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Latitudinal dependence of body condition, growth rate, and stable isotopes of juvenile capelin (*Mallotus villosus*) in the Bering and Chukchi Seas

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Abstract Capelin occupy a key trophic role and have a broad latitudinal distribution in the northeastern Pacific and Arctic Oceans. Understanding their adaptation to a range of conditions is important to predicting how they will respond to climate change. To quantify the variation in body condition in different physical environments, we measured energy density, RNA/DNA ratios, carbon and nitrogen stable isotope ratios in 62 juvenile capelin along the Western Alaskan coast from Bristol Bay to Point Barrow ranging across approximately 14° of latitude. Energy density correlated positively with latitude, whereas RNA/ DNA (instantaneous growth index) was strongly correlated with sea surface temperature, indicating that optimal growth of capelin was achieved at ~9 °C, followed by rapid decreases in RNA/DNA ratios at higher temperatures. δ^{13} C and δ^{15} N had strong, inverse nonlinear relationships with latitude. Depletion of $\delta^{13}C$ seen in capelin North of Bristol Bay may be related to the incorporation of allochthonous basal resources into the diets of juvenile capelin from nearby riverine inputs. Observed enrichment of $\delta^{15}N$ North of Bristol Bay is likely to be related to incorporation of higher trophic level prey items. Given inverse relationship between δ^{13} C and δ^{15} N, these prey items are likely

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Mark B. Barton Mbart034@fiu.edu available due to the increased diversity of basal resources from increased inputs of riverine organic material.

Keywords Capelin · Latitude · Isotopes · Nitrogen · Carbon · Food web · Energy allocation · Growth · Lipid · RNA/DNA

Introduction

Latitudinal variation exists across seascapes and some species found across a wide range of climatic conditions have different feeding and life-history strategies according to the conditions they face (Conover 1988; Post and Parkinson 2001; Shoji et al. 2011; Rypel 2012). Life-history strategy is defined as the allocation of energy throughout a lifetime to optimize growth, survival, and reproduction (Noordwijk and Jong 2014). Energy allocation in juvenile fish is particularly important, when there are concurrent energetic demands to grow (predator avoidance, increase accessibility to prey) and store energy to survive their first winter (Mogensen and Post 2012). Fish condition is often used as a parameter to examine survival potential and recruitment processes (Heintz et al. 2013) and is often expressed as measures of energy density (kJ/g), which is driven by lipid content (Anthony et al. 2000). Stable isotopes have become a popular method for examining dietary sources and trophic structure within species and their ecosystems (DeNiro and Epstein 1978, 1981; Hobson and Welch 1992; Layman et al. 2007a; Peterson and Fry 2014). Many trophic studies focus on carbon and nitrogen isotopes, but they rarely consider simultaneous measures of energetic condition and growth. These latter measures offer important information describing the quality of prey items and consequently diet implications for fish

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condition and optimality of habitats within a system (Sherwood et al. 2007). Thus, food web structures and carbon sources that maximize productivity can be identified by combining isotopic analysis with indices of growth and energetic condition. This is important in the Arctic and sub-Arctic where marine conditions are changing rapidly (Grebmeier et al. 2006), and those changes are likely to affect prey availability and quality.

As coastal waters in the Arctic become increasingly warmer and ice-free due to shifting climate conditions, it is likely that species will extend, restrict, or shift their ranges as a function of changing tolerable conditions, eliciting change in coastal food web structures of northern oceans (Grebmeier et al. 2006; Moline et al. 2008; Eisner et al. 2013). These coastal food webs provide critical subsistence fisheries and serve as a food source for numerous endangered and protected marine mammals and sea birds (Hobson and Welch 1992). With these changes occurring, high-latitude species may see intrusions from lower-latitude species, introducing new predators and competition for resources and space (Gilman et al. 2010; Sorte et al. 2010; Grebmeier 2012; Litzow and Mueter 2014). The lifehistory strategies of species that have wide tolerance of environmental conditions may serve as proxies to examine fish condition with changes in food web structure across extensive spatial domains and offer insight into predicting how these important forage fish may fare in the Arctic as climatic conditions change.

The Western Coast of Alaska is bordered by the Bering Sea in the southwest and the Chukchi Sea in the northwest. The high-nutrient, high-salinity Bering Shelf Water (BSW) and the freshwater runoff and wind-driven Alaska Coastal Current (ACC) run along the entire length of these coasts, and the front between them typically occurs between 30 and 60 km from shore (Stabeno et al. 1995; Steele et al. 2010). The front between these two distinct water masses may generate optimal growth conditions to support an abundance of planktivorous fish in coastal waters (Coyle and Pinchuk 2005; Weingartner et al. 2005; Hopcroft et al. 2010; Eisner et al. 2013). Because of the extensive latitudinal range of this persistent front, it presents the opportunity to study the effects of latitude on trophic interactions and energy allocation strategies of coastal planktivorous fish. This water mass provides gradients of temperature and freshwater influx, two primary conditions that are expected to increase significantly with climate change (Peterson et al. 2002; Chan et al. 2011; Dunton et al. 2014). The proximity of this coastal current to land allows for the examination of fish trophodynamics and condition in response to freshwater/terrestrial influence.

Capelin (*Mallotus villosus*) are an abundant planktivorous forage fish found throughout western Alaskan waters and serve as an important link between lower and higher trophic levels (Gjøsæter and Båmstedt 1998; Johnson et al. 2010; Sherwood et al. 2007). Capelin's wide geographical distribution suggests plasticity in its life-history strategies. Because of capelin's trophic position and value as a prev resource, their ability to adapt to different conditions and prey resources will probably play an important role in the response of larger, economically important or endangered members of higher trophic levels, including fish, marine mammals, and sea birds. Previous research has shown that great variability exists in planktonic assemblages throughout the ACC and BSW (Schell et al. 1998; Dunton et al. 2006; Eisner et al. 2013) and it is likely that distinct planktonic assemblages coincide and interact with distinct food webs. Capelin can be used as a model species for the adaptability of mid trophic levels (forage fish) to changing climatic conditions in the Arctic by examining their energy allocation patterns and trophodynamics across gradients in climate and latitudes in which they are abundant.

Past studies have examined the latitudinal dependence of energy allocation strategies of fish, but to date they have largely been restricted to smaller scales or closed freshwater and have not incorporated stable isotope analysis (SIA) to examine the effects of variations in diet (Shoji et al. 2011; Mogensen and Post 2012; Rypel 2012; Siddon et al. 2013). The use of a combination of energetics and stable isotope data may offer further insight into the adaptability and sensitivity of these coastal food webs to changing conditions throughout the Alaskan coastline. We aim to quantify energy density, growth rates, and isotope ratios among juvenile capelin along the entire west coast of Alaska within a single water body, the ACC/BSW front, to examine how variation in condition and growth rate varies as a function of latitude, temperature, riverine influence, and trophic and dietary composition.

Materials and methods

Fish collections

Juvenile capelin (n = 62) were collected during the summer fall of 2012 using three different vessels spanning 13.9° of latitude (57.5°N–71.4°N; Fig. 1). Capelin were selectively subsampled from expansive fish surveys from stations between 25 and 75 km from the coastline (with exception of the stations in Bristol Bay that were approximately 100 km from the coast) and thus likely to be near the ACC/BSW front (Stabeno et al. 1995; Steele et al. 2010). Capelin analyzed for this project are also limited to those <100 mm length and are assumed to be juveniles based on their size (Vesin et al. 1981; Hop and Gjøsæter 2013) and lack of adult pigmentation. Capelin were sampled using a 198-m-long surface trawl towed behind a

Fig. 1 Map of Alaskan waters showing sampling locations of juvenile capelin collected. Filled shapes represent the Chukchi Sea stations, and open shapes represent Bering Sea stations. Like symbols represent regional groups of fish at similar latitudes. The numbers in parentheses after each river represent the individual annual discharge $(km^3 vear^{-1})$ of each river (Benke and Cushing 2006; Nemeth et al. 2014). This map was produced with the CRAN-R package "Rgooglemaps" (Loecher and Ropkins 2015)



170°0'0"W 160°0'0"W 150°0'0"W 140°0'0"W

54.9-m chartered fishing vessel between August 8th and September 21st, 2012. The trawl had hexagonal mesh on the wings and body, a 1.2 cm mesh cod end liner, and a $50 \text{ m} \times 25 \text{ m}$ mouth (horizontal \times vertical). Each tow lasted for 30 min at approximately 8.3 km h^{-1} at stations within a 103 km² grid along the western Alaskan coast between 60°N and 71.4°N. Four more sites were sampled using the same methods by the NOAA ship Oscar Dyson between 57.5°N and 60°N (August 20th-October 9th, 2012). Capelin along the nearshore were sampled with beach seines near Barrow, AK (\sim 71°N), during August 7-20, 2012. The seine was 37 m long with variable mesh sizes (10 m of 32 mm outer panels, 4 m of 6 mm middle panels, and 9 m of 3.2 mm blunt panel). Each set was round-haul style, paid out of a 7-m skiff following methods used by Johnson et al. (2010). All collections occurred during daylight hours. Physicochemical parameters from offshore stations were averaged from the top 20 m of the water column using a CTD at each sampling station at the time of collection. Water temperature and salinity from beach seine sites were measured from the top 0.5 m of water using a thermometer and refractometer. Fish were measured to fork length (FL) and kept frozen until analyzed in the laboratory.

Individual capelin were randomly subsampled within each site (Table 1). In the laboratory, individual fish wet weights were measured, stomach contents removed, and a sample of white muscle of ~0.01 g was dissected and frozen (-80 °C) for RNA/DNA analysis. Individual capelin were dried to a constant weight using a LECO

Thermogravimetric Analyzer (TGA) 601/701 and homogenized using mortar and pestle until a uniform consistency was reached. Dry homogenates of individual juvenile capelin were stored in a desiccator prior to SIA and bomb calorimetry analyses.

Bomb calorimetry

Energy density of juvenile capelin (kJ g^{-1} dry mass) was quantified using bomb calorimetry. A Parr Instrument 6725 semi-micro bomb calorimeter was used to combust pellets of dry fish homogenate following standard instrument operating protocols from the manufacturer. Precision and accuracy of measurements were assessed by evaluating duplicate benzoic acid standards, replicate samples, and a tissue reference material of Pacific herring or walleye pollock homogenate. Error limits were set for the quality assurance samples, where precision from replicate benzoic acid standards was not allowed to vary by more than 1.5 % coefficient of variation and must have been within 2.0 % of the target value. Sample replicates were not allowed to vary by more than 1.5 standard deviations, and tissue reference samples were not allowed to vary by more than 3.0 % from target reference values.

RNA/DNA analysis

Instantaneous growth rates were estimated from RNA/ DNA ratios following methods outlined in Sreenivasan (2011). A ~ 10 mg of frozen muscle sample was taken

Group name	Symbol	Sample date	Sample gear	Latitude range (°N)	N fish	N stations	Fish length (mm) mean (SD)	SST (°C) mean (SD)
Pt. Barrow		Aug 18–19	Beach Seine	71.09–71.39	11	5	72.8 (8.1)	11.2 (0.2)
Wainwright Inlet	••	Aug 19–20	Surface Trawl	70.50–70.52	11	3	88.4 (7.3)	7.3 (0.6)
Point Hope		Aug 13–14	Surface Trawl	68.54	7	1	89.7 (6.0)	9.9 (-)
Kotzebue Sound	**	Aug 8–10	Surface Trawl	66.57–67.04	11	2	89.5 (8.0)	9.5 (0.5)
Norton Sound	$\nabla \nabla$	Sept 12-14	Surface Trawl	63.97–64.51	8	2	85.4 (10.9)	5.95 (0.0)
Nunivak Island		Sept 10-20	Surface Trawl	60.01-61.98	11	2	82.2 (7.9)	7.1 (0.1)
Bristol Bay	$\diamond \diamond$	Aug 21–26	Surface Trawl	57.52–58.49	3	2	81.3 (7.8)	8.7 (0.9)

Table 1 Details of sample sizes and regional group differentiation

Throughout this paper, "regional group" refers to the seven groups differentiated by symbols

from each fish. RNA/DNA ratios were quantified fluorometrically using one dye and two enzymes (RNase and DNase; Caldarone et al. 2001). Nucleic acids were isolated from the smaller muscle samples and dyed using 75 µL ethidium bromide (5 μ g ml⁻¹) according to the protocol outlined by Caldarone et al. (2001). Total fluorescence at excitation and emission wavelengths of 355 and 600 nm, respectively, was recorded, then the samples were sequentially treated with RNase and DNase, and the resulting reduced fluorescence was measured to obtain RNA and DNA fluorescence, respectively. Standard curves were constructed using serial dilutions of 18 s-28 s rRNA (Sigma R-0889) and calf thymus DNA (Sigma D-4764) standards. DNA concentrations in tissues are stable, but RNA concentrations vary greatly depending on the rate of protein synthesis where a high RNA/DNA ratio indicates a high growth rate (Weber et al. 2003).

Stable isotopes analysis

SIA of carbon and nitrogen is used to examine the origins and type of dietary sources assimilated by fish (DeNiro and Epstein 1978, 1981). All subsamples of dried fish homogenate were weighed to 0.55 ± 0.15 mg. In between every four samples, a standard or duplicate sample was analyzed to examine precision of measurements. Samples were analyzed at the Florida International University SERC Stable Isotope Laboratory using elemental analysis–isotope ratio mass spectrometry (EA-IRMS), with a NA1500 NC (EA) coupled to a Delta C (IRMS). Error based on internal glycine standards ranged 0.09–0.21 ‰ for δ^{15} N and 0.07–0.10 ‰ for δ^{13} C.

Lipid corrections were computed using C/N ratios (4.45 \pm 0.66) following "Eq. 1" outlined by Logan et al. (2008). This equation requires the assumption that the difference between bulk δ^{13} C and lipid-free δ^{13} C approximates 6 ‰ as suggested by McConnaughey and McRoy (1979). Later work by Post et al. (2007) pointed out that these methods are suitable for organisms with 15 % or less lipid content, and that caution should be used at higher lipid contents because they had insufficient samples with such high-lipid content to adequately model the relationship. Thus, using the relationship presented by Post et al. (2007), we estimate that approximately 82 % of our samples contain less than 15 % lipids, 15 % of our samples contain <17.5 % lipid, and the remaining 3 % contain <20 % lipid. Based on this information, we deemed this method of lipid correction appropriate for our samples.

Stomach contents

Prey from the stomach contents of juvenile capelin was identified to species and life-history stage where possible using methods outlined by Sturdevant et al. (2012). Adult *Calanus spp.* were classified by size: small \leq 2.4 mm in length, medium = 2.5–2.9 mm, and large \geq 3 mm. Aggregate wet weights of separated prey groups were measured from each fish. Weights were converted into percent contributions to the total mass of prey found in each stomach to standardize against the unevenness in fullness. Percent contributions from individual capelin were averaged over regional groups (see Data analysis section below) to standardize against unevenness in sample size.

Data analysis

The 62 juvenile capelin from 17 sampling stations were separated into 7 regional groups (Table 1) based on latitude and distance from each other (Fig. 1). General Additive Models (GAMs) were used for qualitative assessment to identify patterns of dependent variables with latitude as the independent variables. Linear regressions were used to assess the relationship between dependent variables (δ^{15} N, δ^{13} C, energy density, RNA/DNA ratios, length, and weight) and to offer a quantitative approach to assessing the relationships of energetics and SIA with latitude.

Results

Energy allocation

Across the sample set, energy density ranged from 20.77 to 26.49 kJ g⁻¹ (Mean \pm SD = 22.85 \pm 1.33 kJ g⁻¹). RNA/ DNA ratios ranged from 13.10 to 31.61 (22.89 \pm 4.33). Linear regression indicated a weak positive correlation between energy density and RNA/DNA ($R^2 = 0.32$, p < 0.0001; Fig. 2a). Energy density increased significantly with latitude, but RNA/DNA did not (energy density: *estimated degrees of freedom [edf] = 4*, $R^2 = 0.34$,



Fig. 2 Linear regression between energy density and RNA/DNA ratios was strongly correlated (A; $R^2 = 0.3163$, p < 0.0001). A linear regression between δ^{13} C and δ^{15} N shows a strong negative correlation (B; $R^2 = 0.6747$, p < 0.0001)



Fig. 3 General additive models for energy density (**a**), RNA/DNA ratios (**b**), δ^{13} C (**c**), and δ^{15} N (**d**) for all juvenile capelin with a smoothing function on latitude. All models created using cubic splines were used with five knots (edf = 4) for energy density and RNA/DNA ratios, and four knots (edf = 3) for δ^{13} C and δ^{15} N. The *dashed lines* represent the 95 % confidence interval. Each model was significant with varying levels of deviance explained (A: $R^2 = 0.34$, p < 0.0001; B: $R^2 = 0.14$, p = 0.01; C: $R^2 = 0.54$, p < 0.0001; D: $R^2 = 0.67$, p < 0.0001). The *bold tick* marks above the x-axis represent latitudes at which fish samples were collected

p < 0.0001, Fig. 3a; RNA/DNA: edf = 4, $R^2 = 0.14$, p = 0.01, Fig. 3b). The models for energy density and RNA/DNA suggest that the Point Barrow group may be an outlier relative to other samples, as the values rapidly decrease at this region group. In addition, capelin from Point Barrow were significantly smaller than all other groups (FL: p = 0.0002; wet weight: p < 0.0001) and were the only fish collected by beach seines in the very near shore. When the Point Barrow group was removed from the analysis, energy density was linearly correlated with latitude ($R^2 = 0.36$, p < 0.0001, Fig. 4a); however, the linear model between RNA/DNA and latitude was not significant ($R^2 = 0.06$,



Fig. 4 Linear regressions of energy density (a), RNA/DNA ratios (b), δ^{13} C (c), and δ^{15} N (d) with latitude. The northernmost sampling stations (Point Barrow group: *filled square*) and the southernmost sampling stations (Bristol Bay group: *open diamond*) were removed as outliers for the analysis. Energy density was strongly correlated ($R^2 = 0.36$, p < 0.0001) with latitude, but RNA/DNA ratios were not ($R^2 = 0.06$, p = 0.05). δ^{13} C is negatively correlated, and δ^{15} N is positively correlated with latitude ($R^2 = 0.55$, p < 0.0001; $R^2 = 0.75$, p < 0.0001; respectively). Station symbology: Wainwright Inlet (*filled circle*), Point Hope (*filled triangle*), Kotzebue Sound (filled diamond), Norton Sound (*inverted triangle*), and Nunivak Island (*open square*)

p = 0.05, Fig. 4b). Furthermore, linear regressions indicated that energy density was strongly positively correlated with fish length (FL: $R^2 = 0.47$, p < 0.0001) and full body wet weight ($R^2 = 0.63$, p < 0.0001).

In contrast, GAM models using sea surface temperature as a predictor indicated that more variability in RNA/DNA ratios was explained by surface temperature than in energy density (RNA/DNA: edf = 3, $R^2 = 0.28$, p < 0.0001; energy density: edf = 4, $R^2 = 0.15$, p < 0.0001, Fig. 5ab); however, both models suggest that these measures vary



Fig. 5 General additive models for energy density (**a**), RNA/DNA ratios (**b**), δ^{13} C (**c**), and δ^{15} N (**d**) for all juvenile capelin with a smoothing function on temperature. All models created using cubic splines were used with five knots (edf = 4) for energy density, and four knots (edf = 3) for RNA/DNA ratios, δ^{13} C, and δ^{15} N. The *dashed lines* represent the 95 % confidence interval. All models were significant with varying levels of deviance explained (A: $R^2 = 0.15$, p < 0.0001; B: $R^2 = 0.28$, p < 0.0001; C: $R^2 = 0.31$, p < 0.0001; D: $R^2 = 0.30$, p < 0.0001). The *bold tick marks* above the x-axis represent latitudes at which fish samples were collected

slightly with increasing temperature until a threshold of approximately 9 °C is reached, at which point they both decreased rapidly. In turn, a GAM using surface temperature with latitude as the predictor indicates that these factors have a nonlinear relationship (edf = 4, $R^2 = 0.29$, p < 0.0001, Fig. 6).

Stable carbon and nitrogen isotopes

Stable δ^{13} C and δ^{15} N isotope ratios of juvenile capelin were analyzed to identify changes in the origins of dietary



Fig. 6 GAM of surface temperature with a smoothing function on latitude. The nonlinear relationship shows that there is no discernible trend between surface temperature and latitude. The *dashed lines* represent the 95 % confidence intervals. Cubic splines were used with five knots (edf = 4), the GAM explains 28.8 % of the deviance, and the relationship was found to be significant (p < 0.0001). The *bold tick marks* above the x-axis represent latitudes at which fish samples were collected

material and trophic position across latitudes. Across this range, δ^{13} C ranged from -22.47 to 17.89 ‰ (mean \pm SD = -20.33 \pm 1.10 ‰) and δ^{15} N ranged from 11.76 to 17.09 ‰ (14.31 \pm 1.29 ‰). The linear correlation between δ^{13} C and δ^{15} N was strongly negative ($R^2 = 0.67$, p < 0.0001; Fig. 2B).

The relationships between δ^{13} C and δ^{15} N with latitude were highly significant, and much of the variability in both SIA measures could be explained by latitude (δ^{13} C: $edf = 3, R^2 = 0.54, p < 0.0001; \delta^{15}$ N: $edf = 3, R^2 = 0.67, p < 0.0001;$ Fig. 3C-D). Linear regressions indicated that δ^{13} C was positively correlated and δ^{15} N was negatively correlated with latitude ($R^2 = 0.55, p < 0.0001; R^2 = 0.75, p < 0.0001,$ respectively; Fig. 4c–d) when anomalies at the northernmost sampling stations (Point Barrow group) and the southernmost sampling stations (Bristol Bay group; Fig. 1) are removed. A similar pattern for each isotope was seen with temperature, suggesting that peak energy content and growth rate at the high latitudes coincide with feeding at higher trophic levels and incorporation of depleted δ^{13} C.

Stomach contents

Juvenile capelin diets were dominated by *Calanus spp*. copepods, with the average proportion of copepods in nonempty stomachs being 79.8 %. Stomach contents in Bristol Bay consisted solely of large- and medium-sized *Calanus spp*. copepods, but stations near the Bering Strait had increased prey diversity, including important contributions from decapod larvae, and smaller contributions of small copepods, cladocerans, and chaetognaths. At the Point Barrow stations, the diets consist completely of small copepods and *Themisto libellula*, a predatory hyperiid



Fig. 7 Average contributions by weight of prey type in the stomach contents of the analyzed capelin for six of the seven regional groups identified in Table 1. Cal = *Calanus spp.* copepod. No stomach content data were available for the Norton Sound group

amphipod (Auel and Werner 2003; Pinchuk et al. 2013). Of the capelin caught at Point Barrow, 56 % of them had empty stomachs, while empty stomachs were not observed at any other station (Fig. 7).

Discussion

Energy allocation

Given that capelin occupy a wide latitudinal range, it is expected that they exhibit plasticity not only in their diet but also the allocation of energy obtained to cope with differences in climatic conditions. When prey resources are limited, we expect to see a trade-off between energy storage and growth rate as energy availability is generally not great enough to allow for concurrent processes of high growth rates and energy storage in juvenile fish (Post and Parkinson 2001). We saw evidence for increased energy provisioning for more severe winters in juvenile capelin in higher latitudes, with a positive correlation between energy density and latitude. This phenomenon has been observed for other fish species where the longer high-Arctic winters require a larger store of energy to survive than shorter, lowlatitude winter locations (Biro et al. 2005).

It is also plausible that though the fish were sampled at approximately the same time of year, season could play a factor in the greater energy densities at higher latitudes. Many species of fish undergo seasonal changes in their energy content as a factor of ontogenetic changes and maturation, particularly pelagic species such as capelin (Vollenweider et al. 2011). In general, fish increase energy during summer periods of high productivity, peak in the fall, and decline overwinter when prey can be scarce and gametes start developing for later spawning. The onset of winter will come much sooner to the fish in the Arctic. Therefore, it could be that higher energy densities observed at higher latitudes are a factor of the accelerated onset of winter, while further to the south there is additional time to prepare.

Increased energy storage at high latitudes could be expected to impair growth rates relative to southern areas. Another reason to expect lower growth rates in the Arctic is that capelin, as are many species of fish, are smaller at age at higher latitudes (Chambers and Leggettt 1987; Chambers et al. 1989; Olsen et al. 2005). However, measurements of RNA/DNA remained relatively constant across latitudes. In concert, higher energy densities in the Arctic and equivalent RNA/DNA across latitudes imply that fish in higher latitudes have greater accessibility to energy resources that can be simultaneously allocated to growth and energy storage. An alternate explanation is that RNA/ DNA ratios in juvenile capelin are driven more by sea surface temperature. Comparisons across a temperature range have been shown to require laboratory-based calibration studies (Caldarone et al. 2001). Without these studies to validate the temperature effect on RNA/DNA and growth relationship, it is conceivable that temperature is confounding our use of RNA/DNA as an index to growth.

However, RNA/DNA ratios varied with temperature and were greatest just above 9 °C, which occurred in the middle range of the Chukchi coast near Wainwright Inlet. This indicates that RNA production is maximized at this temperature. However, it is unclear whether this translates to increased growth. It is possible that increased RNA synthesis is a compensation for reduced efficiency in protein synthesis enzymes (Houlihan et al. 1995; Smith and Ottema 2006). Coincidently, this is also the location where energy density reached a maximum. This could suggest that growing conditions are maximized at 9 °C. High growth rates near the end of the growing season have been associated with increased lipid reserves in high-latitude perch and subsequent winter survival (Huss et al. 2008). A summer-long sampling effort of capelin in the Point Barrow region during icefree periods in 2013 and 2014 shows that capelin were more abundant at 8-9 °C (Barton et al. unpublished data). Capelin distribution, abundance, and diet are impacted by water temperature. In cold years (1-2 °C below average), capelin are distributed across a much broader area in higher numbers in the Bering Sea, and energy-rich Calanus spp. are important diet items (Andrews et al. pers. comm.). In warm years (1-1.5 °C above average), capelin distribution is relatively restricted to the cooler, northern reaches of the Bering Sea, and energy-poor Pseudocalanus spp. and Oikopleura spp. were more abundant in their diet. During August 2012, when these samples were collected, an average temperature anomaly near Point Barrow stations of 5 to 7.5 °C above average was measured, suggesting that an abundance of energy-poor prey items was likely present (Parkinson and Comiso 2013).

Similarly, capelin distribution in Glacier Bay in Southeast Alaska was highly correlated with water temperatures. but in contrast to our results, these capelin were most abundant in the colder (6-7 °C) glacial waters over warmer (7-8 °C) estuarine central bay waters (Arimitsu et al. 2008). In this case, the glacial waters must have offered an advantage over the central bay waters. Glacial runoff brings with it high concentrations of nutrients (Hood and Scott 2008) and creates stratification by forming a freshwater lens, thus promoting plankton blooms to occur and providing an abundant food source for resident capelin. Additionally, Arimitsu et al. (2008) demonstrated that these waters were highly turbid, which may offer protection from sight-based predation, and increased feeding opportunities for capelin. Though we expect that temperature plays an important role in the life-history strategies of capelin, different temperature preferences in Southeast Alaska compared to the Western coast of Alaska in our study suggest that other factors such as predator abundance, prey availability, and turbidity may be more important drivers than temperature.

Stable carbon and nitrogen isotopes

Relationships between δ^{13} C and δ^{15} N can be used to identify a number of factors that describe the variation of assimilated materials obtained through capelin diets as a function of latitude. It is common to observe positive correlations between δ^{13} C and δ^{15} N isotope ratios because both become enriched (δ^{13} C $\approx +1 ~\%$; δ^{15} N $\approx +3.8$ %) with increasing trophic level provided that the basal resources remain the same (Hobson and Welch 1992; Layman et al. 2007a). The negative correlation seen in our results may suggest that within a single species, carbon sources and trophic structure may differ spatially. In order to better understand these differences in dietary composition of juvenile capelin, δ^{13} C and δ^{15} N ratios across the latitudinal range must be investigated individually.

Stable carbon isotopes

Basal resources in the BSW and Anadyr Water (AW) are mostly derived of pelagic production, but in the ACC basal resources consist mostly of terrestrial materials contributed by freshwater runoff (Grebmeier et al. 1988). This concept is consistent with the pattern of δ^{13} C throughout our sampling stations, suggesting that the majority of the

variation in δ^{13} C may be explained by riverine inputs along the path of the ACC. It is expected that the ACC is less depleted in δ^{13} C when entering Bristol Bay as it is derived from relatively marine waters from the Gulf of Alaska moving through Unimak Pass (Kline 1999). As it travels north between 58 and 63°N, a number of rivers discharge a substantial amount of fresh water into the ACC, including the Kvichak, Nushagak, Kuskokwim, and Yukon which cumulatively average 310 km³ year⁻¹ (40 % of the average annual freshwater runoff into the ACC; Fig. 1) (Weingartner et al. 2005: Benke and Cushing 2006). As these rivers discharge, δ^{13} C-depleted labile organic material accumulates and is incorporated into primary producers and eventually secondary consumers like capelin (Dunton et al. 2005; Helfield and Naiman 2016). As the ACC continues northward, it converges and mixes with the benthic-derived marine BSW and the AW (Grebmeier et al. 1988; Dunton et al. 2005), causing isotopic ratios to become less depleted again. The lack of major rivers that drain into the southern Chukchi coast of Alaska limits riverine inputs of terrestrial organic material north of the Bering Strait, and $\delta^{13}C$ continues to become less depleted as the ACC travels further north.

The pattern described above might suggest that δ^{13} C correlates with salinity; however, when this relationship was investigated, we found neither discernible patterns nor any significant correlation. This may be explained by the relative differences of salinity and isotopic content between riverine waters and ACC waters. We posit that the difference in salinity between riverine and coastal waters may be relatively small compared to the difference in abundance of materials that may be incorporated as basal resources, thus explaining why riverine discharge may elicit a change in δ^{13} C, but not in salinity. Given that the ACC is known to be driven by freshwater discharge and has relatively low levels of in situ production (Grebmeier et al. 1988; Weingartner et al. 2005), this phenomenon may serve as a possible explanation for the observed trends.

Freshwater inputs support observations for δ^{13} C between 58 and 70°N, but an anomaly exists at the northernmost regional groups (Wainwright Inlet and Point Barrow) of the data set where δ^{13} C becomes more depleted again. It is possible that this could be attributed to the incorporation of high-lipid content in prey items that are depleting the carbon signature. However, if this were the case we would expect capelin in these regions to show an increase in energy density, which is not the case. A more likely explanation is that the series of small streams and rivers along the Chukchi coast of the North Slope of Alaska, as well as two substantial estuaries (Wainwright Inlet and Peard Bay) that are likely to carry large loads of labile terrestrial carbon from permafrost meltwater runoff into coastal waters, could be responsible for the depletion

of δ^{13} C ratios in capelin at the northernmost stations (Dutta et al. 2006; Schuur et al. 2008). Though this runoff is relatively small compared to major rivers in the Bering Sea, when combined, these terrestrial inputs may be large enough to cause a significant shift in dietary δ^{13} C of capelin.

Stable nitrogen isotope

 δ^{13} C and δ^{15} N have inverse relationships with latitude, where δ^{15} N is less enriched in Bristol Bay, and then rapidly becomes more enriched at Nunivak Island (17.1 ‰). The ratios became gradually less enriched with latitude, reaching a minimum at Point Hope (11.8 %), at which point they become enriched again at the northernmost groups (Wainwright Inlet and Point Barrow). When investigating carbon isotope patterns, we found that our results resembled broad-scale patterns described by Schell et al. (1998); however, the fluctuations we found in nitrogen isotopes do not. The difference between our maximum and minimum $\delta^{15}N$ surpasses the commonly accepted trophic enrichment value of 3.8 % (Hobson and Welch 1992; Post 2002; Hansen et al. 2012) and suggests that capelin along this latitudinal gradient are feeding at different trophic levels. When this pattern is compared to that of δ^{13} C, it becomes clear that δ^{15} N is more enriched where terrestrial inputs are increased. This difference may be attributed to one or both of two scenarios: (1) The prey that capelin feed on are depending on different basal resources; (2) and capelin are feeding on different prey types in relation to latitude.

One possible explanation is that the basal resources vary with latitude and may elicit a cascading effect on the isotopic ratios of capelin. Fractionation of δ^{15} N differs depending on the type of nitrogen compounds used by primary producers. Atmospheric nitrogen (N₂) fixers such as phytoplankton generally have a small range of $\delta^{15}N$ (-2 to 2 %), whereas nitrate, nitrite, ammonia, and ammonium fixers such as benthic marine plants and terrestrial plants have a much greater range (-8 to 3 %), leading to a wide range of possible values of coastal basal resources (Fry 2007). As mentioned previously, the relationship between δ^{13} C and δ^{15} N suggests that δ^{15} N becomes more enriched where terrestrial inputs are highest. It is likely that the increased diversity in basal resources caused by the increased terrestrial inputs led to an increase in trophic level variation, thus supporting more trophic levels than areas with less terrestrial input (Layman et al. 2007b). This suggests that Arctic coastal food webs may gain complexity and productivity as Arctic warming continues to increase the magnitude of freshwater discharge (Peterson et al. 2002).

This logic leads us to consider the types of prey items being consumed as an explanation for patterns in nitrogen

isotopes. The majority of capelin analyzed may be classified into the lower half of the size class defined as juveniles (75-100 mm) by Vesin et al. (1981), a life stage at which gape and stomach size limits consumable prey items and variability in their diet. Dietary composition of these small capelin was dominated by Calanus copepods, with the average proportion of copepods in non-empty stomachs being 79.8 % (Fig. 7). Less common prey items including decapod larvae, chaetognaths, and the hyperiid amphipod T. libellula are only seen at the Point Barrow group. Isotopic ratios of different size copepods are not likely to differ greatly because of their low trophic position (Schell et al. 1998), but decapod larvae, chaetognaths, and T. libellula feed on copepods, thus adding a trophic level between capelin and copepods (Saito and Kiørboe 2001; Auel and Werner 2003). The inclusion of these less common prey types may be responsible for the enrichment observed in δ^{15} N. Feeding data suggest that at the latitudes where $\delta^{15}N$ became suddenly enriched (Nunivak Island and Point Barrow), capelin included less common prey items in their diet that would be expected to feed at higher trophic levels (chaetognaths and T. libellula, respectively; Fig. 7). These results suggest that higher trophic level prey items are present where terrestrial nutrients are abundant and are at least partially responsible for variations in δ^{15} N with latitude.

Fish collected at Point Barrow were anomalous in all dependent variables (δ^{13} C, δ^{15} N, energy density, and RNA/DNA). These fish were collected with beach seines and thus are inhabitants of the very nearshore where conditions can be extremely variable in comparison to the coastal offshore waters where the other fish were collected with surface trawls. If these nearshore samples are comparable with the offshore samples in lower latitudes, we may expect that the energy density of sub-Arctic forage fish in the Arctic may increase as future high-latitude conditions change to resemble current lower-latitude conditions. This may offer an alternate high-quality prey source for Arctic piscivores as climate change continues.

However, it is also possible that these nearshore fish are not comparable with offshore samples and thus were removed from most analyses as they were significantly smaller in size than all other samples. The Point Barrow fish were significantly lower in energy density (p = 0.0004) and RNA/DNA (p = 0.0004) than capelin examined from other areas. When examining the other regional groups, energy density increases from Bristol Bay to Wainwright Inlet and RNA/DNA remains relatively constant; however, at Point Barrow both of these measures fell to some of the lowest values observed. Furthermore, diet data indicate that most of the nearshore fish had empty stomachs (61 %) suggesting that prey are sparsely distributed, and these fish may be undernourished or have difficulty locating prey. The anomalous enrichment of δ^{15} N in these fish may support this premise. During starvation, animals will metabolize their own fat and muscle tissue to survive; in essence, they are eating themselves. This causes trophic fractionation and will make it appear as though they are feeding at a higher trophic level (Vander Zanden and Rasmussen 2001). The combination of poor condition and growth, coupled with evidence of low food availability and potential starvation, suggests that the Arctic nearshore habitats near Barrow may not be an optimal environment for juvenile capelin.

If these nearshore habitats are suboptimal, why were capelin highly abundant? These had the highest surface temperatures (11 °C) and were much higher than the optimal growth temperatures (~9 °C), which is not likely to motivate capelin to inhabit these waters. One explanation for their abundance is that these shallow and turbid nearshore waters offer an advantage over nearby habitats such as refuge from abundant local predators (belugas, seals, and sea birds), which are commonly observed feeding in the nearshore. However, it is also possible that juvenile capelin are blown or advected into suboptimal nearshore areas by strong wind and currents, and therefore, their presence in the Arctic nearshore near Barrow is not by choice.

Conclusions

The combination of growth and condition indices with SIA is a useful method to understand how isotopic food sources contribute to fish production. We suggest that capelin may be an ideal species for such analysis and can provide insight into how the ecosystem may restructure under different climate scenarios, as they have a key trophic position in the ecosystem and a wide latitudinal range. A benefit of using a wide-ranging species such as capelin is that their response to different environmental conditions can be considered a natural experiment; however, as we have pointed out, this also requires additional information to interpret at regional scales. The Chukchi Sea is a sink for biota advected from the Bering Sea, and therefore, communities associated with warm conditions in the Bering Sea may serve as proxies to examine Arctic community level dynamics in the face of warming conditions (Walsh et al. 2004; Woodgate et al. 2012; Coyle et al. 2013). Consequently, examining the response of capelin to a range of physiological conditions may offer a better understanding of how populations of these important forage fish and the community in the Chukchi will respond to climate change. This study has revealed large-scale variation in condition of capelin, but the mechanisms behind this variation need to be further explored.

Varving terrestrial inputs are one mechanism that likely affected capelin condition. Habitats with greater terrestrial inputs are likely to have higher trophic diversity and more basal resources based on ranges of isotopic ratios. This is an important premise as freshwater inputs are expected to increase with climate change, and thus, we may expect to see more complex and more productive Arctic coastal food webs to develop in response. Two important questions arise from our results: (1) Will climate change lead to higher energy density of sub-Arctic forage fish in the Arctic and offer alternate high-quality forage to nearshore piscivores? (2) Are capelin utilizing suboptimal nearshore habitats near Barrow to avoid predation, or are they advected there through their ontogeny? We suggest that additional studies be developed to examine energetics and SIA of capelin in Arctic and sub-Arctic habitats to better elucidate the mechanisms that underlie these patterns.

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