Thermodynamic Assessment of the Mn-O System

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A thermodynamic assessment of the Mn-O system has been made by using thermodynamic models for the Gibbs energy of individual phases. A set of evaluated thermodynamic parameters was obtained by considering the available experimental information, and it gives a reasonable description of the system. The thermodynamic parameters of the system and comparisons between the calculation and experimental measurements, including phase diagrams, are presented.

I. INTRODUCTION

THE phase diagram of the Mn-O system was recently evaluated by Massalski et al.,^[1] mainly based on the reported information of Trömel et al.^[2] and Chen.^[3] The system exhibits a wide liquid miscibility gap on the $Mn-Mn_{1-x}O$ side and an eutectic point on the $Mn_{1-x}O-Mn_3O_4$ side.^[1-4] The manganosite ($Mn_{1-x}O$) is a nonstoichiometric compound with a deficit of Mn. The reported experimental information on the phase diagram is rather scarce and imprecise. However, many thermodynamic investigations have been reported in the literature dealing with various aspects of pure manganese and manganese oxides. The purpose of the present assessment is to obtain a set of revised thermodynamic descriptions of various phases and a thermodynamically consistent phase diagram of the Mn-O system based on the relevant experimental information by means of the CALPHAD technique.

II. EXPERIMENT INFORMATION

The available experimental information on the Mn-O system has been reviewed in the present work. Table I shows the crystal structure data of the solid phases.

A. Manganese (Mn)

The thermodynamic properties of condensed manganese have been compiled several times.^[5-10] The most recent result by Guillermet *et al.*^[10] was adopted in thepresent assessment.</sup>

B. Manganosite $(Mn_{l-x}O)$

Manganosite is a nonstoichiometric compound with the halite type of crystal structure. The oxygen activity has been measured by several authors,^[11–14] but the deviations among these results are large. Bransky and Tallan^[13] measured the excess oxygen in $Mn_{1-x}O$ in the temperature range of 1273 to 1773 K and at oxygen pressures of 10^{-14} to 1 atm by the thermogravimetric method. However, the excess oxygen that they measured is rather

small, especially at the temperatures of 1273 and 1373 K. This may be due to an incorrect assumption regarding the evaporation rate of manganosite during the experiment. Janowski et al.^[14] also studied the excess oxygen by the same method, but the calculated oxygen pressure of the gas phase was too low to keep the manganosite thermodynamically stable. Davies and Richardson^[11] studied the excess oxygen of manganosite by analyzing the composition of the samples quenched from a certain temperature and oxygen pressure. The experimental data are probably rather imprecise, because they found that oxygen pressure in equilibrium with manganosite only varies with the oxygen composition of manganosite but does not change with temperature from 1773 to 1923 K. In 1974, Picard and Gerdanian,^[12] with an improved thermogravimetric technique, measured the oxygen activity at temperatures of 1273, 1323, 1373, and 1423 K. In the present assessment, only Picard's experimental data were used in the optimization, because they gave better agreement with the experimental data of the oxygen solubility in manganosite equilibrated with Mn₃O₄ reported by Fender and Riley.^[15] Information on the Gibbs energy of formation of manganosite. using a solid electrolyte galvanic cell, obtained by Schwerdtfeger^[16] and Simeonov et al.,^[17] was accepted. The heat capacity was studied by Kleinclaus et al.^[18] in the temperature range from 20 to 700 K, and some data from 298 to 700 K were used in the optimization. The heat content measured by Southard and Shomate^[19] was adopted, and $\Delta^{\circ}H_{298}$ and $^{\circ}S_{298}$ were taken from the selected values by Kubaschewski et al.[20]

C. Hausmanite (Mn_3O_4)

Hausmanite is considered as a stoichiometric compound in the present work. Several authors^[22-29] measured the equilibrium oxygen pressure of hausmanite and manganosite, and the experimental data reported by Schaefer,^[22] Ramana Rato and Tare^[23] and Price and Wagner^[24] were used in the optimization. The experimental data of the heat capacity and °S₂₉₈ measured by Robie and Hemingway,^[30] the heat content from 763 to 1768 K studied by Southard and Moore,^[31] and $\Delta^{\circ}H_{298}$ from Kubaschewski *et al.*^[20] were also used. There are several reports^[20,23,31–35] about the values of temperature and enthalpy of a structural transformation from α -Mn₃O₄ to β -Mn₃O₄, and not all of them agree. The values of $\Delta H = 20,810$ J/mol and T = 1445 K, according to a number of measurements,^[20,23,32,33] were accepted.

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Table I. Crystal Structure Data

Phase	Pearson Symbol	Space Group	Strukturbericht Designation	Prototype	Reference
δ-Mn	cI2	Im3m	A2	W	1
γ-Mn	cF4	Fm3m	A2	Cu	1
β-Mn	cP20	P4132	A13	βMn	1
α-Mn	cI58	I43m	A12	α Mn	1
MnO	cF8	Fm3m	B1	NaCl	1
β -Mn ₃ O ₄	cF56	Fd3m	H1,		21
α -Mn ₃ O ₄	tI28	$I4_1/amd$			21
Mn_2O_3	cI80	Ia3	_		21
MnO ₂	tP6	$P4_2/mnm$	C4	TiO ₂	21

D. Bixbyite (Mn_2O_3)

Bixbyite also exists in at least two polymorphic forms. as pointed out as early as 1934 by Dubois.^[36] One form, generally called α , has a body-centered cubic (bcc) structure, while the other, generally called γ , has a spinel structure. According to Masson,^[37] the α form is stable relative to the γ form under all conditions. The γ form was thus not considered in the present work. For the assessment of the Gibbs energy of Mn₂O₃, the following experimental information was used. The heat capacity from 299 to 345 K and $^{\circ}S_{298}$ were measured by Robie and Hemingway,^[30] the heat content from 397 to 1698 K was studied by Orr,^[38] and the Gibbs energy of formation of α -Mn₂O₃ from Mn₃O₄ and oxygen was taken from the experimental measurements by Schaefer,^[22] Chen and Chen^[39] and Huebner and Sato.^[28] The value of $\Delta^{\circ}H_{298}$ is from Kubaschewski et al.^[20]

E. Pyrolusite (MnO₂)

There are large deviations among the experimental measurements of the oxygen pressure in equilibrium with Mn_2O_3 and MnO_2 .^[25,40–43] Because the oxidation of Mn_2O_3 to MnO_2 is very slow at ordinary pressure, most of the experimental measurements concerned were carried out from the decomposition process of MnO_2 , but it has been extremely difficult to attain the actual equilibrium conditions. According to a recent experimental investigation by Chen and Chen,^[39] the generally accepted data reported by Otto^[40] are the most reliable ones and they were now used, together with the reported values of the heat capacity from 298 to 376.5 K and $^{\circ}S_{298}$ by Robie and Hemingway,^[30] the heat content from 407 to 778 K by Moore,^[44] and $\Delta^{\circ}H_{298}$ from Kubaschewski *et al.*^[20]

F. Liquid

The experimental information on the liquid phase is rather limited. The temperature of the invariant equilibria were mainly taken from Trömel *et al.*,^[2] except for the eutectic temperature of manganese and manganosite which was from Chen.^[3] The available manganosite solubility in liquid manganese was taken from the experimental measurements by Jacob^[45] and Simeonov *et al.*^[17]

G. α -, β - and γ -Mn

There is no experimental information available on the oxygen solubility in the α -, β - and γ -Mn phases, and these were treated as pure.

III. THERMODYNAMIC MODELS

The compound-energy model developed by Hillert and co-workers^[46–49] has been successfully applied to metal oxide systems.^[50,51] A detailed thermodynamic description of the model used in the Fe-O system is provided by Sundman.^[50] In the present work, the two-sublattice model is used for the manganosite ($Mn_{1-x}O$) phase and the liquid phase, and the description of the model is the same as the one used in the Fe-O system.^[50] Hausmanite (Mn_3O_4), bixbyite (Mn_2O_3), and pyrolusite (MnO_2) are considered as stoichiometric compounds. The Gibbs energy of these compounds is represented as a function of temperature with five parameters.

A. Manganosite

According to many studies,^[2,11-14] manganese in the manganosite (halite) phase exists in two ionic states, Mn^{+2} and Mn^{+3} , and the fraction of Mn^{+3} increases with the oxygen potential and temperature, but some authors^[3,12] also reported the existence of the Mn^{+1} . In the present assessment, only Mn^{+2} and Mn^{+3} ions are considered. The ideal structure of halite is a simple face-centered cubic (fcc) arrangement of oxygen ions, containing manganese ions in the octahedral interstitial sites. In order to accommodate the trivalent manganese ions, vacant sites are introduced in the octahedral sublattice. Taking this into account, the halite phase may be represented with the following formula:

$$(Mn^{+2}, Mn^{+3}, Va)_1(O^{-2})_1$$

The phase can be represented in a constitutional triangle, and its physical significance was discussed by Sundman.^[50] The compound energy model gives the expression of the Gibbs energy as follows:

$$G_{m} = y_{2}^{\circ}G_{20} + y_{3}^{\circ}G_{30} + y_{v}^{\circ}G_{v0} + RT$$

$$\cdot [y_{2}\ln(y_{2}) + y_{3}\ln(y_{3}) + y_{v}\ln(y_{v})] + {}^{E}G_{m} \quad [1]$$

where the y_2 and y_3 are the fractions of sites in the metallic sublattice occupied by divalent and trivalent manganese ion, respectively, and y_v is the fraction of vacant sites. The y_i fractions are usually called site fractions. The three °G terms describe the Gibbs energy of stoichiometric MnO and hypothetical ionic compounds MnO⁺¹ and O⁻², respectively. According to Sundman,^[50] the parameter ° G_{vo} is set to zero for convenience. The excess Gibbs energy, ${}^{E}G_{m}$, for the halite is described with a Redlich-Kister expression of two terms:

$${}^{E}G_{m} = y_{2} y_{3} [L_{2,3:0}^{0} + (y_{2} - y_{3})L_{2,3:0}^{1}] + y_{2} y_{V} [L_{2,V:0}^{0} + (y_{2} - y_{V})L_{2,V:0}^{1}] + y_{3} y_{V} [L_{3,V:0}^{0} + (y_{3} - y_{V})L_{3,V:0}^{1}]$$
[2]

In the interaction parameters, a colon (:) is used to separate constituents in different sublattices, whereas a comma (,) is used to separate constituents in the same sublattice. In order to extrapolate to a multicomponent system, all parameters involving vacancies introduced for electroneutrality are set to zero. The relation between the mole fraction and the site fraction is expressed by the following formulas:

$$x_{\rm Mn} = \frac{y_2 + y_3}{1 + y_2 + y_3}$$
[3]

$$x_{\rm O} = \frac{1}{1 + y_2 + y_3}$$
[4]

Due to the condition of electroneutrality, the maximum value for y_3 is 2/3, and the model for the halite phase can describe the contents of oxygen x_0 from 0.5 to 0.6.

B. Liquid

For the liquid phase, the ions Mn^{+2} , Mn^{+3} , O^{-2} , and vacancies with an induced charge are considered, and the ionic two-sublattice model^[49] was adopted, which assumes that there is one sublattice for anions and one for cations and hypothetical vacancies. The formula can thus be written as

$$(Mn^{+2}, Mn^{+3})_P(O^{-2}, Va^{-Q})_Q$$

According to the model, the stoichiometric coefficients P and Q vary with the composition in order to make the phase electrically neutral, *i.e.*:

$$P = \sum_{i} y_{i}(-v_{i}) + y_{Va}Q$$
$$Q = \sum_{j} y_{j}v_{j}$$
[5]

where v_i is the valence of anion *i* and v_j the valence of cation *j*.

The Gibbs energy expression used for the liquid is

$$G_{m} = y_{2} y_{0}^{\circ} G_{Mn_{2}O_{2}} + y_{3} y_{0}^{\circ} G_{Mn_{2}O_{3}}$$

+ $Q(y_{2} y_{V}^{\circ} G_{Mn^{+2}:V^{-2}} + y_{3} y_{V}^{\circ} G_{Mn^{+3}:V^{-3}})$
+ $PRT[y_{2} \ln(y_{2}) + y_{3} \ln(y_{3})]$
+ $QRT[y_{0} \ln(y_{0}) + y_{V} \ln(y_{V})] + {}^{E}G_{m}$ [6]

The first and second terms contain the Gibbs energy of 2 moles of MnO and 1 mole Mn_2O_3 , respectively. The third term is the Gibbs energy of pure manganese. The

fourth term represents a hypothetical liquid Mn in an excited state. The excess Gibbs energy term, ${}^{E}G_{m}$, with two Redlich-Kister terms, is

$${}^{E}G_{m} = y_{2} y_{3} y_{V} [L_{2,3:V}^{0} + (y_{2} - y_{3})L_{2,3:V}^{1}] + y_{2} y_{3} y_{0} [L_{2,3:O}^{0} + (y_{2} - y_{V})L_{2,3:V}^{1}] + y_{2} y_{0} y_{V} [L_{2:O,V}^{0} + (y_{0} - y_{V})L_{2:O,V}^{1}] + y_{3} y_{0} y_{V} [L_{3:O,V}^{0} + (y_{0} - y_{V})L_{3:O,V}^{1}]$$
[7]

The four sets of interaction parameters represent the sides of a reciprocal system. However, $L_{2,3:V}^0$, $L_{2,3:V}^1$, $L_{3:O,V}^0$, and $L_{3:O,V}^1$ are put to zero, because they are important only far from the composition of the stable liquid phase and cannot be evaluated with any certainty. The remaining four interaction parameters were used to fit the experimental data.

IV. RESULTS AND DISCUSSION

The assessment was carried out by using the computer software PARROT, developed by Jansson,^[52] included in the Thermo-Calc databank system.^[53] The evaluated parameters of the various phases are listed in Table II, and the calculated phase diagram at 0.21 atm of oxygen is shown in Figure 1. A list of invariant equilbria is given in Table III, and the calculated values of $\Delta^{\circ}H_{298}$ and $^{\circ}S_{298}$ of various compounds are shown in Table IV and compared with the values from literature.

A. Manganosite

The assessment began with stoichiometric manganosite. The calculated heat content and Gibbs energy of formation fit well with the experimental data, which are shown in Figures 2 and 3. Much attention was then paid to reproduce the oxygen activity as a function of composition. The final accepted results are plotted in Figure 4(a) and compared with the experimental data by



Fig. 1-Calculated Mn-O phase diagram at 0.21 atm of oxygen.

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The magnetic contribution to Gibbs energy is described by
G_{\rm m}^{\rm mo} = \mathbf{R}T \ln (\beta + 1)f(\tau), \ \tau = T/Tc
For \tau < 1: f(\tau) = 1 - \left[\frac{79\tau^{-1}}{140p} + \frac{474}{497}\left(\frac{1}{p} - 1\right)\left(\frac{\tau^3}{6} + \frac{\tau^9}{135} + \frac{\tau^{15}}{600}\right)\right] / A
For \tau > 1: f(\tau) = -\left(\frac{\tau^{-5}}{10} + \frac{\tau^{-15}}{315} + \frac{\tau^{-25}}{1500}\right) / A
where A = \left(\frac{518}{1125}\right) + \left(\frac{11,692}{15,975}\right)
\left[\left(\frac{1}{p}\right)-1\right] and p depends on the structure.
Gas (O<sub>2</sub>)
Constituents O<sub>2</sub>
^{\circ}G_{O_2}^{GAS} - H_{O_2}^{SER} = +G(O_2, gas) + RT \ln P
 Ionic-liquid
2 sublattices, sites 2:2
Constituents Mn<sup>+2</sup>, Mn<sup>+3</sup>:O<sup>-2</sup>, Va
^{\circ}G_{Mn^{+}2:0^{-}2}^{IONIC-LIQ} - 2H_{Mn}^{SER} - 2H_{O}^{SER} = +2G(MnO, liq)
^{\circ}G_{Mn+3:\Omega=2}^{\text{IONIC-LIQ}} - 2H_{Mn}^{\text{SER}} - 3H_{\Omega}^{\text{SER}} = +2.5G(\text{MnO}, \text{liq}) - 390,784 + 291.8765T
{}^{\circ}G_{Mn+2;Va}^{\text{IONIC-LIQ}} - H_{Mn}^{\text{SER}} = +G(Mn, \text{liq})
{}^{\circ}G_{Mn}^{IONIC-LIQ} - H_{Mn}^{SER} = +G(Mn, liq) + 87,028
{}^{0}L_{Mn^{+2};0^{-2},Va}^{IONIC-LIQ} = +186,771 - 12.9648T
 {}^{1}L_{Mn^{+}2:O^{-}2,Va}^{IONIC-LIQ} = +219,594 - 132.928T
 {}^{0}L_{Mn^{+2},Mn^{+3};O^{-2}} = +1,344,189 - 803.1101T
 {}^{1}L_{Mn^{+2}Mn^{+3}:O^{-2}}^{IONIC-LIQ} = -11 + 52.122T
 BCC-A2 (δ-Mn)
 2 sublattices, sites 1:3
 Constituents Mn, O:Va
 298.15 < T < 1519
 {}^{\circ}G_{M_0,V_0}^{\text{HBCC-A2}} - H_{M_0}^{\text{SER}} = -3235.3 + 127.85T - 23.7T \ln T - 0.00744271T^2 + 60,000T^{-1}
 1519 < T < 3000
 ^{\circ}G_{Mn:Va}^{HBCC:A2} - H_{Mn}^{SER} = -23,188.83 + 307.7043T - 48T \ln T + 1.265152 \times 10^{30} T^{-9}
 {}^{\circ}G_{O:Va}^{HBCC-A2} - H_{O}^{SER} = +30,000 + G(O_2, gas)
 p = 0.4
 {}^{0}Tc_{Mn:Va}^{BCC-A2} = -580
 {}^{0}\beta_{\text{Mn:Va}}^{BCC-A2} = -0.27
 CBCC-A12 (a-Mn)
 2 sublattices, sites 1:1
 Constituents Mn: Va
 ^{\circ}G_{Mn;Va}^{HCBCC-A12} - H_{Mn}^{SER} = +GH^{SER}(Mn)
 p = 0.28
 {}^{0}Tc_{Mn;Va}^{CBCC-A12} = -285
 {}^{0}\beta_{Mn;Va}^{CBCC-A12} = -0.66
 FCC-Al (y-Mn)
 2 sublattices, sites 1:1
 Constituents Mn,O: Va
 298.14 < T < 1519
  G_{M_0,V_a}^{\text{FCC-A1}} - H_{M_n}^{\text{SER}} = -3439.3 + 131.884T - 24.5177T \ln T - 0.006T^2 + 69.600T^{-1}
  1519 < T < 3000
  ^{\circ}G_{\text{Mn:Va}}^{\text{FCC-A1}} - H_{\text{Mn}}^{\text{SER}} = -26,070.1 + 309.6664T - 48T \ln T - 3.86196 \times 10^{30} T^{-9}
  ^{\circ}G_{O;Va}^{HFCC-AI} - H_{O}^{SER} = +30,000 + G(O_2, gas)
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p = 0.28
{}^{0}Tc_{Mn;Va}^{FCC-A1} = -1620
{}^{0}\beta_{Mn'Va}^{FCC-A1} = -1.86
CUB-A13 (β-Mn)
2 sublattices, sites 1:1
Constituents Mn: Va
298.14 < T < 1519
{}^{\circ}G_{\mathsf{Mn}\cdot\mathsf{Va}}^{\mathsf{CUB-A13}} - H_{\mathsf{Mn}}^{\mathsf{SER}} = -5800.4 + 135.995T - 24.8785T \ln T - 0.00583359T^2 + 70,269T^{-1}
1519 < T < 3000
^{\circ}G_{M_{0};V_{a}}^{CUB-A13} - H_{M_{a}}^{SER} = -28,290.76 + 311.2933T - 48T \ln T - 3.96757 \times 10^{30} T^{-9}
Halite (Mn<sub>1-x</sub>O)
2 sublattices, sites 1:1
Constituents Mn<sup>+2</sup>, Mn<sup>+3</sup>, Va:O<sup>-2</sup>
{}^{\circ}G_{Mn+2;O-2}^{HALITE} - H_{Mn+2}^{SER} - H_{O-2}^{SER} = +G(MnO)
^{\circ}G_{Mn^{+3}:O^{-2}}^{HALITE} - H_{Mn^{+3}}^{SER} - H_{O^{-2}}^{SER} = +G(MnO) - 60,498.07 - 20.03189T
^{\circ}G_{\text{Va}:0^{-2}}^{\text{HALITE}} - H_{0^{-2}}^{\text{SER}} = 0
{}^{0}L_{Mn^{+2}Mn^{+3}:0^{-2}}^{HALITE} = +36,538.5
\alpha-Mn<sub>3</sub>O<sub>4</sub>
2 sublattices, sites 3:4
Constituents Mn:O
^{\circ}G_{\text{Mn};O}^{\text{Mn};O_4-\text{ALPHA}} - 3H_{\text{Mn}}^{\text{SER}} - 4H_O^{\text{SER}} = +G(\text{Mn};O_4)
\beta-Mn<sub>3</sub>O<sub>4</sub>
2 sublattices, sites 3:4
Constituents Mn:O
^{\circ}G_{Mn;O}^{Mn;O4-BETA} - 3H_{Mn}^{SER} - 4H_{O}^{SER} = +G(Mn;O_{4}) + 20,810 - 14.376047T
Mn<sub>2</sub>O<sub>3</sub>
2 sublattices, sites 2:3
Constituents Mn:O
^{\circ}G_{Mn_{2}O_{3}}^{Mn_{2}O_{3}} - 2H_{Mn}^{SER} - 3H_{O}^{SER} = +G(Mn_{2}O_{3})
MnO<sub>2</sub>
2 sublattices, sites 1:2
Constituents Mn:O
{}^{\circ}G_{MnO2}^{MnO2} - H_{Mn}^{SER} - 2H_{O}^{SER} = +G(MnO_2)
Functions:
298.15 < T < 1000
G(O_2, gas) = -6961.74451 - 76,729.7484T^{-1} - 51.0057202T - 22.2710136T \ln T - 0.0101977469T^2
                + 1.32369208 \times 10^{-6} T^{3}
1000 < T < 3300
G(O_2, gas) = -13,137.5203 + 525.809.556T^{-1} + 25.32000332T - 33.627603T \ln T - 0.00119159274T^2
                + 1.35611111 × 10^{-8} T^3
3300 < T < 6000
G(O_2, GAS) = -27,973.4908 + 8,766,421.4T^{-1} + 62.5195726T - 37.9072074T \ln T - 8.50483772 \times 10^{-4}T^2
                  + 2.14409777 × 10^{-8} T^3
298.15 < T < 1519
G(Mn, liq) = +GH^{SER}(Mn) + 17,859.91 - 12.6208T - 4.41929 \times 10^{-21} T^7
1519 < T < 2000
G(Mn, liq) = +GH^{SER}(Mn) + 18,739.51 - 13.2288T - 1.656847 \times 10^{30} T^{-9}
298.14 < T < 1519
GH^{\text{SER}}(\text{Mn}) = -8115.28 + 130.059T - 23.4582T \ln T - 0.00734768T^2 + 69.827T^{-1}
1519 < T < 2000
GH^{\text{SER}}(\text{Mn}) = -28,733.41 + 312.2648T - 48T \ln T + 1.656847 \times 10^{30} T^{-9}
G(MnO, liq) = -252,595 - 293.164T + 15.0712T \ln T
G(\text{MnO}) = -401,784.4 + 268.75251T - 48.2744592T \ln T - 0.00323061T^2 + 257,553.188T^{-1}
G(Mn_{3}O_{4}) = -1,443,167 + 905.71T - 156.211772T \ln T - 0.0179602213T^{2} + 1,106,277.72T^{-1}
G(Mn_2O_3) = -993,634.5 + 584.605647T - 102.148721T \ln T - 0.0181244679T^2 + 595,113T^{-1}
G(MnO_2) = -548,242 + 444.10228T - 72.5000062T \ln T - 0.002720343T^2 + 885,523.74T^{-1}
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Table III.	Invariant	Equilibria	in	the	Assessed	Mn-O	S	ystem	at	0.21	Atr
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				Temperature (K)			
Reaction	Mole Fraction C) in the Respective	Calculated	Experimental			
$L1 = \delta - Mn + Mn_{1-x}O$	1.46×10^{-3}	2.005×10^{-3}	0.5	1515	1515 ^[2,3]		
$L2 = L1 + Mn_{1-x}O$	3.65×10^{-11}	0.4886	0.5	2063	2063 ^[2]		
$Mn_{1-x}O = L2$	0.5	0.5		2123	2123 ^[2,4]		
$L2 = Mn_{1-x}O + \beta - Mn_3O_4$	0.5419	0.5275	0.5714	1833	1833 ^[2]		
β -Mn ₃ O ₄ = L2 + Gas	0.5714	0.5441	1	1846	1853, ^[2] 1840 ^[5]		
$Mn_2O_3 = \alpha - Mn_3O_4 + Gas$	0.6	0.5714	1	1161	1162 ^[39]		
$MnO_2 = Mn_2O_3 + Gas$	0.6667	0.6	1	697	694 ^[40]		
$\alpha - Mn_3O_4 = \beta - Mn_3O_4$	0.5714	0.5714		1448	1445 ^[20,23,32,33]		
γ -Mn = δ -Mn	0	0	~	1383			
β -Mn = γ -Mn	0	0	~	1360			

Table IV. The Assessed and Reported Values of $\Delta^{\circ}H_{298}$ and $^{\circ}S_{298}$

Compounds	Calculated $\Delta^{\circ}H_{298}(\text{KJ/mol})$	Reported $\Delta^{\circ}H_{298}(KJ/mol)$	Calculated °S ₂₉₈ (J/mol/K)	Reported °S ₂₉₈ (J/mol/K)
$\frac{Mn_{1-x}O}{\alpha - Mn_3O_4}$	-385.2 -1387.6	$-385.2^{(20)}$ -1387.6 ⁽²⁰⁾	59.5 163.7	$59.9^{[20]}$ $164.1,^{[30]}$ $154^{[20]}$
Mn_2O_3 MnO_2	-957.6 -520.4	$-957.6^{[20]} \\ -520.4^{[20]}$	117.0 53.1	$\begin{array}{cccc} 113.7, \stackrel{[30]}{}{52.8}, \stackrel{[30]}{}{53.2} \begin{array}{c} 110.5 \stackrel{[20]}{}{52.8}, \stackrel{[30]}{}{53.2} \end{array}$

-20

-22

-24



Fig. 2-Calculated vs experimental heat content of manganosite.

Picard and Gerdanian.^[12] Expressed in oxygen content, the deviations are within the experimental uncertainty. In Figure 4(b), other experimental data are shown, together with the calculated curves. As already mentioned, there are large discrepancies among the reported data. It is suspected that the reliability of the experimental technique and the extreme difficulty in attaining of equilibrium in this system are the main reasons for the discrepancies. Figure 5 shows the oxygen solubility in the manganosite phase compared with the experimental data by Fender and Riley.^[15]



the hausmanite was based on the evaluated parameters of manganosite phase and the selected experimental data. The parameters are given in Table II. Figures 6 and 7 show the calculated heat capacity and enthalpy varying with temperature, both of which have been fitted well to the experimental data. The oxygen potential is presented in Figure 8 in comparison with the experimental data by many authors.^[22-29] The calculated temperature (T =1448 K) and enthalpy ($\Delta H = 20,810 \text{ J/mol}$) of the

2000





Fig. 4—(a) Calculated oxygen activity in manganosite vs experimental data used in the optimization and (b) calculated oxygen activity in manganosite vs other experimental data which were not used in the optimization.

structural transformation of hausmanite also fit with the selected data well.

C. Liquid Phase

With the parameters of solid phases evaluated, the parameters of the liquid phase were optimized to fit all the invariant equilibria and the miscibility gap. More attention was paid to fit the temperatures rather than the compositions. The calculated temperatures of the invariant equilibria are listed in Table III, compared with the values reported. The calculated oxygen solubility in liquid manganese fits well with the experimental data measured



Fig. 5—Calculated oxygen solubility in manganosite with experimental data.



Fig. 6—Calculated vs experimental heat capacity of hausmanite (Mn_3O_4) .

by $Jacob^{[45]}$ and Simeonov *et al.*,^[17] as shown in Figure 9, close to pure manganese.

D. Bixbyite

With the parameters of the hausmanite phase and the experimental information, the Gibbs energy of the bixbyite phase was evaluated with five parameters. The calculated heat capacity, heat content, and oxygen potential *vs* temperature are shown in Figures 10 through 12. In



Fig. 7—Calculated vs experimental heat content of hausmanite (Mn_3O_4) .



Fig. 9—Calculated vs experimental oxygen solubility in liquid manganese.



Fig. 8—Calculated vs experimental oxygen potential of hausmanite (Mn_3O_4) .



Fig. 10-Calculated vs experimental heat capacity of bixbyite (Mn₂O₃).



Fig. 11-Calculated vs experimental heat content of bixbyite (Mn₂O₃).



Fig. 13—Calculated vs experimental heat capacity of pyrolusite (MnO_2) .



Fig. 12—Calculated vs experimental oxygen potential of bixbyite (Mn_2O_3) .



Fig. 14-Calculated vs experimental heat content of pyrolusite (MnO₂).



Fig. 15—Calculated vs experimental oxygen potential of pyrolusite (MnO_2) .

Figure 12, other experimental data by Kim *et al.*,^[25] Ingraham,^[54] Charette and Flengas,^[26] Shenouda and Aziz,^[55] and Hahn and Muan,^[29] which were not used in the optimization, are also included. Good agreement between the calculation and the experimental data can be observed.

E. Pyrolusite

The Gibbs energy of pyrolusite was finally assessed based on the selected experimental data, and the results are shown in Figures 13 through 15 for the heat capacity, heat content, and oxygen potential. The present description of the Gibbs energy of the pyrolusite phase reproduces well the experimental data of the heat capacity by Robie and Hemingway,^[30] the heat content by Moore,^[44] and the oxygen potential only by Otto.^[40] The other experimental results obtained from the decomposition oxygen pressure may be less reliable, as discussed earlier.

V. SUMMARY AND CONCLUSIONS

The present thermodynamic evaluation of the Mn-O system has been made by using thermodynamic models for the Gibbs energy of individual phases. An optimized set of thermodynamic parameters has been obtained on the consideration of related experimental information. For the solid manganese oxides, the parameters of the Gibbs energy have been revised by considering all the available thermodynamic information. Due to the scarce and inconsistent experimental information on some phases, there are some uncertainties in the evaluation of thermodynamic properties still remaining, particularly for the liquid phase. However, with the present assessment, a basic thermodynamic description of the Mn-O system is provided which permits extrapolations and can be included in the description of the higher order systems.

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