

Responses of benthic–pelagic coupling to climate change in a temperate estuary

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Abstract This article reports the first demonstration of the impact of climate change on benthic–pelagic coupling and the biogeochemical cycles of a coastal marine system. Over the last 30 years Narragansett Bay, a 328-km² temperate estuary on the east coast of the United States, has undergone a variety of ecological changes. Building on a robust data set that spans three decades, we present a link between warming (+1.7°C in annual mean water temperature) in the bay and a marked decrease in sediment oxygen consumption, in the fluxes of ammonium and phosphate from sediments to the overlying water, and in sediment denitrification. We attribute this reduction in biogeochemical exchange to a dramatic drop in the standing crop of water-column chlorophyll as the system has shifted from one characterized by a dominant winter–spring bloom to one supported by

more ephemeral and less intense summer–autumn blooms. The recent climate-induced oligotrophication of the bay will be further exacerbated by forthcoming nitrogen reductions due to tertiary sewage treatment.

Keywords Climate change · Benthic–pelagic coupling · Oligotrophication · Denitrification · Nutrient regeneration · Phenology

Introduction

Even if the climate change—eutrophication linkage is well established for the pelagic system...it remains poorly defined for benthic–pelagic coupling and the benthos.

Grall & Chauvaud (2002)

Coastal ecosystems around the world have been subjected to a variety of locally driven anthropogenic changes (e.g., nutrient enrichment: Valiela et al., 1992; Kinney & Roman, 1998; eutrophication: de Jonge et al., 1994; overfishing: Jackson et al., 2001; toxic pollutant loading: Kot-Wasik et al., 2004). Further complicating these local issues are regional and global forces, most notably climate change. Obvious possible effects of climate change on coastal systems include sea level rise, changes in precipitation patterns, and the associated pulsing of freshwater, nutrients, and sediment (Scavia et al., 2002), and perhaps increased

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intensity of storms (Emanuel, 2005). Climate change may also affect coastal systems in more subtle ways by altering the phenology or sequencing of events (Farnsworth et al., 1995; Inouye et al., 2000; Walther et al., 2002; Edwards & Richardson, 2004). To date, the interactions between climate change, sediment biogeochemical cycling, and benthic–pelagic coupling have largely been ignored.

Over the last decade, various ecological changes in Narragansett Bay (Rhode Island, USA) have been reported, including a decrease in mean annual water column chlorophyll concentrations (Li & Smayda, 1998), replacement of boreal demersal fish with demersal decapods (Oviatt, 2004), and earlier development and larger populations of the major predator zooplankton, *Mnemiopsis leidyi* (Sullivan et al., 2001). These changes are thought to be mainly the result of climate variability (North Atlantic Oscillation) and a long-term warming trend (Nixon et al., 2003, 2004; Oviatt, 2004). There is a long history of work documenting the strong linkages between the benthos and the water column in Narragansett Bay, including benthic oxygen uptake and nutrient regeneration (Nixon et al., 1976, 1980), denitrification (Seitzinger et al., 1984; Nowicki & Oviatt, 1990), benthic responses to organic additions of ^{15}N -labeled (Garber, 1982) and ^{14}C -labeled (O'Reilly, 1984; Rudnick & Oviatt, 1986) and unlabeled (Kelly & Nixon, 1984; Grassle et al., 1985) organic matter, and predation pressure by the benthos on the water column (Doering et al., 1986; Doering, 1989; Oviatt, 2004). Building upon this study, we recently revisited the most well-studied site in the bay to measure benthic oxygen uptake, nutrient regeneration, and sediment denitrification over an annual cycle. Our main motivation for this study was to see whether it would be possible to document any changes in the biogeochemical cycling of Narragansett Bay benthos over the last 30 years. Also, the major wastewater treatment plants discharging to the bay are moving to tertiary treatment. This research will provide baseline data for future studies of denitrification and benthic fluxes after tertiary wastewater treatment is implemented. While this management intervention may improve the water quality in the upper bay, the consequences for the much larger mid- and lower bay are unknown.

The purpose of this article is twofold. The main goal is to report, we believe for the first time, a link

between climate-induced changes in benthic–pelagic coupling and sediment biogeochemical cycles. Second, to highlight the possible interactions between climate change and management mitigation of nutrient enrichment.

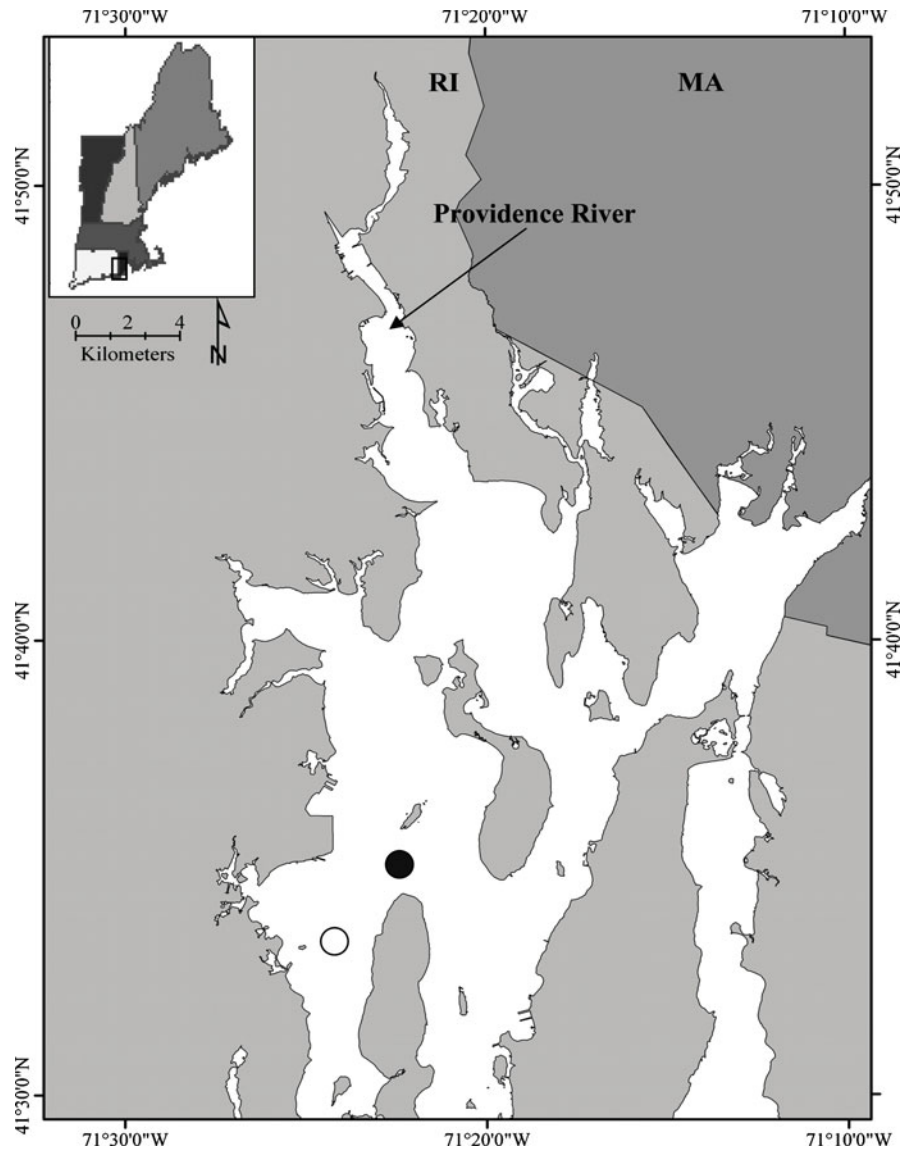
Materials and methods

Site description

Narragansett Bay, including Mount Hope Bay, is a 328-km² phytoplankton-based temperate ecosystem (latitude 41°N) with a mean depth of 8.6 m and a mean flushing rate of 26 days (Pilson, 1985; Nixon et al., 1995). Freshwater input is relatively low (100 m³ s⁻¹), with the result that the mid-bay is generally well mixed and there is only occasional weak vertical stratification (Nixon et al., 2005). Salinity follows a down-bay gradient from ~20 psu at the head to ~32 psu at the mouth of Narragansett Bay. The annual temperature varies from about 0 to 24°C.

Clayey silt and sand-silt-clay sediments comprise the majority of Narragansett Bay sediments (McMaster, 1960). For this study, sediment cores were collected from a station ~7 m deep near the middle of Narragansett Bay (Fig. 1). With an annual mean vertical light attenuation coefficient of -0.55 m^{-1} (Oviatt et al., 2002), only about 2% of surface light reaches the heterotrophic bottom. This mid-bay station has been the site of many studies ranging from in situ experiments on sediment oxygen demand and nutrient regeneration (Nixon et al., 1976) to denitrification measurements using extracted cores (Seitzinger et al., 1984) and larger mesocosms (Nowicki, 1994). Sediment from this site has also been used for various experiments (i.e., assessments of the benthic community (Rudnick et al., 1985; Frithsen, 1989; Ellis, 2002); effects of organic enrichment on the benthos (Garber, 1982; Grassle et al., 1985); benthic predation on the water column (Doering et al., 1986; Oviatt, 2004) in the large mesocosms of the marine ecosystems research laboratory (MERL). Previous study has shown that sediments at this station are predominantly silt-clay (73%) with an organic content of 4.5% (Hale, 1974). At this mid-bay station, the dominant macrofauna include *Mediomastus ambiseta*, *Nephtys incisa*, *Amplisca*, and *Nucula annulata* (Seitzinger et al., 1984;

Fig. 1 Location of the long-term sediment sampling site (*closed circle*) in mid-Narragansett Bay and the location of the long-term water quality site (*open circle*)



Rudnick et al., 1985). This site has been and is currently being monitored for benthic macro infauna and meiofauna species composition and abundance as part of the Environmental Protection Agency northeast coastal assessment program (C.A. Oviatt pers. comm., see <http://www.gso.uri.edu/merl/data.htm>).

Sediment collection and incubation

Over the annual cycle in 2005 and 2006, triplicate cores (78.5 cm² and ~15 cm deep) were collected by SCUBA divers using a round PVC corer. Cores were

maintained in the field at in situ temperatures and transported to an environmental chamber at the Graduate School of Oceanography where they were held in the dark in a water bath at field temperature. The cores were then left uncapped with oxygen bubbling gently through the overlying water overnight (8–12 h).

Throughout this study, the benthic flux measurements were conducted in the dark and separated into two incubations. During the first incubation, water samples were collected for dissolved gas analysis to measure denitrification using the N₂/Ar technique

(Kana et al., 1994; Giblin et al., 1995; Eyre et al., 2002). Then, after a 24-h period of no measurements, the second incubation for nutrient fluxes (NH_4^+ , PO_4^-) was initiated. Sediment oxygen consumption was measured during both incubations. Before each set of flux measurements, the overlying water in the cores was carefully replaced with filtered (1 μm) Narragansett Bay water. Cores were then sealed with gas-tight lids containing magnetic stirrers and the overlying water was continuously stirred (40 rpm). Initial and final samples for Winkler analysis were collected from each core and oxygen was monitored over the course of the incubation using a Hach HQ10 LDO probe. N_2/Ar samples were collected in duplicate at five times during the incubation and preserved with HgCl_2 . Dissolved inorganic nitrogen and phosphorus samples were collected at three points during the second incubation and filtered through glass fiber filters (Whatman GF/F 0.70 μm) using a 60-ml acid-washed polypropylene syringe. The filtrate was collected and stored in 60-ml acid-washed and deionized water-leached polyethylene bottles and stored at -15°C until analysis. An additional core container with no sediments was incubated as a control. When necessary, corrections were made for any changes that occurred, but the water column respiration was usually negligible.

Analytical methods

N_2/Ar was measured using MIMS (Kana et al., 1994; Giblin et al., 1995) including a Pfeiffer Quadrupole Mass Spectrometer and a flow-through silicone capillary membrane inlet (Bay Instruments, Easton, Maryland). Dissolved inorganic nitrogen and phosphorus concentrations were analyzed using a Lachat Instrument QuikChem 8000 flow injection analyzer (Table 1). Oxygen concentrations were determined using the Winkler method with manual titration.

Table 1 Analytical methods used in this study

Parameter	Method reference	Detection limit
N_2/Ar ratio	Kana et al., 1994	0.03%
Ammonium	US EPA, 1983a; Grasshoff, 1976	0.07 μM
Nitrite + Nitrate	US EPA, 1997; Grasshoff, 1976	0.02 μM
Orthophosphate	US EPA, 1983b; Grasshoff, 1976	0.01 μM
Oxygen	Lambert & Oviatt, 1983; Carrit & Carpenter, 1966	0 $\text{mg O}_2 \text{ l}^{-1}$

Data analysis

The N_2 flux for each core was determined by the rate of change in N_2 concentration (Groffman et al., 2006). For the N_2/Ar method, the change in N_2 concentration is determined from the change in N_2/Ar multiplied by the Ar concentration at air saturation (Colt, 1984). N_2 production was then determined from a five-point linear regression (Giblin et al., 1995). Sediment oxygen consumption was determined by subtracting the final concentration from the initial concentration and dividing by the total time of the incubation. Rates were then prorated for the volume and area of the core. Nutrient fluxes were initially calculated as above for oxygen. However, upon further examination, it was noticed that fluxes decreased over time, most likely because of a decrease in the concentration gradient between the sediments and the overlying water. To avoid underestimating the nutrient fluxes, the flux between the initial and middle time points of the incubation was calculated. Differences between the historical and most recently collected data and their relationship with temperature were examined using a two-way ANOVA and least-squared difference (LSD) multiple comparison test.

Results

Oxygen uptake and nutrient fluxes

Sediment oxygen consumption over the annual cycle ranged from a low of 3 $\text{mg m}^{-2} \text{ h}^{-1}$ at 6°C to a high of 27 $\text{mg m}^{-2} \text{ h}^{-1}$ at 23°C (Fig. 2). The low-temperature measurements (4 $\text{mg m}^{-2} \text{ h}^{-1}$ at 5°C) in this study were similar to the historical data, but the annual mean in the past (43 $\text{mg m}^{-2} \text{ h}^{-1}$) was much higher than what was measured here (14 $\text{mg m}^{-2} \text{ h}^{-1}$). Overall, the oxygen consumption measurements in this study showed no significant ($P < 0.01$) relationship with

Fig. 2 Sediment oxygen uptake over the annual temperature cycle for mid-Narragansett Bay sediments. Historical data (*closed circles*) from Nixon et al. (1976, 1980) and unpublished data. Most recent 2005/2006 data (*open circles*) from this study

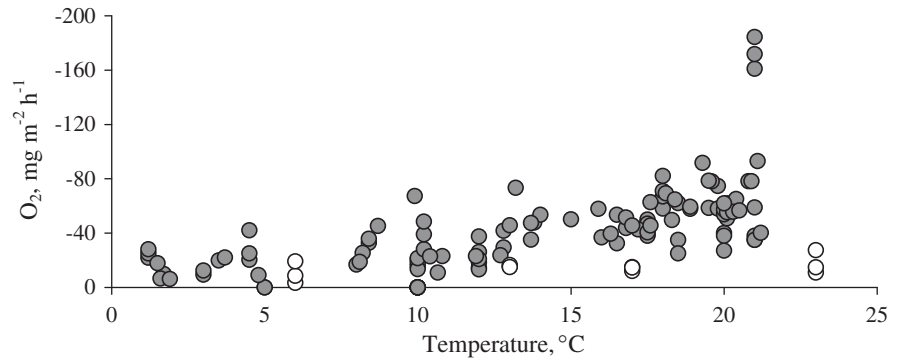


Fig. 3 Ammonia fluxes over the annual temperature cycle for mid-Narragansett Bay sediments. Historical data (*closed circles*) from Nixon et al. (1976, 1980) and unpublished data. Most recent 2005/2006 data (*open circles*) from this study

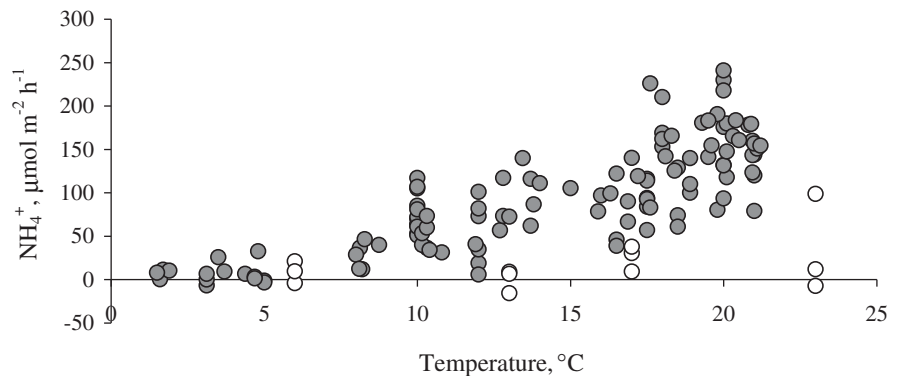
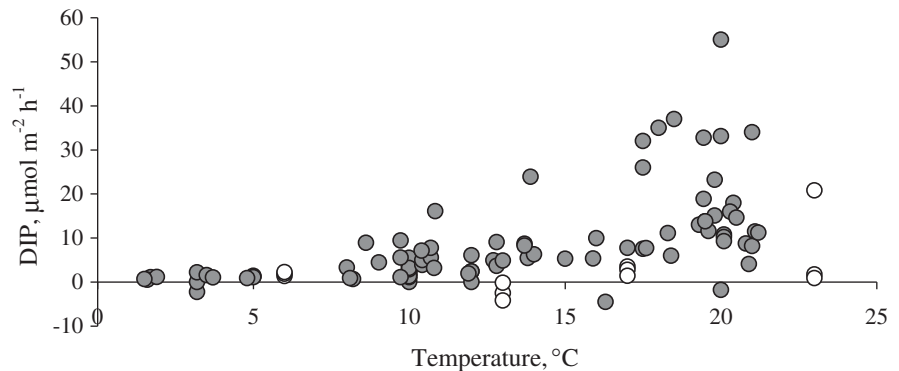


Fig. 4 Dissolved inorganic phosphate (DIP) fluxes over the annual temperature cycle for mid-Narragansett Bay sediments. Historical data (*closed circles*) from Nixon et al. (1976, 1980) and unpublished data. Most recent 2005/2006 data (*open circles*) from this study

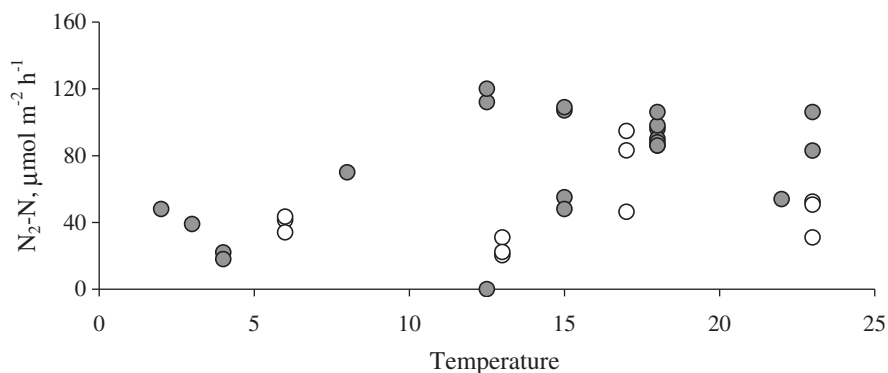


temperature. This is surprising since all of the historical data sets for mid-Narragansett Bay, including measurements made in situ, in extracted cores, and in the large (13 m^3) MERL mesocosms (Oviatt et al., 1993) showed an exponential relationship between oxygen uptake and temperature (Nixon et al., 1976). In the most recent measurements, the rates were also significantly ($P < 0.01$) lower than the historical rates.

The measurements of ammonium flux over the annual cycle in this study also showed small uptakes and releases (Fig. 3). With an influx of $-15 \mu\text{mol}$

$\text{m}^{-2} \text{h}^{-1}$ and an efflux of $98 \mu\text{mol} \text{m}^{-2} \text{h}^{-1}$, the mean ammonium flux ($14 \mu\text{mol} \text{m}^{-2} \text{h}^{-1}$) was significantly ($P < 0.01$) lower than the historical mean ($94 \mu\text{mol} \text{m}^{-2} \text{h}^{-1}$) (Nixon et al., 1980). Previously, ammonium regeneration has exhibited a marked increase with higher temperatures, but this no longer appears to be true. Dissolved inorganic phosphate (DIP) flux ranged from -4 to $21 \mu\text{mol} \text{m}^{-2} \text{h}^{-1}$ with a mean of $2.5 \mu\text{mol} \text{m}^{-2} \text{h}^{-1}$ (Fig. 4). This range is significantly ($P < 0.01$) lower than the historical range (-5 to $55 \mu\text{mol} \text{m}^{-2} \text{h}^{-1}$) and mean ($8 \mu\text{mol} \text{m}^{-2} \text{h}^{-1}$).

Fig. 5 Denitrification over the annual temperature cycle for mid-Narragansett Bay sediments. Historical data (*closed circles*) from Seitzinger et al. (1984) and Nowicki (1994). Most recent 2005/2006 data (*open circles*) from this study



Remarkably, as seen with oxygen and ammonia, DIP fluxes also showed no increase with temperature as they had in the past.

Denitrification rates

Using the N₂/Ar technique, we found denitrification rates ranging from 22 to 100 μmol N₂-N m⁻² h⁻¹. Seitzinger et al. (1984) reported seasonal denitrification rates measured for 1979 at the mid-bay station that ranged from 40 to 109 μmol N₂-N m⁻² h⁻¹. Additional denitrification measurements made by Nowicki (1994) using sediment from this site and a slightly modified technique varied from 18 to 120 μmol N₂-N m⁻² h⁻¹. While the rates in this study are similar to those measured previously, the mean annual rate has significantly ($P < 0.01$) declined from 74 μmol N₂-N m⁻² h⁻¹ in 1979 to 40 μmol N₂-N m⁻² h⁻¹ in 2005/2006. Neither the present nor the historic denitrification rates exhibit a positive relationship with increasing temperature (Fig. 5).

Discussion

The analytical techniques used to measure the concentrations of oxygen, ammonia, and phosphate have not changed substantially between the earlier work and the current study and therefore analytical changes do not confound these comparisons. Unfortunately, it is possible that the denitrification decline may be an artifact of changing methodology. No direct comparisons between the N₂ technique, used by Seitzinger et al. (1984) and Nowicki (1994), and the N₂/Ar technique have yet been reported. However, the oxygen and nutrient regeneration measurements

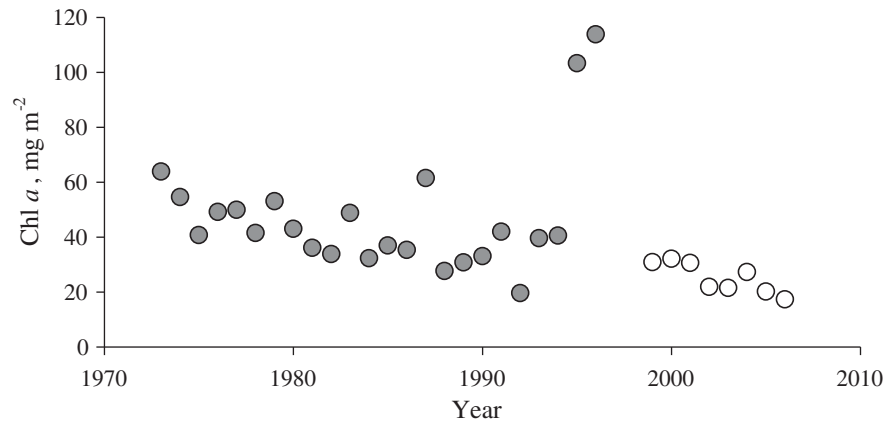
reported previously are not confounded by changing analytical techniques, and their large declines suggest that the decline in denitrification is also real.

Ecological responses to climate change in Narragansett Bay

Narragansett Bay has traditionally been characterized by a strong winter–spring diatom bloom that usually begins in mid-November and reaches peaked concentrations at the end of December or early January (Pratt, 1959). However, over the last three decades the timing and magnitude of the phytoplankton bloom in Narragansett Bay have decreased dramatically. Long-term monitoring of the phytoplankton at a station near the middle of the bay (Fig. 1) shows that mean annual water-column chlorophyll concentrations have decreased by almost 70% since 1970 (Fig. 6). Most of the decline in the annual means is associated with changes in the traditional winter–spring bloom which has arrived later or not at all (Oviatt et al., 2002). The abundance of diatoms during the bloom has also declined (Karentz & Smayda, 1998; Li & Smayda, 1998). Since chlorophyll is tightly coupled with the rate of ¹⁴C uptake in this system (Keller & Riebesell, 1989) as well as in many others (see recent review by Brush et al., 2002), this decline has almost certainly been associated with a marked decline in the rate of primary production.

The cause of the oligotrophication of Narragansett Bay is not well understood, but it cannot be attributed to a reduction in nitrogen inputs as they have remained essentially unchanged for at least 25 years (Nixon et al., 2008). There is some correlative evidence that climate variability (North Atlantic Oscillation) and long-term warming trends may be

Fig. 6 Mean annual water column chlorophyll *a* concentrations at the mid-bay station (see Fig. 1). Historical data (*closed circles*) from Li & Smayda (1998). Most recent data (*open circles*) from the Graduate School of Oceanography/URI (see <http://www.gso.uri.edu/phytoplankton>)



responsible (Hawk, 1998; Borkman, 2002; Oviatt et al., 2002; Oviatt, 2004). Winter warming in the bay is highly significant. Warming has amounted to an increase of about 1.7°C during the 1990s compared with the period between 1890 and 1970 at nearby Woods Hole, MA, where it is known that water temperature is closely correlated with that in Narragansett Bay (Nixon et al., 2003, 2004). Townsend & Cammen (1988) found that year-to-year variation in the spring bloom in the Gulf of Maine was controlled by light and not temperature. Working in Narragansett Bay, however, Borkman (2002) found that warmer winters tended to be cloudier, thus making it difficult to separate temperature and light effects. Experiments with the large MERL mesocosms have shown that slightly warmer temperatures reduce or eliminate winter–spring phytoplankton blooms in Narragansett Bay, presumably because of higher grazing pressure (Keller et al., 1999). Field evidence also confirms that warmer winters are associated with a delayed or complete elimination of the winter–spring bloom (Oviatt et al., 2002).

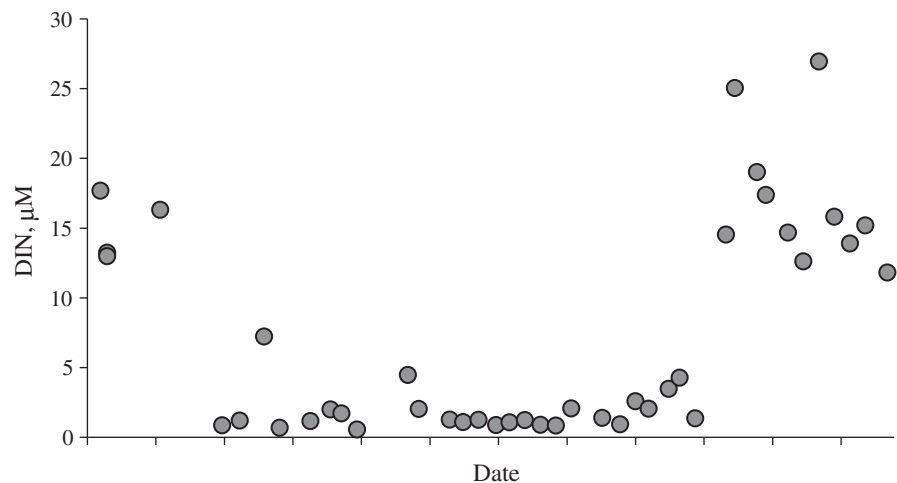
Regardless of the reason, the loss of the winter–spring bloom is particularly alarming for a benthic community that is already food-limited by the end of summer. Tracer and mass balance studies, also carried out in the MERL mesocosms, have shown that a smaller fraction of the summer blooms is deposited on the bottom compared with winter–spring blooms (Rudnick & Oviatt, 1986; Keller et al., 1999). From the field (Rudnick et al., 1985; Craig, 1989; Ellis, 2002), from mesocosm study (Grassle et al., 1985; Beatty, 1991), and from a numerical model (McKenna 1987), there is considerable evidence that certain groups of the meio- and macro-benthic infauna have

historically been food-limited during summer in Narragansett Bay.

Implications for management

There is a common impression among environmental groups, some managers, and some scientists that nuisance blooms of macroalgae (especially *Ulva* sp.) are increasing and, more importantly, that bottom waters in upper Narragansett Bay and the Providence River (actually an estuary at the head of the bay, Fig. 1) are developing more frequent, intense, and widespread hypoxia (e.g., Rhode Island Department of Environmental Management, 2003). Unfortunately, historical data are not really available to address the question of change in either of these conditions. Because of this, many of the major wastewater treatment plants that discharge directly to the bay are in the process of upgrading to tertiary treatment, thus reducing total nitrogen loading to the bay. Nixon et al. (2005) estimated that by 2010 summer nitrogen inputs to Narragansett Bay may decrease between 30% and 40%. Clearly, these reductions will have significant impacts on the new production of the bay because ¹⁴C uptake and bloom formation are strongly limited by nitrogen availability during summer (Furnas et al., 1976; Oviatt et al., 1995), when concentrations of dissolved inorganic nitrogen are very low or undetectable in the surface water of the mid- and lower bay (Fig. 7). It is anticipated that with less phytoplankton production there will be less organic matter deposition to the sediments, less benthic respiration, and therefore fewer hypoxic/anoxic events. However, we have shown here that there have already been large

Fig. 7 Annual surface water-column dissolved inorganic nitrate (DIN) concentrations at the mid-bay station for 2005. Data available at <http://www.gso.uri.edu/phytoplankton/>



reductions in pelagic chlorophyll in the mid-bay and that respiration and nutrient regeneration by the benthos of mid-Narragansett Bay have declined greatly over the past three decades. Others have shown a positive effect of nitrogen loading on benthic biomass (Josefson & Rasmussen, 2000) and a positive relationship between primary and secondary production (Nixon et al., 1986). Since climate change has already begun the oligotrophication of a once highly productive bay, it is possible that reductions of nitrogen input during the growing season will have a significant negative impact on secondary production throughout the mid- and lower bay, a much larger area than the regions currently exposed to frequent hypoxia.

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