Introduction

Planktic foraminifers are marine protozoans with calcareous shells and chambered tests (Plate 1.1), first appearing in the mid-Jurassic approximately 170 million years ago, and populating the global ocean since the mid-Cretaceous (cf. Frerichs et al. 1972; Caron and Homewood 1983). The scientific and economic value of planktic foraminifers is based on their global marine abundance since the Lower Cretaceous ~ 110 Million years ago. Owing to the high preservation potential of their calcareous shell, planktic foraminifers provide information on the past environment and climate. Physical conditions and chemical composition of ambient seawater are reconstructed from faunal assemblages, i.e. the presence or absence of foraminifer species, as well as through the chemical composition of their test calcite, including crystallinity of the test wall, and changes in stable isotope and element ratios.

Test: The foraminifer shell is called a test. Shell and test are often used synonymously. Shell may be used for part(s) of the test, and for fragments of the test.

Planktic—planktonic: Planktic and planktonic may be used synonymously. In the strict Greek meaning the word planktic is possibly correct (Burckhardt 1920; Rodhe 1974). In the international literature both planktic and planktonic are used to the same degree, and either term may be applied based

on personal preference. In benthic foraminifers, the term benthic has largely been used over the past decades, and benthonic has been out of fashion for some time.

Modern planktic foraminifers evolved from the earliest Tertiary including the first spinose species in Earth history soon after the Cretaceous-Paleogene (K/Pg) boundary (Olsson et al. 1999). Most modern species live in the surface to thermocline layer of the open ocean, and in deep marginal seas as the Mediterranean, Caribbean, South China Sea, and Red Sea, Some species descend to waters as deep as several thousand meters in the tropical to temperate ocean. Planktic foraminifers are largely absent from shallow marginal seas, for example the North Sea where reproduction is impeded. The presence and absence of planktic foraminifer species at the regional scale is related to the quality and quantity of food, physical and chemical properties of ambient seawater, and displays an overall latitudinal pattern at the global scale.

Species abundance varies according to seasons as well as on an interannual scale, and on longer time-scales depending on environmental conditions, and affected by climate change. Symbiont-bearing species depend on light and are restricted to the euphotic zone of the surface ocean. Symbiont-barren species may dwell as

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Plate 1.1 The modern planktic foraminifer species *Orbulina universa* seen in transmitted light. The inner trochospiral test of the pre-adult individual is surrounded by the spherical adult test. Spines are protruding from both inner trochospiral and outer spherical test. Pores are

deep as the abyssal ocean, and have been sampled from below 4000 m water depth. Planktic foraminifers are rather marginal to marine biological research including modern biogeochemistry (Sarmiento and Gruber 2006), although they are major producers of marine calcareous particles (i.e. their tests) deposited on the ocean floor forming the globigerina ooze (e.g., Vincent and Berger 1981). Data compilation of a large visible as tiny dark spots on the inner and outer test. Multiple small *circles* on the outer test wall are the apertures of the adult individual. The opening at the inner trochospiral test is caused by dissolution. Scale bar 200 μ m

variety of marine Plankton Functional Types (see text box below) have shown that planktic foraminifers possibly constitute a minor but ubiquitous component of marine planktic biomass (Buitenhuis et al. 2013). In addition, modeling approaches on the planktic foraminifer population dynamics from the 1990s have contributed to a better understanding of planktic foraminifer ecology and application in paleoceanography (e.g., Signes et al. 1993; Žarić et al. 2006; Fraile et al. 2009; Lombard et al. 2011; Roy et al. 2015).

Plankton functional type (PFT): The plankton functional expression type (PFT) is used in modeling, and includes different conceptual categories of organisms as, for example, organisms of similar ecology, and serving similar roles within an ecosystem (Anderson 2005). The PFTs included in the MAREDAT initiative on the ecology and biomass of marine plankton are picophytoplankton, diazotrophs, coccolithophores, Phaeocystis, diatoms, microzooplankton, picoheterotrophs, foraminifers planktic (which range between micro- and mesozooplankton), mesozooplankton, pteropods, and macrozooplankton (Buitenhuis et al. 2013).

By contributing substantially to the fossil record of marine sediments, planktic foraminifers provide indispensable ecologic information used in paleoecologic, paleoceanographic, and stratigraphic research from the Lower Cretaceous (~110 millions years, Ma). Faunistic and biogeochemical (e.g., stable isotopes) information from the calcareous (calcite, CaCO₃) planktic foraminifer tests is used to reconstruct, for example, temperature and salinity of the past surface ocean. Radiocarbon (¹⁴C) gives an absolute age of test formation of late Pleistocene and Holocene sediments. Factors determining the modern faunal composition are applied to the interpretation of the fossil assemblages, for example, by multiple regression techniques (i.e. transfer functions), yielding information (proxy data) on ancient environmental parameters. The chemical composition, i.e. stable isotope and element ratios of the calcareous test (calcite, CaCO₃) provides an assessment of the chemical and physical state of ambient seawater, and is applied to the reconstruction of temperature, and biological productivity of the past marine environment.

Proxy (pl. proxies): A proxy is a measurable feature from which another not directly measurable characteristic can be derived. For example, the test of a planktic foraminifer bears certain stable isotope ratios (e.g., ^{18/16}O), measurable with a mass spectrometer, from which temperature and other parameters of ambient seawater can be reconstructed by applying empirically derived formulae (see, e.g., Fischer and Wefer 1999).

1.1 A Brief History of Planktic Foraminifer Research

Technological improvement of binocular microscopes allowed the French naturalist Alcide d'Orbigny (1826) to describe the first planktic foraminifer species Globigerina bulloides from beach sands of Cuba, but erroneously classifying it with the cephalopods. Alcide d'Orbigny's family lived in the village of Esnandes at the Baie d'Aiguillon north of La Rochelle (France), where Alcide's father Charles Marie d'Orbigny was a renowned 'naturaliste'. Young d'Orbigny was fortunate enough to look at the sediments of the bay, and to find at a rich benthic foraminifer fauna using the first good binocular microscopes available in the 1820s (Vénec-Peyré 2005). D'Orbigny's French contemporary Félix Dujardin (1835), then, correctly described planktic foraminifers as unicellular organisms. Some 30 years later, Owen (1867) suspected the planktic life habit of these organisms. Following the Challenger Expedition from 1872 to 1876, the surface-dwelling habitat planktic of foraminifers was generally recognized thanks to observations provided by John Murray in the Challenger Reports (Brady 1884). Foraminifer biology was described first by Rhumbler (1911). In the first half of the 20th century, foraminifers were widely used for stratigraphic purposes in the search for hydrocarbon reservoirs, and Joseph

Cushman published a plethora of catalogues on foraminifers of all major ocean basins, and from various time-slices (e.g., Cushman 1911; Cushman and Todd 1949).

Distribution and ecology of different living planktic foraminifer species were first studied on plankton samples by Schott (1935). From the 1960s, planktic foraminifers have been used in biostratigraphy to date marine sediments sampled, for example, within the Deep Sea Drilling Programme (DSDP) from 1964 to 1983, followed by the Ocean Drilling Programme (ODP), and the Integrated Ocean Drilling Programme (IODP) from 2003 onward. The taxonomy of modern planktic foraminifers was largely improved by the seminal publication of Frances Parker (1962).

Distribution, ecology, and biology of the live fauna mostly of the western North Atlantic were extensively studied by Bé, Hemleben, Anderson, and co-workers, including graduate students and post-doctoral appointees, between the late 1950s and 1980s. Among these participants were David Caron and Howard Spero who became significant researchers in the field. Other major contributors included Peter Wiebe, Sharon Smith, Susumu Honjo, and Richard Fairbanks at Woods Hole Oceanographic Institution. At about the same time, Esteban Boltovskoy developed new sampling methods, and conducted projects on the production and sedimentation of planktic foraminifers in the South Atlantic. Ecological significance of modern species was applied to paleoecological and paleoceanographic settings to obtain new information on the ancient ocean and Earths' climate. Since the late 1960s, Wolfgang Berger and co-workers supplied ample information in many papers on planktic foraminifer carbonate chemistry and application of proxies to paleoceanography, starting in the eastern north Pacific, and later focusing on the South Atlantic (e.g., Berger 1981; Berger et al. 1989; Kemle-von-Mücke and Hemleben 1999; see also Fischer and Wefer 1999). Population dynamics and carbon turnover of modern planktic foraminifers mostly of the eastern North Atlantic and Indian Ocean including adjacent regions were studied by Christoph Hemleben and co-workers since the late 1960s (e.g., Hemleben 1969; Hemleben and Spindler 1983; Hemleben et al. 1989; Bijma and Hemleben 1994; Schiebel et al. 1995; Schiebel 2002).

In the early 1970s, a joint group guided by O. Roger Anderson, Allan Bé (both Lamont-Doherty Earth Observatory), Christoph Hemleben, and Michael Spindler (both Tübingen University), came together at the Bermuda Biological Station (BBS) in order to culture planktic foraminifers (e.g., Bé et al. 1977; Hemleben et al. 1989). The BBS is close to blue water locations and thus exceptionally suited to experiment with planktic foraminifers. Living foraminifers were sampled by means of SCUBA collection and net tow sampling, and a sophisticated experimental set up in order to maintain viable planktic foraminifers from early ontogenetic stages to maturity was developed. Almost the entire range of all basic planktic foraminifer behavior was observed and recorded. Analyses of planktic foraminifers from laboratory culture have been substantially advanced by Howard Spero and co-workers at the University of California (e.g., Spero 1986; Spero et al. 2015). Culturing of planktic foraminifers also has been conducted at the Bellairs Research Institute at Barbados (e.g., Caron et al. 1982; Spindler et al. 1984), the Caribbean Marine Research Center on Lee Stocking Island, Bahamas (e.g., Spero and Williams 1988; Spero and Lea 1993), the H. Steinitz Marine Biology Laboratory at Eilat, Gulf of Aquaba (e.g., Erez et al. 1991, and references therein), the Caribbean Marine Biological Institute (CARAMBI) at Curacao (e.g., Bijma et al. 1992), the Isla Magueyes Marine Laboratory at Puerto Rico (e.g., Hönisch et al. 2011; Allen et al. 2011, 2012). However, a second generation of any planktic foraminifer species has never been successfully achieved in laboratory culture, which remains one of the major issues to be solved in the future.

Recent work focuses on planktic foraminifer taxonomy, stratigraphy, evolution, ecology, carbonate chemistry, paleoceanography, population dynamics, and biology. Stratigraphy and paleoceanography were among the original scientific interests in planktic foraminifers, due to their economic and scientific value, respectively.



Fig. 1.1 Phylogenetic relationships of the four major groups of modern planktic foraminifers, macroperforate spinose, macroperforate non-spinose, microperforate

Modern techniques of molecular genetics (i.e. DNA sequencing) are currently applied to reveal the taxonomic and phylogenetic relations (Fig. 1.1) of the earlier established morphospecies (Table 1.1) distinguished by their test architecture (e.g., Darling et al. 1997; de Vargas et al. 1999; André et al. 2014). The relation to morphological features of the tests of modern species is reviewed in the fossil species (e.g., Hemleben et al. 1999; Hemleben and Olsson 2006).

Technological development of mass spectrometry analytical systems provides ever more precise measurements of rare elements, stable isotope ratios and 'clumped isotopes'. Based on these advances, new proxies have been developed in paleoceanography (see the review of Katz et al. 2010). Laser ablation-inductively coupled plasma-mass spectrometry (LA-ICP-MS) and secondary ion mass spectrometry (NanoSIMS) allow analyses of single chambers of tests, and hence better interpretation of ontogenetic changes

spinose, and Hastigerinidae, based on a maximum likelihood reconstruction from SSU rDNA. Modified after Aurahs et al. (2009), from Weiner et al. (2012)

in planktic foraminifer ecology. Outer and inner shell architecture is analyzed and visualized at high-resolution using X-ray micro-tomography (e.g., Johnstone et al. 2010). Using refined technology, new knowledge has been gained from planktic foraminifer research, and the field has been substantially advanced, but simultaneously a number of intriguing new questions have been raised. Planktic foraminifer assemblages and test properties have become increasingly valuable proxies, and are applied in monitoring climate and environmental change including the position and strength of marine currents and fronts, oxygenation of the water column, and ocean acidification, among others. In 2010, SCOR (Scientific Committee on Oceanic Research) Working Group 138 was formed to synthesize the current knowledge on 'Modern Planktic Foraminifera and Ocean Changes'.

Investigation of modern and geologically ancient planktic foraminifers have diversified

Genus	Species	Author	Year
Beella	digitata	(Brady)	1879
Berggrenia	pumilio	(Parker)	1962
Bolliella	adamsi	Banner and Blow	1959
Candeina	nitida	d'Orbigny	1839
Dentigloborotalia	anfracta	(Parker)	1967
Gallitellia	vivans	(Cushman)	1934
Globigerina	bulloides	d'Orbigny	1826
	falconensis	Blow	1959
Globigerinella	calida	(Parker)	1962
	siphonifera	(d'Orbigny)	1839
Globigerinita	glutinata	(Egger)	1895
	minuta	(Natland)	1938
	uvula	(Ehrenberg)	1861
Globigerinoides	conglobatus	(Brady)	1879
	ruber	(d'Orbigny)	1839
	sacculifer	(Brady)	1877
Globoquadrina	conglomerata	(Schwager)	1866
Globorotalia	cavernula	Bé	1967
	crassaformis	(Galloway and Wissler)	1927
	hirsuta	(d'Orbigny)	1839
	inflata	(d'Orbigny)	1839
	menardii	(d'Orbigny)	1865
	scitula	(Brady)	1882
	theyeri	Fleisher	1974
	truncatulinoides	(d'Orbigny)	1839
	tumida	(Brady)	1877
	ungulata	Bermudez	1960
Globorotaloides	hexagonus	(Natland)	1938
Globoturborotalita	rubescens	Hofker	1956
	tenella	(Parker)	1958
Hastigerina	digitata	(Rhumbler)	1911
	pelagica	(d'Orbigny)	1839
Neogloboquadrina	dutertrei	(d'Orbigny)	1839
	incompta	(Cifelli)	1961
	pachyderma	(Ehrenberg)	1861
Orbulina	universa	d'Orbigny	1839
Orcadia	riedeli	(Rögl and Bolli)	1973
Pulleniatina	obliquiloculata	(Parker and Jones)	1865
Sphaeroidinella	dehiscens	(Parker and Jones)	1865
Streptochilus	globigerus	(Schwager)	1866

Table 1.1 Modern planktic foraminifer morphospecies sorted by genus, including author and year of first description,and page of detailed description given in Chap. 2

(continued)

Genus	Species	Author	Year
Tenuitella	compressa	(Fordham)	1986
	fleisheri	Li	1987
	iota	(Parker)	1962
	parkerae	(Brönnimann and Resig)	1972
Turborotalita	clarkei	(Rögl and Bolli)	1973
	humilis/cristata	(Brady)/Heron-Allen and Earland 1929	1884
	quinqueloba	(Natland)	1938

Table 1.1 (continued)

substantially since the first discoveries (see, e.g. the reviews and books of Vincent and Berger 1981; Hemleben et al. 1989; Murray 1991; Schiebel and Hemleben 2005; Kucera 2007). An enormous wealth of information is available from textbooks, printed papers, online publications, and various Internet sites (e.g., www.species-identification.org, www.EMIDAS.org, www. eforams.org). Many more researchers and working groups, beyond those referred to above, have added an enormous wealth of knowledge, which is presented in the following topical Chaps. 2–10.

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