RESEARCH

Prachařite, CaSb⁵⁺₂(As³⁺₂O₅)₂O₂·10H₂O, a new mineral from Lavrion, **Greece**

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Abstract

Prachařite, ideally $CaSb^{5+}$ ₂ $(As^{3+}$ ₂ $O_5)$ ₂ O_2 ·10 H_2O , is a new mineral found in underground workings of the Plaka Mine No. 80, Plaka, Lavrion Mining District, Attica, Greece. It occurs as colourless to white, thin tabular hexagonal, in general sharp crystals up to 2.5 mm in diameter, and is associated with pharmacolite, sulphur and very rare smamite $\{Ca_2Sb(OH)_4[H(ASO_4)_2]\cdot 6H_2O\}$ on a matrix composed of sphalerite, galena and carbonate gangue. Prachařite is translucent to transparent, with a glassy lustre, white streak, a good cleavage parallel to {0001} and a distinct cleavage parallel to {1010}. It is non-luminescent, brittle, and has an uneven fracture, a Mohs hardness of 2–2.5 and X-ray density $D_x = 2.848$ g/cm³, $D_{calc} = 2.836 - 2.853$ g/cm³ (for two measured compositions). Optically, it is uniaxial negative, with $\omega = 1.619(1)$ and $\varepsilon = 1.553(1)$. Prachařite is trigonal, space group $\overline{P3c1}$ (no. 165), with $a = 13.951(2)$, $c = 19.899(2)$ Å, $V = 3354.1(10)$ Å³ and $Z = 6$. Strongest lines in the X-ray powder difraction pattern are [*d* in Å (*I*) *hkl*]: 9.894 (100) 002; 6.045 (8) 200; 5.156 (10) 202; 4.946 (11) 004; 3.297 (19) 311, 006, 222; 2.988 (22) 400, 313, 116. Two sets of independent electron probe micro-analyses yielded (wt%): CaO 6.28/7.12, MgO 0.09/-, Zn -/0.01, Sb₂O₅ 39.22/40.19, As₂O₃ 47.59/47.39, SO₃ -/0.02, H₂O 21.65/22.04 (calculated on the basis of ideal composition derived from crystal-structure determination), total 114.83/116.77; the total is reproducibly high due to a loss of a third of all water molecules under the electron beam. The empirical formulae, based on $O=22$ atoms per formula unit, for the two datasets are very similar, $(Ca_{0.93}Mg_{0.02})_{\Sigma 0.95}Sb_{2.02}(AsO_3)_{4.00}$ 10H₂O and $Ca_{1.04}Sb_{2.03}(AsO_3)_{3.92}$ 10H₂O. The ideal formula is $CaSb^{5+}{}_{2}(As^{3+}{}_{2}O_5)_{2}O_2$ 10H₂O, determined with the help of a crystal-structure determination based on single-crystal X-ray diffraction datasets collected at room temperature $(R1 = 2.3\%)$. The atomic arrangement of prachařite is unusual; it is based on two different layers containing a six-membered ring of corner-sharing $SbO₆$ octahedra, an eight-coordinated Ca1 atom in the centre of the ring, two non-equivalent AsO₃ groups corner-linked to form a $(As_2O_5)^{4-}$ diarsenite group, and, on interlayer sites, a seven-coordination Ca2 atom and three water molecules (all only weakly hydrogen-bonded), one of which is only partially occupied (split position). The mineral is named in honour of Dr Ivan Prachař, a long-term researcher of the mineralogy and underground workings of Lavrion.

Keywords Prachařite · New mineral · Crystal structure · Plaka Mine No. 80 · Lavrion Mining District

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Introduction

During some of the authors' (UK, BR, KHF) ongoing studies of the mineralogy and mines of the famous Lavrion mining district in Greece (Bonsall et al. [2011](#page-11-0); Schefer et al. [2019](#page-11-1) and references therein), we encountered, in 2016, a colourless to white supergene mineral on weathered lead–zinc ore, whose thin tabular, six-sided crystal morphology did not remind us of any known species. In fact, a full characterisation of the mineral, including a determination of its crystal structure, showed that it is a new, unusual hydrated oxysalt species containing Ca, Sb(V) and As(III). The mineral and its name were subsequently approved by the International Mineralogical Association's Commission on New Minerals, Nomenclature and Classifcation (IMA 2018-081).

The mineral is named after Dr Ivan Prachař (b. 1957) of Prague, Czech Republic. Dr Prachař has, in close collaboration with one of the authors (KHF), studied the mineralogy of the numerous Lavrion mines for more than 20 years and helped to carefully survey and document all the underground workings (ancient to modern). Several discoveries of rare and unusual minerals in Lavrion are due to his work. The name is pronounced pra-har-ait.

The present contribution provides a full characterisation of prachařite and a comparison with related species. Cotype material is deposited in the Natural History Museum, Vienna, Austria (one cotype specimen and the crystal used for the crystal-structure solution; catalogue no. NHMW-MIN-O357), in the National Museum, Prague, Czech Republic (one cotype specimen, catalogue no. P1P 15/2018), and in the Natural History Museum of Los Angeles County, USA (one cotype specimen, catalogue no. 67242).

Occurrence and paragenesis

Prachařite was found in September 2016 in underground workings of the Plaka Mine No. 80, Plaka, Lavrion Mining District, Lavreotiki, Attica, Greece. The Plaka Mine No. 80 is located in the northern part of the famous Lavrion mining district (Leleu [1966](#page-11-2), [1969;](#page-11-3) Skarpelis [2007](#page-11-4); Voudouris et al. [2008;](#page-11-5) Skarpelis and Argyraki [2009;](#page-11-6) Bonsall et al. [2011](#page-11-0); Scheffer et al. [2019](#page-11-1)). This mine is known for late-stage hydrothermal As-Sb-Pb–Zn(-Ag) ore veins with assemblages of primary silver minerals and a large variety of supergene phases, mainly arsenates (Wendel and Markl [1996;](#page-11-7) Rieck and Rieck [1999;](#page-11-8) Rieck et al. [2018,](#page-11-9) [2020,](#page-11-10) [2022](#page-11-11); Frenzel et al. [2022\)](#page-11-12). About 133 mineral species are currently (May 2023) known from the mine.

On the collected material, prachařite is accompanied by pharmacolite, a few tiny sulphur crystals and very rare smamite ${Ca_2Sb(OH)_4[H(AsO_4)_2]\cdot 6H_2O}$ (the latter only as a unique crystal, adjacent to a prachařite crystal, on

a single, very small specimen; Rieck et al. [2022](#page-11-11)). These supergene phases crystallised on an ore matrix composed of dark brownish sphalerite, galena and carbonate gangue (dolomite and calcite). The vein-type primary ore mineralisation at this locality is composed of sphalerite, galena, pyrite, arsenic, stibarsen, stibnite and other phases. Typical secondary minerals include pharmacolite, picropharmacolite (very common), various Ca arsenates [including the unnamed phase $Ca_5(AsO_4)_2(HAsO_4)_2.5H_2O$, Kolitsch et al. [2014\]](#page-11-13), antimony oxides, sulphur and others.

Morphological, physical and optical properties

Prachařite forms colourless to white, thin tabular hexagonal, in general sharp crystals up to 2.5 mm in diameter, but mostly 1–1.5 mm (Figs. [1](#page-1-0) and [2\)](#page-2-0). Parts of the crystals may appear somewhat crumbly, with recognisable cleavage cracks. The colourless platelets may show a frosted tabular face and a very thin white opaque rim. Prachařite is translucent to transparent, with a glassy lustre, white streak, a good cleavage parallel to {0001} and a distinct cleavage parallel to {1010}. It is non-luminescent, brittle, and has an uneven fracture and a Mohs hardness of 2–2.5. Optically, it is uniaxial negative, with $\omega = 1.619(1)$ and $\varepsilon = 1.553(1)$ (determined in white light) and nonpleochroic. Mass density was not measured directly; the X-ray density D_x is 2.848 g/ cm³ (based on single-crystal data measured in Vienna—see below). For the ideal composition and the Vienna singlecrystal data $D_{\text{calc}} = 2.838$ g/cm³; for the measured compositions (see following section), $D_{\text{calc}} = 2.836 - 2.853$ g/cm³.

Fig. 1 Colourless thin tabular hexagonal crystal of prachařite. Area on cotype specimen P1P 15/2018. White-light photomicrograph Jiří Sejkora

Fig. 2 Colourless thin tabular hexagonal crystal of prachařite associated with white sprays of indistinct acicular pharmacolite crystals. Area on co-type specimen NHMW-MIN-O357. White-light photomicrograph Harald Schillhammer

Chemical composition

Chemical analyses of prachařite were carried out with two electron probe micro-analysers, with very similar results for the standard polished mounts. The cotype material in Prague was quantitatively analysed with a Cameca SX 100 instrument (wavelength-dispersive mode, 15 kV, 4 nA, 10 μm defocussed beam diameter). Contents of Al, Bi, Cl,

Table 1 Results of EPMA chemical analyses of prachařite

Co, Cu, Fe, K, Mn, Na, Ni, P, Pb, S, Si, V and Zn were all below detection limits. The following X-ray lines and reference materials were used: (i) *K*α lines: Ca (wollastonite), Mg (synthetic Mg₂SiO₄), *L*α line: As (lammerite); (ii) *L*β line: Sb (synthetic Sb). Counting times were 20 s on peak and 10 s on each background position. The raw intensities were converted to the concentrations automatically using the *PAP* (Pouchou and Pichoir [1985\)](#page-11-14) matrixcorrection software. The $H₂O$ content could not be determined directly, because of the very meagre amount of pure material. The H₂O content was therefore calculated on the basis of 10 $H₂O$ per formula unit (pfu) from the crystalstructure analysis. Results are given in Table [1](#page-2-1). The high analytical totals after the addition of the calculated H_2O content (114.83) are caused by partial dehydration of the sample (about half of the total water content) under the vacuum conditions of the instrument chamber and due to visible beam damage; dehydration is refected by abundant dehydration fractures in the crystal grains. We also tried to vary measurement conditions, using a beam current of up to 1 nA and a beam diameter of 20 μm, however, with the same results. Finally, we note that the observation of variable transparency of the crystals may indicate possible prior partial dehydration of those crystals that are translucent or not transparent. The material in Vienna was quantitatively analysed with a JEOL "Hyperprobe" JXA 8530F feld-emission gun electron probe micro-analyser (EPMA),

All values are quoted in wt%

Total 116.77

SD standard deviation

^aContent of H₂O was calculated on the basis of ideal composition (10 H₂O pfu) derived from the crystal-structure determination

^bDetection limits (ppm) are: Zn 360, S 100; this may imply that the values for Zn and S are spurious. Mg was measured, but was clearly below the detection limit

Table 2 Powder X-ray difraction data (*d* in Å) for prachařite

$I_{\rm obs}$	d_{obs}	I_{obs} d_{obs}		$d_{\rm calc}$	I_{calc}	hkl	$I_{\rm obs}$ $\,$ $d_{\rm obs}$	$d_{\rm calc}$	$I_{\rm calc}$	hkl
	data 1		data 2				data 2			
3	12.089	16	12.15	12.0850	7	100		1.9031	$\mathbf{1}$	5 1 5
100	9.894	100	9.96	9.8948	100	002		1.9007	$\mathbf{1}$	336
$\sqrt{5}$	7.655	21	7.68	7.6560	13	0 1 2		1.8807	$\mathbf{1}$	$2\;0\;10$
$\overline{4}$	6.980	17	6.92	6.9773	9	$1\ 1\ 0$		1.8569	2	253
				6.5803	2	111	13 1.8479	1.8385	6	319
$\pmb{8}$	6.045	33	6.05	6.0425	17	$2\;0\;0$		1.8350	$\mathbf{1}$	161
6	5.703	24	5.73	5.7022	12	112		1.8131	$\mathbf{1}$	156
10	5.156	30	5.16	5.1570	15	202		1.7385	$\mathbf{1}$	525
11	4.946	10	4.96	4.9474	8	004		1.7264	2	530
$\mathbf{1}$	4.794	5	4.77	4.7934	3	113		1.7213	$\mathbf{1}$	2210
				4.5786	-1	014	1.7170 30	1.7199	6	351
4	4.451	15	4.45	4.4507	7	2 1 1		1.7041	$\mathbf{1}$	3 1 10
$\mathbf{1}$	3.443	\overline{c}	3.862	3.8280	$\mathbf{1}$	024		1.7007	3	532
				3.4887	$\mathbf{1}$	220		1.6945	$\mathbf{1}$	338
		3	3.435	3.4426	$\overline{2}$	115		1.6888	2	419
2	3.352			3.3518	5	310		1.6759	$\mathbf{1}$	620
7	3.305	59	3.314	3.3047	16	3 1 1		1.6702	2	353
19	3.297			3.2983	12	$0\;0\;6$	1.6677 16	1.6691	$\mathbf{1}$	526
4	3.175	14	3.180	3.2901	$\overline{4}$	222		1.6554	$\mathbf{1}$	$4\;0\;10$
				3.1819	2	106		1.6524	3	262
				3.1746	7	3 1 2		1.6491	$\mathbf{1}$	$0\;0\;12$
				3.0839	1	223		1.6315	$\mathbf{1}$	5 1 8
				3.0213	$\overline{2}$	400		1.6300	$\mathbf{1}$	534
22	2.988	95	2.994	2.9882	48	3 1 3		1.5978	$\mathbf{1}$	339
				2.9819	\overline{c}	116		1.5969	$\mathbf{1}$	257
2	2.896	22	2.881	2.8951	5	026		1.5955	$\mathbf{1}$	171
$\overline{\mathcal{A}}$	2.890			2.8896	5	402	1.5935 26	1.5909	$\mathbf{1}$	$0\;2\;12$
$\overline{\mathcal{A}}$	2.851			2.8511	6	224		1.5873	$\mathbf{1}$	624
				2.6740	$\mathbf{1}$	126		1.5852	7	3 1 1 1
$\mathbf{1}$	2.637	12	2.633	2.6372	2	410		1.5824	3	535
				2.6202	$\mathbf{1}$	117	1.5512 3	1.5556	$\mathbf{1}$	173
\overline{c}	2.615			2.6141	5	411	8 1.5297	1.5296	$\mathbf{1}$	356
$\overline{2}$	2.579	31	2.570	2.5785	2	404		1.5242	$\mathbf{1}$	528

Table 2 (continued)

Data 1 = Bruker D8 Advance powder diffractometer, only down to $d = 1.9351 \text{ Å}$; data 2 = Rigaku R-Axis Rapid II (Gandolfi-like data). The strongest difraction maxima are shown in bold

employing JEOL and "Probe for EPMA" (Probe Software, Inc.) analysis software (wavelength-dispersive mode, 10 kV, 20 nA, 20 μm defocussed beam diameter, counting time 10 s on peak and 5 s on background positions). The following natural reference materials and X-ray lines were used: wollastonite (Ca*K*α), kermesite (Sb*L*α, S*K*α), scorodite (As*L*α) and synthetic ZnO (Zn*L*α). If the analytical data were processed assuming an ideal formula, totals of ca. 114–117% were obtained, and the Ca, Sb and As values were all too high (especially As). If a loss of a third of all water molecules (i.e. only the hydrogen-bonded ones, see below) is assumed ("unprocessed" sum is 88%), then the total would be 101% and the Ca:Sb:As-ratio would be nearly perfect.

The calculated empirical formulae, based on $O=22$ atoms per formula unit (apfu), are very similar for the two datasets: $(Ca_{0.93}Mg_{0.02})_{\Sigma 0.95}Sb_{2.02}(AsO_3)_{4.00}.10H_2O$ (Prague material) and $Ca_{1.04}Sb_{2.03}(AsO_3)_{3.92}.10H_2O$ (Vienna material). The ideal formula is $CaSb⁵⁺₂(As³⁺O₃)₄·10H₂O$, which we reformulated as $CaSb⁵⁺₂O₂(As³⁺₂O₅)₂·10H₂O$ to emphasise the presence of $As₂O₅$ dimers in the crystal structure (see below). The ideal formula requires CaO 5.87, $Sb₂O₅$ 33.86, $As₂O₃$ 41.41, H₂O 18.86, total 100.00 wt%.

The calculated Gladstone-Dale compatibility index $(1-K_p/$ K_C ; Mandarino [1981](#page-11-15), [2007](#page-11-16)) is 0.065 (fair) using the empirical formula and single-crystal data from Vienna, and 0.071 (fair) using the empirical formula and single-crystal data from Prague. It is 0.067 using the ideal formula and the single-crystal data from Vienna. The following observation might explain the "fair" category: Prachařite contains corner-shared SbO_6 octahedra; Shannon and Fischer ([2016\)](#page-11-17) noticed that corner-shared octahedral network and chain structures such as perovskites, tungsten bronzes and titaniterelated structures showed systematic deviations between observed and calculated polarisabilities (polarisability analysis is a more reliable measure of the compatibility of a mineral's refractive index, composition and crystal structure) and thus were excluded from their regression analysis.

X‑ray crystallography

X‑ray powder difraction

Powder difraction data were obtained using two diferent methods. The frst data were measured in Los Angeles using a Rigaku R-Axis Rapid II curved imaging plate microdifractometer with monochromatised Mo*K*α radiation. A Gandolf-like motion on the φ and ω axes was used to randomise the sample. Observed *d* values and intensities were derived by profle ftting using JADE 2010 software (Materials Data, Inc.). Data (in Å for $M \circ K_{\alpha}$) are given in Table [2](#page-3-0) ("data 2"). Unit-cell parameters refned from the powder data using JADE 2010 with whole pattern ftting are: *a*=13.952(6), *c*=19.918(10) Å, *V*=3358(3) Å3 .

The second data, also given in Table [2](#page-3-0) ("data 1"), were measured using a Bruker D8 Advance powder difractometer equipped with solid-state LynxEye detector and secondary

Table 3 Crystal data, data collection information and refnement details for prachařite

Crystal data:				
Formula	$CaSb^{5+}(As^{3+},O_5)_{2}O_2 \cdot 10.1H_2O$			
Formula weight	957.73			
Space group, Z	$P\overline{3}c1$ (no. 165), 6			
$a, c (\AA)$	13.951(2), 19.899(2)			
$V(\AA^3)$	3354.1(10)			
$F(000)$, ρ_{calc} (g·cm ⁻³)	2708, 2.845			
$m (mm^{-1})$	8.612			
Absorption correction	multi-scan (Otwinowski et al. 2003)			
SHELX transmission factors	0.358 (min.) -0.480 (max.)			
Crystal dimensions (mm)	$0.15 \times 0.15 \times 0.10$			
Data collection:				
Diffractometer	Nonius KappaCCD system			
λ (Mo– <i>Ka</i>) (Å), <i>T</i> (K)	0.71073, 293			
Crystal-detector distance (mm)	39			
Rotation axes, width $(°)$	φ , ω , 1			
Total number of frames	1325			
Collection time per degree (s)	70			
Collection mode, $2q_{max}$ (°)	full sphere, 70 (used: 64.06)			
h, k, l ranges	$-20 \rightarrow 20, -17 \rightarrow 17, -29 \rightarrow 29$			
Total reflections measured	14,742			
Unique reflections	3898 $(R_{\text{int}} 2.71\%)$			
Refinement (on F^2):				
$R1(F)$, w $R2_{\text{all}}(F^2)$	2.33%, 5.75%			
'Observed' reflections	3057 $[F_0 > 4\sigma(F_0)]$			
Extinctinction coefficient	0.00007(2)			
No. of refined parameters	171			
GooF	1.024			
$(\Delta/\sigma)_{max}$	0.001			
$\Delta\rho_{min}$, $\Delta\rho_{max}$ (e/Å ³)	$-0.54, 0.83$			

Unit-cell parameters were refned from 9806 recorded refections Scattering factors for neutral atoms were employed in the refnement

monochromator producing Cu*K*α radiation housed at the Department of Mineralogy and Petrology, National Museum, Prague, Czech Republic. The instrument was operated at 40 kV and 40 mA. In order to minimise the background, the powder sample was placed on the surface of a fat silicon wafer in ethanol suspension. The powder pattern was collected in the Bragg–Brentano geometry in the range 3–70° 2θ, with a step size of 0.01° and a counting time of 30 s per step (the total duration of the experiment was about three days). The positions and intensities of the difraction maxima were determined and refned using the Pearson VII profle-shape function of the ZDS program package (Ondruš [1993](#page-11-18)) and the following unit-cell parameters were refned by the least-squares program of Burnham [\(1962](#page-11-19)): *a*=13.9541(5), *c*=19.795(1) Å, *V*=3338.0(3) Å³.

 $\sin \hat{\lambda}^2$ for prachafite **Table 4** Fractional atomic coordinates and displacement parameters (in \AA^2) for prachařite $\ddot{\tilde{}}$ d dienle cuino Table 4 Fractional atomic

Table 5 Selected bond distances (\AA) and bond angles (\degree) , and calculated bond-valence sums (in valence units, v.u.) for the coordination polyhedra in prachařite

The second H atom bonded to Ow7 was not detectable. At least three possible positions were detected for H atoms bonded to Ow8; none could be refned. Ow11 and H111 are involved in a bifurcated hydrogen bond. For the partially occupied Ow13 site, two possible H sites were found but no refnement was attempted

Single‑crystal X‑ray difraction

An Ewald sphere of intensity data from a fragment of a welldeveloped transparent prachařite crystal was collected in Vienna with a Nonius KappaCCD four-circle single-crystal difractometer (see Table [3](#page-5-0) for details on data collection and refnement). The crystal structure was solved and refned in space group $\overline{P3c1}$ (no. 165) using SHELX-97 (Sheldrick [2008](#page-11-21)) to $R1 = 2.3\%$ and w $R2_{all} = 5.8\%$.

An additional intensity dataset was collected in Prague using a Rigaku SuperNova single-crystal difractometer (microfocus X-ray tube, Mo*K*a X-radiation, Atlas S2 chargecoupled device detector, 293 K, $2\theta_{\text{max}} = 59.08^{\circ}$, completeness 0.98, $R_{\text{int}} = 5.1\%$). The model obtained is basically identical to the originally obtained model (albeit with standard uncertainties higher by a factor of about 1.5 to 2) and will not be discussed here except in case of notable diferences. The unit-cell parameters, *a*=13.9545(4) Å, *c*=19.7896(5) Å, $V = 3337.33(16)$ Å³, are similar except for a slightly decreased length of the *c* axis (a smaller *c*-axis length is also seen in the powder difraction data of the corresponding material). This is tentatively attributed to variable water contents (see discussion below).

Fig. 3 The fundamental building unit in the crystal structure of prachařite: a heteropolyhedral cluster composed of corner-sharing $SbO₆$ octahedra (yellow), $As₂O₅$ dimers (with eclipsed configuration) built of $AsO₃$ groups (orange) and a central $Ca1O_8$ polyhedron (dark blue). Oxygen ligands are shown as small pale blue circles. Drawing done with ATOMS V. 6.3.1 (Dowty [2006\)](#page-11-22)

Fig. 4 The crystal structure of prachařite in a view along the *c*-axis. Ca-O polyhedra are dark blue, SbO_6 octahedra are yellow, AsO_3 groups are orange, fully occupied Ow atoms are pale blue spheres, and partly occupied Ow13 atoms are green spheres. H atoms have been omitted for clarity. The unit cell is outlined. Drawing done with ATOMS V. 6.3.1 (Dowty [2006\)](#page-11-22)

Table [4](#page-6-0) gives fnal atomic coordinates and displacement parameters for prachařite, while Table [5](#page-7-0) provides selected bond lengths and angles, including calculated bond-valence sums. A CIF-fle is available in the Supplement. The asymmetric unit contains two Ca, one Sb, two As, 13 O and 14 H sites, nine of which could be located and refned with the following restraints: $O-H=0.90(2)$ Å; $H...H=1.50(5)$ Å. Seven of the 13 O sites represent water molecules. One water molecule site (Ow13), located in an interlayer space, is a partially occupied split-site (occupancy \sim 0.40); the refined formula is $CaSb⁵⁺₂(As³⁺O₃)₄~10.1H₂O$. The aforementioned second dataset gave a slightly decreased Ow13 occupancy, ~ 0.25, which might explain the equally slightly decreased length of the *c* axis of the fragment used for that dataset.

Prachařite has a layered structure based on two different layers (Figs. [3](#page-7-1), [4](#page-8-0) and [5\)](#page-9-0). The frst one comprises a six-membered ring of corner-sharing $SbO₆$ octahedra $(**Sb-O**>=1.970$ Å, Table [5](#page-7-0)). The eight-coordinated Ca1 atom occupies the centre of the ring $(Cal-O> = 2.474 \text{ Å}\).$ The $Ca1O_8$ polyhedron is a hexagonal dipyramid that shares each of its horizontal edges with one of the six symmetrically equivalent SbO_6 octahedra. Thus, six ligands are O atoms (O6) and two are water molecules (Ow7). Attached to the SbO_6 octahedra by vertices are As O_3 groups $(<\text{As}1-\text{O}> = 1.779 \text{ Å}, <\text{As}2-\text{O} > = 1.781 \text{ Å}.$ The $(As1O_3)^3$ ⁻ and $(As2O_3)^3$ ⁻ groups are corner-linked to form a $(As_2O_5)^{4-}$ diarsenite group in an eclipsed confguration; thus, the structural formula was formulated as $CaSb⁵⁺₂O₂(As³⁺₂O₅)₂$ 10H₂O. The heteropolyhedral Ca-Sb-As-O-cluster represents the fundamental building unit of prachařite (Figs. [3](#page-7-1) and [4\)](#page-8-0).

The second Ca atom, Ca2, is located between two heteropolyhedral Ca1-Sb-As-O layers (Fig. [5](#page-9-0)). The Ca2 atom shows seven-fold coordination $(Ca2-O>=2.405 \text{ Å}\)$; all ligands are water molecules ($\text{Ow10} \times 3$, $\text{Ow9} \times 3$ and Ow8). Also in the interlayer are located two fully occupied, only hydrogen-bonded water molecules, represented by Ow11 and Ow12, and the already mentioned partially occupied water molecule, Ow13 (Figs. [4](#page-8-0) and [5\)](#page-9-0). The latter is close to a symmetrically equivalent position $[0w13...0w13' = 1.08(2)$ \AA] and shows an increased U_{iso} value (~0.13 \AA^2). Hydrogen bonds are more or less weak, with donor…acceptor distances ranging between 2.74 and 3.22 Å (Table [5\)](#page-7-0). The bond-valence sums, calculated using the recently improved parameters of Gagné and Hawthorne [\(2015\)](#page-11-23), are in very good agreement with the ideal valencies (Table [5\)](#page-7-0).

Raman spectroscopy

Raman spectra (a representative one is shown in Fig. [6\)](#page-9-1) were collected on a crystal lying on its (0001) face in the range $4500-10$ cm⁻¹ using a DXR dispersive Raman Spectrometer (Thermo Scientifc) attached to an Olympus microscope. The Raman signal was excited by the green emission (532 nm) of a diode-pumped solid-state laser and analysed by a charge-coupled device detector. The experimental parameters were: $50 \times$ objective (numerical aperture 0.50), 10 s exposure time, 100 exposures and 1 mW laser power. With a 400 lines/mm difraction grating in the optical pathway, the spectral resolution was in the range $4.4-9.7$ cm⁻¹. The spectra were repeatedly acquired from diferent crystals in order to obtain a representative spectrum with the best signal-to-noise ratio. A possible thermal damage of the measured spots was excluded by visual inspection of each spot after the measurement, by observation of possible decay of spectral features at the start of excitation and checking for thermal downshift of Raman lines. The instrument was set up by a softwarecontrolled calibration procedure using multiple neon emission lines (wavelength calibration), multiple polystyrene Raman bands (laser wavelength calibration) and standardised white-light sources (intensity calibration). Spectral

Fig. 5 The crystal structure of prachařite in a view along the *a*-axis, showing the interlayer content and the eclipsed configuration of the $As₂O₅$ dimer. Legend as in Fig. [4](#page-8-0). Drawing done with ATOMS V. 6.3.1 (Dowty [2006\)](#page-11-22)

manipulations were performed using the Omnic 9 software (Thermo Scientific). The lateral resolution was $\sim 0.7 \mu m$ in confocal mode.

The main bands observed are $(in cm^{-1}, strong bands under$ lined): 3620, 3470, 3383, 3288, 1655, 851, 840, 830, 792, 541, 458, 358, 297, 273, 253, 233, 185, 129, 116, 63 and 24. The dominant bands in the $600-200$ cm⁻¹ region can be assigned to δ As–O–As and AsO₂ vibrations of the $(As_2O_5)^{4-}$ diarsenite group (see above description of crystal structure). Weaker bands in the range $900-750$ cm⁻¹ may be related to stretching (symmetric and antisymmetric) vibrations of As–O bonds, as well as to Sb–O vibrations of the $SbO₆$ octahedron (Vandenborre et al. [1980;](#page-11-24) Baran and Botto [1981;](#page-11-25) Devi and Vidyasagar [1998](#page-11-26); Castro et al. [2009;](#page-11-27) Bahfenne and Frost [2010](#page-11-28); Glamazda et al. [2017\)](#page-11-29). The bands in the $200-10$ cm⁻¹ region are caused by lattice modes. The presence of water is documented by a broad

Fig. 6 Raman spectrum of prachařite (split at 2000 cm–1)

OH-stretching band running from 3800 to 3100 cm^{-1} with recognisable components at 3620, 3470, 3383 and 3288 cm⁻¹; at least four distinct components in this area indicate several structurally non-equivalent water molecules, in agreement with the crystal-structure determination. Following the correlation curve of Libowitzky ([1999\)](#page-11-30), the band component at 3288 cm^{-1} probably corresponds to the hydrogen bond Ow7–H71···Ow13 [2.741(11) Å] (Table [5\)](#page-7-0), while the 3470 and 3383 cm⁻¹ components would correspond to at least fve diferent H bonds with O···O distances between 2.80 and 2.89 Å. The band component at 3620 cm^{-1} is probably related to the very weak hydrogen bond Ow11–H111···O5. A weak band at 1655 cm⁻¹ is assigned to the ν_2 (δ) bending vibrations of water molecules.

Discussion

Relation to other mineral species

No close relationship with other minerals is apparent. Specifcally, there is neither a known mineral species nor a known synthetic compound that both contains Sb^{5+} and As^{3+} . Interestingly, whitecapsite, $H_{16}Fe_5^{2+}Fe_{14}^{3+}Sb_6^{3+}(AsO_4)_{18}O_{16}\cdot 120H_2O$ (Pekov et al. [2014](#page-11-31)) contains both Sb^{3+} and As^{5+} ; its structure is based on a complex heteropolyhedral $[(\Box, Fe^{2+})_6Fe^{3+} _7Sb_3$ $O_8(AsO_4)_9(H_2O)_{30}$ (\square = vacancy) cluster. Smamite, triclinic $Ca_2Sb(OH)_4[H(AsO_4)_2]\cdot 6H_2O$ (Plášil et al. [2020](#page-11-32)), also contains both Sb^{3+} and As^{5+} . Its atomic arrangement is based upon ${Ca_2(H_2O)_6Sb(OH)_4[H(AsO_4)_2]}$ infinite chains consisting of edge-sharing dimers of $Ca(H₂O)₃O₂(OH)₂$ polyhedra which share edges with $Sb(OH)_4O_2$ octahedra. A recently described mineral containing both Sb^{3+} and As^{3+} is lepageite, $Mn^{2+}{}_{3}$ (Fe³⁺₇Fe²⁺₄)O₃[Sb³⁺₅As³⁺₈O₃₄] (Pieczka et al. [2019](#page-11-33)), which was found in the Szklary pegmatite (Lower Silesia, Poland). Seven other minerals are presently known to contain diarsenite groups (listed in chronological order of published crystal-structure determination): paulmooreite, $Pb_2As_2O_5$; gebhardite, $Pb_8(As_2O_5)$, OCl_6 ; fetiasite, $(Fe^{3+},Fe^{2+},Ti)_3(As_2O_5)O_2$; vajdakite, $[(\text{Mo}^{6+}\text{O}_2)_2(\text{H}_2\text{O})_2\text{As}^{3+}{}_2\text{O}_5]$:3H₂O; schneiderhöhnite, Fe²⁺Fe³⁺₃As³⁺₅O₁₃; karibibite, Fe³⁺₃(As³⁺O₂)₄(As³⁺₂O₅) (OH); and bianchiniite, $Ba_2(TiV)(As_2O_5)$, OF. Among these, vajdakite is the only other hydrated phase.

We are not aware of the presence of any other isolated SbO_6 -based Sb_6O_{30} rings in either minerals or synthetic compounds. Although rings of edge-sharing $SbO₆$ octahedra exist in rosiaite, $PbSb₂O₆$ (Basso et al. [1996\)](#page-11-34), and isotypic compounds, they share edges with neighbouring rings to form an interrupted gibbsite-like layer.

Stability and occurrence

Prachařite has formed by weathering of a hydrothermal As-Sb-Pb–Zn(-Ag) ore vein in an estimated temperature range of 10–18 °C (based on long-term observations of the underground microclimate during diferent seasons) and under slightly acidic and slightly oxidising conditions. The E_h (redox potential) range must be fairly narrow since $Sb⁵⁺$ and $As³⁺$ must be simultaneously (meta-?)stable. Although thermodynamically, a mineral containing both Sb^{5+} and $As³⁺$ should not be stable, in reality, this is considered very possible because the stability boundaries Sb^{3+}/Sb^{5+} and $As³⁺/As⁵⁺$ are very close and other arsenite minerals exist that should not be thermodynamically stable, e.g. tooeleite, $Fe^{3+}{}_{6}(As^{3+}O_3)_{4}(SO_4)(OH)_4.4H_2O$ (Juraj Majzlan, pers. comm. to U.K.). Furthermore, it was reported that in a relatively reducing sediment impacted by As- and Sb-rich mining waste, most of the As is reduced to As^{3+} , while Sb^{5+} bound to the oxygen phase persists and represents 58% of the total Sb (Fawcett and Jamieson [2011](#page-11-35)). Other researchers made similar observations (Mitsunobu et al. [2006](#page-11-36); Ritchie et al. [2013](#page-11-37); Fawcett et al. [2015](#page-11-38)). Conditions that would favour the crystallisation of prachařite are assumed to occur also in geologically similar deposits such as the Sainte-Marie-aux-Mines ore district, Haut-Rhin department, France (von Eller and Weil [1966;](#page-11-39) Fluck [1968\)](#page-11-40), which is, as mentioned above, the type locality of the chemically similar mineral smamite, $Ca_2Sb(OH)_4[H(AsO_4)_2]\cdot 6H_2O$, the latter also occuring on one of the studied prachařite specimens.

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