

# Enhancements of Major Aerosol Components Due to Additional HONO Sources in the North China Plain and Implications for Visibility and Haze

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## ABSTRACT

The Weather Research and Forecasting/Chemistry model (WRF-Chem) was updated by including photoexcited nitrogen dioxide (NO<sub>2</sub>) molecules, heterogeneous reactions on aerosol surfaces, and direct emissions of nitrous acid (HONO) in the Carbon-Bond Mechanism Z (CBM-Z). Five simulations were conducted to assess the effects of each new component and the three additional HONO sources on concentrations of major chemical components. We calculated percentage changes of major aerosol components and concentration ratios of gas NO<sub>y</sub> (NO<sub>y,g</sub>) to NO<sub>y</sub> and particulate nitrates (NO<sub>3</sub><sup>-</sup>) to NO<sub>y</sub> due to the three additional HONO sources in the North China Plain in August of 2007. Our results indicate that when the three additional HONO sources are included, WRF-Chem can reasonably reproduce the HONO observations. Heterogeneous reactions on aerosol surfaces are a key contributor to concentrations of HONO, nitrates (NO<sub>3</sub><sup>-</sup>), ammonium (NH<sub>4</sub><sup>+</sup>), and PM<sub>2.5</sub> (concentration of particulate matter of ≤2.5 μm in the ambient air) across the North China Plain. The three additional HONO sources produced a ~5%–20% increase in monthly mean daytime concentration ratios of NO<sub>3</sub><sup>-</sup>/NO<sub>y</sub>, a ~15%–52% increase in maximum hourly mean concentration ratios of NO<sub>3</sub><sup>-</sup>/NO<sub>y</sub>, and a ~10%–50% increase in monthly mean concentrations of NO<sub>3</sub><sup>-</sup> and NH<sub>4</sub><sup>+</sup> across large areas of the North China Plain. For the Bohai Bay, the largest hourly increases of NO<sub>3</sub><sup>-</sup> exceeded 90%, of NH<sub>4</sub><sup>+</sup> exceeded 80%, and of PM<sub>2.5</sub> exceeded 40%, due to the three additional HONO sources. This implies that the three additional HONO sources can aggravate regional air pollution, further impair visibility, and enhance the incidence of haze in some industrialized regions with high emissions of NO<sub>x</sub> and particulate matter under favorable meteorological conditions.

**Key words:** HONO, NO<sub>y</sub>, aerosol component, heterogeneous reaction, WRF-Chem model

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## 1. Introduction

Nitrous acid (HONO) is one of the major sources of the hydroxyl radical (OH), which is the key oxidant in the atmosphere. Direct emissions, gas-phase reactions, and heterogeneous reactions are generally considered as HONO sources, but the detailed formation mechanism of HONO is still under discussion. High HONO concentrations in urban or rural areas are frequently

observed (Su et al., 2008; An et al., 2009; Qin et al., 2009; Yu et al., 2009), but most current air-quality models, such as WRF-Chem, cannot reproduce HONO observations very well, particularly in the daytime, due mainly to the only inclusion of gas-phase production of HONO (Li et al., 2010; An et al., 2011; Li et al., 2011). Recently, Sarwar et al. (2008) added HONO emissions by using the HONO/NO<sub>x</sub> emission ratio of 0.8%, 2NO<sub>2</sub> + H<sub>2</sub>O + aerosol/ground surfaces → HONO

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+ HNO<sub>3</sub>, and HNO<sub>3</sub> (adsorbed) + hν → 0.5HONO (adsorbed) + 0.5NO<sub>2</sub> (adsorbed) into the Community Multiscale Air Quality model (CMAQ) (Foley et al., 2009). HONO simulations were improved, but obvious discrepancies between simulated and observed HONO concentrations still existed in the daytime period (Sarwar et al., 2008). Li et al. (2008) proposed a new production mechanism of HONO through photoexcited NO<sub>2</sub> chemistry ( $R_{\text{NO}_2\text{S}}$ , where S means photoexcited). The importance of  $R_{\text{NO}_2\text{S}}$  is still argued (Crowley and Carl, 1997; Wennberg and Dabdub, 2008; Sarwar et al., 2009; Ensberg et al., 2010; An et al., 2011; Li et al., 2011). For heterogeneous reactions, Li et al. (2010) considered both aerosol and ground surface reactions in WRF-Chem; however, field experiments showed a good correlation between concentrations of particulate matter and HONO (An et al., 2009), or between aerosol surface area and HONO concentrations (Ziemba et al., 2010), suggesting that aerosol surface is the dominant reaction substrate and that stationary sources (e.g., buildings and soils) are likely insignificant (Ziemba et al., 2010). Additionally, Li et al. (2010) used a relatively high emission ratio of 2.3% for HONO/NO<sub>2</sub> to calculate direct emissions of HONO. This could lead to overestimation of HONO concentrations in the air. Recently, Su et al. (2011) have suggested that soil nitrites are a source of HONO emissions, but nitrite concentrations in soils are generally low (Cleemput and Samater, 1996), and related field observations are scarce. Based on the current understanding of HONO sources, An et al. (2011) and Li et al. (2011) incorporated heterogeneous reactions on aerosol surfaces ( $R_{\text{het}}$ ), an improved parameterization scheme (see section 2.2) for direct emissions of HONO ( $R_{\text{em}}$ ), and  $R_{\text{NO}_2\text{S}}$ , which could be important in areas where NO<sub>x</sub> emissions are elevated (e.g., the North China Plain), into WRF-Chem and found that the three additional HONO sources (i.e.,  $R_{\text{NO}_2\text{S}}$ ,  $R_{\text{het}}$ , and  $R_{\text{em}}$ ) significantly improved HONO simulations in comparison with observations from differential optical absorption spectroscopy (DOAS) (Zhu et al., 2009), particularly in the daytime.

The purpose of this study was to compute the contributions of  $R_{\text{NO}_2\text{S}}$ ,  $R_{\text{het}}$ ,  $R_{\text{em}}$ , and the three additional HONO sources ( $R_{\text{NO}_2\text{S}}$ ,  $R_{\text{het}}$ , and  $R_{\text{em}}$ ) to concentrations of major chemical components and to estimate percentage enhancements of major aerosol components (i.e., NO<sub>3</sub><sup>-</sup>, NH<sub>4</sub><sup>+</sup>, SO<sub>4</sub><sup>2-</sup>, and also PM<sub>2.5</sub>) due to the three additional HONO sources in the North China Plain, where emissions of NO<sub>x</sub> and particulate matter (PM) are high (Zhang et al., 2009). Changes were calculated in concentration gas ratios of NO<sub>y</sub> (i.e., NO<sub>yg</sub> = NO + NO<sub>2</sub> + NO<sub>3</sub> + PAN + HNO<sub>3</sub> + HONO + HNO<sub>4</sub> + N<sub>2</sub>O<sub>5</sub>) to NO<sub>y</sub> (i.e., NO<sub>y</sub> = NO<sub>yg</sub> +

particulate nitrates) and particulate nitrates (NO<sub>3</sub><sup>-</sup>) to NO<sub>y</sub> due to the three additional HONO sources. The implications of these results were analyzed and were discussed in section 3.

## 2. Methodology

### 2.1 WRF-Chem model setup

The WRF-Chem version 3.2.1 (Grell et al., 2005; Fast et al., 2006) was utilized in this study. Physical and chemical schemes for simulations followed those of An et al. (2011) and Li et al. (2011). Two nested domains were used in the simulation; domain 3 covered the North China Plain (An et al., 2011; Li et al., 2011). We used 28 vertical model layers from the ground to ~50 hPa, with the first layer 28 m above the ground. Initial and boundary conditions for meteorological fields were obtained from NCEP reanalysis data applied to nudging every 6 h, and initial and boundary conditions for chemical fields were constrained using the output of the Model for Ozone and Related chemical Tracers, version 4 (MOZART-4) (Emmons et al., 2010) every 6 h. Anthropogenic emissions of sulfur dioxide (SO<sub>2</sub>), nitrogen oxides (NO<sub>x</sub>), carbon monoxide (CO), volatile organic compounds (VOCs), PM<sub>10</sub>, PM<sub>2.5</sub>, black carbon (BC), and organic carbon (OC) were taken from Zhang et al. (2009), and those of NH<sub>3</sub> were adopted from Streets et al. (2003). Changes in NH<sub>3</sub> emissions between 2000 and 2006 were insignificant (Zhang et al., 2009). Biogenic emissions were computed on the basis of the work of Guenther et al. (1993) and Simpson et al. (1995). Five model simulations (cases A, B, C, R, and E) were conducted to calculate contributions of each component as well as the three additional HONO sources to concentrations of major chemical components in percentage change concentrations of NO<sub>3</sub><sup>-</sup>, NH<sub>4</sub><sup>+</sup>, SO<sub>4</sub><sup>2-</sup>, and PM<sub>2.5</sub>, and concentration ratios of NO<sub>yg</sub>/NO<sub>y</sub> and NO<sub>3</sub><sup>-</sup>/NO<sub>y</sub> due to the three additional HONO sources in the North China Plain in August 2007. Case R used the Carbon-Bond Mechanism Z (CBM-Z) and the Model for Simulating Aerosol Interactions and Chemistry (MOSAIC) (Zaveri et al., 2008). Cases A, B, C, and E were extensions of case R by inclusion of  $R_{\text{NO}_2\text{S}}$ ,  $R_{\text{em}}$ ,  $R_{\text{het}}$ , and the three additional HONO sources ( $R_{\text{NO}_2\text{S}} + R_{\text{het}} + R_{\text{em}}$ ), respectively (see section 2.2, Table 1).

### 2.2 Parameterization of the three additional HONO sources

For  $R_{\text{het}}$  (NO<sub>2</sub> → 0.5 HNO<sub>3</sub> + 0.5 HONO), we followed the recommendation of Jacob (2000):

$$k_{\text{het}} = S_{\text{a}} \left( \frac{r_{\text{p}}}{D_{\text{g}}} + \frac{4}{u\gamma} \right)^{-1},$$

**Table 1.** Five simulated scenarios used in this research

Case	Photo-excited NO <sub>2</sub> chemistry	HONO emissions	NO <sub>2</sub> heterogeneous production	Gas-phase production
R				✓
A	✓			✓
B		✓		✓
C			✓	✓
E	✓	✓	✓	✓

where  $k_{\text{het}}$  denotes the first-order rate constant and  $S_a$  is the aerosol surface area per unit volume of air.  $S_a$  was derived from aerosol mass concentrations and number density in each bin set by the MOSAIC module. Aerosols considered in MOSAIC are composed of sulfate, nitrate, ammonium, chloride, sodium, other inorganics, organic carbon (OC), elemental carbon (EC), and water.  $r_p$  represents the particle radius (m).  $D_g$  denotes gas molecular diffusion coefficients estimated as  $10^{-5} \text{ m}^2 \text{ s}^{-1}$  (Dentener and Crutzen, 1993).  $u$  is the mean molecular speed of NO<sub>2</sub> ( $\text{m s}^{-1}$ ).  $\gamma$  denotes the uptake coefficient of NO<sub>2</sub>, taken as  $10^{-4}$  (Jacob, 2000).

HONO emissions ( $R_{\text{em}}$ ) were computed using

$$R_{\text{em}} = [0.023 \times f_{\text{DV}} + 0.008 \times (1 - f_{\text{DV}})] \times f_{\text{TS}}$$

where  $f_{\text{DV}}$  stands for the NO<sub>x</sub> emission ratio of diesel vehicles to total vehicles. The averaged  $f_{\text{DV}}$  was 62% in China in 2006 (Li et al., 2011).  $f_{\text{TS}}$  is the NO<sub>x</sub> emission ratio of the traffic source to all anthropogenic sources. The value of 0.8% has been suggested by Kurtenbach et al. (2001) and was used by Aumont et al. (2003), Sarwar et al. (2008), and An et al. (2009). A fraction of 2.3% of the NO<sub>x</sub> emitted in diesel exhaust can be heterogeneously converted to HONO (Gutzwiller et al., 2002), and this fraction was utilized in this research.  $R_{\text{het}}$  and  $R_{\text{NO}_2\text{S}}$  were added to the CBM-Z mechanism, as introduced by Li et al. (2011).

### 3. Results and discussion

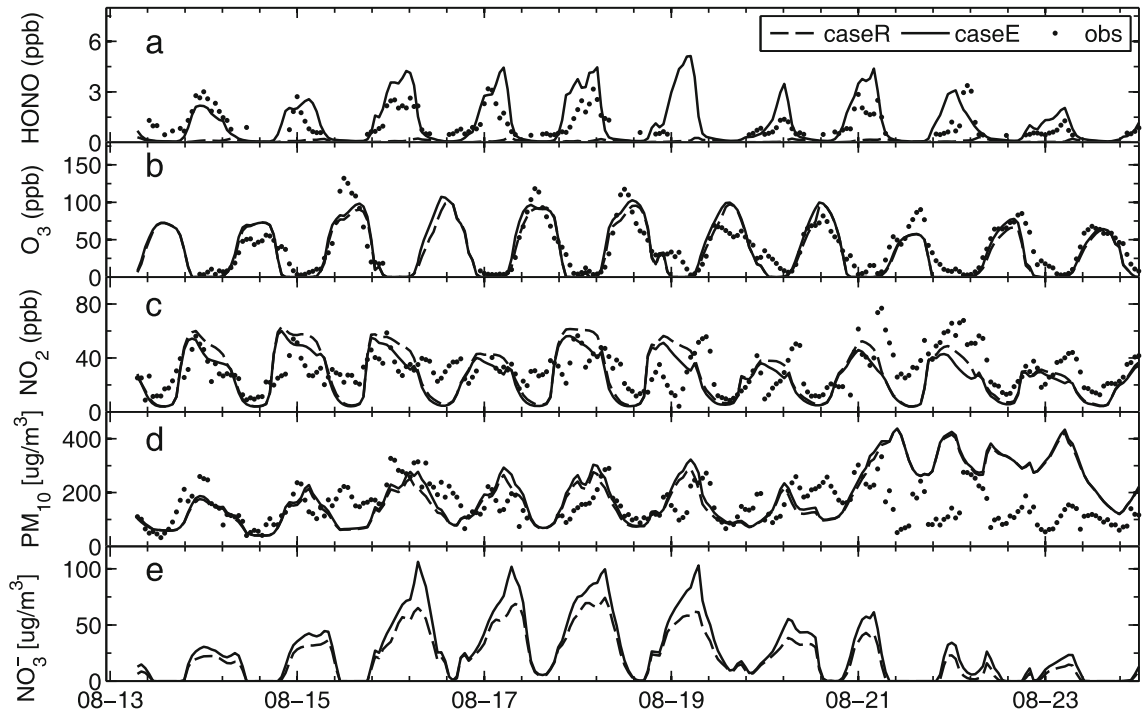
#### 3.1 Comparison of simulations and observations

HONO concentrations measured by DOAS in Beijing on 13–25 August 2007 were obtained from Zhu et al. (2009). Ozone (O<sub>3</sub>), NO<sub>x</sub>, and PM<sub>10</sub> were simultaneously monitored at the 325-m meteorological tower in Beijing (Li et al., 2011). Li et al. (2011) demonstrated good agreement between O<sub>3</sub> and NO<sub>2</sub> measurements from DOAS and the O<sub>3</sub> and NO<sub>2</sub> analyzers in Beijing. The correlation coefficients were 0.97 for O<sub>3</sub> and 0.83 for NO<sub>2</sub>. Compared with gas-phase reactions (case R), the inclusion of  $R_{\text{het}}$ ,  $R_{\text{em}}$ , and  $R_{\text{NO}_2\text{S}}$  in WRF-Chem (case E) led to significant improvements in HONO simulations of measurements at the 325-m

meteorological tower in Beijing (Fig. 1a and Table 2). For example, the root mean square error was reduced by ~50%, and the correlation coefficient improved to 0.97 from 0.19 (Table 2). However, HONO simulations for case E were overestimated by comparison with observations due to overprediction of PM observations (Fig. 1d) and the exclusion of HO<sub>2</sub> heterogeneous reactions on aerosol surfaces. The HO<sub>2</sub> heterogeneous loss on aerosol surfaces ( $2 \text{ HO}_2 \rightarrow \text{H}_2\text{O}_2$ ) considered in Li et al. (2011) lowered HONO simulations when Fig. 1a of this study was compared to Fig. 6a in Li et al. (2011). The uncertainty of the heterogeneous uptake coefficient of HO<sub>2</sub> was high (Thornton and Abbatt, 2005), and the reaction of  $2 \text{ HO}_2 \rightarrow \text{H}_2\text{O}_2$  was not a source of HONO, so the HO<sub>2</sub> heterogeneous reaction was excluded from this research and will be investigated in the near future. The three HONO sources can improve daytime O<sub>3</sub> simulations, especially the peak concentrations. Daytime mean value of O<sub>3</sub> was enhanced by 2.5 ppb, and the maximum hourly value was enhanced by 21 ppb (35%) at 1200 LST 16 August. NO<sub>2</sub> nighttime simulations improved significantly, with the mean value reduced by 5.4 ppb, mainly due to the heterogeneous reaction on aerosol surfaces (Li et al., 2011). PM<sub>10</sub> simulations for cases E and R were reasonable in most cases, except for the period 21–25 August (Fig. 1d). For this period, PM<sub>10</sub> simulations for cases E and R were substantially overestimated (Fig. 1d), and further investigation was needed. The calculated mean nitrate concentration for case R was  $20.13 \mu\text{g m}^{-3}$ , while that for case E was enhanced to  $26.94 \mu\text{g m}^{-3}$  due to the additional HONO sources. NO<sub>3</sub><sup>-</sup> peaks were significantly enhanced, similar to the results of Li et al. (2010).

#### 3.2 Impacts of each and the three additional HONO sources on major gas components

$R_{\text{NO}_2\text{S}}$  was shown to enhance the monthly mean daytime (~0700–1900 LST) HONO by a maximum of 60 ppt (Table 3), being lower than 100 ppt in the Eastern and Western United States (Sarwar et al., 2009).  $R_{\text{NO}_2\text{S}}$  and photolysis of HONO can produce more OH, which reacts with VOCs to yield more HO<sub>2</sub>. More HO<sub>2</sub> can accelerate conversion of NO to NO<sub>2</sub> and result in more O<sub>3</sub> (Table 3). On the other hand, increases



**Fig. 1.** Comparison of simulated concentrations of (a) HONO, (b) O<sub>3</sub>, (c) NO<sub>2</sub>, (d) PM<sub>10</sub>, and (e) NO<sub>3</sub><sup>-</sup> for cases E and R with observations from the 325-m meteorological tower in Beijing on 13–25 August 2007. No observations were available for NO<sub>3</sub><sup>-</sup>.

in OH can lead to HNO<sub>3</sub> enhancements through the reaction of NO<sub>2</sub> + OH → HNO<sub>3</sub>. The reaction of HNO<sub>3</sub> + NH<sub>3</sub> → NH<sub>4</sub>NO<sub>3</sub> yields noticeable increases of NO<sub>3</sub><sup>-</sup> and NH<sub>4</sub><sup>+</sup>, with a maximum of 3.0 μg m<sup>-3</sup> (Table 3).  $R_{em}$  produces greater increases of HONO in the nighttime (2000–0600 LST) than in the daytime (Table 3) due to daytime photolysis of HONO.  $R_{het}$  plays a pivotal role in the formation of HONO, NO<sub>3</sub><sup>-</sup>, NH<sub>4</sub><sup>+</sup> and PM<sub>2.5</sub> in both daytime and nighttime:  $R_{het}$  yielded an increase of 3.4 ppb of maximum monthly mean nighttime HONO, maximum monthly mean daytime NO<sub>3</sub><sup>-</sup> was enhanced to 13.3 μg m<sup>-3</sup> and maximum monthly mean daytime PM<sub>2.5</sub> was enhanced to 18.1 μg m<sup>-3</sup> (Table 3). When the three additional HONO sources were considered, the largest increases in monthly mean daytime were 0.6 ppb for HONO, 4.0 ppt for HO<sub>2</sub>, 17.0 μg m<sup>-3</sup> for NO<sub>3</sub><sup>-</sup>, 5.5 μg m<sup>-3</sup> for NH<sub>4</sub><sup>+</sup>, and 23.9 μg m<sup>-3</sup> for PM<sub>2.5</sub> (Table 3). For O<sub>3</sub> simulations the three additional HONO sources lead

to ~1–20 ppb increases in maximum hourly averages (Fig. 2a) and ~1–4.9 ppb enhancements of monthly daytime averages (Fig. 2b), with maximum values located areas of elevated emissions in the North China Plain or along the Bohai Bay (see discussion in section 3.3).

### 3.3 Impacts of the three additional HONO sources on major aerosol components

Percent changes in concentrations of air pollutants due to the three additional HONO sources were defined as  $(C_E - C_R) \times 100\% / C_R$ , where  $C_E$  and  $C_R$  denote concentrations of air pollutants for cases E and R, respectively. The three additional HONO sources ( $R_{NO_2S}$ ,  $R_{het}$ , and  $R_{em}$ ) resulted in ~10%–50% enhancements of NO<sub>3</sub><sup>-</sup>, ~10%–40% increases of NH<sub>4</sub><sup>+</sup>, ~6%–15% enhancements of SO<sub>4</sub><sup>2-</sup>, and ~3%–12% increase of PM<sub>2.5</sub> in major cities of the North

**Table 2.** Model performance statistics for HONO diurnal-averaged simulations of measurements at the 325-m meteorological tower in Beijing on 13–25 August 2007.

Case	Observed mean (ppb)	Simulated mean (ppb)	NMB (%)	RMSE (ppb)	RC
R	1.02	0.08	-92	1.10	0.19
E	1.02	1.12	29	0.66	0.97

Note: NMB, RMSE, and RC denote the normal mean bias, the root mean square error, and the correlation coefficient, respectively.

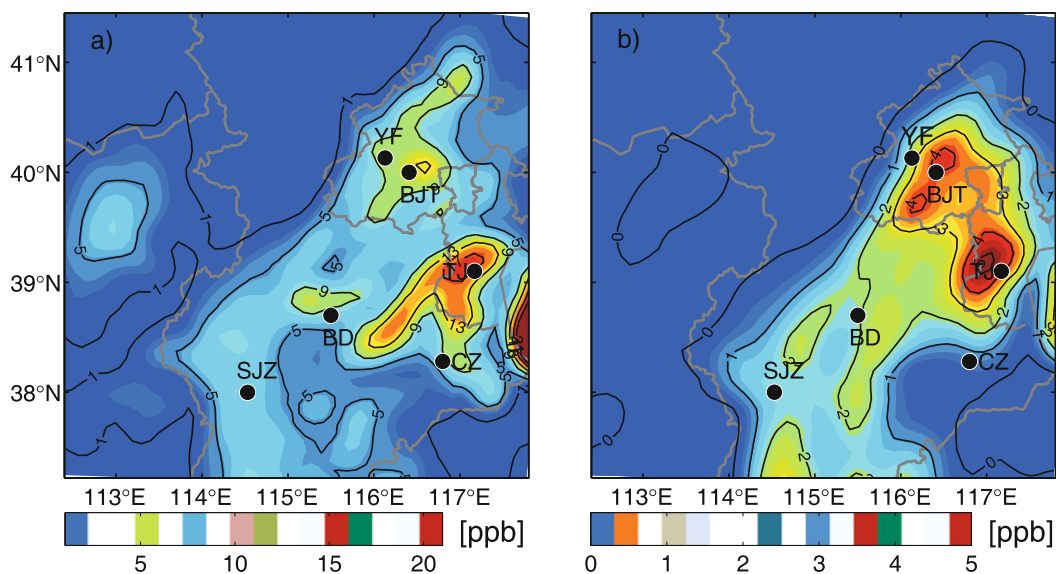
**Table 3.** Maximum monthly mean daytime (0700–1900 LST) and nighttime (2000–0600 LST) enhancements of  $O_3$ , HONO, OH,  $HO_2$ ,  $HNO_3$ ,  $NO_3^-$ ,  $NH_4^+$ ,  $SO_4^{2-}$ ,  $PM_{2.5}$ , and  $NO_3^-/NO_y$  due to photo-excited  $NO_2$  chemistry (case A–case R), direct emissions of HONO (case B–case R), heterogeneous reaction on aerosol surfaces (case C–case R) and the three additional HONO sources (case E–case R). The unit of  $O_3$ , HONO, and  $HNO_3$  is ppb, that of OH and  $HO_2$  is ppt, and that of  $NO_3^-$ ,  $NH_4^+$ ,  $SO_4^{2-}$ , and  $PM_{2.5}$  is  $\mu g m^{-3}$ , and that of  $NO_3^-/NO_y$  is %. A percentage change is shown in parentheses, and a nighttime enhancement is shown behind a slash and its corresponding percentage change is shown both behind a slash and in parentheses

Species	Case A–case R	Case B–case R	Case C–case R	Case E–case R
$O_3$	2.0 (4.2)	0.4 (0.9)	3.4 (7.5)	4.9 (11.7)
HONO	0.06 (70.0)	0.09/1.0 (118.4/>200)	0.4/3.4 (>200/>200)	0.6/4.5 (>200/>200)
OH	0.03 (19.4)	0.01 (7.9)	0.09 (55.7)	0.1 (83.6)
$HO_2$	0.7 (15.6)	0.3 (7.4)	3.2 (59.6)	4.0 (88.9)
$HNO_3$	0.2 (4.6)	0.02 (0.9)	0.5/0.2 (14.5/40.7)	0.8/0.2 (18.7/40.7)
$NO_3^-$	3.0 (7.0)	0.4 (1.6)	13.3/13.2 (47.7/47.6)	17.0/14.1 (53.9/51.0)
$NH_4^+$	1.0 (6.2)	0.1 (1.5)	4.2/3.9 (34.5/35.1)	5.5/4.3 (39.7/38.1)
$SO_4^{2-}$	0.3 (4.0)	0.2 (1.7)	1.5 (13.7)	1.8 (16.6)
$PM_{2.5}$	4.5 (2.7)	0.7 (0.6)	18.1/17.5 (13.4/8.8)	23.9/18.8 (16.5/9.6)
$NO_3^-/NO_y$	3.2/1.1	0.6/0.5	17.9/5.9	19.5/6.3

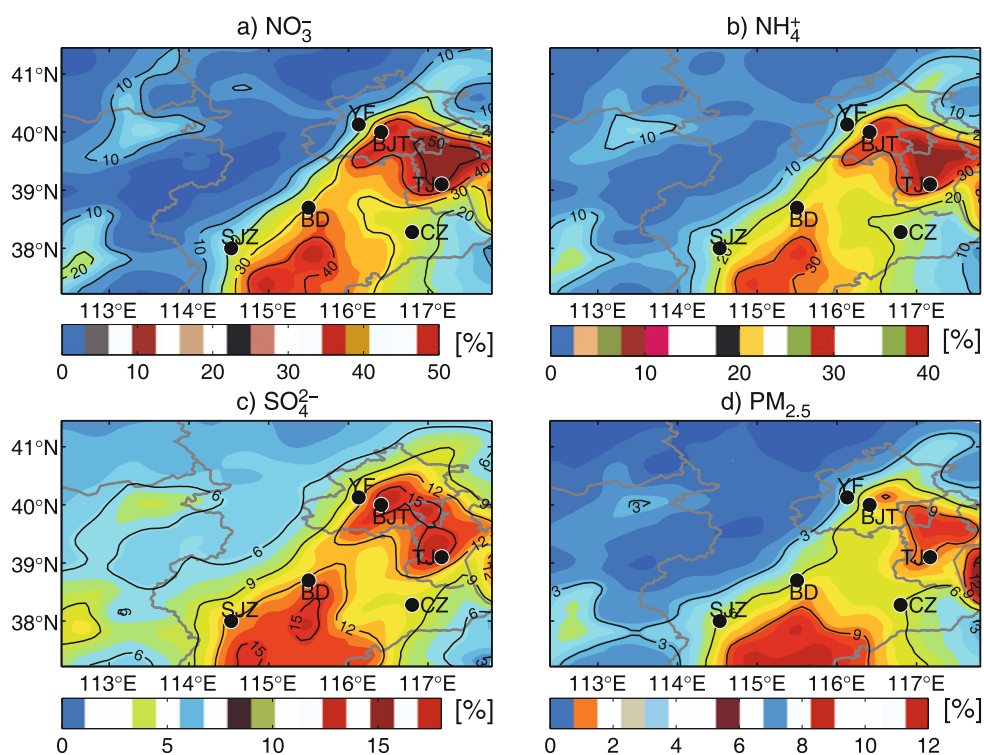
China Plain (i.e., Beijing, Tianjin, Baoding, and Shijiazhuang; Fig. 3) where emissions of  $NO_x$ , PM,  $NH_3$ , and  $SO_2$  are high (Streets et al., 2003; Zhang et al., 2009). Elevated emissions of  $NO_x$  and PM produced more HONO and  $HNO_3$  through the reaction of  $R_{het}$  in both daytime and nighttime (Table 3). High emissions of  $NO_x$  further enhanced HONO through the reaction of  $R_{NO_2S}$  in the daytime and direct emissions ( $R_{em}$ ) in both daytime and nighttime (Table 3). Enhancements of HONO yielded more OH through the photolysis of HONO, finally producing high  $O_3$  (Fig. 2). On the other hand, increases of OH and  $HNO_3$  together with high emissions of  $NH_3$  and  $SO_2$  were favorable for the

production of  $NO_3^-$ ,  $NH_4^+$ ,  $SO_4^{2-}$ , and  $PM_{2.5}$ . By contrast, areas with relatively low emissions of  $NO_x$ , PM,  $NH_3$ , and  $SO_2$  had low percentage increases of  $NO_3^-$ ,  $NH_4^+$ ,  $SO_4^{2-}$ , and  $PM_{2.5}$  (Fig. 3) and minor enhancements of  $O_3$  (Fig. 2) due to minor increases of OH produced through the three additional HONO sources. In the nighttime,  $R_{het}$  produced large enhancements of  $NO_3^-$  and  $NH_4^+$ , with a maximum of  $\sim 40\%$  (figures not shown); however, increases of  $SO_4^{2-}$  were  $< 9\%$ .

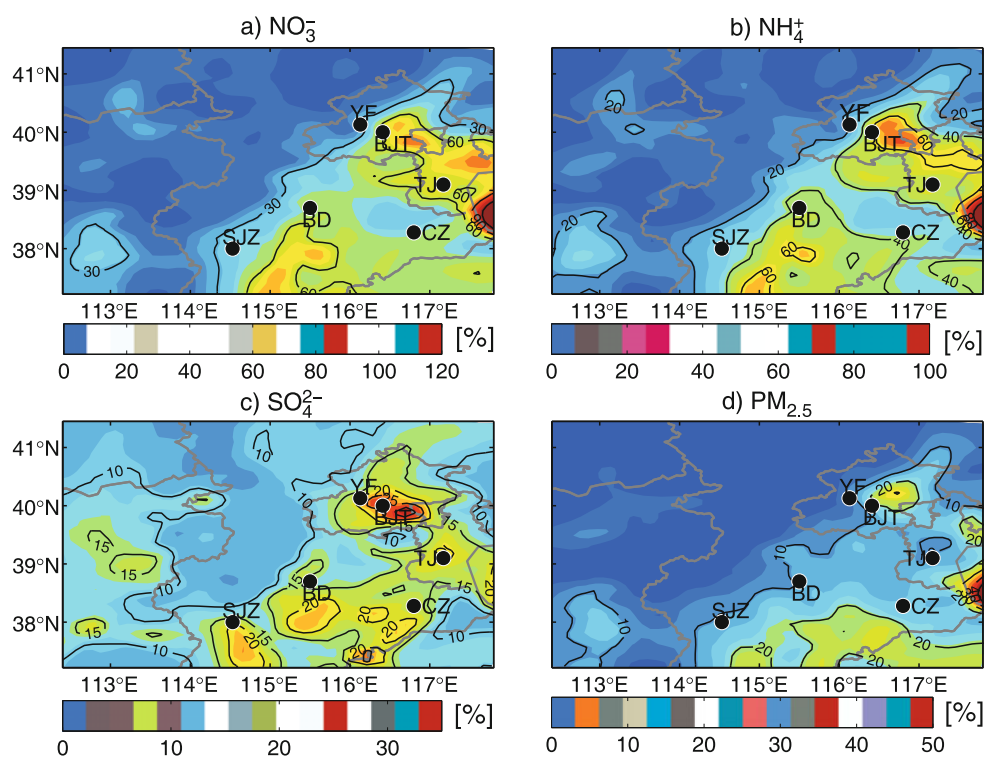
The largest hourly increases in  $NO_3^-$  and  $NH_4^+$  usually ranged from 20% to 60%. Those of  $SO_4^{2-}$  and  $PM_{2.5}$  were generally between 10% and 25% over most areas of the North China Plain, except the Bohai Bay



**Fig. 2.** Largest enhancements of (a) daily maximum 1-h  $O_3$  and (b) monthly mean daytime (0700–1900 LST) enhancements of  $O_3$  due to the three additional HONO sources (case E–case R) in August 2007.



**Fig. 3.** Monthly mean daytime (0700–1900 LST) percentage increases of (a)  $\text{NO}_3^-$ , (b)  $\text{NH}_4^+$ , (c)  $\text{SO}_4^{2-}$ , and (d)  $\text{PM}_{2.5}$  due to the three additional HONO sources in the North China Plain.



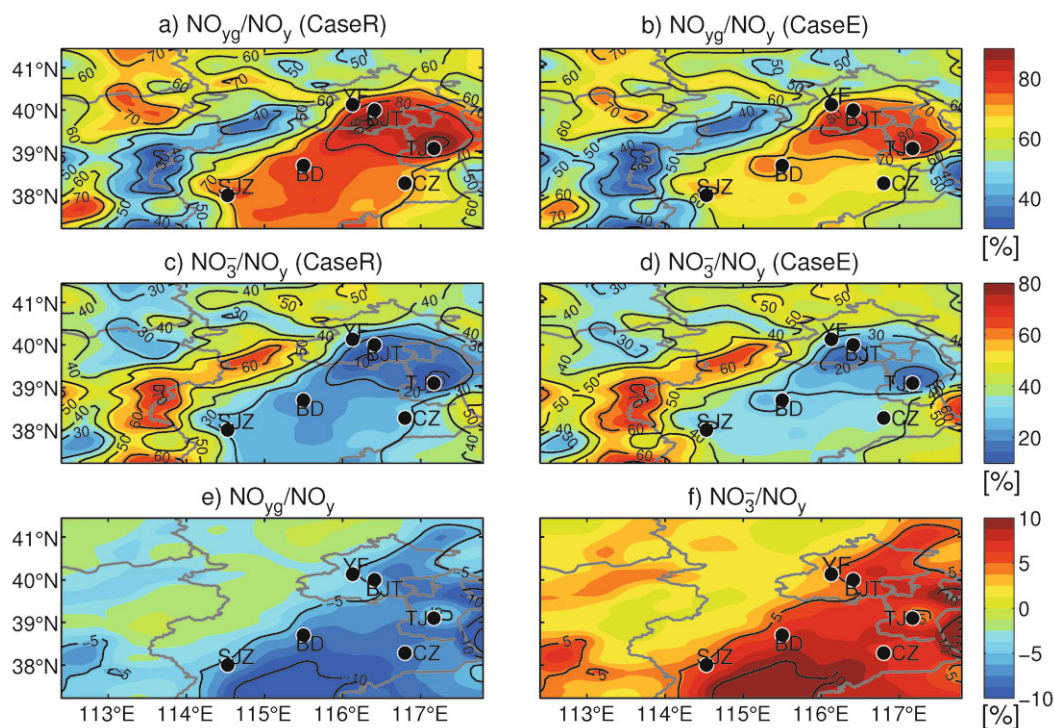
**Fig. 4.** Largest hourly percentage increases of (a)  $\text{NO}_3^-$ , (b)  $\text{NH}_4^+$ , (c)  $\text{SO}_4^{2-}$ , and (d)  $\text{PM}_{2.5}$  due to the three additional HONO sources in the North China Plain in August 2007.

(Fig. 4). For the Bohai Bay, the largest hourly enhancements of  $\text{NO}_3^-$ ,  $\text{NH}_4^+$ , and  $\text{PM}_{2.5}$  >90%, 80%, and 40%, respectively (Fig. 4). Compared with inland areas of the North China Plain, the Bohai Bay has low heights of the atmospheric boundary layer (ABL) and abundant moisture (An et al., 2011). Low heights of the ABL can increase surface concentrations of air pollutants (e.g., PM and  $\text{NO}_x$ ). Elevated PM concentrations and abundant moisture can enhance the aerosol surface area per unit volume through aerosol hygroscopic growth and lead to more production of HONO and  $\text{HNO}_3$  through the reaction of  $R_{\text{het}}$ . Additionally, high  $\text{NO}_x$  concentrations and abundant moisture are favorable for  $R_{\text{NO}_2\text{S}}$  which produces OH and HONO in the daytime. Large increases in  $\text{HO}_x$  ( $\text{OH} + \text{HO}_2$ ) concentrations due to the additional HONO sources and high emissions of  $\text{NH}_3$  (Streets et al., 2003) resulted in significant enhancements of  $\text{NO}_3^-$  and  $\text{NH}_4^+$ , which led to enhancements of  $\text{PM}_{2.5}$  (Table 3).

### 3.4 Impacts of the three additional HONO sources on $\text{NO}_{yg}/\text{NO}_y$ and $\text{NO}_3^-/\text{NO}_y$

For case R, high emissions of  $\text{NO}_x$  led to large monthly mean daytime concentration ratios of  $\text{NO}_{yg}/\text{NO}_y$  in major cities over the North China Plain

(e.g., Beijing, Tianjin, Baoding, and Shijiazhuang, and other locations), with a maximum of 91% in Beijing (Fig. 5a). Spatial distribution of  $\text{NO}_{yg}/\text{NO}_y$  for case E (Figs. 5a and b) is very similar to that for case R (Fig. 5b), but  $\text{NO}_{yg}/\text{NO}_y$  showed no noticeable decreases when Figs. 5a and b were compared. By contrast, large concentration ratios of  $\text{NO}_3^-/\text{NO}_y$ , with a maximum of 71%, were located in relatively low  $\text{NO}_x$  emission areas (Fig. 5c), because nitrates in the air mainly come from chemical reactions. Increases in concentration ratios of  $\text{NO}_3^-/\text{NO}_y$  were obvious when Figs. 5c and d were compared. Differences between Figs. 5a and b and between Figs. 5c and d indicate that the three additional HONO sources produced >5% of the monthly mean daytime decreases in  $\text{NO}_{yg}/\text{NO}_y$  and >5% of the monthly mean daytime increases in concentration ratios of  $\text{NO}_3^-/\text{NO}_y$  in large areas, with high emissions over the North China Plain with a maximum decrease of  $\text{NO}_{yg}/\text{NO}_y$  and a maximum increase of  $\text{NO}_3^-/\text{NO}_y$  of ~20% (Figs. 5e and f). In the nighttime,  $R_{\text{het}}$  and  $R_{\text{em}}$  led to ~3%–6% reductions in  $\text{NO}_{yg}/\text{NO}_y$  and ~3%–6% increments in  $\text{NO}_3^-/\text{NO}_y$  in large areas of the North China Plain (figures not shown). The maximum increases of  $\text{NO}_3^-/\text{NO}_y$  in the daytime were higher than those in the nighttime because  $R_{\text{NO}_2\text{S}}$ ,  $R_{\text{het}}$ , and  $R_{\text{em}}$  in the daytime produced greater in-



**Fig. 5.** Monthly mean daytime (0700–1900 LST) concentration ratios (%) of (a, b)  $\text{NO}_{yg}$  to  $\text{NO}_y$  and (c, d)  $\text{NO}_3^-$  to  $\text{NO}_y$  for cases R and E, and differences (%) in the monthly mean daytime concentration ratios of (e)  $\text{NO}_{yg}$  to  $\text{NO}_y$  and (f)  $\text{NO}_3^-$  to  $\text{NO}_y$  between cases E and R over the North China Plain in August 2007.

creases of  $\text{NO}_3^-$  than  $R_{\text{het}}$  in the nighttime (Table 3). For hourly mean concentration ratios, the largest decreases of  $\text{NO}_{yg}/\text{NO}_y$  and the maximum increases of  $\text{NO}_3^-/\text{NO}_y$  were between 15% and 52% in most areas of the North China Plain, except the Bohai Bay where the largest decreases of  $\text{NO}_{yg}/\text{NO}_y$  and the maximum increases of  $\text{NO}_3^-/\text{NO}_y$  reached 52% (figures not shown) due to low heights of the ABL and rich moisture.

Substantial increases in monthly mean concentrations, maximum hourly mean concentrations of aerosol components, particularly  $\text{NO}_3^-$  and  $\text{NH}_4^+$ , noticeable increases in monthly mean concentration ratios and the maximum hourly mean concentration ratios of  $\text{NO}_3^-/\text{NO}_y$  due to the three additional HONO sources have significant implications. High emissions of  $\text{NO}_x$  and PM produce elevated concentrations of HONO or  $\text{HNO}_3$  through  $R_{\text{em}}$  or  $R_{\text{NO}_2\text{S}}$  or  $R_{\text{het}}$ . Elevated HONO concentrations can yield more OH radicals in the daytime, leading to more formation of  $\text{O}_3$ ,  $\text{NO}_3^-$ ,  $\text{NH}_4^+$ ,  $\text{SO}_4^{2-}$  and  $\text{PM}_{2.5}$  under high emissions of  $\text{NO}_x$ ,  $\text{NH}_3$ , and  $\text{SO}_2$ . At night,  $R_{\text{het}}$  and the reaction of  $\text{HNO}_3 + \text{NH}_3 \rightarrow \text{NH}_4\text{NO}_3$  result in increases of  $\text{NO}_3^-$ ,  $\text{NH}_4^+$ , and  $\text{PM}_{2.5}$ . Compared with that of gas  $\text{HNO}_3$ ,  $\text{NH}_3$ , and  $\text{SO}_2$ , dry deposition velocity of  $\text{NO}_3^-$ ,  $\text{NH}_4^+$ , and  $\text{SO}_4^{2-}$  is small (Zhang et al., 2004), so  $\text{NO}_3^-$ ,  $\text{NH}_4^+$ , and  $\text{SO}_4^{2-}$  are favorable for long-range transport which aggravates regional air pollution. On the other hand, elevated concentrations of  $\text{NO}_3^-$ ,  $\text{NH}_4^+$ , and  $\text{SO}_4^{2-}$  together with their hygroscopic growth can cause visibility impairment and enhance the incidence of haze under stagnant weather conditions.

#### 4. Conclusions

Three additional HONO sources (i.e.,  $R_{\text{NO}_2\text{S}}$ ,  $R_{\text{het}}$ , and  $R_{\text{em}}$ ) were incorporated into WRF-Chem and five simulations were conducted over the North China Plain in August of 2007. Results show that HONO simulations were substantially improved when  $R_{\text{NO}_2\text{S}}$ ,  $R_{\text{het}}$ , and  $R_{\text{em}}$  were included in WRF-Chem (e.g., the root mean square error decreased by  $\sim 50\%$ , and the correlation coefficient improved to 0.97 from 0.19).  $R_{\text{het}}$  significantly enhance concentrations of HONO,  $\text{NO}_3^-$ ,  $\text{NH}_4^+$ , and  $\text{PM}_{2.5}$  in the North China Plain. The three additional HONO sources yielded significant increases in monthly mean concentrations of major aerosol components, particularly  $\text{NO}_3^-$  and  $\text{NH}_4^+$ , with a maximum of  $\sim 50\%$ , in major cities of the North China Plain (i.e., Beijing, Tianjin, Baoding, and Shijiazhuang). The three additional HONO sources produced  $\sim 5\%$ – $20\%$  increases in monthly mean concentration ratios of  $\text{NO}_3^-/\text{NO}_y$ ,  $\sim 15\%$ – $52\%$  increases in the largest hourly mean concentration ratios of

$\text{NO}_3^-/\text{NO}_y$ , and  $\sim 20\%$ – $60\%$  hourly increases of  $\text{NO}_3^-$  and  $\text{NH}_4^+$  in most areas of the North China Plain except the Bohai Bay, where largest hourly enhancements of  $\text{NO}_3^-$ ,  $\text{NH}_4^+$ , and  $\text{PM}_{2.5}$   $> 90\%$ ,  $80\%$ , and  $40\%$ , respectively. These results have important implications for atmospheric projections in that regional air pollution will be aggravated, visibility will be impaired, and the incidence of haze will be increased in some regions with elevated emissions of  $\text{NO}_x$  and PM under favorable weather conditions when the three additional HONO sources are considered.

The uncertainty of  $R_{\text{het}}$  comes mainly from the heterogeneous uptake coefficient of  $\text{NO}_2$  ( $\gamma$ ) and the aerosol surface area per unit volume ( $S_a$ ). The  $\gamma$  value ranges from  $10^{-6}$  to  $10^{-3}$  (Jacob, 2000) and can depend on air temperature, relative humidity, and aerosol components. The modeling study conducted by Bian and Zender (2003) used a  $\gamma$  value of  $4.4 \times 10^{-5}$  for mineral dust. A value of  $10^{-4}$  recommended by Jacob (2000) was used for considered aerosols in this research (see section 2.2). Field and laboratory experiments need to be done to reduce the uncertainty of the  $\gamma$  values. The uncertainty of  $S_a$  is closely related with that of emission inventories, meteorological conditions, and aerosol hygroscopic growth. For HONO emissions we have considered the suggestions of both Kurtenbach et al. (2001) and Gutzwiller et al. (2002), and we have used a reasonable expression ( $R_{\text{em}}$ , see section 2.2). The uncertainty of  $R_{\text{NO}_2\text{S}}$  mainly originates from that of the rate constant for the reaction of electronically excited  $\text{NO}_2$  with water. For this research, the mean value of  $9.1 \times 10^{-14} \text{ cm}^3 \text{ molecule}^{-1} \text{ s}^{-1}$  was used (Li et al., 2011). Further experiments are required to reduce the uncertainty of the rate constant.

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