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## Environmental controls on carbon dioxide flux from black spruce coarse woody debris

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**Abstract** Carbon dioxide flux from coarse woody debris (CWD) is an important source of CO<sub>2</sub> in forests with moderate to large amounts of CWD. A process-based understanding of environmental controls on CWD CO<sub>2</sub> flux ( $R_{CWD}$ ) is needed to accurately model carbon exchange between forests and the atmosphere. The objectives of this study were to: (1) use a laboratory incubation factorial experiment to quantify the effect of temperature ( $T_{CWD}$ ), water content ( $W_C$ ), decay status, and their interactions on  $R_{CWD}$  for black spruce [*Picea mariana* (Mill.) BSP] CWD; (2) measure and model spatial and temporal dynamics in  $T_{CWD}$  for a boreal black spruce fire chronosequence; and (3) validate the  $R_{CWD}$  model with field measurements, and quantify potential errors in estimating annual  $R_{CWD}$  from this model on various time steps. The  $R_{CWD}$  was positively correlated to  $T_{CWD}$  ( $R^2=0.37$ ,  $P<0.001$ ) and  $W_C$  ( $R^2=0.18$ ,  $P<0.001$ ), and an empirical  $R_{CWD}$  polynomial model that included  $T_{CWD}$  and  $W_C$  interactions explained 74% of the observed variation of  $R_{CWD}$ . The  $R_{CWD}$  estimates from the  $R_{CWD}$  model excellently matched the field measurements. Decay status of CWD significantly ( $P<0.001$ ) affected  $R_{CWD}$ . The temperature coefficient ( $Q_{10}$ ) averaged 2.5, but varied by 141% across the 5–42°C temperature range, illustrating the potential shortcomings of using a constant  $Q_{10}$ . The CWD temperature was positively correlated to air temperature ( $R^2=0.79$ ,  $P<0.001$ ), with a hysteresis effect that was correlated to CWD decay status and stand leaf area index. Ignoring this temperature hysteresis introduced errors of –1% to +32% in annual  $R_{CWD}$  estimates. Increasing  $T_{CWD}$  modeling time step from hourly to daily or monthly introduced a 5–11% underestimate in annual  $R_{CWD}$ . The annual  $R_{CWD}$  values in this study were more than two-fold greater than those in a previous study, illustrating the need to incorporate spatial and temporal responses of  $R_{CWD}$  to temperature and water

content into models for long-term  $R_{CWD}$  estimation in boreal forest ecosystems.

**Keywords** Carbon dioxide flux · Coarse woody debris · Boreal forest · Decomposition · Modeling

### Introduction

Boreal forests play a crucial role in the global carbon budget (Kasischke 2000). Wildfires, the dominant disturbance in many boreal forests (Goldammer and Furyaev 1997), produce a large amount of coarse woody debris (CWD) (Bond-Lamberty et al. 2002a), which significantly influences ecosystem structure and function (Harmon et al. 1986). CWD biomass has been quantified for numerous forest ecosystems (Frangi et al. 1997; Harmon et al. 1986; Krankina and Harmon 1995; Siitonen et al. 2000; Spies et al. 1988; Stewart et al. 1994; Stone et al. 1998), but direct measurements of CWD CO<sub>2</sub> flux ( $R_{CWD}$ ) are limited (Chambers et al. 2001; Marra and Edmonds 1994; Yoneda 1975), especially for boreal forests (Bond-Lamberty et al. 2002a).

Bond-Lamberty et al. (2002a) estimated that the  $R_{CWD}$  comprised 1–54% of soil surface CO<sub>2</sub> flux for seven different-aged boreal black spruce stands comprising a fire chronosequence, with the maximum  $R_{CWD}$  occurring in a 12-year-old stand that contained the greatest CWD mass. They used an assumed  $Q_{10}$  to calculate annual  $R_{CWD}$ , but  $Q_{10}$  changes with temperature and water content ( $W_C$ ) for soil surface CO<sub>2</sub> flux (Raich and Schlesinger 1992; Rayment and Jarvis 2000; Strömrgren 2001; Wang et al. 2002) and decaying roots (Chen et al. 2000). The temperature dependence of  $Q_{10}$  for decaying CWD is not well studied. Bond-Lamberty et al. (2002a) also reported that  $R_{CWD}$  was moisture-limited below 43%  $W_C$  based on field measurements. Clearly, a controlled laboratory experiment to quantify the effects of temperature,  $W_C$ , and their interaction on  $R_{CWD}$  is needed.

Empirical models driven by environmental variables are commonly used to calculate annual  $R_{CWD}$  (Bond-Lamberty et al. 2002a). Coarse woody debris tempera-

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ture ( $T_{CWD}$ ) is not routinely measured, and therefore is often assumed to follow air temperature ( $T_A$ ) (Bond-Lamberty et al. 2002a). The hysteresis effects of temperature and  $W_C$  on respiration, resulting from substrate thermal and hydrological conductivity and delayed biological response to environmental factors, are generally ignored (Bond-Lamberty et al. 2002a; Potter et al. 1993; Raich et al. 1991). However, the physical properties of wood might be critical to  $R_{CWD}$  because of its poor thermal conductivity (Campbell and Norman 1998). Furthermore, soil surface  $CO_2$  flux displays nonlinear responses to environmental variables (Rayment and Jarvis 2000; Strömberg 2001; Wang et al. 2002). Presumably,  $R_{CWD}$ , also a product of microbial activity, behaves similarly. Therefore, using different modeling time steps may introduce bias in annual  $R_{CWD}$  estimates. Bond-Lamberty et al. (2002a) assumed that  $T_{CWD}$  followed  $T_A$  on a weekly time step, which has not been tested.

The objectives of this study were to: (1) quantify the effect of  $T_{CWD}$ ,  $W_C$ , decay status, and their interactions on  $R_{CWD}$  for black spruce CWD; (2) measure and model spatial and temporal dynamics in  $T_{CWD}$  for a boreal black spruce fire chronosequence; and (3) validate the  $R_{CWD}$  model and quantify potential errors in estimating annual  $R_{CWD}$ .

## Materials and methods

### Site description and general experiment design

The CWD samples for incubation and temperature measurements were collected from a boreal black spruce fire chronosequence near Thompson, Manitoba (55°48' N 97°52' W). The chronosequence consisted of even-aged stands that originated from stand-killing wildfires in 1989, 1981, 1964, and 1870, containing 177.5±48.1, 58.2±21.6, 13.1±7.7, and 10.5±8.6 ton CWD ha<sup>-1</sup>, respectively (Bond-Lamberty et al. 2002a). The stand ages were 12, 20, 37, and 131 years, respectively, at the time of investigation. All the stands were within 40 km<sup>2</sup>. The stands were dominated by black spruce prior to wildfires, and had well-drained clay soils. Gower et al. (2002) provided detailed descriptions of the stands.

We addressed our objectives using a combination of measurements and modeling.

1. A laboratory factorial incubation experiment was used to quantify the effect of  $T_{CWD}$ ,  $W_C$ , decay status, and their interactions on  $R_{CWD}$  for black spruce CWD.
2. In situ  $T_{CWD}$  and  $T_A$  were measured throughout the growing season to examine spatial and temporal  $T_{CWD}$  dynamics in the chronosequence.
3. The empirical  $R_{CWD}$  model developed from laboratory measurements was compared to field data.

4. The  $R_{CWD}$  model and  $T_{CWD}$  data were applied with various time steps to quantify potential errors in annual  $R_{CWD}$  estimates.

### CWD incubation experiment

CWD samples for the incubation experiment were collected from the 12-year-old stand in November 2000. The stand was composed of trembling aspen (*Populus tremuloides* Michx.), paper birch (*Betula papyrifera* Marsh), jack pine (*Pinus banksiana* Lamb.), black spruce, and willow (*Salix* spp.) seedlings and saplings.

The experimental design of the incubation study was a two-factorial repeated measure design with seven temperatures and three CWD decay classes. The temperatures, selected to include a wide range of temperature between 0 and 45°C, were 5, 12, 17, 20, 26, 32, and 42°C. The decay classes were: class I – a knife could not penetrate into the sample; class II – a knife could slightly penetrate the sample with appreciable resistance; and class III – a knife could fully penetrate the sample with little resistance (Lambert et al. 1980).

A total of 36 samples, 12 from each decay class, were randomly selected in the field and carefully transported to University of Wisconsin – Madison. Each CWD sample was tagged and randomly assigned to one of two Conviron CMP3023 plant growth chambers (Conviron, Winnipeg, MB, Canada), with 6 of the 12 samples from each decay class in each growth chamber. Samples were exposed to 12:12 h of light and darkness each day. Each growth chamber was set to a treatment temperature, and the CWD samples were placed in the growth chamber for 48 h to equilibrate to the new temperature. Each sample was submerged in distilled water of the same treatment temperature for 48 h to achieve fiber saturation. They were removed from the water and incubated between 10 and 30 days, depending on the temperature, to produce a natural  $W_C$  gradient. The  $R_{CWD}$  measurements were started after the samples were removed from the water for 12–24 h, to minimize potential effect of CWD rewetting. The  $R_{CWD}$  was measured three to seven times during the drying period to obtain  $R_{CWD}$  at various  $W_C$  levels. At the end of the drying period a new treatment temperature was assigned to the growth chamber, and the measurement procedure was repeated.

To measure  $R_{CWD}$ , a CWD sample was removed from the growth chamber and immediately inserted into a custom chamber connected to a Li-Cor 6200 infrared gas analyzer (IRGA) (Li-Cor, Lincoln, Neb.). The IRGA chamber consisted of a clear acrylic tube 125 cm in length by 15 cm in diameter, sealed at both ends. Two small internal fans were installed to ensure uniform  $CO_2$  concentration inside the chamber. A plastic sampling tube was connected to the IRGA, and spiraled the entire length of the chamber along the interior wall, ensuring that a representative air sample was obtained. The  $CO_2$  concentration inside the chamber was scrubbed down to about 50  $\mu\text{mol mol}^{-1}$  below ambient, and  $R_{CWD}$  was measured as the  $CO_2$  concentration rose back through ambient. The IRGA configuration used was: total system volume 22,000 cm<sup>3</sup>, flow rate 1,100 cm<sup>3</sup> s<sup>-1</sup>, one measurement every 30 s, five measurements in total. Ambient temperature,  $CO_2$  concentration, and relative humidity were recorded at the time of each measurement. CWD samples were weighed ( $\pm 0.1$  g) at the completion of each  $R_{CWD}$  measurement, to calculate  $W_C$  of each CWD sample, and returned to the appropriate growth chamber. At the end of the

**Table 1** Physical properties of the black spruce coarse woody debris (CWD) for the incubation study. ( $N$  sample size,  $SL_{CWD}$  specific CWD area,  $\bar{x}$  mean)

Decay class	N	Diameter (cm)		Length (cm)		Dry mass (g)		Density (g cm <sup>-3</sup> )		SL <sub>CWD</sub> (cm g <sup>-1</sup> )	
		$\bar{x}$	SD	$\bar{x}$	SD	$\bar{x}$	SD	$\bar{x}$	SD	$\bar{x}$	SD
I	12	8.5	2.3	70.3	7.8	1,676	792	0.41	0.06	1.35	0.45
II	12	7.8	1.2	66.3	10.4	1,084	400	0.34	0.05	1.69	0.36
III	12	8.7	1.1	62.6	9.3	1,044	327	0.28	0.05	1.83	0.35

incubation experiment, each CWD sample was oven-dried at 70°C to a constant mass and weighed ( $\pm 0.1$  g). The sample length and diameters in two perpendicular directions at both ends were measured. We calculated the surface area, volume, density, and specific CWD area ( $SL_{CWD}$ , defined as the CWD surface area to mass ratio) by assuming each sample was a cylinder (Table 1).

Statistical models relating  $R_{CWD}$  to  $T_{CWD}$ ,  $W_C$ , and  $SL_{CWD}$  were developed using SAS software version 8.2 (SAS 2001). The mixed effect procedure (PROC MIXED) was used for the repeated measure analysis, assigning  $SL_{CWD}$ ,  $T_{CWD}$ ,  $W_C$ , and all possible interactions as fixed effects. CWD samples within each  $T_{CWD}$  and  $W_C$  combination were treated as a random effect, and used to test for autocorrelation between measurements. A spatial power covariance matrix was chosen for autocorrelation based on SBC (Schwarz's Bayesian Criterion). Natural logarithm transformation of  $R_{CWD}$  was used to achieve homoscedasticity. A backward elimination procedure was performed to remove insignificant terms ( $\alpha=0.05$ ) in the models. An analysis of covariance (ANCOVA) was conducted to test the effect of decay class on  $R_{CWD}$  by setting  $T_{CWD}$  and  $W_C$  as covariates.

The  $Q_{10}$  was calculated as

$$Q_{10} = (R_{CWD2}/R_{CWD1})^{10/(T_2-T_1)} \quad (1)$$

where  $R_{CWD1}$  and  $R_{CWD2}$  were the  $R_{CWD}$  at temperatures  $T_1$  and  $T_2$ . To explore the relationships among  $Q_{10}$ ,  $T_{CWD}$ , and  $W_C$ , we arbitrarily binned  $W_C$  to the nearest 10%. We calculated  $Q_{10}$  within temperature ranges of 5–20, 20–32, and 32–42°C, and each  $W_C$  level. The  $Q_{10}$  values for the temperature ranges of 20–32 and 32–42°C were statistically combined ( $df=49$ ,  $t=-1.21$ ,  $P=0.234$ ). Analysis of variance (ANOVA) was conducted to test the effects of decay class,  $T_{CWD}$ , and  $W_C$  on  $Q_{10}$ .

#### In situ $T_{CWD}$

The  $T_{CWD}$  was measured continuously in the field at the four stands using type T (copper-constantine) thermocouples connected to Campbell Scientific CR 10X dataloggers (Campbell Scientific, Logan, Utah, USA). Six CWD samples (>5 cm in middle length diameter), two for each decay class, were randomly selected at each stand. Thermocouples were installed in a 2-cm-deep hole in the top of each CWD sample, and the hole was sealed with silicon adhesive. At the 12-year-old stand, we examined the heterogeneity of  $T_{CWD}$  within CWD. Thermocouples were installed in the top and bottom of four CWD samples in each decay class.  $T_A$  at each site was measured at a nearby (<20 m) meteorological station (Gower et al. 2002). Temperatures were measured every minute, and averaged for each half-hour period, from late May through early October 2001.

We examined  $T_{CWD}$  diurnal patterns based on the half-hour interval measurements at the four stands. A paired  $t$ -test was used to test the differences in  $T_{CWD}$  temporal patterns among stands and decay classes. The seasonal differences among stands, decay classes, and heterogeneous positions were tested on daily, weekly, and monthly time steps, respectively. A PROC MIXED procedure was used to develop empirical models of  $T_{CWD}$  against  $T_A$  on a daily time step. A first order autoregressive model was required ( $z=-5.97$ ,  $P<0.001$ ) for autocorrelation in time series temperature measurements.

#### Potential errors in modeling annual $R_{CWD}$

The laboratory empirical  $R_{CWD}$  model was compared to the field measurements of  $R_{CWD}$ ,  $T_{CWD}$ ,  $W_C$ , and CWD physical properties (Bond-Lamberty et al. 2002a). We did not measure  $W_C$  in situ, but, for comparison, employed the  $W_C$  models at the same sites by Bond-Lamberty et al. (2002a), i.e.  $W_C=4.6244-0.3083 T+0.0054 T^2$ ,  $W_C=13.119-0.8816 T+0.0155 T^2$ , and  $W_C=29.062-1.9911 T+0.0349 T^2$  for CWD decay I, II, and III, respectively; where  $T$  was the week of year.

We focused on two sources of error in annual  $R_{CWD}$  estimates: (1) ignoring  $T_{CWD}$  hysteresis to  $T_A$ ; and (2) using various  $T_{CWD}$  modeling time steps. The  $R_{CWD}$  was calculated from the  $R_{CWD}$  model driven by  $W_C$  and  $T_{CWD}$  or  $T_A$ . The  $R_{CWD}$  estimates from  $T_{CWD}$  and  $T_A$  were denoted as  $R_{TCWDijk}$  and  $R_{TAijk}$ , where  $i=1, 2, 3$  for decay classes;  $j=1, 2, 3, 4$  for stands; and  $k=8,760, 365, 52$ , and 12 for hours, days, weeks, and months per year, respectively. The sources of error were evaluated as follows:

Error in annual  $R_{CWD}$  estimates caused by ignoring  $T_{CWD}$  hysteresis to  $T_A$  at  $j$  stand ( $E_{Hj}$ ) was calculated as:

$$E_{Hj} = \left\{ \sum_{i=1}^3 \left[ \left( \sum_{k=1}^{8,760} R_{TCWDijk} - \sum_{k=1}^{8,760} R_{TAijk} \right) / \sum_{k=1}^{8,760} R_{TCWDijk} \right] \right\} / 3 \quad (2)$$

Errors in annual  $R_{CWD}$  due to different modeling time steps were calculated as: (1) hourly versus daily:  $R_{Th}-R_{Td}/R_{Th}$ , (2) hourly versus weekly:  $R_{Th}-R_{Tw}/R_{Th}$ , (3) hourly versus monthly:  $R_{Th}-R_{Tm}/R_{Th}$ , where  $R_{Th}$ ,  $R_{Td}$ ,  $R_{Tw}$ , and  $R_{Tm}$  were annual  $R_{CWD}$  means of the three decay classes in the four stands estimated on hourly, daily, weekly, and monthly time steps, respectively.

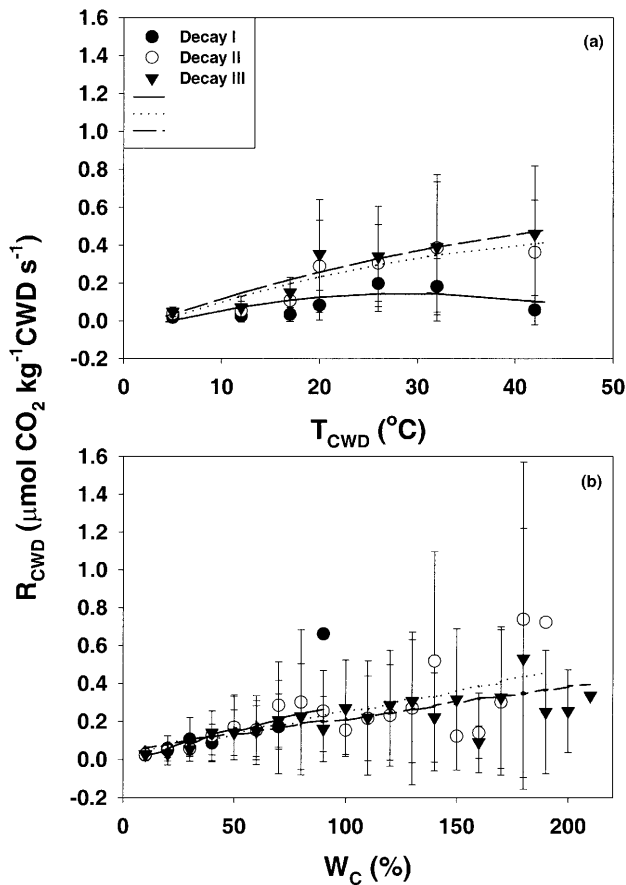
## Results and discussion

### Environmental and substrate controls on $R_{CWD}$ in boreal forests

The  $R_{CWD}$  was positively correlated to  $T_{CWD}$  ( $n=636$ ,  $R^2=0.37$ ,  $P<0.001$ ) and  $W_C$  ( $R^2=0.18$ ,  $P<0.001$ ) (Fig. 1). A polynomial model that included  $T_{CWD}$  and  $W_C$  interactions explained 74% ( $P<0.001$ ) of the observed variation in  $R_{CWD}$  in the laboratory (Table 2).

CWD decay class significantly affected  $R_{CWD}$  ( $df=3$ ,  $615$ ,  $F=271.3$ ,  $P<0.001$ ). The pairwise comparisons indicated that the  $R_{CWD}$  for decay I was significantly less than that for decay II ( $df=615$ ,  $t=-3.70$ ,  $P<0.001$ ) and decay III ( $t=-3.64$ ,  $P<0.001$ ), and the  $R_{CWD}$  for decay II and III did not differ ( $t=-0.41$ ,  $P=0.685$ ). However, the relationship between  $R_{CWD}$  and  $W_C$  did not differ significantly ( $P=0.179$  for the intercepts and  $P=0.389$  for the slopes) among the decay classes (Fig. 1b).  $SL_{CWD}$  was 1.83, 1.69, and 1.35  $cm^2g^{-1}$  for decay class I, II, and III, respectively. The  $SL_{CWD}$  for decay I was significantly less than for decay II ( $P=0.037$ ) and III ( $P=0.004$ ) (Fig. 2a). The  $R_{CWD}$  increased with increasing  $SL_{CWD}$  (Fig. 2b).

Temperature and  $W_C$  are two major environmental factors that influence microbial activity and decomposition rates of organic matter (Boddy 1983a; Chen et al. 2000; Flanagan and Veum 1974; Moore 1986; O'Connell 1990; Wang et al. 2002). Temperature is commonly the more important environmental factor influencing  $R_{CWD}$ , but  $W_C$  strongly interacted with  $T_{CWD}$  on  $R_{CWD}$  across a broad  $W_C$  gradient (from 10 to 190%) (Table 2, Fig. 1b). The  $R_{CWD}$  variability increased with either  $T_{CWD}$  or  $W_C$  until a certain level, and then tended to decline (Fig. 1), which suggested that either temperature or  $W_C$  could become the dominant environmental control on  $R_{CWD}$  in extreme conditions. This result is corroborated by previous studies (Boddy 1983a; Bond-Lamberty et al. 2002a;



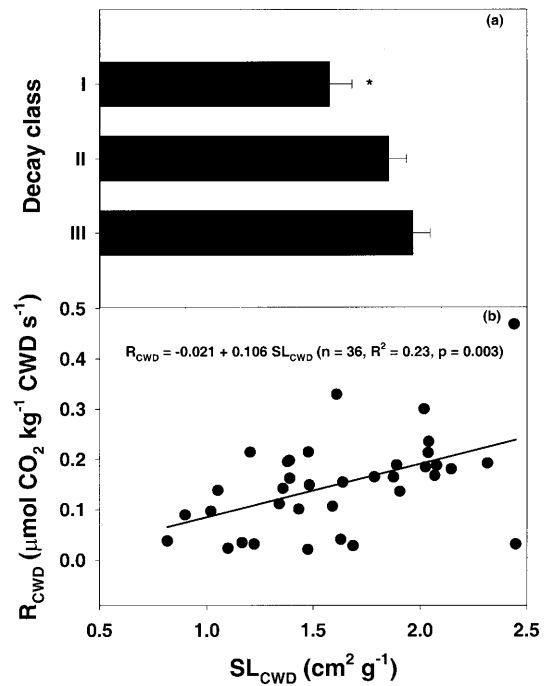
**Fig. 1** Relationships between coarse woody debris (CWD) CO<sub>2</sub> flux ( $R_{CWD}$ ,  $\mu\text{mol CO}_2 \text{ kg}^{-1} \text{ CWD s}^{-1}$ ) and (a) temperature ( $T_{CWD}$ , °C), and (b) water content ( $W_C$ , %). The bars are standard deviations ( $n=2-88$ )

**Table 2** Model relating CO<sub>2</sub> flux ( $R_{CWD}$ ,  $\mu\text{mol CO}_2 \text{ kg}^{-1} \text{ CWD s}^{-1}$ ) to temperature ( $T_{CWD}$ , °C), water content ( $W_C$ ), and specific CWD area ( $SL_{CWD}$ ,  $\text{cm}^2 \text{ g}^{-1}$ ) for black spruce CWD. The model form is:  $\ln(R_{CWD}) = a SL_{CWD} + b SL_{CWD} T_{CWD} + c T_{CWD}^2 + d W_C + e T_{CWD} W_C$  ( $n=636$ ,  $R^2=0.73$ )

Coefficient	Estimate	SE	<i>t</i> -ratio	<i>P</i>
a	-3.272	0.117	-28.07	<0.001
b	0.097	0.008	12.54	<0.001
c	-0.003	0.000	-12.29	<0.001
d	1.214	0.198	6.13	<0.001
e	0.036	0.009	4.03	<0.001

Harmon et al. 1986; Käärrik 1974; Progar et al. 2000; Rayner and Boddy 1988; Yoneda 1980).

$SL_{CWD}$  was a good indicator of CWD decay status, and hence  $R_{CWD}$  (Fig. 2b).  $SL_{CWD}$  increased as decomposition proceeded because of carbon losses through CO<sub>2</sub> fluxes and dissolved organic carbon (Yavitt and Fahey 1985), and increased fragmentation (Harmon et al. 1986) (Fig. 2a). Increased specific area increased water holding capacity and water potential because of greater internal surface and pore space (Boddy 1983b; Bond-Lamberty et al. 2002a; Chambers et al. 2001; Harmon and Sexton



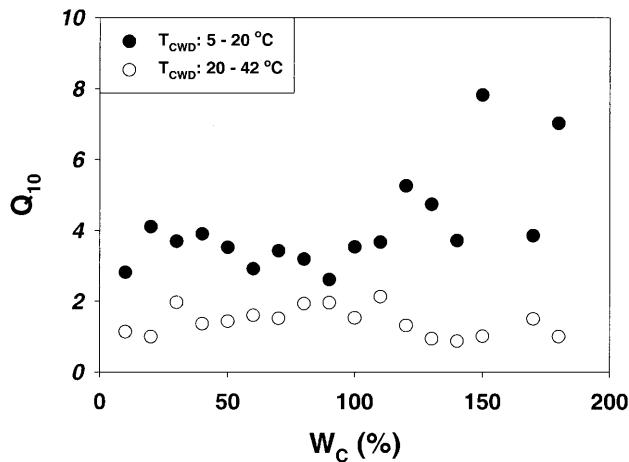
**Fig. 2** (a) Specific CWD area ( $SL_{CWD}$ ,  $\text{cm}^2 \text{ g}^{-1}$ ) by decay class. An asterisk denotes significant difference ( $\alpha=0.05$ ). Error bars are standard deviations ( $n=12$ ). (b) Relationship between CWD CO<sub>2</sub> flux ( $R_{CWD}$ ,  $\mu\text{mol CO}_2 \text{ kg}^{-1} \text{ CWD s}^{-1}$ ) and  $SL_{CWD}$

1995). We speculate that the lower water holding capacity of decay class I CWD may restrict the  $R_{CWD}$ , especially when  $T_{CWD}$  and  $W_C$  are high (Fig. 1a).

#### $R_{CWD}$ model

A polynomial model including  $SL_{CWD}$ ,  $T_{CWD}$ ,  $W_C$ , and their interactions provided a good fit for the laboratory  $R_{CWD}$  measurements ( $n=636$ ,  $R^2=0.73$ ) (Table 2). The  $Q_{10}$  for  $R_{CWD}$  averaged  $2.5 \pm 1.6$  across the environmental conditions and decay classes examined in this study.  $Q_{10}$  was significantly affected by temperature range ( $df=1$ ,  $36$ ,  $F=57.2$ ,  $P<0.001$ ) and averaged 4.1 for 5–20°C and 1.7 for 20–42°C. CWD decay class ( $df=2$ ,  $36$ ,  $F=1.5$ ,  $P=0.249$ ) and  $W_C$  gradient ( $df=16$ ,  $36$ ,  $F=0.8$ ,  $P=0.729$ ) did not significantly affect  $Q_{10}$  (Fig. 3).

Carbon dioxide flux from organic matter is commonly modeled using a simplified  $Q_{10}$  function (Lloyd and Taylor 1994) or using an environment-driven model (Raich and Schlesinger 1992). The  $Q_{10}$  model oversimplifies the respiration process by assuming a constant response of respiration rate to temperature within a certain temperature range, and thus is easy for general application. However, a single  $Q_{10}$  is not appropriate to quantify  $R_{CWD}$  because  $Q_{10}$  changes with temperature (Fig. 3). Such a  $Q_{10}$  temperature dependence has also been reported for soil surface CO<sub>2</sub> flux (Nadelhoffer et al. 1991; Rayment and Jarvis 2000; Stromgren 2001; Wang et al. 2002) and decaying root respiration (Chen et al. 2000). It is important



**Fig. 3** Temperature coefficient ( $Q_{10}$ ) dynamics along water content ( $W_C$ , %) gradient for two temperature ( $T_{CWD}$ , °C) ranges

to notice that using variable  $Q_{10}$  values for various temperature ranges also failed to characterize the significant effect of CWD decay status on  $R_{CWD}$  (Table 2) and introduced appreciable errors in annual  $R_{CWD}$  estimates (discussed below).

Accurately modeling  $R_{CWD}$  requires a process-based understanding of the factors that influence CWD decomposition, and continuous data for each variable. In tropical regions, where temperature is favorable for microbial activity year-round, substrate  $W_C$  is the major environmental factor influencing  $R_{CWD}$  (Chambers et al. 2001). In boreal forests, temperature restricts  $R_{CWD}$  in winter, and the limiting effect of  $W_C$  on  $R_{CWD}$  increases with increasing temperature in the growing season (Bond-Lamberty et al. 2002a). These data suggest that a comprehensive model should be used to simulate  $R_{CWD}$  in boreal forest ecosystems, a finding corroborated by decomposition studies in various ecosystems (Boddy 1983b; Chen et al. 2000; Flanagan and Veum 1974; Marra and Edmonds 1994; O'Connell 1990). The incubation study also demonstrated the importance of quantifying the interactive effect of CWD decay status, temperature, and  $W_C$  on  $R_{CWD}$ .

Standing woody debris decomposes extremely slowly (Harmon et al. 1986; Bond-Lamberty et al. 2002a), but

decomposition accelerates once the woody debris falls and establishes contact with forest floor. The CWD samples in the incubation experiment were not in contact with the soil, although they were in full contact with the ground before being collected. This might restrict the inference of our laboratory study to field application. However, validation of the  $R_{CWD}$  model with field data showed that the difference between the model predictions and measurements of  $R_{CWD}$  averaged  $\pm 2\%$  for decay I and II, and 18% for decay III CWD. The greater error for decay III may be attributed to the fragmentation effect on  $R_{CWD}$  (Harmon et al. 1986). The different decomposition rates are probably attributed to greater microbial colonization and water content of downed than standing CWD.

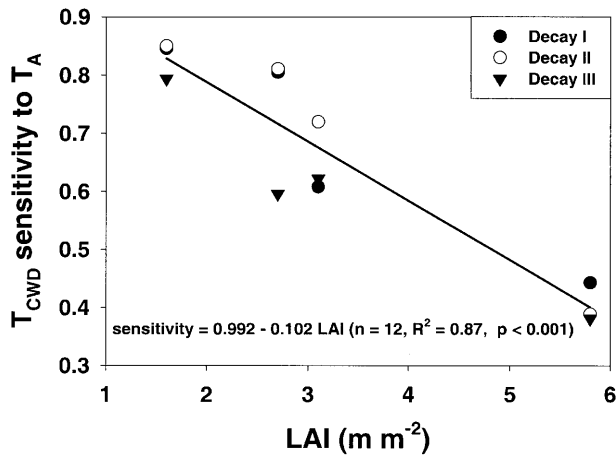
We compared annual  $R_{CWD}$  estimates between the environment-driven models (this study) and a  $Q_{10}$  model (Bond-Lamberty et al. 2002a). The annual  $R_{CWD}$  using the former approach was over two-fold greater than using the later approach. We speculate that the differences were due to effects of  $T_{CWD}$  hysteresis, different modeling time steps, and, more importantly, the  $R_{CWD}$  models used. The annual  $R_{CWD}$  by Bond-Lamberty et al. (2002a) was estimated from separate  $Q_{10}$  and  $W_C$  models, but the revised estimates were from a model that included the interactions between SLCWD,  $T_{CWD}$ , and  $W_C$ , all of which contributed significantly to the  $R_{CWD}$  models (Table 2). We conclude that CWD decay status in various environmental conditions must be considered when estimating  $R_{CWD}$ .

#### Modeling $R_{CWD}$ : temporal-spatial considerations

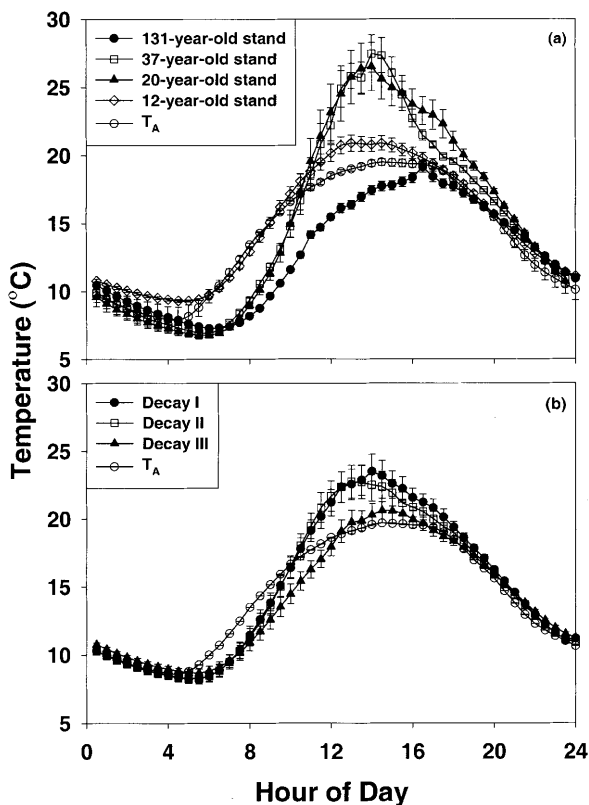
In the field,  $T_{CWD}$  was positively correlated to  $T_A$  ( $n=2,424$ ,  $R^2=0.79$ ,  $P<0.001$ ), but the relationship significantly differed among stands ( $df=3$ , 2,070,  $F=13.1$ ,  $P<0.001$ ) and CWD decay classes ( $df=2$ , 2,070,  $F=4.2$ ,  $P=0.015$ ). The slope of the relationship between  $T_{CWD}$  and  $T_A$  increased from old to young stands, except for decay class III at the 20-year-old stand; the slope was less for decay class III than decay class I and II except, for the 37-year-old stand (Table 3). The  $T_{CWD}$  at old stands with closed canopies responded to  $T_A$  changes

**Table 3** Models relating CWD temperature ( $T_{CWD}$ , °C) to air temperature ( $T_A$ , °C) for the black spruce fire chronosequence. The model form is:  $T_{CWD}=a + b T_A$ . The models are based on daily mean values of  $T_{CWD}$  and  $T_A$

Stand age (years)	Co-efficient	Decay I					Decay II					Decay III				
		Estimate	SE	df	t-ratio	P	Estimate	SE	df	t-ratio	P	Estimate	SE	df	t-ratio	P
12	a	2.672	0.35	2,387	7.62	<0.001	2.310	0.35	2,352	6.55	<0.001	2.773	0.35	2,385	7.90	<0.001
	b	0.846	0.02	2,386	37.35	<0.001	0.850	0.02	2,352	37.29	<0.001	0.794	0.02	2,383	35.00	<0.001
20	a	3.505	0.53	2,360	6.57	<0.001	3.323	0.53	2,379	6.21	<0.001	4.665	0.53	2,360	8.74	<0.001
	b	0.805	0.03	2,356	23.89	<0.001	0.811	0.03	2,374	24.02	<0.001	0.596	0.03	2,356	17.68	<0.001
37	a	4.068	1.55	2,017	2.63	0.009	3.495	1.55	1,998	2.26	0.024	3.918	1.55	1,998	2.54	0.011
	b	0.608	0.10	1,954	5.98	<0.001	0.720	0.10	1,933	7.09	<0.001	0.622	0.10	1,933	6.13	<0.001
131	a	5.860	0.98	1,514	6.00	<0.001	5.484	0.98	1,548	5.60	<0.001	5.958	0.98	1,514	6.10	<0.001
	b	0.443	0.07	1,437	6.50	<0.001	0.389	0.07	1,466	5.69	<0.001	0.381	0.07	1,437	5.59	<0.001

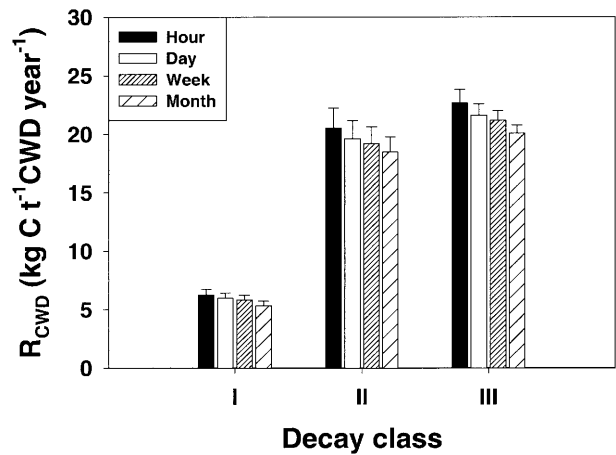


**Fig. 4** Relationship between CWD temperature sensitivity to air temperature (regression slope between  $T_{CWD}$  and  $T_A$  in Table 3) and stand leaf area index ( $LAI$ ) (adapted from Bond-Lamberty et al. 2002b) in the chronosequence



**Fig. 5** Diurnal patterns of CWD temperature ( $T_{CWD}$ , °C) during the growing season by **a** stands, and **b** decay classes. Error bars are the standard deviations ( $n=3$  or  $4$ ).  $T_A$  is air temperature

more slowly than at young stands with open canopies, because the greater leaf area index ( $LAI$ ) influenced the energy balance near the soil surface (Fig. 4). These results indicated that the  $T_{CWD}$  for decay class III at the older stand with the greatest  $LAI$  experienced the greatest hysteresis to  $T_A$ .



**Fig. 6** Modeled CWD  $CO_2$  flux ( $R_{CWD}$ ,  $kg\ C\ t^{-1}\ CWD\ year^{-1}$ ) at hourly, daily, weekly, and monthly time steps for the three decay classes. Error bars are standard errors ( $n=4$ )

There was no significant difference between diurnal  $T_{CWD}$  at the 12-year-old stand and  $T_A$  ( $df=1, 46, F=1.15, P=0.289$ ). The  $T_{CWD}$  amplitude was greater than  $T_A$  from late morning to early evening in the other stands, except for the oldest stand, where the  $T_{CWD}$  was significantly lower during daytime (Fig. 5a). The diurnal  $T_{CWD}$  of decay class III followed that of  $T_A$  ( $df=1, 46, F=3.3, P=0.075$ ), but the  $T_{CWD}$  for decay class I ( $df=1, 46, F=115.6, P<0.001$ ) and II ( $df=1, 46, F=84.8, P<0.001$ ) were significantly higher than  $T_A$  during the daytime (Fig. 5b).

Stand ( $df=3, 738, F=11.9, P<0.001$ ) and decay class ( $df=2, 738, F=3.1, P=0.045$ ) significantly affected weekly-based seasonal dynamics in  $T_{CWD}$ . The mean  $T_{CWD}$  during the growing season (from 1 June to 30 September 2001) were 15.0, 14.4, 13.6, and 12.1°C for 20-, 12-, 37-, and 131-year-old stands, respectively. The seasonal average  $T_{CWD}$  was lower than  $T_A$  in all the stands except the 20-year-old stand where  $T_{CWD}$  and  $T_A$  were similar ( $df=1, 15, F=0.9, P=0.351$ ). Mean  $T_{CWD}$  in the growing season decreased as CWD became more decomposed. There was no significant difference in  $T_{CWD}$  diurnal ( $df=1, 1163, F=1.0, P=0.309$ ) or seasonal ( $df=1, 443, F=0.7, P=0.406$ ) patterns within a specific CWD sample.

Modeling long-term CWD decomposition is challenging (Stone et al. 1998), because biophysical and chemical properties of CWD change as decomposition proceeds in forest ecosystems. Examples of such properties include CWD fragmentation, substrate heterogeneity (Harmon et al. 1986; Rayner and Boddy 1988; Schowalter 1992), density, surface area, charred CWD in fire-prone ecosystems, and nutrient losses (Yavitt and Fahey 1985). For simplicity, we restricted our study to examining the effects of temperature,  $W_C$ , and decay class on  $R_{CWD}$ . The empirical models developed in this study may not provide accurate estimates of  $R_{CWD}$  if other abiotic or biotic variables are more limiting to microbial activity.

Temperature is a major environmental factor influencing  $R_{\text{CWD}}$ . Spatial dynamics of  $T_{\text{CWD}}$  at various decomposition stages under different canopy conditions must be considered in estimating annual  $R_{\text{CWD}}$ .  $T_{\text{CWD}}$  generally follows  $T_{\text{A}}$  (Boddy 1983b), but substrates with such a high thermal capacity as CWD exhibit hysteresis (Campbell and Norman 1998). Ignoring  $T_{\text{CWD}}$  hysteresis in different stands can introduce large errors in annual  $R_{\text{CWD}}$  estimates. If  $T_{\text{A}}$ , instead of  $T_{\text{CWD}}$ , was used in the  $R_{\text{CWD}}$  models with an hourly time step, the annual  $R_{\text{CWD}}$  for the 12-, 20-, 37-, and 131-year-old stands varied by -2, -1, +14, and +32%, respectively. These differences were -4, -5, +9, and +20% when a weekly time step was used.

Modeled annual  $R_{\text{CWD}}$  across the chronosequence on daily, weekly, and monthly time steps averaged 5%, 7%, and 11%, respectively, less than that on an hourly time step (Fig. 6). The use of large time steps in  $R_{\text{CWD}}$  models underestimates annual  $R_{\text{CWD}}$  because of the nonlinear relationship among  $R_{\text{CWD}}$ ,  $T_{\text{CWD}}$ , and  $W_{\text{C}}$  (Table 2). Nonlinear relationships between temperature, autotrophic and heterotrophic respiration have been widely reported (Kozłowski and Pallardy 1997; Landsberg and Gower 1997; Raich and Schlesinger 1992). Thus, the respiration nonlinearity feature should be taken into account in all  $\text{CO}_2$  flux modeling.

In conclusion,  $\text{CO}_2$  flux from CWD was an important  $\text{CO}_2$  input to the atmosphere in the fire-prone boreal forest that we studied. The importance of  $\text{CO}_2$  flux from CWD in most forest ecosystems is unknown because of the paucity of CWD  $\text{CO}_2$  flux data; however,  $R_{\text{CWD}}$  should not be ignored in any forest ecosystem that contains large amount of CWD. Annual estimates of CWD  $\text{CO}_2$  flux can be improved if the spatial-temporal dynamics in decay status, temperature, and  $W_{\text{C}}$  are included in  $R_{\text{CWD}}$  models.

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## References

- Boddy L (1983a) Carbon dioxide release from decomposing wood: effect of water content and temperature. *Biol Biochem* 15:501-510
- Boddy L (1983b) Microclimate and moisture dynamics of wood decomposing in terrestrial ecosystems. *Biol Biochem* 15: 149-157
- Bond-Lamberty B, Wang C, Gower ST (2002a) Annual carbon flux from woody debris for a boreal black spruce fire chronosequence. *J Geophys Res* (in press)
- Bond-Lamberty B, Wang C, Gower ST, Norman JM (2002b) Leaf area dynamics of a boreal black spruce fire chronosequence. *Tree Physiol* (in press)
- Campbell GS, Norman JM (1998) An introduction to environmental biophysics. Springer, Berlin Heidelberg New York, pp 113-128
- Chambers JQ, Schimel JP, Nobre AD (2001) Respiration from coarse wood litter in central Amazon forests. *Biogeochemistry* 52:115-131
- Chen H, Harmon ME, Griffiths RP, Hicks W (2000) Effects of temperature and moisture on carbon respired from decomposing woody roots. *For Ecol Manage* 138:51-64
- Flanagan PW, Veum AK (1974) Relationships between respiration, weight loss, temperature and moisture in organic residues on tundra. In: Holding AJ, Heal OW, Maclean SF Jr, Flanagan PW (eds) Soil organisms and decomposition in tundra. Tundra Biome Steering Committee, Stockholm, Sweden, pp 249-277
- Frangi JL, Richter LL, Barrera MD, Aloggia M (1997) Decomposition of *Nothofagus* fallen woody debris in forests of Tierra del Fuego, Argentina. *Can J For Res* 27:1095-1102
- Goldammer JG, Furyaev VV (1997) Fire in ecosystems of boreal Eurasia. *Kluwer For Sci* (48). Academic Press, Boston, Mass.
- Gower ST, Bond-Lamberty B, Wang C, Litvak M, Goulden M, Trumbore SE, Harden J, Foley J, Norman JM, Veldhuis H, Diak G, Cihlar J, Fernandez R (2002) Measuring and modeling changes in structure and function of a Canadian boreal black spruce forest following wildfire: a chronosequence approach. *J Geophys Res* (in press)
- Harmon ME, Sexton J (1995) Water balance of conifer logs in early stages of decomposition. *Plant Soil* 172:141-152
- Harmon ME, Franklin JF, Swanson FJ, Sollins P, Gregory SV, Lattin JD, Anderson NH, Cline SP, Aumen NJ, Sedell JR, Lienkaemper GW, Cromack K Jr, Cummins KW (1986) Ecology of coarse woody debris in temperate ecosystems. *Adv Ecol Res* 15:133-302
- Käärik AA (1974) Decomposition of wood. In: Dickinson CH, Pugh GJF (eds) Biology of plant litter decomposition. Academic Press, London, pp 129-174
- Kasischke ES (2000) Boreal ecosystems in the global carbon cycle. In: Kasischke ES, Stocks BJ (eds) Fire, climate change, and carbon cycling in the boreal forest. (Ecological studies series) Springer, Berlin Heidelberg New York, pp19-30
- Kozłowski TT, Pallardy SG (1997) The physiological ecology of woody plants. Academic Press, San Diego, Calif.
- Krankina ON, Harmon ME (1995) Dynamics of the dead wood carbon pool in northwestern Russian boreal forests. *Water Air Soil Pollut* 82:227-238
- Lambert RL, Lang GE, Reiners WA (1980) Loss of mass and chemical change in decaying boles of a subalpine fir forest. *Ecology* 61:1460-1473
- Landsberg JJ, Gower ST (1997) Applications of physiological ecology to forest management. Academic Press, San Diego, Calif.
- Lloyd J, Taylor JA (1994) On the temperature dependence of soil respiration. *Funct Ecol* 8:315-323
- Marra JL, Edmonds RL (1994) Coarse woody debris and forest floor respiration in an old-growth coniferous forest on the Olympic Peninsula, Washington, USA. *Can J For Res* 24: 1811-1817
- Moore AE (1986) Temperature and moisture dependence of decomposition rates of hardwood and coniferous leaf litter. *Soil Biol Biochem* 18:427-435
- Nadelhoffer KJ, Giblin AE, Shaver GR, Laundre JA (1991) Effects of temperature and substrate quality on element mineralization in six arctic soils. *Ecology* 72:242-253
- O'Connell (1990) Microbial decomposition (respiration) of litter in Eucalypt forests of south-western Australia: an empirical model based on laboratory incubations. *Soil Biol Biochem* 22: 153-160
- Potter CS, Randerson JT, Field CB, Matson PA, Vitousek PM, Mooney HA, Klooster SA (1993) Terrestrial ecosystem production: a process model based on global satellite and surface data. *Global Biogeochem Cycle* 7: 811-841
- Progar RA, Schowalter TD, Freitag CM (2000) Respiration from coarse woody debris as affected by moisture and saprotroph functional diversity in Western Oregon. *Oecologia* 124: 426-431

- Raich WA, Schlesinger WA (1992) The global carbon dioxide flux in soil respiration and its relationship to vegetation and climate. *Tellus* 44b:81–99
- Raich JW, Rastetter EB, Melillo JM, Kicklighter DW, Steudler PA, Parterson BJ (1991) Potential primary productivity in South America: application of a global model. *Ecol Appl* 1: 399–429
- Rayment MB, Jarvis PJ (2000) Temporal and spatial variation of soil CO<sub>2</sub> efflux in a Canadian boreal forest. *Soil Biol Biochem* 32:35–45
- Rayner ADM, Boddy L (1988) Fungal decomposition of wood: its biology and ecology. Wiley, New York
- SAS (2001) SAS Online Documentation, Version 8.2. Cary, N.C.
- Schowalter TD (1992) Heterogeneity of decomposition and nutrient dynamics of oak (*Quercus*) logs during the first 2 years of decomposition. *Can J For Res* 22:161–166
- Siitonen J, Martikainen P, Punttila P, Rauh J (2000) Coarse woody debris and stand characteristics in mature managed and old-growth boreal mesic forests in southern Finland. *For Ecol Manage* 128:211–225
- Spies TA, Franklin JF, Thomas TB (1988) Coarse woody debris in Douglas-fir forests of western Oregon and Washington. *Ecology* 69:1689–1702
- Stewart GH, Burrows LE (1994) Coarse woody debris on old-growth temperate beech (*Nothofagus*) forests of New Zealand. *Can J For Res* 24:1989–1996
- Stone JN, MacKinnon A, Parminter JV, Lertzman KP (1998) Coarse woody debris decomposition documented over 65 years on southern Vancouver Island. *Can J For Res* 28: 788–793
- Strömberg M (2001) Soil-surface CO<sub>2</sub> flux and growth in a boreal Norway spruce stand. Effects of soil warming and nutrition. PhD thesis. *Acta Univ Agric Suec Silvestria* 220
- Wang C, Bond-Lamberty B, Gower ST (2002) Soil surface CO<sub>2</sub> flux in a boreal black spruce fire Chronosequence. *J Geophys Res* (in press)
- Yavitt JB, Fahey TJ (1985) Chemical composition of interstitial water in decaying lodgepole pine bole wood. *Can J For Res* 15:1149–1153
- Yoneda T (1975) Studies on the rate of decay of wood litter on the forest floor. II. Dry weight loss and CO<sub>2</sub> evolution of decaying wood. *Jpn J Ecol* 25:132–140
- Yoneda T (1980) Studies on the rate of decay of wood litter on the forest floor. III. Effect of moisture content on CO<sub>2</sub> evolution from decaying wood. *Jpn J Ecol* 30:55–62