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## Raman spectroscopic study of synthetic pyrope–grossular garnets: structural implications

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Abstract A study of the effect of substitution of Mg and Ca in garnet solid solution (Grt<sub>ss</sub>) was carried out using Raman spectroscopy to probe changes to the crystal lattice. The garnet solid solutions with composition changing along pyrope (Py; Mg<sub>3</sub>Al<sub>2</sub>Si<sub>3</sub>O<sub>12</sub>) and grossular (Gr; Ca<sub>3</sub>Al<sub>2</sub>Si<sub>3</sub>O<sub>12</sub>) binary were synthesized from glass at 6 GPa and 1400 °C and a second series of Grt<sub>ss</sub> with composition Py40Gr60 were synthesized at 6 GPa but different temperatures from 1000 to 1400 °C. Raman mode assignments were made based on a comparison with the two end members pyrope and grossular, which show consistent result with literature study on single crystals data. The correlation between the Raman mode frequencies and compositional changes along the pyrope-grossular binary suggests a two-mode behavior for Mg and Ca cations in the garnet structure. The full widths at half-maximum of selected Raman modes increase on moving away from the end members and are about double the end-member values in the mid-position, where the frequencies closely linearly change with composition. The frequencies of the translational modes of the  $SiO_4$  tetrahedron (T(SiO<sub>4</sub>)) show large deviations from linearity indicating a strong kinematic coupling with the translational modes of the Ca and Mg

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cations. The anomalies in  $T(SiO_4)$  are linked to mixing unit cell volume, suggesting that the nonlinear mixing volume behavior along the pyrope–grossular binary is related to the resistance of the Si–O bond to expansion and compression, which is caused by substitution of Mg and Ca cations in the dodecahedral sites. Annealing temperature also shows effect on Raman mode frequencies, but the main factor controlling the changes in mode frequencies along pyrope– grossular binary is composition.

Keywords Pyrope–grossular garnet solid solution  $\cdot$  Raman spectroscopy  $\cdot$  FWHM  $\cdot$  Non-ideal mixing  $\cdot$  Short-range ordering

#### Introduction

Garnets are common minerals found in many metamorphic rocks. They also make up an important constituent of the mantle transition zone (e.g., Irifune and Ringwood 1987, 1993). The general garnet formula  $X_3Y_2(SiO_4)_3$ can accommodate a range of divalent cations in the X site (mainly Ca, Mg, Fe and Mn) and trivalent cations in the Y site (mainly Al, Fe, Cr) resulting in large number of possible minerals. In order to simplify our characterization of the relationships between the chemical composition and the physical properties of the garnet minerals we often think of them as pure end members, for example grossular garnet (Gr:  $Ca_3Al_2Si_3O_{12}$ ) or pyrope (Py:  $(Mg_3Al_2Si_3O_{12})$ , with intermediate compositions described in terms of solid solutions of the two (for example: (Ca<sub>0.5</sub>Mg<sub>0.5</sub>)<sub>3</sub>Al<sub>2</sub>Si<sub>3</sub>O<sub>12</sub> is Gr<sub>50</sub>Py<sub>50</sub>). A first-order approach to relating end-member properties to those of intermediate compositions involves using a linear interpolation such as Vegard's law (Vegard 1921). Many structural and chemical properties of complex

garnet solid solutions can be obtained through careful characterization of the corresponding chemically simpler end members. However, deviations from ideal mixing are wellknown for many mixing properties particularly for those binary systems in which there is a significant difference in the size of the cation being substituted, for example the pyrope-grossular solid solution series (e.g., Newton et al. 1977; Geiger et al. 1987; Ganguly et al. 1993; Geiger and Feenstra 1997; Dachs and Geiger 2006; Du et al. 2015, 2016). Understanding these deviations is important since they can lead to local structural heterogeneities caused by substitutions which, for example, control element partitioning behaviors (e.g., van Westrenen et al. 1999). This has resulted in many recent studies that focused on the local structural environment changes with Mg/Ca substitution in pyrope-grossular system, including theoretical calculations (e.g., Bosenick et al. 2000, 2001; Freeman et al. 2006; Kelsey et al. 2008), extended X-ray absorption fine structure (EXAFS) analysis (Oberti et al. 2006; Quartieri et al. 2008), and a nuclear magnetic resonance (NMR) study (Kelsey et al. 2008), showed qualitatively consistent results that the X-O bond lengths vary with composition. However, one would expect that compositional changes should also affect the rotation angle of the SiO<sub>4</sub> tetrahedra; therefore, mechanistic details on the effect of substitution on the crystal structure are still unclear. For example, X-ray absorption spectroscopy studies of pyrope-grossular garnets suggest that the geometries of the local coordination of Ca and Mg cations do not change continuously along the pyrope-grossular binary (Oberti et al. 2006; Quartieri et al. 2008) which might explain the two-mode behavior of the translational Raman mode of X-cation (Ca, Mg) during substitution (Hofmeister and Chopelas 1991a; Kolesov and Geiger 1998) and the bimodal distribution of microstrain peaks along pyrope-grossular binary reported by Du et al. (2016).

Complete and accurate knowledge of how the crystal structure changes with composition along the pyropegrossular binary is essential for correctly computing their thermodynamic and physical properties. Because of the large available calorimetric, volumetric, spectroscopic, and phase equilibrium constraints from experimental studies of the pyrope-grossular garnet solid solutions, many computational studies have attempted atomistic simulations to elucidate mixing properties in order to predict their thermodynamic behavior over a wide range of temperatures and pressures (e.g., Dove 2001; Bosenick et al. 2000, 2001; Vinograd et al. 2004; Baima et al. 2015). However, none of these studies was able to demonstrate quantitative agreement with the available experimental constraints. The most recent Monte Carlo simulation on excess free energy of pyrope-grossular garnets predicated a significant short-range ordering of Mg and Ca cations for garnet solid

solution with composition (Mg<sub>0.5</sub>Ca<sub>0.5</sub>)<sub>3</sub>Al<sub>2</sub>Si<sub>3</sub>O<sub>12</sub> (Vinograd and Sluiter 2006). Although there are no experimental results showing phase transition from cubic Ia-3d to I4122 tetragonal symmetry at 600K as predicted by Vinograd and Sluiter (2006), the description that the two dodecahedral sites become structurally distinct through substitution was consistent with other theoretical and experimental studies (e.g., Bosenick et al. 2000, 2001; Quartieri et al. 2008). In addition, Monte Carlo simulations suggest that the configurational entropy of the pyrope-grossular solid solutions with intermediate compositions might be significantly decreased from the ideal mixing values due to the development of short-range ordering. Therefore, the effect of cation arrangements during substitution in the pyrope-grossular solid solutions should be taken into account for future thermodynamic calculations (Vinograd and Sluiter 2006).

Vibrational spectroscopy provides an excellent probe of the microscopic as well as the macroscopic behavior of minerals. Local effects, including effects of compression, heating, or compositional alteration on specific sites or atomic motions inside the structure, can be examined. Both Raman and infrared spectra of five end-member garnets have been reported (Hofmeister and Chopelas 1991b; Kolesov and Geiger 1998; Mernagh and Liu 1990), but accurate prediction of the mode frequencies of solid solutions by extrapolation from the end members is inadequate due to the nonlinear mixing behavior among garnet systems. The only one reported Raman spectra along the pyrope-grossular binary by Kolesov and Geiger (1998) showed deviation from ideal mixing between the two end members. However, the reported Raman spectra of the two end members also differed largely from the single-crystal Raman spectra (Gillet et al. 1992), leaving some uncertainties with regard to their reported Raman spectra for the pyrope-grossular solid solutions. More work is needed to clear up these uncertainties and to test the nonlinear compositional dependence of the Raman modes along the pyrope-grossular garnet binary.

There has been a long-standing discussion of the behavior of the Mg cation located in the dodecahedral site of pyrope. It could be either dynamically disordered (e.g., Armbruster et al. 1992; Kolsesov and Geiger 1998, 2000) or statically disordered over subsites (e.g., Hofmeister and Chopelas 1991b; Nakatsuka et al. 2011). The large temperature dependence of the line width of the band at ~135 cm<sup>-1</sup> has been interpreted as evidence for thermally disordered of Mg in pyrope (Kolsesov and Geiger 2000). Peak shift and line width variations of phonon spectra are related to local structural changes associated with mixing (Boffa Ballaran and Woodland 2006). It has been shown that the substitution of Mg/Ca inside pyrope–grossular solid solutions will cause different degrees of cation ordering (e.g., Bosenick et al. 2000, 2001; Du et al. 2016). Whether this static ordering of Mg/Ca cations results in changes in the Raman mode frequencies of the pyrope–grossular solid solutions is still unknown.

To address these issues we have collected a new set of Raman data from the synthetic end members pyrope and grossular and a series of solid solutions synthesized under the same pressure and temperature condition. We have also collected spectra of a series of garnets of composition  $Py_{40}Gr_{60}$  synthesized in the same manner, but annealed at different temperatures to test the effect of annealing temperature on microstructure and any possible correlations with the Raman mode frequencies and excess volumes. Systematic line broadening of Raman spectra of solid solutions is quantitative analysis to characterize the local structural changes due to substitution. Specifically, nonlinearity and unexpected trends in the vibrational modes will be closely examined to provide structural information, which brings insights on their volumetric properties.

#### **Experimental methods**

#### Samples synthesis

Pyrope–grossular solid solutions, with eight different compositions  $Py_{100-x}Gr_x$  (x = 0, 10, 20, 40, 60, 80, 90, 100) were synthesized from anhydrous glasses of stoichiometric composition. As previously reported in Du et al. (2015) the glasses were made by quenching molten mixtures of finely ground CaCO<sub>3</sub>, MgO, Al<sub>2</sub>O<sub>3</sub>, and SiO<sub>2</sub> powders of appropriate stoichiometry. Microprobe analysis confirmed that the compositions of the glasses were the same as the starting proportions of the oxides. As previously reported in Du et al. (2015, 2016, 2017), the eight crystalline samples were synthesized at the Lamont-Doherty Earth Observatory (LDEO) in a multi-anvil (MA) device by pressurizing

to 6 GPa and holding at 1400 °C for 0.5 h. Four additional samples with composition Py40Gr60 were synthesized at 6GPa but by holding at 1000, 1100, 1200, and 1400 °C for 48 h in order to explore the effect of annealing temperature on microstructure. We chose the  $Py_{40}Gr_{60}$  composition since this has been shown to exhibit the largest deviation from ideal mixing volume when synthesized at 6 GPa and 1400 °C (Du et al. 2016). The guenched samples were checked by electron microprobe and scanning electron microscopy to confirm that they were of single phase and of the same composition as the starting glasses (Du et al. 2015). The grain size of these garnet samples were found to be about 5-10 µm. The unit cell volume of these garnet crystals were calculated from X-ray diffraction data collected at high-resolution synchrotron light sources (Du et al. 2016, 2017). The synthesis or annealing conditions and unit cell volume of all Py40Gr60 garnets are summarized in Table 1.

#### Raman spectroscopy

Raman spectra were obtained with a Renishaw InVia Reflex laser Raman microprobe at the School of Earth and Space Sciences, Peking University. The system is built around a Leica DM LM microscope combined with an f/1.8 holographic imaging spectrograph and a diode pumped solid-state laser which emits a 50 mW beam at 532 nm (18,797 cm<sup>-1</sup>). Samples were placed beneath the microscope objective and excited by the laser directed through the microscope optics. The scattered light was collected along the same optical path as the incoming laser in 180° backscattering geometry. The laser beam spot was ~1 µm in diameter which is much smaller than the size of the grains in our sample. Ten spectra were collected from each garnet sample from different points on the sample grain giving Raman data from 100 to 1100 cm<sup>-1</sup>. The peak-fitting

Table 1 Synthesis and annealing conditions of garnet solid solution  $Py_{40}Gr_{60}$  and their Raman mode frequencies for translational motion of  $SiO_4$  tetrahedron

Experiment	Temperature (°C)	Heating time (h)	Unit cell parameter <sup>a</sup> (Å/cell)	Raman mode frequency of $T(SiO_4)$ (cm <sup>-1</sup> )	Peak width of $T(SiO_4)$ (cm <sup>-1</sup> )	
TT721	1400	0.5	11.722(1)	187.7(4)	7.27(65)	
GG1000 <sup>b</sup>	>1700	24	11.705(2)	187.7(3)	8.97(50)	
GG1002	1000	48	11.713(1)	188.5(7)	8.00(60)	
GG1004	1100	48	11.732(1)	189.0(5)	6.97(36)	
GG1005	1200	48	11.721(1)	188.2(7)	6.99(52)	
GG1006	1400	48	11.715(1)	187.9(5)	7.92(43)	

<sup>a</sup> Unit cell parameters were measured using synchrotron X-ray diffraction at BNL (Du et al. in press), numbers in the parentheses for unit cell parameters are errors from calculation, and numbers in the parentheses for Raman mode frequency and peak width are standard deviation from 10 measurements on the same samples

<sup>b</sup> Synthesized 1400 °C for 30 min and heated at 1200 °C for 1 day, but ended with blow out; temperature may have been higher than 1700 °C briefly because the Pt capsule melted

routines in Origin pro 2016 were used to obtain peak parameters, Raman mode frequencies, and peak widths [full width at half-maximum (FWHM)] from these spectra. Lorentz peak shape functions were used in a single peak-fitting procedure in line with previous studies (Hofmeister and Chopelas 1991a). Occasionally extra peaks were included in the fitting procedure to account for small shoulders in order to ensure a satisfactory fit to the main peak.

#### Results

#### Raman mode frequencies change with composition

Figure 1 shows the eight Raman spectra collected from our pyrope–grossular solid solution series. Values of the mode frequencies for peaks in these spectra determined using our peak-fitting procedure are contained in Table 2. The Raman spectra of all eight garnets show the same overall appearance containing sharp peaks grouped in three regions: low-frequency peaks (180–420 cm<sup>-1</sup>), medium-frequency peaks (450-700 cm<sup>-1</sup>), and high-frequency peaks (850–1100  $\text{cm}^{-1}$ ). The peak positions of the end members pyrope and grossular agree well with literature data from both synthetic polycrystalline samples (Gillet et al. 1992) and single crystals with same compositions (Kolesov and Geiger 1998) (Table 3). The Raman modes of the end members are well assigned to the vibrational modes allowed by crystal symmetry after Hofmeister and Chopelas (1991a) (Table 2). The observed vibrational modes are grouped into three main regions: the highest frequency modes (A, B, C, D) are dominated by stretching motions of Si–O ((Si–O)<sub>str</sub>), the modes at around 550 cm<sup>-1</sup> (E, F, G, H, I, J) are dominated by bending motions of Si-O ((Si-O)<sub>bend</sub>), and the modes at around 350 cm<sup>-1</sup> (K, L, M, N, O) are dominated by rotation motions of  $SiO_4$  (R(SiO<sub>4</sub>)). In addition, the mode at around 200  $\text{cm}^{-1}$  (S) is due to translational motions of  $SiO_4$  (T(SiO<sub>4</sub>)) (Kolesov and Geiger 1998). Garnet solid solutions with compositions along this binary were similarly assigned as the end members. Several bands between 240 and 350 cm<sup>-1</sup> (P, Q, R) are distinct in the grossular-rich samples, but weak in the pyrope-rich



Fig. 1 Raman spectra of garnets along pyrope–grossular binary. The marked P, Q, R peaks represent the newly appearing Raman modes (translational modes of X-cations) for grossular-rich garnets. E and F peaks mark the Raman modes associated with bending motion of

Si–O bond, showing discontinued changes with composition. And these discontinuities represent the two-mode behavior of Raman spectra of pyrope–grossular garnets

 Table 2
 Raman mode

 frequencies (cm<sup>-1</sup>) of garnet
 solid solutions along pyrope–

 grossular binary recorded at
 ambient condition

Modes	Pyrope	$Py_{90}Gr_{10}$	$\mathrm{Py}_{80}\mathrm{Gr}_{20}$	Py <sub>60</sub> Gr <sub>40</sub>	Py <sub>40</sub> Gr <sub>60</sub>	Py20Gr80	Py <sub>10</sub> Gr <sub>90</sub>	Grossular
T <sub>2g</sub>	1066	1060	1056	1045	1033	1018	1010	1002
A <sub>1g</sub>	927	922	918	907	898	887	884	880
T <sub>2g</sub>	903	884	887		877	857	852	849
T <sub>2g</sub>	869	861	856	844	837	828	826	825
T <sub>2g</sub>	686	675	683					
T <sub>2g</sub>	650	648	646	644	638	633	631	630
T <sub>2g</sub>				606	605	601	593	593
A <sub>1g</sub>	563	563	562	558	555	549	551	550
Eg	522	527		536	528	527	528	528
T <sub>2g</sub>	512	512	514	512	511	509	510	510
T <sub>2g</sub>	493	491	490	486	485	480	479	479
T <sub>2g</sub>	382	406			410	412	417	418
A <sub>1g</sub>	365	366	368	370	373	372	375	375
T <sub>2g</sub>		346			354		334	332
T <sub>2g</sub>	322	323					319	317
T <sub>2g</sub>					281	279	279	279
T <sub>2g</sub>					248	246	247	247 240
Eg	211	207	202	191	188	183	183	183
T <sub>2g</sub>	133	132	128	124				

# Table 3Raman modefrequencies (cm<sup>-1</sup>) of end-member pyrope and grossularrecorded at ambient condition

Ругоре				Grossular				
This study	а	b	с	This study	а	b	с	
1066	1064	1066	1062	1002	1007	1009	1007	
927	927	928	925	880	881	883	880	
903	903	902	899	849	851	852	848	
869	870	871	866	825	826	826	827	
686				630	629	631	630	
650	651	650	648	593	590			
563	564	563	562	550	549	551	550	
522	524		525	528	526	532	529	
512	513	512	510	510	509	510	512	
493	492	492	490	479	478	478	483	
382	380	383	379	418	416	416	420	
365	366	364	365	375	374	375	373	
				332	330	330	333	
322	324	322	318	317	317		319	
				279	278	278	280	
				248	246	247	247	
211	211	211	208	183	178	178	181	
133			135					

*a* Polarized single-crystal spectra on natural samples (Hofmeister and Chopelas 1991a), *b* unpolarized spectra on synthetic polycrystalline samples (Gillet et al. 1992), *c* polarized single-crystal spectra on synthetic samples (Kolesov and Geiger 1998)

samples. They decrease in intensity towards the intermediate solid solution compositions until they are no longer observable in the  $Py_{60}Gr_{40}$  spectrum (Fig. 1). They have previously been identified as modes associated with X-site cations (Ca in this case) in the dodecahedral site with the decrease in intensity being due to substitution of Mg for Ca and the two-mode behavior caused by the large size difference between the mixing cations Mg and Ca (Hofmeister and Chopelas 1991a). In the pyrope-rich samples, we see peaks due to Mg-translational modes (~133 cm<sup>-1</sup>). The frequencies associated with the Mg and Ca-translational modes do not change linearly with composition between the two end-member garnets as was previously observed (Kolesov and Geiger 1998).

Figure 2 shows the compositional dependencies of the internal mode frequencies of  $SiO_4$ ,  $(Si-O)_{str}$  and  $(Si-O)_{bend}$  along the pyrope–grossular binary. The stretching motions of Si–O bonds show a strong dependence on garnet composition (Fig. 2a), and the bending motions of Si–O bonds

show a relatively weak dependence on garnet composition (Fig. 2b). The internal mode frequencies of  $SiO_4$  in pyroperich garnets are higher than the grossular-rich garnets. This trend has been found for garnets in general (Hofmeister and Chopelas 1991a) that the internal modes of  $SiO_4$  tetrahedra show simple dependencies on lattice parameters. Two  $(Si-O)_{bend}$  bands (E and F) appear in two different ranges for pyrope-rich garnets (646–686 cm<sup>-1</sup>) and grossular-rich garnets (593–646 cm<sup>-1</sup>) (Table 2; Fig. 2b) indicating a discontinuous change with composition.

Selected internal vibrational mode frequencies and widths as a function of composition along pyrope– grossular binary are shown with details in Fig. 3. The



Fig. 2 Compositional dependence of frequencies of Raman mode corresponding with internal motions of SiO<sub>4</sub>-tetrahedra: **a** Si–O stretching and **b** Si–O bending. The Raman modes marked as E and F in Fig. 1 (600–700 cm<sup>-1</sup>) show discontinuity along pyrope–grossular binary





Fig. 3 Selected Raman mode frequencies and their peak widths change with composition along pyrope–grossular binary.  $\mathbf{a}$  (Si–O) stretching,  $\mathbf{b}$  (Si–O) bending,  $\mathbf{c}$  SiO<sub>4</sub> rotation, and  $\mathbf{d}$  SiO<sub>4</sub> translation.

The *empty circles* are data from Kolesov and Geiger (1998), and the *gray circles* are  $Py_{40}Gr_{60}$  garnets synthesized at different temperatures

frequencies for internal modes are seen to decrease linearly with Ca content in garnet composition (Fig. 3a1 and b1), as previously observed (Kolesov and Geiger 1998). The stretching motions show stronger dependence on composition and hence unit cell volume than the bending modes as expected (Hofmeister and Chopelas 1991a) with (Si-O)<sub>str</sub> varying at an average rate of ~130  $\pm$  20 cm<sup>-1</sup>/Å compared to an average garnet value of about 120 cm<sup>-1</sup>/Å and the (Si–O)<sub>bend</sub> varying at an average rate of 63 cm<sup>-1</sup>/Å compared to an average garnet value of about ~90 cm<sup>-1</sup>/Å. The higher rate of change for the (Si–O)<sub>str</sub> modes is due to the fact that the force constants for stretching modes are much smaller than those bending modes thus amplifying the effect.

The Raman mode frequency (*L*) associated with the rotation of the  $SiO_4$ -tetrahedron ( $R(SiO_4)$ ) is roughly independent of composition (Fig. 3c1), and the translational modes of the  $SiO_4$ -tetrahedron ( $T(SiO_4)$ ) (S) changes nonlinearly with composition (Fig. 3d1), indicating the coupling of this vibration with translational mode of the X-site cations.

#### Raman mode frequencies vs. synthesis temperatures

Figure 4 shows the Raman spectra from the  $Py_{40}Gr_{60}$  garnet solid solutions synthesized at 6 GPa and a range of temperatures from 1000 to 1400 °C. The Raman spectra of these five  $Py_{40}Gr_{60}$  garnets have the same overall appearance. The selected Raman mode frequencies determined from these spectra using out peak-fitting method are shown in Fig. 5. The Raman mode associated with translational motions of the SiO<sub>4</sub>-tetrahedra and its peak width (FWHM) for all of these five garnet samples are summarized in Table 1 and Fig. 6. We can see that there is a weak-positive correlation between translational mode frequency and unit cell parameter (Fig. 6a) and that the Raman mode lines become broader with increasing annealing temperature (Fig. 6b).

#### Discussion

Similar compositional dependence of selected Raman modes for the pyrope-grossular garnets was reported by Kolesov and Geiger (1998), although their Raman mode frequencies associated with bending of Si-O bond and rotation of SiO<sub>4</sub>-tetrahedron are both smaller than that we report in this study (Fig. 3b1, c1). Kolesov and Geiger (1998) reported Raman mode frequencies of end-member garnets from both single-crystal and polycrystalline samples and that the single-crystal values are larger than those from the polycrystalline samples and agree well with our data and other literature studies (Table 3). The Raman laser spot size (~1 µm) that we used was smaller than the grain size (5-10 µm) of our samples. This could be the reason why the Raman data presented here is closer to the data collected from single crystals and Raman frequency values from pyrope-grossular garnet solid solutions are larger than those from polycrystalline samples as reported by Kolesov and Geiger (1998).

As shown by Bosenick et al. (1995), the grossular-rich garnets that were used by Kolesov and Geiger (1998) were synthesized from seeded glass, where the garnet seed was hydrothermally synthesized from garnet glass with 7 wt% distilled water. There was no report of the water content of these garnets in their paper and whether



**Fig. 4** Raman spectra of  $Py_{40}Gr_{60}$  garnets synthesized at different temperatures from 1000 to 1400 °C, showing same overall appearance



Fig. 5 Selected Raman mode frequencies of Py40Gr60 garnets annealed at different temperatures

there may be an effect on the vibrational spectra from any OH included in their material. The garnet solid solutions used in the previous study were synthesized at temperatures from 1050 to 1400 °C (Bosenick et al. 1995). As mentioned above, the synthesis temperature has a slight effect on the Raman mode frequencies. On the other hand, the garnet samples used in this study were all synthesized anhydrously from dry glass at the same temperature (1400 °C). Therefore, the Raman mode frequencies from our samples should be more consistent and provide a better data set for calculation of thermodynamic parameters via lattice dynamic models.

The low-frequency mode at around  $210 \text{ cm}^{-1}$  in pyrope that shifts gradually to  $180 \text{ cm}^{-1}$  in grossular is related to the translational motions of the SiO<sub>4</sub>-tetrahedra rather than the translational motions of the X-cation (Mg, Ca) (Kolesov and Geiger 1998). Our new observation shows that the translational motions of the SiO<sub>4</sub> tetrahedra deviate strongly from linearity along the solid solution series (Fig. 3d1). Since these Raman mode frequencies are close to those associated with the translational motions of the Mg and Ca cations (T (Mg, Ca)), the large deviation suggests a strong kinematic coupling between the T(SiO<sub>4</sub>) and T (Mg, Ca) modes. This implies that the frequency of the  $T(SiO_4)$  mode depends on the bond strengths of surrounding polyhedral rather than on the internal  $SiO_4$  bonds. Therefore, the decrease in frequency of the T(SiO<sub>4</sub>) and T (Mg, Ca) modes from pyrope to grossular is caused by expansion of the lattice due to substitution of the smaller  $Mg^{2+}$  cation with the much larger  $Ca^{2+}$  cation. The negative deviation from linearly changing with composition of the Raman modes of  $T(SiO_4)$  indicates an increased volume for intermediate compositions. We have observed similar mixing volume deviations previously from unit cell volume measurements the behavior of which as a function of composition we compare with the Raman mode frequency data in Fig. 7. The large excess volume change with composition is caused by large size mismatch between Ca (1.12 Å) and Mg (0.89 Å) (Shannon 1976), which then causes changes to the rotational orientation of the SiO<sub>4</sub> tetrahedra and changes in the Si–O bond lengths.

Figure 3 shows plots of the peak width (FWHM) of Raman modes as a function of composition (Table 4). The Raman modes of the synthetic end members, pyrope and grossular, are very sharp indicating that the cations are well ordered. The Raman modes of garnet solid solutions show peak broadening, reaching a maximum near the  $Py_{50}Gr_{50}$  composition (Fig. 3a2, b2, and c2). For solid solution garnets line broadening in the vibrational spectra can be caused by defects or disorder (Boffa Ballaran et al. 1999). Similar Raman mode peak broadening has been observed for the pyrope-majorite solid solutions, which was interpreted as cation disorder due to the  $Mg^{2+} + Si^{4+} = 2Al^{3+}$  substitution. This mode broadening was interpreted as a characteristic precursor effect for structural phase transitions, which was confirmed in the pyrope-majorite system by X-ray diffraction as well as vibrational spectroscopy (Liu et al. 2017). We did not observe any structural phase transition for the pyrope-grossular system during our previous X-ray diffraction studies (Du et al. 2016, 2017). We did, however, observe similar X-ray diffraction peak broadening which was found to be related to the arrangement of Ca and Mg





cations in the dodecahedral sites (Du et al. 2016, 2017). When one atom (Mg or Ca) is replaced by the other on the dodecahedral site, the local structure relaxes around it resulting in structural variations on the same length scale as phonons. The line broadening that we present here for Si-O stretching, bending, and rotation modes of SiO<sub>4</sub> may be due to the effect of these substitutions on the Raman mode frequencies. However, the line width of the Raman mode associated with the translational motion of  $SiO_4$  (T(SiO\_4)) shows a large difference compared to the other three modes (Fig. 3d2). Garnet solid solutions with compositions close to 50:50 ( $Py_{40}Gr_{60}$  and  $Py_{60}Gr_{40}$ ) show narrower peaks than those with compositions close to the two end members. As discussed above, the Raman modes of  $T(SiO_4)$  show strong coupling with cation (Mg and Ca) substitutions, therefore we might expect a much broader peak width for  $T(SiO_4)$ . However, we observed the opposite. The peak widths of garnets with intermediate compositions are narrower than those with compositions close to the end members, indicating that the physical origins of the peak broadening for  $T(SiO_4)$  is different comparing with  $R(SiO_4)$ . If we took line broadening as indicative of the range of local distortions, the local structural heterogeneities in a crystal will cause strain gradients. The microstrain calculated from X-ray diffraction data varies in a similar double-peak pattern as the peak width of the Raman mode  $T(SiO_4)$  (Fig. 8), suggesting that they both probe the same distortions of the crystal structure. As mentioned above, the frequency of the  $T(SiO_4)$  mode is largely affected by the surrounding X-cations, because it is dependent on whether it is strongly compressed by the larger cation (Ca) in the neighboring dodecahedron or weekly compressed by the smaller cation (Mg). Along this line, we would expect that the arrangement of Mg and Ca cation in the dodecahedral site will affect the Raman mode frequency of



**Fig. 7** Deviation of translational mode frequencies of  $SiO_4$  tetrahedra from linearity along pyrope–grossular binary (*empty symbols*). The *solid symbols* are excess volume along this binary reported by Du et al. (2016). The corresponding changes on rotation of  $SiO_4$  tet-

rahedra caused by substitution of Mg and Ca cations show good correlation with unit cell volume changes along pyrope–grossular binary. The *dashed line* is polynomial fitting to the excess volume data

 $T(SiO_4)$ . If we took the peak broadening of the  $T(SiO_4)$ mode as a sign of disordering then the much narrower peaks for  $Py_{60}Gr_{40}$  and  $Py_{40}Gr_{60}$  indicate that these two garnets have a more ordered structure, which is consistent with previous studies using NMR spectroscopy and synchrotron X-ray diffraction (e.g., Bosenick et al. 1995, 1999; Du et al. 2016). As proposed by Bosenick et al. (1995), garnet solid solutions with intermediate composition along the pyrope-grossular binary show larger degree of short-range ordering than those with compositions close to the end members. It was also highlighted through an ab initio calculation that the energies of particular atomic configurations are closely related to the third nearest neighbor (3NN) arrangement of the Mg and Ca cations, which are edge sharing with the same SiO<sub>4</sub> tetrahedron resulting in the pairing of the same cations (Mg–Mg and Ca–Ca) in the S3NN site being unfavorable from an energy point of view due to a higher configurational energy (Freeman et al. 2006). If we assume that the broadening of the X-ray diffraction peaks and the Raman peaks associated with the  $T(SiO_4)$  mode both indicate more microstrain, then this would indicate that garnets

with composition close to the end members have more cation pairings (Mg–Mg and Ca–Ca) in the third nearest neighbor position than garnets with intermediate compositions (e.g.,  $Py_{60}Gr_{40}$  and  $Py_{40}Gr_{60}$ ) which thus have less possibility of randomly generated Mg–Mg or Ca–Ca (S3NN) pairings giving an intrinsic degree of short-range ordering (e.g., Bosenick et al. 1995, 2000, 2001; Dove et al. 2000; Vinograd et al. 2004; Lavrentiev et al. 2006; Vinograd and Sluiter 2006). This less "strained" structure may have relatively large excess volume because Ca–Mg pairs pack less efficiently than strain-causing Mg–Mg and Ca–Ca pairs.

The short-range ordering of Mg and Ca in the pyrope– grossular garnets was reported to decrease with increasing synthesis temperature (Bosenick et al. 1999). Garnets synthesized at higher temperatures might be expected to exhibit larger microstrain as more strain-causing S3NN pairs are formed resulting in broader Raman peaks as high synthesis temperature corresponds to a more disordered cation distribution. The Raman spectra of  $Py_{40}Gr_{60}$  synthesized at different temperatures show that the peak width of the T(SiO<sub>4</sub>) Raman mode increases with increasing

Table 4	Selected Raman peak
widths (I	FWHM) of garnets
along py	rope-grossular binary

	Pyrope	$Py_{90}Gr_{10}$	$Py_{80}Gr_{20}$	$Py_{60}Gr_{40}$	$Py_{40}Gr_{60}$	$Py_{20}Gr_{80}$	Py10Gr90	Grossular
(Si–O)str.	12.7(1)	17.3(1)	17.4(1)	19.4(1)	19.9(2)	21.8(2)	16.2(1)	11.5(1)
(Si-O)bend	11.3(1)	13.2(3)	13.4(4)	14.1(9)	12.9(3)	12.4(2)	7.7(1)	5.7(2)
R(SiO <sub>4</sub> )	19.7(6)	20.4 (10)	21.4(3)	27.6(8)	22.3(5)	23.1(4)	12.7(4)	8.1(6)
T(SiO <sub>4</sub> )	7.9(4)	10.1(3)	12.9(8)	6.7(7)	6.6(3)	8.4(4)	4.1(1)	2.2(1)

Numbers in the parentheses are errors from peak-fitting procedure



**Fig. 8** Peak width of Raman mode corresponding to the translational motion of SiO<sub>4</sub> tetrahedra. The *open symbols* are microstrain data calculated from synchrotron X-ray diffraction pattern (Du et al. 2017) with equation  $B_{\text{garnet}} \cos \theta = \frac{K\lambda}{(L) > \cos \theta} \cos \theta + 4\eta \tan \theta \cos \theta = \frac{K\lambda}{(L)} + 4\eta \sin \theta$ , where  $\lambda$  is the wavelength of the X-ray, and  $\theta$  is the Bragg angle, *K* is a constant of value approximately 0.9 (Langford and Wilson 1978), *L* 

synthesis temperature, which corresponds to a more disordered structure (Fig. 6). Therefore, unlike internal modes of SiO<sub>4</sub>, which change linearly with composition, the Raman mode of  $T(SiO_4)$  deviates largely from linearity and shows strong correlation with the microstrain inside the garnet structure and so may be a good indicator of the degree of cation ordering in solid solutions.

#### Conclusions

An analysis of Raman mode frequencies of garnets along the pyrope–grossular solid solution series shows that the vibrational spectra change with composition and annealing temperatures. The Raman modes associated with the stretching and bending motions of Si–O bonds and the rotational motions of the SiO<sub>4</sub> tetrahedra show simple linear dependences on composition, but the translational modes of SiO<sub>4</sub> tetrahedra show a nonlinear trend due to stronger coupling with translational modes of the X (Ca and Mg) cations.

The peak widths of the selected Raman modes  $(A_{1g})$  show a nonlinear dependence on composition, and reach a maximum near the intermediate garnets. The arrangement of Mg and Ca has a large effect on the translational motion of the SiO<sub>4</sub> tetrahedra. The peak widths of translational modes of T(SiO<sub>4</sub>) are slightly "M" shaped vs. composition and may result from different short-range ordering status of Mg and Ca cations during substitution.

is the crystal size (diffraction domain size),  $\frac{K\lambda}{(L)\cos\theta}$  is the size effect on XRD peak broadening (Scherrer 1918),  $4\eta \tan\theta$  is the width of the peak due to microstrain (Stokes and Wilson 1944), and  $\eta$  is the microstrain. Similar two-peaked pattern indicates the local distortional strain by Mg–Ca size misfit affects the translational mode of SiO<sub>4</sub> tetrahedra. The *dashed line* is polynomial fitting to the microstrain data

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### References

- Armbruster T, Geiger CA, Lager GA (1992) Single crystal X-ray refinement of almandine-pyrope garnets at 298 and 100 K. Am Min 77:512–523
- Baima J, Ferrabone M, Orlando R, Erba A, Dovesi R (2015) Thermodynamics and phonon dispersion of pyrope and grossular silicate garnets from ab initio simulations. Phys Chem Min 43:137–149
- Boffa Ballaran T, Woodland AB (2006) Local structure of ferric iron-bearing garnets deduced by IR-spectroscopy. Chem Geol 225:360–372
- Boffa Ballaran T, Carpenter MA, Geiger CA, Koziol AM (1999) Local structural heterogeneity in garnet solid solutions. Phys Chem Mineral 26:554–569
- Bosenick A, Geiger CA, Schaller T, Sebald A (1995) A <sup>29</sup>Si MAS NMR and IR spectroscopic investigation of synthetic pyropegrossular garnet solid solutions. Am Min 80:691–704
- Bosenick A, Geiger CA, Phillips BL (1999) Local Ca-Mg distribution of Mg-rich pyrope-grossular garnets synthesized at different temperatures revealed by <sup>29</sup>Si MAS NMR spectroscopy. Am Min 84:1422–1432
- Bosenick A, Dove MT, Geiger CA (2000) Simulation studies on the pyrope-grossular garnet solid solution. Phys Chem Min 27:398–418

- Bosenick A, Dove MT, Heine V, Geiger CA (2001) Scaling of thermodynamic mixing properties in garnet solid solutions. Phys Chem Min 28:177–187
- Dachs E, Geiger CA (2006) Heat capacities and entropies of mixing of pyrope-grossular (Mg<sub>3</sub>Al<sub>2</sub>Si<sub>3</sub>O<sub>12</sub>-Ca<sub>3</sub>Al<sub>2</sub>Si<sub>3</sub>O<sub>12</sub>) garnet solid solutions: a low-temperature calorimetric and a thermodynamic investigation. Am Min 91:894–906
- Dove MT (2001) Computer simulations of solid solutions. In: Geiger CA (ed) Solid solutions in silicate and oxide systems of geological importance. EMU notes in mineralogy, vol 3. Eötvös University Press, Budapest, pp 225–250
- Dove MT, Bosenick A, Myers ER, Warren MC, Redfer SAT (2000) Modeling in relation to cation ordering. Phase Trans 71:205–226
- Du W, Clark SM, Walker D (2015) Thermo-compression of pyropegrossular garnet solid solution: non-linear compositional dependence. Am Min 100:215–222
- Du W, Clark SM, Walker D (2016) Excess mixing volume, microstrain, and stability of pyrope-grossular garnets. Am Min 101:193–204
- Du W, Li X, Li B (2017) Microstrain in pyrope-grossular garnet solid solution at high pressure: a case study of  $Py_{90}Gr_{10}$  and  $Py_{10}Gr_{90}$  up to 15 GPa. Phys Chem Min 44(6):377–388
- Freeman CL, Allan NL, van Westrenen W (2006) Local cation environments in the pyrope-grossular Mg<sub>3</sub>Al<sub>2</sub>Si<sub>3</sub>O12–Ca<sub>3</sub>Al<sub>2</sub>Si<sub>3</sub>O<sub>12</sub> garnet solid solution. Phys Rev B 74:134203
- Ganguly J, Cheng W, O'Neill HC (1993) Syntheses, volume, and structural changes of garnets in the pyrope-grossular join; implications for stability and mixing properties. Am Min 78:583–593
- Geiger CA, Feenstra A (1997) Molar volumes of mixing of almandine-pyrope and almandine-spessartine garnets and the crystal chemistry and thermodynamic-mixing properties of the aluminosilicate garnets. Am Min 82:571–581
- Geiger CA, Newton RC, Kleppa OJ (1987) Enthalpy of mixing of synthetic almandine-grossular and almandine-pyrope garnets from high-temperature solution calorimetry. Geochim Cosmochim Acta 51:1755–1763
- Gillet P, Fiquet G, Malezieux JM, Geiger CA (1992) High-pressure and high-temperature Raman spectroscopy of end-member garnets: pyrope, grossular and andradite. Eur J Min 4:651–664
- Hofmeister AM, Chopelas A (1991a) Vibrational spectroscopy of endmember silicate garnets. Phys Chem Min 17:503–526
- Hofmeister AM, Chopelas A (1991b) Thermodynamic properties of pyrope and grossular from vibrational spectroscopy. Am Min 76:880–891
- Irifune T, Ringwood AE (1987) Phase transformation in primitive MORB and pyrolite compositions to 25 GPa and some geophysical implications. In: Manghnani MH, Syono Y (eds) High pressure research in mineral physics, vol 39. TERRAPUB Tokyo/ American Geophysical Union, Washington, DC, pp 235–246
- Irifune T, Ringwood AE (1993) Phase transformations in subducted oceanic crust and buoyancy relationships at depths of 600–800 km in the mantle. Earth Planet Sci Lett 117(1–2):101–110
- Kelsey KE, Stebbins JF, Du L, Mosenfelder JL, Asimow PD, Geiger CA (2008) Cation order/disorder behavior and crystal chemistry of pyrope-grossular garnets: an 17O 3QMAS and 27Al MAS NMR spectroscopic study. Am Min 93:134–143

- Kolesov BA, Geiger CA (1998) Raman spectra of silicate garnets. Phys Chem Min 25:142–151
- Kolesov BA, Geiger CA (2000) Low-temperature single-crystal Raman spectrum of pyrope. Phys Chem Min 27:645–649
- Langford JI, Wilson AJC (1978) Scherrer after sixty years: a survey and some new results in the determination of crystallite size. J Appl Crystallogr 11:102–113
- Lavrentiev MY, van Westrenen W, Allan NL, Freeman CL, Purton JA (2006) Simulation of thermodynamic mixing properties of garnet solid solutions at high temperatures and pressures. Chem Geol 225:336–346
- Liu Z, Du W, Shinme T, Gréaux S, Zhou C, Arimoto T, Kunimoto T, Irifune T (2017) Garnets in the majorite-pyrope system: symmetry, lattice microstrain, and order-disorder of cations. Phys Chem Min 44:237–245
- Mernagh TP, Liu L (1990) Pressure dependence of Raman spectra from the garnet end-members pyrope, grossularite and almandine. J Raman Spec 21:305–309
- Nakatsuka A, Shimokawa M, Nakayama N, Ohtaka O, Arima H, Okube M, Yoshiasa A (2011) Static disorders of atoms and experimental determination of Debye temperature in pyrope: low- and high-temperature single-crystal X-ray diffraction study. Am Min 96:1593–1605
- Newton RC, Charlu TV, Kleppa OJ (1977) Thermochemistry of high pressure garnets and clinopyroxenes in the system CaO-MgO-Al<sub>2</sub>O<sub>3</sub>-SiO<sub>2</sub>. Geochim Cosmochim Acta 41:369–377
- Oberti R, Quartieri S, Dalconi MC, Boscherini F, Iezzi G, Boiocchi M, Eeckhout SG (2006) Site preference and local geometry of Sc in garnets: part I. Multifarious mechanisms in the pyrope-grossular join. Am Min 91:1230–1239
- Quartieri S, Boscherini F, Dalconi C, Iezzi G, Meneghini C, Oberti R (2008) Magnesium K-edge EXAFS study of bond-length behavior in synthetic pyrope-grossular garnet solid solutions. Am Min 93:495–498
- Scherrer P (1918) Bestimmung der Grösse und der inneren Struktur von Kolloidteilchen mittels Röntgenstrahlen. Nachrichten der Kgl Gesellschaft der Wissenschaftern zu Göttingen 26:98–100
- Shannon RD (1976) Revised effective ionic radii. Acta Crystallogr A32:751–767
- Stokes AR, Wilson AC (1944) The diffraction of X rays by distorted crystal aggregates. Proc Phys Soc 56:174–181
- van Westrenen W, Blundy J, Wood B (1999) Crystal-chemical controls on trace element partitioning between garnet and anhydrous silicate melt. Am Min 84:838–847
- Vegard L (1921) Die Konstitution der Mischkristalle und die Raumfüllung der Atome. Zeitschrift für Physik 5:17–26
- Vinograd VL, Sluiter MHF (2006) Thermodynamics of mixing in pyrope-grossular, Mg<sub>3</sub>Al<sub>2</sub>Si<sub>3</sub>O<sub>12</sub>-Ca<sub>3</sub>Al<sub>2</sub>Si<sub>3</sub>O<sub>12</sub>, solid solution from lattice dynamics calculations and Monte Carlo simulations. Am Min 91:1815–1830
- Vinograd VL, Sluiter M, Winkler B, Putnis A, Hålenius U, Gale JD, Becker U (2004) Thermodynamics of mixing and ordering in pyrope-grossular solid solution. Min Mag 68:101–121