**ORIGINAL PAPER**



# **Evaluating the Microstructure Evolution Behaviors of Saturated Sandstone Using NMR Testing Under Uniaxial Short‑Term and Creep Compression**

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# **Abstract**

Understanding the micromechanical mechanism of the rock creep process is of great importance for studying the macroscopic time-dependent behavior of rocks. In this study, the evolution characteristics of the microstructure (cracks and pores) of saturated sandstones under short term and creep uniaxial compression conditions were investigated with the nuclear magnetic resonance (NMR) technique. The samples were frst loaded to diferent stress levels and creep stages and then completely unloaded for NMR testing. Based on the testing results, the macroscopic deformation behavior, moisture migration law, pore size distribution, porosity, and microstructure change of the each sample under the short-term loading or diferent stages of creep were quantitatively analyzed. After that, by introducing a nonlinear elasto-viscoplastic damage creep model (EVP) by Zhao et al. (18:04017129, 2018), the relationships between the macroscopic irreversible strains and microscopic porosity increments were established. Overall, it was observed that: (1) regardless of the stress level, the magnitudes of the axial and lateral critical strains of samples at the onset of the accelerating creep stage are both relatively constant, and the axial strain is almost comparable to that at the peak stress in the short-term test, while the lateral strain is larger than that of the shortterm test. (2) During the mechanical tests, the moisture in the samples migrates from large pores into small pores, and after mechanical tests, the porosities of the samples increase, in which the small pores always account for a larger proportion. (3) Corresponding to the three creep stages, the porosity of the sample increases slightly after the transient stage, increases to a constant value that is largely independent of stress after the steady stage, and increases signifcantly after the creep failure. In particular, compared to the initial porosity of 6.7%, the average porosities of samples taken to the onset of the tertiary stage and creep failure is 7.49% and 8.71%, increasing by 16.7% and 29.8%, respectively. (4) The porosity growth of sandstone during the brittle creep is mainly driven by the microscopic subcritical crack growth along the grain boundaries.

**Keywords** Saturated sandstone · Microstructure · Porosity · Nuclear magnetic resonance (NMR) · Creep test



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# **1 Introduction**

The time-dependent behavior of rock, i.e., creep, is a fundamental and pervasive phenomenon in geotechnical engineering, such as deep tunnels, mining, and slopes, and thus usually has a critical effect on the long-term safety of supporting structures (Chu et al. [2021;](#page-20-0) Guayacán-Carrillo et al. [2018;](#page-21-0) Liu et al. [2020](#page-21-1); Wu et al. [2020](#page-22-0)). Sandstone, commonly encountered in engineering, generally has a large proportion of pores and strong creep behavior, especially in a saturated state (Ma et al. [2019;](#page-21-2) Ngwenya et al. [2001;](#page-21-3) Tang et al. [2018](#page-21-4); Yu et al. [2020](#page-22-1)). Extensive laboratory creep tests on diferent types of saturated sandstones have shown that creep deformation can account for 20–50% of the total deformation (Tsai et al. [2008;](#page-21-5) Yang and Jiang [2010;](#page-22-2) Zhang et al. [2015](#page-22-3)), and in extreme cases can even exceed 80% (Tian et al. [2015](#page-21-6); Zheng et al. [2015\)](#page-22-4). Similar to most soft rocks, the creep behavior of sandstone usually exhibits a typical three-stage characteristic over time, namely, primary or transient creep, secondary or steady-state creep, and tertiary or accelerating creep (Yang et al. [2014a,](#page-22-5) [b](#page-22-6); Yu et al. [2019](#page-22-7)). However, unlike compact rocks, such as mudstone or argillaceous slate, the creep behaviors of sandstone, due to its porous nature, are caused not only by intergranular cracking but also by the collapse of large intergranular pores (Wong and Baud [2012](#page-21-7); Wu et al. [2000\)](#page-21-8). Therefore, for the brittle creep analysis of sandstone, the major challenge is to understand how and to what extent microscopic (grain-scale) time-dependent pore (or crack) growth processes are linked to the observed macroscopic mechanical behaviors (Brantut et al. [2014a](#page-20-1), [b](#page-20-2); Brantut et al. [2013\)](#page-20-3).

Mechanistically, whether a creep process or a general loading process occurs, the macroscopic deformation and failure of rocks mainly result from the coalescence of many microscopic, grain-scale cracks and collapse of voids (Abdallah et al. [2021;](#page-20-4) Cong et al. [2020;](#page-20-5) Fan et al. [2020](#page-20-6); Paterson and Wong [2005](#page-21-9)). Previous theoretical studies on rock brittle creep have demonstrated that the time-dependent behavior of rock under constant stress conditions can be essentially attributed to chemically activated subcritical crack growth processes, which are referred to as stress corrosion reactions (Amitrano and Helmstetter [2006](#page-20-7); Brantut et al. [2014a,](#page-20-1) [b\)](#page-20-2). To explain this physical mechanism, several micromechanical approaches have been proposed to establish the link between subcritical crack growth and the macroscopic time-dependent behavior of rocks (Brantut et al. [2012;](#page-20-8) Costin [1983;](#page-20-9) Li et al. [2020;](#page-21-10) Main [2000](#page-21-11)). For example, based on fracture mechanics, N. Brantut et al. [\(2012](#page-20-8)) proposed a micromechanical model that combines the sliding wing cracking laws of Ashby and Sammis ([1990](#page-20-10)) and the subcritical crack propagation laws of Charles [\(1958](#page-20-11)) to describe the macroscopic brittle creep behavior of rocks. Nevertheless, the real propagation process and spatial distribution of microscopic cracks and pores induced by stress corrosion still need to be quantitatively analyzed via laboratory tests, which is also a prerequisite for micromechanical analysis (Fan et al. [2018;](#page-20-12) Li and Shao [2016](#page-21-12)).

In terms of creep tests on sandstone, most of previous studies have placed emphases on analyzing the macroscopic mechanical behavior of samples under diferent environmental and stress conditions, such as high temperature, fuid pressure, and chemical corrosion (Ma et al. [2019;](#page-21-2) Noël et al. [2019;](#page-21-13) Tang et al. [2018;](#page-21-4) Yang et al. [2018;](#page-22-8) Yang and Hu [2018;](#page-22-9) Yu et al. [2019;](#page-22-7) Zheng et al. [2015](#page-22-4); Zhu et al. [2019](#page-22-10)). Correspondingly, the creep deformation, permeability, failure mode, and other macroscopic characteristics of diferent sandstone samples have been systematically investigated as well (Brantut et al. [2014a](#page-20-1), [b;](#page-20-2) Wong and Baud [2012](#page-21-7)). However, for microstructural analysis of sandstone during creep, by contrast, only a few studies based on acoustic emission (AE) monitoring, optical microscopy, and SEM have been reported previously (Baud and Meredith [1997](#page-20-13); Heap et al. [2009](#page-21-14); Ngwenya et al. [2001](#page-21-3); Yang et al. [2014a,](#page-22-5) [b\)](#page-22-6). For example, through microscopic optical analysis on four types of sandstones after creep failure, Ngwenya et al. ([2001](#page-21-3)) found that macroscopic brittle creep behavior was characterized by dilatant microcracking, which resulted in shear localization along narrower fracture zones. Using AE monitoring and optical microscopy simultaneously, Heap et al. [\(2009,](#page-21-14) [2015](#page-21-15)) investigated the spatial development of sample damage during brittle and ductile creep and the microcrack density following creep deformation. Their results shown that the aggregation and localization of damage increased as creep progresses, with the greatest changes occurring in the tertiary stage. Moreover, compared to the initial undeformed state, approximately 24–37% increases in crack density at the onset of tertiary stage during brittle creep. Similarly, using SEM, Yang et al.  $(2014a, b)$  $(2014a, b)$  $(2014a, b)$  $(2014a, b)$  investigated the crack distributions of red sandstone subjected to creep tests with diferent fuid pressures, and found that with increasing pore pressure, the arrangement of mineral grains becomes denser and the microcracks decrease in number. In addition, combining the stress–strain curve and wave velocity characteristics under uniaxial compression, Cao et al. ([2014\)](#page-20-14) analyzed the microstructure change of sandstone creep and determined the critical stresses for secondary and tertiary creep. However, the growth of microcracks and the evolution of pores at diferent creep stages and creep stresses were not quantitatively analyzed.

Regarding laboratory tests, both optical microscopy and SEM are 2D observation methods, which cannot quantitatively analyze the microstructure evolution of a whole sample due to their observation range limitations (Fan et al. [2018;](#page-20-12) Wu et al. [2000](#page-21-8)). For AE monitoring, although it is a three dimensional testing method that can monitor the entire creep process in real time, it can only locate damage inside the sample through the signals originated from large cracks, but not quantify the size and distribution of microscopic cracks or pores (Weng et al. [2019\)](#page-21-16). In contrast, the low-feld NMR technique, a recently developed method, can measure the microscopic pore and crack features of rocks with obvious considerable advantages in effectiveness, accuracy (approximately 0.1 nm), and speed (Cai et al. [2014](#page-20-15); Shen et al. [2020](#page-21-17)). In particular, through inversion of watercontaining signals of pores or cracks at diferent scales, the pore size distribution and porosity of samples can also be obtained by the NMR technique (Lin et al. [2020\)](#page-21-18). Based on this technique, many microscopic studies associated with short-term mechanical tests of sandstone under diferent conditions have been conducted, which mainly involve static loading, dynamic impact, freezing and thawing, and chemical corrosion (Gao et al. [2017;](#page-20-16) Hao et al. [2020;](#page-21-19) Lyu et al. [2020;](#page-21-20) Shen et al. [2020;](#page-21-17) Wang et al. [2020a,](#page-21-21) [b;](#page-21-22) Weng et al. [2018](#page-21-23), [2021;](#page-21-24) Xue et al., [2018\)](#page-22-11). However, the application of NMR technique to the creep experiments of rock is relatively rare, especially for detecting the damage changes during the creep stage. To date, only Cheng et al. [\(2021a](#page-20-17), [b\)](#page-20-18) have undertaken a preliminary investigation on the microcracking characteristics of sandstones after creep using this technique. However, in their work, the pore characteristics of samples after experiencing diferent stages of creep at diferent stress levels were not investigated and the moisture migration laws within samples during instantaneous loading and creep loading were also not concerned.

Accordingly, in this paper, the microstructure (cracks and pores) characteristics of saturated sandstone samples under short-term loading and long-term creep processes are

systematically investigated by the NMR testing method. In the mechanical tests, the sandstone samples frst were loaded to eight diferent levels of stress, then underwent diferent stages of creep under each constant stress, and fnally were completely unloaded. In the NMR tests, the sandstone samples were measured three times before loading, immediately after unloading, and after being resaturated. Macroscopic creep deformation behaviors (including the creep strain, steady-state creep rate, and critical creep time) under different conditions were systematically studied. The changes in the NMR  $T<sub>2</sub>$  spectrum distribution and porosity of the samples after being subjected to diferent creep stresses and creep stages were quantifed and analyzed. On this basis, the moisture migration law, microscopic crack and pore evolution, and microdamage mechanisms of the sample during short-term loading and creep were investigated in combination with SEM observations. Finally, by introducing a nonlinear elasto-viscoplastic damage creep model proposed by Zhao et al.  $(2018)$  $(2018)$ , the elasto-viscoplastic strains of the sample under each stress level were separated, and creep damage was conducted by considering the porosity.

#### <span id="page-2-0"></span>**2 Experimental Methodology**

#### **2.1 Sample Preparation**

The rock used in this study is red sandstone from a hydropower station slope in Yunnan Province, and was cored from a single block without macroscopic cracks to ensure the reliability of the test results. Figure [1a](#page-3-0) shows the microscopic photographs of the red sandstone sample, Fig. [1](#page-3-0)b shows the grain size distribution, and Fig. [1](#page-3-0)c and d show its SEM microphotographs. Clearly, this red sandstone is a granular cementitious material with medium-sized grains of approximately 0.01–0.3 mm. Moreover, according to X-ray difraction (XRD), its mineral contents are listed in Table [1.](#page-3-1)

In the study, the dimensions of the samples were 50 mm in diameter and 100 mm in height. To reduce the differences between the samples, all the sandstone samples were cored from the same rock mass. Moreover, to further ensure the uniformity of the samples, the density, P-wave velocity and initial porosity of the prepared sandstone samples were first measured prior to performing the mechanical tests. Figure [2](#page-4-0)a shows the lateral and vertical P-wave velocities of the 40 samples employed in this research, which were measured with the ultrasonic pulse apparatus. The average P-wave velocity of all samples is approximately 2400 m/s. Figure [2b](#page-4-0) shows the corresponding  $T_2$  spectra of those samples tested by the NMR technique. In fact, the NMR  $T_2$  spectrum is the water content distribution of different sizes of pores and the area of the curve is the porosity of the sample. Therefore, the average



<span id="page-3-0"></span>**Fig. 1 a** Photographs of Yunnan red sandstone specimen, **b** grain size distribution, **c** and **d** microphotographs taken by SEM, under×50 and  $\times$  500 magnifications, respectively

<span id="page-3-1"></span>**Table 1** Mineral contents and petrophysical parameters of the red sandstone

Mineral contents based on XRD Petrophysical parameters				
Ouartz, $%$	48.12	P-wave velocity, m/s	2400	
Feldspar, %	30.51	Initial porosity, %	6.7	
Calcite, %	7.92	UCS, MPa	48.32	
Hematite, %	3.07	Poisson's ratio	0.35	
Clay, $%$	6.64	Elastic modulus, MPa	8.75	
Others, %	3.74			

initial porosity of this red sandstone is 6.7%. In addition, to further understand the petrophysical properties of this red sandstone, uniaxial compressive strength (UCS) tests were performed on three saturated samples. The average UCS, elastic modulus *E* and Poisson's ratio *ν* are summarized in Table [1](#page-3-1).

#### **2.2 Experimental Scheme and Procedure**

In this paper, two types of tests were carried out: mechanical tests and NMR tests. The mechanical tests include the short-term uniaxial compression test and creep compression test. The short-term loading test refers to the monotonic uniaxial compression test with the loading rate of 0.2 MPa/s (see Fig. [3a](#page-4-1)). The end point of each short-term loading level is equal to the preset creep stress, and once the stress has reached the target value, it is immediately unloaded to zero for NMR tests. Table [2](#page-5-0) shows the detailed testing scheme of the mechanical test, where as expected, the factors of the stress level and creep stage are taken into account. Specifcally, to account for the infuence of the stress level, the uniaxial compression stress  $\sigma_0$  was divided into eight levels (i.e.,  $\sigma_0$ =0.48–0.95UCS). Similarly, to consider the influence of diferent creep stages, for each stress level, the samples underwent creep to the end of the transient stage (S1), onset of the accelerating stage (S2), and fnal failure (S3), where the positions of S1, S2 and S3 are shown in Fig. [3b](#page-4-1). However, when  $\sigma_0$ =0.48UCS and 0.58UCS, the samples



<span id="page-4-0"></span>**Fig. 2 a** Initial P-wave velocity and **b**  $T_2$  distribution of the chosen samples



<span id="page-4-1"></span>**Fig. 3** Illustrations of **a** two kinds of loading tests and **b** diferent creep stages and the meanings of the states of S0–S3

could not reach the accelerating stage due to the lower stress, and even at the steady stage, the creep rates are very small. Therefore, only the transient creep stage was considered for these two stress levels. In addition, when  $\sigma_0 = 0.95$ UCS, the transient creep stage of the samples could not be precisely captured since the total creep durations are quite short. As a consequence, the tests of only the steady and accelerating creep stages are listed in Table [2](#page-5-0) for this stress level.

For the mechanical testing process, short-term tests were frst conducted on samples P1 to P8. After that, since the creep behavior of the sandstone under diferent stress levels is unknown prior to the tests, the infection points associated with the three creep stages (i.e., S1, S2 and S3 in Fig. [3b](#page-4-1)) cannot be easily and accurately determined from the creep curves. Therefore, during the creep testing, we frst carried out creep tests on the samples C1-1, C2-1, C3-3, C4-3, C5-3, C6-3, C7-3 and C8-3 to characterize the entire creep process and behavior of the sandstone under each stress level. Based on the testing results, the strains, strain rates, and creep time at the infection points of S1 and S2 were preliminarily estimated, providing a reference target value for subsequent creep tests. Then, creep tests were sequentially carried out on samples C3-2 to C8-2 and C3-1 to C7-1, corresponding to creep stages to S2 and S1, respectively. However, due to the sample variability, both the axial strain and strain rate of the above samples during creep were monitored in real time, and once the estimated values at S1 and S2 were reached, the samples were slowly unloaded.

The detailed research procedures consists of the following steps: (1) Sample grouping and initial NMR testing. According to the test scheme listed in Table [2,](#page-5-0) the sandstone samples should be divided into 8 groups before testing, with 5 samples in each group, for a total of 40 samples. For each group, four samples were designated for testing, and one was reserved as a backup. Before the mechanical testing, all the

<span id="page-5-0"></span>**Table 2** Testing scheme of the mechanical tests

Load level $\sigma_0$	Short-term test	Creep test			
	Load to S0	Creep to S1	Creep to S2	Creep to S3	
0.48UCS, 22.93 MPa	P1	$C1-1$			
0.58UCS, 28.03 MPa	P <sub>2</sub>	$C2-1$			
0.69UCS, 33.12 MPa	P <sub>3</sub>	$C3-1$	$C3-2$	$C3-3$	
0.74UCS, 35.67 MPa	P <sub>4</sub>	$C4-1$	$C4-2$	$C4-3$	
0.79UCS, 38.22 MPa	P5	$C5-1$	$C5-2$	$C5-3$	
0.84UCS, 40.76 MPa	<b>P6</b>	$C6-1$	$C6-2$	$C6-3$	
0.90UCS, 43.31 MPa	P7	$C7-1$	$C7-2$	$C7-3$	
0.95UCS, 45.86 MPa	P8	۰	$C8-2$	$C8-3$	

UCS is the uniaxial compressive strength, and S0-S[3](#page-4-1) are shown in Fig. 3

<span id="page-5-1"></span>**Fig. 4** Mechanical test photo of the sample

samples were saturated by a vacuum saturator for 48 h and tested by the NMR system to obtain the initial porosity,  $T_2$ spectra and pore volume. (2) Mechanical testing. Short-term tests and creep tests were performed to obtain the short-term and creep behaviors of sandstones subjected to diferent levels of stress. Notably, the samples for each stress level and creep stage were independent. To prevent the evaporation of water inside the sample, the surface of the sample during mechanical tests was always tightly sealed by the waterproof membrane, and its axial and lateral deformations were recorded simultaneously, as seen in Fig. [4](#page-5-1). (3) NMR testing: After the samples were unloaded, NMR tests were carried out immediately to investigate the moisture migration in samples during the loading and creep process. Afterwards, the samples were resaturated for 48 h in vacuum saturator and retested by the NMR system. The porosity,  $T_2$  spectrum and pore volume under varying stress levels and creep stages were studied and compared with those of the initial state.

# **2.3 Experimental Setup**

In this study, the short-term uniaxial compression test and creep uniaxial compression test were conducted on an MTS 815 testing machine, which is capable of applying a stable servo-controlled axial load. The maximum axial load capacity of the testing machine was 2600 kN. The axial and lateral deformation of the samples were measured with linear voltage diferential transducers (LVDTs) with an accuracy of 10−3 mm. During creep tests, the initial loading rate was the same as that of the instantaneous test, which was set to 0.2 MPa/s. After the target stress was reached, the stress was held constant using a servo-control system with a control accuracy of  $\pm$  0.2 kN. In addition, the room temperature was maintained at 25 °C.

NMR tests were performed using a MacroMR12-150H-I system (produced by Suzhou Niumag Analytical Instrument Corporation, Suzhou, China) with a constant magnetic feld strength of 0.28 T and a resonance frequency of 12.534 MHz (Weng et al. [2021\)](#page-21-24). The main components of the system were a controller, a spectrometer subsystem, a RF unit and a gradient unit. The basic principle of NMR technology has been introduced in detail in previous literatures (Li et al. [2018](#page-21-25); Xue et al. [2018](#page-22-11)). Notably, the transverse relaxation time,  $T_2$ , is generally determined by three diferent relaxation types (bulk relaxation, surface relaxation, and difusion relaxation) for a porous material. However, for water-saturated rocks, the bulk and difusion relaxations can be disregarded (Chen et al. [2018](#page-20-19); Liu et al. [2018](#page-21-26)). Therefore, the measured NMR  $T_2$  value is mainly related to surface relaxation, which can be simply expressed as follows (Li et al. [2018;](#page-21-25) Weng et al. [2018](#page-21-23); Zhou et al. [2018a,](#page-22-13) [b\)](#page-22-14):



$$
\frac{1}{T_2} \approx \frac{1}{T_{2S}} = \rho \left(\frac{S}{V}\right)_{\text{Pore}},\tag{1}
$$

where  $T_2$  is the transverse relaxation time (ms);  $T_{2S}$  is the transverse surface relaxation time (ms);  $\rho_2$  is the surface relaxivity (μm/s), which generally equals 0.05 μm/ms for sandstones; and *S/V* is the surface-to-volume ratio of the pore (m−1), which is equal to 2/*r* (*r* is the radius of pores) here because the pores in the rock material are assumed to be cylindrical (Li et al. [2018\)](#page-21-25). Therefore, substituting all the given quantities into Eq.  $(1)$  $(1)$ , the relation between the pore radius and the relaxation time  $T_2$  of the sandstone is rewritten as:  $r = 0.1T_2$  ( $\mu$ m). Clearly,  $T_2$  is proportional to the pore size; the smaller  $T_2$ , the smaller the pore size.

#### **3 Experimental Results**

#### **3.1 Creep Mechanical Behavior**

Figure [5](#page-7-0) shows the stress–strain relationships of the sandstone under diferent creep stress levels, which were obtained from samples C1-1, C2-1, C3-3, C4-3, C5-3, C6-3, C7-3 and C8-3. As shown in Fig. [5](#page-7-0)a, the axial and lateral loading curves under the eight stress levels show good consistency, indicating that the selected samples have good uniformity and that the problem of sample variability is eliminated. Figure [5b](#page-7-0) shows the volume stress–strain curves, along which points *C* and *D* are the microcrack initiation point and the volumetric strain infection point, respectively. In general, the stresses ( $\sigma_{\rm C}$  and  $\sigma_{\rm D}$ ) at these two points are the thresholds for stable and unstable crack growth. For this saturated sandstone,  $\sigma_{\rm C}$  and  $\sigma_{\rm D}$  are approximately equal to 20.8 MPa and 34 MPa, corresponding to 43%UCS and 70%UCS, respectively. The  $\sigma_C$  was calculated with crack volumetric strain (CVS) method (Nicksiar and Martin [2013](#page-21-27)). With regard to rock creep,  $\sigma_C$  is generally considered to be a threshold for judging creep stability or failure, above which the samples will finally collapse due to plastic or viscoplastic deformation (Brantut et al. [2014a,](#page-20-1) [b;](#page-20-2) Heap et al. [2009\)](#page-21-14). For this reason, in some studies (Chandler [2013;](#page-20-20) Deng et al. [2016](#page-20-21); Zhou et al.  $2018a$ , [b\)](#page-22-14),  $\sigma_C$  is also regarded as the long-term strength. However, if the applied creep stress is equal to or slightly greater than this value, the creep failure time of sample is likely to be very long. Therefore, in this paper most of the creep stresses are larger than this value to ensure the distinct three-stage creep behavior of the samples within an acceptable time scale in laboratory.

As shown in Fig. [6](#page-8-0), when the creep stress is greater than or equal to 33.12 MPa (i.e., 69%UCS), the creep failure of samples can be observed, with the maximum and minimum creep time of 9,200 and 72 min. Moreover, the fgure shows that when  $\sigma_0$  = 22.93 MPa and 28.03 MPa, basically only the

<span id="page-6-0"></span>transient creep stage is observed, while for  $\sigma_0 \geq 33.12$  MPa, the samples exhibit a typical three-stage creep characteristic. Therefore, corresponding to the following NMR porosity test, the creep tests at  $\sigma_0 \geq 35.12$  MPa were performed on samples C3-2 to C8-2 and C3-1 to C8-1, where samples C3-1 to C8-1 only creep to the end of the transient creep stage, while samples C3-1 to C8-1 creep to the onset of the accelerating creep stage.

Figure [7](#page-9-0) shows the creep curves of all the samples tested under  $\sigma_0 \geq 35.12$  MPa, in which the curves for creep to the fnal failure as shown in Fig. [6](#page-8-0). As shown in Fig. [7](#page-9-0), for each stress level, the three groups of creep curves associated with the three creep stages are quite close, especially for the axial strain, which demonstrates once again that the samples employed in this research have good consistency and uniformity. From the above creep curves, the axial and lateral strains at the initial short-term loading (S0), the end of the transient stage (S1) and the onset of the accelerating stage  $(S2)$  can be readily obtained, as shown in Fig. [8.](#page-10-0) This figure shows that as the axial stress increases, the axial and lateral strains at S0 and S1 increase at diferent proportions, while the increases in the axial and lateral strains at S3 are quite small and almost negligible. Therefore, the strain diferences between S0 and S1 as well as between S1 and S2, namely, the strains occurring during the transient and steady creep stages, are related to the magnitude of stress. Figure [9](#page-10-1) shows the relationship between the average creep strains and axial stress in these two stages. Clearly, as the stress increases, the axial and lateral transient strains increase slightly, while the steady-state creep strain decreases sharply, with some local fuctuations.

According to the creep curves in Fig. [7,](#page-9-0) all the axial, lateral and volumetric critical strains at S2 for diferent stress levels are plotted in Fig. [10](#page-10-2), and the corresponding average values are summarized in Table [3.](#page-11-0) The magnitudes of the three critical strains are relatively constant, approximately 0.65%, −0.42% and −0.2%, regardless of the stress level. This demonstrates that there is a critical damage threshold at the occurrence of the tertiary creep stage. Similar fndings have also been reported in several previous experimental studies (Heap et al. [2009](#page-21-14); Wang et al. [2020a,](#page-21-21) [b](#page-21-22)). In addition, as can be further observed from Fig. [10](#page-10-2), the magnitudes of the axial strains at S2 are almost comparable to those at the peak stress in the short-term uniaxial compression test, while the lateral and volumetric strains at S2 are slightly larger than those in the short-term test.

In addition to creep deformation, the steady creep rate and two critical creep time, namely, the time-to-onset of the accelerating stage  $t_{s2}$  and time-to-failure  $t_f$ , are of great importance for the prediction of engineering failure, especially in slope analysis. The steady creep rate refers to the strain rate at the secondary creep stage and is calculated from the slopes of the straight sections of creep curves.

<span id="page-7-0"></span>**Fig. 5** Stress–strain curves for the creep tests under diferent stress levels: **a** axial and lateral curves and **b** volume curves. Points *C* and *D* are the microcrack initiation point and the volumetric strain infection point, respectively



Table [3](#page-11-0) summarizes all the axial and lateral steady creep rates for the diferent axial stress levels, and the relationship between them and the axial stress is plotted in Fig. [11.](#page-11-1) From Table [3,](#page-11-0) it is seen that both the axial and lateral steady creep rates are extremely sensitive to the applied stress. Taking the axial strain rate as an example, at  $\sigma_0$  = 33.12 MPa, the average steady creep rate is  $1.6 \times 10^{-7}$ /min, while as the stress increases to 45.86 MPa, steady creep rate increases to



<span id="page-8-0"></span>**Fig. 6** Creep curves under diferent stress levels corresponding to Fig. [5](#page-7-0): **a** and **b** axial strain, **c** and **d** lateral strain, and **e** and **f** volumetric strain; **b**, **d** and **f** are the local magnifcations of **a**, **c** and **e**, respectively

 $1.98 \times 10^{-5}$ /min. This implies that the increment of the stress of a few MPa induces an increment in the strain rate of up to several orders of magnitude. Furthermore, as shown in Fig. [11,](#page-11-1) the relationships of the creep strain rates in the axial and lateral directions with the stress also can be described well by the power law, i.e., axial strain rate  $\dot{\epsilon}_a = 4.09 \times 10^{-25}$  $\sigma_0^{15.21}$  and lateral strain rate  $\dot{\epsilon}_l = 9.87 \times 10^{-28} \sigma_0^{17.08}$ , which are fairly consistent with the fndings from the creep tests

previously performed on other sandstones (Brantut et al., [2013](#page-20-3); Yu et al., [2019\)](#page-22-7). In addition, by comparing the slopes of the two ftted curves in Fig. [11](#page-11-1), it is seen that the lateral creep strain rate is more sensitive to the stress level than the axial strain.

Table [3](#page-11-0) lists the two creep times  $(t_{s2}$  and  $t_f$ ) and their ratio  $(t_{s2}/t_f)$  for the samples in Fig. [6](#page-8-0) that underwent creep to failure, and  $t_{s2}$  and  $t_f$  versus axial stress are



<span id="page-9-0"></span>**Fig. 7** Creep curves for all the samples that undergo creep to the three different stages under different stress levels: **a**  $\sigma_0 = 33.12$  MPa, **b**  $\sigma_0$ =35.67 MPa, **c**  $\sigma_0$ =38.22 MPa, **d**  $\sigma_0$ =40.76 MPa, **e**  $\sigma_0$ =43.31 MPa, and **f**  $\sigma_0$ =45.86 MPa

also plotted in Fig. [12.](#page-11-2) Unlike the steady creep rates, as the axial stress increases,  $t_{s2}$  and  $t_f$  decrease sharply and  $t_{s2}/t_f$  decreases slightly, suggesting that under higher stresses, the creep duration is shorter, and the accelerating creep occurs earlier. Moreover, as shown in Fig. [12,](#page-11-2) the relationships between the creep stress and two creep times can be well ftted with the exponential functions, namely,  $t_{s2} = 3.55 \times 10^9 \exp(-0.39\sigma_0)$  and



<span id="page-10-0"></span>**Fig. 8** Axial and lateral strains during the initial short-term loading (S0), the end of transient creep (S1) and the onset of accelerating creep (S2), where the symbols denote experimental data and the solid lines donate the corresponding average values



<span id="page-10-1"></span>**Fig. 9** Relationship between the axial and lateral transient and steady creep strains and axial stress

 $t_f = 2.23 \times 10^9$ exp(-0.374 $\sigma_0$ ). Similar fitting relationships have been proposed by Tang et al. [\(2018](#page-21-4)).

# **3.2 Pore Evolution Analysis**

As mentioned in [Experimental Methodology,](#page-2-0) the porosity and pore distribution of the samples after the mechanical tests were measured by NMR. However, with respect to pore evolution analysis, a key representative parameter of the rock pore structure, the  $T_2$  cutoff, should be interpreted because it can refect the degree of mobile fuid in porous media (Lyu et al. [2020](#page-21-20)). As specifed in NMR theory, the pores in terms of  $T_2$  less than the  $T_2$  cutoff are generally



<span id="page-10-2"></span>**Fig. 10** Typical stress–strain curves for the uniaxial compression test, and the relative positions of the axial, lateral and volumetric strains at the onset of the accelerating creep stage

regarded as small pores, in which water is bound and difficult to displace; conversely, the pores with  $T_2 > T_2$  cutoff are large pores, in which the water is free and mobile (Gao et al. [2019;](#page-20-22) Liu et al. [2018](#page-21-26); Lu et al. [2020](#page-21-28)). In this study, the  $T_2$  cutoff of the red sandstone measured by centrifugation is 3.9 ms (i.e., pore radius  $r = 0.39 \,\mu$ m). Hence, when a pore of the red sandstone are larger than 0.39 μm, it is regarded as a large pores, and a pore no larger than 0.39 μm is regarded as a small pore.

#### **3.2.1 Moisture Migration in Pores During Loading**

As mentioned in [Experimental Methodology](#page-2-0), to investigate the moisture migration law during mechanical loading, samples were tested by the NMR technique immediately after unloading. Here, we take the results of samples C5-1 and C5-3 as an example, which were subjected to short-term loading and creep loading at  $\sigma_0$ =38.22 MPa, respectively. Sample C5-3 underwent creep to the onset of the accelerating stage. Figure  $13a$  and b show the NMR  $T_2$  spectra of the two samples, where the Y-axis is the porosity component, corresponding to the signal intensity or amplitude used in the literature.

Figure [13](#page-11-3) shows that the main changes in the porosity of the samples with diferent mechanical loadings are concentrated in the first two peaks. The initial  $T_2$  spectra are somewhat diferent from those tested immediately after unloading. Whether a short-term test or a creep test was performed, during loading, the water content decreases in the large pores and increases in the small pores. However, the total water content in the sample during mechanical tests is almost constant since its surface was always tightly sealed <span id="page-11-0"></span>**Table 3** Average strain at the onset of the accelerating creep stage, average steady creep rate and two creep times



 $0.14$ 



<span id="page-11-1"></span>**Fig. 11** Relationships between the axial and lateral steady creep rates and creep stress







<span id="page-11-3"></span>**Fig. 13** NMR  $T_2$  distributions of two samples subjected to **a** shortterm loading and **b** creep loading, at  $\sigma_0 = 38.22$  MPa

test, the change in porosity between the small pores and

<span id="page-11-2"></span>**Fig. 12** Time to the secondary and tertiary creep stages under diferent constant creep stresses

large pores is approximately 15.4%, which is much larger than the 4.2% observed in the short-term test. This indicates that the moisture migration in the creep test is more obvious than that in short-term loading. In addition, as observed in Fig. [11](#page-11-1), after samples were resaturated, the areas of the *T*2 spectrum were enlarged compared with the initial state,

by the waterproof membrane. Therefore, the water in the samples migrated progressively from the large pores into the smaller pores during loading. In particular, for the creep and the increase in porosity during the creep test was also much larger than that during the short-term test. We further discuss this fnding below.

#### **3.2.2 Pore Evolution Under Diferent Creep Stresses**

As previously mentioned, all samples after unloading were resaturated and retested by the NMR technique. According to the testing scheme, the  $T_2$  spectra of all the samples under diferent stress levels and testing states are illustrated in Fig. [14](#page-12-0). Here, the testing states are the initial state, after short-term loading (S0), after transient creep (S1), after transient and steady creep (S2), and after creep failure (S3). Moreover, each  $T_2$  spectrum curve shown above is obtained from a single sample, which is in accordance with the creep curves in Fig. [7.](#page-9-0) However, as previously clarifed, the heterogeneity between samples is small and thus does not impede the comparative analysis of the  $T<sub>2</sub>$  curves under different stress levels and diferent test states. In addition, during brittle failure, localized shear bands and macroscopic cracks usually appear in a sample, making it fracture into several parts. Therefore, in this study, NMR tests of the samples

after failure were performed on the relatively intact parts of the damaged samples.

Figure [14](#page-12-0) shows that the overall shapes of the  $T_2$  curves after the mechanical tests are essentially unchanged, with four peaks, similar to the initial unformed state. Based on the relationship between  $T_2$  and the pore radius (see Eq. [\(1](#page-6-0))), the pore radii for the four peaks approximately correspond to the intervals of  $1-500$  nm,  $0.5-10.2$  µm,  $10.2-22$  µm and 22–95 µm. In terms of each testing state, it is seen from Fig.  $14a$  that the  $T_2$  peak areas after short-term loading are all enlarged with increasing stress, whose main changes are concentrated near the frst two peaks. Here, for clarity, the differences in the  $T_2$  distribution between the initial state and after short-term loading for each stress level are shown in Fig. [15](#page-13-0). The diference in the porosity components within the range of smaller pores  $(T_2 < T_2$  cutoff) is apparently greater than that within the range of large pores ( $T_2 > T_2$  cutoff), indicating that a larger number of small pores form during the short-term loading. Moreover, with the increase of the stress level, the increasing rate of total porosity increases sharply, indicating that the damage degree becomes more serious under the high loading ratio. The similar fndings have also been observed by Cheng et al. ([2021a](#page-20-17), [b](#page-20-18)).



<span id="page-12-0"></span>**Fig. 14** *T*2 relaxation distribution of samples under diferent axial stresses and at diferent testing states **a** after short-term loading, **b** after transient creep, **c** after transient and steady creep, and (d) after short-term and creep failure



<span id="page-13-0"></span>**Fig. 15** Differences in the  $T_2$  distributions with different stress levels between the initial state and that after short-term loading

Second, Fig. [14b](#page-12-0) and c show that the peak areas of the  $T_2$ distribution after transient creep also gradually enlarge with increasing stress, while after transient and steady creep, the  $T_2$  curves almost overlap with each other. Furthermore, as observed from Fig. [14](#page-12-0)d, all the peak areas of the  $T_2$  curves after failure are substantially enlarged, especially after creep failure, and are much larger than those after short-term failure. However, under diferent stresses, the discrepancies between the curves after creep failure are small, especially at  $T_2$ >40 ms. By comparing the curve of the short-term peak load (i.e.,  $\sigma_0$  = 48.32 MPa) with that after short-term failure, it can be concluded that during the post peak damage stage, the proportion of small pores increases sharply, while the large pores, despite the increase, remain essentially unchanged at  $T_2 > 20$  ms. In addition, Fig. [14](#page-12-0)d also presents the average pore size distribution for the samples taken to the onset of accelerating creep stage. Compared to the peak load, the two spectra are very close, especially

within the range of small pores. This indicates that the two critical states have a similar pore distribution. This is also basically in agreement with the results refected in the creep strain above. However, this fndings is not surprising since both states represent a point that is well into the expansion and cracking phase but prior to crack coalescence and damage localization (Brantut et al. [2013](#page-20-3); Heap et al. [2009](#page-21-14); Wu et al. [2000](#page-21-8)).

Corresponding to Fig. [14](#page-12-0), the porosity of a sample under each stress level and testing state is obtained by accumulating the porosity components on the  $T<sub>2</sub>$  spectra. Table [4](#page-13-1) lists the obtained porosity  $\phi$  and the corresponding porosity diference Δ*ϕ*. Figure [16](#page-13-2) shows the relationship between the porosity and axial stress for the samples after diferent stages of loading. Clearly, the porosities for samples S0 and S1 increase gradually with increasing axial stress. For a certain stress level, the differences in porosity (i.e.,  $\phi_{S1}$ - $\phi_{S0}$ ) between the two testing states are fairly small, especially at



<span id="page-13-2"></span>**Fig. 16** Relationships between the porosity and axial stress at the four testing states



 $\phi_{Si}$  Initial porosity, equal to 6.7%,  $\phi_{SO}$  porosity after short-term loading,  $\phi_{S1}$  porosity after transient creep,  $\phi_{S2}$  porosity at the onset of the accelerating creep stage, and  $\phi_{S2}$  porosity after failure

<span id="page-13-1"></span>**Table 4** Porosity and porosity diference for various stages under diferent stress levels

a lower axial stress. This indicates that the damage in the sample from S0 to S1 is not obvious. Meanwhile, in terms of the increasing gradient between every two subsequent stress levels, both curves at  $\sigma_0$  < 33.12 MPa increase slowly, while for  $\sigma_0$  > 33.12 MPa (0.7UCS), remarkable increases can be observed. This may be because unstable cracks between intergranular particles begin to form between grains begin and dislocate as  $\sigma_0$  becomes larger than  $\sigma_{\rm D}$  (Wen et al. [2018](#page-21-29)).

In addition, as observed from Table [4](#page-13-1) and Fig. [16,](#page-13-2) the porosities for the samples after transient and steady creep (S2) are relatively constant and remain nearly unafected by the stress. This again indicates that there is a critical damage state of the sample at the onset of accelerating creep and it occurs at a similar level in all cases. Moreover, the porosities of the samples after creep failure are substantially larger than those at S0, S1 and S2, and with increasing stress, the values tend to decrease (see Fig. [16](#page-13-2)). This demonstrates that the damage of the sample during the tertiary creep stage is more serious at a lower axial stress. The main reason for this result may be that when the stresses are lower, it takes longer to reach the fnal failure state, allowing more microcrack propagation before macroscopic shear bands or fractures form. A similar observation has been noticed by Brantut et al. [\(2014a](#page-20-1), [b\)](#page-20-2), who indicated that a larger total strain arises at the fnal creep failure when the creep stresses are lower.

To further account for the porosity changes in diferent pore sizes, Fig. [17](#page-14-0) shows the relationship of the porosity for the small and large pores with axial stress, where, as noted earlier, the small and large pores are demarcated by the  $T_2$ cutof. Comparing Figs. [16](#page-13-2) and [17](#page-14-0), it can be clearly observed that the changes in the two porosity components  $\phi_s$  and  $\phi_l$ with axial stress are basically the same as those of all the pores. Moreover,  $\phi_s$  at each stress level is greater than  $\phi_l$ .



<span id="page-14-0"></span>**Fig. 17** Relationships between the porosity components ( $\phi_s$  and  $\phi_l$ ) and the axial stress

#### **3.2.3 Pore Evolution Under Diferent Creep Stages**

To further investigate the pore evolution of diferent creep stages at a given stress level, corresponding to Fig. [14,](#page-12-0) Fig. [18](#page-15-0) redraws the  $T_2$  distribution of the samples under four test states of stress from  $\sigma_0$ =33.12–45.86 MPa. For each stress level, the peak area of the  $T_2$  spectrum obviously increases throughout the testing process, which means that the porosity of the samples increases from short-term loading (S0) to long-term creep failure (S3). More specifically, to quantitatively evaluate the  $T<sub>2</sub>$  spectrum changes of the samples after loading for diferent durations, taking  $\sigma_0$ =40.76 MPa as an example (see Fig. [18](#page-15-0)f), the differences in the  $T<sub>2</sub>$  spectra for samples between the initial state and after the four-stage mechanical tests are plotted in Fig. [19](#page-16-0)a. After the four-stage mechanical tests, the curves of the diference in porosity component remain basically above zero, indicating that the porosity components corresponding to different  $T_2$  increase, especially within the range of 0.1–100 ms. However, along the  $T_2$  axis, the above curves show obvious fuctuations, indicating that the increments in the porosity component at each pore size are quite different. By comparing the four test states, it can be seen that the above curves of S0, S1 and S2 show a similar trend but substantially difer from that after creep failure, whose peak area is much larger than those of other three curves.

In this regard, as shown in Fig. [19](#page-16-0)b, the porosity increments for the small and large pores in the sample after the four-stage tests are determined. The porosity increments of the two kinds of pores increase gradually from S0 to S3, and for each testing state, the increment corresponding to the small pores is substantially larger than that of the large pores. This is consistent with the observation obtained by Cheng et al. [\(2021a,](#page-20-17) [b\)](#page-20-18), although the samples in their study only experienced transient creep or partial steady-state creep. Furthermore, the total porosity increment for all the pores increases with the testing states as well, increasing by 0.39%, 0.52%, 0.77% and 2.05%, respectively. Clearly, the diference in the porosity increment between S0 and S1 is not signifcant and is much smaller than that between S1 and S2. However, the porosity increment after creep failure (S3) is much greater than the earlier states. More specifcally, from Table [4](#page-13-1), we can observed that the average total porosities for all the samples taken to S2 and S3 are 7.49% and 8.71%, respectively. Correspondingly, the percentage increases relative to the initial porosity are 11.8% and 29.8%, respectively. Clearly, the porosity increment at S3 is almost 2.5-fold that at S2. Moreover, in the increased porosities, the porosity of the small pores account for a large proportion: 60% for S2 and 62% for S3. Above results indicate that the total porosity increases sharply from the onset of the accelerating stage to fnal creep failure, especially the proportion contributed by the small pores.



<span id="page-15-0"></span>**Fig. 18** NMR *T*2 distribution of the samples at diferent testing states and stress levels, where **a**–**f** correspond to the eight stress levels listed in Table [2](#page-5-0)

In addition, comparing the closest stress level between Cheng et al.  $(2021a, b)$  $(2021a, b)$  $(2021a, b)$  $(2021a, b)$  $(2021a, b)$  and ours (i.e., the creep stress is about 0.8UCS), the maximum porosity increment in their work is about 0.386%, which is signifcantly less than the 0.75% at S2 in our tests, but slightly larger than 0.37% at S1 (see  $\sigma_0$  = 38.22 MPa in Table [4](#page-13-1)). The difference may be due to three main factors: frst, the pore size distribution and the initial porosity of the sandstones in the two studies are diferent. Second, the samples in Cheng's test were unsaturated, so the creep effect would be generally smaller (Deng et al. [2016](#page-20-21); Yu et al. [2019](#page-22-7)). Third, the duration of creep in Cheng's work seems insufficient for samples to undergo creep to the onset of the accelerating stage. Therefore, at the same stress level, we obtain a signifcantly larger porosity increment than theirs.

#### **3.3 Microstructural Observations**

Corresponding to the pore evolution analysis, microstructural observations with SEM are further conducted.



<span id="page-16-0"></span>**Fig. 19** For  $\sigma_0 = 40.76$  MPa, **a** differences in the  $T_2$  distribution between the initial state and after the four-stage tests, and **b** the corresponding porosity increments during the four-stage mechanical tests

Figure [20](#page-17-0) shows the SEM images of the three samples after experiencing creep to S1, S2 and S3, where the microcracks along the grain boundaries are intergranular crack and within the grains are transgranular crack. As noted earlier, the minimum pore size of this sandstone is greater than 5 nm. Therefore, by comparing Fig. [20](#page-17-0) with Fig. [19,](#page-16-0) it can be seen that the porosity of the sample is mainly contributed by the intergranular cracks and micropores. Considering Fig. [20](#page-17-0) a, c and e, we can see that the connection between crystal granules inside the samples becomes looser as the samples creep from S1 to S3. Specifcally, as shown in Fig. [20](#page-17-0)a and b, after transient creep, the intergranular connections in the sample are still tight, similar to the initial unformed state (see Fig. [1\)](#page-3-0), with only some localized intergranular loosening. In contrast, for the sample after transient and steady creep (see Fig. [20c](#page-17-0) and d), the intergranular loosening is more pronounced and small amounts of crystal grains with lower strength, such as feldspar and clay, begin to break. However, after the sample fails (see Fig. [20e](#page-17-0) and f),

widespread loosening of the connections at the intergranular boundaries can be observed, and with the coalescence of intergranular cracks, large numbers of macrocracks appear in the sample. Similar observations were noted by Heap et al. ([2009](#page-21-14)) from optical photomicrographs of DDS sandstone as well. Therefore, it can be concluded that the brittle creep of the sandstone is driven by microscopic subcritical crack growth within granules and along granular boundaries, and that the increase in porosity is mainly due to the latter. This is remarkably diferent from ductile creep in porous rocks, where a large number of large particles are crushed, resulting in the pore space to collapse and small particles to slide into the pore space (Brantut et al. [2014a,](#page-20-1) [b](#page-20-2); Heap et al. [2015](#page-21-15); Wong and Baud [2012](#page-21-7)).

# **4 Discussion**

Within the framework of damage theory, damage is generally considered to be the formation of pore space (including pores and cracks) inside rock samples (Abdallah et al. [2021](#page-20-4); Weng et al. [2018](#page-21-23); Zheng et al. [2016\)](#page-22-15). In classical damage analysis, irreversible strain or irreversible work is often used to evaluate the degree of damage in samples (Amitrano and Helmstetter [2006;](#page-20-7) Baud and Meredith [1997](#page-20-13); Liu and Dai [2018](#page-21-30)). However, with the application of NMR, porosity is also regarded as an efective indicator to evaluate the damage in samples (Li et al. [2019;](#page-21-31) Weng et al. [2019](#page-21-16); Zhou et al. [2018a,](#page-22-13) [b\)](#page-22-14). Therefore, based on the NMR testing results, the relationships between the irreversible strain and porosity increment for the samples after mechanical tests can be established. Meanwhile, the creep damage associated with diferent creep loadings can be quantitatively analyzed and discussed in combination with porosity.

To characterize rock creep deformation, numerous creep damage constitutive models have been proposed in previous works (Cheng et al. [2021a](#page-20-17), [b](#page-20-18); Fahimifar et al. [2015;](#page-20-23) Ma et al. [2012;](#page-21-32) Wang and Cai [2020;](#page-21-33) Wang et al. [2019](#page-21-34); Yang et al. [2014a](#page-22-5), [b\)](#page-22-6). However, in this study a nonlinear elastoviscoplastic creep constitutive model proposed by Zhao et al. ([2018](#page-22-12)) is employed (see Fig. [21\)](#page-17-1), taking into account its advantage in separating out the diferent strain components. In this model, the total strain consists of the following four components: instantaneous elastic strain, instantaneous plastic strain, viscoelastic strain, and viscoplastic strain. To describe the three stages of creep, the Kachanov creep damage law is adopted to account for the weakening of the viscoplastic viscosity coefficient with time, namely,  $\eta_{\text{vn2}}(t) = \eta_0(1-\omega)$ , where  $\omega$  is the creep damage variable and  $\eta_0$  is the initial viscosity coefficient. At the beginning of creep (i.e., at  $t=0$ ),  $\omega=0$ , while at creep failure (at  $t=t_f$ ),  $\omega$  = 1. Furthermore, under an axial stress  $\sigma_0$ , the one-dimensional creep equation of this model is given as follows:

<span id="page-17-0"></span>**Fig. 20** SEM images of the samples after being subjected to diferent creep stages at  $\sigma_0$ =40.76 MPa, **a** and **b** after transient creep, **c** and **d** after transient and steady creep, and **e** and **f** after creep failure. **b**, **d** and **f** are the magnifed images of the rectangles shown in **a**, **c** and **e**





<span id="page-17-1"></span>**Fig. 21** EVP creep model proposed by Zhao et al. [\(2018](#page-22-12))

$$
\varepsilon_{\rm a}(t) = \frac{\sigma_0}{E_{\rm ie}} + \left(\frac{\sigma_0}{kE_{\rm vp}} + \frac{\sigma_0 - \sigma_s}{E_{\rm vp}}\right) \left[1 - \exp\left(-\frac{E_{\rm vp}t}{\eta_{\rm vp1}}\right)\right] + \frac{\sigma_0 - \sigma_s}{E_{\rm ip}},
$$

$$
+ \frac{(\sigma_0 - \sigma_s)t_{\rm f}(\gamma + 1)}{\eta_0\gamma} \left[1 - \left(1 - \frac{t}{t_{\rm f}}\right)^{\gamma/(\gamma + 1)}\right]
$$
(2)

where  $\varepsilon_a$  is the axial strain;  $E_i$  is the instantaneous elastic parameter;  $E_{\text{ip}}$  is the instantaneous plastic parameter;  $E_{\text{ve}}$ and  $\eta_{ve}$  are the viscoelastic parameters;  $E_{vp}$  and  $\eta_{vp1}$  are the viscoplastic parameters;  $\sigma_s$  is the long-term strength;  $t_f$  is the time-to-failure at a constant stress level; *γ* is the material constant related with the damage variable  $\omega$ ; and *k* is the constant factor to account for the simplifed relationship between the viscoelastic parameters and viscoplastic parameters, i.e.,  $E_{ve} = kE_{vp}$  and  $\eta_{ve} = k\eta_{vp1}$ . Detailed descriptions of the above parameters are given in Zhao et al. ([2018](#page-22-12)).

To analyze the whole process of damage in the three stages of creep, as shown in Fig. [22](#page-18-0), the samples that



<span id="page-18-0"></span>**Fig. 22** Comparison between the experimental data and EVP creep model prediction at fve diferent levels of axial stress

underwent creep to the final failure under five stress levels are chosen for analysis in this study. Based on the above creep model, the experimental data are all fitted (see Fig. [22\)](#page-18-0), and the corresponding model parameters are obtained as follows:  $E_{\text{ie}} = 7.75 \text{ GPa}$ ,  $E_{\text{ve}} = 65.33 \text{ GPa}$ ,  $\eta_{ve} = 4.35 \times 10^4$  GPa·min,  $E_{ip} = 27.56$  GPa,  $E_{vp} = 217.76$ GPa,  $\eta_{\text{vp1}} = 14.5 \times 10^4 \text{ GPa·min}, \eta_0 = 1.28 \times 10^4 \text{ GPa·min},$  $\sigma_s$ =20.8 MPa,  $k$ =0.3 and  $\gamma$ =1.23. Figure [22](#page-18-0) shows that the predictions with the EVP model are in good agreement with the experimental data, indicating that the creep model can accurately describe the creep behavior of the samples. According to the ftting results, the calculated elastic, plastic, viscoelastic, and viscoplastic strains at the onset of accelerating creep stage at each stress level are shown in Table [5.](#page-18-1)

As noted above, both porosity and irreversible strain can be used to evaluate the damage of a sample. Therefore, after short-term loading, as shown in Fig. [23](#page-18-2), a notable positive correlation between the porosity increment  $\phi_{\text{SO}}$ - $\phi_{\text{Si}}$ and instantaneous plastic strain can be observed, and with increasing stress, both quantities increase dramatically, indicating that the damage degree increases as well. Meanwhile, with the ftted experimental data, a linear function can be



<span id="page-18-2"></span>**Fig. 23** Relationship between the instantaneous plastic strain and porosity increment resulting from short-term loading

used to describe the above relationship, which is expressed as:

$$
\Delta \phi = 3 \times 10^{-5} + 6.6 \varepsilon_{\text{ip}}, \ R^2 = 0.99, \tag{3}
$$

where  $\Delta \phi = \phi_{\text{SO}} - \phi_{\text{Si}}$ , (%) and  $\varepsilon_{\text{in}}$  is the instantaneous plastic strain  $(10^{-6})$ .

Similarly, Fig. [24](#page-19-0) shows the relationship between the viscoplastic strain and porosity increment  $\phi_{S2}$ - $\phi_{S0}$  at the onset of the accelerating creep stage. However, diferent from plastic strain, the two quantities decrease sharply as the stress increases since the total strain and porosity at S2 in all cases are basically identical, as discussed previously. Likewise, the relationship between the two quantities also shows a good positive linear correlation, namely:

$$
\Delta \phi = 0.0019 + 4.23 \varepsilon_{vp}, \ R^2 = 0.98, \tag{4}
$$

where  $\Delta \phi = \phi_{S2} - \phi_{S0}$ , (%) and  $\varepsilon_{vp}$  is viscoplastic strain,  $(10^{-6})$ .

Furthermore, based on the above fits and the porosity results in Table [5,](#page-18-1) the creep damage variable  $\omega_{s2}$  and the

<span id="page-18-1"></span>**Table 5** Axial strain composition at the onset of the accelerating creep stage

Stress level $\sigma_0 \epsilon_{\rm ie}$ (10 <sup>-6</sup> ) (MPa)		$\varepsilon_{\rm in}$ (10 <sup>-6</sup> )	$\varepsilon_{\rm ve}$ $(10^{-6})$	$\varepsilon_{\rm vp}$ $(1^{-6})$	$\varepsilon_{\text{tot}\_\text{cal}}$ $(10^{-6})$	$\varepsilon_{\text{tot}\_\text{mea}}$ $(10^{-6})$
33.12	4407	158	388	1307	6260	6380
35.67	4637	324	412	1100	6473	6264
38.22	4900	452	441	651	6444	6569
40.76	5226	612	470	366	6674	6488
43.31	5553	771	500	185	7009	6762

 $\varepsilon_{\rm ie}$  instantaneous elastic strain,  $\varepsilon_{\rm ip}$  instantaneous plastic strain,  $\varepsilon_{\rm ve}$  viscoelastic strain,  $\varepsilon_{\rm vp}$  viscoplastic strain,  $\varepsilon$ <sub>tot\_cal</sub> calculated total strain, and  $\varepsilon$ <sub>tot\_mea</sub> measured total strain</sub>



<span id="page-19-0"></span>**Fig. 24** Relationship between the viscoplastic strain and porosity increment at the onset of the accelerating creep stage



<span id="page-19-1"></span>**Fig. 25** Relationship between the creep damage variable  $\omega_{S2}$  and the porosity increment ratio  $R_{S2}$  at the onset of the accelerating creep stage

porosity increment ratio  $R_{S2}$  at the onset of the accelerating creep stage are obtained, and the relationship between them is shown in Fig. [25.](#page-19-1) Here,  $R_{S2} = (\phi_{S2} - \phi_{S0})/(\phi_{S3} - \phi_{S0})$ , which is the increment ratio of the porosity occurring in the primary and secondary creep stages to that occurring in the entire creep process. Clearly, when  $\phi_{S2} = \phi_{S0}$ ,  $R_{S2}$  equals zero, and when  $\phi_{S2} = \phi_{S3}$ ,  $R_{S2}$  equals one. Accordingly,  $R_{S2}$  reflects the creep damage at S2, similar to  $\omega_{s2}$ .  $R_{S2}$  is slightly smaller than  $\omega_{\gamma}$ , which may be due to the fitting error of the EVP creep model and the limitation of the porosity index. Nevertheless, the relationship between the two quantities also can be described by the following linear equation.

$$
R_{S2} = -0.224 + 0.937\omega_{S2}, \ R^2 = 0.96 \tag{5}
$$

In addition,  $\omega_{\rm s2}$  and  $R_{\rm s2}$  decrease with increasing stress, indicating that less creep damage occurs during S0 to S2 at higher stresses. This is intuitive, since the instantaneous damage induced by loading is greater at a higher stress, whereas the total damage at the onset of tertiary creep is consistent among all the cases studied. Generally, it can be concluded from the above analysis that the change in the porosity of the rock samples is directly related to the creep damage process and that the porosity increment is a reliable indicator for evaluating the degree of rock damage.

# **5 Conclusion**

In this study, a series of mechanical and NMR tests were carried out to explore the microcracking and pore evolution of saturated red sandstone during short-term loading and longterm creep under axial compression. Based on the experimental results, the following conclusions can be drawn.

(1) As the stress increases, the axial and lateral transient creep strains increase slightly, while the steady creep strain drops sharply, albeit with some local fuctuations. The magnitudes of the axial and lateral critical strains at the onset of the accelerating creep stage are both relatively constant regardless of the stress level, and the axial critical strain is almost comparable to that at the peak stress in the shortterm test, while the lateral critical strain is larger than that of the short-term test. Both the axial and lateral steady creep rates are extremely sensitive to the applied stress, and an increase of a few megapascals leads to an increase in the strain rate by up to several orders of magnitude. Conversely, as the stress increases, the two critical creep times (i.e., the time-to-onset of accelerating stage and the time-to-failure) decrease sharply, and the relationships between them and stresses can be described by the negative power law and the exponential law, respectively.

(2) During mechanical loading, the moisture migrates from the large pores  $(> 0.39 \mu m)$  into the small pores  $(< 0.39 \,\mathrm{\mu m})$  in the sample, and this phenomenon is more obvious in the creep test than in the short-term test for the same stress level.

(3) After the short-term and creep mechanical tests, the overall shape of the NMR  $T_2$  spectrum of a sample is basically the same as the initial state, but the corresponding  $T_2$ peak area and porosity both increase with the stress level. During the two mechanical tests, the increasing proportion of the small pores within the samples is much greater than that of the large pores. Regarding the three creep stages, the porosity increment caused by the transient stage is very small, while those caused by both the steady stage and

accelerating stage are closely dependent on the stress level, and the higher the stress level is, the smaller the corresponding porosity increments both are. However, regardless of the stress level, the porosity of the sample at the onset of the accelerating stage is always relatively constant, indicating that the onset of the accelerating stage corresponds to a critical damage state of the sample. In addition, compared to the initial porosity of 6.7%, the average porosities of samples taken to the onset of the tertiary stage and creep failure is 7.49% and 8.71%, i.e., the percentages increases are 11.8% and 29.8%, respectively. The average total porosity increment at S3 is almost 2.5-fold that at S2.

(4) SEM analysis shows that the brittle creep of sandstone is mainly driven by microscopic subcritical crack growth within grains and along granular boundaries. The increase in the porosity of the samples is mainly caused by this microcracking.

(5) The creep damage analysis based on porosity shows that increases in the porosity and the irreversible plastic and viscoplastic strains satisfy good linear relationships at all stress levels. The porosity increment is a reliable indicator for evaluating the degree of creep damage.

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**Data Availability** All data generated during the study are available from the corresponding author by request.

#### **Declarations**

**Conflict of interest** The authors declare that they have no conficts of interest.

# **References**

- <span id="page-20-4"></span>Abdallah Y, Sulem J, Bornert M, Ghabezloo S, Stefanou I (2021) Compaction banding in high-porosity carbonate rocks: 1. experimental observations. J Geophys Res 126(1):e2020JB020538. [https://doi.](https://doi.org/10.1029/2020JB020538) [org/10.1029/2020JB020538](https://doi.org/10.1029/2020JB020538)
- <span id="page-20-7"></span>Amitrano D, Helmstetter A (2006) Brittle creep, damage, and time to failure in rocks. J Geophys Res. [https://doi.org/10.1029/2005J](https://doi.org/10.1029/2005JB004252) [B004252](https://doi.org/10.1029/2005JB004252)
- <span id="page-20-10"></span>Ashby MF, Sammis CG (1990) The damage mechanics of brittle solids in compression. Pure Appl Geophys 133(3):12. [https://doi.org/10.](https://doi.org/10.1007/bf00878002) [1007/bf00878002](https://doi.org/10.1007/bf00878002)
- <span id="page-20-13"></span>Baud P, Meredith PG (1997) Damage accumulation during triaxial creep of darley dale sandstone from pore volumometry and acoustic emission. Int J Rock Mech Min Sci 34(3):24.e21-24.e10. [https://doi.org/10.1016/S1365-1609\(97\)00060-9](https://doi.org/10.1016/S1365-1609(97)00060-9)
- <span id="page-20-8"></span>Brantut N, Baud P, Heap MJ, Meredith PG (2012) Micromechanics of brittle creep in rocks. Geophys Res 117(B8):12. [https://doi.org/](https://doi.org/10.1029/2012JB009299) [10.1029/2012JB009299](https://doi.org/10.1029/2012JB009299)
- <span id="page-20-3"></span>Brantut N, Heap MJ, Meredith PG, Baud P (2013) Time-dependent cracking and brittle creep in crustal rocks: a review. J Struct Geol 52:17–43. <https://doi.org/10.1016/j.jsg.2013.03.007>
- <span id="page-20-1"></span>Brantut N, Heap MJ, Baud P, Meredith PG (2014a) Mechanisms of time-dependent deformation in porous limestone. J Geophys Res 119(7):5444–5463.<https://doi.org/10.1002/2014JB011186>
- <span id="page-20-2"></span>Brantut N, Heap MJ, Baud P, Meredith PG (2014b) Rate- and strain-dependent brittle deformation of rocks. J Geophys Res 119(3):1818–1836.<https://doi.org/10.1002/2013JB010448>
- <span id="page-20-15"></span>Cai C, Li G, Huang Z, Shen Z, Tian S, Wei J (2014) Experimental study of the effect of liquid nitrogen cooling on rock pore structure. J Nat Gas Sci Eng 21:507–517. [https://doi.org/10.1016/j.](https://doi.org/10.1016/j.jngse.2014.08.026) [jngse.2014.08.026](https://doi.org/10.1016/j.jngse.2014.08.026)
- <span id="page-20-14"></span>Cao Y, Deng J, Yu B, Tan Q, Ma C (2014) Analysis of sandstone creep and wellbore instability prevention. J Nat Gas Sci Eng 19:237–243. <https://doi.org/10.1016/j.jngse.2014.05.013>
- <span id="page-20-20"></span>Chandler NA (2013) Quantifying long-term strength and rock damage properties from plots of shear strain versus volume strain. Int J Rock Mech Min Sci 59:105–110. [https://doi.org/10.1016/j.](https://doi.org/10.1016/j.ijrmms.2012.12.006) [ijrmms.2012.12.006](https://doi.org/10.1016/j.ijrmms.2012.12.006)
- <span id="page-20-11"></span>Charles RJ (1958) Static fatigue of glass. J Appl Phys 29(11):4
- <span id="page-20-19"></span>Chen M, Dai J, Liu X, Qin M, Pei Y, Wang Z (2018) Diferences in the fuid characteristics between spontaneous imbibition and drainage in tight sandstone cores from nuclear magnetic resonance. Energy Fuels 32(10):10333–10343. [https://doi.org/10.](https://doi.org/10.1021/acs.energyfuels.8b01396) [1021/acs.energyfuels.8b01396](https://doi.org/10.1021/acs.energyfuels.8b01396)
- <span id="page-20-17"></span>Cheng H, Zhang Y, Zhou X (2021a) A nonlinear creep model for rocks considering damage evolution based on the modifed Nishihara model. Int J Geomech. [https://doi.org/10.1061/](https://doi.org/10.1061/(ASCE)GM.1943-5622.0002071) [\(ASCE\)GM.1943-5622.0002071](https://doi.org/10.1061/(ASCE)GM.1943-5622.0002071)
- <span id="page-20-18"></span>Cheng H, Zhou X, Pan X, Berto F (2021b) Damage analysis of sandstone during the creep stage under the diferent levels of uniaxial stress using NMR measurements. Fatigue Fract Eng Mater Struct 44(3):719–732. [https://doi.org/10.1111/fe.13389](https://doi.org/10.1111/ffe.13389)
- <span id="page-20-0"></span>Chu Z, Wu Z, Liu Q, Liu B, Sun J (2021) Analytical solution for lined circular tunnels in deep viscoelastic burgers rock considering the longitudinal discontinuous excavation and sequential installation of liners. J Eng Mech. [https://doi.org/10.1061/](https://doi.org/10.1061/(ASCE)EM.1943-7889.0001912) [\(ASCE\)EM.1943-7889.0001912](https://doi.org/10.1061/(ASCE)EM.1943-7889.0001912)
- <span id="page-20-5"></span>Cong Y, Wang Z, Zheng Y, Zhang L (2020) Efect of unloading stress levels on macro- and microfracture mechanisms in brittle rocks. Int J Geomech 20(6):04020066. [https://doi.org/10.1061/](https://doi.org/10.1061/(ASCE)GM.1943-5622.0001647) [\(ASCE\)GM.1943-5622.0001647](https://doi.org/10.1061/(ASCE)GM.1943-5622.0001647)
- <span id="page-20-9"></span>Costin LS (1983) A microcrack model for the deformation and failure of brittle rock. J Geophys Res 88(B11):9485–9492. [https://](https://doi.org/10.1029/JB088iB11p09485) [doi.org/10.1029/JB088iB11p09485](https://doi.org/10.1029/JB088iB11p09485)
- <span id="page-20-21"></span>Deng HF, Zhou ML, Li JL, Sun XS, Huang YL (2016) Creep degradation mechanism by water-rock interaction in the red-layer soft rock. Arab J Geosci.<https://doi.org/10.1007/s12517-016-2604-6>
- <span id="page-20-23"></span>Fahimifar A, Karami M, Fahimifar A (2015) Modifcations to an elasto-visco-plastic constitutive model for prediction of creep deformation of rock samples. Soils Found 55(6):1364–1371. <https://doi.org/10.1016/j.sandf.2015.10.003>
- <span id="page-20-12"></span>Fan LF, Gao JW, Wu ZJ, Yang SQ, Ma GW (2018) An investigation of thermal efects on micro-properties of granite by X-ray CT technique. Appl Therm Eng 140:505–519. [https://doi.org/10.](https://doi.org/10.1016/j.applthermaleng.2018.05.074) [1016/j.applthermaleng.2018.05.074](https://doi.org/10.1016/j.applthermaleng.2018.05.074)
- <span id="page-20-6"></span>Fan L, Gao J, Du X, Wu Z (2020) Spatial gradient distributions of thermal shock-induced damage to granite. J Rock Mech Geotech Eng 12(5):917–926. [https://doi.org/10.1016/j.jrmge.2020.](https://doi.org/10.1016/j.jrmge.2020.05.004) [05.004](https://doi.org/10.1016/j.jrmge.2020.05.004)
- <span id="page-20-16"></span>Gao F, Wang Q, Deng H, Zhang J, Tian W, Ke B (2017) Coupled effects of chemical environments and freeze–thaw cycles on damage characteristics of red sandstone. Bull Eng Geol Env 76(4):1481–1490.<https://doi.org/10.1007/s10064-016-0908-0>
- <span id="page-20-22"></span>Gao H, Wang C, Cao J, He M, Dou L (2019) Quantitative study on the stress sensitivity of pores in tight sandstone reservoirs of Ordos basin using NMR technique. J Petrol Sci Eng 172:401–410. <https://doi.org/10.1016/j.petrol.2018.09.083>
- <span id="page-21-0"></span>Guayacán-Carrillo L-M, Sulem J, Seyedi DM, Ghabezloo S, Armand G (2018) Size efect on the time-dependent closure of drifts in Callovo-Oxfordian claystone. Int J Geomech 18(10):04018128. [https://doi.org/10.1061/\(ASCE\)GM.1943-5622.0001213](https://doi.org/10.1061/(ASCE)GM.1943-5622.0001213)
- <span id="page-21-19"></span>Hao L, Zuliang Z, Imo-Imo EK, Yong S, Xinrong L, Dongmin Y (2020) Experimental investigation of the permeability and mechanical behaviours of chemically corroded limestone under diferent unloading conditions. Rock Mech Rock Eng 53(4):1587– 1603.<https://doi.org/10.1007/s00603-019-01961-y>
- <span id="page-21-14"></span>Heap MJ, Baud P, Meredith PG, Bell AF, Main IG (2009) Timedependent brittle creep in Darley Dale sandstone. J Geophys Res. <https://doi.org/10.1029/2008JB006212>
- <span id="page-21-15"></span>Heap MJ, Brantut N, Baud P, Meredith PG (2015) Time-dependent compaction band formation in sandstone. J Geophys Res 120(7):4808–4830.<https://doi.org/10.1002/2015JB012022>
- <span id="page-21-12"></span>Li XZ, Shao ZS (2016) Micro–macro modeling of brittle creep and progressive failure subjected to compressive loading in rock. Environ Earth Sci 75(7):583.<https://doi.org/10.1007/s12665-016-5365-3>
- <span id="page-21-25"></span>Li J, Kaunda RB, Zhou K (2018) Experimental investigations on the efects of ambient freeze-thaw cycling on dynamic properties and rock pore structure deterioration of sandstone. Cold Reg Sci Technol 154:133–141. [https://doi.org/10.1016/j.coldregions.2018.](https://doi.org/10.1016/j.coldregions.2018.06.015) [06.015](https://doi.org/10.1016/j.coldregions.2018.06.015)
- <span id="page-21-31"></span>Li S, Huo R, Yoshiaki F, Ren D, Song Z (2019) Effect of acid-temperature-pressure on the damage characteristics of sandstone. Int J Rock Mech Min Sci 122:104079. [https://doi.org/10.1016/j.ijrmms.](https://doi.org/10.1016/j.ijrmms.2019.104079) [2019.104079](https://doi.org/10.1016/j.ijrmms.2019.104079)
- <span id="page-21-10"></span>Li X, Shao Z, Qi C (2020) An analytical micro-macro model of stress drops during brittle creep in rocks. Eng Fract Mech 223:106794. <https://doi.org/10.1016/j.engfracmech.2019.106794>
- <span id="page-21-18"></span>Lin Y, Zhou K, Li J, Ke B, Gao R (2020) Weakening laws of mechanical properties of sandstone under the efect of chemical corrosion. Rock Mech Rock Eng 53(4):1857–1877. [https://doi.org/10.1007/](https://doi.org/10.1007/s00603-019-01998-z) [s00603-019-01998-z](https://doi.org/10.1007/s00603-019-01998-z)
- <span id="page-21-30"></span>Liu Y, Dai F (2018) A damage constitutive model for intermittent jointed rocks under cyclic uniaxial compression. Int J Rock Mech Min Sci 103:289–301. [https://doi.org/10.1016/j.ijrmms.2018.01.](https://doi.org/10.1016/j.ijrmms.2018.01.046) [046](https://doi.org/10.1016/j.ijrmms.2018.01.046)
- <span id="page-21-26"></span>Liu Z, Yang H, Wang W, Cheng W, Xin L (2018) Experimental study on the pore structure fractals and seepage characteristics of a coal sample around a borehole in coal seam water infusion. Transp Porous Media 125(2):289–309. [https://doi.org/10.1007/](https://doi.org/10.1007/s11242-018-1119-x) [s11242-018-1119-x](https://doi.org/10.1007/s11242-018-1119-x)
- <span id="page-21-1"></span>Liu H, Li L, Li S, Yang W (2020) The time-dependent failure mechanism of rocks and associated application in slope engineering: An explanation based on numerical investigation. Math Probl Eng 2020:1680265. <https://doi.org/10.1155/2020/1680265>
- <span id="page-21-28"></span>Lu Y, Wang L, Ge Z, Zhou Z, Deng K, Zuo S (2020) Fracture and pore structure dynamic evolution of coals during hydraulic fracturing. Fuel 259:116272. <https://doi.org/10.1016/j.fuel.2019.116272>
- <span id="page-21-20"></span>Lyu C, Ning Z, Cole DR, Wang Q, Chen M (2020) Experimental investigation on T2 cutofs of tight sandstones: Comparisons between outcrop and reservoir cores. J Petrol Sci Eng 191:107184. [https://](https://doi.org/10.1016/j.petrol.2020.107184) [doi.org/10.1016/j.petrol.2020.107184](https://doi.org/10.1016/j.petrol.2020.107184)
- <span id="page-21-32"></span>Ma L, Liu X, Fang Q, Xu H, Xia H, Li E, Li W (2012) A new elastoviscoplastic damage model combined with the generalized Hoek-Brown failure criterion for bedded rock salt and its application. Rock Mech Rock Eng 46(1):53–66. [https://doi.org/10.1007/](https://doi.org/10.1007/s00603-012-0256-8) [s00603-012-0256-8](https://doi.org/10.1007/s00603-012-0256-8)
- <span id="page-21-2"></span>Ma D, Duan H, Li X, Li Z, Zhou Z, Li T (2019) Efects of seepageinduced erosion on nonlinear hydraulic properties of broken red sandstones. Tunn Undergr Space Technol 91:102993. [https://doi.](https://doi.org/10.1016/j.tust.2019.102993) [org/10.1016/j.tust.2019.102993](https://doi.org/10.1016/j.tust.2019.102993)
- <span id="page-21-11"></span>Main IG (2000) A damage mechanics model for power-law creep and earthquake aftershock and foreshock sequences. Geophys J Int 142(1):11.<https://doi.org/10.1046/j.1365-246x.2000.00136.x>
- <span id="page-21-3"></span>Ngwenya BT, Main IG, Elphick SC, Crawford BR, Smart BGD (2001) A constitutive law for low-temperature creep of water-saturated sandstones. J Geophys Res 106(B10):21811–21826. [https://doi.](https://doi.org/10.1029/2001JB000403) [org/10.1029/2001JB000403](https://doi.org/10.1029/2001JB000403)
- <span id="page-21-27"></span>Nicksiar M, Martin CD (2013) Crack initiation stress in low porosity crystalline and sedimentary rocks. Eng Geol 154:64–76. [https://](https://doi.org/10.1016/j.enggeo.2012.12.007) [doi.org/10.1016/j.enggeo.2012.12.007](https://doi.org/10.1016/j.enggeo.2012.12.007)
- <span id="page-21-13"></span>Noël C, Pimienta L, Violay M (2019) Time-dependent deformations of sandstone during pore fuid pressure oscillations: Implications for natural and induced seismicity. J Geophys Res 124(1):801–821. <https://doi.org/10.1029/2018JB016546>
- <span id="page-21-9"></span>Paterson MS, Wong TF (2005) Experimental rock deformation-the brittle feld, 2nd edn. Springer, Berlin
- <span id="page-21-17"></span>Shen Y, Wang Y, Wei X, Jia H, Yan R (2020) Investigation on mesodebonding process of the sandstone-concrete interface induced by freeze-thaw cycles using NMR technology. Constr Build Mater 252:118962.<https://doi.org/10.1016/j.conbuildmat.2020.118962>
- <span id="page-21-4"></span>Tang SB, Yu CY, Heap MJ, Chen PZ, Ren YG (2018) The infuence of water saturation on the short- and long-term mechanical behavior of red sandstone. Rock Mech Rock Eng 51(9):2669–2687. [https://](https://doi.org/10.1007/s00603-018-1492-3) [doi.org/10.1007/s00603-018-1492-3](https://doi.org/10.1007/s00603-018-1492-3)
- <span id="page-21-6"></span>Tian H, Chen W, Yang D, Gong Z (2015) Experimental and numerical analysis of the time-dependent behaviour of argillaceous red sandstone under high in situ stress. Bull Eng Geol Env 74(2):567–575
- <span id="page-21-5"></span>Tsai LS, Hsieh YM, Weng MC, Huang TH, Jeng FS (2008) Timedependent deformation behaviors of weak sandstones. Int J Rock Mech Min Sci 45(2):144–154. [https://doi.org/10.1016/j.ijrmms.](https://doi.org/10.1016/j.ijrmms.2007.04.008) [2007.04.008](https://doi.org/10.1016/j.ijrmms.2007.04.008)
- <span id="page-21-33"></span>Wang M, Cai M (2020) A grain-based time-to-failure creep model for brittle rocks. Comput Geotech 119:103344. [https://doi.org/10.](https://doi.org/10.1016/j.compgeo.2019.103344) [1016/j.compgeo.2019.103344](https://doi.org/10.1016/j.compgeo.2019.103344)
- <span id="page-21-34"></span>Wang R, Li L, Simon R (2019) A model for describing and predicting the creep strain of rocks from the primary to the tertiary stage. Int J Rock Mech Min Sci 123:104087. [https://doi.org/10.1016/j.](https://doi.org/10.1016/j.ijrmms.2019.104087) [ijrmms.2019.104087](https://doi.org/10.1016/j.ijrmms.2019.104087)
- <span id="page-21-21"></span>Wang C, Liu J, Chen L, Liu J, Wang L (2020a) Mechanical behaviour and damage evolution of Beishan granite considering the transient and time-dependent efects of excavation unloading. Eur J Environ Civil Eng.<https://doi.org/10.1080/19648189.2020.1784292>
- <span id="page-21-22"></span>Wang F, Cao P, Wang Y, Hao R, Meng J, Shang J (2020b) Combined efects of cyclic load and temperature fuctuation on the mechanical behavior of porous sandstones. Eng Geol 266:105466. [https://](https://doi.org/10.1016/j.enggeo.2019.105466) [doi.org/10.1016/j.enggeo.2019.105466](https://doi.org/10.1016/j.enggeo.2019.105466)
- <span id="page-21-29"></span>Wen T, Tang H, Ma J, Wang Y (2018) Evaluation of methods for determining crack initiation stress under compression. Eng Geol 235:81–97.<https://doi.org/10.1016/j.enggeo.2018.01.018>
- <span id="page-21-23"></span>Weng L, Wu Z, Li X (2018) Mesodamage characteristics of rock with a pre-cut opening under combined static–dynamic loads: a nuclear magnetic resonance (NMR) investigation. Rock Mech Rock Eng 51(8):2339–2354.<https://doi.org/10.1007/s00603-018-1483-4>
- <span id="page-21-16"></span>Weng L, Wu Z, Liu Q (2019) Evaluating damage and microcracking behavior of granite using NMR testing under diferent levels of unconfined compression. Int J Geomech 19(1):04018186. [https://](https://doi.org/10.1061/(ASCE)GM.1943-5622.0001335) [doi.org/10.1061/\(ASCE\)GM.1943-5622.0001335](https://doi.org/10.1061/(ASCE)GM.1943-5622.0001335)
- <span id="page-21-24"></span>Weng L, Wu Z, Liu Q, Chu Z, Zhang S (2021) Evolutions of the unfrozen water content of saturated sandstones during freezing process and the freeze-induced damage characteristics. Int J Rock Mech Min Sci 142:104757. [https://doi.org/10.1016/j.ijrmms.2021.](https://doi.org/10.1016/j.ijrmms.2021.104757) [104757](https://doi.org/10.1016/j.ijrmms.2021.104757)
- <span id="page-21-7"></span>Wong T-F, Baud P (2012) The brittle-ductile transition in porous rock: a review. J Struct Geol 44:25–53. [https://doi.org/10.1016/j.jsg.](https://doi.org/10.1016/j.jsg.2012.07.010) [2012.07.010](https://doi.org/10.1016/j.jsg.2012.07.010)
- <span id="page-21-8"></span>Wu XY, Baud P, Wong T (2000) Micromechanics of compressive failure and spatial evolution of anisotropic damage in Darley Dale sandstone. Int J Rock Mech Min Sci 37(1):143-160. [https://doi.](https://doi.org/10.1016/S1365-1609(99)00093-3) [org/10.1016/S1365-1609\(99\)00093-3](https://doi.org/10.1016/S1365-1609(99)00093-3)
- <span id="page-22-0"></span>Wu K, Shao Z, Qin S (2020) An analytical design method for ductile support structures in squeezing tunnels. Archiv Civ Mech Eng 20(3):91. <https://doi.org/10.1007/s43452-020-00096-0>
- <span id="page-22-11"></span>Xue DJ, Zhou HW, Liu YT, Deng LS, Zhang L (2018) Study of drainage and percolation of nitrogen-water fooding in tight coal by NMR imaging. Rock Mech Rock Eng 51(11):3421–3437. [https://](https://doi.org/10.1007/s00603-018-1473-6) [doi.org/10.1007/s00603-018-1473-6](https://doi.org/10.1007/s00603-018-1473-6)
- <span id="page-22-9"></span>Yang SQ, Hu B (2018) Creep and long-term permeability of a red sandstone subjected to cyclic loading after thermal treatments. Rock Mech Rock Eng 51(10):2981–3004. [https://doi.org/10.1007/](https://doi.org/10.1007/s00603-018-1528-8) [s00603-018-1528-8](https://doi.org/10.1007/s00603-018-1528-8)
- <span id="page-22-2"></span>Yang SQ, Jiang Y (2010) Triaxial mechanical creep behavior of sandstone. Mining Sci Technol 20(3):339–349. [https://doi.org/10.](https://doi.org/10.1016/s1674-5264(09)60206-4) [1016/s1674-5264\(09\)60206-4](https://doi.org/10.1016/s1674-5264(09)60206-4)
- <span id="page-22-5"></span>Yang SQ, Jing HW, Cheng L (2014a) Infuences of pore pressure on short-term and creep mechanical behavior of red sandstone. Eng Geol 179(10):10–23. [https://doi.org/10.1016/j.enggeo.2014.06.](https://doi.org/10.1016/j.enggeo.2014.06.016) [016](https://doi.org/10.1016/j.enggeo.2014.06.016)
- <span id="page-22-6"></span>Yang W, Zhang Q, Li S, Wang S (2014b) Time-dependent behavior of diabase and a nonlinear creep model. Rock Mech Rock Eng 47(4):1211–1224. <https://doi.org/10.1007/s00603-013-0478-4>
- <span id="page-22-8"></span>Yang S, Hu B, Ranjith P, Xu P (2018) Multi-step loading creep behavior of red sandstone after thermal treatments and a creep damage model. Energies 11(1):10. <https://doi.org/10.3390/en11010212>
- <span id="page-22-7"></span>Yu C, Tang S, Tang CA, Duan D, Zhang Y, Ma T (2019) The effect of water on the creep behavior of red sandstone. Eng Geol 253:64– 74. <https://doi.org/10.1016/j.enggeo.2019.03.016>
- <span id="page-22-1"></span>Yu J, Yao W, Duan K, Liu X, Zhu Y (2020) Experimental study and discrete element method modeling of compression and permeability behaviors of weakly anisotropic sandstones. Int J Rock Mech Min Sci 134:104437. [https://doi.org/10.1016/j.ijrmms.2020.](https://doi.org/10.1016/j.ijrmms.2020.104437) [104437](https://doi.org/10.1016/j.ijrmms.2020.104437)
- <span id="page-22-3"></span>Zhang Y, Shao J, Xu W, Jia Y, Zhao H (2015) Creep behaviour and permeability evolution of cataclastic sandstone in triaxial rheological

tests. Eur J Environ Civ Eng 19(4):496–519. [https://doi.org/10.](https://doi.org/10.1080/19648189.2014.960103) [1080/19648189.2014.960103](https://doi.org/10.1080/19648189.2014.960103)

- <span id="page-22-12"></span>Zhao Y, Zhang L, Wang W, Wan W, Ma W (2018) Separation of elastoviscoplastic strains of rock and a nonlinear creep model. Int J Geomech 18(1):04017129. [https://doi.org/10.1061/\(ASCE\)GM.](https://doi.org/10.1061/(ASCE)GM.1943-5622.0001033) [1943-5622.0001033](https://doi.org/10.1061/(ASCE)GM.1943-5622.0001033)
- <span id="page-22-4"></span>Zheng H, Feng X, Hao X (2015) A creep model for weakly consolidated porous sandstone including volumetric creep. Int J Rock Mech Min Sci 78:99–107. [https://doi.org/10.1016/j.ijrmms.2015.](https://doi.org/10.1016/j.ijrmms.2015.04.021) [04.021](https://doi.org/10.1016/j.ijrmms.2015.04.021)
- <span id="page-22-15"></span>Zheng Z, Sun W, Fish J (2016) Micropolar efect on the cataclastic fow and brittle-ductile transition in high-porosity rocks. J Geophys Res 121(3):1425–1440.<https://doi.org/10.1002/2015JB012179>
- <span id="page-22-13"></span>Zhou K, Liu T, Hu Z (2018a) Exploration of damage evolution in marble due to lateral unloading using nuclear magnetic resonance. Eng Geol 244:75–85. [https://doi.org/10.1016/j.enggeo.2018.08.](https://doi.org/10.1016/j.enggeo.2018.08.001) [001](https://doi.org/10.1016/j.enggeo.2018.08.001)
- <span id="page-22-14"></span>Zhou X, Huang X, Berto F (2018b) A three-dimensional long-term strength criterion of rocks based on micromechanical method. Theoret Appl Fract Mech 97:409–418. [https://doi.org/10.1016/j.](https://doi.org/10.1016/j.tafmec.2017.07.003) [tafmec.2017.07.003](https://doi.org/10.1016/j.tafmec.2017.07.003)
- <span id="page-22-10"></span>Zhu W, Li S, Li S, Niu L (2019) Influence of dynamic disturbance on the creep of sandstone: An experimental study. Rock Mech Rock Eng 52(4):1023–1039. [https://doi.org/10.1007/](https://doi.org/10.1007/s00603-018-1642-7) [s00603-018-1642-7](https://doi.org/10.1007/s00603-018-1642-7)

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