sign was included in the definition of the dimensionless specific discharge  $q_D$ . It is

necessary to emphasize two important dimensionless variables, i.e.,  $B_D$  and  $\beta_D$ .  $B_D$  is a dimensionless variable representing the leakage and a larger  $B_D$  means smaller leakage.  $\beta_D$  is a dimensionless parameter representing the non-Darcian effect and a larger  $\beta_D$  indicates greater turbulence. With these definitions, the problem proposed here can be expressed in dimensionless form as:

$$-\frac{\partial[r_D q_D]}{r_D \partial r_D} - \frac{s_D}{B_D} = \frac{\partial s_D}{\partial t_D}, \quad (7)$$

$$s_D(r_D, 0) = 0, \quad (8)$$

$$s_D(\infty, t_D) = 0, \quad (9)$$

$$\frac{1}{2}(r_D q_D)|_{r_D \rightarrow 0} = 1. \quad (10)$$

The boundary condition considering wellbore storage can also be expressed in dimensionless format as:

$$\frac{1}{2}(r_D q_D)|_{r_D=r_{wD}} + \frac{r_{cD}^2}{4S} \frac{ds_{wD}(t_D)}{dt_D} = 1. \quad (11)$$

The Forchheimer equation will be changed to:

$$q_D + \beta_D q_D^2 = -\frac{\partial s_D}{\partial r_D}. \quad (12)$$

### Solutions for the problem

#### Numerical solution

The finite difference method was used to simulate the problem investigated here. It should be pointed out that

only the case with wellbore storage was considered when doing the simulation. The case without wellbore storage can be easily solved by setting very small values (close to zero) to the dimensionless radii  $r_w$  and  $r_c$ . First, one needs to discretize the dimensionless spatial domain  $[r_{wD}, r_{eD}]$ , where  $r_{eD}$  is a very large value which is used to approximate the outer boundary condition at which the drawdown is equal to zero. The value of  $r_D$  at the  $i$ th node can be referred as  $r_i$ , where the subscript represents the node of interest. For any node of  $r_i$ ,  $r_{wD} < r_i < r_{eD}$ ,  $i = 1, 2, \dots, N$ , and  $r_0 = r_{wD}$  and  $r_{N+1} = r_{eD}$ . As the drawdown changes rapidly near the pumping well, the dimensionless domain was discretized logarithmically (Wu 2002a; Mathias et al. 2008):

$$r_i = (r_{i-1/2} + r_{i+1/2})/2, \quad i = 1, 2, \dots, N, \tag{13}$$

with respect to

$$\log_{10}(r_{i+1/2}) = \log_{10}(r_{wD}) + i \left[ \frac{\log_{10}(r_{eD}) - \log_{10}(r_{wD})}{N} \right], \tag{14}$$

$i = 0, 1, \dots, N$ .

The governing equation, Eq. (7), can also be reduced to the following differential equation with respect to the dimensionless time  $t_D$ :

$$\frac{ds_i}{dt_D} \approx \frac{r_{i-1/2}q_{i-1/2} - r_{i+1/2}q_{i+1/2}}{r_i(r_{i+1/2} - r_{i-1/2})} - \frac{s_i}{B_D}, \tag{15}$$

$i = 1, 2, \dots, N$ ,

in which  $s_i$  and  $q_i$  are the dimensionless drawdown  $s_D$  and dimensionless specific discharge  $q_D$  at node  $i$ , respectively. From the Forchheimer equation, one has:

$$q_D = \frac{1}{2\beta_D} \left\{ -1 + \left[ 1 - 4\beta_D \frac{\partial s_D}{\partial r_D} \right]^{1/2} \right\}. \tag{16}$$

Therefore, the dimensionless specific discharge that goes into and out of each cell can be expressed as:

$$q_{i-1/2} \approx \frac{1}{2\beta_D} \left\{ -1 + \left[ 1 + 4\beta_D \left( \frac{s_{i-1} - s_i}{r_i - r_{i-1}} \right) \right]^{1/2} \right\}, \quad i = 2, 3, \dots, N, \tag{17}$$

$$q_{i+1/2} \approx \frac{1}{2\beta_D} \left\{ -1 + \left[ 1 + 4\beta_D \left( \frac{s_i - s_{i+1}}{r_{i+1} - r_i} \right) \right]^{1/2} \right\}, \quad i = 1, 2, \dots, N - 1, \tag{18}$$

. At the inner and outer boundaries, one has:

$$q_{1-1/2} \approx \frac{1}{2\beta_D} \left\{ -1 + \left[ 1 + 4\beta_D \left( \frac{s_{wD} - s_1}{r_1 - r_{wD}} \right) \right]^{1/2} \right\}, \tag{19}$$

$$q_{N+1/2} \approx \frac{1}{2\beta_D} \left\{ -1 + \left[ 1 + 4\beta_D \left( \frac{s_N - 0}{r_{eD} - r_N} \right) \right]^{1/2} \right\}, \tag{20}$$

where  $s_{wD}$  is the dimensionless drawdown inside the well, which can be approximated as follows by considering Eq. (11):

$$\frac{ds_{wD}}{dt_D} \approx \frac{4S}{r_{cD}^2} \left( 1 - \frac{1}{2} r_{wD} q_{1-1/2} \right). \tag{21}$$

After these preparations, the problem can be solved with the MATLAB software by using the stiff integrator ODE15s (Mathias et al. 2008). A MATLAB program named LeakyForch was developed to do this calculation. For all the simulations,  $N$  was chosen to be 5000 and  $r_{eD}$  was chosen to be  $10^8$ . These values are found to be sufficiently large for this study.

**The solution of Birpinar and Sen (2004)**

Birpinar and Sen (2004) proposed an analytical solution for the same problem by using the Forchheimer equation. They obtained the analytical solution for the type curves in leaky aquifers by using the volumetric approach proposed by Sen (1985), and on the basis of their previous study (Sen 1989). It is notable that the solution of Sen (1989) for the type curves in confined aquifers was obtained by the Boltzmann transform. Here the point is that the Boltzmann transform cannot be used to derive the type curves for non-Darcian flow problems, as pointed out by several scientists (e.g., Camacho-V and Vasquez-C 1992; Mathias et al. 2008; Wen et al. 2008a). In this study, the numerical solution will be used to verify the analytical solution obtained by Birpinar and Sen (2004) for the same problem. As stated by Birpinar and Sen (2004), the analytical solution (see Eq. (11) in their study) obtained by the volumetric approach can be changed to the following equation by using the dimensionless variables defined in this study:

$$s_D = \int_{\frac{r_D}{4B_D}}^{\infty} \frac{\exp\left(-x - \frac{r_D^2}{4B_D} \frac{1}{x}\right)}{\left(1 + \frac{2\sqrt{\pi}\beta_D\sqrt{x}}{r_D}\right)x} dx + \frac{2\beta_D}{r_D} \int_{\frac{r_D}{4B_D}}^{\infty} \frac{\exp\left(-2x - \frac{r_D^2}{4B_D} \frac{1}{x}\right)}{\left(1 + \frac{2\sqrt{\pi}\beta_D\sqrt{x}}{r_D}\right)^2 x} dx, \tag{22}$$

in which  $x$  is a dummy variable. The aforementioned equation can be integrated easily by using a MATLAB program.

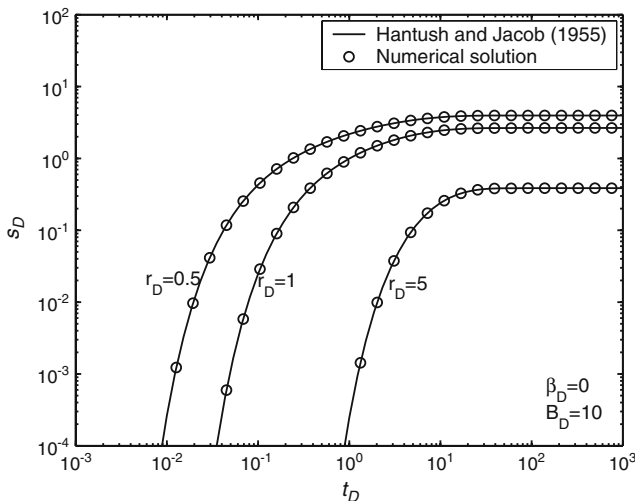
**Results and discussion**

In order to check the numerical solution proposed in this study, first the numerical solution was compared with the analytical solution of Hantush and Jacob (1955) for the Darcian flow case ( $\beta_D=0$ ). Three typical dimensionless distances were chosen as  $r_D=0.5, 1, \text{ and } 5$  and the dimensionless leakage parameter was given as 10. As shown in Fig. 2, the numerical solution agrees perfectly with the analytical solution during the entire pumping period. This indicates that the numerical error caused by the finite difference method probably can be ignored and the numerical solution in this study is sufficiently accurate.

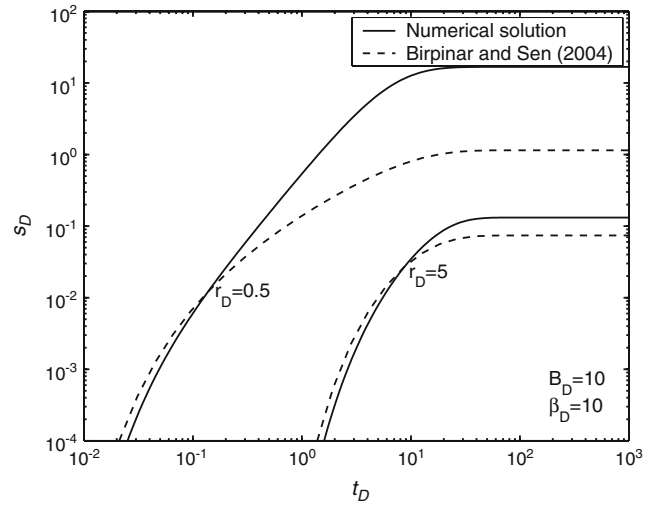
**Dimensionless drawdown without wellbore storage**

Figure 3 shows the comparison of the numerical solution used in this study and the solution of Birpinar and Sen (2004). The analytical solution of Birpinar and Sen (2004) was calculated from Eq. (22). The parameters used in Fig. 3 were given as  $\beta_D=10, B_D=10, r_D=0.5$  and 5. As can be seen from Fig. 3, it is evident that the solution obtained by Birpinar and Sen (2004) is not reliable. The method of the volumetric approach associated with the Boltzmann transform overestimates the dimensionless drawdown at early times and underestimates the dimensionless drawdown at late times. It can also be found that when the dimensionless distance is larger, the discrepancy between the analytical solution and the numerical solution is smaller, as expected.

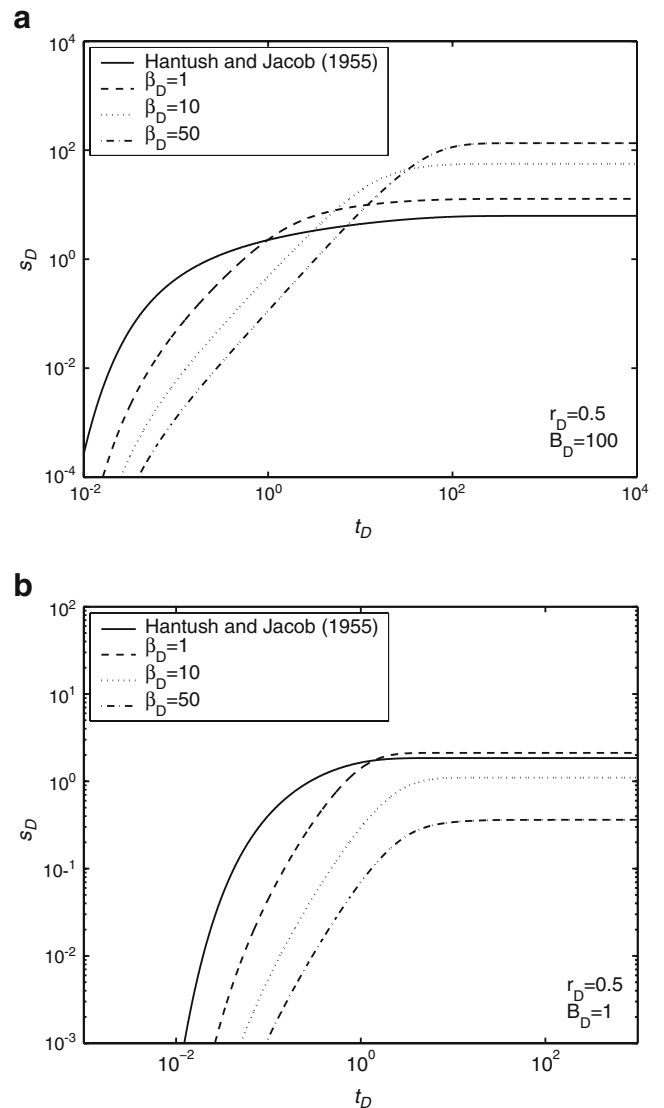
The sensitivity analysis of the dimensionless turbulent factor  $\beta_D$  on the dimensionless drawdown is shown in Fig. 4. Two typical dimensionless leakage parameters  $B_D=100$  and 1 were used in this figure, as shown in Fig. 4a and b. The other parameters were given as:  $r_D=0.5, \beta_D=1, 10$  and 50. The solution of Hantush and Jacob (1955) has also been depicted as a reference. Interestingly enough, the features of these two figures are quite



**Fig. 2** Comparison of the numerical solution and the solution of Hantush and Jacob (1955) for the Darcian flow case



**Fig. 3** Comparison of the numerical solution and the solution of Birpinar and Sen (2004) with  $B_D=10, \beta_D=10, r_D=0.5$  and 5

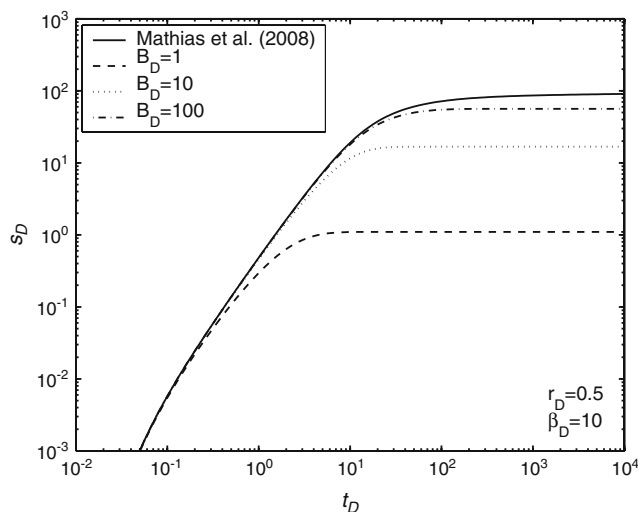


**Fig. 4** Drawdown-time behavior for different values of the turbulent factor  $\beta_D$  with  $r_D=0.5, \beta_D=1, 10$  and 50. **a**  $B_D=100$ ; **b**  $B_D=1$



different. For the case  $B_D=100$ , a relatively large value for the dimensionless leaky parameter which indicates a relatively small leakage effect, a larger  $\beta_D$  results in a smaller dimensionless drawdown at early times and leads to a larger dimensionless drawdown at late times. This phenomenon can be explained as follows. The Forchheimer equation can be rewritten as:  $q = K_\beta(\partial s/\partial x)$ , in which  $K_\beta = k/(1 + \beta|q|)$ , and  $K_\beta$  can be regarded as apparent “Darcian” hydraulic conductivity. If  $\beta_D$  is larger, which also means a larger  $\beta$ , then  $K_\beta$  is smaller. At early times, the flow has not approached the steady state and the specific discharge  $q$  increases from zero to the stationary  $q_0$  at steady state. A smaller  $K_\beta$  indicates that it will take a longer time for the flow to approach the steady state. Thus, at the same time, a smaller  $K_\beta$  has a smaller specific discharge. Consequently, a smaller drawdown can be found from the equation  $q = K_\beta(\partial s/\partial x)$ . At late times, the flow approaches the steady state and the specific discharge does not change any more, which is only proportional to the pumping rate  $Q$ . From the equation  $q = K_\beta(\partial s/\partial x)$ , one can see a smaller  $K_\beta$  will result in a larger drawdown. As reflected in Fig. 4a, a larger  $K_\beta$  results in a larger dimensionless drawdown at late times. For the case  $B_D=1$  shown in Fig. 4b, a relatively small value for the dimensionless leaky parameter which indicates a relatively great leakage effect, a larger  $\beta_D$  results in a smaller dimensionless drawdown during the entire pumping period. This might be because a large leakage from the aquitard will result in the system approaching the quasi steady state earlier. Therefore, the “early-time” behavior of the type curve is very short and not long enough to be shown in the figure. Figure 4b might only reflect the “large-time” behavior of the type curve; therefore, it is no wonder that a larger  $\beta_D$  results in a smaller drawdown during the entire pumping period, as shown in Fig. 4b.

It is also worthwhile to analyze the effect of the leakage on the drawdown. Figure 5 is about the dimensionless drawdown versus dimensionless time for



**Fig. 5** Drawdown-time behavior for different values of the leakage factor  $B_D$  with  $r_D=0.5$ ,  $\beta_D = 10$ ,  $B_D=1, 10$  and  $100$

different dimensionless leakage parameters. The parameters were given as:  $r_D=0.5$ ,  $\beta_D=10$ ,  $B_D=1, 10$  and  $100$ . The case without the leakage (Mathias et al. 2008) has also been depicted in this figure as a reference. As shown in Fig. 5, all the curves for different  $B_D$  values approach the same asymptotic value at early times, as well as the case of Mathias et al. (2008). This is because water from the leakage has not arrived at the confined aquifer at early times; therefore, the leakage has little impact on the drawdown at early times. As the pumping time increases, great differences have been found for different  $B_D$ . When  $B_D$  is larger, the dimensionless drawdown is larger. This is because a larger  $B_D$  means a smaller leakage effect, thus a larger dimensionless drawdown at late times. It can also be found that when  $B_D$  is larger, the curve is closer to that obtained by Mathias et al. (2008), as expected.

### Type curves

It is useful to generate a series of type curves for the hydrology community to conduct well testing analysis. When conducting a pumping test, the position of the observation well  $r$  and the aquifer thickness  $m$  are usually known. This means one can obtain  $r_D$  before analyzing the pumping test data. For a specific value of  $r_D$ , one can obtain a series of type curves for different dimensionless parameters. Taking  $r_D=0.5$  as an example, the corresponding type curves for different values of the dimensionless parameters are shown in Fig. 6. For a specific pumping test, if  $r_D$  happens to equal 0.5, one can use these type curves to estimate the aquifer parameters (an example will be given in the following section). The type curves for different  $r_D$  values can also be obtained by the MATLAB program, but it is impossible to report all the type curves in this paper. For the hydrology community use, one can get the MATLAB program upon request.

### Dimensionless drawdown with wellbore storage

When the wellbore storage is considered, the dimensionless drawdowns inside the well for different turbulent factor  $\beta_D$  and different leakage parameter  $B_D$  were analyzed. As shown in Figs. 7 and 8, all the curves for different turbulent factor  $\beta_D$  and different leakage parameter  $B_D$  approach the same asymptotic value at early times while significant differences have been found at late times. Similar to the features of the drawdowns in the aquifer without the wellbore storage, a larger  $\beta_D$  (or  $B_D$ ) results in a larger drawdown inside the well. The drawdowns in the aquifer when considering the wellbore storage have also been analyzed. As the features are similar to those without the wellbore storage, it is not repeated here.

### Sensitivity analysis

In order to assess the influence of different parameters on the results, the sensitivity analysis was done on different parameters, i.e.,  $B_D$ ,  $\beta_D$ ,  $r_{wD}$ , and  $r_{cD}$ . This analysis is

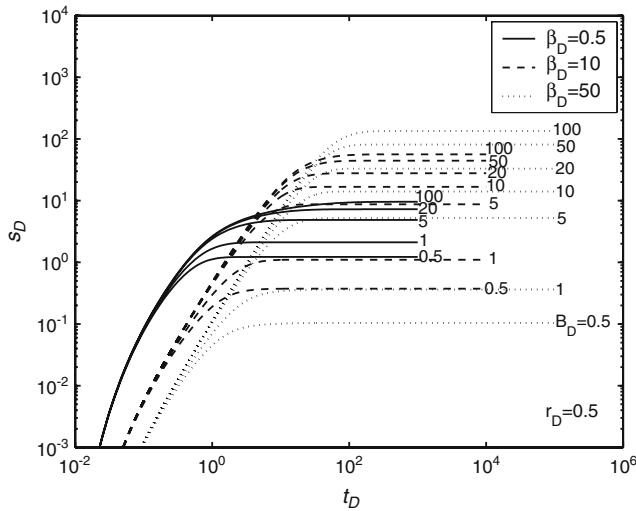


Fig. 6 Type curves for  $r_D=0.5$ ,  $B_D=0.5$ ,  $\beta_D = 0.5, 10$  and  $50$

useful in assessing how a conceptual model responds to the change in certain variables. As proposed by Liou and Yeh (1997) and Huang and Yeh (2007), the sensitivity of a dependent variable can be defined as:

$$X_{i,j} = \frac{\partial R_i}{\partial P_j}, \tag{23}$$

in which  $X_{i,j}$  is the sensitivity coefficient of the  $j$ th parameter  $P_j$  at the  $i$ th time.  $R_i$  is the dependent variable at the  $i$ th time, which is the dimensionless drawdown in this study. Huang and Yeh (2007) have proposed a normalized sensitivity method to assess the effect of different variables on the dependent variable, which is defined as:

$$X'_{i,j} = P_j \frac{\partial R_i}{\partial P_j}, \tag{24}$$

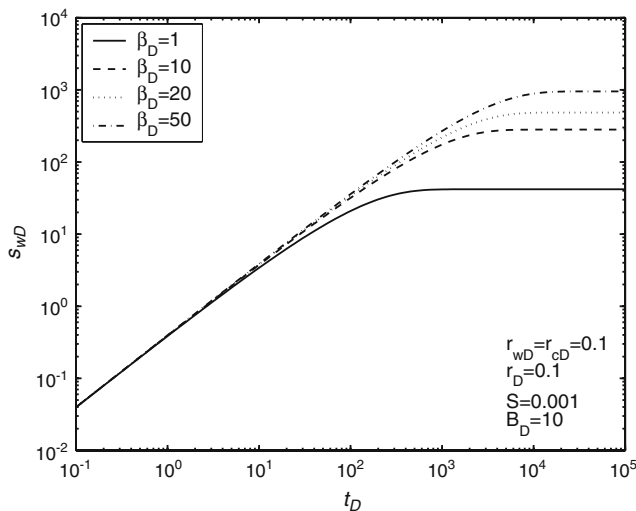


Fig. 7 Drawdown-time behavior inside the well when considering the wellbore storage for different values of the turbulent factor  $\beta_D$  with  $r_D=0.1$ ,  $r_{wD}=r_{cD}=0.1$ ,  $S=0.001$ ,  $B_D=10$ ,  $\beta_D = 1, 10, 20$  and  $50$

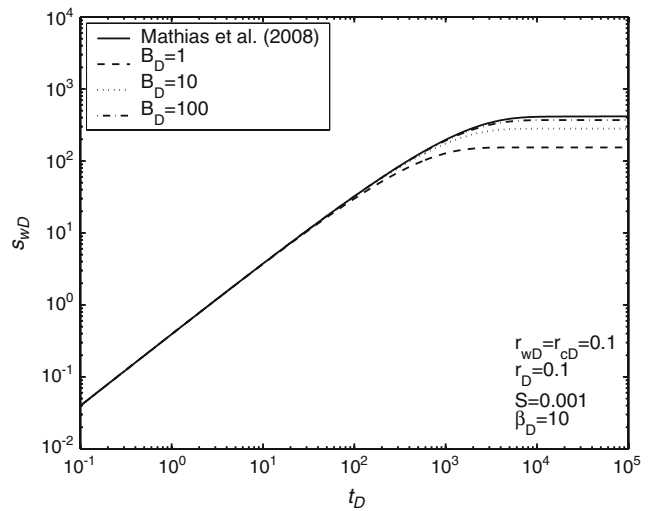


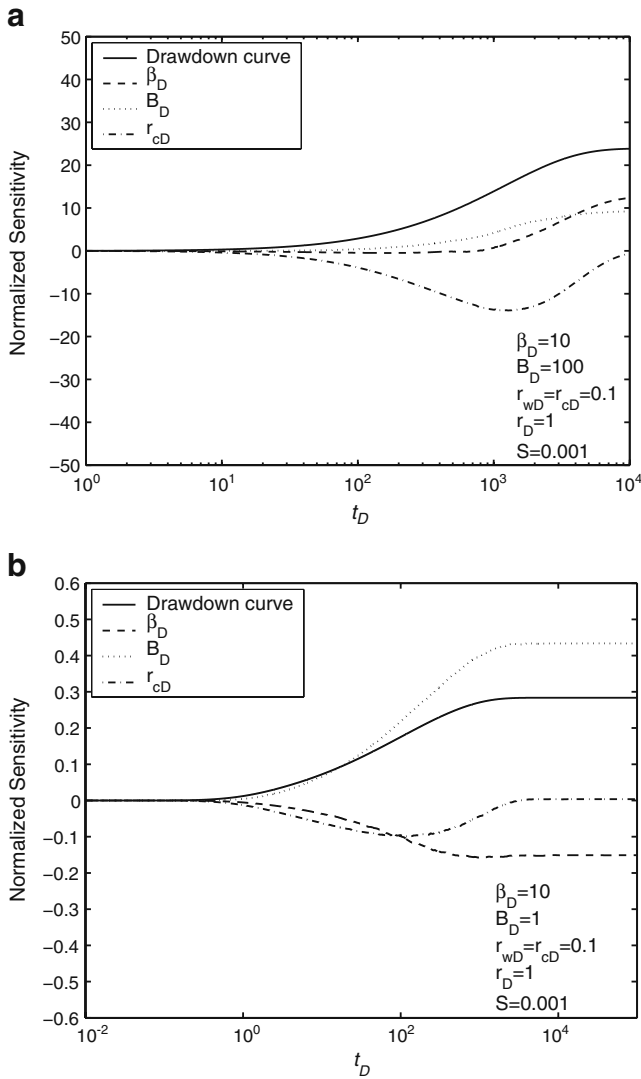
Fig. 8 Drawdown-time behavior inside the well when considering the wellbore storage for different values of the leakage factor  $B_D$  with  $r_D=0.1$ ,  $r_{wD}=r_{cD}=0.1$ ,  $S=0.001$ ,  $\beta_D = 10$ ,  $B_D = 1, 10$  and  $100$

where  $X'_{i,j}$  is the normalized sensitivity of the  $j$ th parameter  $P_j$  at the  $i$ th time. Notice that there is a partial derivative on the right hand of the Eq. (24), which is always difficult to obtain for a specific case. A finite difference formula will be used to approximate this differentiation (Yeh 1987), that is

$$\frac{\partial R_i}{\partial P_j} = \frac{R_i(P_j + \Delta P_j) - R_i(P_j)}{\Delta P_j}, \tag{25}$$

in which  $\Delta P_j$  is a small increment, which will be chosen as  $10^{-2} \times P_j$  (Yang and Yeh 2009). The two most important dimensionless parameters used in this study are  $\beta_D$  and  $B_D$ . When the wellbore storage is considered,  $r_{cD}$  is also an important parameter which reflects the water stored in the wellbore. Thus, normalized sensitivities of these parameters will be analyzed in the following. The hypothetical data used were  $\beta_D = 10$ ,  $r_{wD}=r_{cD} = 0.1$ ,  $r_D=1$ ,  $S=0.001$ , and  $B_D=100$  (or 1). The reason for not analyzing the sensitivities of  $r_{wD}$  and  $S$  are as follows: (1)  $r_{wD}$  is the dimensionless radius of the well screen which is assumed to be equal to  $r_{cD}$  in this study; (2) the storage coefficient  $S$  has been used in the definition of the dimensionless variable  $t_D$ .

Figure 9a and b plot the dimensionless time-drawdown curve and the normalized sensitivities of the parameters  $\beta_D$ ,  $B_D$  and  $r_{cD}$  in semi-log scales. Figure 9a is for a case with a relatively small leakage  $B_D=100$ , while Fig. 9b is for a case with a relatively large leakage  $B_D=1$ . It can be seen that the normalized sensitivity of the drawdown with respect to  $B_D$  has a positive effect from both parts a and b of Fig. 9. This feature is quite distinct in Fig. 9b and this is because when  $B_D$  is larger, the leaky rate is smaller, resulting in a larger drawdown. For  $B_D=1$ , the leakage is very large, then this positive phenomenon is evident. A relative change in  $\beta_D$  has a positive effect for  $B_D=100$  in Fig. 9a, while it has a negative effect for  $B_D=1$  in Fig. 9b. This is consistent with the features of Fig. 4a and b. The



**Fig. 9** Drawdown-time curve and the normalized sensitivities of the parameters  $\beta_D$ ,  $B_D$  and  $r_{cD}$ . **a**  $B_D=100$ ; **b**  $B_D=1$

normalized sensitivity with respect to  $r_{cD}$  produces a negative effect at early times and stabilizes at zero after a relatively large dimensionless time  $t_D$  (say,  $10^4$ ). This can be easily understood by considering the physics of the wellbore storage. Comparing parts a and b of Fig. 9, one can see that  $r_{cD}$  produces the largest normalized sensitivity in the magnitude when  $B_D$  is relatively large ( $B_D=100$ ), while  $B_D$  produces the largest normalized sensitivity in the magnitude when  $B_D$  is relatively small ( $B_D=1$ ). Those results indicate that the drawdown is very sensitive to the change in  $r_{cD}$  when  $B_D$  is large, while it is sensitive to the change in  $B_D$  when  $B_D$  is relatively small.

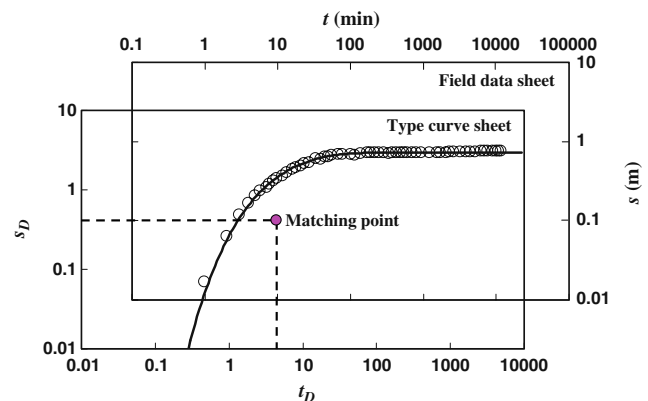
It seems that a relative change of these three parameters has no effect on the drawdown at early times (say,  $t_D < 10^0$ ), as shown in Fig. 9a. The truth is that the value of the dimensionless drawdown at early times is very small (see Fig. 4), consequently, the corresponding normalized sensitivity is a very small value around zero.

**Application**

The numerical solution developed in this study can be used to determine the aquifer parameters associated with the pumping test data. In order to show the applications of the solutions, the data from Ahmad (1998) was used, who did a pumping test in the Chaj-Doab area near Gujrat water distributory in Pakistan. The pumping rate is fixed at  $Q=3.77\text{m}^3/\text{min}$ , the thickness of the aquifer is measured as  $m=76\text{m}$ , and the observation well is about  $r=122\text{m}$  away from the pumping well. Therefore, the corresponding dimensionless distance  $r_D$  is equal to 1.61. As stated by Ahmad (1998), the aquifer in that area can be classified as a mixture of leaky and unconfined aquifers (Birpinar and Sen 2004). Therefore, it is possible to use the type curves developed in this study to estimate the aquifer parameters. The observation data of time-drawdown was plotted in log-log scale as shown in Fig. 10. After a trial and error process, the curve with  $r_D=1.6053$ ,  $B_D=25$  and  $\beta_D=0.4$  matched the observation data perfectly, as shown in Fig. 10. A matching point was chosen in the common area so that one could obtain the corresponding coordinates:  $s=0.1\text{m}$ ,  $t=10\text{min}$ ,  $s_D=0.36$  and  $t_D=5.68$ . When the matching procedure is done, one can do the following calculations.

1. From the definitions of the dimensionless drawdown  $s_D = \frac{4\pi km}{Q}s$ , one can obtain the apparent hydraulic conductivity  $k = \frac{s_D Q}{4\pi ms} = \frac{0.36 \times 3.77}{4 \times 3.14 \times 76 \times 0.1} = 0.014[\text{m}/\text{min}]$ .
2. When  $k$  is known, the storage coefficient  $S$  can be calculated as:  $S = \frac{kt}{mt_D} = \frac{0.014 \times 10}{5.68 \times 76} = 3.24 \times 10^{-4}$ .
3. With the dimensionless definition  $\beta_D = \frac{\beta Q}{4\pi m^2}$ , one has:  $\beta = \frac{4\pi m^2 \beta_D}{Q} = \frac{4 \times 3.14 \times 76^2 \times 0.4}{3.77} = 7697[\text{min}/\text{m}]$ .
4. Finally, the leakage parameter  $B$  can be calculated as:  $B = \frac{B_D m^2}{k} = \frac{25 \times 76^2}{0.014} = 1.031 \times 10^7$ . As  $B$  is defined as  $m \times m_1/k_1$ , if the thickness of the aquitard is known, the hydraulic conductivity of the aquitard can be obtained subsequently.

As discussed in section Type curves, for a specific given distance  $r_D$ , one can generate a series of type curves. When



**Fig. 10** Type curve and field data matching sheets. The field data were obtained from Birpinar and Sen (2004)



the pumping data are available, one can estimate the aquifer parameters through the matching processes described in the previous with the type curves and the observed data.

## Summary and conclusions

A numerical solution for non-Darcian flow to a well in leaky aquifers has been obtained by using the finite difference method in this study. The Forchheimer equation has been used to describe the flow in the aquifer. The impacts of the turbulent factor and the leakage parameter on the drawdowns have been analyzed. The results were also compared with the solution of Birpinar and Sen (2004) who investigated a similar problem with a volumetric approach. The solutions obtained in this study can also be used to estimate the aquifer parameters when the groundwater flow is non-Darcian. Several findings can be presented from this study:

1. The method of the volumetric approach associated with the Boltzmann transform (Birpinar and Sen 2004) overestimates the dimensionless drawdown at early times and underestimates the dimensionless drawdown at late times. When the dimensionless distance is larger, the discrepancy between the analytical solution and the numerical solution becomes smaller.
2. For a relatively larger dimensionless leakage parameter, e.g.,  $B_D=100$ , a larger turbulent factor  $\beta_D$  results in a smaller drawdown at early times and a larger drawdown at late times. While for a relatively smaller dimensionless leakage parameter, e.g.,  $B_D=1$ , a larger  $\beta_D$  results in a smaller drawdown during the entire pumping period.
3. The impact of the leakage on the drawdown is similar to that of Darcian flow; the leakage has little impact on the drawdown at early times, while a greater leakage parameter  $B_D$  leads to a larger drawdown at late times.
4. The drawdown is very sensitive to the change in  $r_{cD}$  reflecting the importance of wellbore storage when  $B_D$  is large, while it is sensitive to the change in  $B_D$  when  $B_D$  is relatively small.

**Acknowledgements** This research was partially supported by the National Natural Science Foundation of China (Grant Nos. 41002082, 50779067), the National Basic Research Program of China (Grant No. 2010CB428802), and the Special Fund for Basic Scientific Research of Central Colleges, China University of Geosciences (Wuhan; Grant No. CUGL090301). The constructive comments of two anonymous reviewers and the Editor are also gratefully acknowledged. The authors also sincerely thank the Technical Editorial Advisor, Sue Duncan, for carefully checking the manuscript.

## References

Ahmad N (1998) Evaluation of groundwater resources in the upper middle part of Chaj-Doak area, Pakistan. PhD Thesis, Istanbul Technical Univ., Turkey

- Basak P (1977) Non-penetrating well in a semi-infinite medium with non-linear flow. *J Hydrol* 33:375–382
- Basak P (1978) Analytical solutions for two-regime well flow problems. *J Hydrol* 38:147–159
- Birpinar ME, Sen Z (2004) Forchheimer groundwater flow law type curves for leaky aquifers. *J Hydrol Eng* 9(1):51–59
- Bordier C, Zimmer D (2000) Drainage equations and non-Darcian modeling in coarse porous media or geosynthetic materials. *J Hydrol* 228:174–187
- Camacho-V RG, Vasquez-C M (1992) Comment on “Analytical solution incorporating nonlinear radial flow in confined aquifers” by Zekai Sen. *Water Resour Res* 28(12):3337–3338
- Choi ES, Cheema T, Islam MR (1997) A new dual-porosity/dual permeability model with non-Darcian flow through fractures. *J Petrol Sci Eng* 17(3–4):331–344
- Darcy H (1856) *Les Fontaines publiques de la ville de Dijon* [The public fountains of the city of Dijon]. Dalmont, Paris
- Dudgeon CR (1966) An experimental study of the flow of water through coarse granular media. *Houille Blanche* 7:785–801
- Escande L (1953) Experiments concerning the filtration of water through rock mass. Proceedings of Minnesota International Hydraulics Convention, September 1953, New York
- Ewing RE, Lazarov RD, Lyons SL, Papavassiliou DV, Pasciak J, Qin G (1999) Numerical well model for non-Darcy flow through isotropic porous media. *Comp Geosci* 3:185–204
- Ewing RE, Lin Y (2001) A mathematical analysis for numerical well models for non-Darcy flows. *App Num Math* 39(1):17–30
- Forchheimer PH (1901) *Wasserbewegung durch Boden* [Movement of water through soil]. *Zeitschr Ver deutsch Ing* 49:1736–1749 and 50:1781–1788
- Giorgi T (1997) Derivation of the Forchheimer law via matched asymptotic expansions. *Transp Porous Med* 29(2):191–206
- Harr ME (1962) *Ground water and seepage*. McGraw-Hill, New York, 410 pp
- Hantush MS, Jacob CE (1955) Non-steady radial flow in an infinite leaky aquifer. *Trans Am Geophys Union* 36(1):95–100
- Huang YC, Yeh HD (2007) The use of sensitivity analysis in on-line aquifer parameter estimation. *J Hydrol* 335(3–4):406–418
- Irmay S (1958) On the theoretical derivation of Darcy and Forchheimer formulas. *Trans Am Soc Geophys Union* 39:702–707
- Izbash SV (1931) *O filtracii v kropnozernstom materiale* [Groundwater flow in the material kropnozernstom?]. *Izv. Nauchnoissled, Inst. Gidrotechniki (NIIG), Leningrad, USSR*
- Li J, Huang G, Wen Z, Zhan H (2008) Experimental study on non-Darcian flow in two kinds of media with different diameters (in Chinese with English abstract). *J Hydraul Eng* 39(6):726–732
- Liou TS, Yeh HD (1997) Conditional expectation for evaluation of risk groundwater flow and transport: one-dimensional analysis. *J Hydrol* 199:378–402
- Mathias S, Butler A, Zhan H (2008) Approximate solutions for Forchheimer flow to a well. *J Hydraul Eng* 134(9):1318–1325
- Muskat M (1937) *The flow of homogeneous fluids through porous media*, 2nd edn. McGraw-Hill, New York, Edwards, Ann Arbor, MI
- Polubarinova-Kochina P (1962) *Theory of ground water movement*. Translated by J.M. De Wiest. Princeton University Press, Princeton, NJ
- Rose HE (1951) Fluid flow through beds of granular material: some aspects of fluid flow. Arnold, London, pp 136–162
- Sen Z (1985) Volumetric approach to type curves in leaky aquifers. *J Hydraul Eng* 111(3):467–484
- Sen Z (1987) Non-Darcian flow in fractured rocks with a linear flow pattern. *J Hydrol* 92:43–57
- Sen Z (1988) Type curves for two-region well flow. *J Hydraul Eng* 114(12):1461–1484
- Sen Z (1989) Nonlinear flow toward wells. *J Hydraul Eng* 115(2):193–209
- Sen Z (1990) Nonlinear radial flow in confined aquifers toward large-diameter wells. *Water Resour Res* 26(5):1103–1109
- Sen Z (2000) Non-Darcian groundwater flow in leaky aquifers. *Hydrolog Sci J* 45(4):595–606

- Slepicka F (1961) Hydraulic function of cylindrical well in an artesian aquifer with regard to new research on flow through porous media. Proceedings of the 9th World Congress of the International Association of Hydraulic Research, vol 1, Dubrovnik, Czechoslovakia, 395 pp
- Sorek S, Levi-Hevroni D, Levy A, Ben-Dor G (2005) Extensions to the macroscopic Navier–Stokes equation. *Trans Porous Med* 61:215–233
- Whitaker S (1996) The Forchheimer equation: a theoretical development. *Transp Porous Med* 49(2):1573–1634
- Wilkinson JK (1956) The flow of water through rockfill and application to the design of dams. Proceedings of the 2nd Australian-New Zealand Conference on Soil Mechanics and Foundation Engineering, Christchurch, New Zealand, 1956, 141 pp
- Wen Z, Huang G, Zhan H (2006) Non-Darcian flow in a single confined vertical fracture toward a well. *J Hydrol* 330:698–708
- Wen Z, Huang G, Zhan H (2008a) An analytical solution for non-Darcian flow in a confined aquifer using the power law function. *Adv Water Resour* 31:44–55
- Wen Z, Huang G, Zhan H (2008b) Non-Darcian flow to a well in an aquifer–aquitard system. *Adv Water Resour* 31:1754–1763
- Wen Z, Huang G, Zhan H, Li J (2008c) Two-region non-Darcian flow toward a well in a confined aquifer. *Adv Water Resour* 31:818–827
- Wen Z, Huang G, Zhan H (2009) A numerical solution for non-Darcian flow to a well in a confined aquifer using the power law function. *J Hydrol* 364:99–106
- Wu YS (2002a) Numerical simulation of single-phase and multi-phase non-Darcian flow in porous and fractured reservoirs. *Transp Porous Med* 49:209–240
- Wu YS (2002b) An approximate analytical solution for non-Darcy flow toward a well in fractured media. *Water Resour Res* 38(3), 1023. doi:10.1029/2001WR000713
- Yang SY, Yeh HD (2009) Radial groundwater flow to a finite diameter well in a leaky confined aquifer with a finite-thickness skin. *Hydrol Process* 23:3382–3390
- Yeh HD (1987) Theis' solution by nonlinear least-squares and finite-difference Newton's method. *Ground Water* 25:710–715