

# Evolution of calc-alkaline to alkaline magmatism through Carboniferous convergence to Permian transcurrent tectonics, western Chinese Tianshan

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**Abstract** Continuous magmatic activity occurred in the western Chinese Tianshan, Central Asia, from the Carboniferous to the Permian, i.e. before and after the Late Carboniferous amalgamation of Junggar and the Yili Blocks. Zircon U–Pb LA-ICPMS and Ar–Ar data reveal a coincidence in time between regional wrench faulting and granitoid emplacement. Permian post-collisional granitoids crop out within or at the margins of large-scale dextral strike-slip shear zones, some of them show synkinematic fabrics. The whole rock geochemical features of the Early–Middle Permian granitoids indicate an evolution from high-K calc-alkaline towards alkaline series. In other places of the North Tianshan, alkaline magmatism occurred together with deep marine sedimentation within elongated troughs controlled by wrench faults. Therefore, in contrast with previous interpretations that forwarded continental rift or mantle plume hypotheses, the coexistence of diverse magmatic sources during the same tectonic episode suggests that post-collisional lithosphere-scale transcurrent

shearing tightly controlled the magmatic activity during the transition from convergent margin to intraplate anorogenic processes.

**Keywords** Granitoids · Zircon U–Pb geochronology · Geochemistry · Transcurrent shear zones · Late Palaeozoic · Chinese Tianshan

## Introduction

The tectonic assemblage of the Central Asian Orogenic Belt was suggested to occur during Palaeozoic time by the closure of Paleasian Ocean and multi-phase subduction–accretion of various micro-continents, ancient island arcs and fragments of oceanic island (Coleman 1989; Dobretsov et al. 1995; Jahn et al. 2000, 2004; Jahn 2004; Xiao et al. 2004; Windley et al. 2007).

The Chinese Tianshan Belt which is a part of this orogenic belt, separates the Tarim Basin to the south from the Junggar Basin to the north (Fig. 1). Classically, it is considered that the western Chinese Tianshan (west of Meridian 88°E; Fig. 1) was built by two collisions, during the Early Palaeozoic between Tarim and Central Tianshan blocks, and during the Late Palaeozoic between Central Tianshan and North Tianshan (Windley et al. 1990; Allen et al. 1993; Gao et al. 1998; Chen et al. 1999; Carroll et al. 2001; Zhou et al. 2001), but there is no agreement about the end of convergence. Generally, the convergent orogeny is considered to be terminated before the Early Permian (Wang et al. 1994; Liu et al. 1996; Xiao et al. 2004, 2006; Gao and Klemd 2003; Klemd et al. 2005; Gao et al. 2006; Li et al. 2006a; Charvet et al. 2007), and this phase of amalgamation was followed by regional-scale Late Palaeozoic (Permian) wrench faulting (Allen et al. 1995; Allen and Vincent 1997; Laurent-Charvet

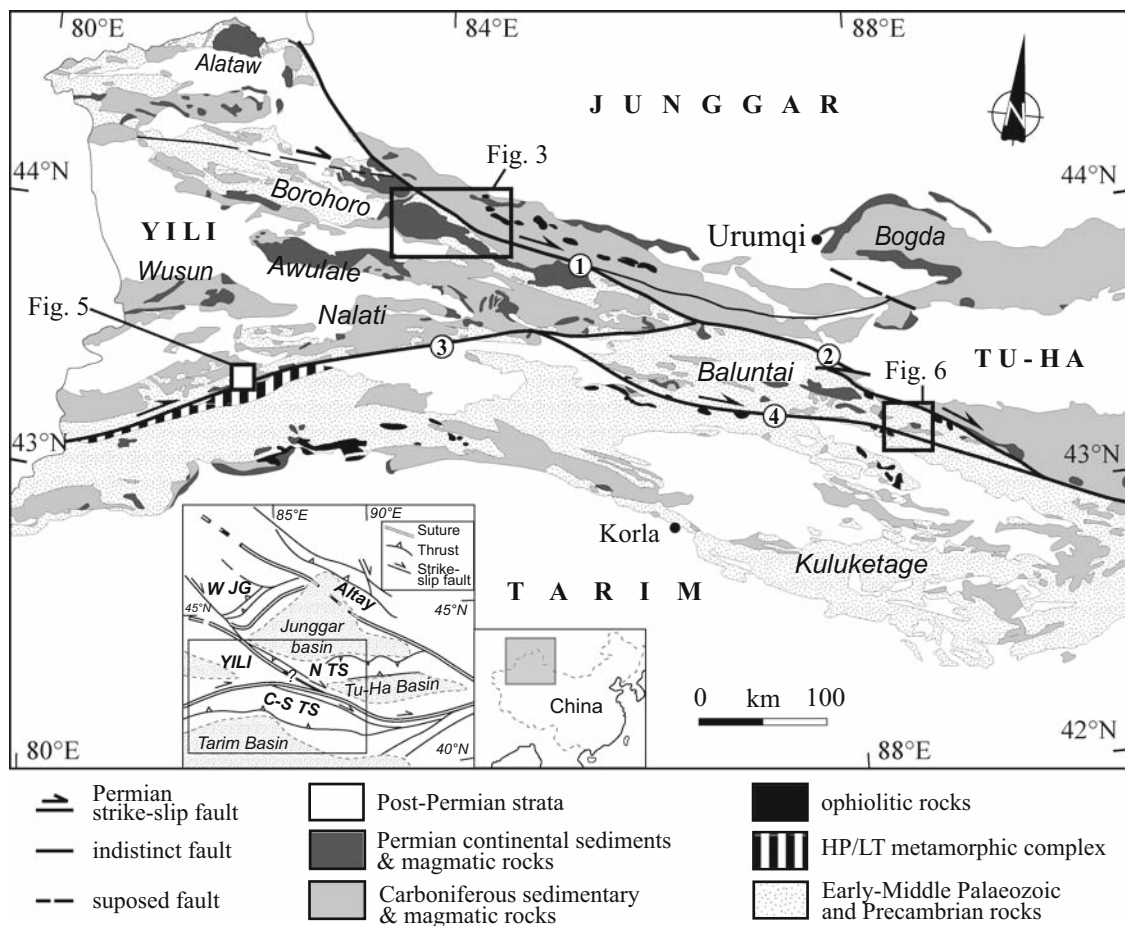
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**Fig. 1** Simplified sketch map of the western Chinese Tianshan Belt (modified from XBGMR 1993; insets after Zhang et al. 1993, Allen and Vincent 1997), showing the occurrences of Carboniferous to Permian igneous rocks and the main wrench faults. Numbers 1–4

correspond to the main faults, 1 North Tianshan Fault (NTF), 2 Main Tianshan Shear Zone (MTSZ), 3 Qingbulak-Nalati Fault (QNF), 4 Sangshuyuanzi Fault (SF), W JG Western Junggar, N TS North Tianshan, C-S TS Central-South Tianshan

et al. 2002, 2003). Some authors argued that continent amalgamation was accomplished in the pre-Carboniferous and that a post-collisional continental rifting occurred during the Carboniferous to Permian (Che et al. 1996; Xia et al. 2004b). Others suggest that the collision in southwestern Tianshan occurred as late as Late Permian-Triassic (Li et al. 2002, 2005; Zhang et al. 2007).

The magmatic affinities and emplacement ages provide keys for understanding the geodynamic evolution of the orogen, but in Chinese Tianshan, the coexistence of Carboniferous to Permian calc-alkaline and alkaline magmatic rocks has inferred contrasting interpretations. It has been widely accepted that the Carboniferous magmatism is subduction-related and/or syncollisional, and the Permian one is post-collisional. However, Carboniferous-Permian volcanic rocks were also interpreted as a result of continental rifting (Che et al. 1996; Xia et al. 2004b); or parts of a Large Igneous Province (Xia et al. 2004a, 2006) although these views are not supported by compelling tectonic or stratigraphic evidences.

Recent studies suggest that the genetic links may exist between Permian wrench faulting and post-collision magmatism, and therefore could be a clue for understanding the evolution of the Chinese Tianshan during this period. This article presents our field observations, new zircon U–Pb LA-ICPMS dating and geochemical data from Permian igneous rocks of the western Chinese Tianshan. Combining with previous data, we discuss the petrogenesis of Permian igneous rocks and suggest that most petrologic features were controlled by the location of wrench faults that allowed local asthenosphere uplift and transition from calc-alkaline to alkaline magmas to occur.

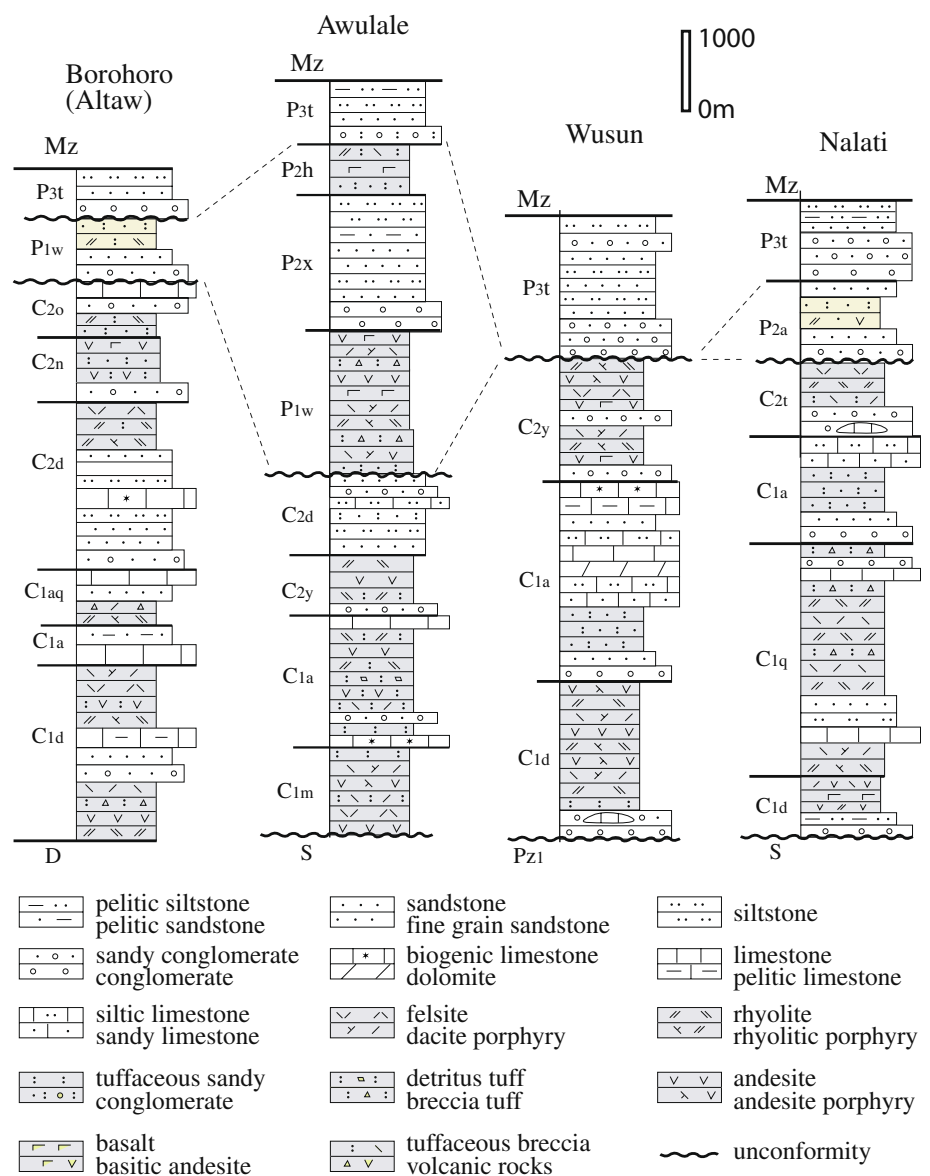
### Geologic setting

The Late Palaeozoic evolution of the Chinese Tianshan is characterized by polyphase deformation and transition from continental active margin to intraplate tectonics. The pre-Carboniferous subduction finally induced a continental

collision between the Yili Block and Tarim in Early Carboniferous (~345 Ma) along the HP/LT metamorphic zone in SW of the Tianshan Belt (Fig. 1; Allen et al. 1993; Gao et al. 1998; Gao and Klemd 2003; Wang et al. 2007c). During the Carboniferous, massive limestone, pelitic limestone and sandstone were deposited (XBGMR 1993). Meanwhile, in the NW part, extensive volcanic rocks (mainly andesite and rhyolite) erupted indicating a long-term magmatic activity (Figs. 1, 2). Between the latest Carboniferous and Early-Middle(?) Permian, in the North Tianshan (Fig. 1), the formation of an ophiolitic mélange including blocks of Late Carboniferous oceanic crustal rocks and turbidites suggests that the amalgamation of continental blocks came to the end (Windley et al. 1990, 2007; Allen et al. 1993; Wang et al. 2006a).

During the Permian, large-scale dextral ductile shearing occurred along the North Tianshan Fault (Allen and Vincent 1997; or Dzhungarian Fault by Zhang et al. 1993) and Main Tianshan Shear Zone (Laurent-Charvet et al. 2003), and along the Qingbulak-Nalati Fault (Windley et al. 1990) and Sangshuyuanzi Fault (Yin and Nie 1996) (faults 1, 2, 3 and 4 in Fig. 1, respectively). They are developed in a LP-HT thermal regime as testified by the widespread development of synkinematic andalusite (Wang et al. 2006a). The age of ductile deformation is constrained at ca. 290–245 Ma by <sup>40</sup>Ar/<sup>39</sup>Ar dating of synkinematic feldspar (Zhou et al. 2001) and biotite (Yin and Nie 1996; Laurent-Charvet et al. 2002, 2003; de Jong et al. 2008, this volume). In Central Asia as a whole, a Permian transcurrent regime followed the earlier N–S-directed convergence (Bazhenov et al. 1999,

**Fig. 2** Simplified stratigraphic columns of Carboniferous–Permian volcanic and sedimentary rocks in northwestern Chinese Tianshan (modified from 1:200,000 geological maps of Xinjiang region). Abbreviations of formations: *C<sub>1d</sub>* lower Carboniferous Dahalajunshan Fm., *C<sub>1m</sub>* Meiluokahe Fm., *C<sub>1q</sub>* Qiergustao Fm., *C<sub>1a</sub>* Aqialehe Fm., *C<sub>2y</sub>* upper Carboniferous Yishijilike Fm., *C<sub>2d</sub>* Dongtujin Fm., *C<sub>2n</sub>* Naogaitu Fm., *C<sub>2o</sub>* Oyiman Fm., *C<sub>2t</sub>* Tuergong Fm., *P<sub>1w</sub>* lower Permian Wulang Fm., *P<sub>2x</sub>* Xiaoshansayi Fm., *P<sub>2h</sub>* Hamist Fm., *P<sub>2a</sub>* Aikendaban Fm., *P<sub>3t</sub>* Tielimuke Fm., *D* Devonian, *S* Silurian, *P<sub>z1</sub>* Early Palaeozoic strata, *Mz* Mesozoic strata



2003; Allen et al. 2006; Van der Voo et al. 2006; Wang et al. 2007b). During this period, a major paleogeographic change occurred in the Chinese Tianshan, and terrestrial sediments associated with sub aerial volcanic rocks deposited unconformably or disconformably over Carboniferous sedimentary and volcanic rocks. In contrast, in restricted areas such as the south of Bogda Mountain (Fig. 1), deep water sediments accumulated in elongated troughs where they overlie collapse breccia (olistoliths) and pillow basalt (Shu et al. 2005).

### A synthesis of previous data

Carboniferous and Permian igneous rocks are widespread in the Chinese Tianshan Belt (Fig. 1). In order to get a unified understanding on the geochemistry and chronology of these magmatic rocks, the following paragraphs summarise the previous data, except for undated or poorly time-constrained rocks that will not be considered. According to their geochemical features, three types of magmatic rocks are defined, namely calc-alkaline (CA), alkaline (A) and transitional series (TR, calc-alkaline rocks with prominently high REE contents) (see Table 1).

#### Carboniferous magmatism

Carboniferous volcanic rocks consist of mafic, intermediate and felsic rocks (XBGMR 1993) that are exposed in the Borohoro, Awulale, Nalati and Wusun mountains (Figs. 1, 2). The Lower Carboniferous Dahalajunshan Formation ( $C_{1d}$ ) is well known for its extensive distribution and remarkable thickness. It is mainly composed of andesite, rhyolite, felsite, tuff breccia and minor basalt, associated with limestone and sandstone. The volcanic rocks show significant depletion in Nb and Ta, moderate depletion of Hf and Zr, and prominent enrichment in Rb and Th (Li et al. 1998; Ma and Wang 2000; Yang et al. 2003; Zhang and Li 2006; A et al. 2006; Guo and Zhu 2006; Li et al. 2006b; Shao et al. 2006). The intermediate to acidic volcanic rocks of the Lower Carboniferous Akeshake Fm. ( $C_{1a}$ ) and Upper Carboniferous Tuergong Fm. ( $C_{2t}$ ) in the Nalati range have similar geochemical compositions (Wang et al. 2007a), they show positive  $\varepsilon_{Nd}(t)$  values (+0.32 to +4.90), and variable  $^{87}Sr/^{86}Sr(i)$  ratios (0.7015–0.7068) (Zhu et al. 2005, 2006a; Guo and Zhu 2006; Qian et al. 2006). Recent SHRIMP zircon U–Pb dating provided consistent Carboniferous ages (363–313 Ma, Zhu et al. 2005, 2006b; Zhai et al. 2006). These volcanic rocks are defined as CA series in Table 1.

Carboniferous adakites were recognized in Alataw, Borohoro and Baluntai areas (Fig. 1) and show high Sr, Eu compositions,  $\varepsilon_{Nd}(t)$  values (+1.5 to +10.0), and low Y,

Yb contents and  $^{37}Sr/^{86}Sr(i)$  ratios (<0.7070) (Wang et al. 2003, 2006b; Zhao et al. 2003b, 2006). The volcanic rocks of the Upper Carboniferous Yishijilike Fm. ( $C_{2y}$ ) in the Awulale range (Figs. 1, 2) display enrichment in incompatible elements and a moderate Ta and Nb negative anomaly (Liu et al. 2006). The Lower Carboniferous andesite and rhyolite of Borohoro and Wusun mountains show high Ti/Y (>500), Ce/Y (>3) ratios, and relatively low total Fe (5.8–7.8 ppm) (Che et al. 1996; Xia et al. 2004b); the Carboniferous basalt, andesite and rhyolite in Baluntai and Bogda areas have lower Ti/Y (<500), Ce/Y (<3), high Total Fe (6.4–11 ppm) (Xia et al. 2004b). These “rift-related” volcanic rocks are characterized with positive  $\varepsilon_{Nd}(t)$  values (+4.1 to +9.7) and consistent  $^{87}Sr/^{86}Sr(i)$  ratios (0.7034–0.7059) suggesting origin from a mantle source and weak crustal contamination (Xia et al. 2004b). These rocks are not directly dated, and therefore are not listed in Table 1.

Carboniferous plutonic rocks in Wusun, Nalati and west of Borohoro mountains (Fig. 1) were formed during 352–308 Ma interval (zircon U–Pb TIMS by Xu et al. 2006 and LA-ICP-MS by Wang et al. 2006a). Geochemistry indicates that they are mainly I-type granites associated with minor S-type granites (Wang et al. 1993; Li et al. 1995), they display enrichment in Rb, Th and depletion in Nb, Ta, Zr and Hf (Xu et al. 2006). In the east of Alataw area, I-type granodiorite and K-granite at ca. 307–290 Ma ( $^{40}Ar/^{39}Ar$ ) display a weak LREE enrichment and strong enrichment in K, Rb; the  $\varepsilon_{Nd}(t)$  values (0 to +7), and  $^{87}Sr/^{86}Sr(i)$  ratios (0.7020–0.7110) (Chen et al. 1994, 2000; Zhou et al. 1994, 1995, 1996) are characteristic of the calc-alkaline series and therefore they are correlated to type CA (Table 1).

#### Permian magmatism

The Permian of the western Chinese Tianshan is characterized by plant-bearing conglomeratic red-beds that unconformably overlie older strata (XBGMR 1993). The relatively small-sized granitoids that are generally referred to as “Late Hercynian plutons” on 1/200,000 maps (e.g. XBGMR 1973) are distributed widely in the study area (Fig. 1). On the basis of recent U–Pb zircon dating, they predominantly belong to Permian with only a few Late Carboniferous rocks. In the southern Tianshan, diorite and porphyritic granite were formed during ca. 298–284 Ma (U–Pb dating on zircon), they are rich in K, Rb, Th, and depleted in Ta, Nb and Zr; alkali feldspar granite was emplaced during 265–260 Ma, and shows enrichment in Nb, Y, LREE, and lower Ti, Sr and Ba contents (Jiang et al. 1999). In Alataw area, the S-type K-granite and monzonite were formed during 298–271 Ma ( $^{40}Ar/^{39}Ar$  and zircon U–Pb La-ICPMS), they have flat REE distribution patterns,

**Table 1** A compilation of geochronological and geochemical data of Carboniferous to Permian igneous rocks in western Chinese Tianshan

Rocks	Localities	Age		Methods	Main geochemical		References	Types
		Ma			(La/Yb) <sub>N</sub>	$\epsilon_{Nd(T)}$		
Andesite	Borohoro	363.2 ± 5.7		Zircon U–Pb SHRIMP	4.9–12.6		Zhai et al. (2006) and Wang et al. (2007a)	CA
Basalt	Nalati	353.7 ± 4.5		Zircon U–Pb SHRIMP	4.1–5.8	+0.3 to +3.1	Zhu et al. (2006a)	CA
Granodiorite	W Borohoro	352 ± 2		Zircon U–Pb TIMS	3.02		Xu et al. (2006)	CA
Granodiorite	Wusun	348 ± 1		Zircon U–Pb TIMS	2.21		Xu et al. (2006)	CA
Trachyte andesite	Nalati	312.8 ± 4.2		Zircon U–Pb SHRIMP	3.5–5.9	+2.7 to +4.9	Zhu et al. (2005) and Guo and Zhu (2006)	CA
Granodiorite	Nalati	308 ± 1		Zircon U–Pb TIMS	0.89		Xu et al. (2006)	CA
Granodiorite	E Alataw	307 ± 3		Ar/Ar plateau	2.2–4.5	+0.1 to +6.9	Chen et al. (1994, 2000)	CA
Porphyritic granite	N Alataw	298.4 ± 5.7		Zircon U–Pb	3.15–5.96	+3.5 to +3.6	Liu et al. (2005)	CA
Diorite	S Tianshan	298		Zircon U–Pb	34	–6.5	Jiang et al. (1999)	A
Biotite granite (B101)	Borohoro	294 ± 7		Zircon U–Pb	7.2		This study	TR
K-granite	S Alataw	292.4 ± 4.9		Zircon U–Pb	2.75–3.5	+5	Liu et al. (2005)	CA
K-granite	W Alataw	290 ± 5		Ar/Ar plateau	2.6–6.2	+0.9 to +3.2	Liu et al. (2005)	TR
Diabase porphyre	Bogda	288.9 ± 4.7		Zircon U–Pb	4.5–5.9		Shu et al. (2005)	A
Biotite granite (B94)	Borohoro	285.3 ± 7.3		Zircon U–Pb	16.8		This study	A
Granitic dyke	Alataw	285 ± 13		Zircon U–Pb	≈1	+2.1 to +2.5	Liu et al. (2005)	CA
Porphyritic granite	S Tianshan	284.4 ± 1.5		Zircon U–Pb	11–25	–4.1 to –6.3	Jiang et al. (1999)	A
K-granite (B102)	Borohoro	280 ± 5		Zircon U–Pb	5.7		This study	TR
Two-mica granite	S Tianshan	280–270		Zircon U–Pb	2–3	–4.4 to –4.7	Jiang et al. (1999)	CA
Rhyolite	Alataw	270.7 ± 6.5		Zircon U–Pb	4–5	+3 to +5.2	Liu et al. (2005)	TR
K-granite (B95)	Borohoro	266 ± 6		Zircon U–Pb	20.9		This study	A
Alkali-granite	S Tianshan	265–260		Zircon U–Pb	1.1–6.8	–4.9 to –6.2	Jiang et al. (1999)	TR
Quartz albitophre	Alataw	260 ± 5		Ar/Ar plateau	1.94–3.6	+1.6 to +3.3	Zhao et al. (2003b)	CA
Trachytic basalt	Nalati	271–261		Whole rock K–Ar	2.9–4.9		Chen et al. (2004a, b)	TR
Trachyte andesite	Nalati	271–261		Whole rock K–Ar	9.1–16		Chen et al. (2004a, b)	A

CA Calc-alkaline series, A alkaline series, TR transitional series



depletion in Ba, Nb, Sr and Ti, enrichment in Rb, K, U and Th, and positive  $\varepsilon_{\text{Nd}}(t)$  values (+2.1 to +5.2) (Chen et al. 1994, 2000; Zhou et al. 1994, 1995, 1996; Liu et al. 2005).

Lower-Middle Permian volcanic rocks are locally distributed in Awulale, Nalati and Alataw areas (XBGMR 1993; Figs. 1, 2). The volcanic and hypabyssal rocks of Awulale and Nalati areas were dated at 296–260 Ma ( $^{40}\text{Ar}/^{39}\text{Ar}$ ; Zhao et al. 2003b). The Wulang Fm. ( $P_{1w}$ ) that develops in the Awulale and Alataw areas is composed of tuffaceous breccia, crystal-bearing tuff, amygdaloidal basalt, porphyritic augite-andesite, rhyolitic porphyrite and dacitic porphyrite (Fig. 2). In Awulale range, Middle Permian Hamist Fm. ( $P_{2h}$ ) consists of basalt, rhyolite, sodic dacite, albitophyre and tuffaceous mudstone that bears fish fossils (XBGMR 1993) (Fig. 2). The albitophyre (260  $\pm$  5 Ma,  $^{40}\text{Ar}/^{39}\text{Ar}$  plateau age) show high Sr, low Yb and Y contents, less variable Nd isotopic ratios ( $^{143}\text{Nd}/^{144}\text{Nd}(i) = 0.5124\text{--}0.5125$ ;  $\varepsilon_{\text{Nd}}(t) = +1.57$  to +3.26), and low Sr isotopic ratios ( $^{87}\text{Sr}/^{86}\text{Sr}(i) = 0.7051\text{--}0.7054$ ) (Xiong et al. 2001). To the Northeast of Nalati range, volcanic rocks develop within terrestrial red beds of Aikendaban Fm. ( $P_{2a}$ ) (Fig. 2), felsite and trachyte yield whole rock K–Ar age of ca. 270–260 Ma, associated trachytic basalt show slight LREE enrichment depletion in Zr, Hf, Nb, Ta and Ba; while trachyte andesite is significantly rich in LREE and depleted in Sr (Chen et al. 2004a, b). To the south of Bogda range, where collapse breccia and deep-water sediments overlie alkaline pillow basalt, a mafic feeding dyke yields Early Permian zircons (289  $\pm$  5 Ma, U–Pb) (Shu et al. 2005).

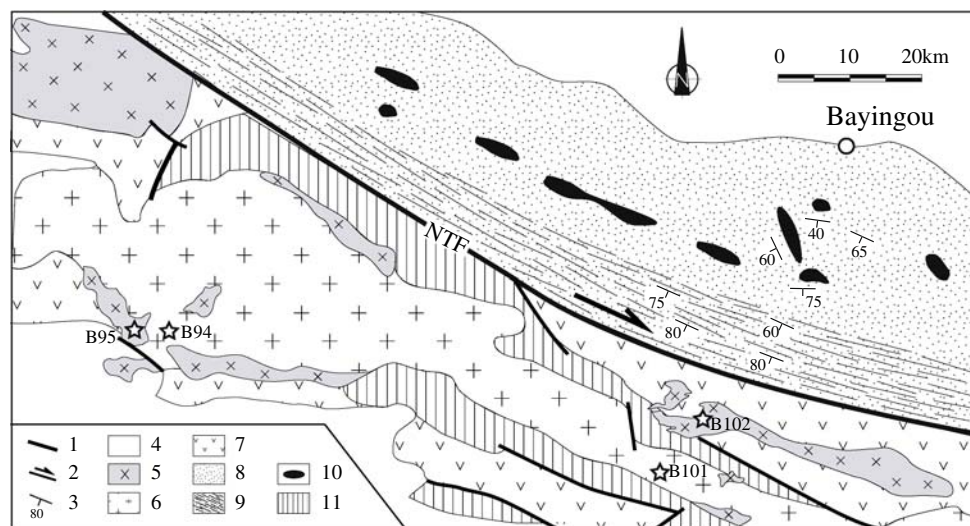
The above-mentioned magmatic rocks are diversely correlated to CA, A and TR types (Table 1). Thus, they could neither be simply related to an extensional intraplate setting, nor to a subduction regime. It appears that the complex geochemical features present in the study area can not unambiguously constrain the geodynamic setting at a given period. The Carboniferous magmatic rocks are mostly of CA type, whereas diverse suites (CA, A and TR) coexist during the Permian. Therefore, it is necessary to take into account every field and structural features before proposing an evolutionary model. In order to get more constraints on the chronology and petrogenesis of the Permian magmatism, several key sections across the main wrench faults were investigated in the following sections.

### The Borohoro plutons

The Borohoro plutons develop to the south of the North Tianshan Fault (NTF) in a poorly accessible area. They show a spindle shape extending parallel to the NTF (Figs. 1, 3). They were previously correlated with the Carboniferous on 1:1,500,000 and 1:200,000 geological maps (XBGMR 1973, 1975, 1993), but remained undated and their geochemical compositions were unknown.

### Petrology

The Borohoro plutons are mainly composed of greyish-white biotite granite and pink K-granite; both are characterized by



**Fig. 3** Structural sketch maps of Borohoro area (modified from XBGMR 1973, 1975), showing the spatial relationship of granitic plutons with ductile dextral North Tianshan Fault. For location, see Fig. 1. The sampling localities are marked with *open pentacles* followed by sample numbers, and their GPS coordinates are listed in Table 2. 1 fault, 2 North Tianshan Fault zone, 3 foliation/bedding and

dip angle, 4 Meso-Cenozoic, 5 Permian pink granite, 6 Permian dark granite, 7 Early Carboniferous volcanoclastic rocks, 8 weakly deformed Carboniferous turbidite, 9 mylonitic and metamorphic turbidite, 10 blocks of ultra-mafic rocks, 11 Silurian-Devonian undeformed sedimentary rocks

coarse grain texture, and typically consist of plagioclase, K-feldspar, quartz, biotite and hornblende; Biotite often occurs as euhedral grains and rarely as amphibole pseudomorphs. In K-granite, pyroxene relics are preserved in poikilitic hornblende. Accessory minerals include apatite, zircon and ilmenite.

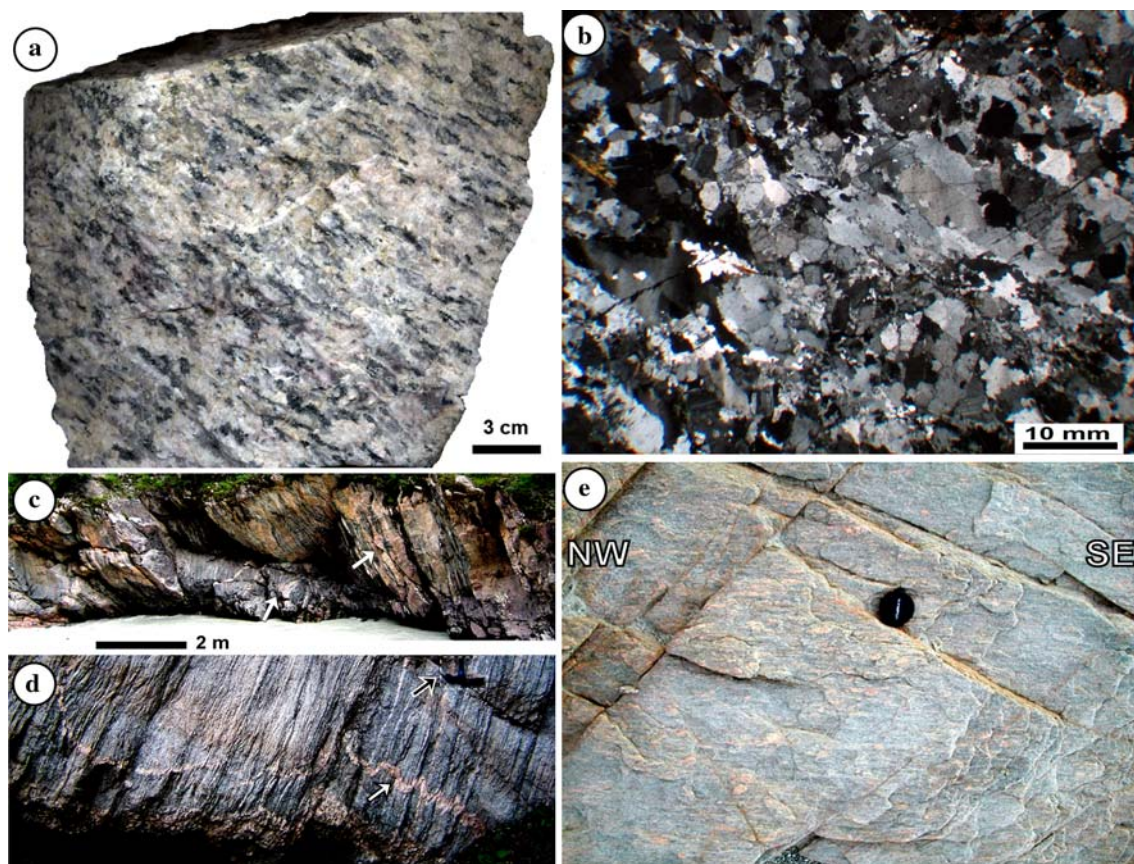
#### Internal structure

A deformation fabric can be recognized in most rocks of the Borohoro plutons, it comprises shallow linear and steep planar textures. The lineation is defined by aligned biotite, hornblende, and feldspar grains and elongated quartz aggregates, the planar fabric is marked by ribbons of mafic minerals (Fig. 4a). Micro-texture does not show significant internal deformation in euhedral grains of plagioclase, biotite, hornblende and interstitial ovoid quartz, although locally quartz domains display minor sutured boundaries and weak undulose extinction, and occasional fracturing

and disaggregation of the grain corners (Fig. 4b). These are typical high-temperature fabrics developed in condition of pre-critical melt percentage (e.g. Tribe and D'Lemos 1996) suggesting a still magmatic state deformation. Although no systematic investigation has been undertaken, the planar fabric strikes N120°–N130°, slightly oblique to the strike of the NTF, but parallel to the intrusion margin.

#### Relationships with country rocks and the NTF

The Silurian-Devonian (?) and Early Carboniferous host rocks of the Borohoro plutons (Fig. 3) display an unusual kind of contact metamorphism. Spotted slate often develops within the contact zones while hornfels is rare. The host rock is characterized by steeply dipping foliation that bears a shallow dipping stretching lineation with andalusite porphyroblasts. This foliation represents the effect of syn-tectonic thermal metamorphism due to synkinematic granite intrusion associated with a ductile shear zone



**Fig. 4** **a** Photograph of hand specimen of Borohoro biotite K-granite showing the linear fabric defined by alignment of biotite, hornblende, feldspar and elongated quartz, and the planar fabric marked by ribbons of mafic minerals; **b** microphotograph of Borohoro K-granite, pale grains are mostly quartz, feldspar and brown minerals are biotite and minor hornblende, dark grains are mainly extinct quartz and

feldspar; **c, d** field photographs in Kekesu section showing boudinaged or folded K-granite dykes (directed by white arrows) intruding the mylonitic gabbro, black arrow in **d** shows the location of a hammer for scale; **e** mylonitized K-granite in Gangou section, shallow dipping lineation included in sub-vertical SE extending foliation



(Wang et al. 2006a). Along the northern margin of the plutons, metre- to decimetre-scale sheets of pink K-granite dykes intrude Carboniferous turbidite. Although some are randomly oriented, most of granitic dykes have an east-west trend, sub-parallel to the NTF, suggesting a structural control on magma emplacement as well.

### Kekesu and Gangou sections

In order to make a comparison, two sections were also investigated in the Qingbulak-Nalati and Sangshuyuanzi shear zones, which are briefly described below. In the Kekesu section (Fig. 5), massive pink granites are composed of K-feldspar, plagioclase, quartz and minor mafic minerals that are often chloritized and/or epidotized, they intrude cataclastic or mylonitized gabbro of Devonian to Carboniferous age and probable Proterozoic orthogneiss. Locally, K-granite dykes are boudinaged or folded, showing axial plane parallel to the SW–NE trending gabbro foliation (Fig. 4c, d). Within this shear zone, Carboniferous granite ( $313 \pm 4$  Ma, U–Pb age on zircon) underwent ductile deformation and yield Middle Permian biotite ( $263.4 \pm 0.6$  Ma,  $^{40}\text{Ar}/^{39}\text{Ar}$  plateau age) related to Ar loss

during mylonitisation (Wang et al. 2007c). The boundaries of the K-granite are generally parallel to the Qingbulak-Nalati fault zone (Fig. 5), indicating a close relationship between the emplacement of granite and the shear zone.

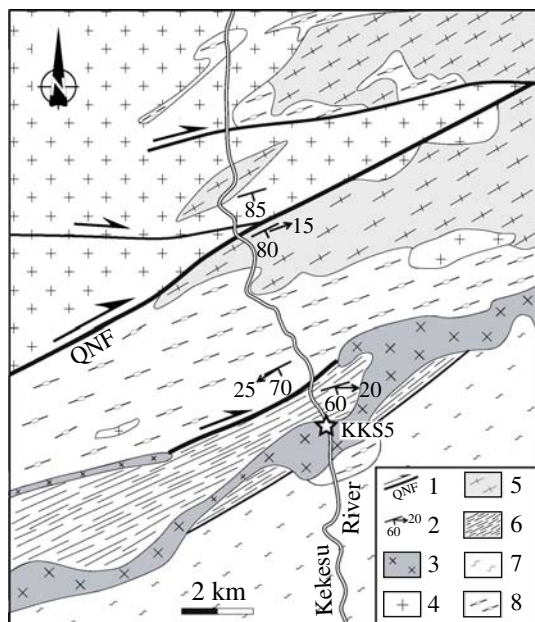
In Gangou area, the Sangshuyuanzi Fault extends parallel to the SE–NW Main Tianshan Shear Zone (MTSZ) which is connected westwards with the NTF (Figs. 1, 6). K-granite and synchronous dykes showing a SE–NW elongated shape intrude Proterozoic orthogneiss and Silurian–Carboniferous strata. Both K-granite and its country rock display a well developed mylonitic foliation that strikes N130°E (Fig. 4e), and typically bears a low dip stretching lineation ( $15^\circ$ – $20^\circ$  SE). Sigmoid K-feldspar and asymmetrical biotite pressure shadow reveal a dextral ductile deformation (Allen et al. 1995; Allen and Vincent 1997; Laurent-Charvet et al. 2002). Although the granite underwent significant mylonitisation, a coarse to medium grain texture can be recognized, K-feldspar, quartz, plagioclase, minor biotite and sometimes amphibole are the main minerals.

### New geochronological and geochemical data

Two biotite granite and two K-granite samples were collected from the Borohoro plutons, one K-granite and one biotite K-granite samples were collected from the Kekesu and the Gangou sections, respectively. Sampling locations are shown in Figs. 3, 5, 6 and Table 2.

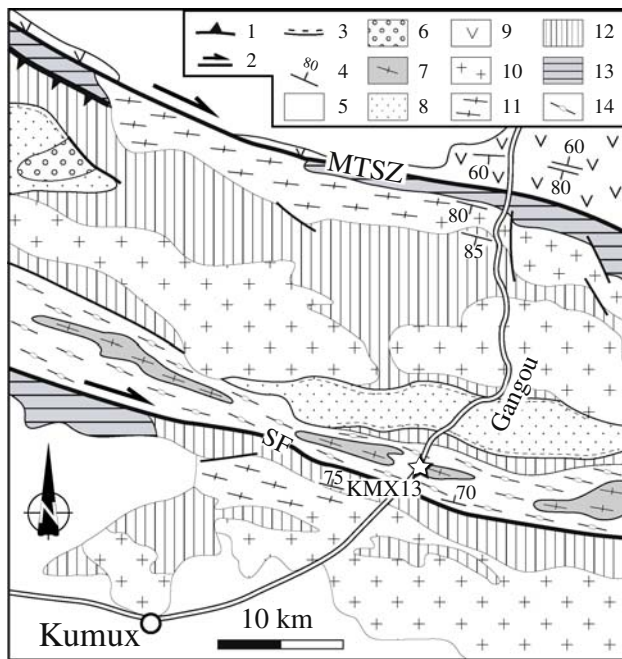
### Analytical techniques

The selected samples were crushed and milled into rock powder, zircon crystals are enriched using heavy liquids and magnetic separator, and finally selected by handpicking under binocular microscope. Euhedral and colorless zircon grains were selected for laser ablation U–Pb dating that was carried out at the University of Tasmania, Australia, using a Hewlett Packard HP 4500 quadrupole Inductively Coupled Plasma Mass Spectrometer (ICP-MS) coupled with a 213 nm NewWave Merchantek UP213 Nd-YAG Laser. Preserved rock powder was used for whole rock geochemical analyses by ICP-AES for major elements and by ICPMS for incompatible and rare earth elements at Centre de Recherches Pétrographiques et Géo-chimiques (CRPG-CNRS, Nancy, France). The limit of determination is less than 0.07% for major elements, and less than 0.5 ppm for most of trace elements, and up to 1.5–6 ppm for Ba, Cr, Sr, Zn and Zr. Analytical uncertainties are given as 2% for major elements, and 5 or 10% for trace element concentrations around 20 ppm, and the precision for REE is estimated at 5 or 10% depending on the chondrite-normalised concentrations are >10 ppm or lower, respectively.



**Fig. 5** Structural sketch map of Kekesu River section (modified from XBGMR 1979). For location, see Fig. 1. The sampling locality is marked with *open pentacles* followed by sample numbers, and reference Table 2 for the GPS coordinate. 1 Qingbulake-Nalati fault zone, 2 foliation with dip angle and lineation with pitch angle, 3 Permian A-type K-granite, 4 Carboniferous undeformed I-type granite, 5 Devonian?-Carboniferous strongly deformed granite and gabbro, 6 mylonite, 7 greenschist facies meta-sedimentary rocks, 8 Proterozoic gneiss





**Fig. 6** Structural sketch map of Gangou section (modified from XBGMR 1959). For location, see Fig. 1. The sampling locality is marked with *open pentacle* followed by sample number, and the GPS coordinate can be found in Table 2. 1 Early Palaeozoic thrust, 2 strike-slip faults, where MTSZ refers to the Mains Tianshan Shear Zone, and SF represents the Sangshuyuanzi Fault, 3 unconformity, 4 foliation/bedding and dip angle, 5 Cenozoic, 6 Permian molassic deposits, 7 Permian foliated K-granite, 8 Carboniferous sedimentary rocks, 9 Carboniferous volcaniclastic rocks, 10 Early Palaeozoic undeformed granite, 11 Early Palaeozoic foliated granite, 12 Ordovician-Devonian rocks, 13 ophiolitic melange, 14 Proterozoic orthogneiss

#### LA-ICPMS dating

Six granitoids samples were dated, and twelve zircon single grains were analysed for each sample. The results are shown in Table 2 and plotted on reversed Concordia diagrams (Fig. 7). Although different in U and Th absolute compositions, all zircons show consistently concordant Th/U ratios of  $>0.2$ , indicating a magmatic origin (Vavra et al. 1996). Four granitoids from Borohoro Mountain were dated. The formation time of the granodiorite sample B101 can be constrained by nine zircons yielding a concordant age of  $294 \pm 7$  Ma (Fig. 7a), while the rest three zircons provide older ages of 312–589 Ma that probably represent inherited zircon derived from Palaeozoic country rocks. Zircons of granodiorite sample B94 can be divided into two groups, wherein the uranium contents of high U zircons are ten times more than those of low U ones (Table 2). The first low U zircon was rejected due to lots of common Pb, five low U zircons and one high U zircon yield consistent apparent ages defining a concordant age at  $272.8 \pm 6$  Ma, and the last

low U zircon together with the other four high U zircons yield a concordant age at  $293 \pm 0.5$  Ma (Fig. 7b); The latter one is consistent with the age of the sample B101 at the uncertainty level, but the former one is significantly younger, indicating that the granodiorite was emplaced during two distinct episodes. All zircons together give an average age at  $285.3 \pm 7.3$  Ma.

For the K-granite sample B95, the first zircon display high common Pb content throughout the analysis inducing young age and large analytical uncertainty ( $225 \pm 10$  Ma), the second one also yield a rather younger age of  $247 \pm 4$  Ma but shows no analytical anomaly, this is probably due to statistical outlier. The last zircon yielding an age of  $340 \pm 3$  Ma is likely an inherited one mixed with older core. The rest nine zircons provide concordant ages ranging from 257–277 Ma that are used to calculate a concordant age at  $266 \pm 6$  Ma, indicative of the formation time of the K-granite (Fig. 7c). On the Concordia diagram (Fig. 7d) of the K-granite B102, except one point that was rejected due to discordant apparent age, the other eleven points were used in the calculation of concordant age at  $280 \pm 5$  Ma.

Seven zircons from the Kekesu K-granite KKS5 yield concordant apparent ages that are used to calculate a weighted average U–Pb age at  $277 \pm 3$  Ma. Older ages of 312–375 Ma are also obtained from two zircons that were probably derived from foliated/mylonitized country rocks. The other three points are discordant because of either Pb loss in high U zircon (the first point), or high proportion of common Pb (the last one) (Fig. 7e) and consequently provide geologically meaningless ages.

From the Gangou K-granite KMX13, eight points concordantly define a U–Pb age at  $252 \pm 4$  Ma (Fig. 7f), and can be interpreted as the intrusion time of the K-granite. The last four zircons yield significantly older ages ranging from 312 to 357 Ma (Table 2; Fig. 7f), which may reflect the incorporation of zircon xenocrysts derived from Carboniferous country rocks. The occurrence of Devonian-Carboniferous strata and granitic rocks around the K-granite (Fig. 6; XBGMR 1993; Xu et al. 2006) confirms this interpretation.

#### Geochemical composition

Only four granitic rocks from Borohoro Range were analysed for whole rock geochemistry. Table 3 comprises the already published geochemical data and our new results. Borohoro granites are significantly rich in alkali (7–9.7 wt%) with relatively high content of  $K_2O$  with respect to  $Na_2O$ . Relative low aluminum contents (13.8–14.1 wt%; Aluminum Saturation Index  $\approx 1$ ) indicate that they belong to metaluminous or slightly aluminous I-type granitoid series (Fig. 8).

**Table 2** Zircon U–Pb LA-ICPMS analysis results of the granitoids from the western Chinese Tianshan

No	Compositions			Atomic ratios			Apparent ages (Ma)			Disc (%)						
	Th (ppm)	U (ppm)	Th/U	$^{207}\text{Pb}/^{235}\text{U}$	$^{206}\text{Pb}/^{238}\text{U}$	$^{207}\text{Pb}/^{206}\text{Pb}$	$^{207}\text{Pb}/^{235}\text{U}$	$^{206}\text{Pb}/^{238}\text{U}$	$^{207}\text{Pb}/^{206}\text{Pb}$	$^{207}\text{Pb}/^{235}\text{U}$	$^{206}\text{Pb}/^{238}\text{U}$	$^{207}\text{Pb}/^{206}\text{Pb}$				
Borohoro granodiorite B94 (GPS:N43°50.30', E83°31.55')																
1	262	424	0.62	0.32	0.06	0.0020	0.012	0.068	0.068	283	44	249	13	880	367	13
2	138	232	0.60	0.30	0.02	0.0006	0.003	0.053	0.053	266	14	268	4	317	136	4
3	207	402	0.52	0.31	0.02	0.0006	0.003	0.054	0.054	274	12	270	4	389	113	4
4	120	320	0.38	0.27	0.02	0.0007	0.004	0.049	0.049	247	14	268	4	156	175	4
5	73	163	0.45	0.30	0.02	0.0006	0.003	0.050	0.050	263	15	275	4	195	154	4
6	72	150	0.48	0.33	0.02	0.0008	0.005	0.056	0.056	286	19	282	5	453	182	5
7	62	126	0.49	0.33	0.03	0.0009	0.005	0.056	0.056	292	22	298	6	449	207	6
8	2,671	5,013	0.53	0.30	0.01	0.0003	0.001	0.051	0.051	270	6	276	2	248	58	2
9	1,504	3,339	0.45	0.32	0.01	0.0003	0.001	0.052	0.052	283	5	291	2	280	51	2
10	2,480	5,844	0.42	0.32	0.01	0.0004	0.001	0.052	0.052	284	5	293	2	267	44	3
11	3,498	6,418	0.55	0.33	0.01	0.0004	0.001	0.053	0.053	293	5	296	2	323	44	2
12	3,035	5,215	0.58	0.35	0.01	0.0007	0.002	0.055	0.055	303	11	297	4	419	97	4
Borohoro K-granite B95 (GPS:N43°50.30', E83°31.55')																
1	138	137	1.01	2.82	0.16	0.0017	0.014	0.352	0.352	1,361	45	356	10	3,714	62	10
2	79	178	0.45	0.33	0.02	0.0007	0.004	0.063	0.063	293	15	250	4	695	127	4
3	3,098	5,976	0.52	0.46	0.01	0.0003	0.002	0.080	0.080	387	6	266	2	1,186	37	2
4	1,691	3,324	0.51	0.43	0.01	0.0005	0.003	0.076	0.076	365	11	266	3	1,087	67	3
5	97	182	0.53	0.34	0.02	0.0006	0.004	0.061	0.061	296	17	264	4	634	153	4
6	124	279	0.44	0.34	0.02	0.0006	0.004	0.058	0.058	296	17	264	3	522	138	4
7	1,270	2,716	0.47	0.31	0.01	0.0003	0.001	0.054	0.054	274	6	266	2	354	52	2
8	118	181	0.65	0.31	0.03	0.0013	0.007	0.058	0.058	277	26	267	8	514	252	8
9	698	1,998	0.35	0.35	0.01	0.0004	0.002	0.060	0.060	308	7	271	3	621	67	3
10	8,271	10,479	0.79	0.32	0.01	0.0002	0.001	0.054	0.054	282	5	270	2	381	42	2
11	1,445	3,798	0.38	0.32	0.01	0.0003	0.001	0.053	0.053	281	5	277	2	331	50	2
12	203	527	0.38	0.39	0.01	0.0005	0.002	0.052	0.052	331	11	339	3	276	95	3
Borohoro granodiorite B101 (GPS:N43°41.34', E84°24.86')																
1	99	231	0.43	0.26	0.02	0.001	0.004	0.044	0.044	234	17	274	6	–86	218	6
2	317	717	0.44	0.40	0.02	0.001	0.003	0.061	0.061	343	16	289	4	644	118	4
3	184	349	0.53	0.36	0.02	0.001	0.004	0.056	0.056	308	19	289	5	451	171	6
4	140	270	0.52	0.37	0.03	0.001	0.004	0.059	0.059	323	20	290	6	575	164	6
5	156	562	0.28	0.34	0.02	0.001	0.004	0.054	0.054	294	18	289	5	363	165	5
6	60	112	0.54	0.34	0.04	0.001	0.007	0.054	0.054	296	30	298	9	388	289	9

Table 2 continued

No	Compositions			Atomic ratios			Apparent ages (Ma)			Disc (%)				
	Th (ppm)	U (ppm)	Th/U	$^{207}\text{Pb}/^{235}\text{U}$	$^{206}\text{Pb}/^{238}\text{U}$	$^{207}\text{Pb}/^{206}\text{Pb}$	$^{207}\text{Pb}/^{235}\text{U}$	$^{206}\text{Pb}/^{238}\text{U}$	$^{207}\text{Pb}/^{206}\text{Pb}$					
7	208	467	0.45	0.35	0.02	0.052	0.003	307	15	301	5	274	136	5
8	319	1,070	0.30	0.37	0.01	0.048	0.001	322	11	303	3	416	91	3
9	332	1,567	0.21	0.61	0.02	0.050	0.001	486	13	317	5	1,336	67	5
10	169	347	0.49	0.38	0.02	0.049	0.001	326	18	311	5	392	149	5
11	284	564	0.50	0.38	0.02	0.050	0.001	325	16	315	5	376	137	5
12	390	1,111	0.35	0.81	0.03	0.096	0.001	605	16	590	7	594	80	7
Borohoro K-granite B102 (GPS: N43°46.10', E84°26.2')														
1	57	94	0.60	0.61	0.04	0.044	0.001	483	27	279	6	1,827	129	6
2	36	89	0.41	0.28	0.02	0.043	0.001	249	20	270	6	377	210	6
3	42	81	0.52	0.35	0.03	0.044	0.001	302	24	278	6	786	188	6
4	54	95	0.57	0.27	0.03	0.044	0.001	245	23	275	6	166	245	6
5	38	72	0.53	0.33	0.04	0.045	0.001	292	28	281	7	659	231	8
6	50	87	0.58	0.25	0.03	0.044	0.001	223	25	279	6	307	275	6
7	44	72	0.61	0.28	0.03	0.045	0.001	249	26	281	7	272	275	7
8	64	92	0.70	0.24	0.03	0.044	0.001	220	27	280	7	–153	342	7
9	90	128	0.70	0.30	0.03	0.045	0.001	263	23	284	6	366	220	6
10	31	67	0.47	0.30	0.04	0.045	0.001	267	29	285	7	305	295	7
11	58	95	0.61	0.34	0.04	0.046	0.001	294	28	292	6	647	236	7
12	29	66	0.44	0.39	0.04	0.048	0.001	335	26	301	7	857	194	8
Kekesu K-granite KKS5 (GPS: N42°45.98', E81°55.44')														
1	1,694	3,255	0.52	0.32	0.01	0.037	0.0003	279	6	235	2	778	50	2
2	1,131	889	1.27	0.54	0.02	0.044	0.0005	441	13	280	3	1,487	68	3
3	333	809	0.41	0.36	0.01	0.044	0.0004	311	10	278	2	674	87	3
4	1,754	998	1.76	0.31	0.01	0.044	0.0004	271	9	275	2	341	83	2
5	536	718	0.75	0.46	0.03	0.045	0.0011	383	25	286	7	1,209	145	7
6	746	691	1.08	0.43	0.02	0.045	0.0006	361	15	285	4	977	99	4
7	752	832	0.90	0.33	0.01	0.044	0.0004	290	11	280	3	468	94	3
8	616	1,464	0.42	0.36	0.01	0.045	0.0004	311	7	283	2	649	56	2
9	458	758	0.60	0.45	0.02	0.046	0.0005	378	15	288	3	1,048	95	3
10	503	434	1.16	0.37	0.03	0.050	0.0011	323	20	315	7	558	152	7
11	182	648	0.28	0.42	0.02	0.060	0.0006	358	11	375	4	372	81	4
12	746	501	1.49	3.49	0.33	0.074	0.0034	1,526	76	458	20	3,760	92	19

Table 2 continued

No	Compositions			Atomic ratios			Apparent ages (Ma)			Disc (%)						
	Th (ppm)	U (ppm)	Th/U	$^{207}\text{Pb}/^{235}\text{U}$	$^{206}\text{Pb}/^{238}\text{U}$	$^{207}\text{Pb}/^{206}\text{Pb}$	$^{207}\text{Pb}/^{235}\text{U}$	$^{206}\text{Pb}/^{238}\text{U}$	$^{207}\text{Pb}/^{206}\text{Pb}$							
Gangou biotite K-granite KMX13 (GPS: N:42°31.22', E:88°31.08')																
1	837	627	1.34	0.27	0.01	0.039	0.001	0.052	0.002	244	10	248	3	306	100	3
2	242	240	1.01	0.30	0.02	0.040	0.001	0.056	0.004	264	15	251	4	458	141	4
3	658	698	0.94	0.32	0.02	0.040	0.001	0.061	0.003	282	13	254	5	654	119	5
4	277	577	0.48	0.32	0.02	0.040	0.001	0.063	0.004	284	16	255	5	709	128	5
5	617	451	1.37	0.32	0.01	0.041	0.001	0.060	0.003	281	11	258	4	603	97	5
6	120	227	0.53	0.30	0.03	0.041	0.001	0.057	0.006	267	25	258	6	499	233	6
7	58	122	0.48	0.29	0.04	0.041	0.001	0.054	0.007	261	32	259	8	361	314	8
8	370	317	1.17	0.33	0.03	0.042	0.001	0.061	0.005	287	21	262	7	656	173	7
9	123	214	0.58	0.33	0.03	0.049	0.001	0.050	0.005	288	23	311	6	204	217	6
10	207	325	0.63	0.36	0.02	0.054	0.001	0.051	0.003	314	17	337	6	259	149	6
11	852	874	0.98	0.40	0.02	0.057	0.001	0.053	0.002	338	12	356	6	317	90	6
12	757	554	1.37	0.43	0.02	0.057	0.001	0.057	0.003	362	17	358	7	474	121	7

The samples B101 and B102 are characterized by enrichment in Rb, Th and K compared to REE, and low contents of Nb and Ta (Table 3), indicative of calc-alkaline series. However, the samples B94 and B95 show lower Ba (125–134 ppm) and Sr (20–40 ppm), characteristic of transitional type as defined above. A similar feature, e.g. co-existence of both calc-alkaline and transitional series, occurs for Alataw and southern Tianshan granitoids (Jiang et al. 1999; Liu et al. 2005). This feature can be observed on the expanded REE and trace elements spider diagram normalized to the primitive mantle (Sun and McDonough 1989) (Fig. 9a), in which, all the granitic rocks of Borohoro, Alataw and southern Tianshan as well as Nalati volcanic rocks exhibit a regular decrease with increasing compatibility for HFS elements. In addition, depletions of Nb, Ta, Y and Yb can be observed for all these magmatic rocks. Otherwise, the samples B94 and B95 exceptionally show enrichments of Ce, Zr and Hf.

The samples B101 and B102 display total REE content of 90–188 ppm, slight LREE enrichment ( $\text{La}_N/\text{Yb}_N = 5.7\text{--}7.2$ ), weak negative Eu anomaly ( $\delta\text{Eu} = 0.6\text{--}0.88$ ). They have low Ce/Pb (1.7–3.7) and Nb/La ( $\approx 0.4$ ) ratios, and Th/La ratio of 0.4–0.7 (Table 3). These features, consistent with calc-alkaline granites, are similar to those of Alataw granitoids, southern Tianshan two-mica granite and potassium granite (Table 3). In contrast, the samples B94 and B95 have higher REE contents (231–236 ppm), higher  $\text{La}_N/\text{Yb}_N$  (16.8–20.9) and Ce/Pb (7.7–6.9) ratios, and lower Nb/La ( $\approx 0.16$ ) and Th/La ( $\approx 0.15$ ) ratios (Table 3). These features are those of alkaline or transitional granites, and are comparable to the characteristic of the trachytic andesite from Nalati and granodiorite from southern Tianshan (Table 3). On the Chondrite-normalized REE patterns (Fig. 9b; Pearce 1982), samples B101 and B102 of Borohoro plutons, and Alataw granitoids show lesser REE fractionation by comparison with the samples B94 and B95 as well as the volcanic rocks of the Nalati and Bogda areas.

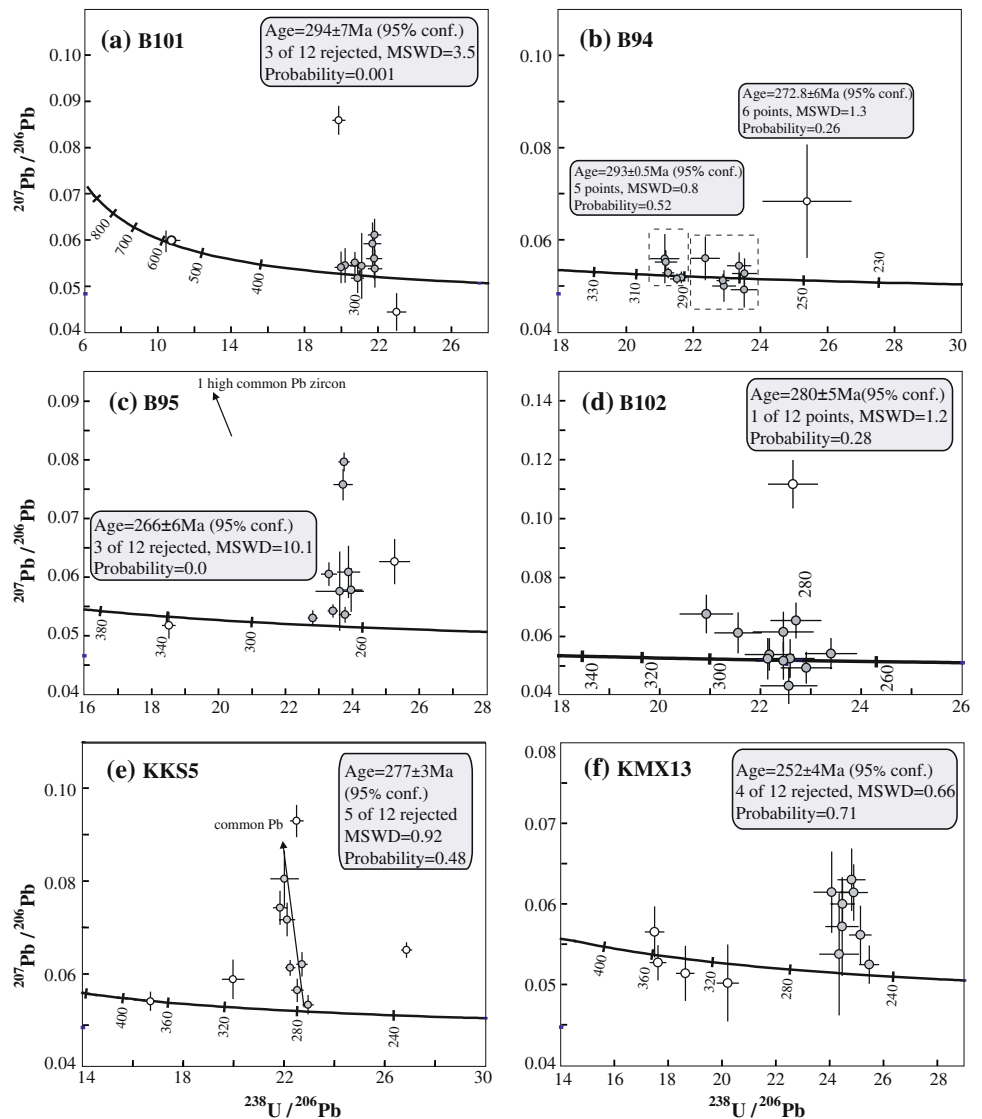
## Discussion

### Persistence and evolution of magmatism through Carboniferous to Permian

On the basis of our study of the Borohoro plutons, the high-K calc-alkaline granites formed during 294–280 Ma, whereas the transitional plutonism occurred between 285 and 266 Ma, and continued until 250 Ma in other places. Combining the previous results synthesized above, the magmatic activity in western Chinese Tianshan was a permanent process that started as early as the beginning of Carboniferous and lasted up to the end of Permian. Geochemical compositions suggest that the Carboniferous



**Fig. 7** Inversed Concordia diagrams of zircon U–Pb LA-ICPMS dating results on the granitoids from the Western Chinese Tianshan. *Filled gray circles* refer to the plots used to calculate the concordant ages, and *open circles* represent rejected plots



magmatic rocks are mainly calc-alkaline although a few transitional series may occasionally appear (Table 1); in contrast, the Permian magmatic rocks are characterised by the coexistence of calc-alkaline, alkaline and transitional series, indicating a transition of magma chemistry from calc-alkaline to alkaline.

#### Petrogenesis of Carboniferous-Permian magmatic rocks and geodynamic implications

Some authors favor a continental rift (Che et al. 1996; Xia et al. 2004b) or an intra-plate environment (Liu et al. 2006) for the Carboniferous volcanic rocks, and a post-collision setting for the Carboniferous granites (Xu et al. 2006). Nevertheless, their geochemical and isotopic features indicate that subduction play a prominent role, such as suggested by the occurrence of adakite, high-Mg andesite, and Nb-poor andesitic basalt (Wang et al. 2003, 2006b;

Zhao et al. 2003b, 2006). Moreover, stratigraphic features of Carboniferous volcanic rocks and the shallow water sedimentary association (Fig. 2) suggest a rather long period of eruption on a continental margin. Elsewhere in Central Asia, Carboniferous subduction-related and/or syncollisional igneous rocks with calc-alkaline affinities are reported in the South Tianshan belt of Kyrgyzstan (Hamrabaev and Simon 1984; cited by Solomovich and Trifonov 2002) and around the Tu-Ha basin of the Chinese eastern Tianshan (Li et al. 2001, 2006a, b; Sun et al. 2006). Therefore, Carboniferous magmatic rocks have most probably been generated in an active margin setting.

Usually, the Permian high-K calc-alkaline and alkaline igneous rocks were considered to be post-collisional, e.g. in the South Tianshan of Kyrgyzstan (Solomovich and Trifonov 2002; Solomovich 2007). Particularly, the Permian high-K calc-alkaline and transitional granites of the Borohoro plutons also have subduction components according

**Table 3** A synthesis of Whole rock geochemistry of Permian magmatic rocks in Borohoro, Alataw, Southern Tianshan, Nalati and Bogda areas, West Chinese Tianshan

Locations	Borohoro			Alataw			Nalati (Aikendaban Fm.)						
	Sample no. Age Rocks	B94 285 Ma Granodiorite	B95 266 Ma K-granite	B101 294 Ma Granodiorite	B102 280 Ma K-granite	kwsy12 298 Ma Granite	kwsy3 285 Ma Granitic dyke	kwsy27 292 Ma Granite	kwsy17 271 Ma Rhyolite	al-39 P2 Shoshonite (?)	al-45 P1 Trachyte andesite	al-38 P1 Trachytic andesite	al-4 P1 Trachyte
References	This study				Liu et al. (2006)					Chen et al. (2004a, b)			
SiO <sub>2</sub>	71.4	71.35	72.79	69.39	73.05	76.97	77.16	74.45	41.69	53.46	54.79	59.99	
TiO <sub>2</sub>	0.23	0.22	0.24	0.47	0.29	0.08	0.1	0.25	0.91	1.31	1.19	0.94	
Al <sub>2</sub> O <sub>3</sub>	14.01	13.78	14.06	13.93	13.39	12.43	12.36	11.95	18.29	15.2	17.23	15.63	
Fe <sub>2</sub> O <sub>3</sub>	2.66	2.54	1.77	3.9	2.43	0.93	1.38	2.28	4.14	3.13	3.31	2.6	
MnO	0.05	0.06	0.05	0.08	0.05	0.01	0.02	0.04	0.23	0.16	0.13	0.15	
MgO	0.1	0.09	0.51	0.49	0.52	0.04	0.03	0.14	4.78	7.46	3.92	3.37	
CaO	0.88	0.8	2.09	2.03	1.65	0.44	0.44	1.31	20.43	5.01	5.31	0.76	
Na <sub>2</sub> O	3.9	3.78	3.72	4.02	3.7	3.91	3.85	3.35	0.33	4.75	4.22	4.1	
K <sub>2</sub> O	5.75	5.65	3.3	3.72	4.37	4.57	4.8	3.93	0.11	1.75	4.26	5.72	
P <sub>2</sub> O <sub>5</sub>	0.04	0.05	0.08	0.12	0.07	0.01	0.01	0.06	0.2	0.33	0.34	0.47	
LOI	0.29	0.33	0.78	0.66	0.68	1.09	0.6	1.07	5.55	2.95	2.33	3.15	
Total	99.3	98.64	99.37	98.82	100.21	100.47	100.75	98.82	99.96	99.99	100.03	99.98	
ALK	9.65	9.43	7.02	7.74	8.07	8.48	8.65	7.28	0.44	6.5	8.48	9.82	
ASI	0.98	1.00	1.04	0.97	0.97	1.02	1.00	0.98	0.48	0.81	0.81	1.09	
La	52.07	54.77	18.35	33	31.88	18.16	40.31	38.73	18.82	30.28	50.62	19.75	
Ce	103.6	106.2	37.36	71.59	68.45	49.72	92.39	83.2	38	68.92	171.8	40.4	
Pr	11.69	12.11	4.222	8.925	8.63	7.28	12.13	10.49	4.47	7.93	13.87	5.83	
Nd	42.58	43.49	15.48	35.96	31.94	27.56	44.23	39.79	21.49	31.24	49.6	22.18	
Sm	6.416	6.315	3.36	8.295	7.7	7.82	11.96	10.2	5.72	5.85	9.05	6.47	
Eu	0.468	0.471	0.637	2.323	0.82	0.1	0.18	1.53	1.71	1.79	2.17	1.34	
Gd	4.373	4.104	3.05	7.651	7.7	8.25	12.21	9.76	7.03	5.94	9.64	6.97	
Tb	0.65	0.572	0.498	1.238	1.36	1.76	2.26	1.66	1.03	0.79	1.19	0.99	
Dy	3.588	3.171	2.914	7.538	7.89	11.84	12.59	10.02	6.07	3.81	6.06	5.48	
Ho	0.676	0.591	0.574	1.48	1.67	2.76	2.76	1.92	1.38	0.86	1.2	1.39	
Er	1.992	1.729	1.666	4.222	5.01	9.01	8.45	6.04	3.38	2.11	2.71	3.82	
Tm	0.307	0.263	0.26	0.623	0.79	1.51	1.24	0.87	0.48	0.32	0.4	0.58	

Table 3 continued

Locations	Borohoro			Alataw			Nalati (Aikendaban Fm.)					
	Sample no.	Age	Rocks	B102	kwsy12	kwsy3	kwsy27	kwsy17	al-39	al-45	al-38	al-4
	B94	B95	B101	B102	kwsy12	kwsy3	kwsy27	kwsy17	al-39	al-45	al-38	al-4
	285 Ma	266 Ma	294 Ma	280 Ma	298 Ma	285 Ma	292 Ma	271 Ma	P2	P1	P1	P1
	Granodiorite	K-granite	Granodiorite	K-granite	Granite	Granitic dyke	Granite	Rhyolite	Shoshonite (?)	Trachyte andesite	Trachytic andesite	Trachyte
References	This study											
	Liu et al. (2006)											
Yb	2.218	1.878	1.838	4.159	5.25	10.18	8.07	5.72	2.62	1.49	2.14	3.31
Lu	0.376	0.321	0.283	0.641	0.8	1.62	1.13	0.87	0.39	0.25	0.33	0.54
ΣREE	231	235.99	90.49	187.65	179.89	157.57	149.91	220.8	112.59	161.58	320.78	119.05
(La/Yb) <sub>N</sub>	16.84	20.92	7.16	5.69	4.20	1.24	3.46	4.69	4.85	13.73	15.98	4.03
δEu	0.26	0.27	0.60	0.88	0.11	0.01	0.02	0.15	0.82	0.92	0.71	0.61
V	1.611	1.903	18.02	14.89	-	-	-	-	238.4	172.1	165.9	137.5
Cr	5.857	9.21	10.65	7.22	-	-	-	-	232.9	282.3	130.1	38.96
Co	0.734	0.796	2.815	3.26	4.46	1.01	0.45	1.82	14.72	19.43	18.96	15.79
Ni	-	3.726	4.38	3.479	-	-	-	-	49.29	111.8	53.94	30.32
Cu	-	-	6.963	6.343	4.01	47.06	2.73	14.78	45.6	95.42	112	8859
Zn	36.14	52.38	40.96	72.43	-	-	-	-	159.1	129.6	90.69	99.28
Ga	18.04	18.41	17.68	18.56	21.38	12.94	13.46	33.26	23.79	27.52	23.64	22.76
Ge	1.175	1.109	1.374	1.434	-	-	-	-	-	-	-	-
Rb	73	67.33	162	120.2	187.58	225.01	203.52	103.55	-	-	-	-
Sr	40.44	19.81	231.6	251.8	133.18	14.81	15.81	126.2	49	440	680	55
Y	18.13	15.93	18.06	43.3	45.88	83.23	76.59	52.03	39.78	19.23	23.38	42.59
Zr	410.7	414.5	110.7	533.8	194.92	143.23	215.37	265.18	77.9	193	292	219
Nb	8.563	8.768	8.131	11.58	17.01	13.35	15.27	13.63	9.22	10.7	15.97	13.21
Cs	9.741	8.138	7.14	6.533	8.06	5.08	1.36	8.63	-	-	-	-
Ba	134.3	124.6	330.3	1152	407.87	47.96	60.26	881.47	76	312	826	934
Hf	9.486	9.55	3.21	10.56	7.02	7.64	9.22	8.59	2.12	3.25	6.19	6.03
Ta	0.854	0.608	1.447	1.206	3.93	2.02	1.44	0.87	0.7	0.36	1.32	0.62
Th	7.68	7.499	12.19	11.54	19.82	42.71	21.92	12.22	-	-	-	-
Pb	13.5221	15.539	22.251	19.322	19.58	18.07	22.18	18.7	25.14	25.97	25.7	27.77
U	1.965	1.295	2.516	2.674	2.94	7.81	5.21	3.94	-	-	-	-
Th/Ta	8.99297	12.334	8.4243	9.5688	5.04326	21.1436	15.2222	14.04598	-	-	-	-

Table 3 continued

Locations	Borohoro				Alataw				Nalati (Aikendaban Fm.)			
	B94 285 Ma Granodiorite	B95 266 Ma K-granite	B101 294 Ma Granodiorite	B102 280 Ma K-granite	kwsy12 298 Ma Granite	kwsy3 285 Ma Granitic dyke	kwsy27 292 Ma Granite	kwsy17 271 Ma Rhyolite	a1-39 P2 Shoshonite (?)	a1-45 P1 Trachyte andesite	a1-38 P1 Trachytic andesite	a1-4 P1 Trachyte
References	Chen et al. (2004a, b)											
Ce/Pb	7.66	6.83	1.68	3.71	3.50	2.75	4.17	4.45	1.51	2.65	6.68	1.45
Nb/La	0.16	0.16	0.44	0.35	0.53	0.74	0.38	0.35	0.49	0.35	0.32	0.67
Th/La	0.15	0.14	0.66	0.35	0.62	2.35	0.54	0.32	–	–	–	–
Locations	Southern Tianshan											
Sample no.	1											
Age	5											
Rocks	284 Ma Porphyritic granite											
References	Jiang et al. (1999)											
SiO <sub>2</sub>	59.69	70.6	74.3	75.5	75.5	75.5	75.5	75.5	51.43	52.08	49.68	71.13
TiO <sub>2</sub>	0.97	0.33	0.3	0.1	0.1	0.1	0.1	0.1	1.09	1.42	1.38	0.21
Al <sub>2</sub> O <sub>3</sub>	16.44	13.91	14.4	12.27	12.27	12.27	12.27	12.27	15.37	14.76	15.25	12.48
Fe <sub>2</sub> O <sub>3</sub>	4.2	0.81	0.72	1.74	1.74	1.74	1.74	1.74	4.35	5.16	4.79	2.36
MnO	0.04	0.05	0.01	0.06	0.06	0.06	0.06	0.06	0.13	0.16	0.18	0.09
MgO	2.41	0.46	0.41	0.18	0.18	0.18	0.18	0.18	6.35	7.25	6.86	0.58
CaO	4.71	1.52	0.99	0.3	0.3	0.3	0.3	0.3	8.52	8.58	8.46	1.13
Na <sub>2</sub> O	4.74	3.7	4	3.86	3.86	3.86	3.86	3.86	2.86	2.36	2.62	3.52
K <sub>2</sub> O	2.95	4.7	4.48	4.9	4.9	4.9	4.9	4.9	0.73	0.82	0.83	5.56
P <sub>2</sub> O <sub>5</sub>	0.47	0.13	0.05	0.06	0.06	0.06	0.06	0.06	0.23	0.18	0.24	0.08
LOI	0.92	1.3	0.35	0.24	0.24	0.24	0.24	0.24	1.04	0.86	1.13	0.85
Total	100.27	99.8	100.7	99.9	99.9	99.9	99.9	99.9	99.68	100.4	99.83	99.01
ALK	7.69	8.4	8.48	8.76	8.76	8.76	8.76	8.76	3.59	3.18	3.45	9.08
ASI	0.84	1.00	1.09	1.00	1.00	1.00	1.00	1.00	0.73	0.72	0.74	0.90
La	81.57	68.39	5.74	22.5	22.5	22.5	22.5	22.5	24.79	24.43	18.55	34.32
Ce	151.19	136.72	12.44	57.44	57.44	57.44	57.44	57.44	52.22	51.75	39.57	72.93
Pr	17.16	15.41	1.22	7.28	7.28	7.28	7.28	7.28	7.66	6.89	5.62	11.69
Nd	59	51.52	4.91	27.9	27.9	27.9	27.9	27.9	28.35	26.73	20.23	39.62
Sm	9.98	102	1.74	7.97	7.97	7.97	7.97	7.97	5.66	6.32	4.85	11.75
Eu	2.05	1.33	0.52	0.36	0.36	0.36	0.36	0.36	2.14	2.13	1.64	1.32
Gd	5.71	8.52	1.41	6.54	6.54	6.54	6.54	6.54	7.09	6.76	4.88	8.46
Tb	0.95	1.5	0.32	0.25	0.25	0.25	0.25	0.25	0.95	0.87	0.62	1.61
Locations	Bogda											
Sample no.	2961											
Age	2962											
Rocks	289 Ma Dacite dyke											
References	Shu et al. (2005)											
SiO <sub>2</sub>	59.69	70.6	74.3	75.5	75.5	75.5	75.5	75.5	51.43	52.08	49.68	71.13
TiO <sub>2</sub>	0.97	0.33	0.3	0.1	0.1	0.1	0.1	0.1	1.09	1.42	1.38	0.21
Al <sub>2</sub> O <sub>3</sub>	16.44	13.91	14.4	12.27	12.27	12.27	12.27	12.27	15.37	14.76	15.25	12.48
Fe <sub>2</sub> O <sub>3</sub>	4.2	0.81	0.72	1.74	1.74	1.74	1.74	1.74	4.35	5.16	4.79	2.36
MnO	0.04	0.05	0.01	0.06	0.06	0.06	0.06	0.06	0.13	0.16	0.18	0.09
MgO	2.41	0.46	0.41	0.18	0.18	0.18	0.18	0.18	6.35	7.25	6.86	0.58
CaO	4.71	1.52	0.99	0.3	0.3	0.3	0.3	0.3	8.52	8.58	8.46	1.13
Na <sub>2</sub> O	4.74	3.7	4	3.86	3.86	3.86	3.86	3.86	2.86	2.36	2.62	3.52
K <sub>2</sub> O	2.95	4.7	4.48	4.9	4.9	4.9	4.9	4.9	0.73	0.82	0.83	5.56
P <sub>2</sub> O <sub>5</sub>	0.47	0.13	0.05	0.06	0.06	0.06	0.06	0.06	0.23	0.18	0.24	0.08
LOI	0.92	1.3	0.35	0.24	0.24	0.24	0.24	0.24	1.04	0.86	1.13	0.85
Total	100.27	99.8	100.7	99.9	99.9	99.9	99.9	99.9	99.68	100.4	99.83	99.01
ALK	7.69	8.4	8.48	8.76	8.76	8.76	8.76	8.76	3.59	3.18	3.45	9.08
ASI	0.84	1.00	1.09	1.00	1.00	1.00	1.00	1.00	0.73	0.72	0.74	0.90
La	81.57	68.39	5.74	22.5	22.5	22.5	22.5	22.5	24.79	24.43	18.55	34.32
Ce	151.19	136.72	12.44	57.44	57.44	57.44	57.44	57.44	52.22	51.75	39.57	72.93
Pr	17.16	15.41	1.22	7.28	7.28	7.28	7.28	7.28	7.66	6.89	5.62	11.69
Nd	59	51.52	4.91	27.9	27.9	27.9	27.9	27.9	28.35	26.73	20.23	39.62
Sm	9.98	102	1.74	7.97	7.97	7.97	7.97	7.97	5.66	6.32	4.85	11.75
Eu	2.05	1.33	0.52	0.36	0.36	0.36	0.36	0.36	2.14	2.13	1.64	1.32
Gd	5.71	8.52	1.41	6.54	6.54	6.54	6.54	6.54	7.09	6.76	4.88	8.46
Tb	0.95	1.5	0.32	0.25	0.25	0.25	0.25	0.25	0.95	0.87	0.62	1.61



Table 3 continued

Locations		Southern Tianshan					Bogda				
Sample no.	1	5	16–6	29–12	2961	2962	302	2989			
Age	298 Ma	284 Ma	280–270 Ma	264 Ma	P1	P1	289 Ma	P1			
Rocks	Granodiorite	Porphyritic granite	Two-mica granite	Alkali feldspar granite	Basalt	Basalt	Dacite dyke	Rhyolite			
References	Jiang et al. (1999)				Shu et al. (2005)						
Dy	11.32	34.84	6.26	10.45	6.02	5.55	3.79	8.72			
Ho	0.72	1.7	0.5	1.5	1.49	1.27	0.84	2.12			
Er	1.87	4.48	1.51	4.07	3.38	3.15	2.39	5.73			
Tm	0.27	0.7	0.2	0.63	0.59	0.56	0.35	1.11			
Yb	1.69	4.44	1.47	4.14	3.68	3.35	2.27	5.47			
Lu	0.23	0.8	0.22	0.73	0.43	0.44	0.33	0.63			
ΣREE	343.71	432.35	38.46	151.76	144.5	140.2	105.93	205.48			
(La/Yb) <sub>N</sub>	34.62	11.05	2.80	3.90	4.8	5.2	5.9	4.5			
δEu	0.76	0.06	0.98	0.15	1.03	0.99	1.02	0.39			
V	–	–	–	–	–	–	–	–			
Cr	17.8	22.04	22.59	21.04	–	–	–	–			
Co	–	–	–	–	–	–	–	–			
Ni	5.7	8.03	5.9	6.48	–	–	–	–			
Cu	–	–	–	–	–	–	–	–			
Zn	–	–	–	–	–	–	–	–			
Ga	–	–	–	–	–	–	–	–			
Ge	–	–	–	–	–	–	–	–			
Rb	91.6	179.56	142.76	246.59	–	–	–	–			
Sr	396.7	154.42	87.67	52.8	–	–	–	–			
Y	15.28	54.64	12.43	39.18	–	–	–	–			
Zr	65.87	303.06	39.48	83.97	25.48	24.86	21.35	41.22			
Nb	14.27	20.13	8.11	23.79	–	–	–	–			
Cs	2.87	5.89	4.1	3.25	–	–	–	–			
Ba	1197	623.9	295.08	70.4	–	–	–	–			
Hf	2.11	8.23	1.5	4.18	–	–	–	–			
Ta	0.73	2.09	0.46	2.03	–	–	–	–			
Th	31.88	33.18	13.02	50.25	–	–	–	–			
Pb	34.48	40.6	40.74	38.32	–	–	–	–			
U	3.45	3.73	1.81	8.13	–	–	–	–			
Th/Ta	43.671	15.8755981	28.30434783	24.7537	–	–	–	–			
Ce/Pb	4.38	3.37	0.31	1.50	–	–	–	–			

Table 3 continued

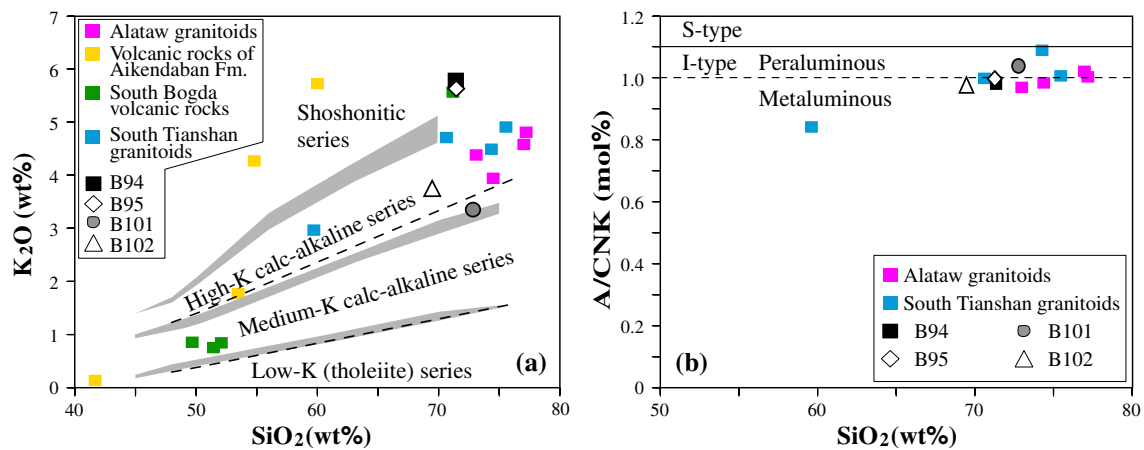
Locations		Southern Tianshan				Bogda			
Sample no.	1	5	16-6	29-12	2961	2962	302	2989	
Age	298 Ma	284 Ma	280–270 Ma	264 Ma	P1	P1	289 Ma	P1	
Rocks	Granodiorite	Porphyritic granite	Two-mica granite	Alkali feldspar granite	Basalt	Basalt	Dacite dyke	Rhyolite	
References	Jiang et al. (1999)				Shu et al. (2005)				
Nb/La	0.17	0.29	1.41	1.06	–	–	–	–	
Th/La	0.39	0.49	2.27	2.23	–	–	–	–	

ALK (all alkaline) =  $\text{Na}_2\text{O} + \text{K}_2\text{O}$ , ASI (alumina saturation index) = molar  $\text{Al}_2\text{O}_3/(\text{Na}_2\text{O} + \text{K}_2\text{O} + \text{CaO})$ ,  $\Sigma\text{REE}$  total rare earth elements,  $\delta\text{Eu}(\text{Eu anomaly}) = \text{Eu}_N \times 2/(\text{Sm}_N + \text{Gd}_N)$ , subscript  $N$  refers to chondrite normalized value

to the low Ce/Pb and Nb/La ratios (Table 3) indicating the influence of subduction-derived fluids on the sub-continental mantle wedge (Hofmann et al. 1986); moreover, the high Th/La ratios suggest influence of subducted sediments (Plank 2005). It is also the cases for the volcanic rocks of Aikedaban Fm., for Alataw granitoids except the granitic dyke, and for the granodiorite and porphyritic granite of the southern Tianshan. However, it seems not necessary to conclude that these magmatic rocks were produced in an active margin since subduction components may be inherited from the subcontinental mantle lithosphere in a post-collisional setting. Thus, the high-K calc-alkaline and transitional granites of Borohoro plutons show a compositional signature similar to either mature continental arcs or post-collisional settings, but limited geochemical arguments are not conclusive to discriminate between an active margin and a post-collisional environment for these magmatic rocks.

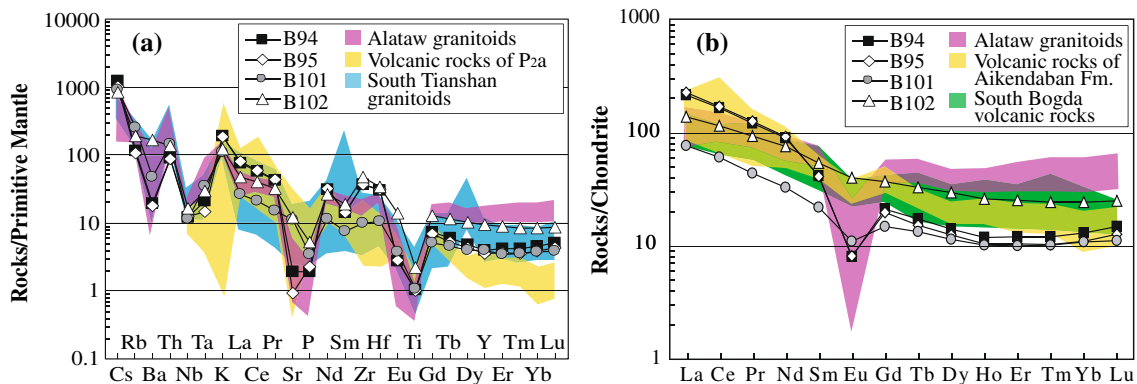
The Permian magmatic rocks are also considered to form during a continental rifting. However, In Awulale area, Permian volcanic and intrusive rocks as well as red molasses deposits are well developed, our structural investigation reveals that neither normal faults nor strike-slip faults predominate. Otherwise, considering the extensive distribution of Permian and Carboniferous magmatic rocks in northern Chinese Tianshan, the hypothesis of Large Igneous Province was proposed by Xia et al. (2004a, 2006) and correlated with a mantle superplume. Using geochemical and isotopic data, these authors suggest an asthenospheric magma source similar to that of Hawaii and Iceland. Nevertheless, the diversity of magma sources, the predominance of Carboniferous calc-alkaline magmas and the presence of ophiolitic mélanges do not fit this interpretation. In addition, the extremely long duration of magmatic activity (>110 Ma from 360 to 250 Ma) is rather unusual for mantle plumes over the world that generally result in continental break-up within less than 20–30 Ma. Moreover, no other criteria for identifying ancient mantle plume, such as crustal uplift prior to volcanism, radial dyke swarms, volcanic chain or thick plateau basalt (e.g. Campbell 2001), are documented in Tianshan, or even in Central Asia. Thus, on the basis of field observation, lithological features, geochemical and geochronological data, another interpretation may be proposed for the generation of Permian magmatism in Chinese Tianshan.

The Late Carboniferous to Permian magmatism of the northwestern Chinese Tianshan displays a compositional variation from highly fractionated calc-alkaline to alkaline, indicating the contribution of multiple magma sources in a transitional tectonic regime (Black and Liégeois 1993; Liégeois et al. 1998; Bonin 2004). This feature often occurs in large-scale shear zones in a post-collisional setting where the previously metasomatised mantle wedge,



**Fig. 8** **a** Classification of Permian magmatic rocks of the West Tianshan in the potassium–silica diagram after Rickwood (1989); **b** compositional plots of the Permian granitic rocks of the West

Tianshan in the A/CNK (molecular ratio  $Al_2O_3/[CaO + Na_2O + K_2O]$ ) versus Silica diagram. See Table 3 for the references



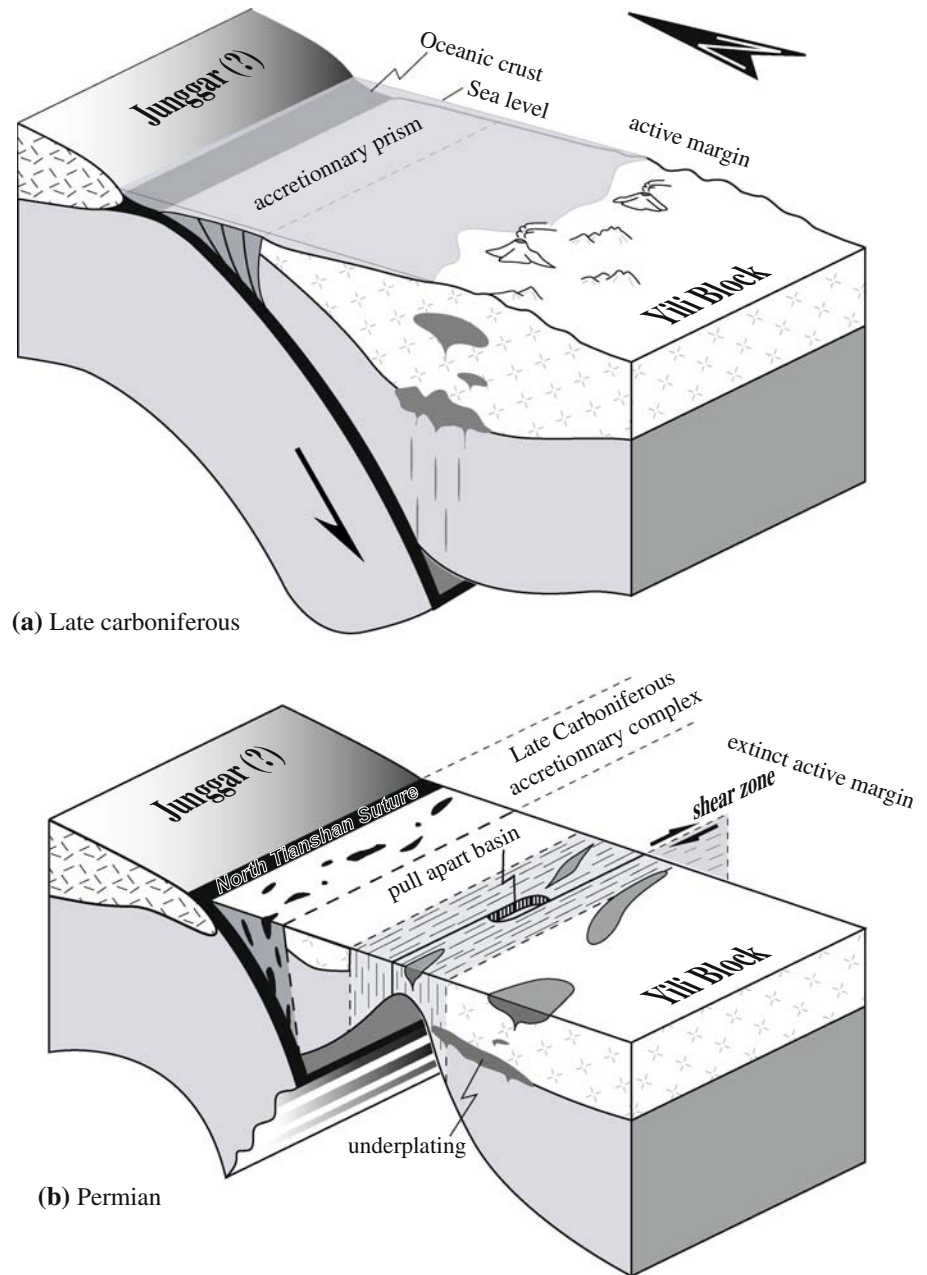
**Fig. 9** **a** Primitive mantle normalized trace elements abundances and **b** Chondrite normalized rare earth elements patterns for the granitoids from the Borohoro plutons, northwestern Chinese Tianshan. Normalized values are after Sun and McDonough (1989). See Table 3 for references

underplated mafic material, and uplifted asthenosphere, are possible sources for magmas in a high regional heat flow (Cluzel et al. 1991, 1994; D'lemos et al. 1992; Hutton and Reavy 1992; Tommasi et al. 1994; Saint Blanquat et al. 1998; Steenken et al. 2000; Weinberg et al. 2004; Rosenberg 2004; Njanko et al. 2006; Oyhantçabal et al. 2007). According to available geological maps and our field observations, the Borohoro plutons and granites of Kekesu and Gangou areas were emplaced within or at the margins of transcurrent shear zones, some have an “en cornue” shape (Figs. 1, 3, 5, 6) developing high-temperature fabrics, and therefore may be considered as synkinematic. The North Tianshan Fault, the Qingbulak-Nalati Fault and the Sangshuyuanzi Fault are major ductile dextral strike-slip faults (Allen et al. 1995; Allen and Vincent 1997; Wang et al. 2006a) that crosscut the lithosphere according to seismic data (Zhao et al. 2003a; Wang et al. 2004). The cooling age of these shear zones is constrained between 285 and 245 Ma on the basis of  $^{40}Ar/^{39}Ar$  dating of biotite from mylonites (Yin and Nie 1996; Zhou et al. 2001;

Laurent-Charvet et al. 2003; de Jong et al. 2008, this volume), confirming a close temporal relationship between pluton intrusion and ductile wrench tectonics. Therefore, Permian large-scale wrench faulting likely induced and controlled the emplacement of synchronous magma with diverse sources. This interpretation is supported by numerical modeling by Weinberg et al. (2004), which reveals that some fragile, dilatational spaces with low mean pressure can be produced during crustal-scale transcurrent shearing because of the changes in the rheological properties of rocks, favouring magmas ascent and pluton emplacement. The prominent post-collisional fluid activity and gold mineralization associated with transcurrent shear zones in Chinese Tianshan and Central Asia as a whole (Mao et al. 2004; Yakubchuk 2004; Han et al. 2006; Xiao et al. 2006) are also consistent with syn-kinematic magmatism.

Post-collisional evolution from calc-alkaline to alkaline magmatism usually occurs during a gradual transition between syn-orogenic convergence and post-orogenic

**Fig. 10** Schematic model showing **a** Late Carboniferous subduction and active continental margin; **b** Permian post-orogenic dextral shearing, uplifted asthenosphere and thinned lithosphere, where multi-originated magmatic rocks and pull-apart basin formed



intra-continental extension as observed in western Mediterranean area (Coulon et al. 2002; Duggen et al. 2005; Lustrino et al. 2007), Pan-African Central Cameroon and Uruguay (Njanko et al. 2006; Oyhantçabal et al. 2007). This transition may be brought on after continental collision or subduction extinction. Taking into account all the available geological data (e.g. Windley et al. 1990; Allen et al. 1993; Gao et al. 1998; Xiao et al. 2004; Charvet et al. 2007), kinematic observations (Laurent-Charvet et al. 2002, 2003; Wang et al. 2006a) and paleomagnetic evidences (Van der Voo et al. 2006; Wang et al. 2007b), the Carboniferous convergence was

followed by the Permian large-scale post-orogenic dextral transcurrent tectonics. We propose that the transcurrent tectonics at least partly induced the synkinematic magmatism and provoked the Early-Middle Permian transition from high-K calc-alkaline to alkaline magmatism (Fig. 10). The occurrence of Middle to Late Permian thick red molasse (XBGMR 1993; Cui et al. 1996; Carroll et al. 1995) is a rather strong evidence for post-orogenic setting; in contrast, the localised occurrence of Permian deep seated sediments (turbidites) closely associated with alkaline mafic rocks (e.g. southern Bogda area) infers that Permian strike-slip tectonics generated narrow zones of



thinned lithosphere (pull-apart basins) in which volcanic rocks were generated by the melting of uplifted asthenosphere. Such lithospheric thinning could be correlated to an east-west extension as proposed by Wartes et al. (2002). It is worth noting that the thinned area partly involved the Carboniferous mantle wedge. In this case, previously metasomatised mantle rocks coexist with uplifted juvenile asthenosphere, and source or magma mixing may easily occur. Within the Permian shear zones, source mixing is likely to be lesser than on the boundaries and “real” alkaline magmas may be generated; in contrast, away from the Permian shear zones, calc-alkaline magmatism (the so-called post-collisional plutons) may persist in the absence of juvenile asthenospheric mantle (Fig. 10b). However, this proposition should be convinced by further studies.

## Conclusions

Geochemical and geochronological data indicate that Borohoro high-K calc-alkaline granites formed between 294 and 280 Ma, while transitional and alkaline granites of Borohoro plutons, Kekesu K-granite and Gangou K-granite were produced during 285–252 Ma. Considering the present study and previous data, it appears that igneous activity during Carboniferous to Permian in western Chinese Tianshan was a permanent process showing an evolution from calc-alkaline to alkaline, the transition occurred during the Early-Middle Permian.

Early to Middle Permian magmatic association of high-K calc-alkaline, transitional and alkaline granites suggests a post-collisional setting. Previously proposed continental rift and supra mantle plume are unlikely for the genesis of these magmatic rocks because the lack of field evidence other than geochemistry. Conversely, distinctive spatial and temporal relationships between pluton emplacement and dextral strike-slip faulting are observed. Although additional detailed studies are needed, it may be proposed that the Early-Middle Permian magmatism, at least in part, occurred during regional-scale transcurrent tectonics, as a transitional event from Carboniferous convergence to Middle-Late Permian anorogenic environment.

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