# **Chapter 8 Flux-Based O3 Risk Assessment for Japanese Temperate Forests**

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**Abstract** Ground-level ozone  $(O_3)$  levels are expected to increase over the twenty-first century, particularly in the region of East Asia. We performed an  $O_3$ flux-based risk assessment of C sequestering capacity in an old cool temperate deciduous forest, consisting of O3-sensitive Japanese beech (*Fagus crenata*), and in a warm temperate deciduous and evergreen forest dominated by  $O<sub>3</sub>$ -tolerant Konara oak (*Quercus serrata*), based on long-term CO2 flux observations. Light-saturated

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gross primary production, as a measure of C sequestering capacity, declined earlier in the late-growth season with increasing cumulative  $O_3$  uptake, suggesting an earlier autumn senescence in the O<sub>3</sub>-sensitive beech forest, but not in the O<sub>3</sub>-tolerant oak forest.

**Keywords** C sequestration •  $O_3$  flux •  $O_3$ -sensitive forest •  $O_3$ -tolerant forest • Deciduous temperate forests

### **8.1 Introduction**

In Japan, which is located on the edge of East Asia, a continuous increase in  $O<sub>3</sub>$ concentration has been observed since the latter half of the 1980s, partly due to pollutants advected from foreign sources, particularly in East Asia (Richter et al. [2005;](#page-8-0) Ohara et al. [2008;](#page-8-1) Nagashima et al. [2003;](#page-8-2) Akimoto [2003\)](#page-7-0). Since 2000, the flux tower sites of the Forestry and Forest Products Research Institute (FFPRI) have monitored  $CO<sub>2</sub>$ , energy, and water vapor fluxes in several Japanese forests with different tree species. Although  $O_3$  is known to be a detrimental air pollutant for trees, as it reduces the photosynthetic rate, increases the respiration rate, and accelerates leaf senescence (Matyssek et al. [2010](#page-8-3)), few studies have been conducted to assess the  $O_3$  risk in forests directly. Thus, we have a unique chance to assess the  $O_3$  risk in Japanese forests based on the long-term  $CO<sub>2</sub>$  flux observations.

 $O_3$  flux-based risk assessment is more essential than  $O_3$  exposure-based risk assessment for evaluating the physiological effects of  $O_3$  on plants (Emberson et al.  $2000$ ; Matyssek et al.  $2007$ ). To estimate  $O_3$  fluxes at the forest level, canopy-level stomatal conductance is required, as well as  $O<sub>3</sub>$  concentration over the forest, which is generally estimated by the Penman–Monteith (P–M) equation, based on energy and water flux over a forest (Gerosa et al. [2005](#page-7-2); Monteith [1981](#page-8-5)). However, the use of this approach is valid only when the entire evaporation process in the canopy takes place through stomatal transpiration. In this context, it is difficult to conduct a continuous estimation of the canopy-level stomatal conductance using the P–M approach in temperate deciduous forests, which have regular rainfall and a period when the canopy is not closed, in spring and autumn (Biftu and Gan [2000\)](#page-7-3).

To solve this problem, we have developed a novel approach (Kitao et al. [2014](#page-8-6)) combining the P−M approach with a semi-empirical photosynthesis-dependent stomatal model (Ball–Woodrow–Berry [BWB] model; Ball et al. [1987\)](#page-7-4), where photosynthesis, relative humidity, and  $CO<sub>2</sub>$  concentration are assumed to regulate stomatal conductance. Based on the BWB model, leaf-level stomatal conductance  $(g_s)$  is estimated as:

$$
g_s = g_{\min} + a A_g \text{rh} / C_s \tag{8.1}
$$

where  $g_{\text{min}}$  denotes the minimum conductance in the dark,  $a$  is an empirical scaling parameter,  $A_g$  the gross photosynthetic rate, rh the relative humidity, and  $C_s$  the leaf surface  $CO<sub>2</sub>$  concentration.

In the present study, we applied the modified BWB model to estimate canopylevel stomatal conductance  $(G_s)$  taking into account the non-linear response of canopy stomatal conductance to relative humidity (Fares et al. [2013](#page-7-5)) as:

$$
G_{\rm s} = G_{\rm min} + ab^{\rm th} \text{GPP} / C_{\rm s} \tag{8.2}
$$

<span id="page-2-0"></span>where  $G<sub>min</sub>$  denotes the canopy-level minimum conductance in the dark, *a* and *b* are empirical scaling parameters, and GPP is the gross primary production (= net ecosystem exchange [NEE] – respiration of ecosystem  $[R_{\text{eco}}]$ ). We first determined the coefficients in the equation, using  $G_s$  derived by the  $(P-M)$  equation when the canopy was closed from June to August for the beech forest, and  $G_s$  derived by the (P–M) equation from June to September for the oak forest without rain (>1 mm within 24 h) for each year. Then we estimated the stomatal conductance over the canopy of temperate deciduous forests continuously, using this novel approach combining the P–M approach with a photosynthesis-dependent stomatal model (Kitao et al. [2014\)](#page-8-6).

Based on the above approach, we performed flux-based assessments of  $O_3$  effects on photosynthetic  $CO<sub>2</sub>$  uptake in a cool temperate deciduous forest, consisting of O3-sensitive deciduous broadleaf tree species, Japanese beech (*Fagus crenata* Blume; 70–80 years old), and a warm temperate mixed deciduous and evergreen broadleaf forest, dominated by  $O_3$ -tolerant deciduous broadleaf tree species, Konara oak (*Quercus serrata* Thunb. ex. Murray; approximately 30 years old) (Yasuda et al. [2012;](#page-8-7) Kominami et al. [2012](#page-8-8); Yamaguchi et al. [2011;](#page-8-9) Kitao et al. [2015\)](#page-8-10).

### **8.2 Study Sites**

The Appi forest meteorology research site  $(40^{\circ} 00' N, 140^{\circ} 56' E, 825 m$  above sea level) is located on the Appi highland in Iwate Prefecture, Japan (details of which are described in Yasuda et al. [2012](#page-8-7)) (Fig. [8.1\)](#page-3-0). The site is located in a secondary cool temperate deciduous broadleaf forest primarily consisting of the Japanese beech (*Fagus crenata* Blume), which was approximately 80 years old. The canopy height was measured to be 19–20 m in 2009. There is not much vegetation on the forest floor, and evergreen trees are rarely observed. It snows heavily from November to May, with the snow depth reaching 2 m. The annual mean temperature was 5.9 °C (in 2000–2006), annual precipitation was 1,869 mm (in 2007–2009), and annual mean solar radiation was 12.7 MJ m<sup>-2</sup> day<sup>-1</sup> (in 2000–2006). The soil was classified as moderately moist brown forest soil.

The Yamashiro forest hydrology research site (34° 47′ N, 135° 50′ E, 220 m above sea level) is situated in the southern part of Kyoto Prefecture, Japan (details of which are described in Kominami et al. [2012\)](#page-8-8) (Fig. [8.2](#page-3-1)). The site is located in a warm temperate mixed deciduous and evergreen broadleaf forest, which is built upon weathered granite. After an invasion by pine wilt disease in the 1980s, Konara oak (*Quercus serrata* Thunb. ex. Murray) has taken over and the forest is now

<span id="page-3-0"></span>

**Fig. 8.1** Cool temperate deciduous forest consisting of Japanese beech at the Appi tower site

<span id="page-3-1"></span>

**Fig. 8.2** Warm temperate deciduous and evergreen mixed forest, predominantly consisting of Konara oak, at the Yamashiro tower site

regenerated. The tree biomass (diameter at breast height [DBH]  $\geq$ 3 cm) was estimated at 51 Mg C ha<sup>-1</sup> in 1999, dominated by Konara oak, classified as a deciduous broadleaf tree species (66% of biomass), and *Ilex pedunculosa* Miq. (an evergreen broadleaf tree species; 28% of biomass) (Goto et al. [2003\)](#page-8-11). The canopy height ranged from 6 to 20 m with an average of 12 m. The annual mean temperature was 14.7 °C (in 2000–2002), annual precipitation was 1,095 mm (in 2000–2002), and annual mean solar radiation was 11.9 MJ m<sup>-2</sup> day<sup>-1</sup> (in 2000–2002).

# **8.3 Estimation of Light-Saturated GPP**

Light-saturated GPP was derived from the relationship between GPP and photosynthetic photon flux density (PPFD). The data were rejected when the friction velocity (u\*) was below 0.25 m s−<sup>1</sup> , but in cases of precipitation, the data were included. We regressed the relation between GPP and PPFD as follows:

<span id="page-4-0"></span>

$$
GPP = \alpha GPP_{\text{max}} PPPFD / (GPP_{\text{max}} + \alpha PPPFD)
$$
 (8.3)

where  $\alpha$  denotes the ecosystem quantum yield and GPP<sub>max</sub> the maximum GPP. We derived  $\alpha$  and GPP<sub>max</sub> for the pooled data at 2-week intervals from the onset of budbreak. We set the GPP at a PPFD of 1,500 µmol  $m^{-2} s^{-1}$  based on the equation as the light-saturated GPP (Fig. [8.3](#page-4-0)).

The maximum light-saturated GPP of the beech forest varied among the years, but the environmental factors determining the inter-annual variation have not yet been fully identified (Yasuda et al. [2012](#page-8-7)). A survey of each individual tree suggested that the oak forest was still growing, as the total biomass increased from 2004 to 2009 (Kominami et al. [2012](#page-8-8)). Therefore, to investigate the  $O_3$  effects on the seasonal changes in foliar photosynthetic maturation and senescence, we used a relative unit of GPP (GPP\_rel), which is calculated as:

> $GPP$   $_{-}$  rel  $=$  (light  $-$  saturated  $GPP$ ) /(the maximum light  $-$  saturated GPP during the growth period for each year)

The light-saturated GPP was estimated as the GPP at PPFD of 1,500 µmol  $m^{-2} s^{-1}$ based on the GPP light-response curves derived from the pooled data for 2-week intervals from the budbreak. We categorized the growth season as spring–summer (April or May to July) for leaf maturation and summer–autumn (August to October or November) for leaf senescence stages.

### **8.4 Estimation of Cumulative O<sub>3</sub> Uptake (COU)**

The approach used to estimate stomatal  $O<sub>3</sub>$  fluxes involves several steps, according to Cieslik ([2004](#page-7-6)) and Gerosa et al. [\(2003\)](#page-7-7). The aerodynamic resistance  $(R_a)$  is calculated from measured micrometeorological parameters such as friction velocity and sensible

heat flux by using the Monin–Obukhov similarity theory (see Gerosa et al. [2003,](#page-7-7) for calculation details), while the quasi-laminar layer resistance  $(R_b)$  is calculated by using the parameterization proposed by Hicks et al. ([1987](#page-8-12)). We calculated the surface resistance  $(R_c)$  from the stomatal  $(R_{ST})$  and non-stomatal resistance to  $O_3(R_{NS})$  as follows:

$$
R_{\rm c} = 1 / \left( 1 / \, R_{\rm ST} + 1 / \, R_{\rm NS} \right) \tag{8.4}
$$

$$
R_{\rm NS} = 1 / \left( LAI / r_{\rm ext} + 1 / \left( R_{\rm inc} + R_{\rm gs} \right) \right) \tag{8.5}
$$

where LAI is the leaf area index  $(m^2 m^{-2})$ ,  $r_{ext}$  denotes the external leaf resistance,  $R_{\text{inc}}$  the in-canopy resistance, and  $R_{\text{gs}}$  the ground surface resistance. The maximum LAI in the beech forest was estimated from the amounts of leaf litter (Yasuda et al. [2012\)](#page-8-7). Based on the field observations, we assumed that the LAI of the beech forest increased from 0 to the maximum within 1 month from the budbreak, and then decreased from the maximum to 0, also within 1 month, before the end of the foliage period. In contrast, seasonal changes in the LAI in the oak forest were measured using a plant canopy analyzer (LAI-2000; Li-Cor, Lincoln, NE, USA). The external leaf resistance ( $r_{\text{ext}}$ ) is set at 2,500 s m<sup>-1</sup>, the in-canopy resistance ( $R_{\text{inc}}$ ) is defined as b LAI h/u\*, where h is the canopy height and b an empirical constant taken as 14 m<sup>-1</sup>, and *R*<sub>gs</sub> is set at 200 s m<sup>-1</sup> (Erisman et al. [1994](#page-7-8); Simpson et al. [2012\)](#page-8-13).

The stomatal  $O_3$  flux ( $F_{ST}$ ) is obtained as follows:

$$
F_{\rm ST} = C_{\rm m} * R_{\rm c} / \left[ \left( R_{\rm a} + R_{\rm b} + R_{\rm c} \right) R_{\rm ST} \right] \tag{8.6}
$$

where  $C<sub>m</sub>$  denotes the ozone concentration at the measurement height. The stomatal resistance  $(R_{ST})$  is calculated by multiplying the stomatal resistance for water vapor flux (noted as  $R_s$ ) by 1.65 (Gerosa et al. [2003](#page-7-7)).  $R_s$  is expressed as  $1/G_s$ , where  $G_s$  is determined from the modified BWB model described above (Eq. [8.2](#page-2-0)). We estimated  $C<sub>m</sub>$  at the flux sites using the  $O<sub>3</sub>$  concentration data monitored by the nearest air pollution stations (Komatsu et al. [2015](#page-8-14)).

The COU in the forest ecosystem was calculated by summing up the ozone stomatal flux  $(F_{ST})$  during the daytime (PPFD >0) from the budbreak to a given period as follows:

$$
COU = \sum F_{ST} \Delta t,\tag{8.7}
$$

where the mean daily sum of  $F_{ST}$  for the days of each month when *F*st data were available was substituted for the rest days when  $F_{ST}$  data were unavailable.

# **8.5 Effects of Cumulative O<sub>3</sub> Uptake (COU) on C Sequestering Capacity in the Beech and Oak Forests**

To quantitatively evaluate the influence of the major explanatory factor(s) on the photosynthetic capacity, multiple regression analysis was used. We initially set four explanatory factors affecting potential photosynthetic performance

<span id="page-6-0"></span>

(GPP\_rel): leaf age, photoperiod, air temperature, and COU (Bauerle et al. [2012\)](#page-7-9). GPP\_rel in the spring–summer period could be explained by three explanatory factors, i.e., leaf age, photoperiod, and COU, in both the beech and oak forests. Increased COU, in addition to the photoperiod and leaf age, showed a positive effect on leaf maturation (Fig. [8.4a, b](#page-6-0)). As young leaves generally have less sensitivity to  $O_3$  than old leaves (Kitao et al. [2015;](#page-8-10) Bohler et al. [2010\)](#page-7-10), a stimulating effect of  $O_3$  on photosynthesis at the early stage of leaf development needs to be further investigated. In contrast, GPP\_rel in the summer–autumn period in the beech forest could be explained by two environmental factors, i.e., leaf age and COU. Thus, increased COU may accelerate age-dependent leaf senescence in the  $O<sub>3</sub>$ -sensitive beech forest (Fig. [8.4a\)](#page-6-0), whereas leaf senescence was not influenced by COU but was primarily influenced by air temperature in the  $O<sub>3</sub>$ -tolerant oak forest (Fig. [8.4b\)](#page-6-0). The lower  $O_3$  sensitivity during leaf senescence in the oak forest may be attributed to the lower COU (due to the lower  $G_s$ ) than in the beech forest (data not shown), as reported for the seedlings of Konara oak grown under free-air  $O<sub>3</sub>$  fumigation (Kitao et al. [2015\)](#page-8-10).

## **8.6 Conclusion**

In summary, our findings, based on long-term  $CO<sub>2</sub>$  flux observations, indicate that the photosynthetic C sequestering capacity of the forest ecosystem in an  $O_3$ -sensitive forest is potentially affected by the present-level  $O_3$ .  $O_3$ -induced earlier leaf senescence has been reported in several previous studies using open-top chambers and free-air fumigation systems (Karnosky et al. [2005;](#page-8-16) Kitao et al. [2009](#page-8-17); Yamaji et al. [2003;](#page-8-18) Calatayud et al. [2011\)](#page-7-11). In the present study, it is noteworthy that such accelerated leaf senescence induced by  $O_3$  was apparently detected in the  $O_3$ -sensitive beech species at the real-world forest level. As the atmospheric  $O_3$  concentrations are predicted to increase, particularly in East Asia, including Japan (Akimoto [2003;](#page-7-0) Karnosky et al. [2005](#page-8-16); Ashmore [2005](#page-7-12); Koike et al. [2013](#page-8-19)), earlier leaf senescence induced by elevated  $O_3$  could cause further adverse effects on forest C sequestering in the future, particularly in forests consisting of  $O_3$ -sensitive species.

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