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¹⁵N abundance of soils and plants along an experimentally induced forest nitrogen supply gradient

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Abstract ¹⁵N abundances of soils and a grass species (*Deschampsia flexuosa* (L.) Trin.) were analysed in a forest fertilization experiment 10 years after the last fertilization. Nitrogen had been given as urea, at seven doses, ranging from 0 to 2400 kg N ha⁻¹. Previously, we have shown that plants in systems experiencing large losses of N become enriched with ¹⁵N. This was explained by the fact that processes leading to loss of N, e.g. ammonia volatilization, nitrification followed by leaching or denitrification and denitrification itself, tend to fractionate against ¹⁵N. In this experiment, ¹⁵N abundance increased with dose of N applied in both grass and soil total-N, but more so in the grass. This was interpreted to be due to the grass sampling small but active pools of N subject to losses. In contrast, soil total-N largely consists of inactive N that does not immediately exchange with pools of N from which fractionating losses occur. Hence, soil total-N shows a large pretreatment ¹⁵N memory effect, and is, therefore, an integrator of the long-term N balance. When short-term changes (years, decades) in N balances are monitored using variations in ¹⁵N abundance, plants are more suitable indicators of such change than is soil total-N.

Key words ¹⁵N · Forests · N cycling · Soil N pools

Introduction

Elevated levels of nitrogen deposition, and adverse effects thereof on terrestrial and aquatic ecosystems, are of increasing concern (Aber et al. 1989; Schulze 1989; Tamm 1991). Högberg (1990b, 1991) proposed studies of the abundance of the stable nitrogen isotope ¹⁵N to monitor changes in the N balance of ecosystems. Processes leading to loss of N, i.e. NH₃ volatilization, nitrifi-

cation followed by denitrification or leaching, and denitrification, discriminate against the heavier ¹⁵N (Shearer and Kohl 1986). This leaves the remaining N enriched in ¹⁵N as suggested by Vitousek et al. (1989) and experimentally verified by Högberg (1990b, 1991) and Högberg and Johannisson (1993) studying plants in a forest fertilization trial.

When fertilizer N has a ¹⁵N abundance lower than that of soil-derived N, nitrogen fertilization should lower the ¹⁵N abundance of N in the system (Kohl et al. 1973). However, as stated above, when inputs of N are large, this initial decrease is followed by an enrichment in ¹⁵N (Högberg 1991) as a result of loss of N (Högberg and Johannisson 1993).

It has generally been concluded that the ¹⁵N abundance of vegetation reflects the ¹⁵N abundance of N sources available to the plants (Shearer and Kohl 1986). Commonly, plants have lower ¹⁵N abundances than soil total-N (Gebauer and Schulze 1991; Handley and Raven 1992; Virginia and Delwiche 1982; Vitousek et al. 1989), but exceptions have been reported (Pate et al. 1993). Gebauer and Schulze (1991) showed that the ¹⁵N abundance of both soil total-N and N in fine roots increased with depth in a spruce forest, but also that the ¹⁵N abundance of roots was 0.5–4.2 δ ¹⁵N (‰) lower than that of soil total-N.

The relation between the ¹⁵N abundances of plants and soils is evidently crucial when evaluations of N balances and the development of N saturation are to be based on plant samples (Högberg 1990b, 1991). For example, Sutherland et al. (1993) found a greater variability in ¹⁵N abundance of plants than of soils in durum wheat fields, and attributed this to the greater susceptibility of the small mineral-N pool to isotopic fractionation than the larger inert soil organic-N pool. In this paper we show how a wide range of defined N additions to a forest have affected ¹⁵N abundances of both plants and soils, and based on these observations, make inferences about soil N pool dynamics.

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Materials and methods

The experimental site is located in a Norway spruce (*Picea abies* Karst.) forest on a gentle slope, at Nyby in northern Sweden (64°20'N, 18°27'E, 370 m a.s.l.). The soil is a podsolized sandy-silty till supporting a rather poor tree stand. In 1966, when the stand was 142 years old, the mean tree height was only 16 m. The field layer is dominated by *Vaccinium myrtillus* L., while the grass species *Deschampsia flexuosa* (L.) Trin. is sparsely distributed (Kellner 1993). Soil pH_(H2O) varies between 4.0–4.5 and 4.0–4.2 in the mor layer and in the upper 5 cm of the mineral soil, respectively (H.-Ö. Nohrstedt, pers. comm.). The C/N ratio of the H-layer of the mor is around 32. Deposition of N is below 5 kg ha⁻¹ yr⁻¹.

We studied the experiment D67 maintained by the Research Institute SkogForsk. Two blocks of plots, with seven 30 × 30 m plots each, had been fertilized with urea at seven levels (NO–N6) in the years of 1966, 1971, 1976 and 1981. Total additions of N varied between 0 and 2400 kg ha⁻¹. The experimental design is summarized in Table 1. We have not measured the ¹⁵N abundance of fertilizer-N, but know that the fertilizer was supplied by a Swedish manufacturer, Supra, which has been reported to produce urea significantly depleted in ¹⁵N (Högberg 1990b). Low annual additions of urea (\bar{x} = 36 kg N ha⁻¹), supplied by Supra, caused a consistent decline in ¹⁵N abundance in a nearby experiment during the period 1971–1989 (Högberg 1991).

Fertilized spruce grew consistently better than non-fertilized spruce, and responded positively to each fertilization without obvious signs of saturating responses to N. There was, however, no additional growth effect of N applications at levels greater than N4 (F. Pettersson, pers. comm.). In 1989, the annual increment of basal area on fertilized plots was 120–240% of that on control (NO) plots. Nitrogen fertilization also increased the coverage of *D. flexuosa* at the expense of *V. myrtillus*.

Ten cores (10 cm diameter) of the mor-layer were taken at random from the inner 20 × 20 m of each plot in late August 1991, i.e. 10 years after the last fertilization. The H-layer was used for analysis. A composite sample of a few grams of vegetative shoots of the grass *D. flexuosa* was taken from each of the plots (see Högberg et al. 1986). The samples were oven-dried (70°C, 48 h). Subsamples of soils and grass were ground in a ball mill and analysed for total-N and ¹⁵N abundance (Haystead, 1983). Results are reported in per cent N and in $\delta^{15}\text{N}$ (‰): $\delta^{15}\text{N} = 1000 \times (R_{\text{sample}} - R_{\text{standard}}) / R_{\text{standard}}$ (‰), where $R = \text{mass } 29 / \text{mass } 28$, and the standard used was pure N₂ with a ¹⁵N abundance of -2.6 ‰ ¹⁵N (‰) relative to atmospheric N₂ (Mariotti 1983).

The dominating field layer species, *V. myrtillus*, assimilates little NO₃⁻ (Gebauer et al. 1988; Högberg et al. 1990; Lee and Stewart 1978; but see Ingestad 1973). In contrast, *D. flexuosa*, although preferentially utilizing NH₄⁺ (Gigon and Rorison 1972), also readily assimilates NO₃⁻ (Högberg et al. 1986; Lee and Stewart 1978) and was, therefore, used as an indicator of ¹⁵N of the pools of N available to plants. Whether NH₄⁺ or NO₃⁻ is taken up

also depends on their concentrations and mobility in the soil system (see e.g. Marschner et al. 1991).

Data were analysed by the GLM procedure (e.g. ¹⁵N_{soil} = block; covariate N-dose) of MINITAB (PC Version 8.0, Microsoft Corp., 1991). Partial correlation between ¹⁵N_{soil} and ¹⁵N_{grass} and between ¹⁵N_{soil} and N concentration in the soil, was calculated after elimination of treatment and block effects. It was tested whether the (theoretical) partial correlation was zero by means of the test statistic $t = r_1 (n-4)^{1/2} / (1-r_1^2)^{1/2}$, where r_1 is the coefficient of correlation between residuals. The statistic t has approximately a t -distribution with $n-4$ degrees of freedom under the null-hypothesis (a modification of Quade 1989). In the case of the soil samples, means per plot were used in the analyses.

Results

Ten years after the last fertilization, treatment effects on ¹⁵N abundance were still evident in the soil and in the grass. The ¹⁵N abundance of the soil and of the grass increased with dose of N fertilizer applied ($P = 0.014$ and $P = 0.000$, respectively; Fig. 1). The grass had a lower ¹⁵N abundance than the soil, but the increase in ¹⁵N was larger in the grass than in the soil ($P = 0.000$). Nevertheless, when treatment effects on ¹⁵N abundance were partialled out or accounted for, variations in ¹⁵N of the grass were highly correlated with the ¹⁵N abundance of the soil ($r_1 = 0.806$, $P = 0.002$). For example, relatively high ¹⁵N abundance in the grass in treatment N1 in block 1 corresponded to a high ¹⁵N abundance in the soil. The difference between the grass and the soil was significantly smaller in block 1 ($P = 0.004$). The standard deviation of the mean for each plot varied between 0.6–1.7 ‰ ¹⁵N (‰) in the soil, i.e. it was large compared to the absolute values.

The concentration of N in the soil increased with level of fertilizer application (Fig. 2, $P = 0.000$). A block effect was found ($P = 0.039$); block 1 having a somewhat higher N concentration than block 2. There was no sig-

Table 1 Experimental design of the experiment D67 at Nyby

Number of blocks	2	
Number of treatments	7	
Number of fertilizations	4	
Years of fertilization	1966, 1971, 1976, 1981	
N applied (kg N ha ⁻¹)	Per occasion	Total
N0	0	0
N1	120	480
N2	180	720
N3	240	960
N4	360	1440
N5	480	1920
N6	600	2400

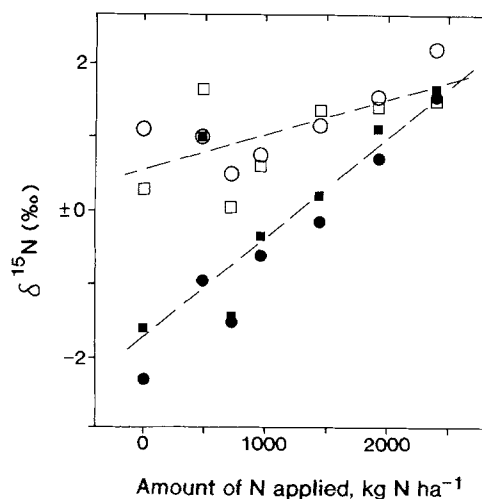


Fig. 1 ¹⁵N abundances of soils (open symbols) and of the grass species *Deschampsia flexuosa* (closed symbols) in 1991, in relation to amount of N applied during the period 1966–1981 to plots at Nyby. Squares, block 1; circles, block 2

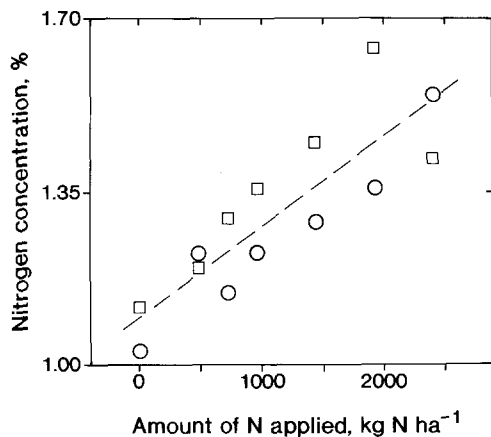


Fig. 2 Nitrogen concentrations of soils (H-layer) in 1991, in relation to amount of N applied during the period 1966–1981 to plots at Nyby. Symbols as in Fig. 1

nificant treatment effect on the N concentration in the grass. When treatment effects were partialled out, neither the ^{15}N abundance of the grass nor that of the soil were correlated with the N concentration of the soil ($r_1 = -0.031$ and $r_1 = 0.012$, respectively).

Discussion

Most N in soils is in the form of organic N (Tamm 1991), and the greater portion of soil organic N is fixed in stable forms (Jansson 1958; Paul and Clark 1988; Paul and Juma 1981). This humus N may constitute a passive organic phase existing more or less beside the pathways of the mineralization-immobilization turnover cycle (Jansson 1958). The active organic phase would then be more readily mineralizable N of plant and microbial origin plus N in the living microbial biomass. Nitrogen transformations, e.g. mineralization, nitrification and immobilization, within the biologically active pools of microbial N, NH_4^+ and NO_3^- are considerable although the actual concentrations of NH_4^+ and, especially, NO_3^- are low in forest soils (Davidson et al. 1992).

After fertilization urea is hydrolysed to NH_3 , which readily reacts to form NH_4^+ . Estimates of volatilization losses of NH_3 -N during the first few weeks after fertilizer application vary from <5% to >40% of the N applied in urea fertilized forests (Nason et al. 1988). NH_3 is, furthermore, chemically incorporated into stable organic compounds in the mor layer (Foster et al. 1985). Volatilization of NH_3 fractionates against ^{15}N (Shearer and Kohl 1986), leading to ^{15}N enrichment of the soil NH_4^+ pool. Except for the first few weeks following fertilization, losses by NH_3 volatilization should be negligible from the acid soils at Nyby.

In this experiment, the last fertilization took place 10 years before the study. The ^{15}N abundance of plant N should primarily reflect the ^{15}N abundances of NH_4^+ and NO_3^- in the soil solution or on exchange sites plus active organic pools of N. Losses of soil-N in urea-fertil-

ized forests should be from the $\text{NH}_3/\text{NH}_4^+$ pool directly following fertilization, and thereafter from the NO_3^- pool, and therefore affect the isotopic composition of these pools, as well as that of pools of N exchanging with them (see Davidson et al. 1992). In contrast, the ^{15}N abundance of soil total-N is an integrated measure of the ^{15}N abundance of all pools of N, which involves dilution of treatment effects on ^{15}N abundance by a large pool of old unreactive stable organic N. Hence, the change in ^{15}N abundance of the soil is less evident than that of the grass (Fig. 1).

Theoretically, the plants could acquire a higher ^{15}N abundance than soil total-N at high levels of N application (Fig. 1), as a result of increasingly higher ^{15}N enrichment of active pools of N. This has now been observed in a nearby experiment similar to the Nyby trial (L. Högbom & C. Johannisson, pers. comm.). Plants and soils have also been studied along a steep natural N supply gradient c. 50 km from our site in Nyby; the difference in ^{15}N abundance between plants and soils diminished with increasing supply of N there also (M. Högbom & P. Högbom, pers. comm.). Recently, Pate et al. (1993) reported that in Australian woodland disturbed by fire, resin extractable NO_3^- and NO_2^- in the xylem sap was enriched with ^{15}N relative to soil total-N. The ^{15}N abundance of plants increased with increasing utilization of NO_3^- , and some plants reached a higher ^{15}N abundance than soil total-N. This contradicts the simple view that one only has to consider that products should become progressively depleted (Högbom 1990a), which would be appropriate in a closed system without losses.

We have assumed that fertilizer-N had a lower ^{15}N abundance than soil total-N. The fact that the lowest ^{15}N abundance of soil total-N is found in N2, from which treatment the ^{15}N abundance progressively increases with amount of N applied, supports this assumption.

Exchange of N between active and passive pools of N, should, with time, reduce the magnitude of the ^{15}N enrichment of the active pools of N. As the experiment was sampled 10 years after the last treatment, this could explain why the difference in ^{15}N abundance in the grass between the lowest and the highest dose of N applied was only one-third of the difference found at Norrleden (Högbom 1990b), which was studied under conditions of experimental steady-state, i.e. during the 18th year of annual fertilizations. Estimates of the microbially available pool of N (Nordgren 1992) at Nyby showed no differences between treatments in 1993 (A. Nordgren, pers. comm.). Neither did we observe any increase in N concentration in the grass with increasing dose of N applied, which is in contrast to the large increase in soil N concentration (Fig. 2). Hence, the isotopic labelling of the pools of N available to plants is largely reflecting the N dynamics during and shortly after experimental treatments, and not the current N dynamics.

In conclusion, applications of unlabelled N to a forest ecosystem induced increases in ^{15}N abundances of

both soils and plants, which were detectable 10 years after the last treatment. Workers interested in monitoring short-term (years, decades) changes in N balances of forests by use of variations in ^{15}N abundance (Högberg 1990b) should observe the plants, which sample small, but active pools of soil N involved in losses from the system as well as other fractionating processes, rather than soil total-N, which mainly represents passive N carrying a large ^{15}N memory effect. Those interested in the long-term N balance should analyze the soils.

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