

Integral formulas for computing a third-order gravitational tensor from volumetric mass density, disturbing gravitational potential, gravity anomaly and gravity disturbance

Michal Šprlák · Pavel Novák

Received: 28 May 2014 / Accepted: 3 October 2014 / Published online: 22 October 2014
© Springer-Verlag Berlin Heidelberg 2014

Abstract A new mathematical model for evaluation of the third-order (disturbing) gravitational tensor is formulated in this article. Firstly, we construct corresponding differential operators for the components of the third-order (disturbing) gravitational tensor in a spherical local north-oriented frame. We show that the differential operators may efficiently be decomposed into an azimuthal and an isotropic part. The differential operators are even more simplified for a certain class of isotropic kernels. Secondly, the differential operators are applied to the well-known integrals of Newton, Abel-Poisson, Pizzetti and Hotine. In this way, 40 new integral formulas are derived. The new integral formulas allow for evaluation of the components of the third-order (disturbing) gravitational tensor from density distribution, disturbing gravitational potential, gravity anomalies and gravity disturbances. Thirdly, we investigate the behaviour of the corresponding integral kernels in the spatial domain. The new mathematical formulas extend the theoretical apparatus of geodesy, i.e. the well-known Meissl scheme, and reveal important properties of the third-order gravitational tensor. They may be exploited in geophysical studies, continuation of gravitational field quantities and analysing the gradiometric-geodynamic boundary value problem.

Keywords Gravitational tensor · Differential operator · Integral equation · Meissl scheme

1 Introduction

Currently available sensors in geodesy and geophysics allow for measuring scalar, vectorial and second-order tensorial quantities of the Earth's gravitational potential field, see, e.g. (Torge 1989; Schwarz and Li 1997; Chelton et al. 2001; Seiber 2003; Forsberg and Olesen 2010; Rummel 2010; Timmen 2010; Torge and Müller 2012). Such observables have widely been exploited by geodesists in various parametrization methods, such as harmonic analysis, integral formulas or radial basis function approach, for the purpose of the Earth's gravitational field modelling. Theoretical aspects of these quantities have also been extensively studied and well understood (Meissl 1971; Rummel and van Gelderen 1995; Rummel 1997; Grafarend 2001; Bölling and Grafarend 2005; Freeden and Schreiner 2009).

On the other hand, new sensors proposed for observing a third-order gravitational tensor are currently under development. This fact may be illustrated by the project Dulkyn (<http://www.dulkyn.ru/>) which aims at developing a system that would eventually observe directional derivatives of the gravitational potential up to the third order together with their temporal variations on the Earth surface (Balakin et al. 1997). In addition, the gravity-dedicated satellite mission called OPTIMA (OPTical Interferometry for global MAass change detection from space), which should measure third-order derivatives of the Earth's gravitational potential, was proposed (Brieden et al. 2010).

Although not currently observable, third-order potential derivatives have already been discussed in various contexts in geodesy. Moritz (1967) investigated parameters of the disturbing gravitational field up to the third-order disturbing gravitational tensor. He showed that gravitational tensors of order three and higher were independent of orientation on platforms without inertial stabilization. Thus, any instrument

M. Šprlák (✉) · P. Novák
New Technologies for the Information Society, Faculty of Applied Sciences, University of West Bohemia, Univerzitní 22, 306 14 Plzeň, Czech Republic
e-mail: michal.sprlak@gmail.com; sprlak@gmail.com; sprlak@ntis.zcu.cz

capable of observing such quantities would provide a purely gravitational signal. [Cunningham \(1970\)](#) derived spherical harmonic series for gravitational tensors of an arbitrary order in a geocentric reference frame. This study, which is of particular interest to satellite geodesy and celestial mechanics, was further extended by [Metris et al. \(1999\)](#) and [Petrovskaya and Vershkov \(2010\)](#). Higher-order vertical derivatives of a harmonic function for its analytical continuation in space were also considered by [Velkoborský \(1982\)](#). [Rummel \(1986\)](#), [Koop \(1993\)](#) and [Albertella et al. \(2000\)](#) used a third-order gravitational tensor for error analysis of gradiometric observations. [Wang \(1989\)](#) provided expressions for the third-order derivatives of the normal gravitational potential generated by an equipotential ellipsoid for analysing sensitivity of second-order gradients measured by an airborne gradiometer. Similarly, [Ardalan and Grafarend \(2001\)](#) expanded the normal gravitational potential in the Taylor series up to the third-order potential derivatives in analysing Bruns's formula. [Grafarend \(1997\)](#) derived a functional relationship between curvature and torsion of a plumb-line to the second- and third-order derivatives of the gravitational potential.

[Nagy et al. \(2000\)](#) provided directional derivatives of the gravitational potential up to third-order for the mass prism. [Keller and Sharifi \(2005\)](#) analysed satellite gradiometry using a pair of satellites and investigated the influence of higher-order gravitational tensors. [Tóth \(2005\)](#) investigated the gradiometric-geodynamic boundary value problem and its linearisation, and provided expressions of the third-order gravitational tensor in a spherical local reference frame. Making use of the same expressions, [Tóth and Földvály \(2005\)](#) studied projection of the second-order gravitational tensor observed by the gradiometric mission GOCE (Gravity field and steady-state Ocean Circulation Explorer, [ESA 1999](#)) onto a geocentric sphere approximating its orbit. [Casotto and Fantino \(2009\)](#) provided general equations for gravitational tensors up to the fourth order by tensor analysis. They also derived explicit expressions for gravitational tensors in the local and global reference frames up to the third order. [Csapó et al. \(2009\)](#) exploited third-order derivatives of the gravitational potential for evaluation of the second-order vertical gradient which is not directly observable by a torsion balance apparatus. [Sali \(2009\)](#) modelled the third-order vertical derivative of the gravitational potential to reduce observation noise in aerial gradiometry data caused by aircraft motion.

Computational aspects of harmonic synthesis up to third-order derivatives of Legendre functions were also discussed ([Fantino and Casotto 2009](#); [Fukushima 2012, 2013](#)). Motivated by higher sensitivity of the third-order derivatives to short-wavelength structures of potential fields ([Jacoby and Smilde 2009](#)), their exploitation has also been suggested for geophysical exploration purposes, see, e.g. ([Troshkov and Shalaev 1968](#); [Smith et al. 1998](#); [Fedi and Florio 2001](#); [Thurston et al. 2002](#); [Abdelrahman et al. 2003](#); [Hafez et al.](#)

[2006](#); [Pajot et al. 2008](#); [Veryaskin and McRae 2008](#); [Beiki 2010](#); [Klokočnák et al. 2010](#); [Eppelbaum 2011](#)).

Higher-order potential tensors have been applied also in other sciences. [Jordan et al. \(1995\)](#) constructed potential energy surfaces from third-order potential derivatives for studies in molecular chemistry. [Sun et al. \(2011\)](#) provided a mathematical formulation, which splits Hamiltonian into both kinetic and potential energy. Such formulation may be defined by an operator containing third-order potential derivatives and allows for solving related problems of astronomy and celestial mechanics.

In this article, we provide a rigorous mathematical model for transformation of the third-order gravitational tensor to other parameters of the Earth's gravitational field. Namely, we derive differential operators for the components of the third-order (disturbing) gravitational tensor in the spherical local north-oriented frame (LNOF). We also provide integral formulas that relate the volumetric mass density distribution, disturbing gravitational potential, gravity anomaly and gravity disturbance to the components of the third-order (disturbing) gravitational tensor.

The presented study reveals important properties of the third-order gravitational tensor and further extends the Meissl scheme of physical geodesy, i.e. the well-known theoretical paradigm interrelating various parameters of the Earth's gravitational field. The new formulas may be exploited in geophysical studies, continuation of gravitational field quantities in space, e.g. by Taylor expansions or gravity transport ([Marussi 1985](#)), and in analysing the gradiometric-geodynamic boundary value problem ([Tóth 2005](#)). Further applications of the new formulas may emerge in the future after successful development of instruments sensing directly the third-order gravitational gradients.

The article is organised as follows: the nomenclature is defined in Sect. 2; in Sect. 3, we define differential operators for the components of the third-order (disturbing) gravitational tensor in LNOF; in Sect. 4, we apply the differential operators to the integrals of Newton, Abel-Poisson, Pizzetti and Hotine and in Sect. 5, we investigate properties of corresponding integral kernels. Finally, main contributions of the article are summarized in the conclusions.

2 Nomenclature

Before we start mathematical derivations, we define the nomenclature used throughout the article. Generally, a gravitational tensor and a gravitational gradient are synonyms, both composed of components. Components of the gravitational tensor will be referred to as gravitational gradients since this terminology has widely been used in geodesy. Whenever important, an order of gravitational gradients will also be specified to indicate the order (or rank) of the respec-

gravitational tensor. We consider only gravitational gradients referred to the spherical LNOF. Such a reference frame has a moving origin and it is defined by a right-handed orthogonal basis with the following orientation of axes: the x -axis points to the north, the y -axis points to the west and the z -axis is directed radially outward.

In the article, we will use angular geocentric spherical coordinates with the triad (r, Ω) , where r is the geocentric radius and $\Omega = (\varphi, \lambda)$ represents a pair of angular spherical geocentric coordinates, i.e. spherical latitude φ and longitude λ , to designate the position of an evaluation point. The evaluation point stands for a point in which an integral formula is evaluated. The primed spherical triad (r', Ω') stands for the position of an integration volume element, i.e. an infinitesimal tesseroidal volume $r'^2 dr' d\Omega'$. For a constant value of $r' = R$, where R represents the radius of a mean geocentric sphere approximating the geoid, the spherical triad (R, Ω') designates the position of an integration surface element, i.e. an infinitesimal spherical surface $R^2 d\Omega'$.

Except for the angular geocentric spherical coordinates, we define the angular spherical polar coordinates $\alpha = \alpha(\Omega, \Omega')$, $\alpha' = \alpha'(\Omega, \Omega')$ and $\psi = \psi(\Omega, \Omega')$, see Fig. 1. The symbol α abbreviates the direct azimuth, i.e. the azimuth between an evaluation point and an integration element. It is measured at the evaluation point clockwise from the north. The symbol α' stands for the backward azimuth, i.e. the azimuth between an integration element and an evaluation point as measured at the integration element clockwise from the north. The symbol ψ abbreviates the spherical distance between an evaluation point and an integration element.

We also introduce the following substitutions—unitless variables:

$$t = t(r, r') = \frac{r'}{r}, \quad r' \leq r, \tag{1}$$

$$u = u(\Omega, \Omega') = \cos \psi, \tag{2}$$

$$g = g(t, u) = \sqrt{1 - 2tu + t^2}, \tag{3}$$

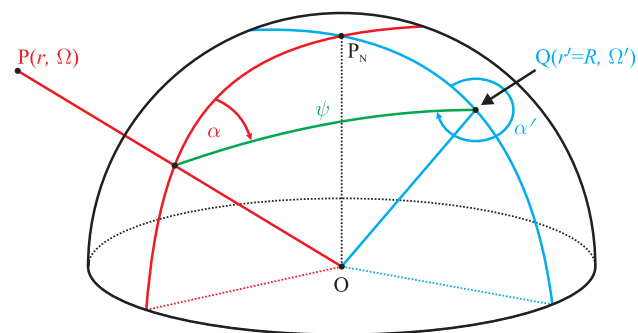


Fig. 1 Graphical illustration of the angular spherical polar coordinates, i.e. the direct azimuth α , the backward azimuth α' and the spherical distance ψ , between the points P and Q . P designates an evaluation point and Q is the centre of an integration surface element. P_N is the North pole and O is the centre of the mean geocentric sphere with radius R

where t is the attenuation factor, u is a cosine of the spherical distance and g represents the normalised Euclidean distance. These substitutions allow for simpler representations of mathematical formulas.

3 Differential operators for third-order gravitational gradients

A third-order gravitational tensor is generally represented by 27 gravitational gradients $V_{\mu\nu\tau}$; $\mu, \nu, \tau \in \{x, y, z\}$. Due to continuity of the gravitational field, the third-order gravitational tensor is symmetric with respect to the pairs of indices (μ, ν) , (ν, τ) and (μ, τ) , i.e. it holds $V_{\mu\nu\tau} = V_{\nu\mu\tau}$, $V_{\mu\nu\tau} = V_{\mu\tau\nu}$ and $V_{\mu\nu\tau} = V_{\tau\nu\mu}$. By such a symmetry, the third-order gravitational tensor can be defined only by ten gravitational gradients, see, e.g. (Tóth 2005). Another simplification comes from the Laplace differential equation that holds for the gravitational potential in mass-free space and by which $V_{xxv} + V_{yyv} + V_{zzv} = 0$, $v \in \{x, y, z\}$, see (Casotto and Fantino 2009). In such a case, the third-order gravitational tensor is composed of seven independent components.

Throughout the article, we will provide explicit expressions for ten gravitational gradients of the third order. These may further be divided into the four groups, i.e.

- horizontal–horizontal–horizontal (HHH) if $\mu, \nu, \tau \in \{x, y\}$,
- horizontal–horizontal–vertical (HHV) if $\mu, \nu \in \{x, y\}$ and $\tau = z$,
- horizontal–vertical–vertical (HVV) if $\mu \in \{x, y\}$ and $\nu, \tau = z$,
- and vertical–vertical–vertical (VVV) if $\mu, \nu, \tau = z$.

Obviously, the third-order gravitational tensor is composed of four HHH, three HHV, two HVV and one VVV gravitational gradients.

The gravitational tensor of the third order is obtained by the successive application of the gradient operator to the gravitational potential. Therefore, the ten gravitational gradients are defined by differential operators. Such differential operators in terms of the spherical geocentric coordinates (r, Ω) read as follows (Tóth 2005; Casotto and Fantino 2009):

$$D^{xxx} = -\frac{1}{r^2} \left(\frac{2}{r} \frac{\partial}{\partial \varphi} - 3 \frac{\partial^2}{\partial r \partial \varphi} - \frac{1}{r} \frac{\partial^3}{\partial \varphi^3} \right), \tag{4}$$

$$D^{xxy} = -\frac{1}{r^2 \cos \varphi} \left(\frac{2 \tan^2 \varphi}{r} \frac{\partial}{\partial \lambda} + \frac{\partial^2}{\partial r \partial \lambda} + \frac{2 \tan \varphi}{r} \frac{\partial^2}{\partial \varphi \partial \lambda} + \frac{1}{r} \frac{\partial^3}{\partial \varphi^2 \partial \lambda} \right), \tag{5}$$

$$D^{xxz} = -\frac{1}{r} \left(\frac{1}{r} \frac{\partial}{\partial r} - \frac{\partial^2}{\partial r^2} + \frac{2}{r^2} \frac{\partial^2}{\partial \varphi^2} - \frac{1}{r} \frac{\partial^3}{\partial r \partial \varphi^2} \right), \tag{6}$$

$$\mathcal{D}^{xyy} = -\frac{1}{r^2} \left(\frac{1}{r \cos^2 \varphi} \frac{\partial}{\partial \varphi} - \frac{\partial^2}{\partial r \partial \varphi} + \frac{\tan \varphi}{r} \frac{\partial^2}{\partial \varphi^2} - \frac{2 \tan \varphi}{r \cos^2 \varphi} \frac{\partial^2}{\partial \lambda^2} - \frac{1}{r \cos^2 \varphi} \frac{\partial^3}{\partial \varphi \partial \lambda^2} \right), \quad (7)$$

$$\mathcal{D}^{xyz} = \frac{1}{r^2 \cos \varphi} \left(\frac{2 \tan \varphi}{r} \frac{\partial}{\partial \lambda} - \tan \varphi \frac{\partial^2}{\partial r \partial \lambda} + \frac{2}{r} \frac{\partial^2}{\partial \varphi \partial \lambda} - \frac{\partial^3}{\partial r \partial \varphi \partial \lambda} \right), \quad (8)$$

$$\mathcal{D}^{xzz} = \frac{1}{r} \left(\frac{2}{r^2} \frac{\partial}{\partial \varphi} - \frac{2}{r} \frac{\partial^2}{\partial r \partial \varphi} + \frac{\partial^3}{\partial r^2 \partial \varphi} \right), \quad (9)$$

$$\mathcal{D}^{yyy} = \frac{1}{r^2 \cos \varphi} \left(\frac{2}{r \cos^2 \varphi} \frac{\partial}{\partial \lambda} - 3 \frac{\partial^2}{\partial r \partial \lambda} + \frac{3 \tan \varphi}{r} \frac{\partial^2}{\partial \varphi \partial \lambda} - \frac{1}{r \cos^2 \varphi} \frac{\partial^3}{\partial \lambda^3} \right), \quad (10)$$

$$\mathcal{D}^{yyz} = -\frac{1}{r} \left(\frac{1}{r} \frac{\partial}{\partial r} - \frac{2 \tan \varphi}{r^2} \frac{\partial}{\partial \varphi} - \frac{\partial^2}{\partial r^2} + \frac{\tan \varphi}{r} \frac{\partial^2}{\partial r \partial \varphi} + \frac{2}{r^2 \cos^2 \varphi} \frac{\partial^2}{\partial \lambda^2} - \frac{1}{r \cos^2 \varphi} \frac{\partial^3}{\partial r \partial \lambda^2} \right), \quad (11)$$

$$\mathcal{D}^{yzz} = -\frac{1}{r \cos \varphi} \left(\frac{2}{r^2} \frac{\partial}{\partial \lambda} - \frac{2}{r} \frac{\partial^2}{\partial r \partial \lambda} + \frac{\partial^3}{\partial r^2 \partial \lambda} \right), \quad (12)$$

$$\mathcal{D}^{zzz} = \frac{\partial^3}{\partial r^3}. \quad (13)$$

The superscripts on the left-hand sides of Eqs. (4)–(13) indicate gravitational gradients that are obtained by applying the differential operators. Note that these differential operators were originally derived in a right-handed spherical LNOF with the x -axis pointing eastwards and the y -axis pointing north.

The differential operators of Eqs. (4)–(13) are composed of derivatives with respect to the spherical coordinates r , φ and λ up to the third order. These are of a particular interest when the gravitational potential is expanded in a spherical harmonic series. For parametrization of the Earth's gravitational field, e.g. by integral transformations, least-squares collocation or radial basis function approach, other representations, such as those in the spherical polar coordinates (r, ψ, α) or in terms of (t, u, α) , are more appropriate. These lead to decomposition of the differential operators in Eqs. (4)–(13) into an azimuthal and an isotropic part and significantly simplify subsequent mathematical derivations and numerical evaluation.

At least two different ways may be followed to get the differential operators for the third-order gravitational gradients in terms of (r, ψ, α) or (t, u, α) . In the first approach, one starts with Eqs. (4)–(13) and applies the chain rule up to the third-order derivatives with respect to r , φ and λ . In the second approach, recursive relations between the differential operators for the second-order gravitational gradients, which are summarized in Appendix A, and their third-order

counterparts are sought. Such a transition involves only first-order differentiation with respect to r , φ and λ which may be handled more easily by the chain rule again.

Our experience has shown that the second approach was less laborious; therefore, it will be followed in this article below. To start with the second approach, we need the recursive relation between the differential operators for the second- and third-order gravitational gradients. This relation is defined by the following proposition:

Proposition 1 *The recursive relationships between the differential operators for the second- and third-order gravitational gradients are:*

$$\begin{aligned} \mathcal{D}^{xxx} &= \frac{1}{r} \left(2\mathcal{D}^{xz} + \frac{\partial}{\partial \varphi} \mathcal{D}^{xx} \right), \quad \mathcal{D}^{xxy} = \frac{1}{r} \left(\mathcal{D}^{yz} + \frac{\partial}{\partial \varphi} \mathcal{D}^{xy} \right), \\ \mathcal{D}^{xxz} &= \frac{\partial}{\partial r} \mathcal{D}^{xx}, \quad \mathcal{D}^{xyy} = \frac{1}{r} \frac{\partial}{\partial \varphi} \mathcal{D}^{yy}, \\ \mathcal{D}^{xyz} &= \frac{\partial}{\partial r} \mathcal{D}^{xy}, \quad \mathcal{D}^{xzz} = \frac{\partial}{\partial r} \mathcal{D}^{xz}, \\ \mathcal{D}^{yyy} &= \frac{1}{r} \left[2(\mathcal{D}^{yz} - \tan \varphi \mathcal{D}^{xy}) - \frac{1}{\cos \varphi} \frac{\partial}{\partial \lambda} \mathcal{D}^{yy} \right], \\ \mathcal{D}^{yyz} &= \frac{\partial}{\partial r} \mathcal{D}^{yy}, \quad \mathcal{D}^{yzz} = \frac{\partial}{\partial r} \mathcal{D}^{yz}, \quad \mathcal{D}^{zzz} = \frac{\partial}{\partial r} \mathcal{D}^{zz}. \end{aligned} \quad (14)$$

One may prove Proposition 1 by applying Eq. (14) to those of Eqs. (71)–(76). After performing differentiations and algebraic manipulations, we get expressions for the differential operators of the third-order gravitational gradients defined by Eqs. (4)–(13). Alternatively, Proposition 1 may be proven by the recursive formula of Casotto and Fantino (2009, Eq. 24) requiring the Christoffel symbols of the second kind.

It can be seen from Eq. (14) that the differential operators for the HHV, HVV and VVV gravitational gradients are obtained by the first-order radial derivative of the corresponding second-order differential operators. Generally, the differential operators for the HHH gravitational gradients are linear combinations of two terms: the first term may be either zero, another second-order differential operator or a combination of second-order differential operators; the second term is a first-order derivative either with respect to φ or λ of a second-order differential operator.

We can now formulate representation of the differential operators for the third-order gravitational gradients in terms of the spherical polar triad (r, ψ, α) and in terms of (t, u, α) :

Proposition 2 *The differential operators for the third-order gravitational gradients in terms of the spherical polar triad (r, ψ, α) and in terms of (t, u, α) are:*

$$\begin{aligned} \mathcal{D}^{xxx} &= \frac{1}{r^3} \left(\cos \alpha \mathcal{D}_3^1 + \cos 3\alpha \mathcal{D}_3^2 \right), \\ \mathcal{D}^{xxy} &= -\frac{1}{r^3} \left(\frac{1}{3} \sin \alpha \mathcal{D}_3^1 + \sin 3\alpha \mathcal{D}_3^2 \right), \\ \mathcal{D}^{xxz} &= \frac{1}{r^3} \left(\mathcal{D}_3^3 + \cos 2\alpha \mathcal{D}_3^4 \right), \end{aligned}$$

$$\begin{aligned}
 \mathcal{D}^{xyy} &= \frac{1}{r'^3} \left(\frac{1}{3} \cos \alpha \mathcal{D}_3^1 - \cos 3\alpha \mathcal{D}_3^2 \right), \\
 \mathcal{D}^{xyz} &= -\frac{1}{r'^3} \sin 2\alpha \mathcal{D}_3^4, \\
 \mathcal{D}^{xzz} &= \frac{1}{r'^3} \cos \alpha \mathcal{D}_3^5, \quad \mathcal{D}^{yyy} = -\frac{1}{r'^3} \left(\sin \alpha \mathcal{D}_3^1 - \sin 3\alpha \mathcal{D}_3^2 \right), \\
 \mathcal{D}^{yyz} &= \frac{1}{r'^3} \left(\mathcal{D}_3^3 - \cos 2\alpha \mathcal{D}_3^4 \right), \\
 \mathcal{D}^{yzz} &= -\frac{1}{r'^3} \sin \alpha \mathcal{D}_3^5, \quad \mathcal{D}^{zzz} = \frac{1}{r'^3} \mathcal{D}_3^6,
 \end{aligned} \tag{15}$$

where

$$\begin{aligned}
 \mathcal{D}_3^1 &= \frac{3r'^3}{r^2} \left[\frac{1}{4r} \left(2 + \frac{1}{\sin^2 \psi} \right) \frac{\partial}{\partial \psi} - \frac{\partial^2}{\partial r \partial \psi} \right. \\
 &\quad \left. - \frac{1}{4r} \left(\cot \psi \frac{\partial^2}{\partial \psi^2} + \frac{\partial^3}{\partial \psi^3} \right) \right] \\
 &= -3t^3 \sqrt{1-u^2} \left(\frac{\partial}{\partial u} + t \frac{\partial^2}{\partial t \partial u} + u \frac{\partial^2}{\partial u^2} - \frac{1-u^2}{4} \frac{\partial^3}{\partial u^3} \right), \tag{16}
 \end{aligned}$$

$$\begin{aligned}
 \mathcal{D}_3^2 &= -\frac{r'^3}{4r^3} \left[\left(\frac{1}{\sin^2 \psi} + 2 \cot^2 \psi \right) \frac{\partial}{\partial \psi} \right. \\
 &\quad \left. - 3 \cot \psi \frac{\partial^2}{\partial \psi^2} + \frac{\partial^3}{\partial \psi^3} \right] = \frac{t^3 \sqrt{(1-u^2)^3}}{4} \frac{\partial^3}{\partial u^3}, \tag{17}
 \end{aligned}$$

$$\begin{aligned}
 \mathcal{D}_3^3 &= -\frac{r'^3}{r} \left(\frac{1}{r} \frac{\partial}{\partial r} + \frac{\cot \psi}{r^2} \frac{\partial}{\partial \psi} - \frac{\partial^2}{\partial r^2} - \frac{\cot \psi}{2r} \frac{\partial^2}{\partial r \partial \psi} \right. \\
 &\quad \left. + \frac{1}{r^2} \frac{\partial^2}{\partial \psi^2} - \frac{1}{2r} \frac{\partial^3}{\partial r \partial \psi^2} \right) \\
 &= t^3 \left[3t \frac{\partial}{\partial t} + 2u \frac{\partial}{\partial u} + tu \frac{\partial^2}{\partial t \partial u} + t^2 \frac{\partial^2}{\partial t^2} \right. \\
 &\quad \left. - (1-u^2) \left(\frac{\partial^2}{\partial u^2} + \frac{t}{2} \frac{\partial^3}{\partial t \partial u^2} \right) \right], \tag{18}
 \end{aligned}$$

$$\begin{aligned}
 \mathcal{D}_3^4 &= \frac{r'^3}{r^2} \left(\frac{\cot \psi}{r} \frac{\partial}{\partial \psi} - \frac{\cot \psi}{2} \frac{\partial^2}{\partial r \partial \psi} - \frac{1}{r} \frac{\partial^2}{\partial \psi^2} + \frac{1}{2} \frac{\partial^3}{\partial r \partial \psi^2} \right) \\
 &= -t^3 (1-u^2) \left(\frac{\partial^2}{\partial u^2} + \frac{t}{2} \frac{\partial^3}{\partial t \partial u^2} \right), \tag{19}
 \end{aligned}$$

$$\begin{aligned}
 \mathcal{D}_3^5 &= -\frac{r'^3}{r} \left(\frac{2}{r^2} \frac{\partial}{\partial \psi} - \frac{2}{r} \frac{\partial^2}{\partial r \partial \psi} + \frac{\partial^3}{\partial r^2 \partial \psi} \right) \\
 &= t^3 \sqrt{1-u^2} \left(2 \frac{\partial}{\partial u} + 4t \frac{\partial^2}{\partial t \partial u} + t^2 \frac{\partial^3}{\partial t^2 \partial u} \right), \tag{20}
 \end{aligned}$$

$$\mathcal{D}_3^6 = r'^3 \frac{\partial^3}{\partial r^3} = -t^4 \left(6 \frac{\partial}{\partial t} + 6t \frac{\partial^2}{\partial t^2} + t^2 \frac{\partial^3}{\partial t^3} \right). \tag{21}$$

The proof of Proposition 2 may be performed in several steps. Firstly, the differential operators of Eq. (77) are substituted into those of Eq. (14). Secondly, we make use of the following expressions, see, e.g. (Heiskanen and Moritz 1967, p. 113):

$$\frac{\partial}{\partial \varphi} = -\cos \alpha \frac{\partial}{\partial \psi}, \quad \frac{1}{\cos \varphi} \frac{\partial}{\partial \lambda} = -\sin \alpha \frac{\partial}{\partial \psi}, \tag{22}$$

and, e.g. (Winch and Roberts 1995):

$$\begin{aligned}
 \frac{\partial \sin 2\alpha}{\partial \varphi} &= 2 \cot \psi \cos 2\alpha \sin \alpha, \\
 \frac{\partial \cos 2\alpha}{\partial \varphi} &= -2 \cot \psi \sin 2\alpha \sin \alpha, \\
 \frac{\partial \cos 2\alpha}{\partial \lambda} &= \frac{2 \cos \varphi' \sin 2\alpha \cos \alpha'}{\sin \psi}.
 \end{aligned} \tag{23}$$

After performing some algebraic operations, we get the expressions for the differential operators of the third-order gravitational gradients in terms of the triad (r, ψ, α) . Such a representation is defined by Eq. (15) and the first equalities of \mathcal{D}_3^i ; $i = 1, \dots, 6$ of Eqs. (16)–(21). Thirdly, Eqs. (16)–(21) are expressed in terms of the substitutions t and u , see Eqs. (1) and (2), applying again the chain rule relations:

$$\begin{aligned}
 \frac{\partial}{\partial r} &= -\frac{t^2}{r'} \frac{\partial}{\partial t}, \quad \frac{\partial}{\partial \psi} = -\sqrt{1-u^2} \frac{\partial}{\partial u}, \\
 \frac{\partial^2}{\partial r^2} &= \frac{t^3}{r'^2} \left(2 \frac{\partial}{\partial t} + t \frac{\partial^2}{\partial t^2} \right), \quad \frac{\partial^2}{\partial r \partial \psi} = \frac{t^2 \sqrt{1-u^2}}{r'} \frac{\partial^2}{\partial t \partial u}, \\
 \frac{\partial^2}{\partial \psi^2} &= -u \frac{\partial}{\partial u} + (1-u^2) \frac{\partial^2}{\partial u^2}, \\
 \frac{\partial^3}{\partial r^3} &= -\frac{t^4}{r'^3} \left(6 \frac{\partial}{\partial t} + 6t \frac{\partial^2}{\partial t^2} + t^2 \frac{\partial^3}{\partial t^3} \right), \\
 \frac{\partial^3}{\partial r^2 \partial \psi} &= -\frac{t^3 \sqrt{1-u^2}}{r'^2} \left(2 \frac{\partial^2}{\partial t \partial u} + t \frac{\partial^3}{\partial t^2 \partial u} \right), \\
 \frac{\partial^3}{\partial r \partial \psi^2} &= \frac{t^2}{r'} \left[u \frac{\partial^2}{\partial t \partial u} - (1-u^2) \frac{\partial^3}{\partial t \partial u^2} \right], \\
 \frac{\partial^3}{\partial \psi^3} &= \sqrt{1-u^2} \left[\frac{\partial}{\partial u} + 3u \frac{\partial^2}{\partial u^2} - (1-u^2) \frac{\partial^3}{\partial u^3} \right].
 \end{aligned} \tag{24}$$

Note that the subscript on the left-hand sides of Eqs. (16)–(21) represents the respective order of the gravitational tensor which distinguishes the operators of Eqs. (16)–(21) from those of Eqs. (78)–(81).

Proposition 2 demonstrates that the differential operators for the third-order gravitational gradients may efficiently be decomposed into the azimuthal and isotropic parts. The azimuthal part is represented by trigonometric functions of multiples of the direct azimuth α . The maximum multiple of the direct azimuth α corresponds to the order of the gravitational tensor. The isotropic part is defined by the six differential operators \mathcal{D}_3^i ; $i = 1, \dots, 6$ that are rotationally symmetric about the radial direction, i.e. azimuthally isotropic. We will refer to such differential operators as isotropic.

Moreover, the inspection of Eq. (15) reveals that for one group of the gravitational gradients (either HHH, HHV, HVV or VVV), isotropic differential operators occur repeatedly. For example, the four HHH gravitational gradients are expressed by the isotropic differential operators \mathcal{D}_3^1 and \mathcal{D}_3^2 . It may also be seen that some pairs of the differential

operators, i.e. $(D^{xxx}, D^{yyy}), (D^{xyy}, D^{xxy}), (D^{xzz}, D^{yzz})$ and (D^{xxz}, D^{yyz}) , are symmetric with respect to the rotation of the direct azimuth α . The first three pairs are interrelated by the rotation of $\pi/2$, e.g. we get D^{yyy} by considering $\alpha + \pi/2$ for the differential operator D^{xxx} . The last pair of the differential operators is rotationally symmetric by rotating α by angle $\pm\pi/2$.

The differential operators of Proposition 2 may be simplified for a certain class of isotropic kernels to which they are applied. The more concise definition is defined by the following proposition:

Proposition 3 *For the isotropic kernels of the series form:*

$$\mathcal{K}(t, u) = \sum_{n=0}^{\infty} t^{n+1} h_n P_{n,0}(u), \tag{25}$$

with h_n being eigenvalues of the isotropic kernels, the differential operators for the third-order gravitational gradients may be expressed as:

$$\begin{aligned} D^{xxx} &= \frac{1}{r^3} \left(\cos \alpha D_3^{1*} + \cos 3\alpha D_3^{3*} \right), \\ D^{xxy} &= -\frac{1}{r^3} \left(\frac{1}{3} \sin \alpha D_3^{1*} + \sin 3\alpha D_3^{3*} \right), \\ D^{xxz} &= \frac{1}{r^3} \left(D_3^{0*} + \cos 2\alpha D_3^{2*} \right), \\ D^{xyy} &= \frac{1}{r^3} \left(\frac{1}{3} \cos \alpha D_3^{1*} - \cos 3\alpha D_3^{3*} \right), \\ D^{xyz} &= -\frac{1}{r^3} \sin 2\alpha D_3^{2*}, \quad D^{xzz} = -\frac{4}{3r^3} \cos \alpha D_3^{1*}, \\ D_3^{yyy} &= -\frac{1}{r^3} \left(\sin \alpha D_3^{1*} - \sin 3\alpha D_3^{3*} \right), \\ D^{yyz} &= \frac{1}{r^3} \left(D_3^{0*} - \cos 2\alpha D_3^{2*} \right), \quad D^{yzz} = \frac{4}{3r^3} \sin \alpha D_3^{1*}, \\ D^{zzz} &= -\frac{2}{r^3} D_3^{0*}, \end{aligned} \tag{26}$$

where

$$D_3^{0*} = D_3^3 = -\frac{1}{2} D_3^6 = \frac{t^4}{2} \frac{\partial^3}{\partial t^3} \left(t^2 \cdot \right), \tag{27}$$

$$D_3^{1*} = D_3^1 = -\frac{3}{4} D_3^5 = -\frac{3t^3 \sqrt{1-u^2}}{4} \frac{\partial^3}{\partial t^2 \partial u} \left(t^2 \cdot \right), \tag{28}$$

$$D_3^{2*} = D_3^4 = -\frac{t^2(1-u^2)}{2} \frac{\partial^3}{\partial t \partial u^2} \left(t^2 \cdot \right), \tag{29}$$

$$D_3^{3*} = D_3^2 = \frac{t^3 \sqrt{(1-u^2)^3}}{4} \frac{\partial^3}{\partial u^3}. \tag{30}$$

To prove Proposition 3, the isotropic differential operators D_3^2, D_3^4, D_3^5 and D_3^6 of Eqs. (16)–(21) are rewritten in the form of Eqs. (27)–(30) that follows from the rules of differentiation. It can be seen that the isotropic differential operators of Eqs. (27)–(30) have been renamed as indicated by different numbers in the superscript followed by an asterisk. The number in the superscript corresponds to the order of

differentiation with respect to u . Note that the symbol $(t^2 \cdot)$ under differentiation in Eqs. (27)–(29) stands for multiplication of t^2 and the isotropic kernel of Eq. (25). To complete the proof of the proposition, we have to show that $D_3^3 = -\frac{1}{2} D_3^6$ and $D_3^1 = -\frac{3}{4} D_3^5$. These equalities follow by applying the isotropic differential operators D_3^1, D_3^3, D_3^5 and D_3^6 of Eqs. (16), (18), (20) and (21) to the kernel of Eq. (25). Note that singular terms of the type $\frac{u}{\sqrt{1-u^2}} P_{n,m}(u)$ for $m = 1, 2$ occur after the application of D_3^1 and D_3^3 . These can be removed by the following identity (Arfken 1968, p. 456):

$$\begin{aligned} \frac{u}{\sqrt{1-u^2}} P_{n,m}(u) &= \frac{1}{2m} \left[(n+m)(n-m+1) P_{n,m-1}(u) \right. \\ &\quad \left. + P_{n,m+1}(u) \right]. \end{aligned} \tag{31}$$

Proposition 3 illustrates another symmetry of the differential operators for the third-order gravitational gradients. It can be seen from Eq. (26) that multiples of the direct azimuth α and the numbers in the superscripts of the corresponding isotropic differential operators agree. The simplification of Proposition 3 is based on the fact that only four isotropic differential operators are required compared to the six isotropic differential operators of Proposition 2. One has to keep in mind that this holds only if the differential operators act on the isotropic kernels of the form defined by Eq. (25).

4 Application to the integrals of Newton, Abel-Poisson, Pizzetti and Hotine

Having the differential operators in hand, we may relate the third-order gravitational gradients to various parameters of the Earth’s gravitational field. We now focus on the application of the differential operators to the integrals of Newton, Abel-Poisson, Pizzetti (also known as extended Stokes’s integral) and Hotine. The first of these transformations relates the volumetric mass density distribution ρ to the gravitational potential V by the volume integral, see, e.g. (Kellogg 1929; Martinec 1998):

$$V(r, \Omega) = G \int_{\Omega'} \int_{r'} \rho(r', \Omega') \mathcal{K}^N(t, u) r' dr' d\Omega', \tag{32}$$

where G stands for the Newtonian gravitational constant. Note that the first power of r' in Eq. (32), as opposed to the second power usually presented in geodetic literature, appears due to definition of the Newtonian kernel \mathcal{K}^N , see below. The other three surface integral transformations, i.e. Abel-Poisson (Heiskanen and Moritz 1967), Pizzetti (1911) and Hotine (1969), indicated by the respective superscripts \mathcal{P}, \mathcal{S} and \mathcal{H} , may concisely be written as:

$$T(r, \Omega) = \frac{1}{4\pi} \int_{\Omega'} T^j(R, \Omega') \mathcal{K}^j(t, u) d\Omega', \quad j \in \{\mathcal{P}, \mathcal{S}, \mathcal{H}\}, \quad (33)$$

where $T^{\mathcal{P}} = T$ is the disturbing gravitational potential, $T^{\mathcal{S}} = R\Delta g$, with Δg being the gravity anomaly and $T^{\mathcal{H}} = R\delta g$ with δg standing for the gravity disturbance.

The integral kernels for the four integral formulas of Eqs. (32) and (33) in the closed form read as follows:

$$\mathcal{K}^{\mathcal{N}}(t, u) = \frac{t}{g}, \quad (34)$$

$$\mathcal{K}^{\mathcal{P}}(t, u) = \frac{t(1-t^2)}{g^3}, \quad (35)$$

$$\mathcal{K}^{\mathcal{S}}(t, u) = t + \frac{2t}{g} - 3tg - t^2u \left[5 + 3 \ln \left(\frac{1-tu+g}{2} \right) \right], \quad (36)$$

$$\mathcal{K}^{\mathcal{H}}(t, u) = \frac{2t}{g} - \ln \left(\frac{t-u+g}{1-u} \right). \quad (37)$$

These correspond to the infinite series, i.e. the spectral forms:

$$\mathcal{K}^j(t, u) = \sum_{n=0}^{\infty} t^{n+1} h_n^j P_{n,0}(u), \quad j \in \{\mathcal{N}, \mathcal{P}, \mathcal{S}, \mathcal{H}\}, \quad (38)$$

where the corresponding eigenvalues h_n^j are:

$$h_n^{\mathcal{N}} = 1, \quad \forall n \geq 0, \quad (39)$$

$$h_n^{\mathcal{P}} = 2n + 1, \quad \forall n \geq 0, \quad (40)$$

$$h_0^{\mathcal{S}} = h_1^{\mathcal{S}} = 0, h_n^{\mathcal{S}} = \frac{2n + 1}{n - 1}, \quad \forall n \geq 2, \quad (41)$$

$$h_n^{\mathcal{H}} = \frac{2n + 1}{n + 1}, \quad \forall n \geq 0. \quad (42)$$

The symbol $P_{n,m}(u)$ stands for the non-normalised Legendre function of the first kind of degree n and order m .

Mathematical relation between the third-order (disturbing) gravitational gradients and the mass density distribution, disturbing gravitational potential, gravity anomaly and gravity disturbance is formulated in the following proposition:

Proposition 4 *Integral formulas, that relate mass density distribution to the third-order gravitational gradients, are:*

$$V_{\mu\nu\tau}(r, \Omega) = G \int_{\Omega'} \int_{r'} \rho(r', \Omega') \mathcal{K}^{\mu\nu\tau, N}(t, u) r'^{-2} dr' d\Omega', \quad \mu, \nu, \tau \in \{x, y, z\}. \quad (43)$$

Integral formulas transforming the disturbing gravitational potential, gravity anomaly and gravity disturbance to the third-order disturbing gravitational gradients are of the form:

$$T_{\mu\nu\tau}(r, \Omega) = \frac{1}{4\pi R^3} \int_{\Omega'} T^j(R, \Omega') \mathcal{K}^{\mu\nu\tau, j}(t, u) d\Omega', \quad \mu, \nu, \tau \in \{x, y, z\}, \quad j \in \{\mathcal{P}, \mathcal{S}, \mathcal{H}\}. \quad (44)$$

The sub-integral kernels $\mathcal{K}^{\mu\nu\tau, j} = \mathcal{D}^{\mu\nu\tau} \mathcal{K}^j$; $\mu, \nu, \tau \in \{x, y, z\}$; $j \in \{\mathcal{N}, \mathcal{P}, \mathcal{S}, \mathcal{H}\}$ of Eqs. (43) and (44) are defined as follows:

$$\begin{aligned} \mathcal{K}^{xxx, j} &= \cos \alpha \mathcal{K}^{1*, j} + \cos 3\alpha \mathcal{K}^{3*, j}, \\ \mathcal{K}^{xxy, j} &= -\frac{1}{3} \sin \alpha \mathcal{K}^{1*, j} - \sin 3\alpha \mathcal{K}^{3*, j}, \\ \mathcal{K}^{xxz, j} &= \mathcal{K}^{0*, j} + \cos 2\alpha \mathcal{K}^{2*, j}, \\ \mathcal{K}^{xyy, j} &= \frac{1}{3} \cos \alpha \mathcal{K}^{1*, j} - \cos 3\alpha \mathcal{K}^{3*, j}, \\ \mathcal{K}^{xyx, j} &= -\sin 2\alpha \mathcal{K}^{2*, j}, \quad \mathcal{K}^{xzz, j} = -\frac{4}{3} \cos \alpha \mathcal{K}^{1*, j} \\ \mathcal{K}^{yyx, j} &= -\sin \alpha \mathcal{K}^{1*, j} + \sin 3\alpha \mathcal{K}^{3*, j}, \\ \mathcal{K}^{yyz, j} &= \mathcal{K}^{0*, j} - \cos 2\alpha \mathcal{K}^{2*, j}, \quad \mathcal{K}^{yzz, j} = \frac{4}{3} \sin \alpha \mathcal{K}^{1*, j}, \\ \mathcal{K}^{zzz, j} &= -2\mathcal{K}^{0*, j}, \end{aligned} \quad (45)$$

where the isotropic kernels read:

$$\mathcal{K}^{i*, j}(t, u) = \mathcal{D}_3^{i*} \mathcal{K}^j(t, u); \quad i \in \{0, 1, 2, 3\}. \quad (46)$$

Proposition 4 may be proven by the action of the differential operators of Eq. (26) to Newton's, Abel-Poisson's, Pizzetti's and Hotine's integrals, see Eqs. (32) and (33). Application of the more concise differential operators of Proposition 3 is permitted because the kernels for these integrals, see Eq. (38), are of the form of Eq. (25). Note that $r' = R$ in Eq. (26) when these are applied to the integrals of Abel-Poisson, Pizzetti and Hotine. In total, 40 new integral formulas, of which 28 are independent due to the Laplace differential equation, are embedded in Eqs. (43) and (44). Obviously, the sub-integral kernels of Eqs. (45) possess the same properties as we have already discussed for the third-order differential operators, see the paragraphs below Propositions 2 and 3.

Connection between the gravitational field quantities of Eqs. (44)–(46) is depicted in Fig. 2 in terms of a Meissl-type scheme in the spatial domain. The scheme summarizes possible applications of Eqs. (44)–(46). The direct problem (upward continuation) allows for computing disturbing gravitational gradients from the disturbing gravitational potential, gravity anomaly or gravity disturbance. On the other hand, the disturbing gravitational potential, gravity anomaly or gravity disturbance are estimated from the disturbing gravitational gradients by solving the inverse problem (downward continuation). Similar scheme may be depicted also for the volumetric mass density and the gravitational gradients, i.e. Eqs. (43), (45) and (46), and would provide another extension of the Meissl scheme in the spatial domain.

Explicit expressions for the isotropic kernels of Eq. (46) are required for numerical evaluation of the integral formulas

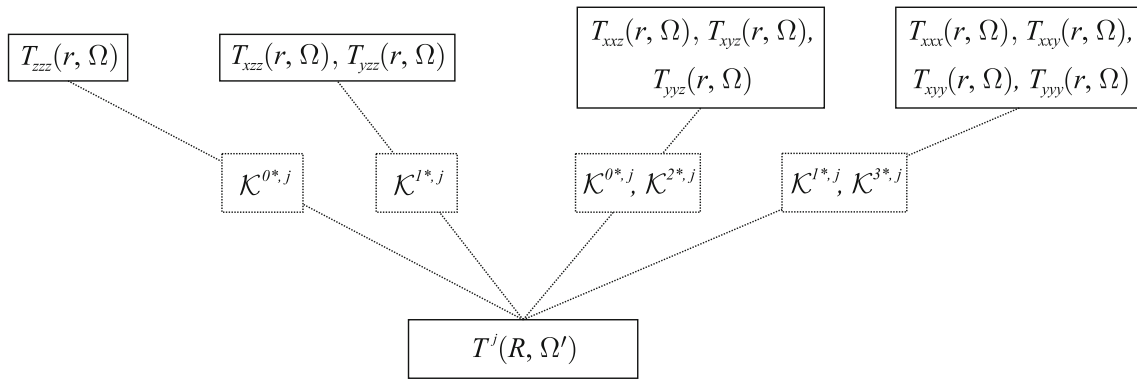


Fig. 2 Relation between the disturbing gravitational gradients and the disturbing gravitational potential, gravity anomaly and gravity disturbance. The disturbing gravitational gradients at the evaluation point with spherical coordinates (r, Ω) are located at the upper level. The symbol

$T^j(R, \Omega')$, $j \in \{\mathcal{P}, \mathcal{S}, \mathcal{H}\}$, at the lower level substitutes the disturbing gravitational potential, gravity anomaly and gravity disturbance on the surface of the mean geocentric sphere of radius R . The corresponding isotropic kernels are located at the middle level

of Eq. (43) and (44). Their spectral forms are defined in the next proposition:

Proposition 5 *Spectral forms of the isotropic kernels are:*

$$\mathcal{K}^{0*,j}(t, u) = \frac{1}{2} \sum_{n=0}^{\infty} t^{n+4} h_n^j (n+1)(n+2)(n+3) P_{n,0}(u), \quad (47)$$

$$\mathcal{K}^{1*,j}(t, u) = -\frac{3}{4} \sum_{n=0}^{\infty} t^{n+4} h_n^j (n+2)(n+3) P_{n,1}(u), \quad (48)$$

$$\mathcal{K}^{2*,j}(t, u) = -\frac{1}{2} \sum_{n=0}^{\infty} t^{n+4} h_n^j (n+3) P_{n,2}(u), \quad (49)$$

$$\mathcal{K}^{3*,j}(t, u) = \frac{1}{4} \sum_{n=0}^{\infty} t^{n+4} h_n^j P_{n,3}(u). \quad (50)$$

Looking at the general definition of the isotropic kernels in Eq. (46), we may prove Proposition 5 by applying the isotropic differential operators of Eqs. (27)–(30) to Eq. (38). Proposition 5 demonstrates a very general definition of the isotropic kernels in the spectral form. Corresponding expressions for $j \in \{\mathcal{N}, \mathcal{P}, \mathcal{S}, \mathcal{H}\}$ may easily be obtained by inserting the eigenvalues of Eqs. (39)–(42) into Eqs. (47)–(50).

The closed (spatial) forms of the isotropic kernels must be given separately for each of the superscripts i and j . These are summarized in the next proposition:

Proposition 6 *The closed (spatial) forms of the isotropic kernels are:*

(a) for $j = \mathcal{N}$:

$$\mathcal{K}^{0*,\mathcal{N}}(t, u) = \frac{3t^4}{2g} \left\{ 2 + \frac{t}{g^2} \left[3(2u - 3t) + \frac{3t(t-u)(4t-3u)}{g^2} - \frac{5t^2(t-u)^3}{g^4} \right] \right\}, \quad (51)$$

$$\mathcal{K}^{1*,\mathcal{N}}(t, u)$$

$$= -\frac{9t^5 \sqrt{1-u^2}}{4g^3} \left\{ 4 + \frac{t}{g^2} \left[8u - 9t + \frac{5t(t-u)^2}{g^2} \right] \right\}, \quad (52)$$

$$\mathcal{K}^{2*,\mathcal{N}}(t, u) = -\frac{15t^6(1-u^2)}{2g^5} \left[1 + \frac{t(u-t)}{g^2} \right], \quad (53)$$

$$\mathcal{K}^{3*,\mathcal{N}}(t, u) = \frac{15t^7 \sqrt{(1-u^2)^3}}{4g^7}, \quad (54)$$

(b) for $j = \mathcal{P}$:

$$\mathcal{K}^{0*,\mathcal{P}}(t, u) = \frac{3t^4}{2g} \left\{ 2 + \frac{t}{g^2} \left[22u - 49t + \frac{3t[(t-u)(54t-25u) + 2u^2]}{g^2} - \frac{5t^2(t-u)^2(37t-25u)}{g^4} + \frac{70t^3(t-u)^4}{g^6} \right] \right\}, \quad (55)$$

$$\mathcal{K}^{1*,\mathcal{P}}(t, u) = -\frac{9t^5 \sqrt{1-u^2}}{4g^3} \left\{ 12 + \frac{t}{g^2} \left[64u - 87t + \frac{5t(t-u)(29t-23u)}{g^2} - \frac{70t^2(t-u)^3}{g^4} \right] \right\}, \quad (56)$$

$$\mathcal{K}^{2*,\mathcal{P}}(t, u) = -\frac{15t^6(1-u^2)}{2g^5} \left\{ 5 + \frac{t}{g^2} \left[17u - 19t + \frac{14t(t-u)^2}{g^2} \right] \right\}, \quad (57)$$

$$\mathcal{K}^{3*,\mathcal{P}}(t, u) = \frac{105t^7(1-t^2)\sqrt{(1-u^2)^3}}{4g^9}, \quad (58)$$

(c) for $j = \mathcal{S}$:

$$\mathcal{K}^{0*,\mathcal{S}}(t, u) = 3t^4 \left\{ \frac{11}{g} - g(12-g) - t^2 \right\}$$

$$\begin{aligned}
 & -tu \left[31 + 12 \ln \left(\frac{1+g-tu}{2} \right) \right] \\
 & + \frac{t}{2g^3} \left[20u - 27t + \frac{3t(t-u)(9t-7u)}{g^2} \right. \\
 & \left. - \frac{10t^2(t-u)^3}{g^4} \right], \tag{59}
 \end{aligned}$$

$$\begin{aligned}
 \mathcal{K}^{1*,\mathcal{S}}(t,u) = & \frac{9t^5\sqrt{1-u^2}}{4} \left\{ 3 \left[13 + 4 \ln \left(\frac{1+g-tu}{2} \right) \right. \right. \\
 & \left. \left. - \frac{4(g+1)^2}{g(1+g-tu)} \right] \right. \\
 & \left. - \frac{1}{g^3} \left[15 + \frac{t(19u-21t)}{g^2} + \frac{10t^2(t-u)^2}{g^4} \right] \right\}, \tag{60}
 \end{aligned}$$

$$\begin{aligned}
 \mathcal{K}^{2*,\mathcal{S}}(t,u) = & -\frac{3t^5(1-u^2)}{2} \left\{ \frac{4t}{1+g-tu} \left[1 + \frac{1}{g^3} \right. \right. \\
 & \left. \left. + \frac{(g+1)^3}{g^2(1+g-tu)} \right] + \frac{1}{g^5} \left[13t + \frac{10t^2(u-t)}{g^2} \right] \right\}, \tag{61}
 \end{aligned}$$

$$\begin{aligned}
 \mathcal{K}^{3*,\mathcal{S}}(t,u) = & \frac{3t^7\sqrt{(1-u^2)^3}}{4} \left\{ \frac{10}{g^7} + \frac{1}{1+g-tu} \left(\frac{3}{g^5} + \frac{1}{1+g-tu} \right) \right. \\
 & \left. + \frac{1}{g^3(1+g-tu)^2} \left[\frac{4g+3}{g} + \frac{2(g+1)^4}{1+g-tu} \right] \right\}, \tag{62}
 \end{aligned}$$

(d) for $j = \mathcal{H}$:

$$\begin{aligned}
 \mathcal{K}^{0*,\mathcal{H}}(t,u) = & \frac{t^4}{2g} \left\{ 6 + \frac{t}{g^2} \left[30u - 47t \right. \right. \\
 & \left. \left. + \frac{3t(t-u)(23t-17u)}{g^2} - \frac{30t^2(t-u)^3}{g^4} \right] \right\} \tag{63}
 \end{aligned}$$

$$\begin{aligned}
 \mathcal{K}^{1*,\mathcal{H}}(t,u) = & \frac{3t^3\sqrt{1-u^2}}{2} \left\{ \frac{1}{1-u} - \frac{g+t}{g(g+t-u)} \right. \\
 & \left. - \frac{t^2}{2g^3} \left[19 + \frac{3t(15u-17t)}{g^2} + \frac{30t^2(t-u)^2}{g^4} \right] \right\}, \tag{64}
 \end{aligned}$$

$$\begin{aligned}
 \mathcal{K}^{2*,\mathcal{H}}(t,u) = & t^3(1-u^2) \left\{ \frac{1}{(1-u)^2} - \frac{1}{g^2(g+t-u)} \right. \\
 & \left. \times \left[\frac{t^2}{g} + \frac{(g+t)^2}{g+t-u} \right] - \frac{3t^3}{2g^5} \left[9 + \frac{10t(u-t)}{g^2} \right] \right\}, \tag{65}
 \end{aligned}$$

$$\begin{aligned}
 \mathcal{K}^{3*,\mathcal{H}}(t,u) = & \frac{t^3\sqrt{(1-u^2)^3}}{4} \left\{ \frac{30t^4}{g^7} - \frac{2}{(1-u)^3} + \frac{1}{g^3(g+t-u)} \right. \\
 & \left. \times \left[\frac{3t^3}{g^2} + \frac{3t^2(g+t)}{g(g+t-u)} + \frac{2(g+t)^3}{(g+t-u)^2} \right] \right\}. \tag{66}
 \end{aligned}$$

Each of Eqs. (51)–(66) may be proven by summing up the corresponding infinite series representation defined in Proposition 5. We describe details of such a derivation for one of the isotropic kernels, e.g. $\mathcal{K}^{2*,\mathcal{S}}$. First, the corresponding eigenvalue of Eq. (41) is inserted into Eq. (49) and the degree-dependent term, i.e. the term $h_n^{\mathcal{S}}(n+3)$, is decomposed into partial fractions. Second, based on the partial fraction decomposition in Eq. (92), the original infinite summation is divided into three separate series. Each of the three series is defined by closed expressions defined by Eqs. (110), (111) and (125). Thirdly, making use of the closed expressions and

performing some algebraic operations, we get the isotropic kernel $\mathcal{K}^{2*,\mathcal{S}}$ in closed form, see Eq. (61). Closed forms of the other isotropic kernels may be obtained by repeating the same steps. Due to brevity, we left this task for an interested reader who may find all necessary formulas in Appendices B and C.

Proposition 6 may be complemented by inspecting limits of the isotropic kernels for some special values of u and t . We have inspected four special cases that are concisely summarized in the next proposition:

Proposition 7 Limits of the isotropic kernels $\mathcal{K}^{i*,j}$; $i \in \{0, 1, 2, 3\}$; $j \in \{\mathcal{N}, \mathcal{P}, \mathcal{S}, \mathcal{H}\}$ are:

(a) for $u = 1$ and $t < 1$:

$$\begin{aligned}
 \mathcal{K}^{0*,\mathcal{N}}(t,1) = & \frac{3t^4}{(1-t)^4}, \quad \mathcal{K}^{0*,\mathcal{P}}(t,1) = \frac{3t^4(1+7t)}{(1-t)^5}, \\
 \mathcal{K}^{0*,\mathcal{S}}(t,1) = & 3t^5 \left[\frac{t(38-88t+73t^2-21t^3)}{(1-t)^4} - 12 \ln(1-t) \right], \\
 \mathcal{K}^{0*,\mathcal{H}}(t,1) = & \frac{t^4(3+6t-4t^2+t^3)}{(1-t)^4}, \\
 \mathcal{K}^{1*,j}(t,1) = & \mathcal{K}^{2*,j}(t,1) = \mathcal{K}^{3*,j}(t,1) = 0, \quad j \in \{\mathcal{N}, \mathcal{P}, \mathcal{S}, \mathcal{H}\}, \tag{67}
 \end{aligned}$$

(b) for $u = -1$ and $t < 1$:

$$\begin{aligned}
 \mathcal{K}^{0*,\mathcal{N}}(t,-1) = & \frac{3t^4}{(1+t)^4}, \quad \mathcal{K}^{0*,\mathcal{P}}(t,-1) = \frac{3t^4(1-7t)}{(1+t)^5}, \\
 \mathcal{K}^{0*,\mathcal{S}}(t,-1) = & 3t^5 \left[\frac{t(38+88t+73t^2+21t^3)}{(1+t)^4} + 12 \ln(1+t) \right], \\
 \mathcal{K}^{0*,\mathcal{H}}(t,-1) = & \frac{t^4(3-6t-4t^2-t^3)}{(1+t)^4}, \\
 \mathcal{K}^{1*,j}(t,-1) = & \mathcal{K}^{2*,j}(t,-1) = \mathcal{K}^{3*,j}(t,-1) = 0, \\
 & j \in \{\mathcal{N}, \mathcal{P}, \mathcal{S}, \mathcal{H}\}, \tag{68}
 \end{aligned}$$

(c) for $u = 1$ and $t = 1$:

$$\begin{aligned}
 \mathcal{K}^{1*,\mathcal{N}}(1,1) = & \mathcal{K}^{3*,\mathcal{N}}(1,1) = \mathcal{K}^{0*,\mathcal{P}}(1,1) = \mathcal{K}^{1*,\mathcal{P}}(1,1) \\
 = & \mathcal{K}^{2*,\mathcal{P}}(1,1) = \mathcal{K}^{1*,\mathcal{S}}(1,1) = \mathcal{K}^{3*,\mathcal{S}}(1,1) \\
 = & \mathcal{K}^{1*,\mathcal{H}}(1,1) = \mathcal{K}^{3*,\mathcal{H}}(1,1) = \infty, \\
 \mathcal{K}^{0*,\mathcal{N}}(1,1) = & \mathcal{K}^{2*,\mathcal{N}}(1,1) = \mathcal{K}^{0*,\mathcal{S}}(1,1) \\
 = & \mathcal{K}^{2*,\mathcal{S}}(1,1) = \mathcal{K}^{0*,\mathcal{H}}(1,1) = \mathcal{K}^{2*,\mathcal{H}}(1,1) = -\infty, \\
 \mathcal{K}^{3*,\mathcal{P}}(1,1) = & 0, \tag{69}
 \end{aligned}$$

(d) for $u = -1$ and $t = 1$:

$$\begin{aligned}
 \mathcal{K}^{0*,\mathcal{N}}(1,-1) = & \frac{3}{16}, \quad \mathcal{K}^{0*,\mathcal{P}}(1,-1) = -\frac{9}{16}, \\
 \mathcal{K}^{0*,\mathcal{S}}(1,-1) = & \frac{165}{4} + 36 \ln 2, \quad \mathcal{K}^{0*,\mathcal{H}}(1,-1) = -\frac{1}{2}, \\
 \mathcal{K}^{1*,j}(1,-1) = & \mathcal{K}^{2*,j}(1,-1) = \mathcal{K}^{3*,j}(1,-1) = 0, \\
 & j \in \{\mathcal{N}, \mathcal{P}, \mathcal{S}, \mathcal{H}\}. \tag{70}
 \end{aligned}$$

Proposition 7 may be proven by finding the limits directly or by making use of the L'Hospital rule. Equations (67)–(70) reveal that the isotropic kernels are bounded for $t < 1$. On the other hand, the isotropic kernels are unbounded for $t = 1$ except for $\mathcal{K}^{3*,\mathcal{P}}$.

5 Investigation of the sub-integral kernels

To get an insight into the integral transformations of Proposition 4, we have further investigated properties of the

sub-integral kernels $\mathcal{K}^{\mu\nu\tau,j}$; $\mu, \nu, \tau \in \{x, y, z\}$; $j \in \{\mathcal{N}, \mathcal{P}, \mathcal{S}, \mathcal{H}\}$. The sub-integral kernels have been evaluated for two different scenarios: (1) assuming $R = 6,371$ km and $r = 6,621$ km, and (2) $R = r = 6,371$ km. In the first case, the constant height of 250 km above the mean geocentric sphere approximating the geoid is considered. Thus, it may correspond to altitude of an Earth-orbiting satellite that would eventually observe the third-order gravitational gradients, e.g. such as the proposed mission OPTIMA (Brieden et al. 2010). The second case corresponds to a terrestrial system for observing the third-order gravitational gradients, e.g.

Fig. 3 Behaviour of the sub-integral kernels: **a** $\mathcal{K}^{xxx,\mathcal{S}}$, **b** $\mathcal{K}^{xxz,\mathcal{S}}$, **c** $\mathcal{K}^{yyy,\mathcal{S}}$, **d** $\mathcal{K}^{xyz,\mathcal{S}}$, **e** $\mathcal{K}^{zzz,\mathcal{S}}$, **f** $\mathcal{K}^{zzz,\mathcal{S}}$ in the spatial domain $\alpha \in [0^\circ, 360^\circ]$, $\psi \in [0^\circ, 7^\circ]$. The kernels were evaluated for $R = 6,371$ km and $r = 6,621$ km. Values of the direct azimuth α are indicated along the circumference of each of the polar plots. The spherical distance ψ is measured from the centre of the polar plots (indicated by the dashed concentric circles). Black lines inside the plots indicate zero crossings of the sub-integral kernels

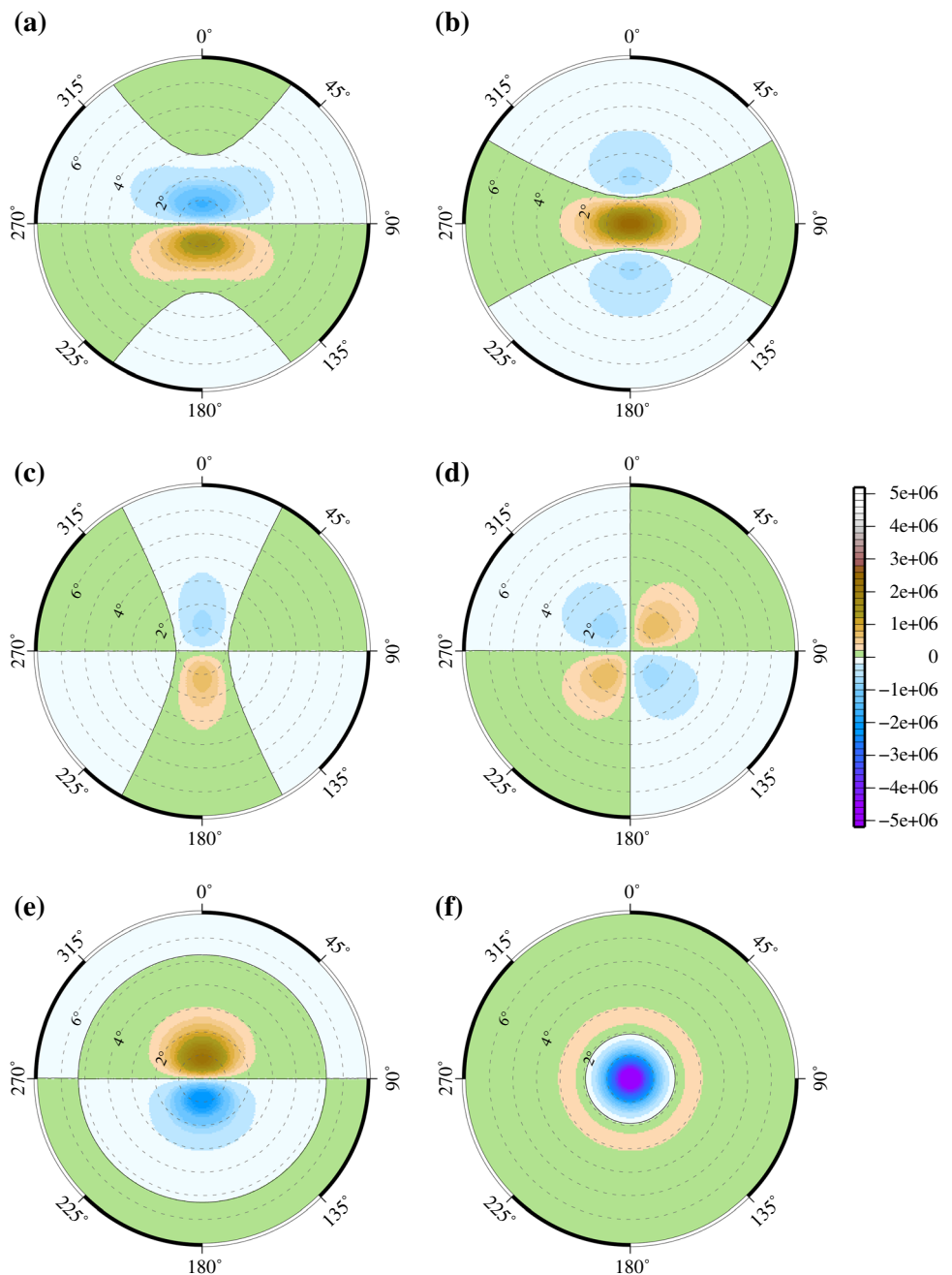
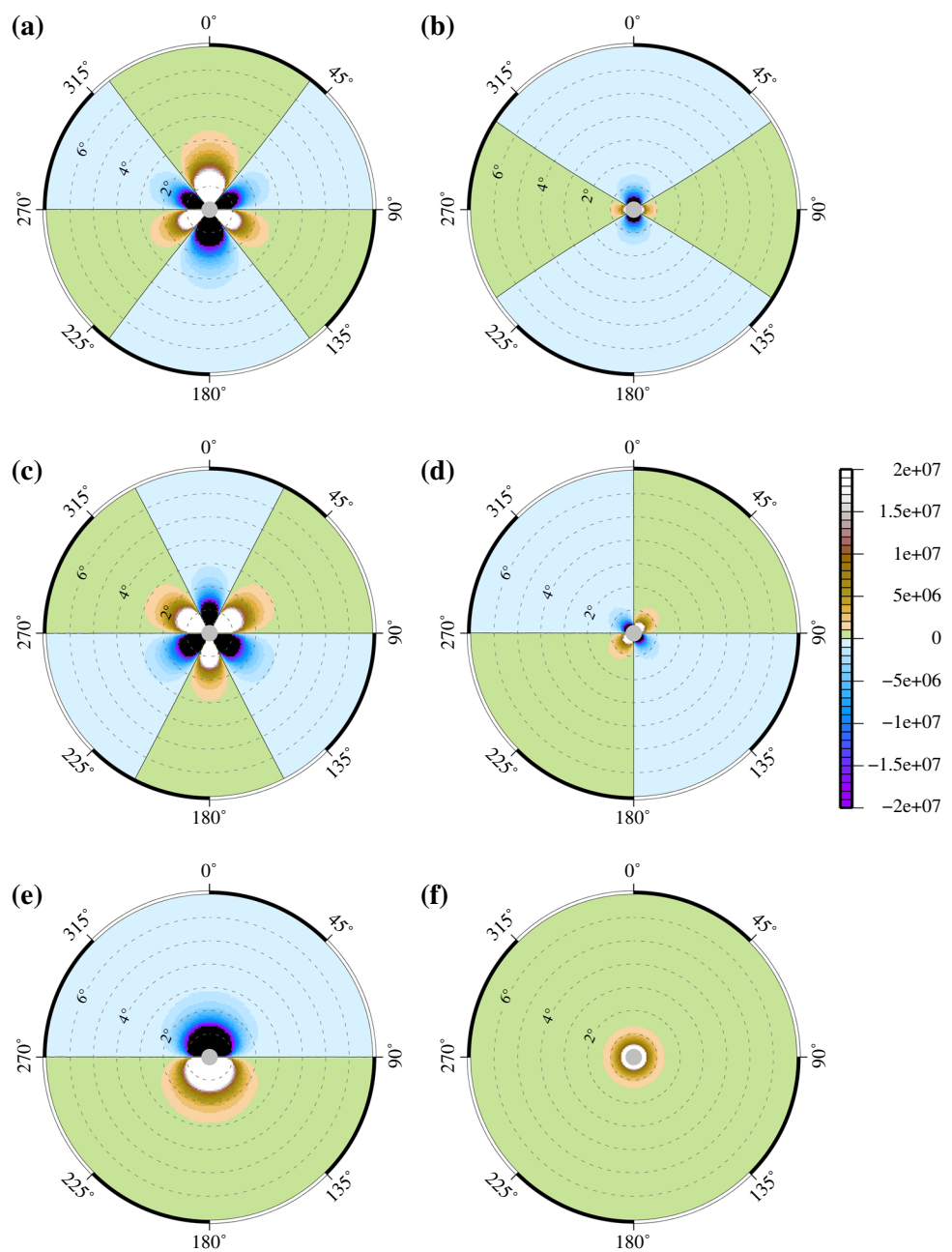


Fig. 4 Behaviour of the sub-integral kernels: **a** $\mathcal{K}^{xxx,S}$, **b** $\mathcal{K}^{xxz,S}$, **c** $\mathcal{K}^{xyy,S}$, **d** $\mathcal{K}^{xyz,S}$, **e** $\mathcal{K}^{xzz,S}$, **f** $\mathcal{K}^{zzz,S}$ in the spatial domain $\alpha \in [0^\circ, 360^\circ]$, $\psi \in [0.3^\circ, 7^\circ]$. The kernels were evaluated for $R = r = 6,371$ km. Values of the direct azimuth α are indicated along the circumference of each of the polar plots. The spherical distance ψ is measured from the centre of the polar plots (indicated by the *dashed concentric circles*). *Black lines* inside the plots indicate zero crossings of the sub-integral kernels



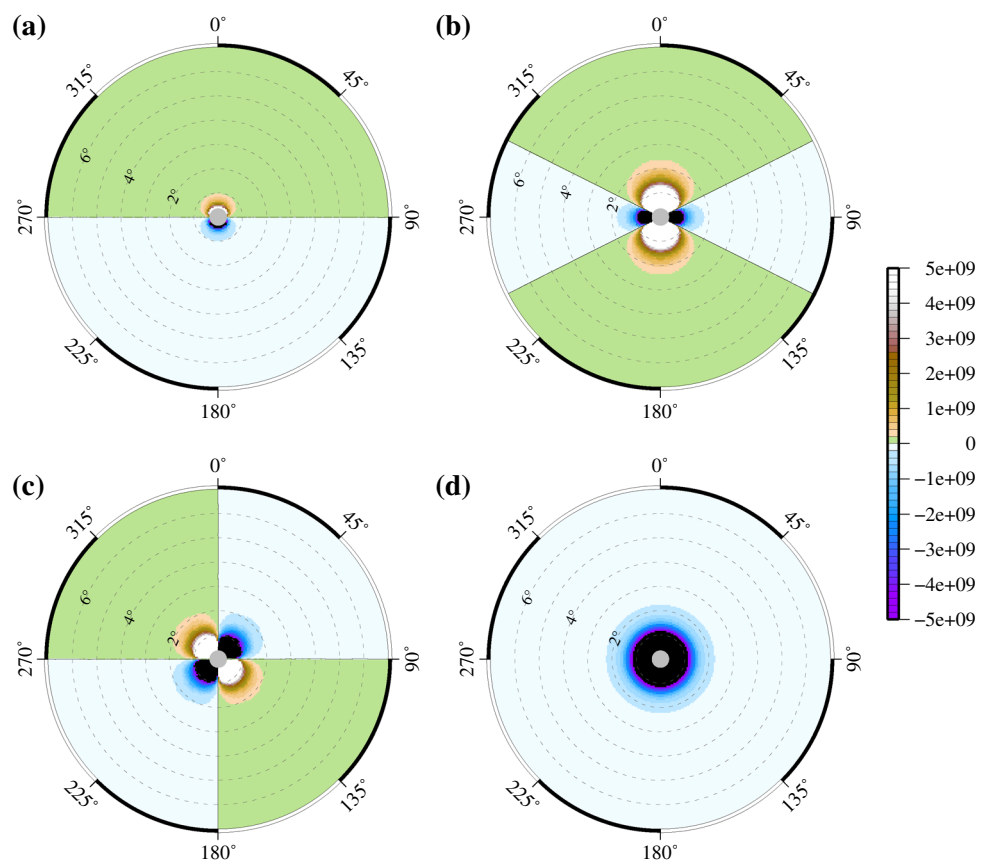
as proposed by the Dulkyn project (<http://www.dulkyn.ru/>, Balakin et al. 1997).

Figure 3 illustrates the sub-integral kernels $\mathcal{K}^{xxx,S}$, $\mathcal{K}^{xxz,S}$, $\mathcal{K}^{xyy,S}$, $\mathcal{K}^{xyz,S}$, $\mathcal{K}^{xzz,S}$ and $\mathcal{K}^{zzz,S}$ for the first scenario. The kernels are depicted as functions of the direct azimuth α (indicated by the values along the circumference of each plot) and spherical distance ψ (indicated by the dashed concentric circles). Figures are restricted to the limited spherical domain bounded by $\alpha \in [0^\circ, 360^\circ]$ and $\psi \in [0^\circ, 7^\circ]$.

Among all 40 sub-integral kernels, only six of them are sufficient to demonstrate their spatial properties due to two reasons: First, the sub-integral kernels $\mathcal{K}^{\mu\nu\tau,j}$; $\mu, \nu, \tau \in$

$\{x, y, z\}$ behave similarly when mutually comparing between each $j \in \{\mathcal{N}, \mathcal{P}, \mathcal{S}, \mathcal{H}\}$. One could also observe that magnitudes of the sub-integral kernels for $j \in \{\mathcal{N}, \mathcal{S}, \mathcal{H}\}$ are of the same order. On the other hand, these are two orders lower as compared to the sub-integral kernels for $j = \mathcal{P}$. Among the four values of j , we have opted to further investigate those for $j = \mathcal{S}$. Second, for a particular value of j , the sub-integral kernels possess azimuthal symmetries as mentioned in Sect. 3 below Proposition 2. It holds that the pairs of sub-integral kernels $(\mathcal{K}^{xxx,j}, \mathcal{K}^{yyy,j})$, $(\mathcal{K}^{xyy,j}, \mathcal{K}^{xxy,j})$ and $(\mathcal{K}^{xzz,j}, \mathcal{K}^{yzz,j})$ are interrelated by the azimuthal rotation of $\pi/2$. For example, one has to rotate the polar plot of the sub-integral kernel

Fig. 5 Behaviour of the sub-integral kernels: **a** $\mathcal{K}^{xxx,\mathcal{P}}$, **b** $\mathcal{K}^{xxz,\mathcal{P}}$, **c** $\mathcal{K}^{xyz,\mathcal{P}}$, **d** $\mathcal{K}^{zzz,\mathcal{P}}$ in the spatial domain $\alpha \in [0^\circ, 360^\circ]$, $\psi \in [0.3^\circ, 7^\circ]$. The kernels were evaluated for $R = r = 6,371$ km. Values of the direct azimuth α are indicated along the circumference of each of the polar plots. The spherical distance ψ is measured from the centre of the polar plots (indicated by the dashed concentric circles). Black lines inside the plots indicate zero crossings of the sub-integral kernels



$\mathcal{K}^{xxx,j}$ by $\pi/2$ counterclockwise to obtain $\mathcal{K}^{yyy,j}$. Moreover, the pair of sub-integral kernels $(\mathcal{K}^{xxz,j}, \mathcal{K}^{yyz,j})$ is azimuthally symmetric by rotating α by $\pm\pi/2$.

Figure 3 reveals several interesting properties of the sub-integral kernels. It can be seen that the kernel $\mathcal{K}^{zzz,\mathcal{S}}$ is isotropic, while the others possess corresponding azimuthal dependencies. Obviously, the kernels reach maxima or minima in a close neighbourhood of the evaluation point (located in the centre of the polar plots) and decay with the increasing spherical distance. One may also observe that the magnitude of the sub-integral kernels $\mathcal{K}^{xxx,\mathcal{S}}$, $\mathcal{K}^{xyy,\mathcal{S}}$, $\mathcal{K}^{xyz,\mathcal{S}}$ and $\mathcal{K}^{xzz,\mathcal{S}}$ equals zero at the evaluation point. On the other hand, the sub-integral kernels $\mathcal{K}^{xxz,\mathcal{S}}$ and $\mathcal{K}^{zzz,\mathcal{S}}$ reach their extreme values at the evaluation point.

On the surface of the mean geocentric sphere, i.e. for the second scenario with $R = r = 6,371$ km, we have to distinguish between two different cases of the sub-integral kernels. The kernels for $j \in \{\mathcal{N}, \mathcal{S}, \mathcal{H}\}$ degenerate into those illustrated by Fig. 4. Note that only the sub-integral kernels for $j = \mathcal{S}$ are depicted. Moreover, we restrict to the spatial domain bounded by $\alpha \in [0^\circ, 360^\circ]$ and $\psi \in [0.3^\circ, 7^\circ]$ due to the singularity at the evaluation point. For handling the singular behaviour in practical calculations see, e.g. (Klees 1996; Klees and Lehmann 1998). The omitted domain, for which it holds $\alpha \in [0^\circ, 360^\circ]$ and $\psi \in [0^\circ, 0.3^\circ]$, is indi-

cated by the grey dots located in the centre of each polar plot. We also note that the magnitude of the sub-integral kernels exceeds the maximum and minimum values of the colour bar scale. Therefore, one may see white and black colours in the polar plots revealing strong increase and decrease of the sub-integral kernels. Similar to the first scenario depicted in Fig. 3, it is obvious from Fig. 4 that only the kernel $\mathcal{K}^{zzz,\mathcal{S}}$ is isotropic, while the other kernels have also azimuthal dependencies. The kernels $\mathcal{K}^{xxx,\mathcal{S}}$, $\mathcal{K}^{xyy,\mathcal{S}}$ and $\mathcal{K}^{xzz,\mathcal{S}}$ reach higher magnitudes for larger values of the spherical distance as compared to $\mathcal{K}^{xxz,\mathcal{S}}$, $\mathcal{K}^{xyz,\mathcal{S}}$ and $\mathcal{K}^{zzz,\mathcal{S}}$. In particular, this is seen for $\psi \in [1^\circ, 3^\circ]$.

The second case of the sub-integral kernels is represented for $j = \mathcal{P}$. On the mean geocentric sphere, the isotropic kernel $\mathcal{K}^{3*,\mathcal{P}} = 0$ for $-1 \leq u \leq 1$ that may be proven by inserting $t = 1$ into Eq. (58). Considering this in Eq. (45) of Proposition 4, we get $\mathcal{K}^{xxx,\mathcal{P}} = 3 \mathcal{K}^{xyy,\mathcal{P}} = -3/4 \mathcal{K}^{xzz,\mathcal{P}}$ revealing another symmetry of the sub-integral kernels. Consequently, only four sub-integral kernels, which are depicted in Fig. 5, are sufficient to illustrate their complete spatial behaviour for $j = \mathcal{P}$. In correspondence with Figs. 3 and 4, we observe in Fig. 5 the isotropic behaviour of the kernel $\mathcal{K}^{zzz,\mathcal{P}}$ and the azimuthal dependence for $\mathcal{K}^{xxx,\mathcal{P}}$, $\mathcal{K}^{xxz,\mathcal{P}}$ and $\mathcal{K}^{xyz,\mathcal{P}}$. The kernels $\mathcal{K}^{xxz,\mathcal{P}}$, $\mathcal{K}^{xyz,\mathcal{P}}$ and $\mathcal{K}^{zzz,\mathcal{P}}$ are more significant compared to $\mathcal{K}^{xxx,\mathcal{S}}$, especially for $\psi \in [1^\circ, 2^\circ]$.

6 Conclusions

A new mathematical model for evaluation of the third-order (disturbing) gravitational gradients is formulated in this article. Required differential operators in the spherical local north-oriented frame were derived first, see Sect. 3. The operators were subsequently applied to integral equations of Newton, Abel-Poisson, Pizzetti and Hotine that link the volumetric mass density, disturbing gravitational potential, gravity anomaly and gravity disturbance with the (disturbing) gravitational potential. The new integral equations can then be used for transformations between third-order (disturbing) gravitational gradients and the volumetric mass density, disturbing gravitational potential, gravity anomaly and gravity disturbance, respectively; see Proposition 4. Spatial properties of the new integral kernels were then studied in Sect. 5, where several plots illustrating their characteristic features and properties are shown.

The new mathematical formulas extend the theoretical apparatus of geodesy, i.e. the well-known Meissl scheme of physical geodesy, and reveal important properties of the third-order gravitational tensor. They may be exploited in future geophysical studies, continuation of gravitational field quantities and analysing the gradiometric-geodynamic boundary value problem.

Acknowledgments Michal Šprlák was supported by the Project EXLIZ-CZ.1.07/2.3.00/30.0013, which is co-financed by the European Social Fund and the state budget of the Czech Republic. Pavel Novák was supported by the Project 209/12/1929 of the Czech Science Foundation. Thoughtful and constructive comments of Dr. Stefano Casotto and two anonymous reviewers are gratefully acknowledged. Thanks are also extended to the editor-in-chief Prof. Roland Klees and the responsible editor Prof. Christopher Jekeli for handling our manuscript.

Appendix A: Differential operators for the second-order gravitational gradients

In Sect. 3, we construct the differential operators for the third-order (disturbing) gravitational gradients. For this purpose, we exploit also the differential operators for the second-order gravitational gradients. The six second-order differential operators in terms of angular spherical geocentric coordinates (r, Ω) read as follows (Reed 1973; Koop 1993):

$$\mathcal{D}^{xx} = \frac{1}{r} \left(\frac{\partial}{\partial r} + \frac{1}{r} \frac{\partial^2}{\partial \varphi^2} \right), \tag{71}$$

$$\mathcal{D}^{xy} = -\frac{1}{r^2 \cos \varphi} \left(\tan \varphi \frac{\partial}{\partial \lambda} + \frac{\partial^2}{\partial \varphi \partial \lambda} \right), \tag{72}$$

$$\mathcal{D}^{xz} = -\frac{1}{r} \left(\frac{1}{r} \frac{\partial}{\partial \varphi} - \frac{\partial^2}{\partial r \partial \varphi} \right), \tag{73}$$

$$\mathcal{D}^{yy} = \frac{1}{r} \left(\frac{\partial}{\partial r} - \frac{\tan \varphi}{r} \frac{\partial}{\partial \varphi} + \frac{1}{r \cos^2 \varphi} \frac{\partial^2}{\partial \lambda^2} \right), \tag{74}$$

$$\mathcal{D}^{yz} = \frac{1}{r \cos \varphi} \left(\frac{1}{r} \frac{\partial}{\partial \lambda} - \frac{\partial^2}{\partial r \partial \lambda} \right), \tag{75}$$

$$\mathcal{D}^{zz} = \frac{\partial^2}{\partial r^2}. \tag{76}$$

Alternatively, the second-order differential operators of Eqs. (71)–(76) may be expressed in terms of the spherical polar coordinates (r, ψ, α) , see, e.g. (Wolf 2007; Šprlák et al. 2014):

$$\begin{aligned} \mathcal{D}^{xx} &= \mathcal{D}_2^1 + \cos 2\alpha \mathcal{D}_2^2, & \mathcal{D}^{xy} &= -\sin 2\alpha \mathcal{D}_2^2, \\ \mathcal{D}^{xz} &= \cos \alpha \mathcal{D}_2^3, & \mathcal{D}^{yy} &= \mathcal{D}_2^1 - \cos 2\alpha \mathcal{D}_2^2, \\ \mathcal{D}^{yz} &= -\sin \alpha \mathcal{D}_2^3, & \mathcal{D}^{zz} &= \mathcal{D}_2^4, \end{aligned} \tag{77}$$

where

$$\mathcal{D}_2^1 = \frac{1}{r} \left[\frac{\partial}{\partial r} + \frac{1}{2r} \left(\frac{\partial^2}{\partial \psi^2} + \cot \psi \frac{\partial}{\partial \psi} \right) \right], \tag{78}$$

$$\mathcal{D}_2^2 = \frac{1}{2r^2} \left(\frac{\partial^2}{\partial \psi^2} - \cot \psi \frac{\partial}{\partial \psi} \right), \tag{79}$$

$$\mathcal{D}_2^3 = \frac{1}{r} \left(\frac{1}{r} \frac{\partial}{\partial \psi} - \frac{\partial^2}{\partial r \partial \psi} \right), \tag{80}$$

$$\mathcal{D}_2^4 = \frac{\partial^2}{\partial r^2}. \tag{81}$$

Note that the subscript in Eqs. (78)–(81) refers to the order of the gravitational gradients.

Appendix B: Explicit decomposition of degree-dependent terms

To arrive at the new isotropic kernels $\mathcal{K}^{i^*,j}$; $i \in \{0, 1, 2, 3\}$; $j \in \{\mathcal{N}, \mathcal{P}, \mathcal{S}, \mathcal{H}\}$ in the closed form, see Sect. 4, an explicit decomposition of some degree-dependent terms is required. Such a decomposition is completely defined by the following equations:

$$h_n^{\mathcal{N}} (n+1)(n+2)(n+3) = 6 + 11n + 6n^2 + n^3, \tag{82}$$

$$h_n^{\mathcal{N}} (n+2)(n+3) = 6 + 5n + n^2, \tag{83}$$

$$h_n^{\mathcal{N}} (n+3) = 3 + n, \tag{84}$$

$$h_n^{\mathcal{N}} = 1, \tag{85}$$

$$h_n^{\mathcal{P}} (n+1)(n+2)(n+3) = 6 + 23n + 28n^2 + 13n^3 + 2n^4, \tag{86}$$

$$h_n^{\mathcal{P}} (n+2)(n+3) = 6 + 17n + 11n^2 + 2n^3, \tag{87}$$

$$h_n^{\mathcal{P}} (n+3) = 3 + 7n + 2n^2, \tag{88}$$

$$h_n^{\mathcal{P}} = 1 + 2n, \tag{89}$$

$$\begin{aligned} h_n^{\mathcal{S}} (n+1)(n+2)(n+3) \\ = \frac{72}{n-1} + 66 + 43n + 15n^2 + 2n^3, \end{aligned} \tag{90}$$

$$h_n^{\mathcal{S}} (n+2)(n+3) = \frac{36}{n-1} + 30 + 13n + 2n^2, \tag{91}$$

$$h_n^{\mathcal{S}} (n+3) = \frac{12}{n-1} + 9 + 2n, \tag{92}$$

$$h_n^S = \frac{3}{n-1} + 2, \tag{93}$$

$$h_n^{\mathcal{H}}(n+1)(n+2)(n+3) = 6 + 17n + 11n^2 + 2n^3, \tag{94}$$

$$h_n^{\mathcal{H}}(n+2)(n+3) = -\frac{2}{n+1} + 8 + 9n + 2n^2, \tag{95}$$

$$h_n^{\mathcal{H}}(n+3) = -\frac{2}{n+1} + 5 + 2n, \tag{96}$$

$$h_n^{\mathcal{H}} = -\frac{1}{n+1} + 2. \tag{97}$$

Note that the symbols h_n^j ; $j \in \{N, \mathcal{P}, S, \mathcal{H}\}$ stand for the eigenvalues of the Newton, Abel-Poisson, Pizzetti and Hotine kernels. These are defined by Eqs. (39)–(42).

Appendix C: Sums of infinite series

To derive the new isotropic kernels $\mathcal{K}^{i*,j}$; $i \in \{0, 1, 2, 3\}$; $j \in \{N, \mathcal{P}, S, \mathcal{H}\}$ in closed form, see Sect. 4, we also need closed form expressions for some infinite series. Complete set of these summation rules reads as follows:

$$S_1(t, u) = \sum_{n=0}^{\infty} t^{n+4} \frac{1}{n+1} P_{n,0}(u) = t^3 \ln\left(\frac{g+t-u}{1-u}\right), \tag{98}$$

$$S_2(t, u) = \sum_{n=0}^{\infty} t^{n+4} P_{n,0}(u) = \frac{t^4}{g}, \tag{99}$$

$$S_3(t, u) = \sum_{n=0}^{\infty} t^{n+4} n P_{n,0}(u) = \frac{t^5(u-t)}{g^3}, \tag{100}$$

$$S_4(t, u) = \sum_{n=0}^{\infty} t^{n+4} n^2 P_{n,0}(u) = \frac{t^5}{g^3} \left[u - 2t + \frac{3t(t-u)^2}{g^2} \right], \tag{101}$$

$$S_5(t, u) = \sum_{n=0}^{\infty} t^{n+4} n^3 P_{n,0}(u) = \frac{t^5}{g^3} \left[u - 4t + \frac{9t(t-u)(2t-u)}{g^2} - \frac{15t^2(t-u)^3}{g^4} \right], \tag{102}$$

$$S_6(t, u) = \sum_{n=0}^{\infty} t^{n+4} n^4 P_{n,0}(u) = \frac{t^5}{g^3} \left[u - 8t + \frac{12t(t-u)(7t-u) + 9tu^2}{g^2} - \frac{90t^2(t-u)^2(2t-u)}{g^4} + \frac{105t^3(t-u)^4}{g^6} \right], \tag{103}$$

$$S_7(t, u) = \sum_{n=1}^{\infty} t^{n+4} \frac{1}{n+1} P_{n,1}(u) = t^3 \sqrt{1-u^2} \left[\frac{1}{1-u} - \frac{g+t}{g(g+t-u)} \right], \tag{104}$$

$$S_8(t, u) = \sum_{n=1}^{\infty} t^{n+4} P_{n,1}(u) = \frac{t^5 \sqrt{1-u^2}}{g^3}, \tag{105}$$

$$S_9(t, u) = \sum_{n=1}^{\infty} t^{n+4} n P_{n,1}(u) = \frac{t^5 \sqrt{1-u^2}}{g^3} \left[1 + \frac{3t(u-t)}{g^2} \right], \tag{106}$$

$$S_{10}(t, u) = \sum_{n=1}^{\infty} t^{n+4} n^2 P_{n,1}(u) = \frac{t^5 \sqrt{1-u^2}}{g^3} \left[1 + \frac{3t(3u-4t)}{g^2} + \frac{15t^2(t-u)^2}{g^4} \right], \tag{107}$$

$$S_{11}(t, u) = \sum_{n=1}^{\infty} t^{n+4} n^3 P_{n,1}(u) = \frac{t^5 \sqrt{1-u^2}}{g^3} \left[1 + \frac{3t(7u-13t)}{g^2} + \frac{45t^2(t-u)(3t-2u)}{g^4} - \frac{105t^3(t-u)^3}{g^6} \right], \tag{108}$$

$$S_{12}(t, u) = \sum_{n=2}^{\infty} t^{n+4} \frac{1}{n+1} P_{n,2}(u) = t^3(1-u^2) \left[\frac{1}{(1-u)^2} - \frac{t^2}{g^3(g+t-u)} - \frac{(g+t)^2}{g^2(g+t-u)^2} \right], \tag{109}$$

$$S_{13}(t, u) = \sum_{n=2}^{\infty} t^{n+4} P_{n,2}(u) = \frac{3t^6(1-u^2)}{g^5}, \tag{110}$$

$$S_{14}(t, u) = \sum_{n=2}^{\infty} t^{n+4} n P_{n,2}(u) = \frac{3t^6(1-u^2)}{g^5} \left[2 + \frac{5t(u-t)}{g^2} \right], \tag{111}$$

$$S_{15}(t, u) = \sum_{n=2}^{\infty} t^{n+4} n^2 P_{n,2}(u) = \frac{3t^6(1-u^2)}{g^5} \left[4 + \frac{5t(5u-6t)}{g^2} + \frac{35t^2(t-u)^2}{g^4} \right], \tag{112}$$

$$S_{16}(t, u) = \sum_{n=3}^{\infty} t^{n+4} \frac{1}{n+1} P_{n,3}(u) = t^3 \sqrt{(1-u^2)^3} \left[\frac{2}{(1-u)^3} - \frac{3t^3}{g^5(g+t-u)} - \frac{3t^2(g+t)}{g^4(g+t-u)^2} - \frac{2(g+t)^3}{g^3(g+t-u)^3} \right], \tag{113}$$

$$S_{17}(t, u) = \sum_{n=3}^{\infty} t^{n+4} P_{n,3}(u) = \frac{15t^7 \sqrt{(1-u^2)^3}}{g^7}, \tag{114}$$

$$S_{18}(t, u) = \sum_{n=3}^{\infty} t^{n+4} n P_{n,3}(u) = \frac{15t^7 \sqrt{(1-u^2)^3}}{g^7} \left[3 + \frac{7t(u-t)}{g^2} \right], \quad (115)$$

$$S_{19}(t, u) = \sum_{n=2}^{\infty} t^{n+4} \frac{1}{n-1} P_{n,0}(u) = t^4 \left[1 - g - tu - tu \ln \left(\frac{1+g-tu}{2} \right) \right], \quad (116)$$

$$S_{20}(t, u) = \sum_{n=2}^{\infty} t^{n+4} P_{n,0}(u) = S_2(t, u) - t^4(1+tu), \quad (117)$$

$$S_{21}(t, u) = \sum_{n=2}^{\infty} t^{n+4} n P_{n,0}(u) = S_3(t, u) - t^5 u, \quad (118)$$

$$S_{22}(t, u) = \sum_{n=2}^{\infty} t^{n+4} n^2 P_{n,0}(u) = S_4(t, u) - t^5 u, \quad (119)$$

$$S_{23}(t, u) = \sum_{n=2}^{\infty} t^{n+4} n^3 P_{n,0}(u) = S_5(t, u) - t^5 u, \quad (120)$$

$$S_{24}(t, u) = \sum_{n=2}^{\infty} t^{n+4} \frac{1}{n-1} P_{n,1}(u) = t^5 \sqrt{1-u^2} \left[\frac{(g+1)^2}{g(1+g-tu)} - 2 - \ln \left(\frac{1+g-tu}{2} \right) \right], \quad (121)$$

$$S_{25}(t, u) = \sum_{n=2}^{\infty} t^{n+4} P_{n,1}(u) = S_8(t, u) - t^5 \sqrt{1-u^2}, \quad (122)$$

$$S_{26}(t, u) = \sum_{n=2}^{\infty} t^{n+4} n P_{n,1}(u) = S_9(t, u) - t^5 \sqrt{1-u^2}, \quad (123)$$

$$S_{27}(t, u) = \sum_{n=2}^{\infty} t^{n+4} n^2 P_{n,1}(u) = S_{10}(t, u) - t^5 \sqrt{1-u^2}, \quad (124)$$

$$S_{28}(t, u) = \sum_{n=2}^{\infty} t^{n+4} \frac{1}{n-1} P_{n,2}(u) = \frac{t^6(1-u^2)}{1+g-tu} \left[1 + \frac{1}{g^3} + \frac{(g+1)^3}{g^2(1+g-tu)} \right], \quad (125)$$

$$S_{29}(t, u) = \sum_{n=3}^{\infty} t^{n+4} \frac{1}{n-1} P_{n,3}(u) = \frac{t^7 \sqrt{(1-u^2)^3}}{1+g-tu} \left[\frac{3}{g^5} + \frac{1}{1+g-tu} + \frac{4g+3}{g^4(1+g-tu)} + \frac{2(g+1)^4}{g^3(1+g-tu)^2} \right]. \quad (126)$$

Note that some of the summation rules may be found, e.g. in (Pick et al. 1973; Moritz 1980; Martinec 2003; Šprlák and Novák 2014).

References

- Abdelrahman E-SM, El-Araby HM, El-Araby TM, Abo-Ezz ER (2003) A least-squares derivatives analysis of gravity anomalies due to faulted thin slabs. *Geophysics* 68:535–543
- Albertella A, Migliaccio F, Sansó F, Tschering CC (2000) The space-wise approach—overall scientific data strategy. In: Sünkel H (ed) From Eötvös to Milligal, Final Report, ESA Study, ESA/ESTEC Contract No 13329/98/NL/GD, pp 267–298
- Ardalan AA, Grafarend EW (2001) Ellipsoidal geoidal undulations (ellipsoidal Bruns formula). *J Geod* 75:544–552
- Arfken G (1968) *Mathematical methods for physicists*. Academic Press, New York
- Balakin AB, Daishev RA, Murzakhanov ZG, Skochilov AF (1997) Laser-interferometric detector of the first, second and third derivatives of the potential of the Earth gravitational field. *Izvestiya vysshikh uchebnykh zavedenii, seriya Geologiya i Razvedka* 1:101–107
- Beiki M (2010) Analytic signals of gravity gradient tensor and their application to estimate source location. *Geophysics* 75:159–174
- Bölling C, Grafarend EW (2005) Ellipsoidal spectral properties of the Earth's gravitational potential and its first and second derivatives. *J Geod* 79:300–330
- Brieden P, Müller J, Flury J, Heinzel G (2010) The mission OPTIMA—novelties and benefit. *Geotechnologien, Science Report*, No 17, Potsdam, Germany, pp 134–139
- Casotto S, Fantino E (2009) Gravitational gradients by tensor analysis with application to spherical coordinates. *J Geod* 83:621–634
- Chelton DB, Ries JC, Haines BJ, Fu L-L, Callahan PS (2001) Satellite altimetry. In: Fu L-L, Cazenave A (eds) *Satellite altimetry and earth sciences: a handbook of techniques and applications*. International geophysics series, vol 69. Academic Press, San Diego, pp 1–131
- Csapó G, Égető C, Laky S, Tóth G, Ulmann Z, Völgyesi L (2009) Test measurements by Eötvös torsion balance and gravimeters. *Civ Eng* 53(2):75–80
- Cunningham L (1970) On the computation of the spherical harmonic terms needed during the numerical integration of the orbital motion of an artificial satellite. *Celest Mech* 2:207–216
- Eppelbaum LV (2011) Review of environmental and geological micro-gravity applications and feasibility of its employment at archaeological sites in Israel. *Int J Geophys* 2011:1–9
- ESA (1999) Gravity field and steady-state ocean circulation mission. ESA SP-1233(1), Report for Mission Selection of the Four Candidate Earth Explorer Missions, ESA Publication Division
- Fantino E, Casotto S (2009) Methods of harmonic synthesis for global geopotential models and their first-, second- and third-order gradients. *J Geod* 83:595–619
- Fedi M, Florio G (2001) Detection of potential field source boundaries by enhanced horizontal derivative method. *Geophys Prospect* 49:40–58
- Forsberg R, Olesen AV (2010) Airborne gravity field determination. In: Xu G (ed) *Sciences of geodesy-I: advances and future directions*. Springer, Berlin, pp 83–104
- Freedon W, Schreiner M (2009) Spherical functions of mathematical geosciences. A scalar, vectorial, and tensorial setup. *Advances in geophysical and environmental mechanics and mathematics*. Springer, Berlin
- Fukushima T (2012) Numerical computation of spherical harmonics of arbitrary degree and order by extending exponent of floating

- point numbers: II first-, second-, and third-order derivatives. *J Geod* 86:1019–1028
- Fukushima T (2013) Recursive computation of oblate spheroidal harmonics of the second kind and their first-, second-, and third-order derivatives. *J Geod* 87:303–309
- Grafarend EW (1997) Field lines of gravity, their curvature and torsion, the Lagrangian and the Hamilton equations of the plumbline. *Annali di Geofisica* 40:1233–1247
- Grafarend EW (2001) The spherical horizontal and spherical vertical boundary value problem—vertical deflections and geoid undulations—the completed Meissl diagram. *J Geod* 75:363–390
- Hafez M, El-Qady G, Awad S, Elsayed E-S (2006) Higher derivative analysis for the interpretation of self-potential profiles at southern part of Nile delta, Egypt. *Arab J Sci Eng* 31:129–145
- Heiskanen WA, Moritz H (1967) *Physical geodesy*. Freeman and Co., San Francisco
- Hotine M (1969) *Mathematical geodesy*. ESSA Monograph No 2, US Department of Commerce, Washington, DC, USA
- Jacoby W, Smilde PL (2009) *Gravity interpretation: fundamentals and application of gravity inversion*. Springer, Berlin
- Jordan MJT, Thompson KC, Collins MA (1995) The utility of higher order derivatives in constructing molecular potential energy surfaces by interpolation. *J Chem Phys* 103:9669–9675
- Keller W, Sharifi MA (2005) Satellite gradiometry using a satellite pair. *J Geod* 78:544–557
- Kellogg OD (1929) *Foundations of potential theory*. Dover Publications, New York
- Klees R (1996) Numerical calculation of weakly singular surface integrals. *J Geod* 70:781–797
- Klees R, Lehmann R (1998) Calculation of strongly singular and hyper-singular surface integrals. *J Geod* 72:530–546
- Klokočník J, Kostelecký J, Pešek I, Novák P, Wagner CA, Sebera J (2010) Candidates for multiple impact craters: popigai and chixulub as seen by EGM08, a global 5' × 5' gravitational model. *Solid Earth Discuss* 2:69–103
- Koop R (1993) *Global gravity field modelling using satellite gravity gradiometry*. Publications on Geodesy, Netherlands Geodetic Commission, No 38, Delft, The Netherlands
- Martinec Z (1998) Boundary value problems for gravimetric determination of a precise geoid. *Lecture notes in earth sciences*, vol 73. Springer, Berlin
- Martinec Z (2003) Green's function solution to spherical gradiometric boundary-value problems. *J Geod* 77:41–49
- Marussi A (1985) *Intrinsic geodesy*. Springer, Berlin
- Meissl P (1971) A study of covariance functions related to the Earth's disturbing potential. Report No 151, Department of Geodetic Science, The Ohio State University, Columbus, Ohio, USA
- Metris G, Xu J, Wyrzyszcak (1999) Derivatives of the gravity potential with respect to rectangular coordinates. *Celest Mech Dyn Astron* 71:137–151
- Moritz H (1967) *Kinematical geodesy*. Report No 92, Department of Geodetic Science, Ohio State University, Columbus, Ohio, USA
- Moritz H (1980) *Advanced physical geodesy*. Herbert Wichmann Verlag, Karlsruhe
- Nagy D, Papp G, Benedek J (2000) The gravitational potential and its derivatives for the prism. *J Geod* 74:552–560
- Pajot G, de Viron O, Diament M, Lequentrec-Lalancette M-F, Mikhailov V (2008) Noise reduction through joint processing of gravity and gravity gradient data. *Geophysics* 73:123–134
- Petrovskaya MS, Vershkov AN (2010) Construction of spherical harmonic series for the potential derivatives of arbitrary orders in the geocentric Earth-fixed reference frame. *J Geod* 84:165–178
- Pick M, Pícha J, Vyskočil V (1973) *Theory of the Earth's gravity field*. Elsevier, Amsterdam
- Pizzetti P (1911) Sopra il calcolo teorico delle deviazioni del geoido dall' ellissoide. *Atti della Reale Accademia della Scienze di Torino* 46:331
- Reed GB (1973) Application of kinematical geodesy for determining the short wavelength components of the gravity field by satellite gradiometry. Report No 201, Department of Geodetic Science, Ohio State University, Columbus, Ohio, USA
- Rummel R (1986) Satellite gradiometry. In: Sünkel H (ed) *Mathematical and numerical techniques in physical geodesy*. Lecture notes in earth sciences, vol 7. Springer, Berlin, pp 317–363
- Rummel R (1997) Spherical spectral properties of the Earth's gravitational potential and its first and second derivatives. In: Sansó F, Rummel R (eds) *Geodetic boundary value problems in view of the one centimeter geoid*. Lecture notes in earth sciences, vol 65. Springer, Berlin, pp 359–404
- Rummel R (2010) GOCE: Gravitational gradiometry in a satellite. In: Freeden W, Nashed MZ, Sonar T (eds) *Handbook of Geomathematics*. Springer, Berlin, pp 93–103
- Rummel R, van Gelderen M (1995) Meissl scheme—spectral characteristics of physical geodesy. *Manuscripta Geodaetica* 20:379–385
- Sali WB (2009) Removing noise due to aircraft motion from gravity gradient data. MSc Thesis, University of Witwatersrand, Johannesburg, South Africa
- Schwarz KP, Li Z (1997) An introduction to airborne gravimetry and its boundary value problems. In: Sansó F, Rummel R (eds) *Geodetic boundary value problems in view of the one centimeter geoid*. Lecture notes in earth sciences, vol 65. Springer, Berlin, pp 312–358
- Seeber G (2003) *Satellite geodesy*, 2nd edn. De Gruyter, Berlin
- Smith RS, Thurston JB, Dai TF, MacLeod IN (1998) *iSPITM*—the improved source parameter imaging method. *Geophys Prospect* 46:141–151
- Sun W, Wu X, Huang G-Q (2011) Symplectic integrators with potential derivatives to third order. *Res Astron Astrophys* 11:353–368
- Šprlák M, Novák P (2014) Integral transformations of gradiometric data onto a GRACE type of observable. *J Geod* 88:377–390
- Šprlák M, Sebera J, Valko M, Novák P (2014) Spherical integral formulas for upward/downward continuation of gravitational gradients onto gravitational gradients. *J Geod* 88:179–197
- Thurston JB, Smith RS, Guillon JC (2002) A multimodel method for depth estimation from magnetic data. *Geophysics* 67:555–561
- Timmen L (2010) Absolute and relative gravimetry. In: Xu G (ed) *Sciences of geodesy-I: advances and future directions*. Springer, Berlin, pp 1–48
- Torge W (1989) *Gravimetry*. De Gruyter, Berlin
- Torge W, Müller J (2012) *Geodesy*, 4th edn. De Gruyter Inc., Berlin
- Tóth G (2005) The gradiometric-geodynamic boundary value problem. In: Jekeli C, Bastos L, Fernandes L (eds) *Gravity, geoid and space missions, IAG symposia*, vol 129. Springer, Berlin, pp 352–357
- Tóth G, Földvály L (2005) Effect of geopotential model errors on the projection of GOCE gradiometer observables. In: Jekeli C, Bastos L, Fernandes L (eds) *Gravity, geoid and space missions, IAG symposia*, vol 129. Springer, Berlin, pp 72–76
- Troshkov GA, Shalaev SV (1968) Application of the Fourier transform to the solution of the reverse problem of gravity and magnetic surveys. *Can J Explor Geophys* 4:46–62
- Velkoborský P (1982) Use of Poisson's integral in calculating higher vertical derivatives of harmonic functions—Part 1. *Studia Geophysica et Geodaetica* 26:3–11
- Veryaskin A, McRae W (2008) On the combined gravity gradient components modeling for applied geophysics. *J Geophys Eng* 5:348–356
- Wang YM (1989) Determination of the gravity disturbance by processing the gravity gradiometer surveying system—processing procedure and results. Scientific Report No 8, Department of Geodetic

- Science and Surveying, Ohio State University, Columbus, Ohio, USA
- Winch DE, Roberts PH (1995) Derivatives of addition theorem for Legendre functions. *J Aust Math Soc Ser B Appl Math* 37:212–234
- Wolf KI (2007) Kombination globaler Potentialmodelle mit terrestrische Schweredaten für die Berechnung der zweiten Ableitungen des Gravitationspotentials in Satelitenbahnhöhe. Deutsche Geodätische Kommission, Reihe C, No 603, München, Germany